

Intermediate-depth earthquakes in subduction zones: insights from high pressure experiments

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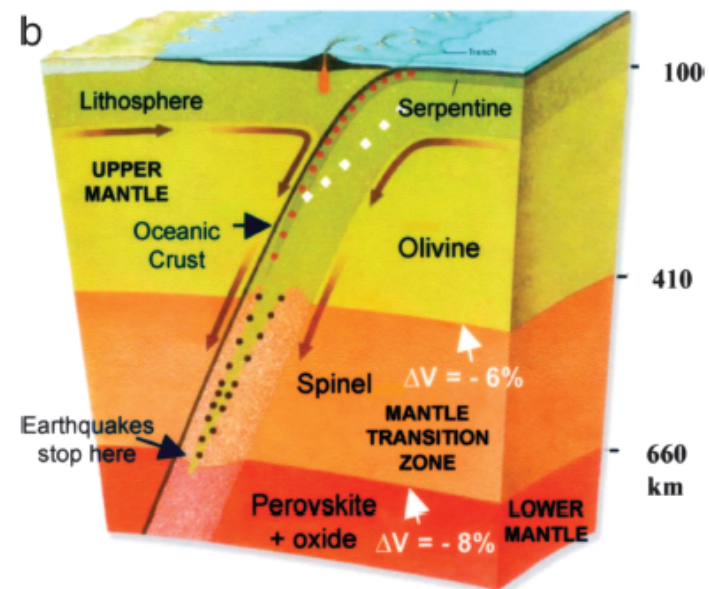
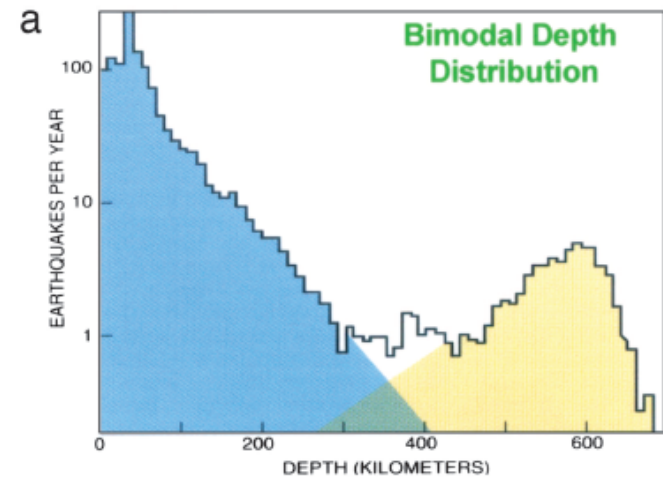
7 - UC Riverside, USA

Earthquakes at depths

Earthquakes after ca. 50 km depth are not explained by friction theories.

... are these earthquakes induced *somehow* by mineralogical reactions ?

(and how do we test for this experimentally?)

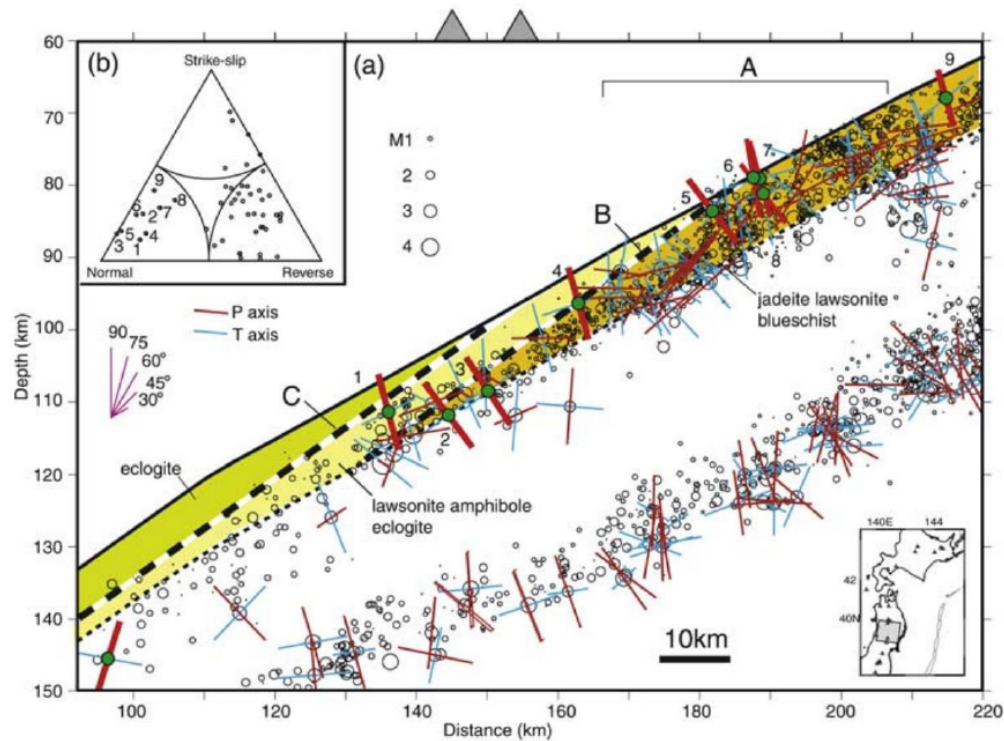


Green, PNAS 2007

Earthquakes at intermediate depths and mineralogical reactions

Candidates:

- reactions in the oceanic crust (eclogitisation)
- dehydration reactions in the mantle
- *Small grain sizes produced*
- *Fluid release -> decrease in effective pressure*
- *...??*



Kita et al, 2006

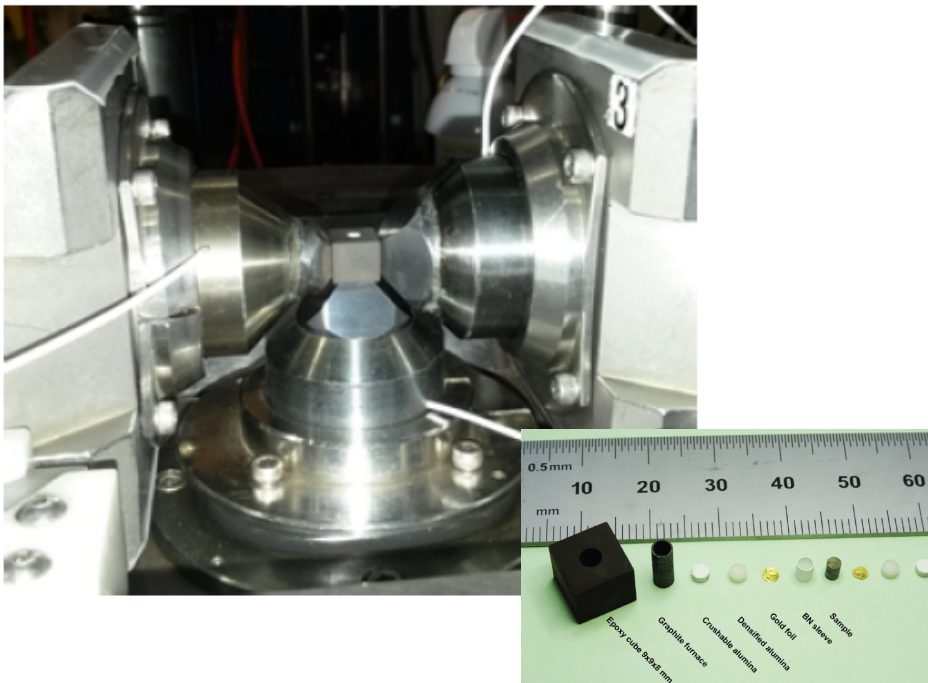


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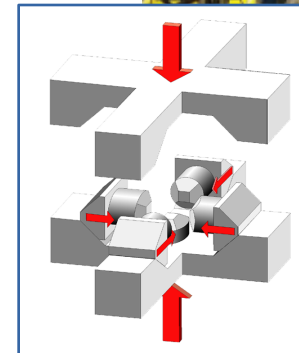
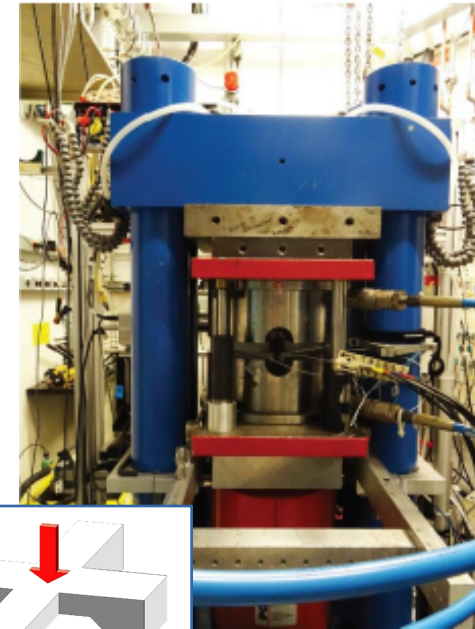
Investigate brittle-like behavior potential under
high pressures

Generating large pressures (and deformation)

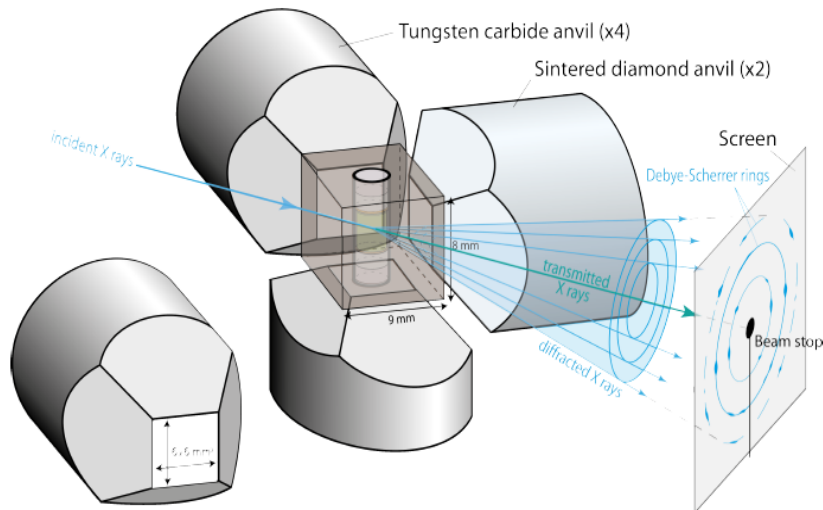
Experimental deformation under high pressures ($> 2\text{GPa}$):
Deformation-DIA multianvil press



T. Ferrand 2017



+ follow in-situ stress and mineralogical reaction progress :

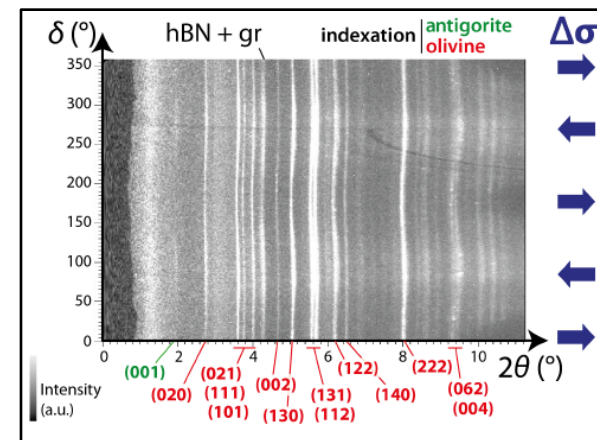
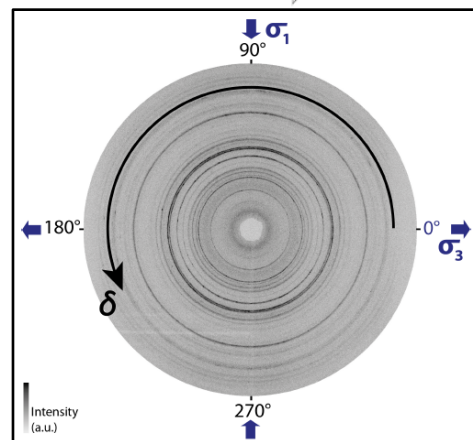


Access to

- crystals and phases
- stresses rather than F/S
- phases proportions
- and reaction progress

In-situ monitoring
using synchrotron X-
rays diffraction
(GSE-CARS, APS)

T. Ferrand, *phD thesis*.





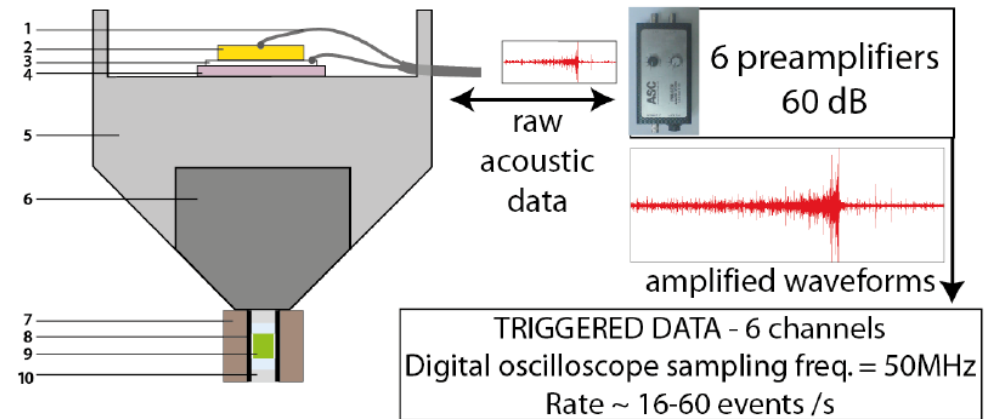
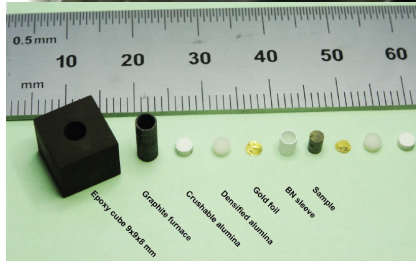
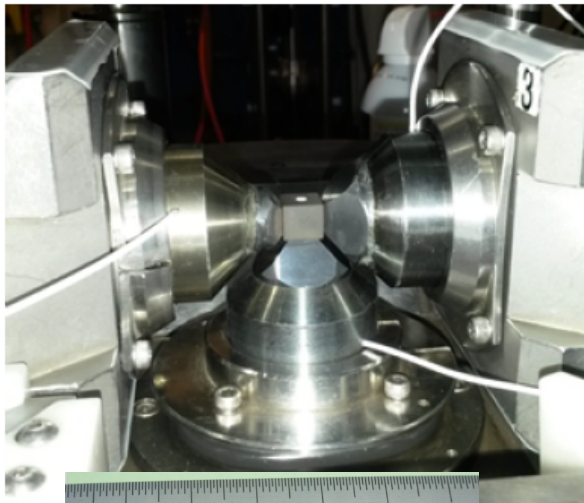
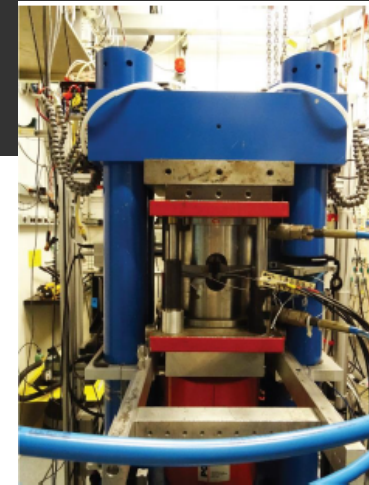
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... listening to rocks under high pressures?

= add acoustic emissions
monitoring

Listening to rocks under high pressures?

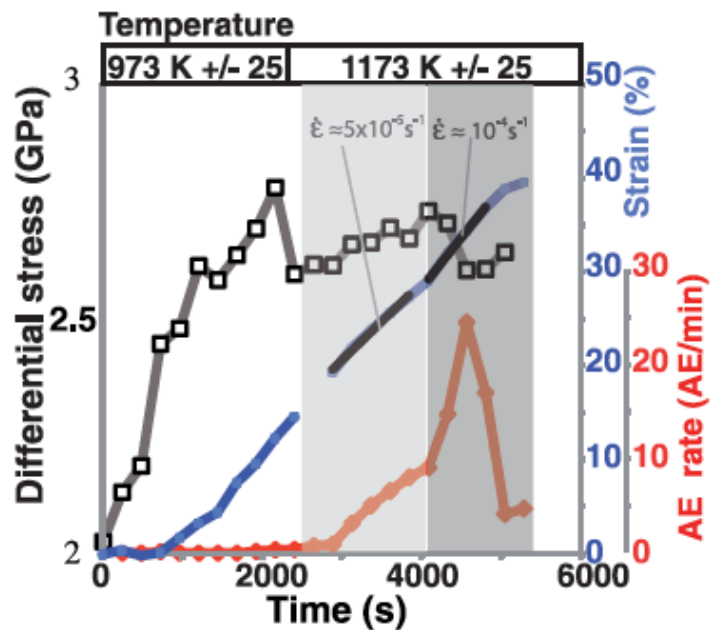
Coupling Deformation-DIA (HP-HT) and acoustic emission (AE) recording



Gasc et al, 2011

« Proof of concept » study

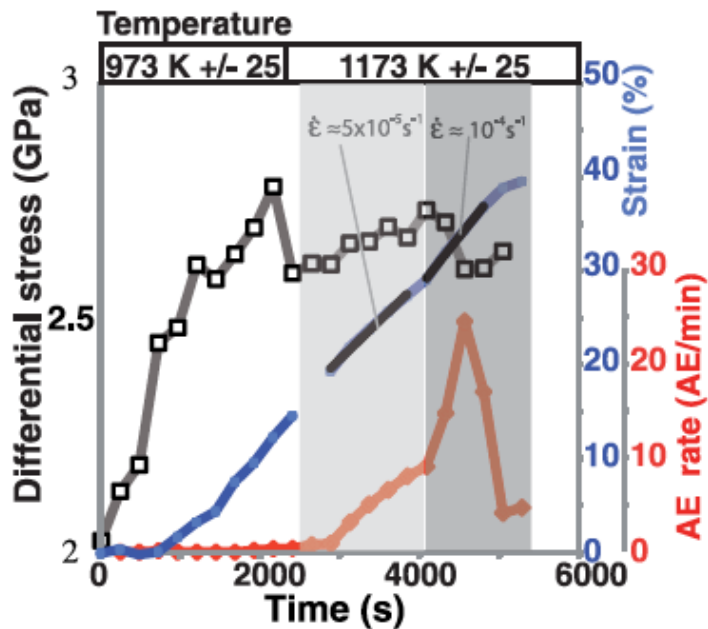
Metastable olivine \rightarrow spinel in Mg_2GeO_4



Schubnel et al, Science 2013

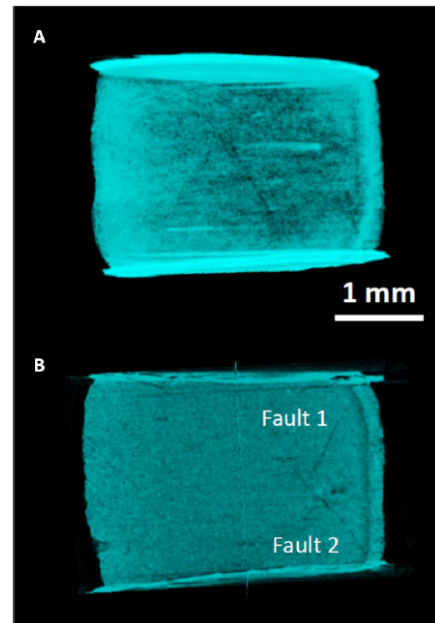
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Metastable olivine \rightarrow spinel in Mg_2GeO_4



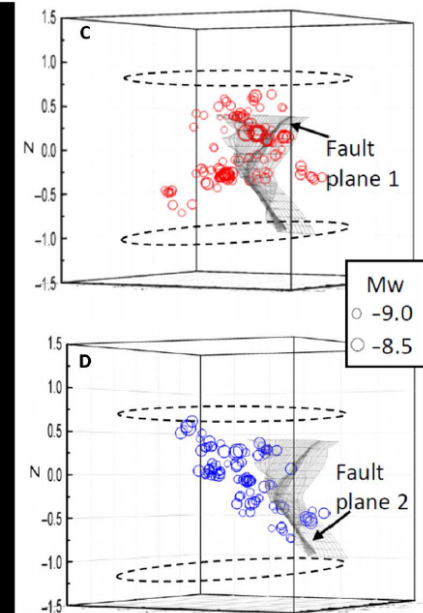
Schubnel et al, Science 2013

Post-experiment
tomography



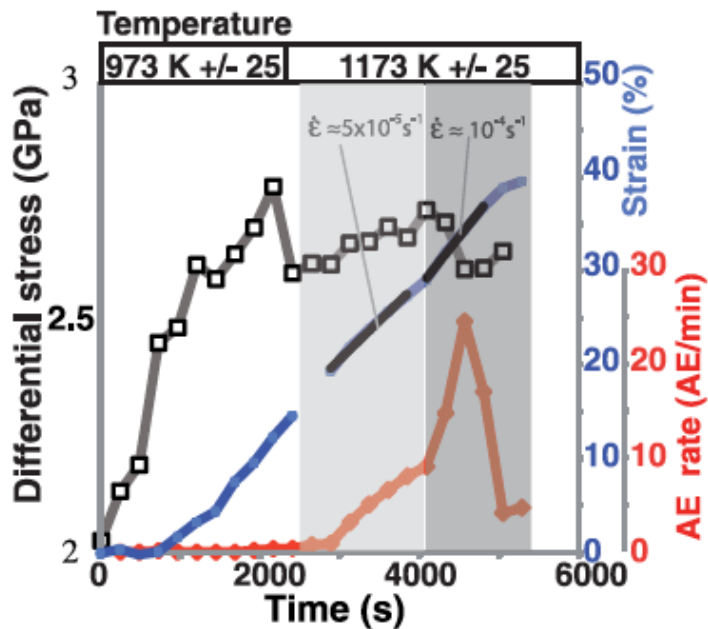
Wang et al, Sci. Adv. 2017

Relocated AEs

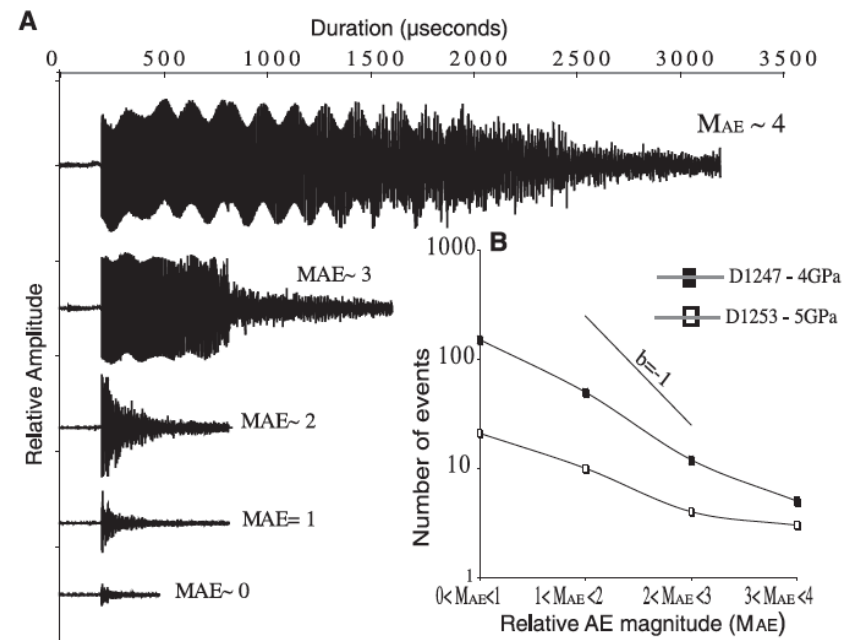


« Proof of concept » study

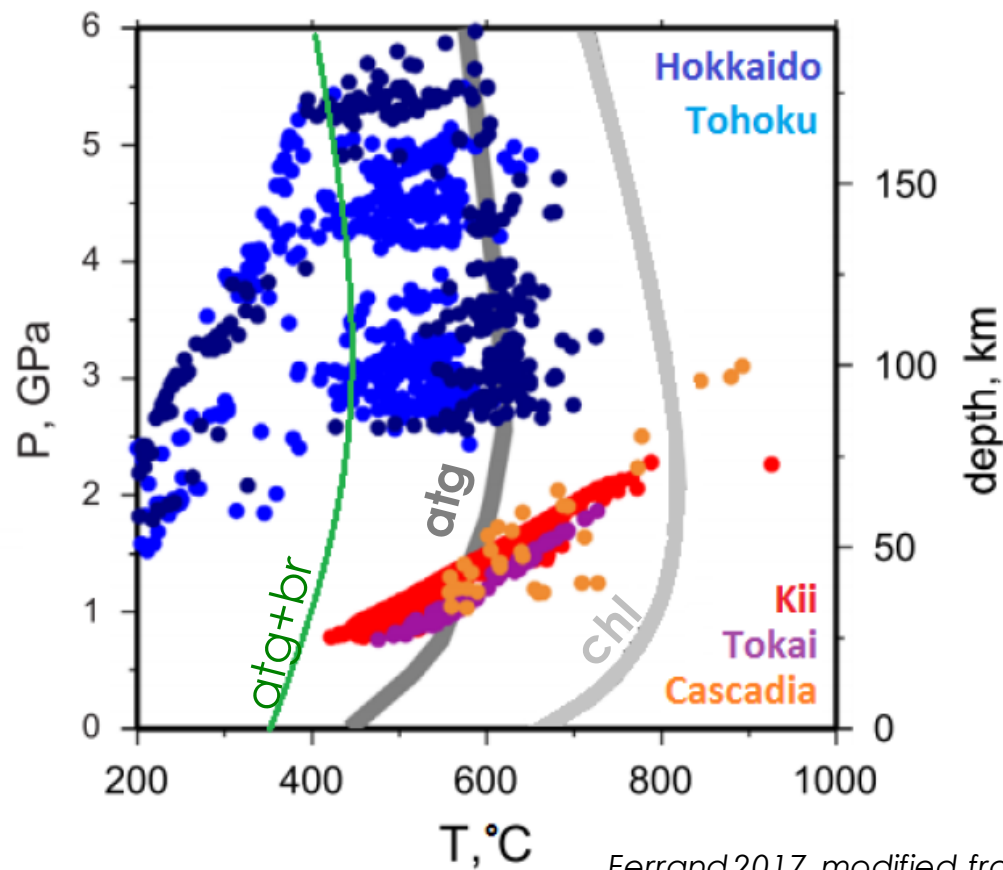
Metastable olivine \rightarrow spinel in Mg_2GeO_4



Schubnel et al, Science 2013

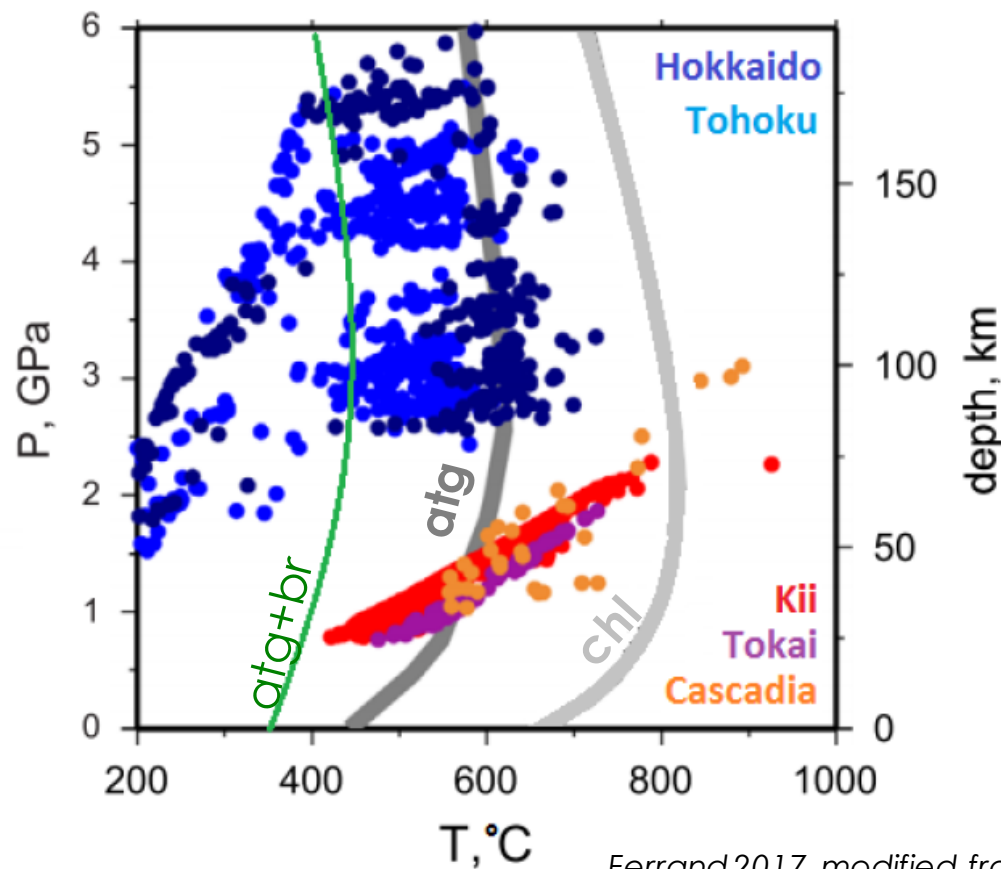


Lower Wadati-Benioff plane and dehydration reactions in the slab mantle



Ferrand 2017, modified from Abers et al, 2013

Lower Wadati-Benioff plane and dehydration reactions in the slab mantle



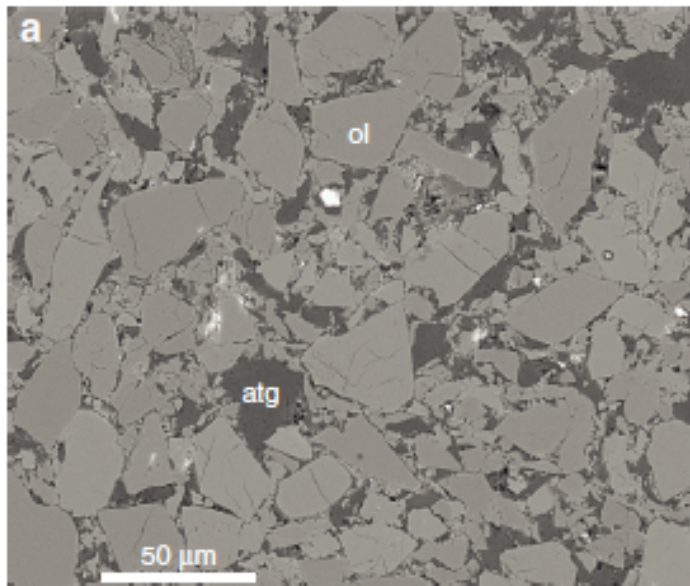
Ferrand 2017, modified from Abers et al, 2013

Mechanical stability of serpentinite upon dehydration and stress ?

- Dobson et al 2002, Jung Green and Dobrzhinetskaya, 2004 fluid release causes « dehydration embrittlement »
- Gasc et al. 2011, Hilairet and Reynard, unpub., Chernak and Hirth, 2011, Okazaki and Hirth 2016, stable.

Experiments

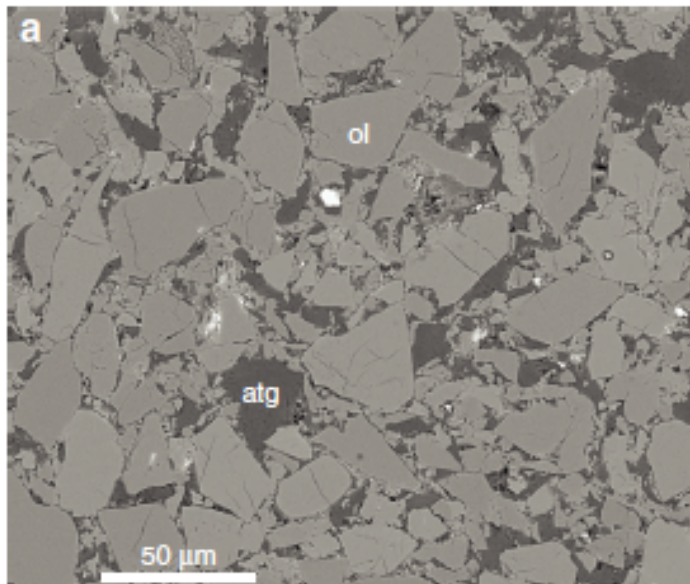
- ❑ Cold pressed antigorite + olivine aggregates
(50:50, 20:80, 5:95, 0:100)



Starting material / T. Ferrand.

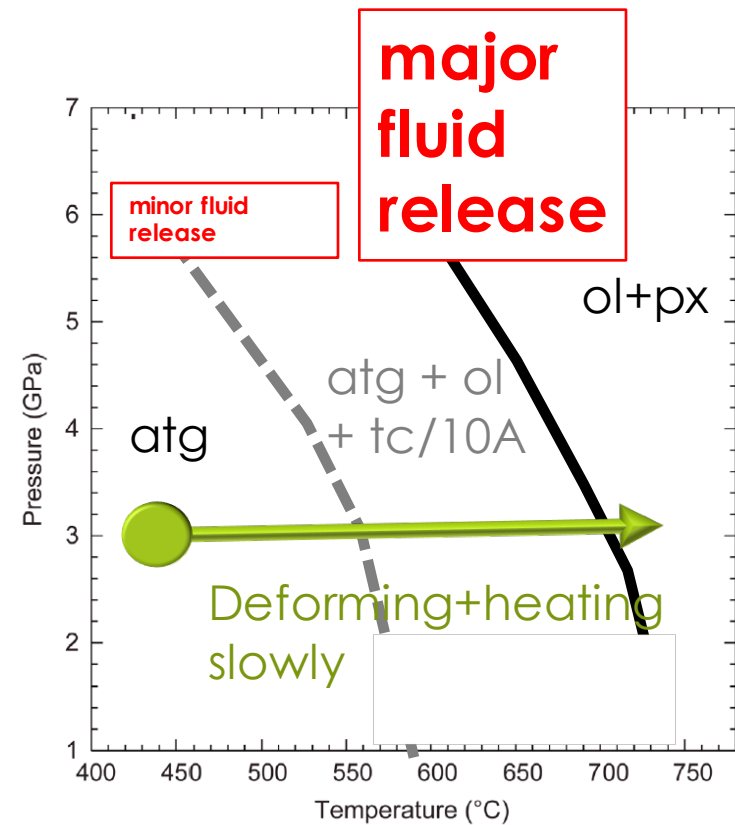
Experiments

- ❑ Cold pressed antigorite + olivine aggregates (50:50, 20:80, 5:95, 0:100)
- ❑ $\dot{T} / \dot{\epsilon}$ is about 1200 K (below Okazaki and Hirth 2016, Chernak and Hirth 2011, and Gasc et al, 2017)



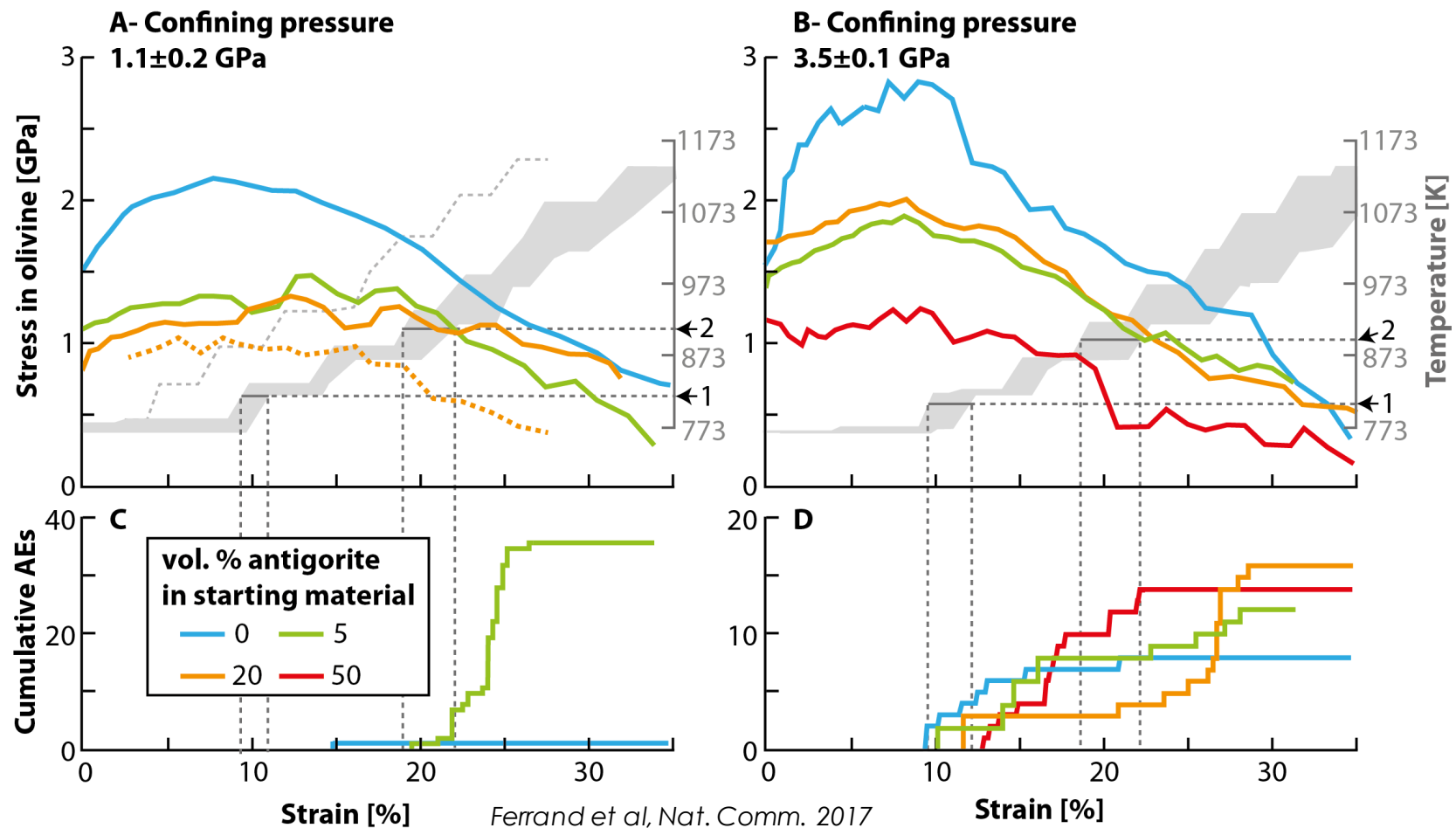
Starting material / T. Ferrand.

Experiments (vs. nature)...

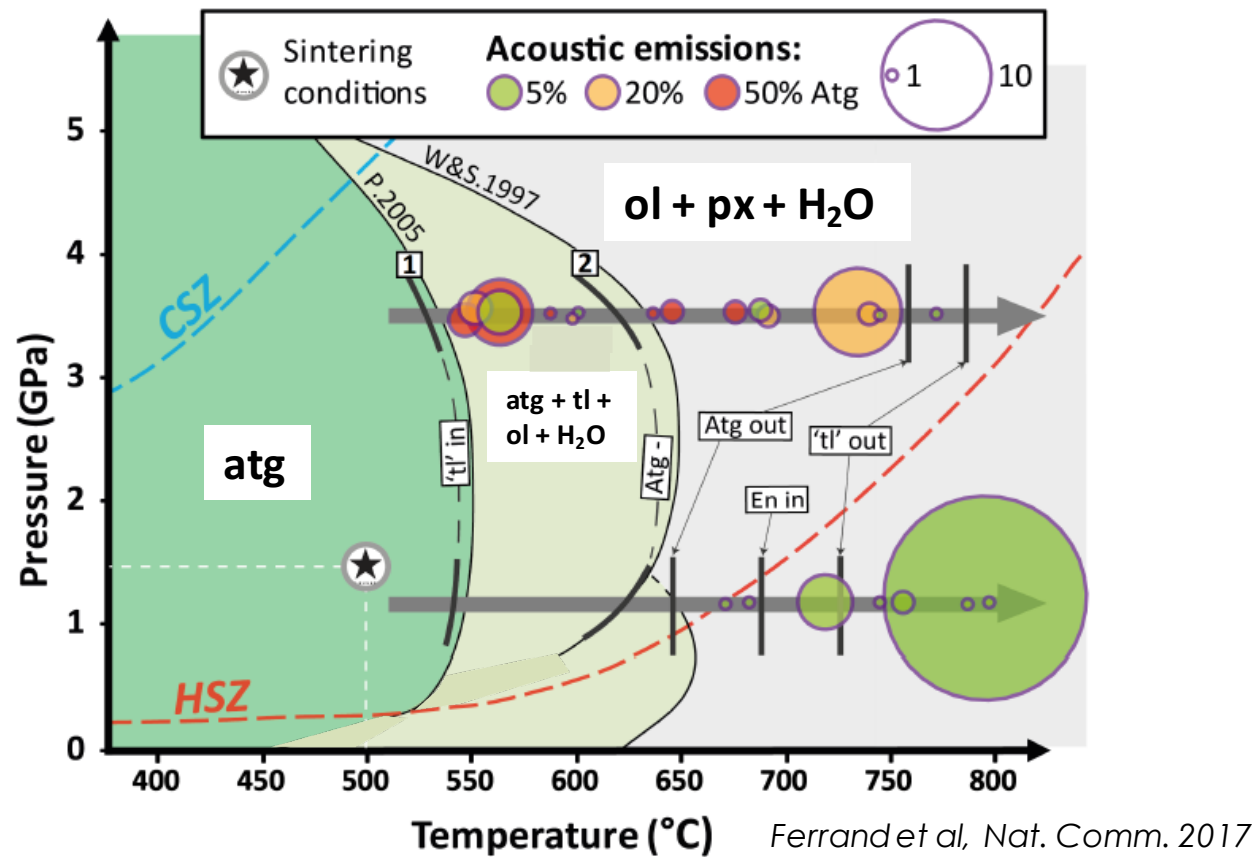


modified from Chollet et al., JGR 2011

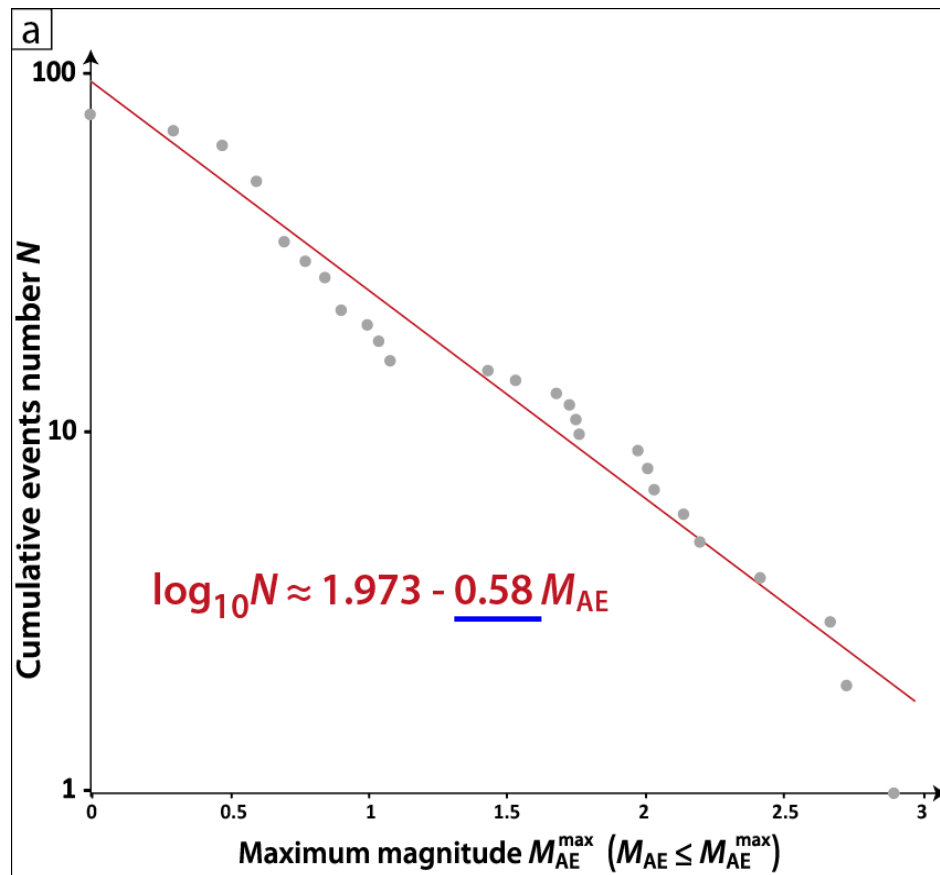
Ol + atg aggregates and 100% olivine: Stresses and AE records



... vs. phase P-T fields



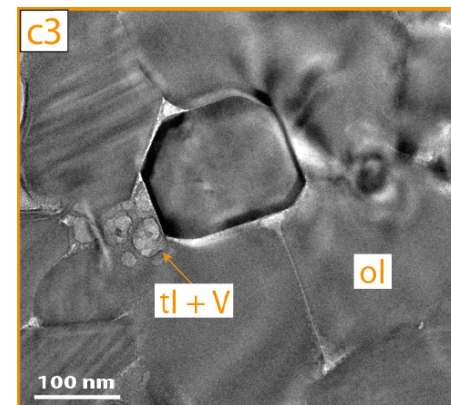
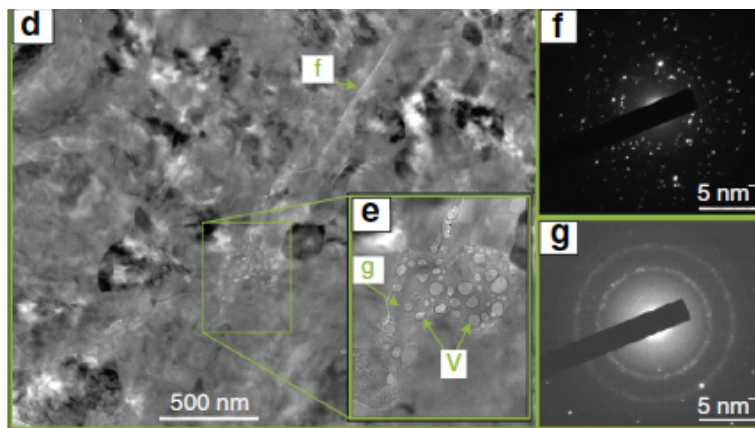
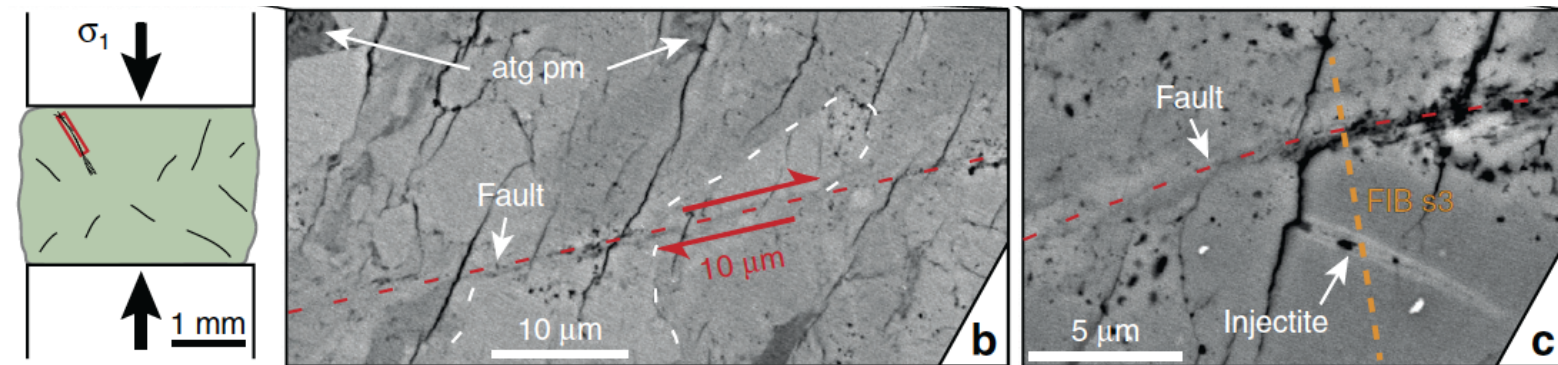
AE statistics



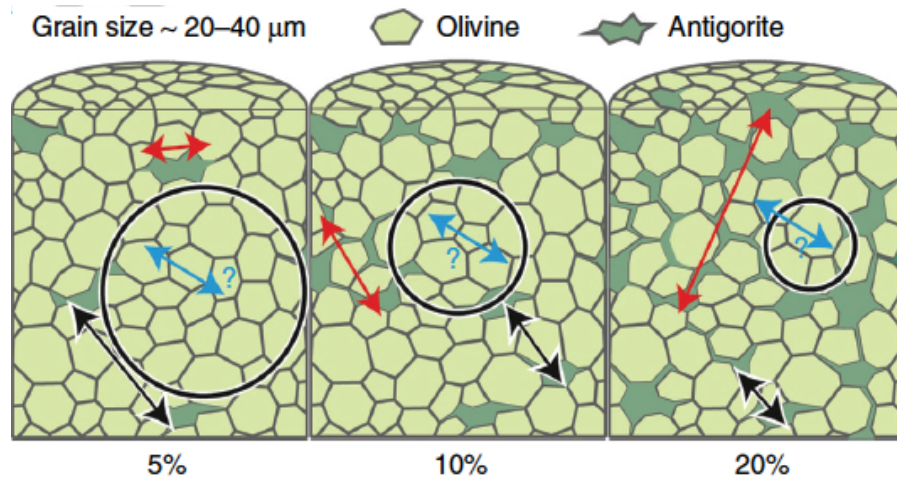
Ferrand et al, Nat. Comm. 2017

microstructures

5% antigorite 1.1 GPa



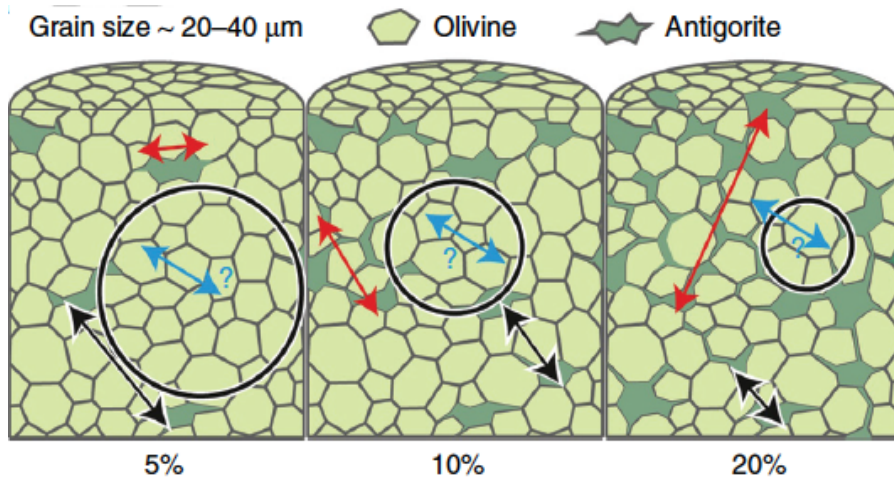
A dehydration-driven stress transfer model based on percolation theory



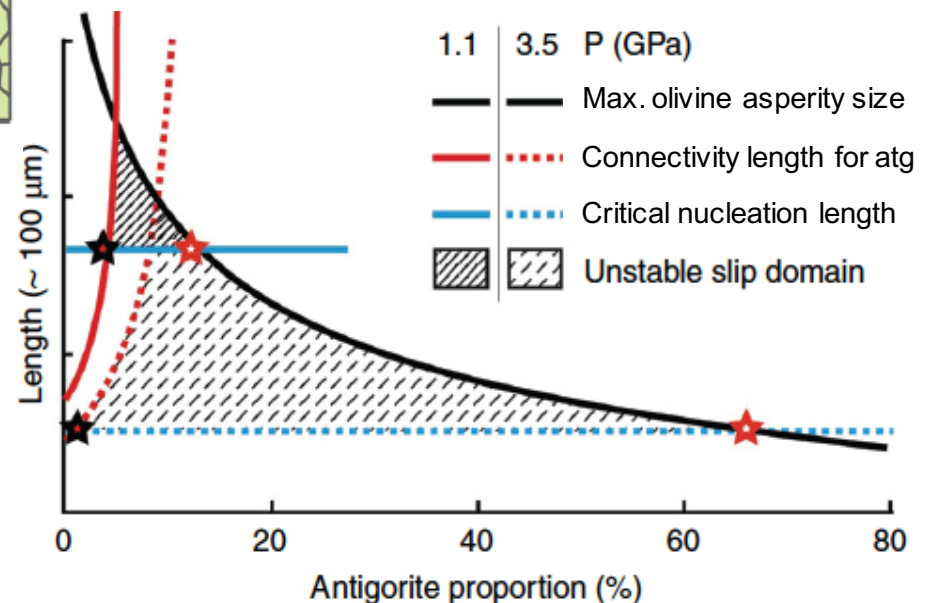
Antigorite stress goes to zero close
to dehydration

Connectivity length for antigorite
Maximum olivine asperity size
Critical nucleation length

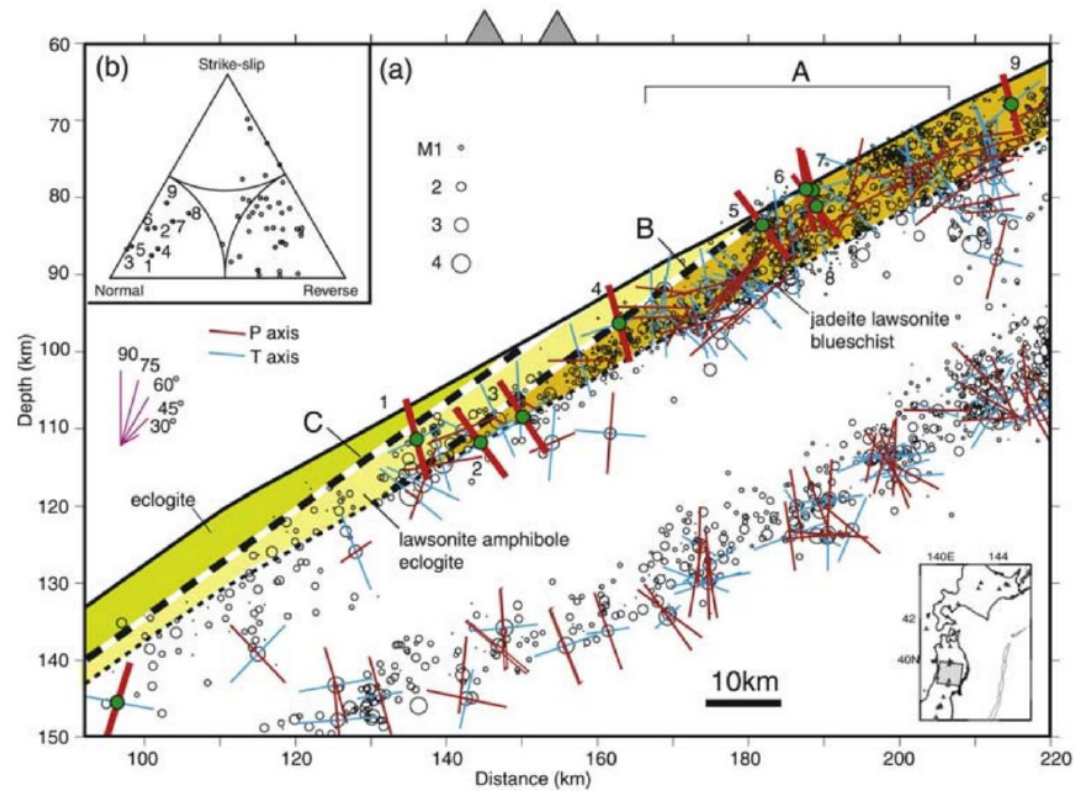
A dehydration-driven stress transfer model based on percolation theory



- Olivine 'asperities' need to be larger than the critical fracture nucleation length
- Atg clusters need to be large enough to transfer a significant amount of stress
- trigger \neq failing material



What about oceanic crust eclogitisation ?

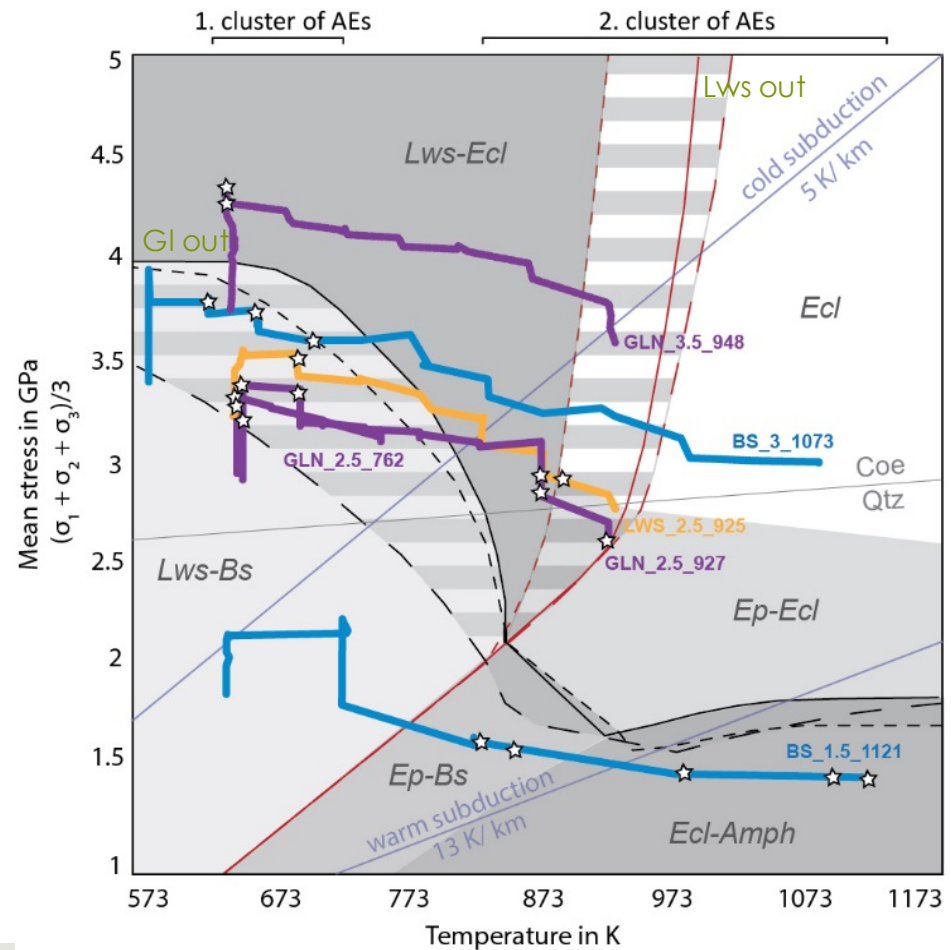


Kita et al, GRL 2006

What about oceanic crust eclogitisation ?

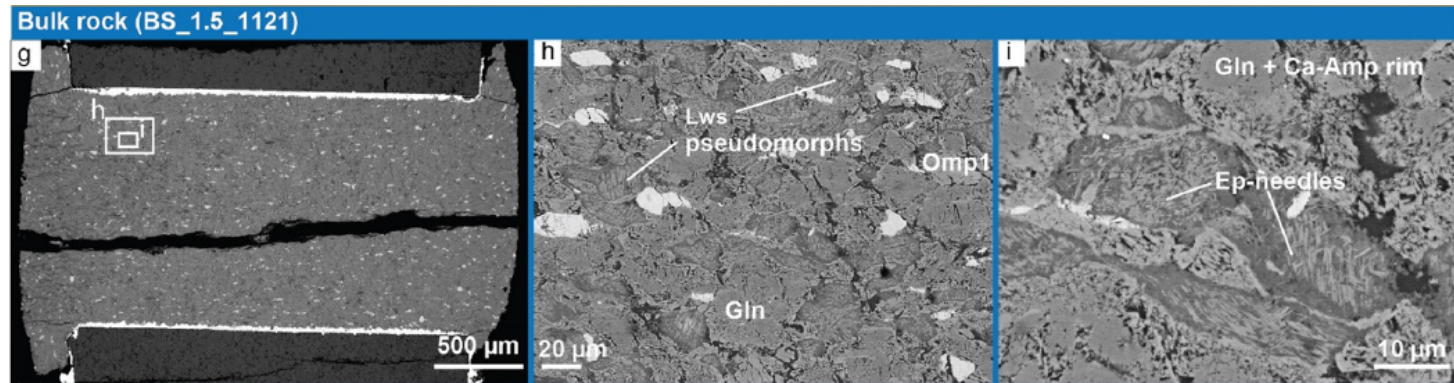
Incel et al, EPSL 2017

- lawsonite + glaucophane in various proportions
- natural blueschist
- AE are not specifically associated with a major fluid release (ie. lws breakdown).

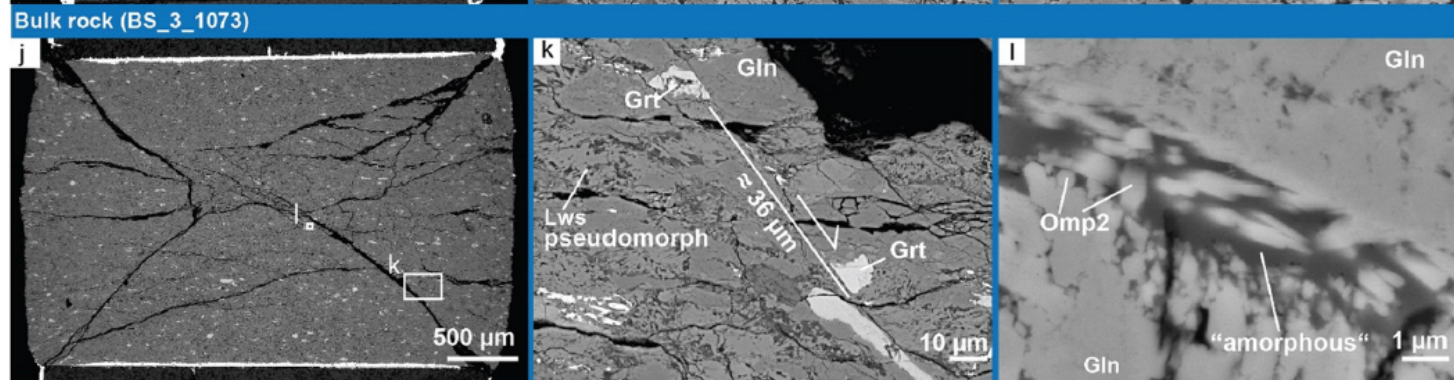


Microstructures and AEs

1.5 GPa:
 only lws
 breakdown
 AEs, no visible
 failure



3 GPa:
 omphacite grain
 growth (in situ +
 post- microstr)
 correlated with
 AEs, failure



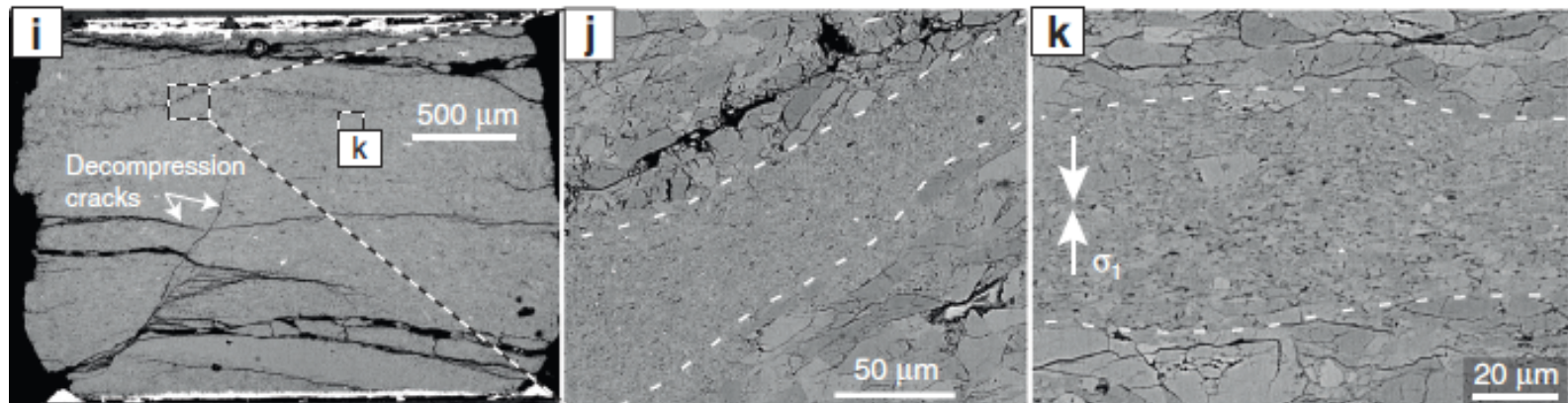
Incel et al, EPSL 2017

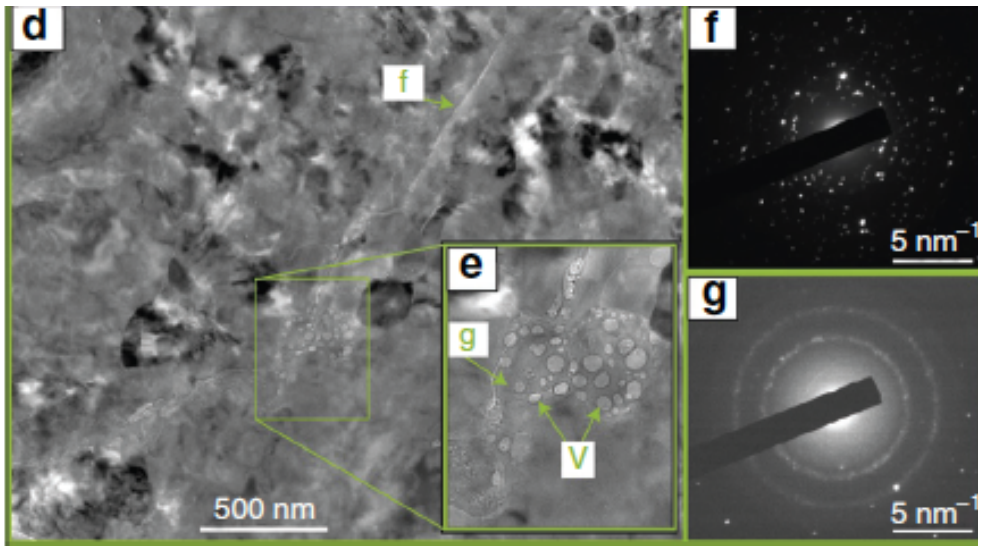
Summary and open questions

- Coupling of in-situ reaction, stresses and AE monitoring was crucial in this work
- Dehydration with large water releases (serpentine, lawsonite) do not (necessarily) induce sample failures
- Rather, distribution of stresses within heterogeneous aggregates may be one important key (grains sizes, minerals with different mechanical properties).
- Some mineralogical reactions may be privileged as a trigger for dynamic sample failure, not the ones previously thought of.

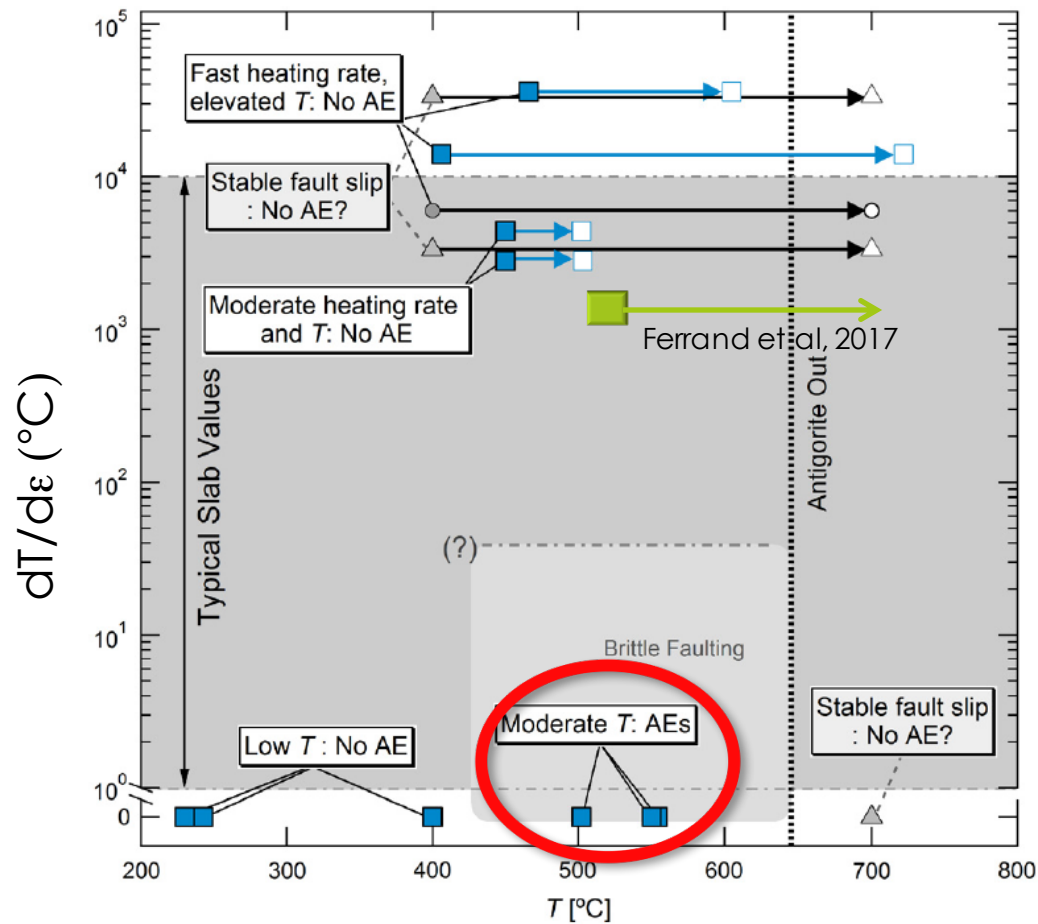


Olivine 100%, 3.5 GPa





The scaling problem



Gasc et al, EPSL 2017

■ This study

▲ Chernak and Hirth, Geology 2011

● Okazaki and Hirth, Nature 2016

□ AEs only within Atg stability field

□ = experimental inverse brittle-to-ductile transition (e.g. Proctor et al, JGR 2016)