

Seismic Evidence for Magma Assisted Continental Rifting

Michael Kendall
University of Bristol

or

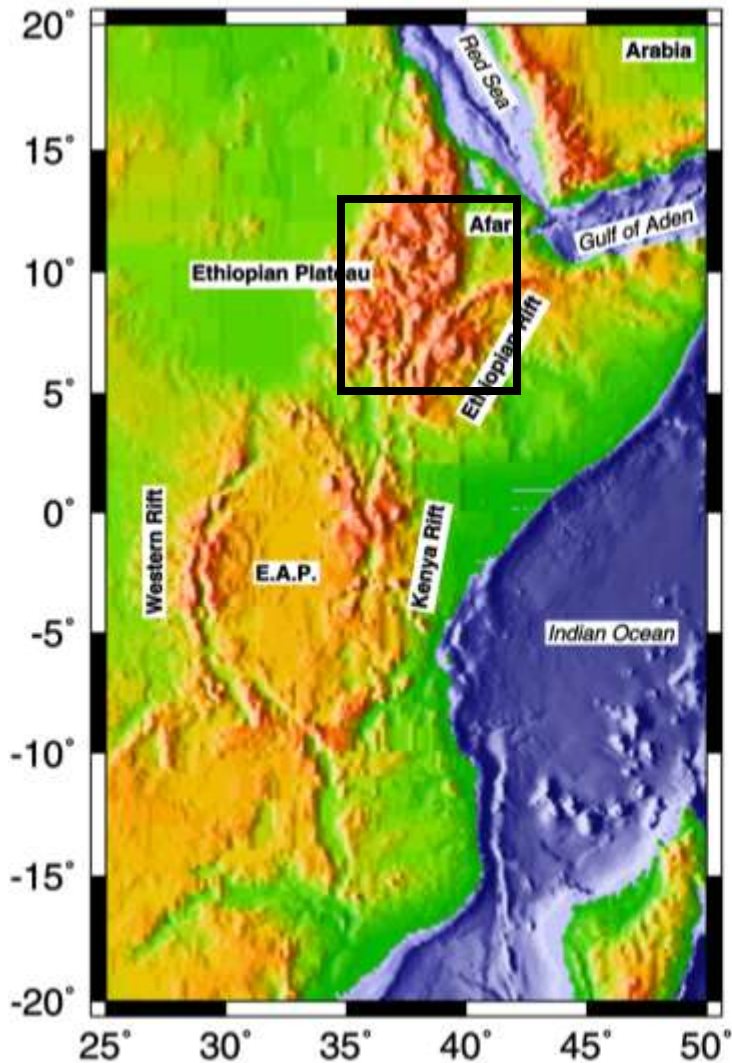
10 years of fieldwork in Ethiopia



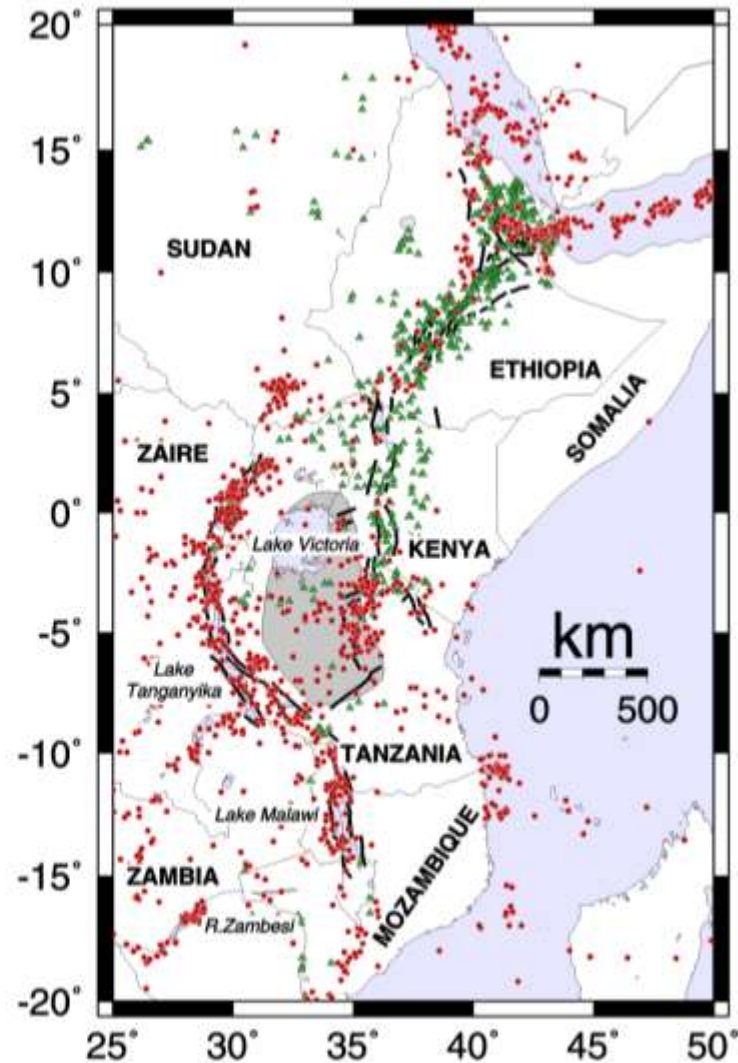
- Atalay Ayele, Elias Lewi – Addis Ababa University
- Mike Kendall, **Anna Stork**, James Wookey – Bristol
- Nick Johnson, Kathy Whaler – Edinburgh
- Ghebrebrhan Ogubazghi, Berhe Goitom – Eritrea Inst of Tech
- **James Hammond**, Ian Bastow – Imperial College
- Graham Stuart, **Dave Thompson**, Tim Wright – Leeds
- Cindy Ebinger, **Manahloh Belachew** – Rochester
- **Catherine Rychert**, Derek Keir – NOC Southampton



East-Africa Rift System: juvenile continental rifting to oceanic spreading

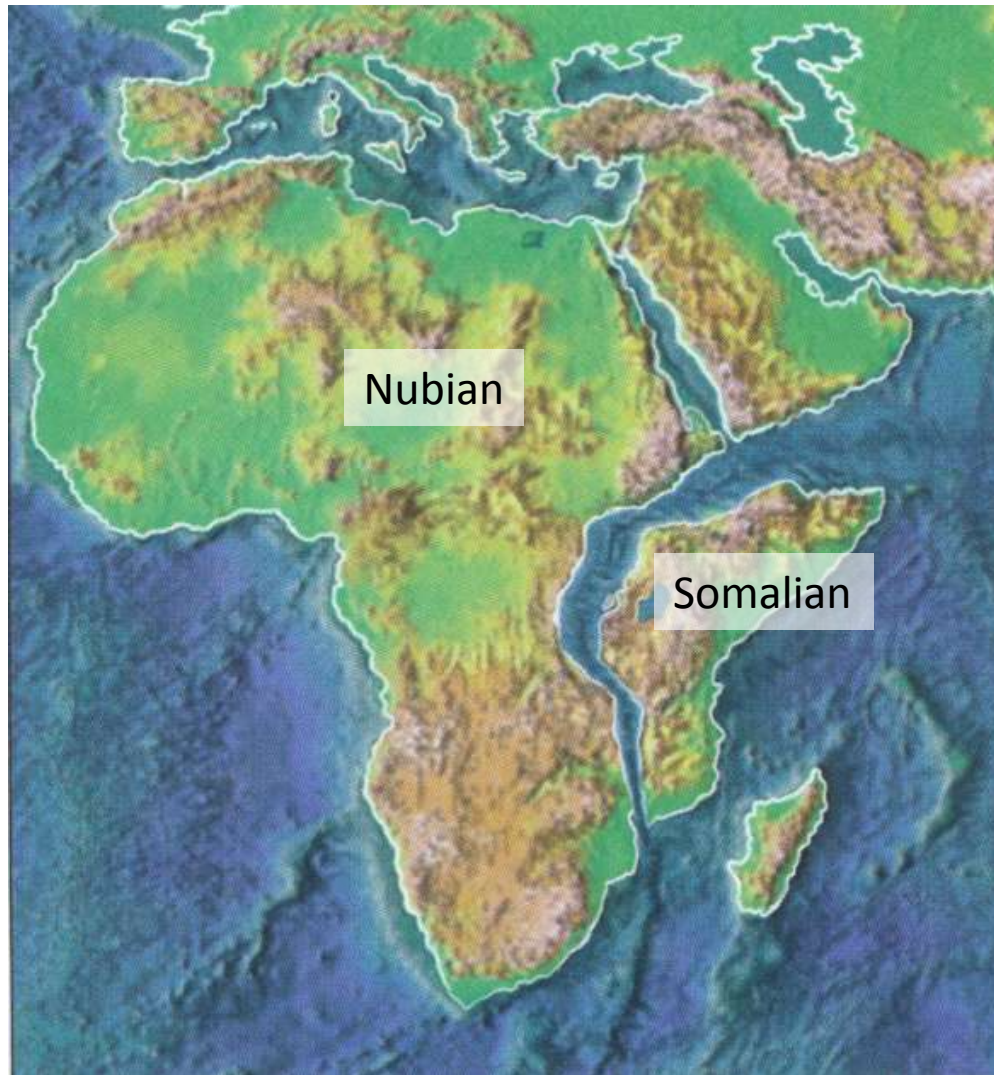


East Africa - Topography
(E.A.P. - East African Plateau)



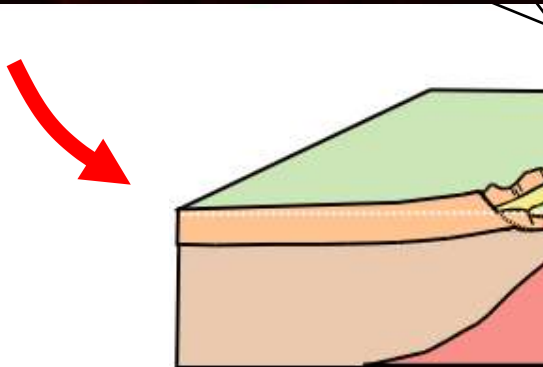
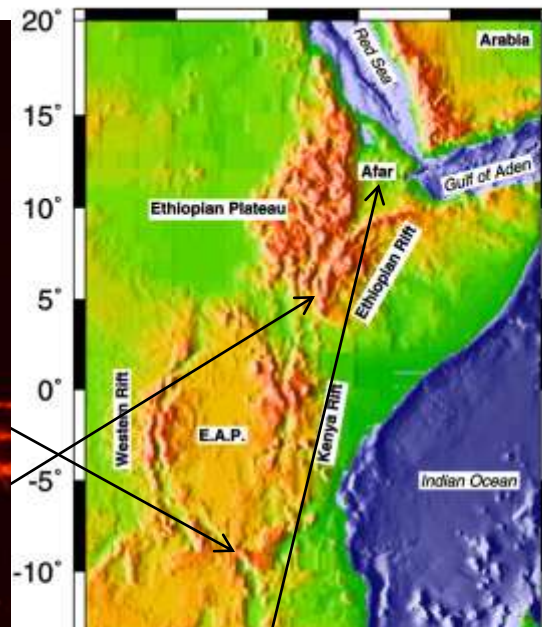
East Africa - Seismicity (Dots) and Volcanoes (Triangles).
(The Nyanza craton is shown by the shaded pattern)

Birth of an Ocean



- Continental rifting - fundamental component of plate tectonics
- Rift valley has played a key role in the evolution of humans – climate, habitability
- How and why do continents break apart?





Fundamental process – but how does it work

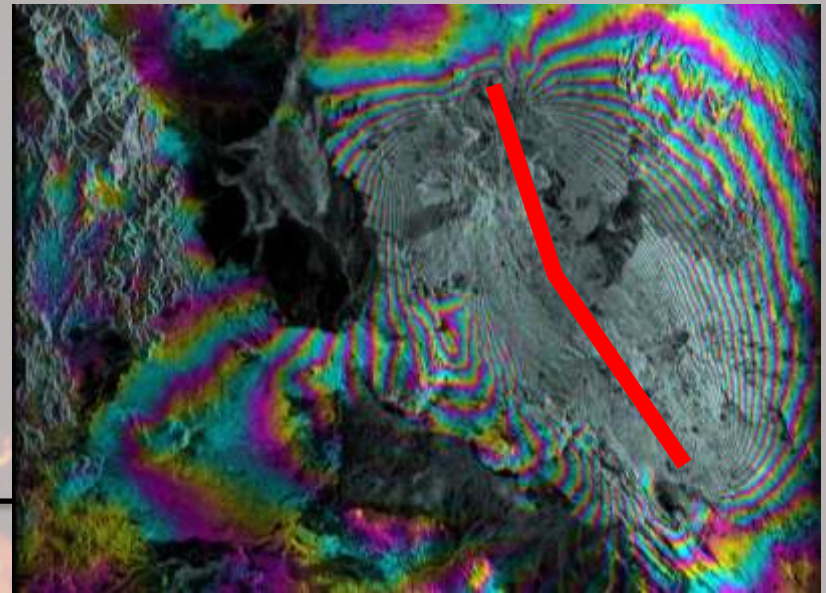
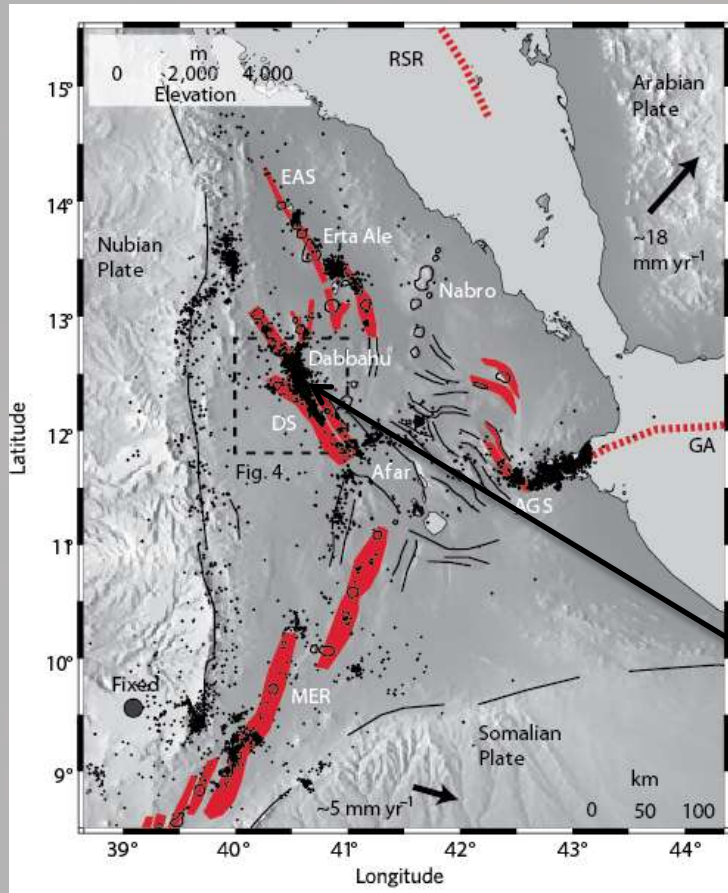
The volcano-seismic crisis in Afar - Sept, 2005



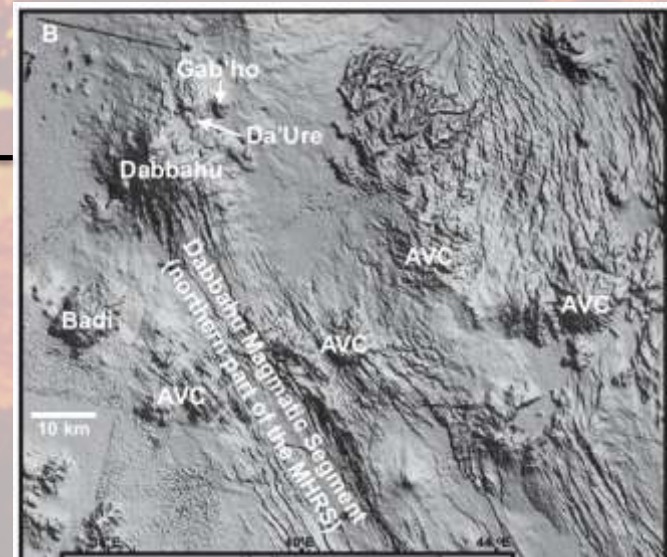
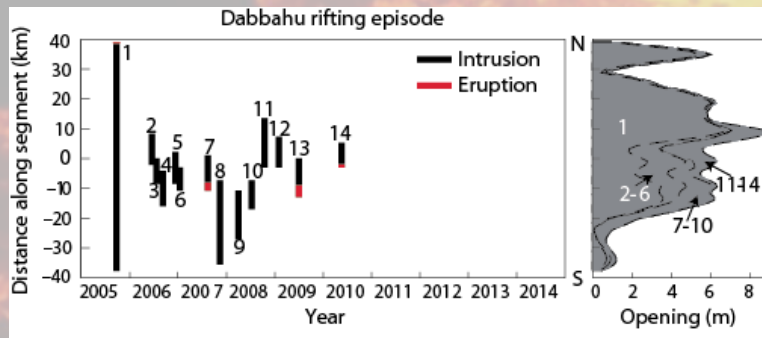
Afar (Sept 2005): a “once in a generation” event.

- 60 km of plate boundary opened by up to 8 metres.
- 2.5 km³ of new crust created.
- First rifting event above sea level in the era of satellite geodesy.
- First eruption of rhyolite in Africa in over a century.
- UK/Ethiopian/US team first on the ground.





Wright et al., 2006

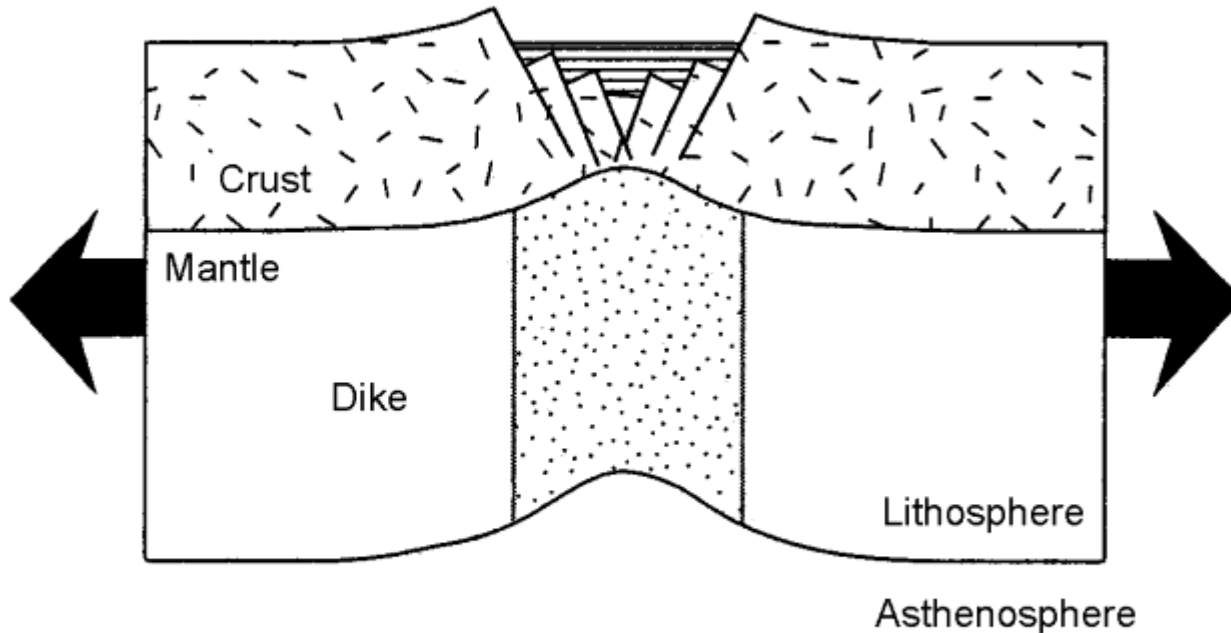


Field et al., 2013

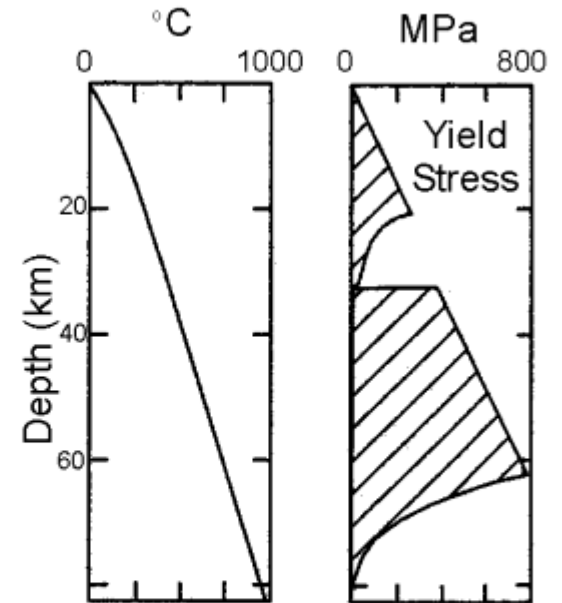
Wright et al., 2012

Breaking a plate.

Tectonic Stretching



Broad zone of stretching

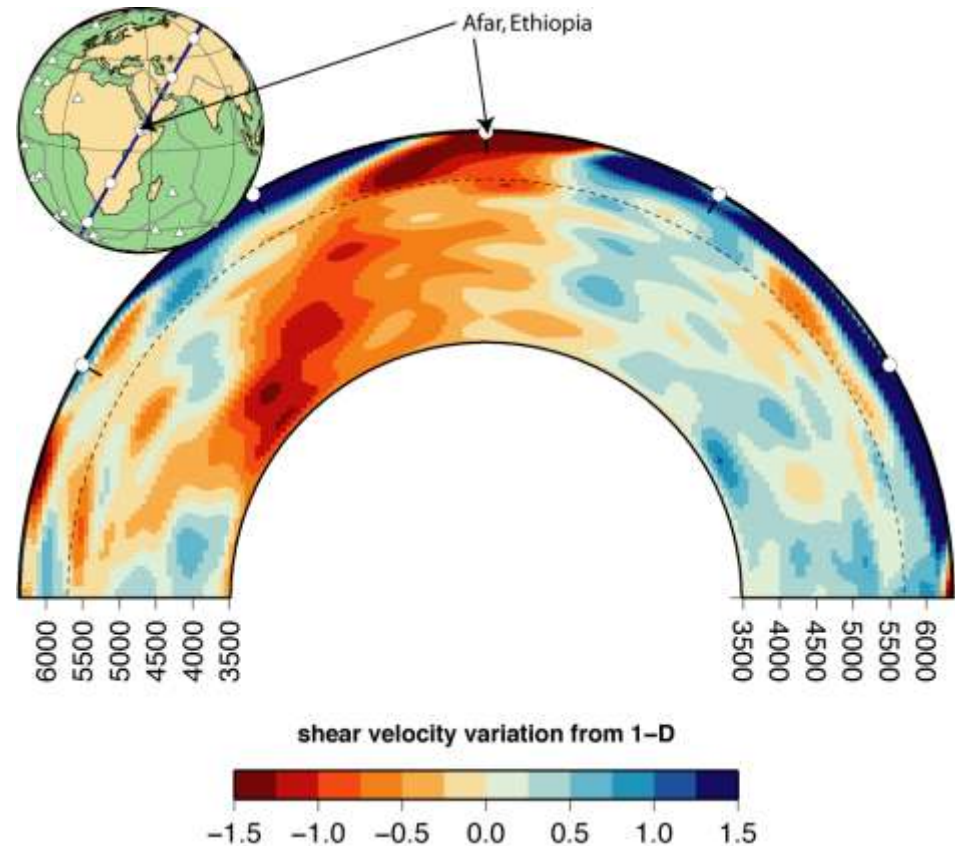
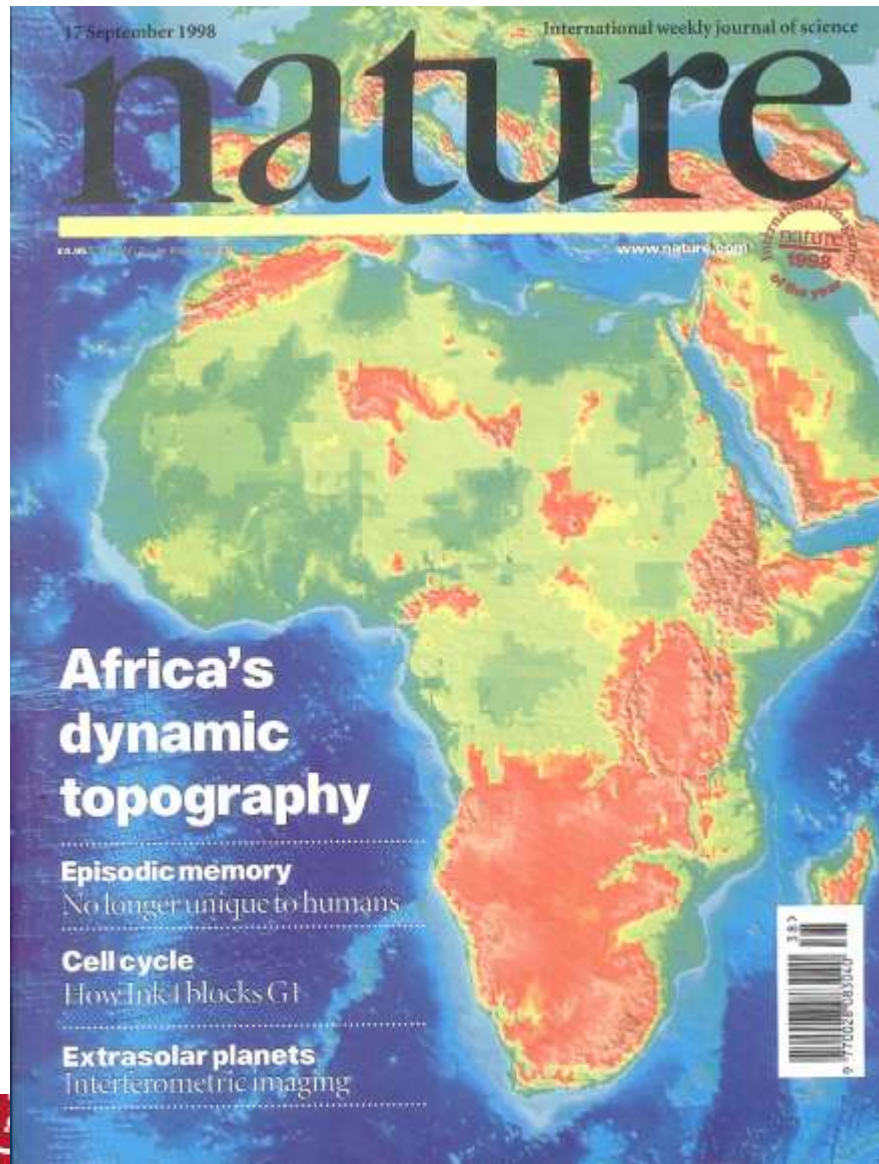


Buck, 2004

So why is Africa rifting?

- Forces available for rifting
 - Distant subduction
 - Gravitational potential energy
 - Basal traction – compression vs extension
 - Preexisting weaknesses and lithospheric thin spots

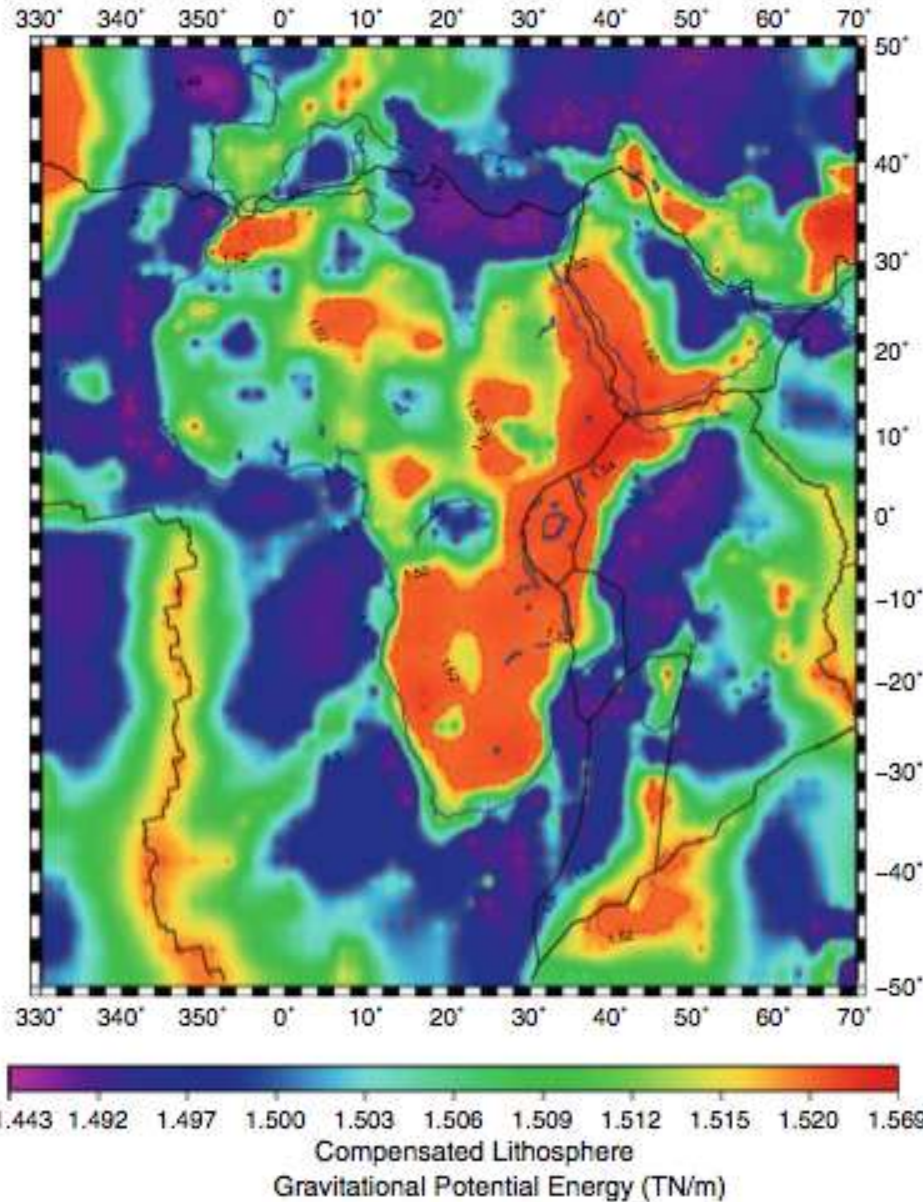
Africa: superplume ... superswell



Ritsema and Allen 2003

Lithgow-Bertelloni and Silver 1998

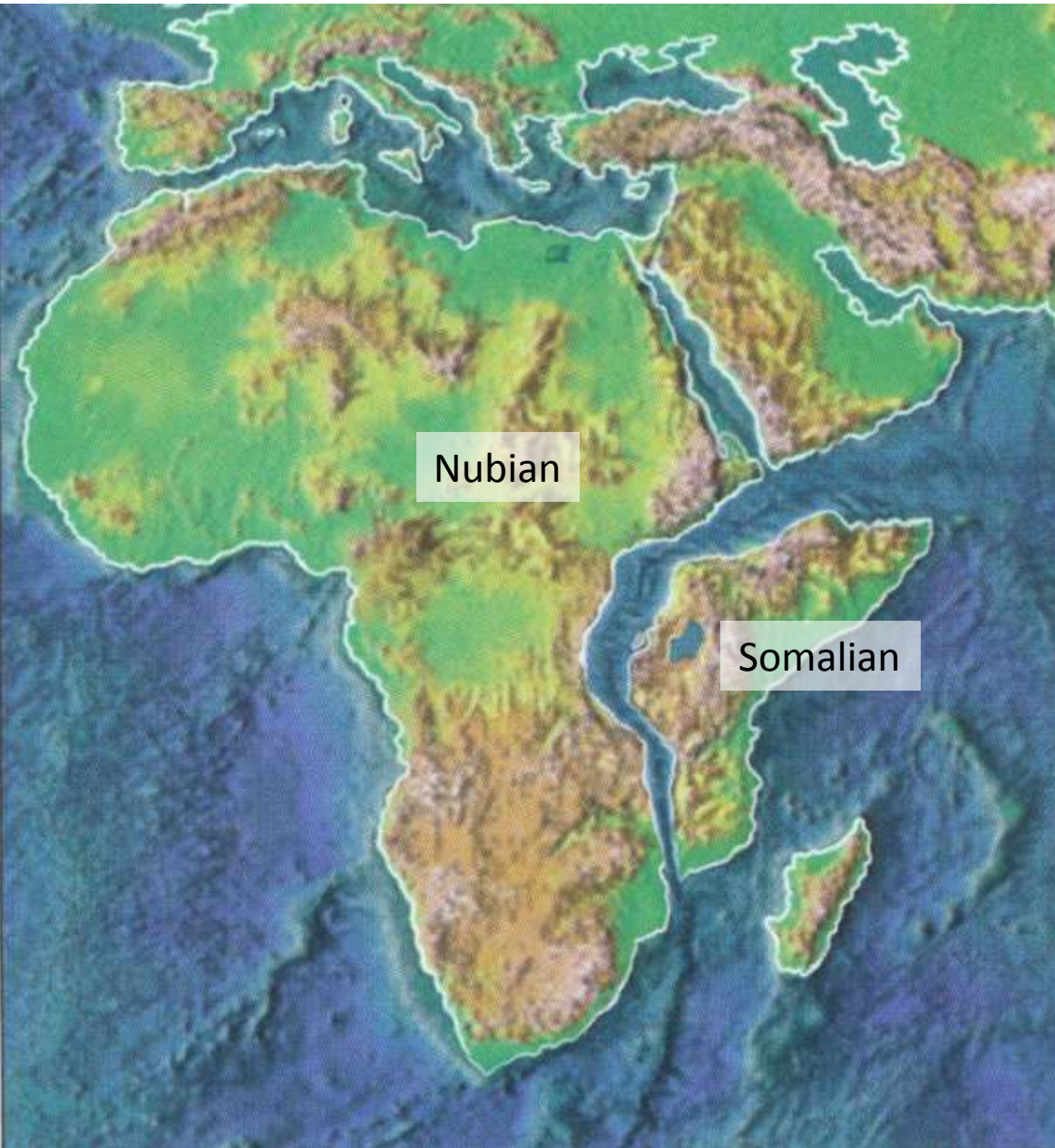
Gravitational potential energy



“ Once quantified, it appears that deviatoric stresses alone are not sufficient to overcome the strength of the continental lithosphere in the Eastern rift.”
Stamps et al., 2010

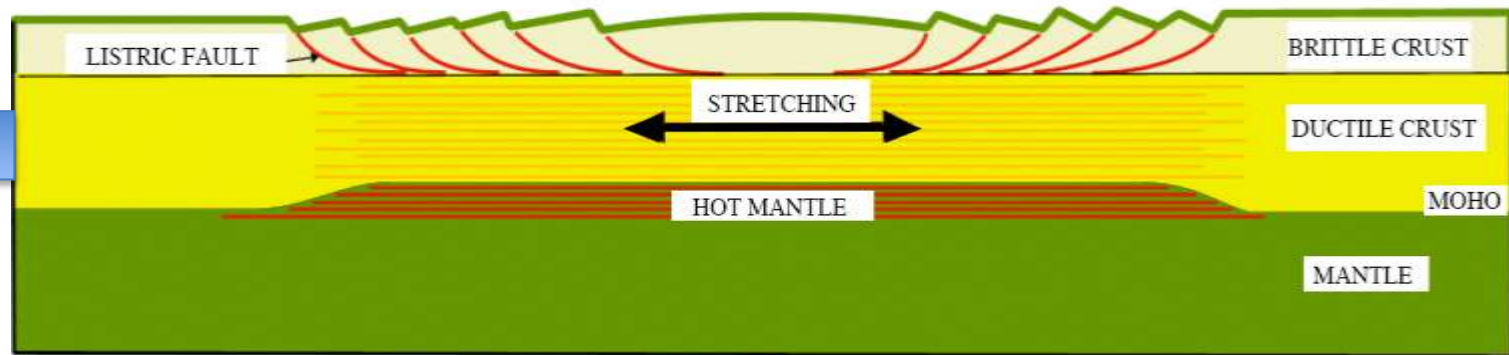
Plate Force Paradox:

Plate forces are up to an order of magnitude too small to break thick cold continental lithosphere (Buck 2004).



- Continental rifting - fundamental component of plate tectonics
- How and why do continents break apart?

(a) Pure Shear



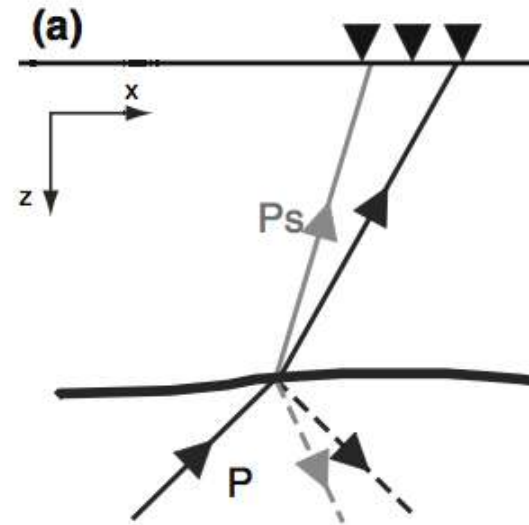
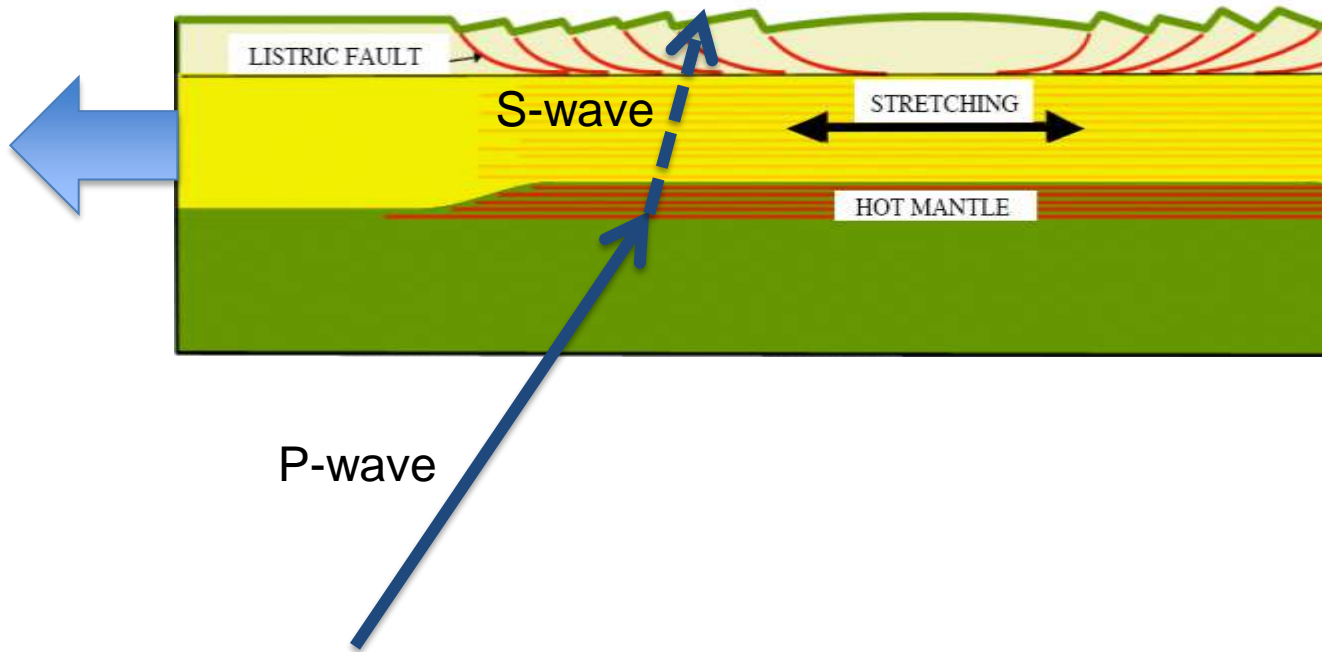
McKenzie 1978

Which seismic methods can be used to address this question?

Seismic methods for imaging

- Receiver functions – discontinuity structure;
 V_p/V_s ratios

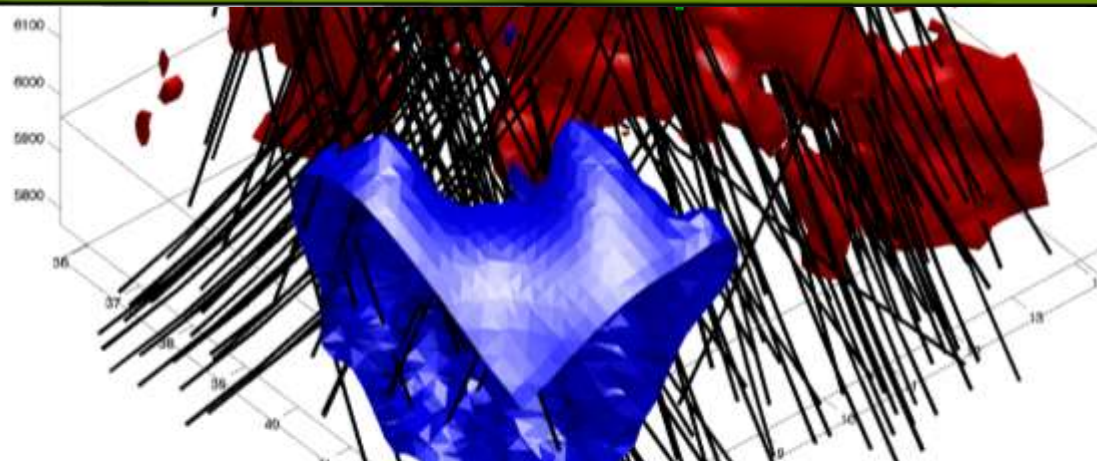
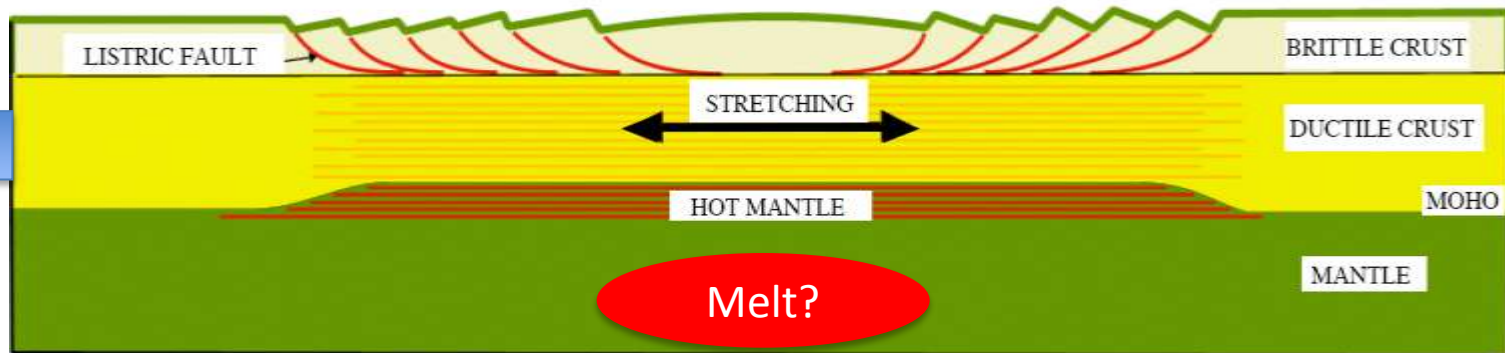
(a) Pure Shear



Seismic methods for imaging

- Tomography – velocity structure; thermal anomalies; partial melt

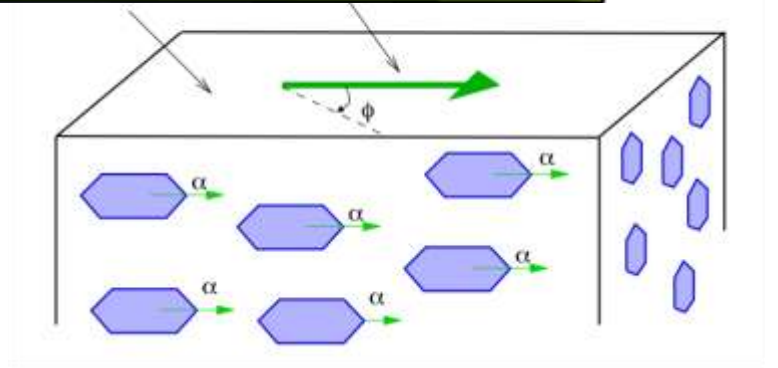
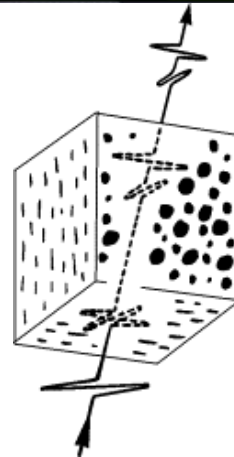
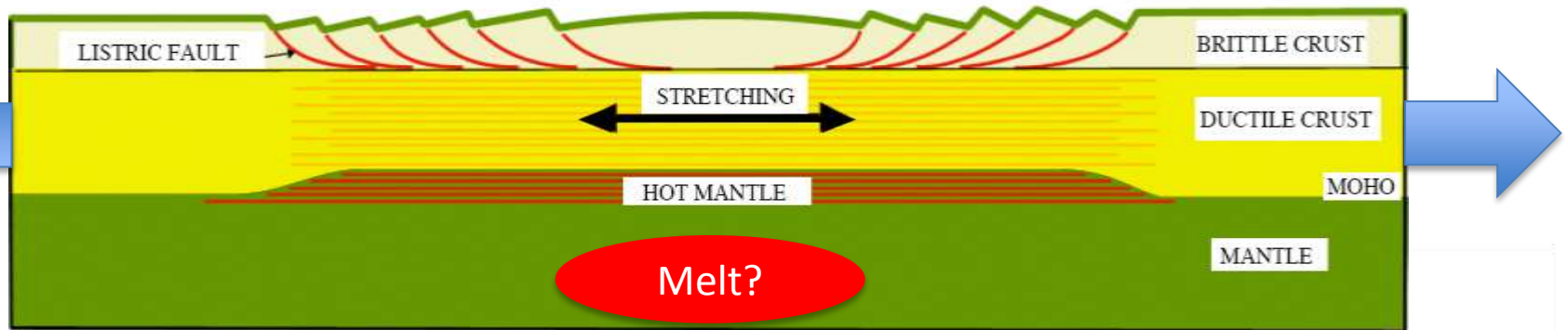
(a) Pure Shear



Seismic methods for imaging

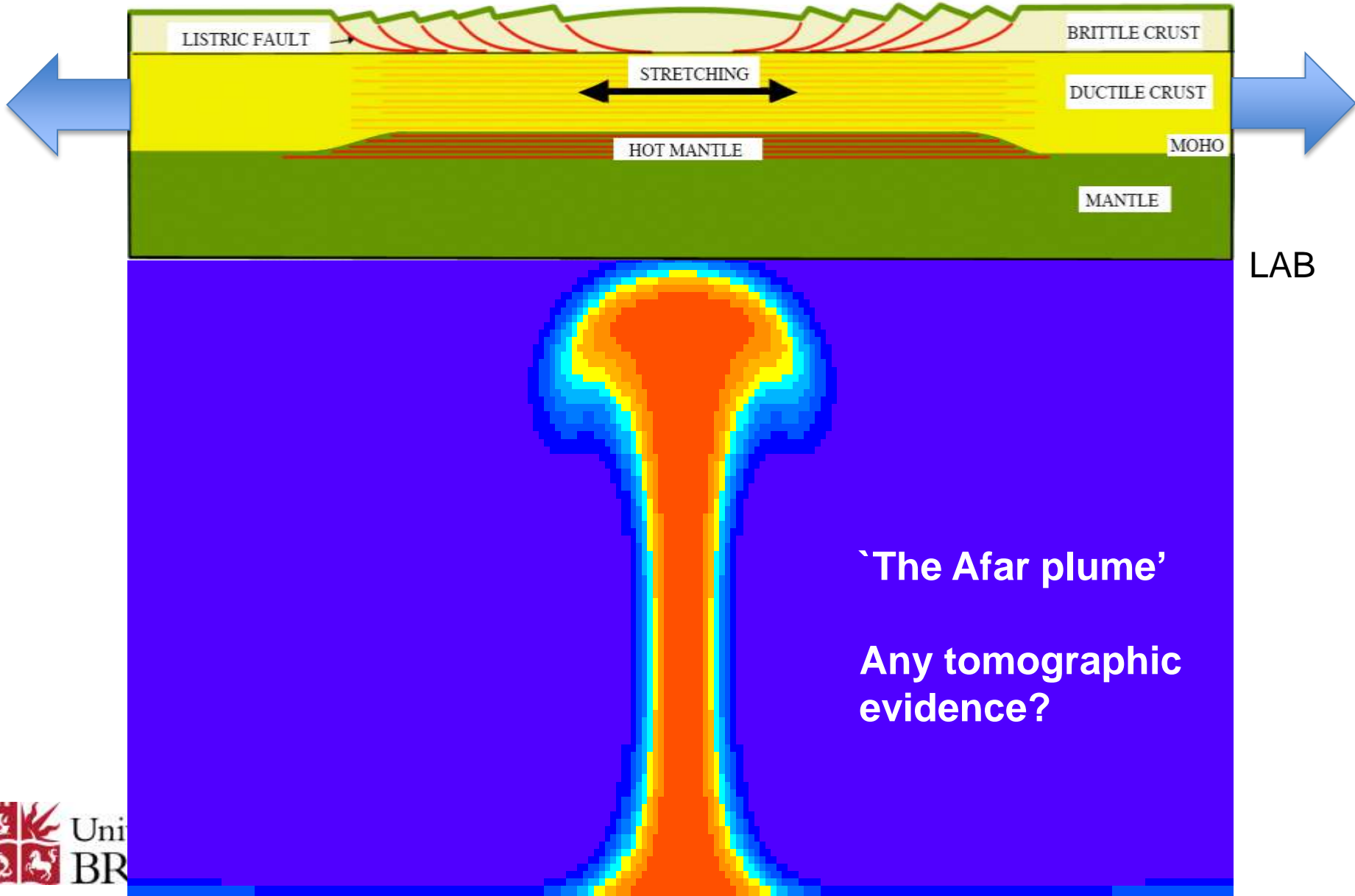
- Seismic anisotropy = mantle flow; aligned melt

(a) Pure Shear



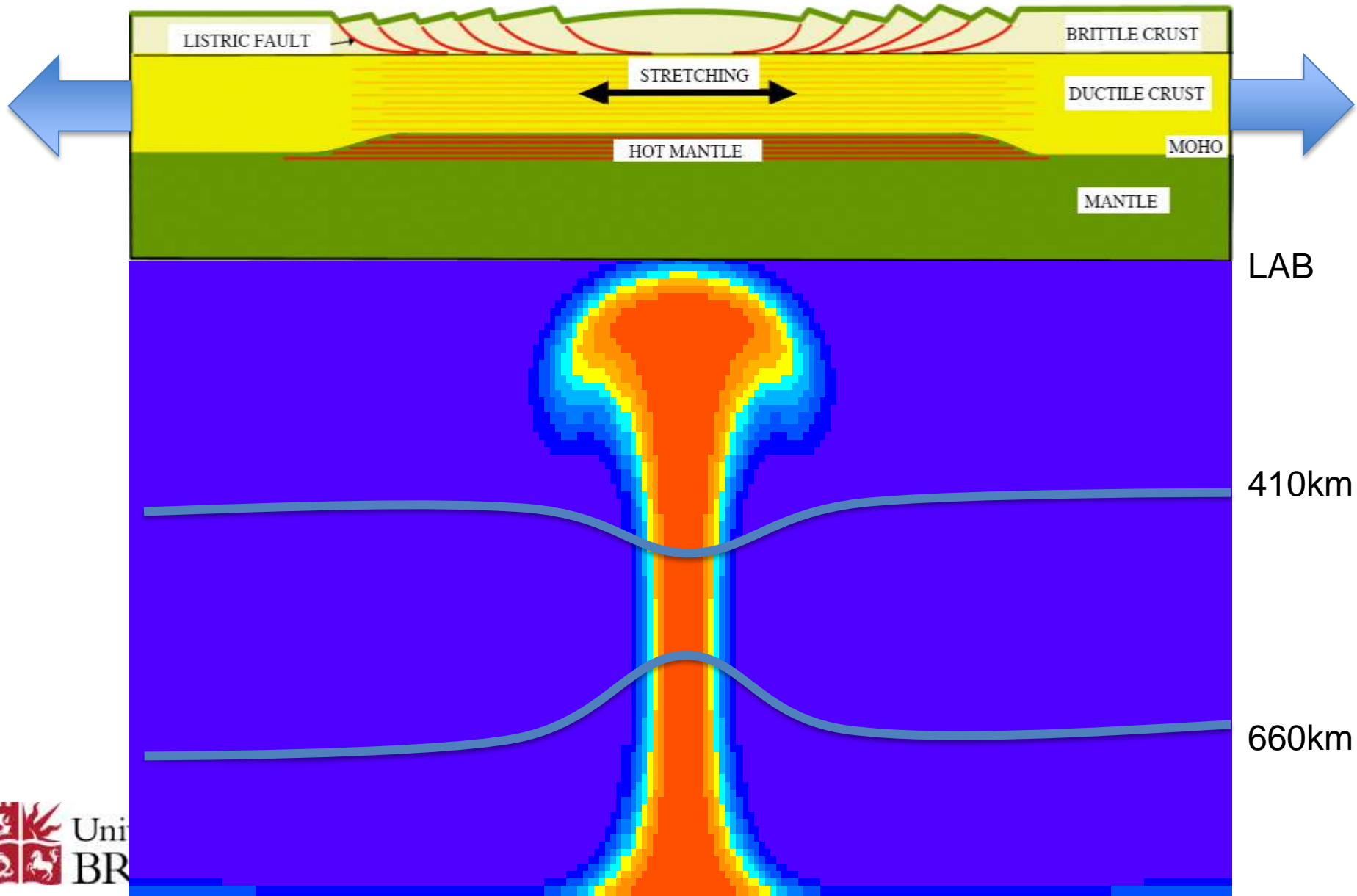
Mantle plume

(a) Pure Shear



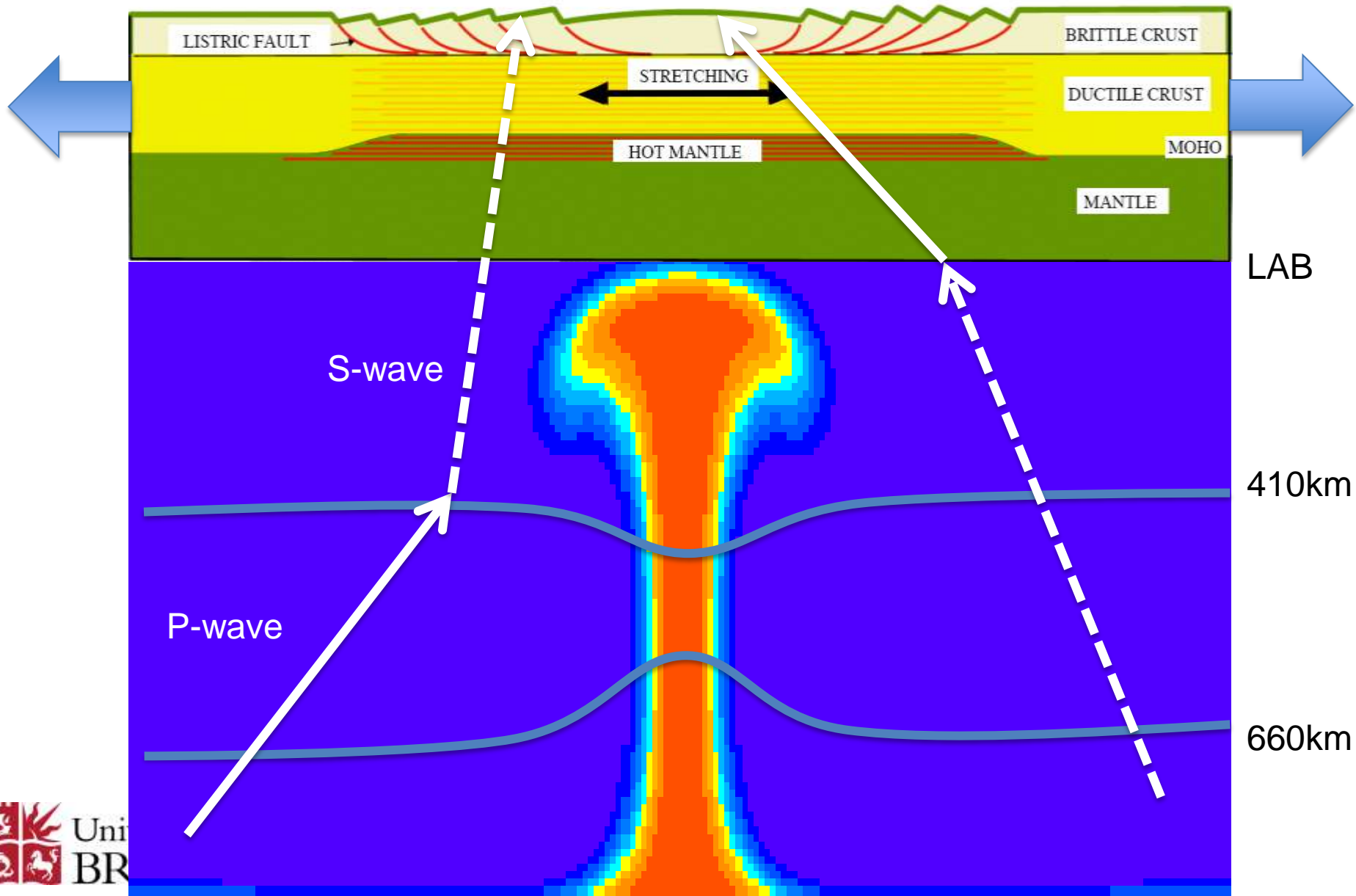
Mantle plume

(a) Pure Shear



Mantle plume

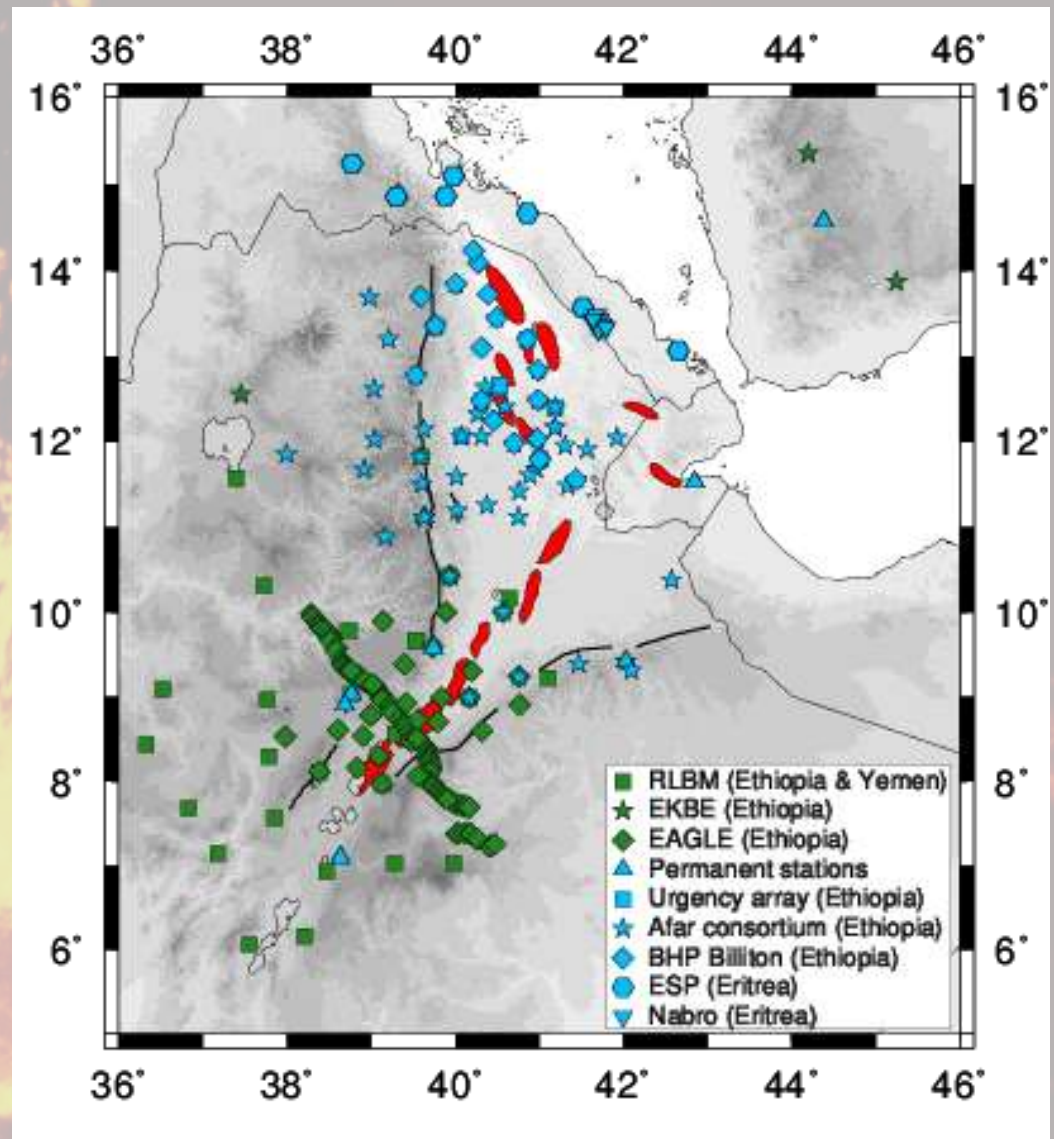
(a) Pure Shear



Fieldwork



- **RLBM (Horn of Africa)**
 - 5 stations: Jun 1999 - Dec 2002
- **EKBE (MER)**
 - 35 stations: Feb 2000 - Dec 2002
- **EAGLE (MER)**
 - 55 stations: Oct 2001 - Feb 2003
- **Permanent stations (IRIS, GEOSCOPE)**
 - 7 stations: Jul 1993 - Present
- **Urgency array (Afar)**
 - 9 stations: Oct 2005 - Ma, 2007
- **Afar consortium (NERC & NSF)**
 - 51 stations: Mar 2007 – Oct 2009
- **Danakil Seismic Project**
 - 12 stations: Oct 2009 – Feb 2013
- **Eritrea Seismic Project**
 - 6 stations: Jun 2011 – Oct 2012
- **Nabro Urgency Project**
 - 8 stations: Aug 2011 – Oct 2012



Seismic station on Nabro Volcano – Ethiopia/Eritrea Border



Seismic station – Biye Kabobe – Ethiopia/Somalia Border



Dust storm in the Danakil Depression

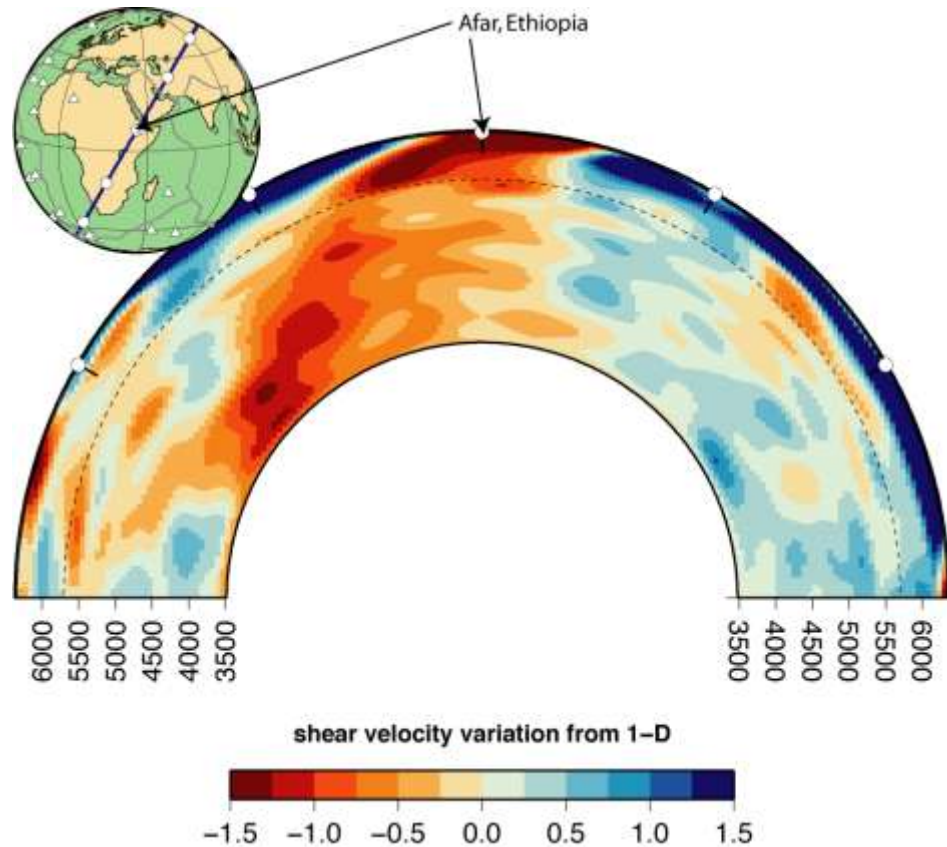


Results

- Tomography
- Receiver Functions
- Seismic anisotropy

Architecture of a superplume

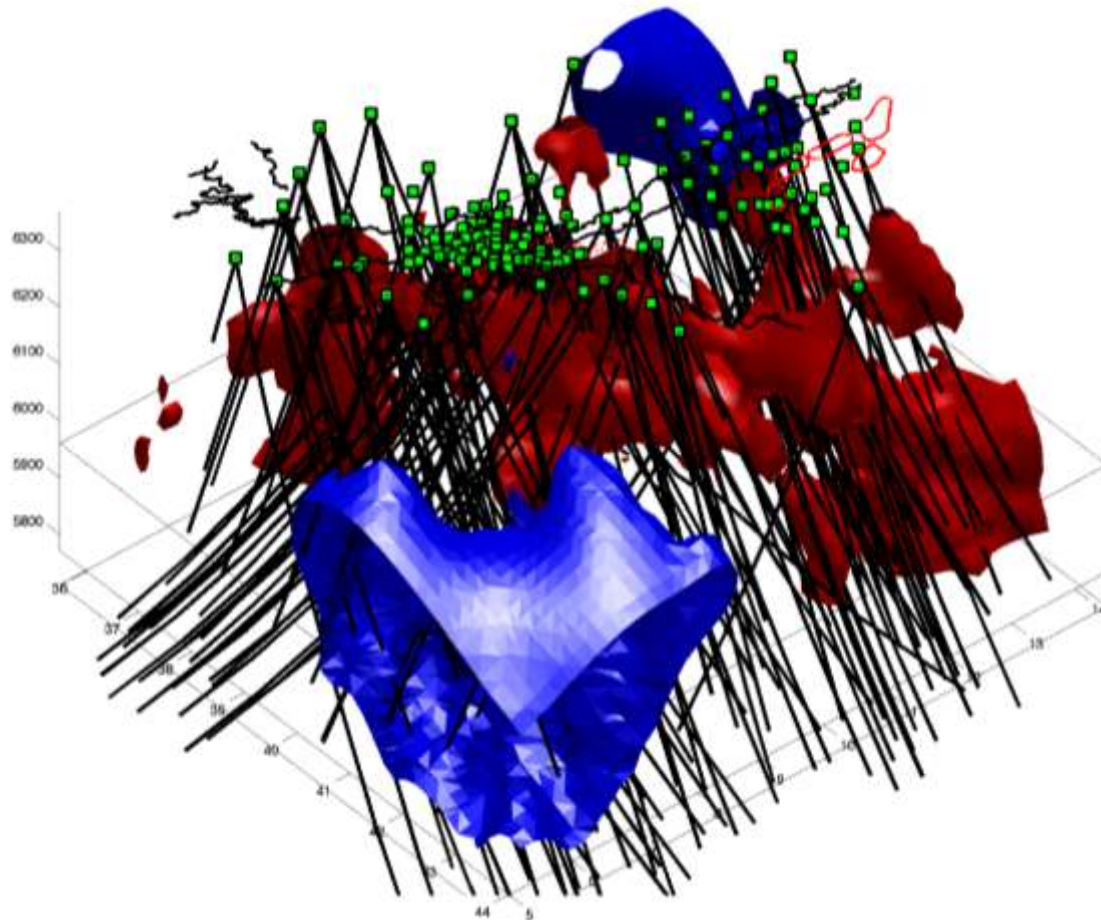
Chemical versus thermal?



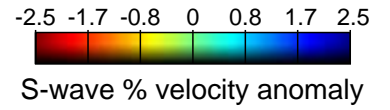
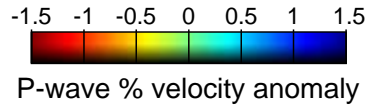
P- and S-wave travel-time tomography

Main Ethiopian Rift

- Bastow et al. (GJI, 2005; 2008); Hammond et al. (2013)
- Broad low velocity sheet-like anomaly that cuts through the pan-African fabric; not a conventional plume-like upwelling.

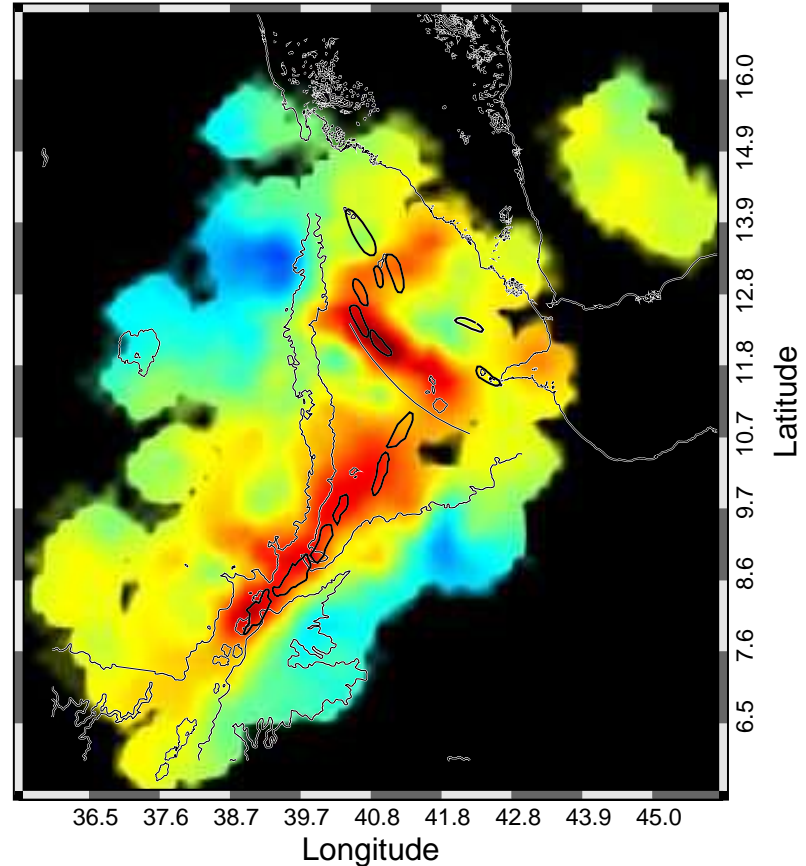
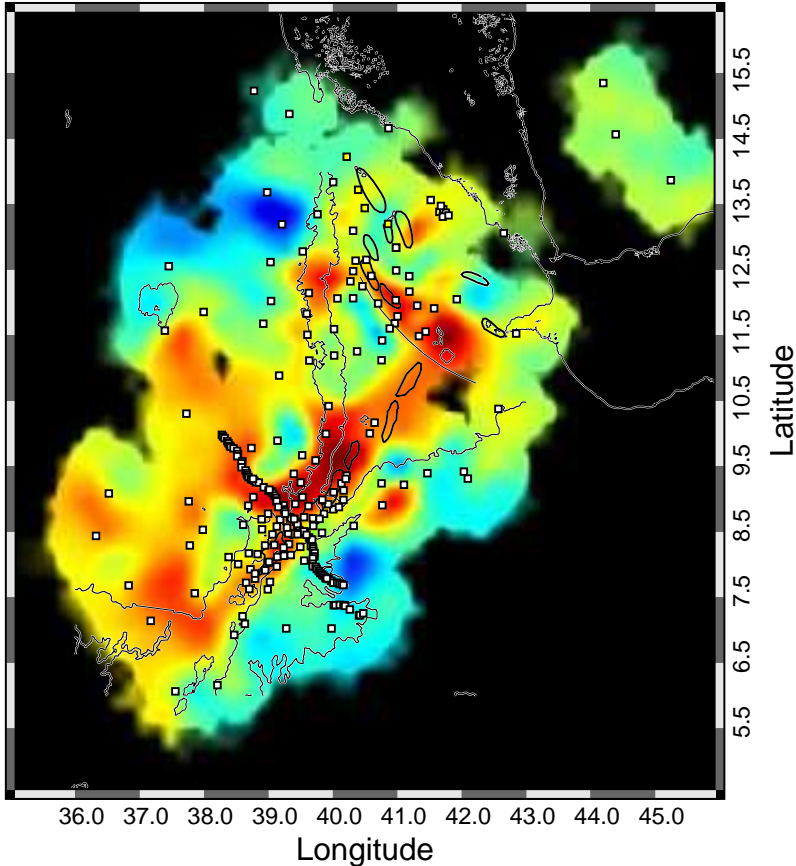


50 – 150 km



depth =
75 km

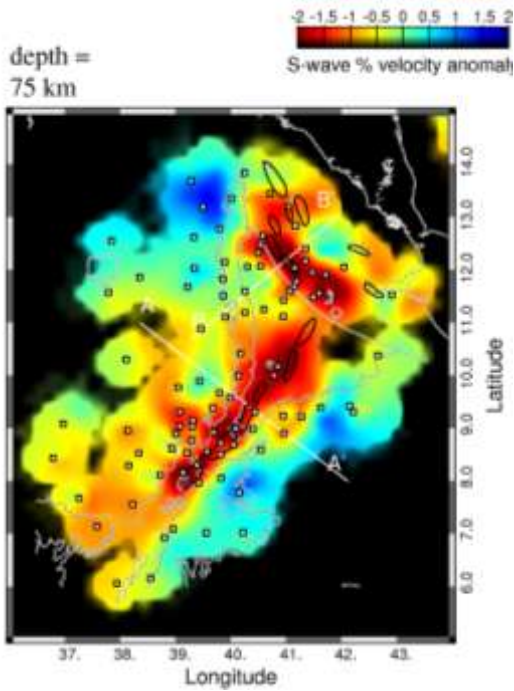
depth =
75 km



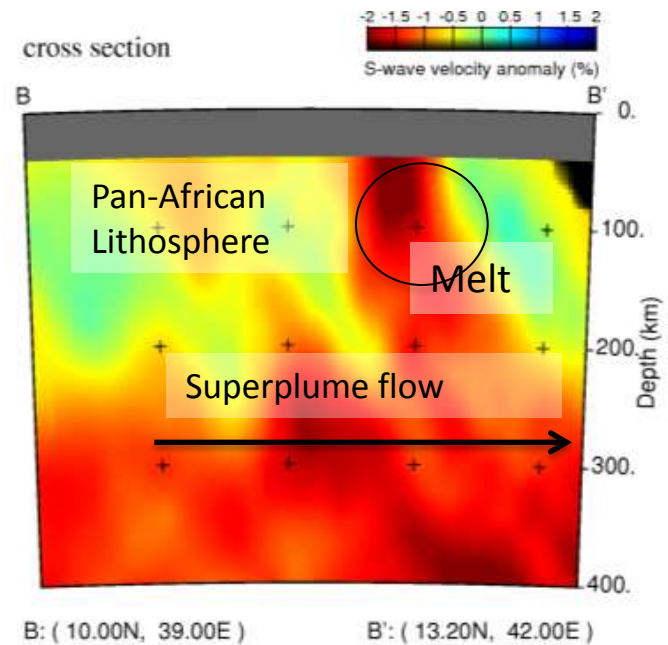
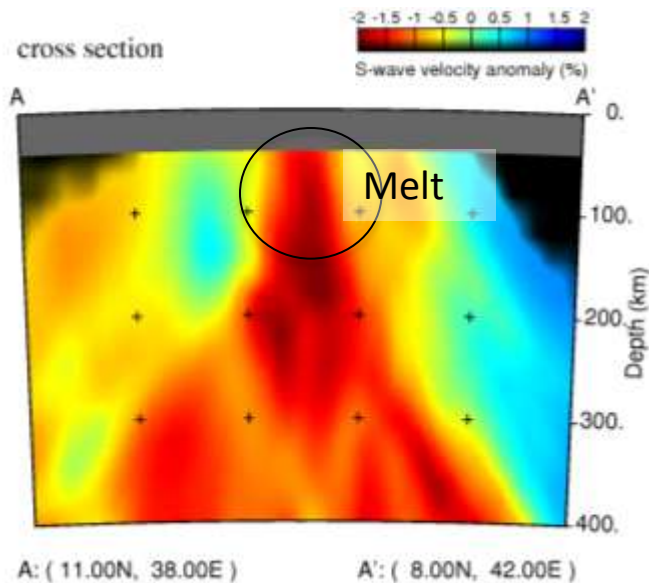
P-wave model – 10132 arrivals

S-wave model – 10811 arrivals

Seismic tomography



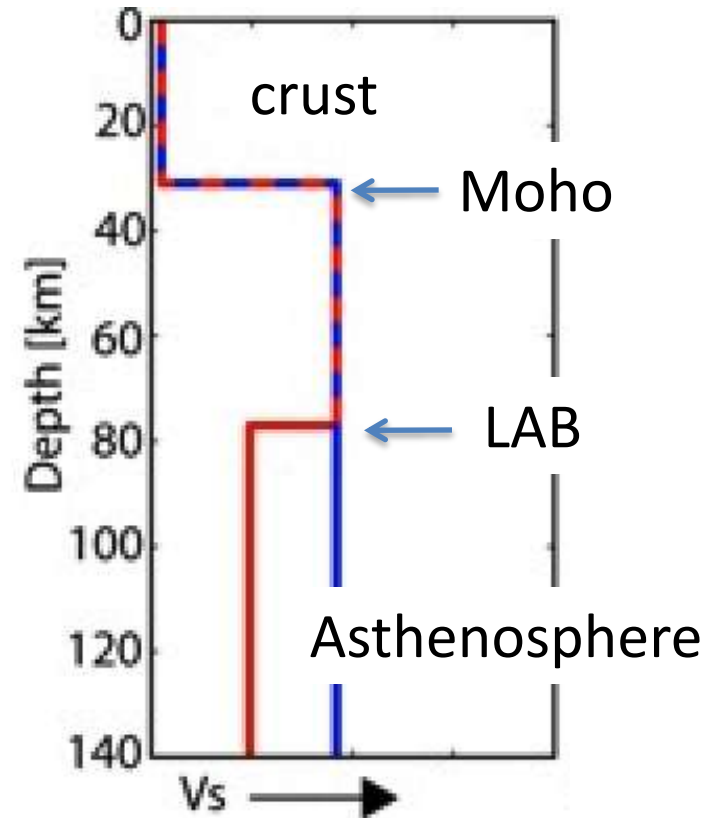
- Focused low seismic velocity anomalies in top 150km – Melt
- Broad tabular low seismic velocity anomaly to depths of at least 400 km
- Seismically fast Pan-African lithosphere
- Seismic velocities are best explained by high temperatures and melt in elongate inclusions. **Huge absolute delay time and R values (V_p/V_s) -> melt**
- Latest absolute delay times of anywhere in the world.



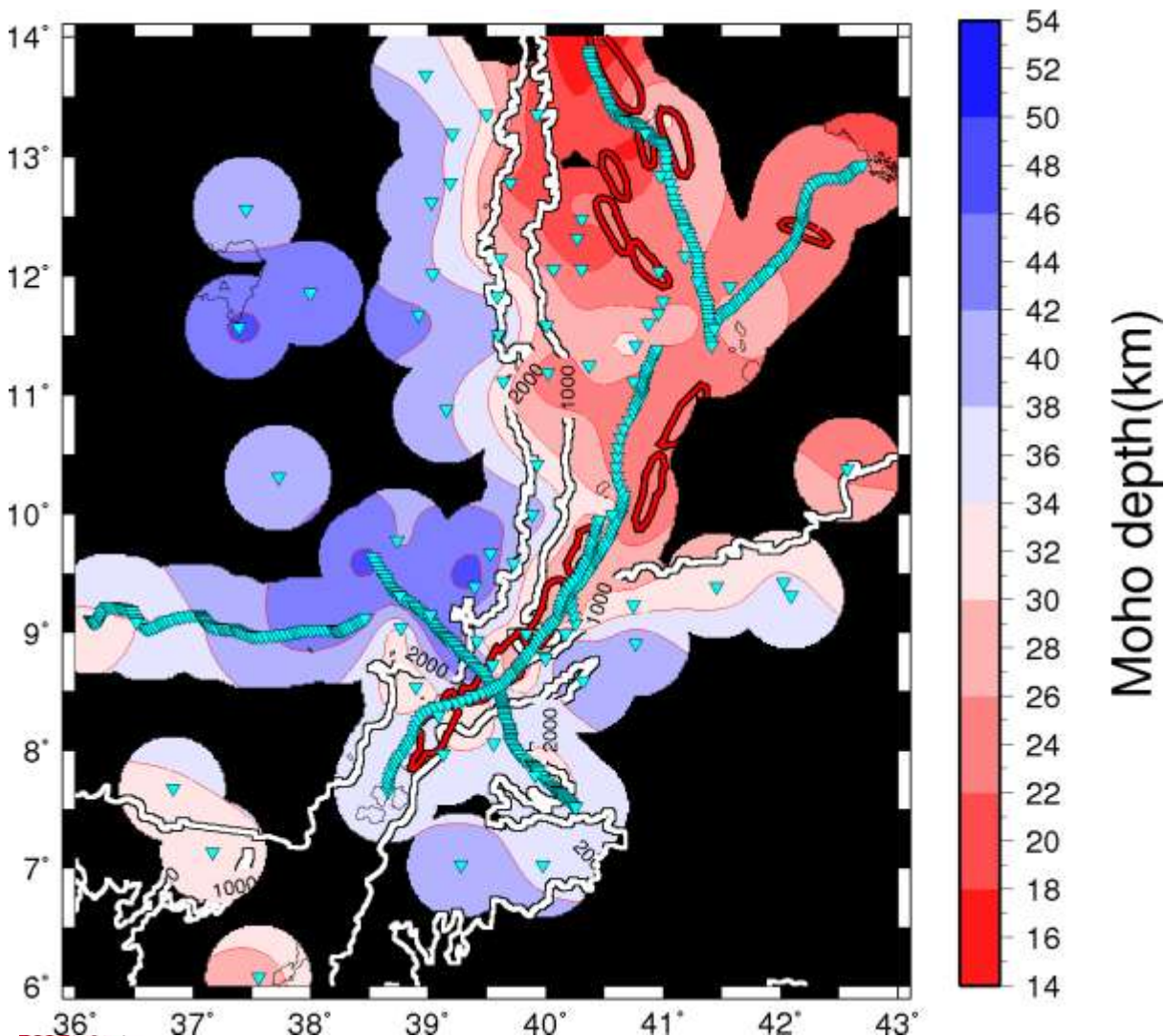
The Stratigraphy of the Lithosphere

Use receiver functions

- Image variations in crust:
 - Moho depth
 - V_p/V_s ratios
- Image variations in the lithosphere (tectonic plate)
 - lithosphere-asthenosphere boundary (LAB)



Receiver Functions: Crustal thickness



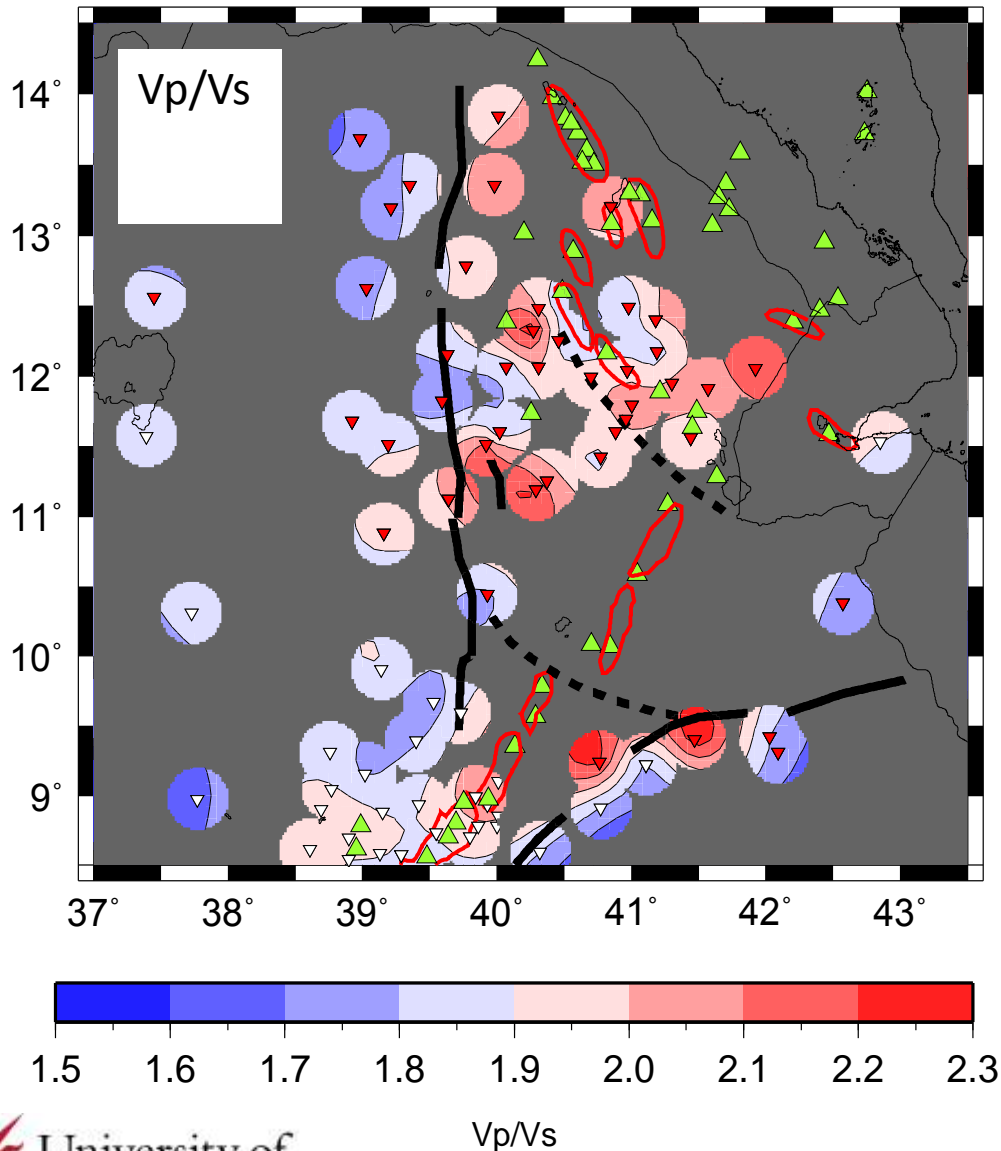
- Dramatic variations across Ethiopia: 10-50km

- Thinnest crust in northern Afar

- Thickest beneath northern plateau

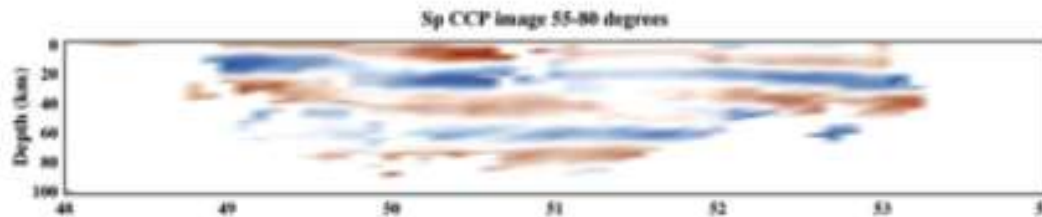
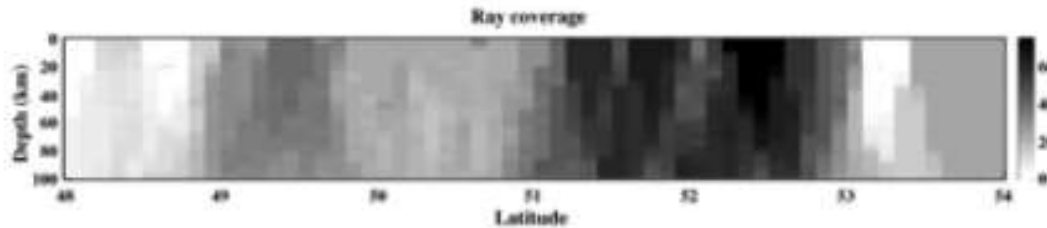
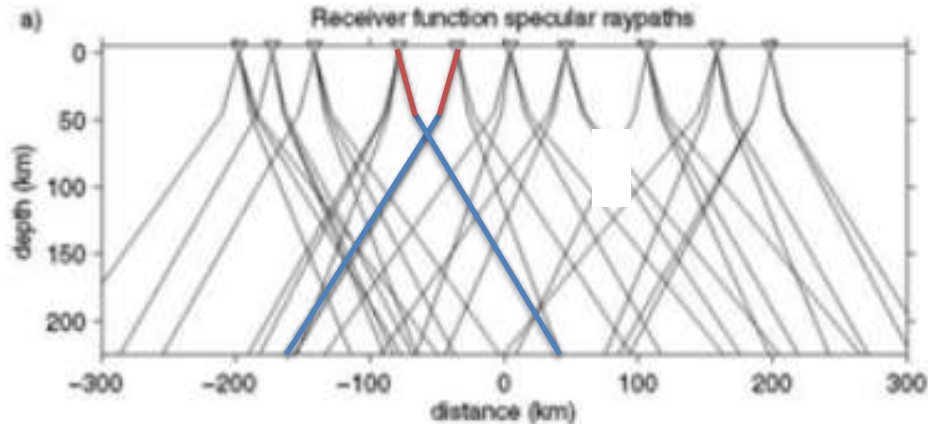
- Sharp variations at rift flanks

Receiver Functions: V_p/V_s ratios



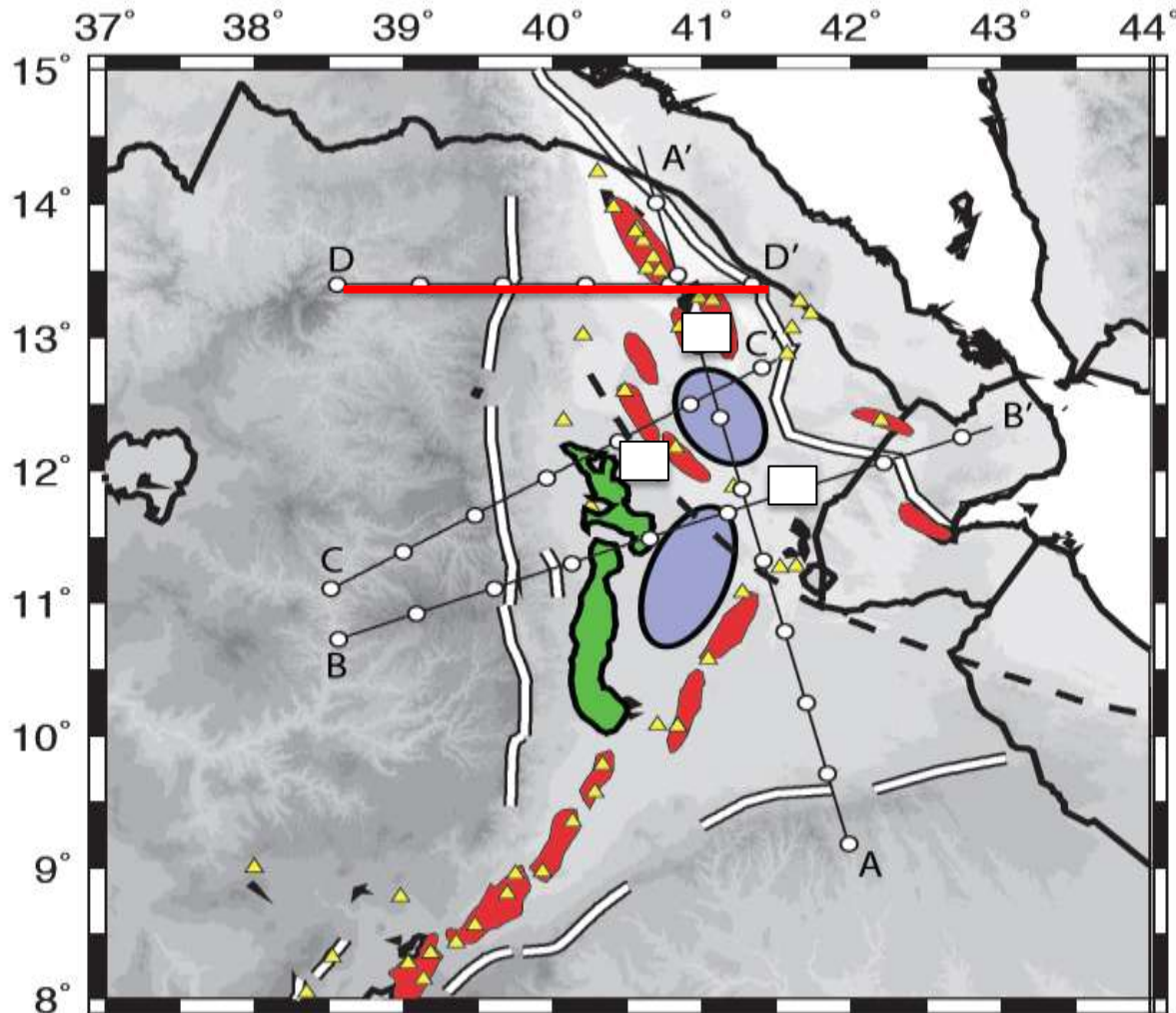
- Vary from 1.6 beneath the plateaus
- Up to 2.3 in parts of Afar
- $V_p/V_s > 2.0$ means melt
- Sharp variations at rift flanks

Crustal Structure: CCP migration

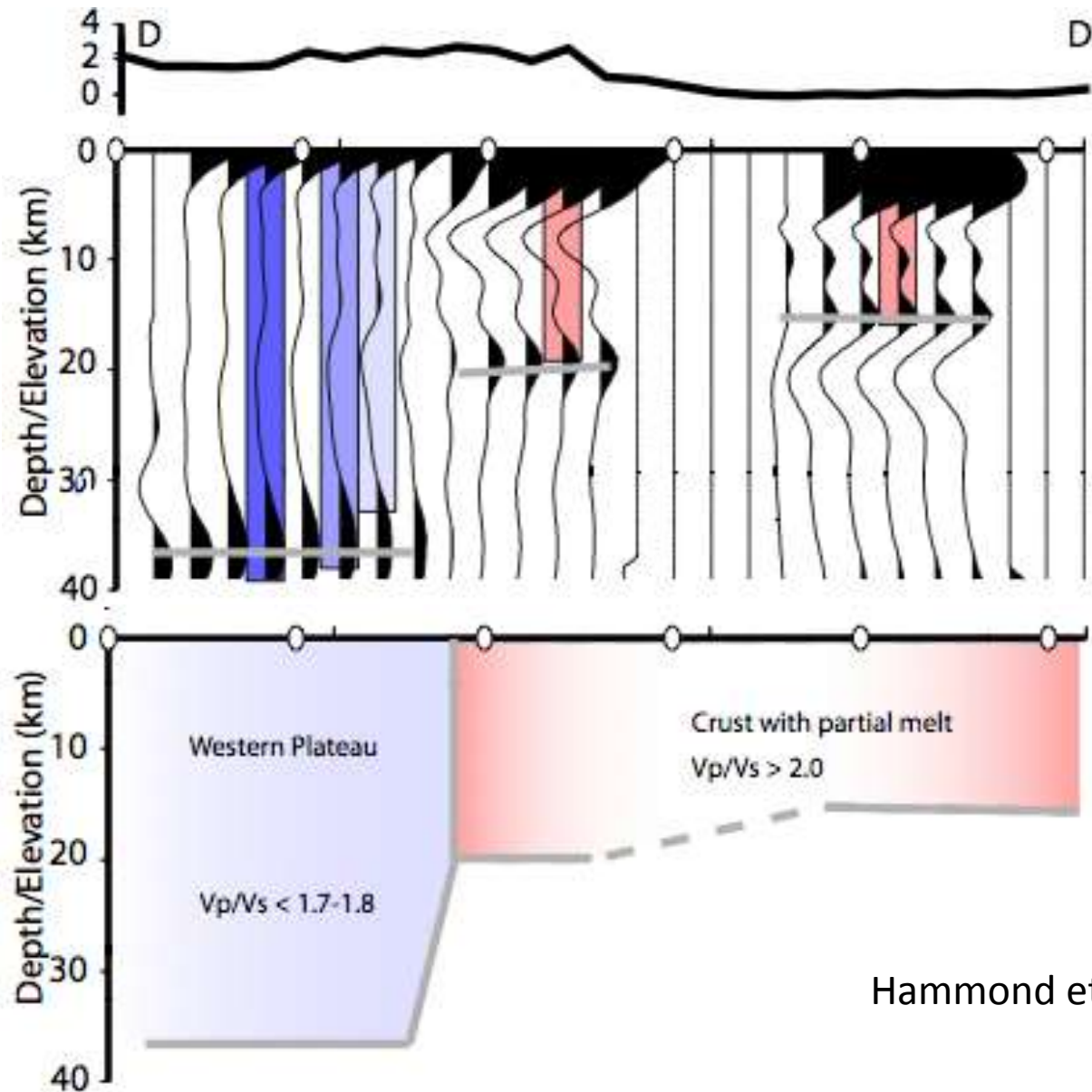


Angus et al., 2010

Crustal Structure: CCP migration



Crustal Structure: CCP migration

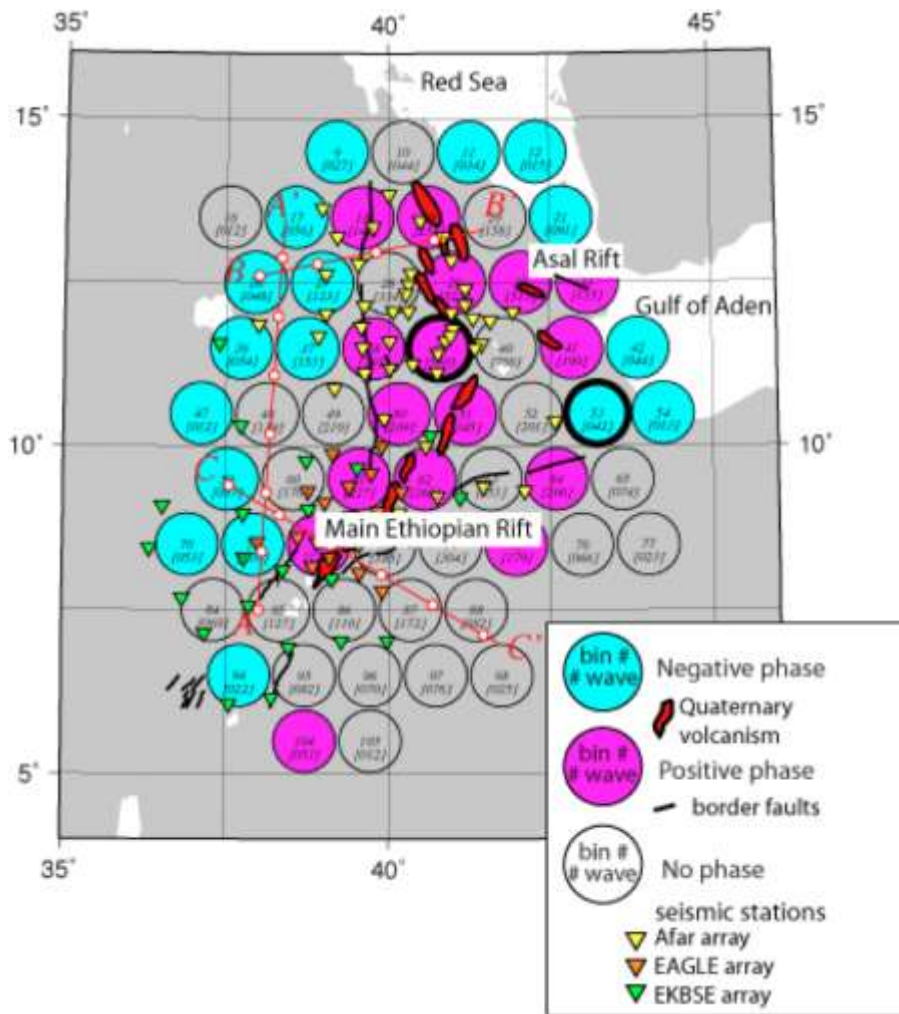


Hammond et al. 2011

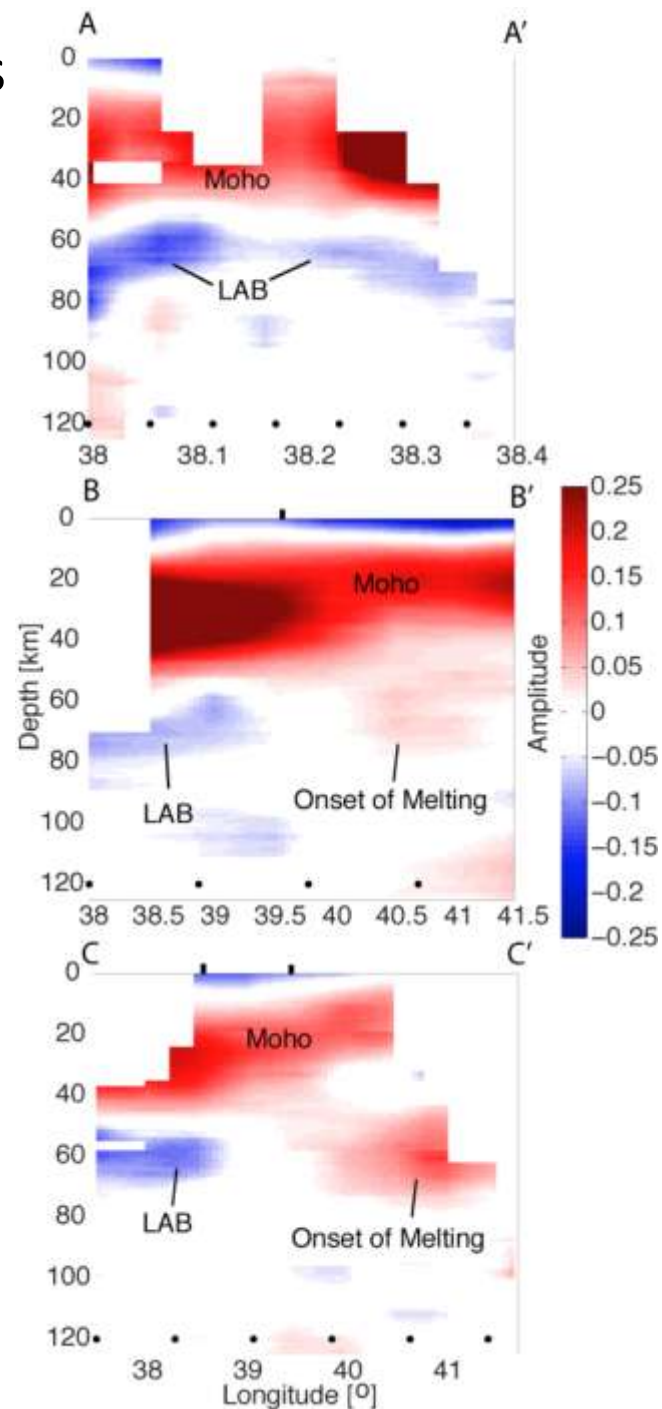
The Lithosphere-Asthenosphere Boundary

- S-P conversion
- Common Conversion Point migration
- Clear differences between plateau and rift

The LAB beneath Ethiopia: migrated images

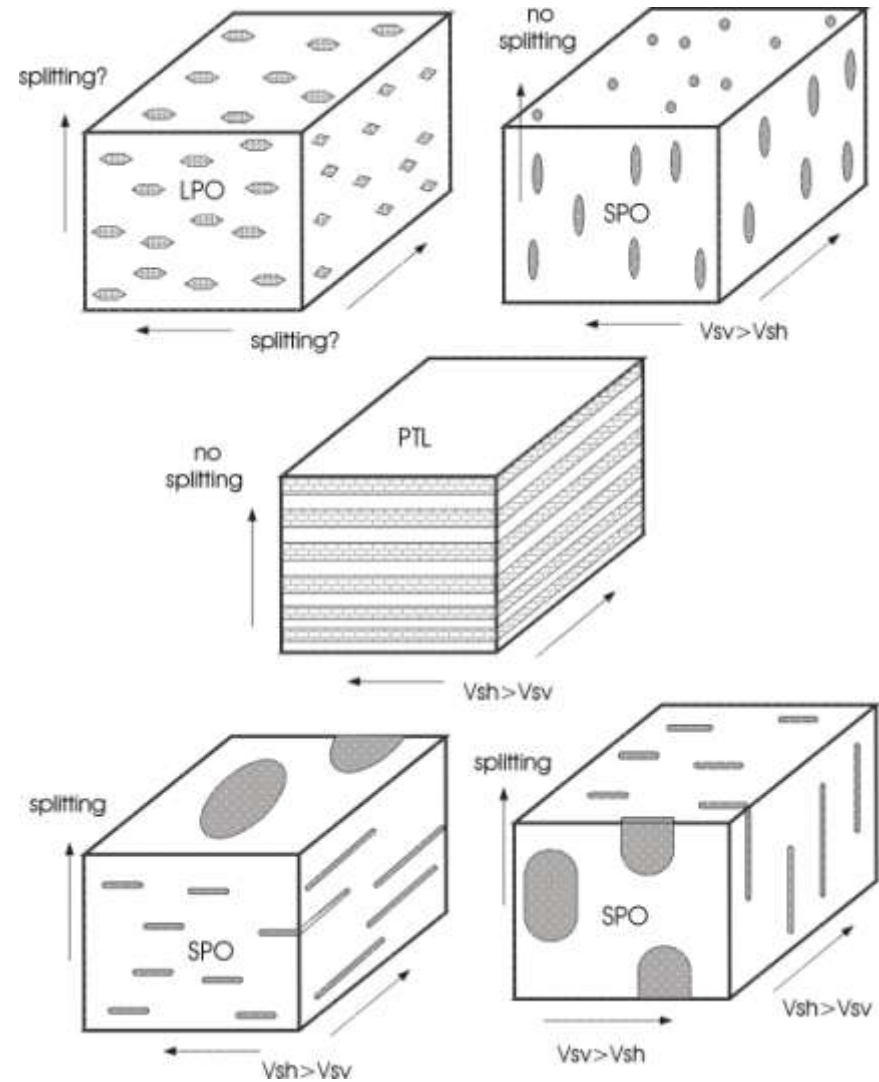
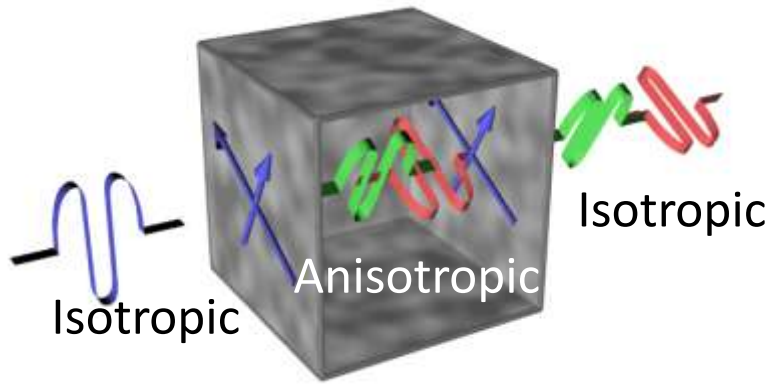


- At ~75km S-wave receiver functions show a velocity decrease beneath the Plateau and a velocity increase beneath most of Afar



(Rychert et al., 2012)

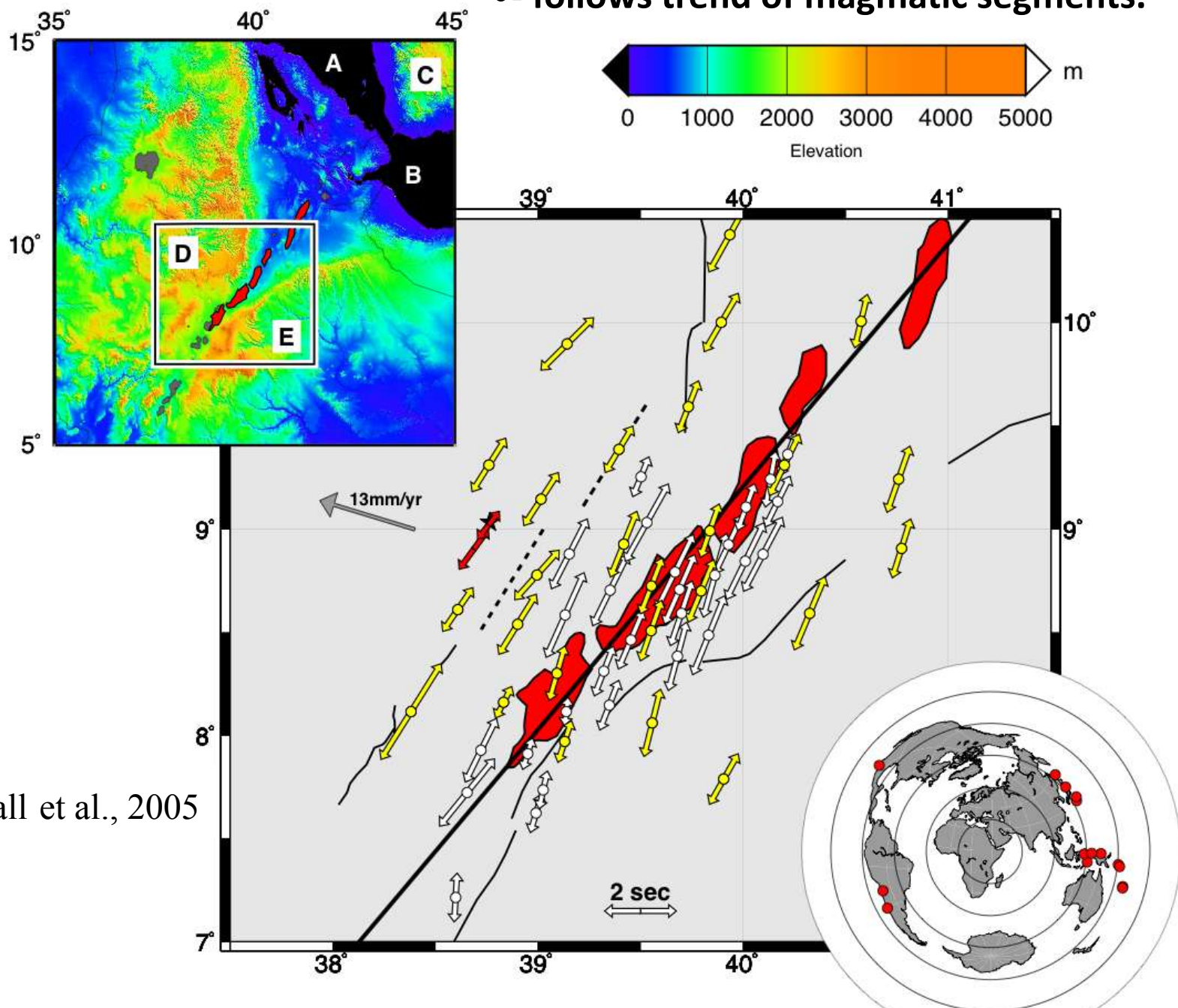
Seismic anisotropy: *Shear-wave splitting analysis.*



Unravelling mechanisms for anisotropy

• Counter-clockwise rotation within rift valley

• - follows trend of magmatic segments.

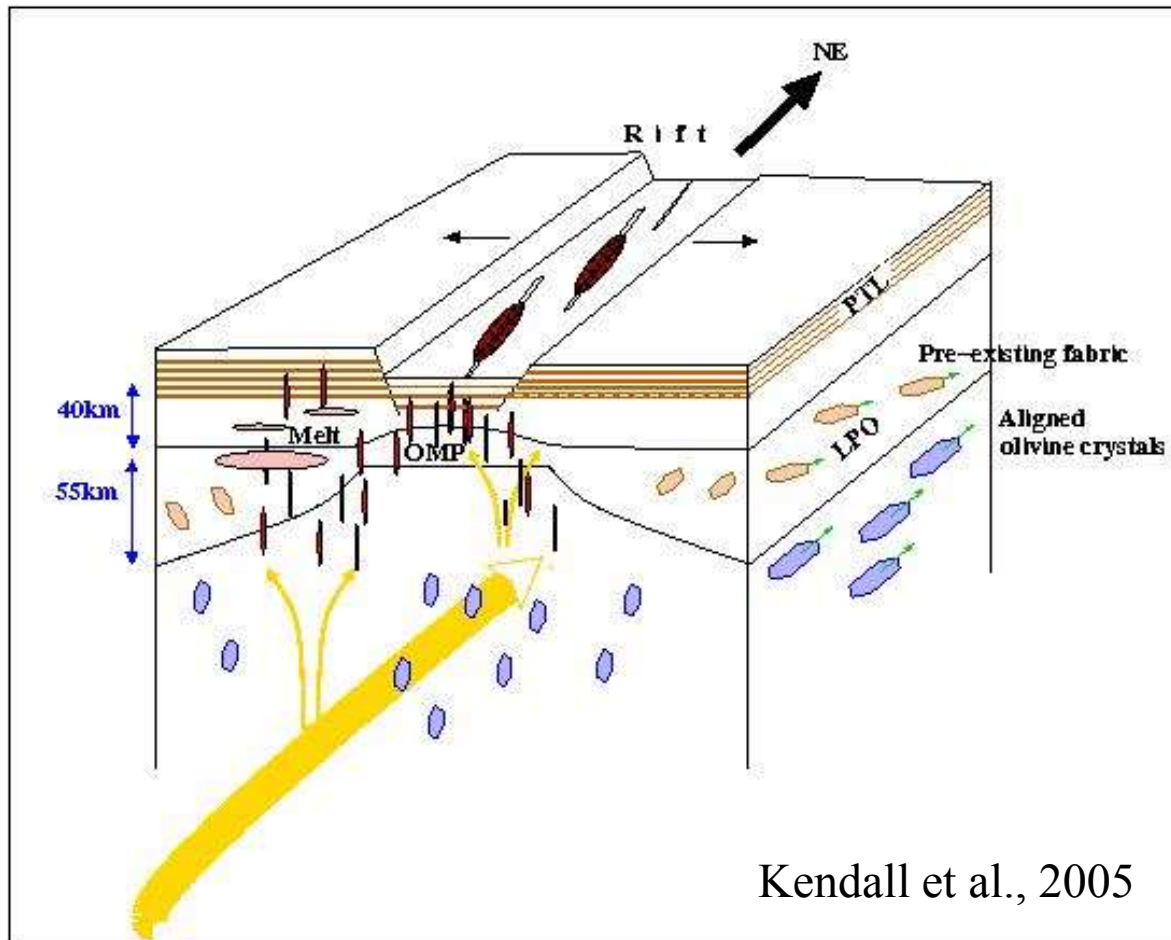


Kendall et al., 2005



Mechanisms for anisotropy:

- Working model for anisotropy beneath East Africa Rift

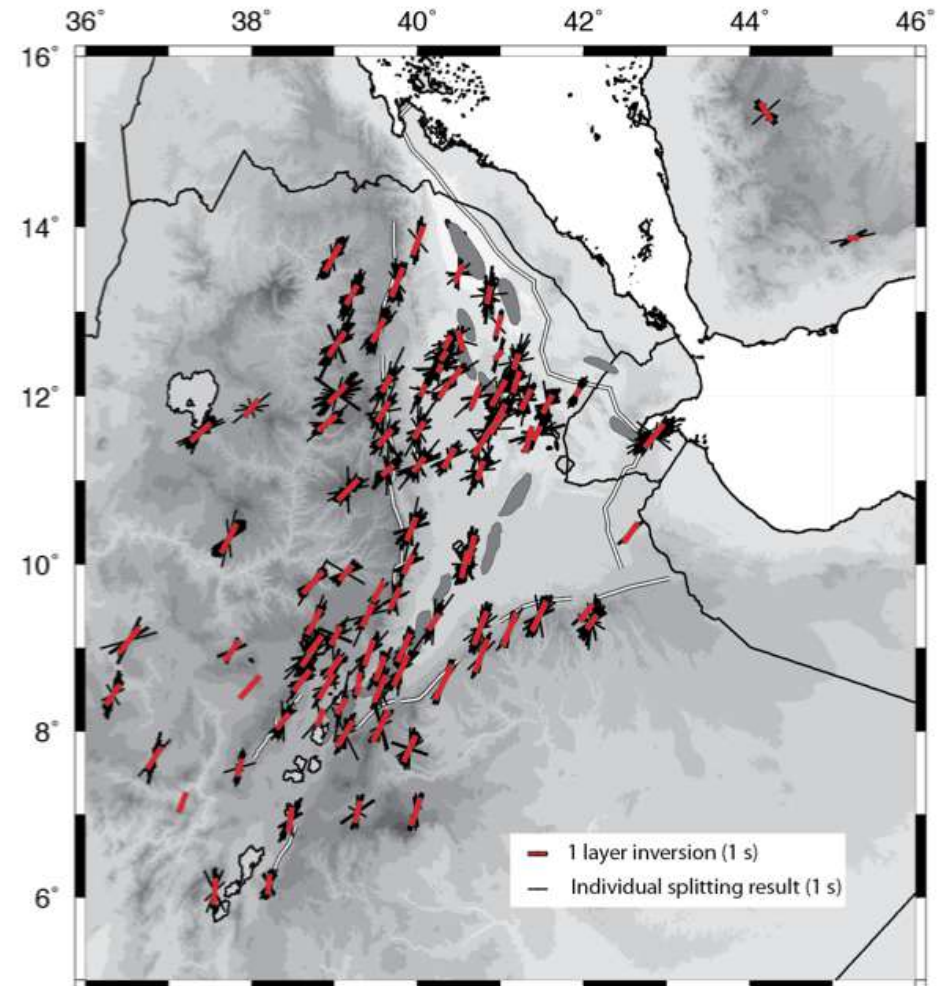
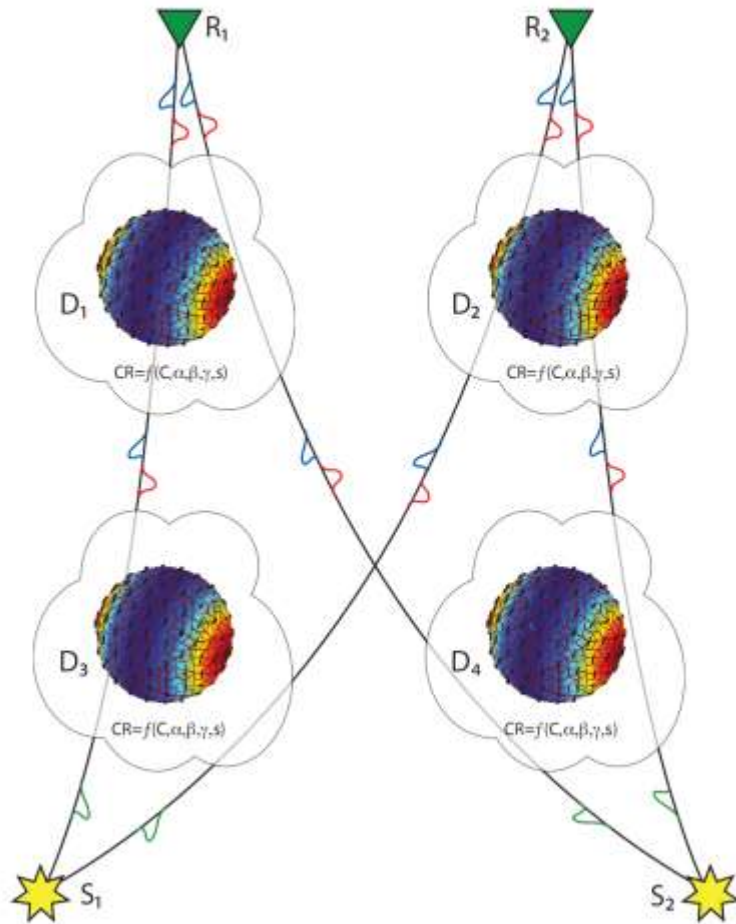


- Large-scale flow beneath eastern Africa associated with super-swell.
- Melt focused at plate boundaries - leads to oriented vertical pockets of melt.
- Contribution from pan-African fabric in lithosphere away from Main Ethiopian rift and continental slivers within Afar.

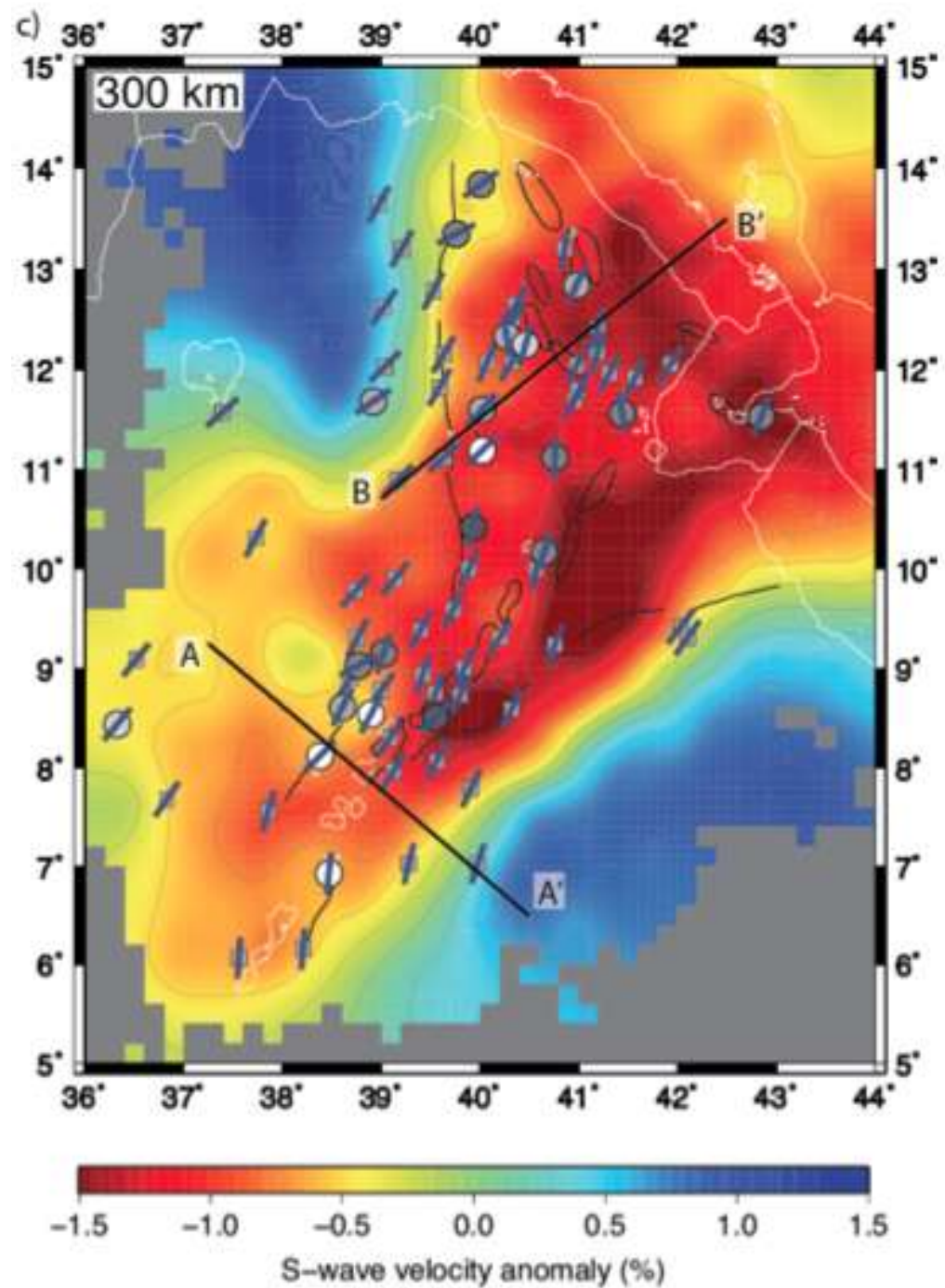
- See also Obrebski et al., 2010; Gao et al., 2010; Bastow et al., 2010

Shear-wave splitting tomography

- Based on Wookey (2012)
- Two layer inversion across multiple tectonic domains

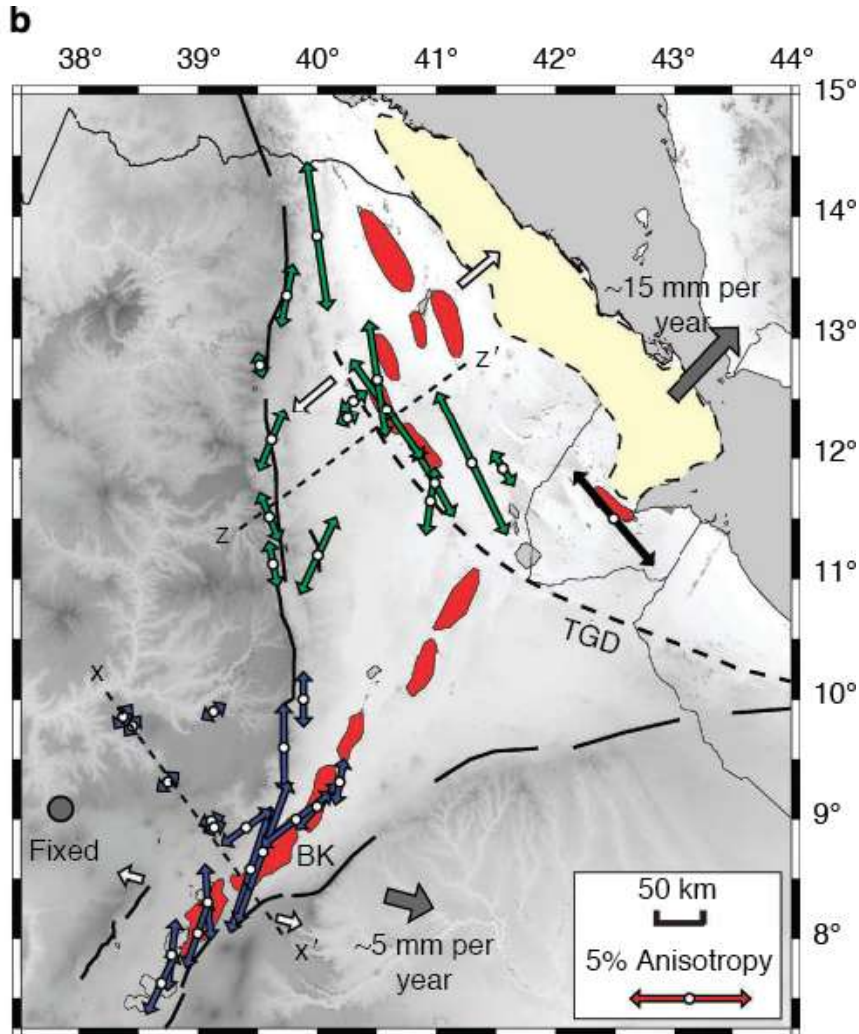


Lower layer aligned with
density driven mantle flow

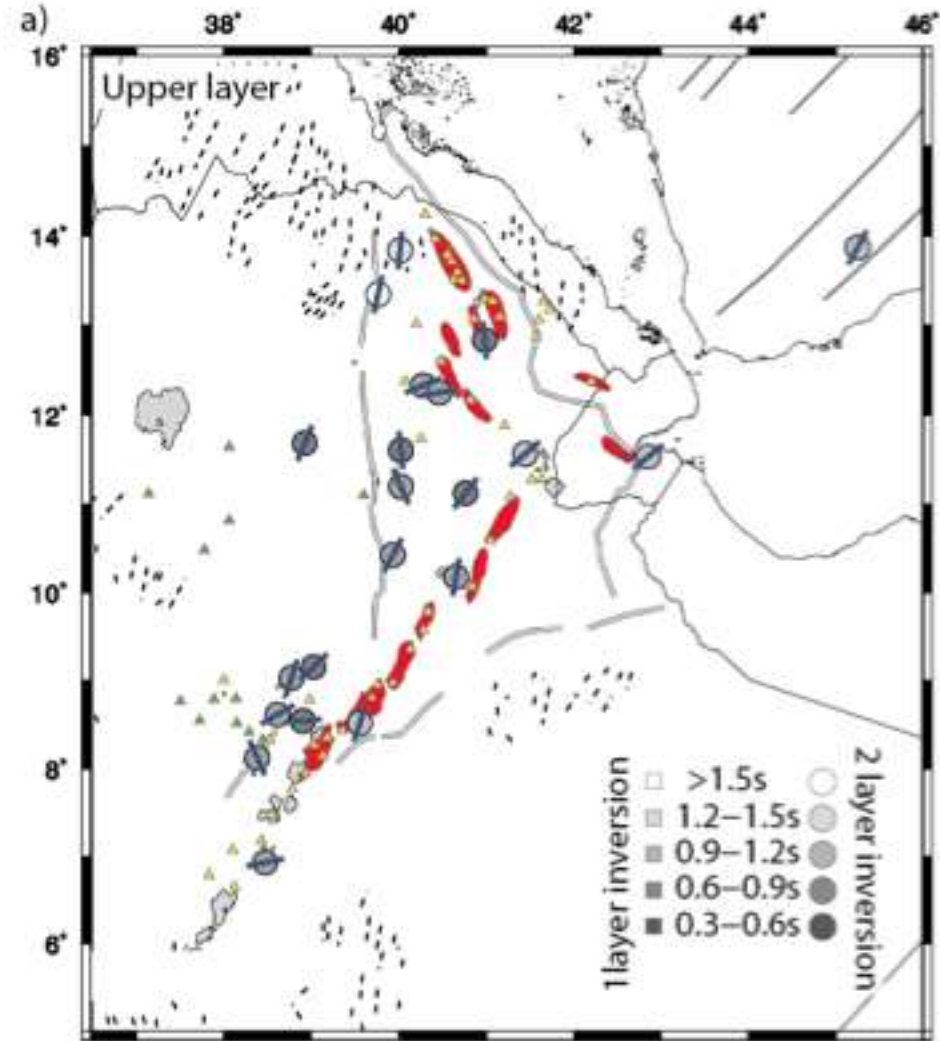


Hammond et al., 2013

Upper layer – fossil fabric and melt-induced anisotropy

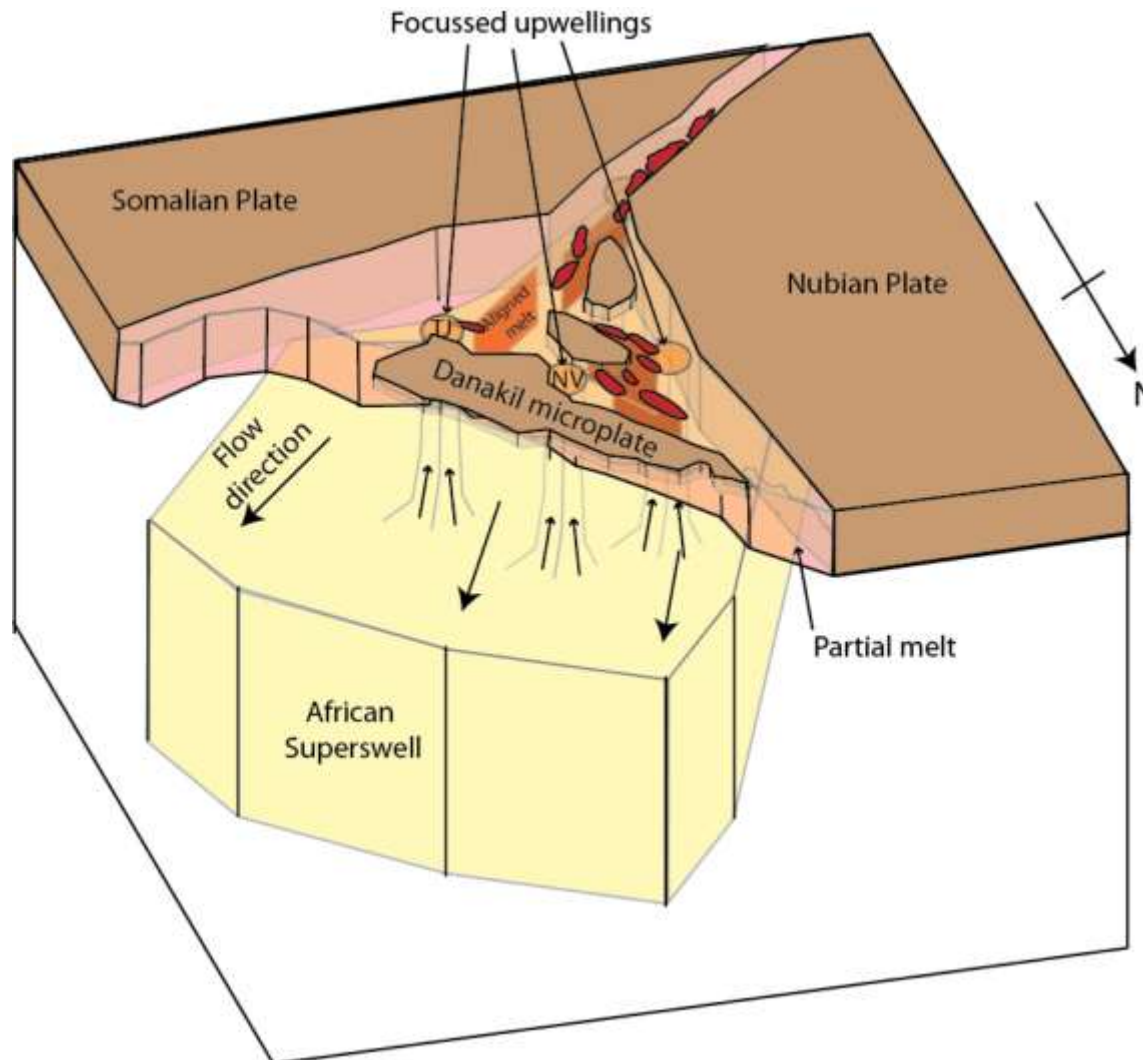


Crustal anisotropy; Keir et al., 2011



Upper layer; Hammond et al., 2013

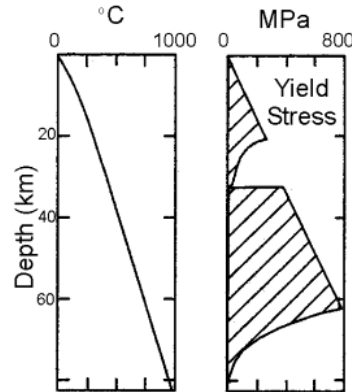
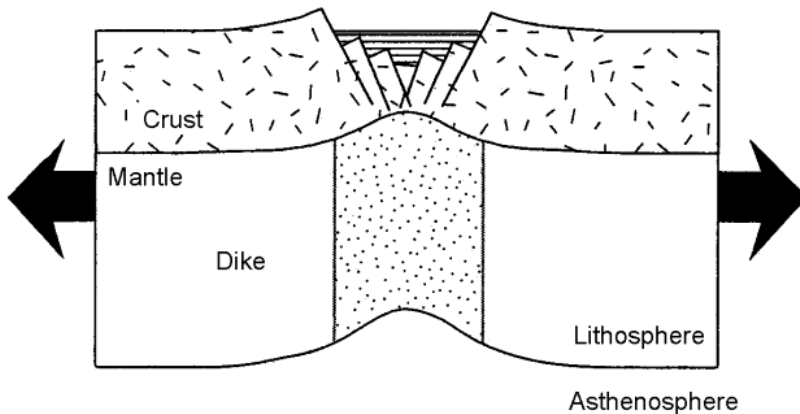
Stratified upper-mantle beneath the EAR



- Broad low-velocity region of density driven flow
- Focuses upwellings punctuating lithosphere and crust
- Microplate isolation in Afar

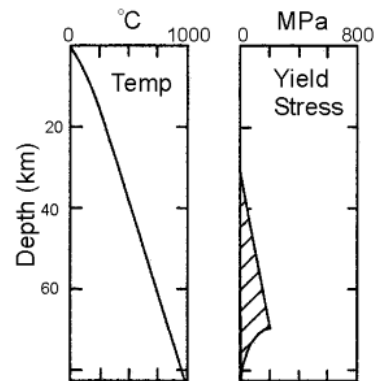
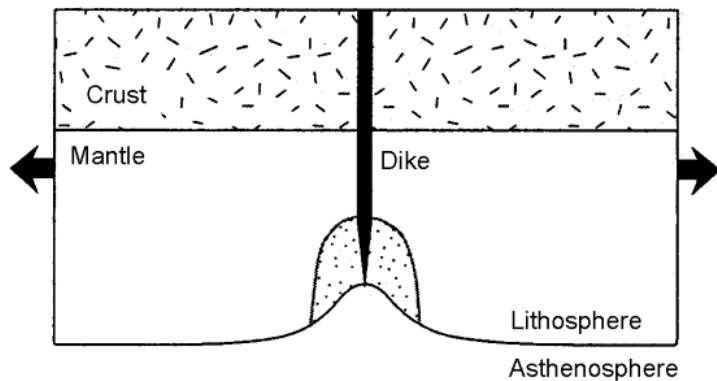
Magma Assisted Rifting

Tectonic Stretching

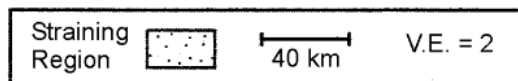


- Melt injection leads to much lower yield stresses.

Magmatic Extension



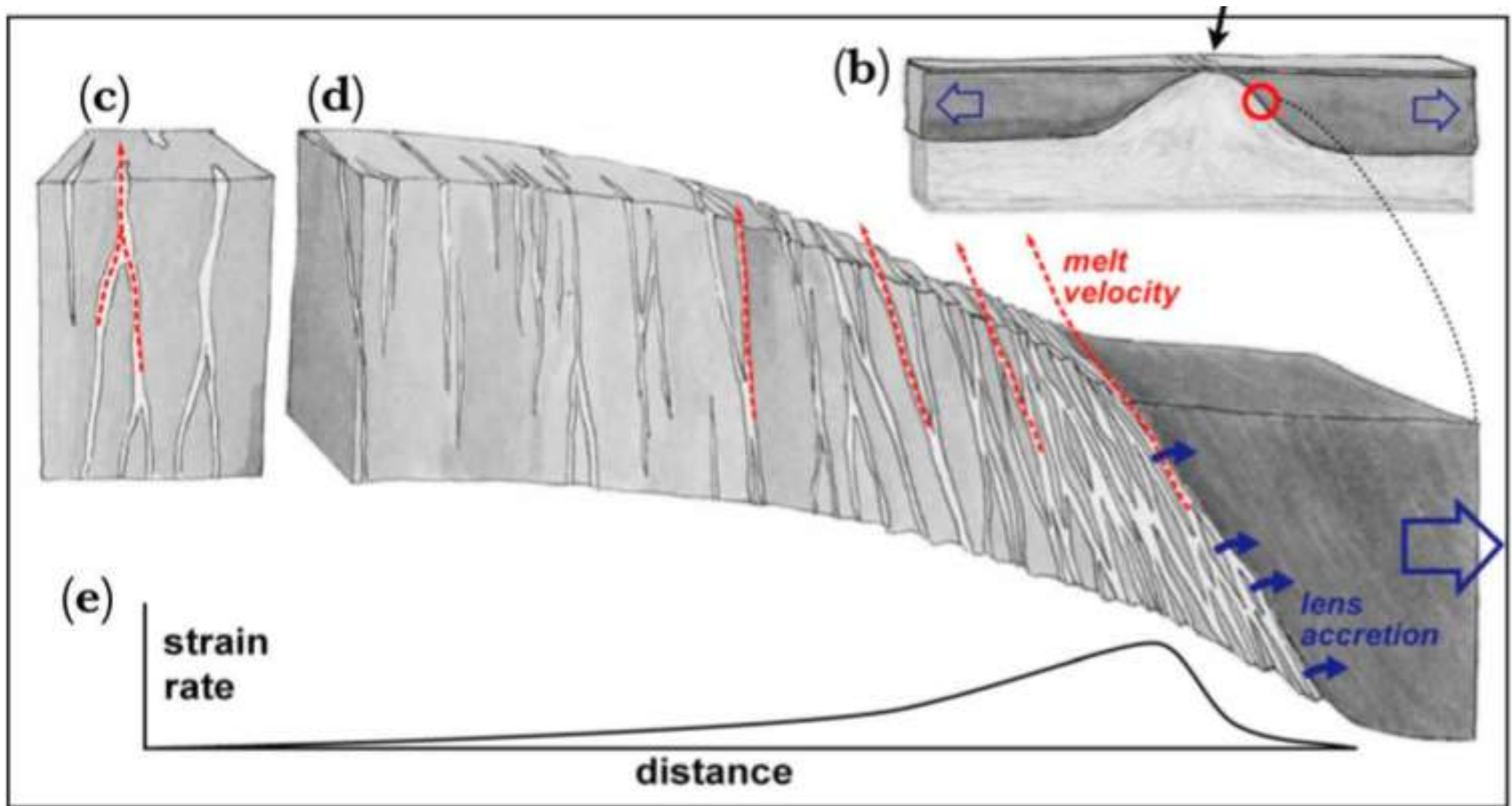
- Deformation (strain) is focused at plate boundaries.



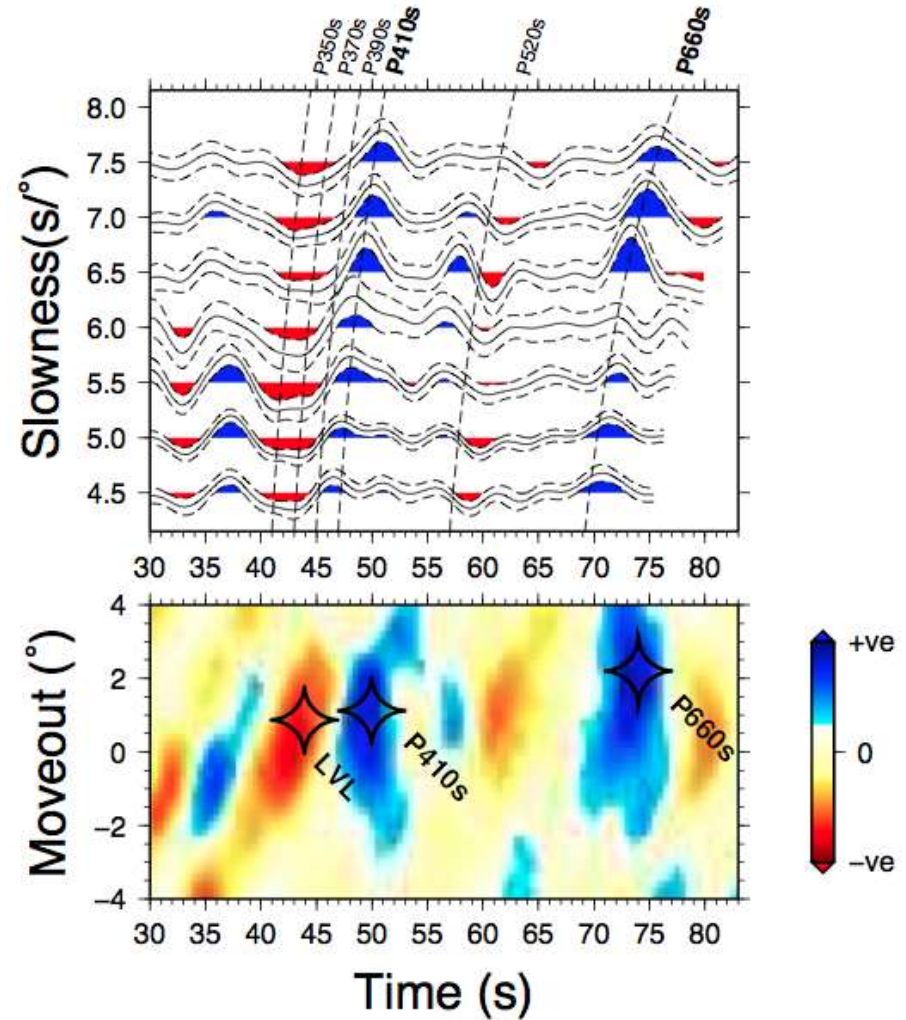
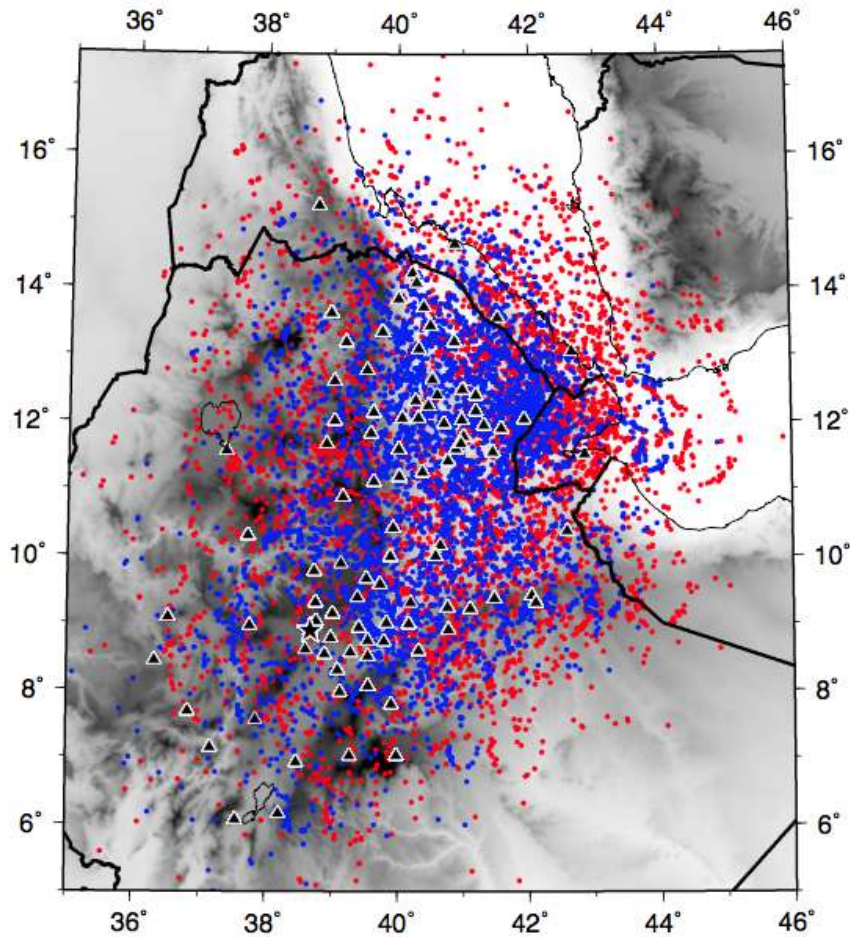
(After Buck, 2004)

Stress driven melt segregation

- most effective at flanks (marginal LAB)



Evidence for plume in the mantle transition zone?



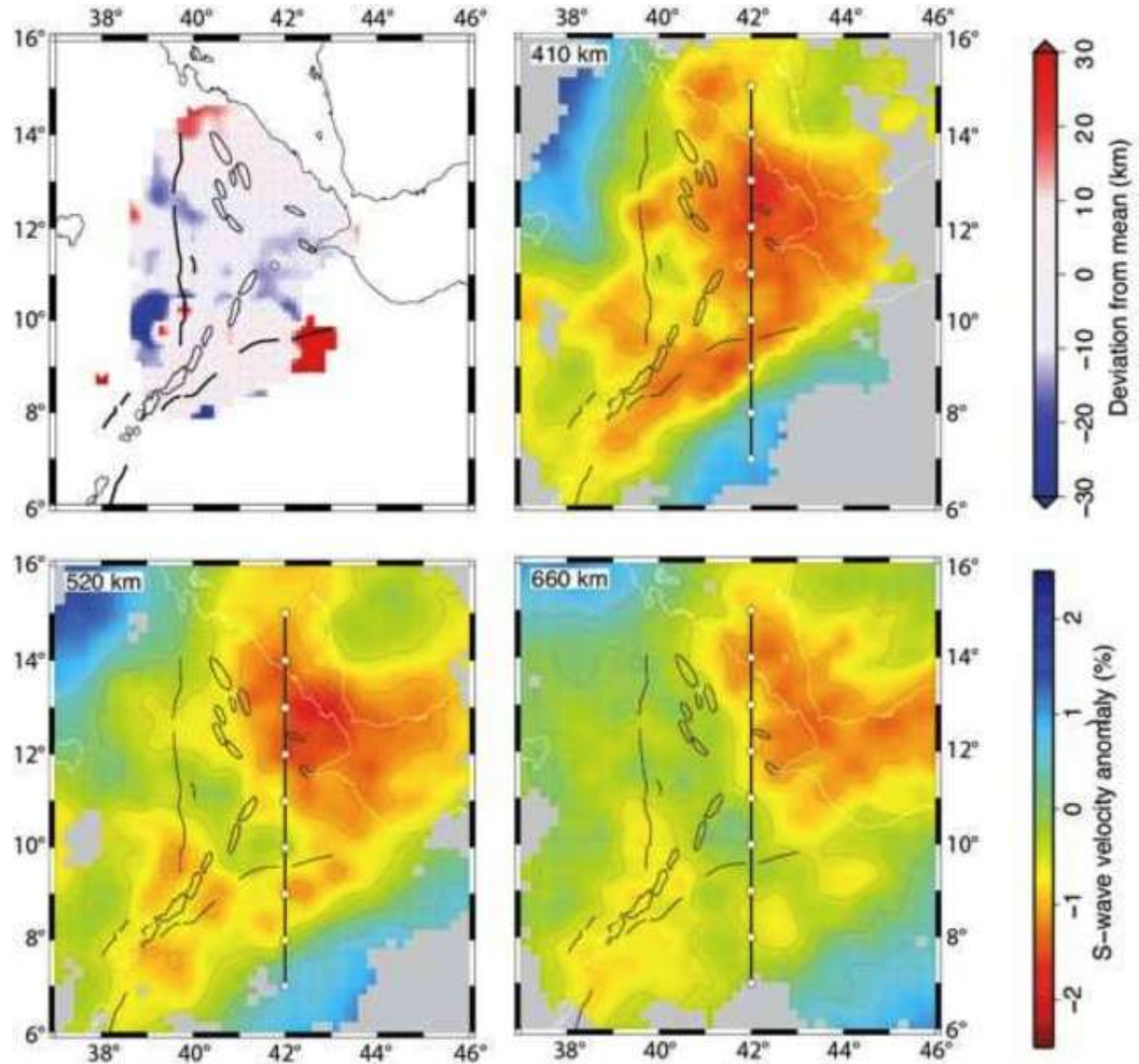
Thompson et al., 2013

Mantle transition zone

Little topography on the 410 and 660 km

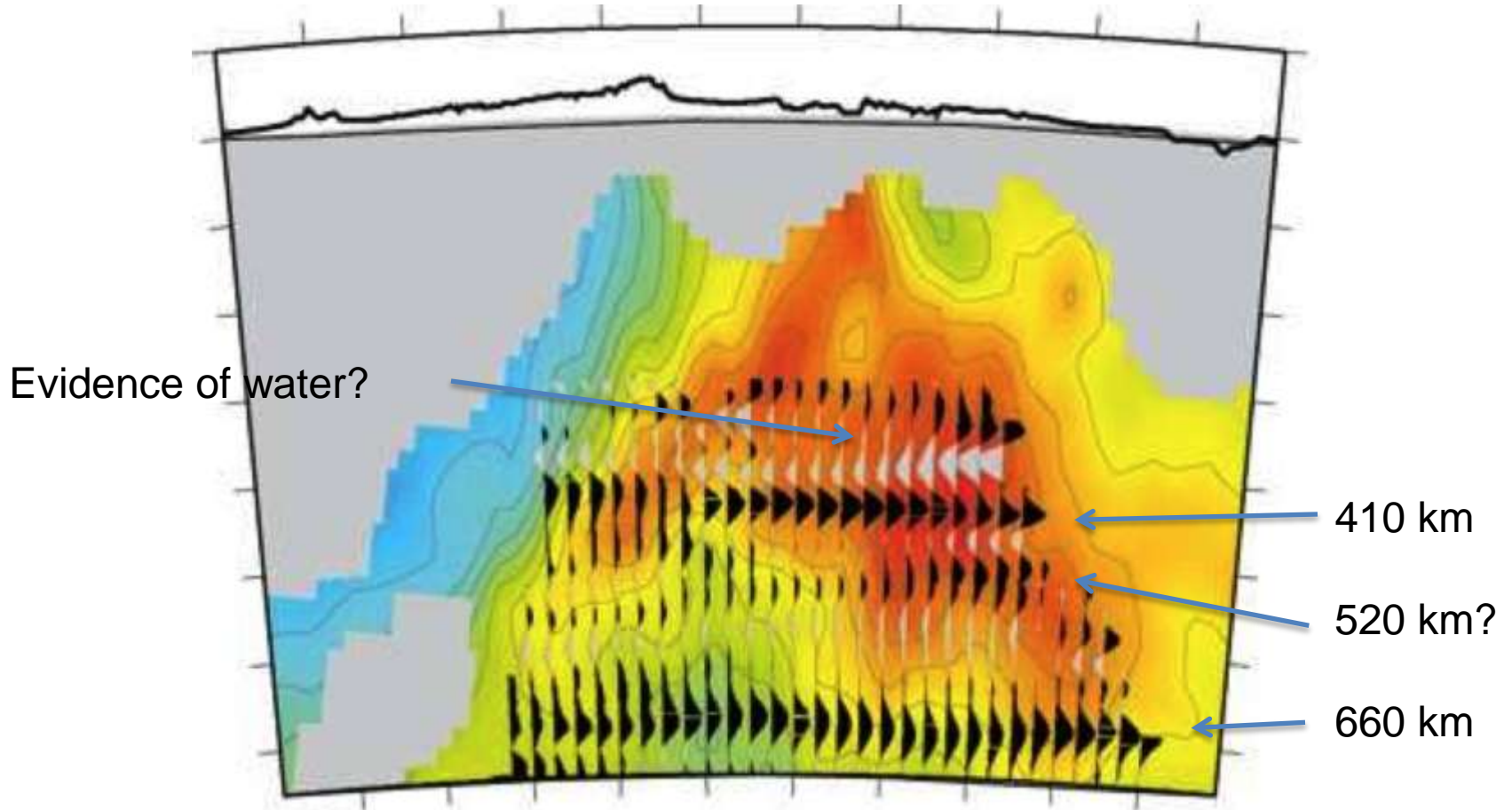
Localised upwelling in the MTZ

Little evidence for thermal anomaly



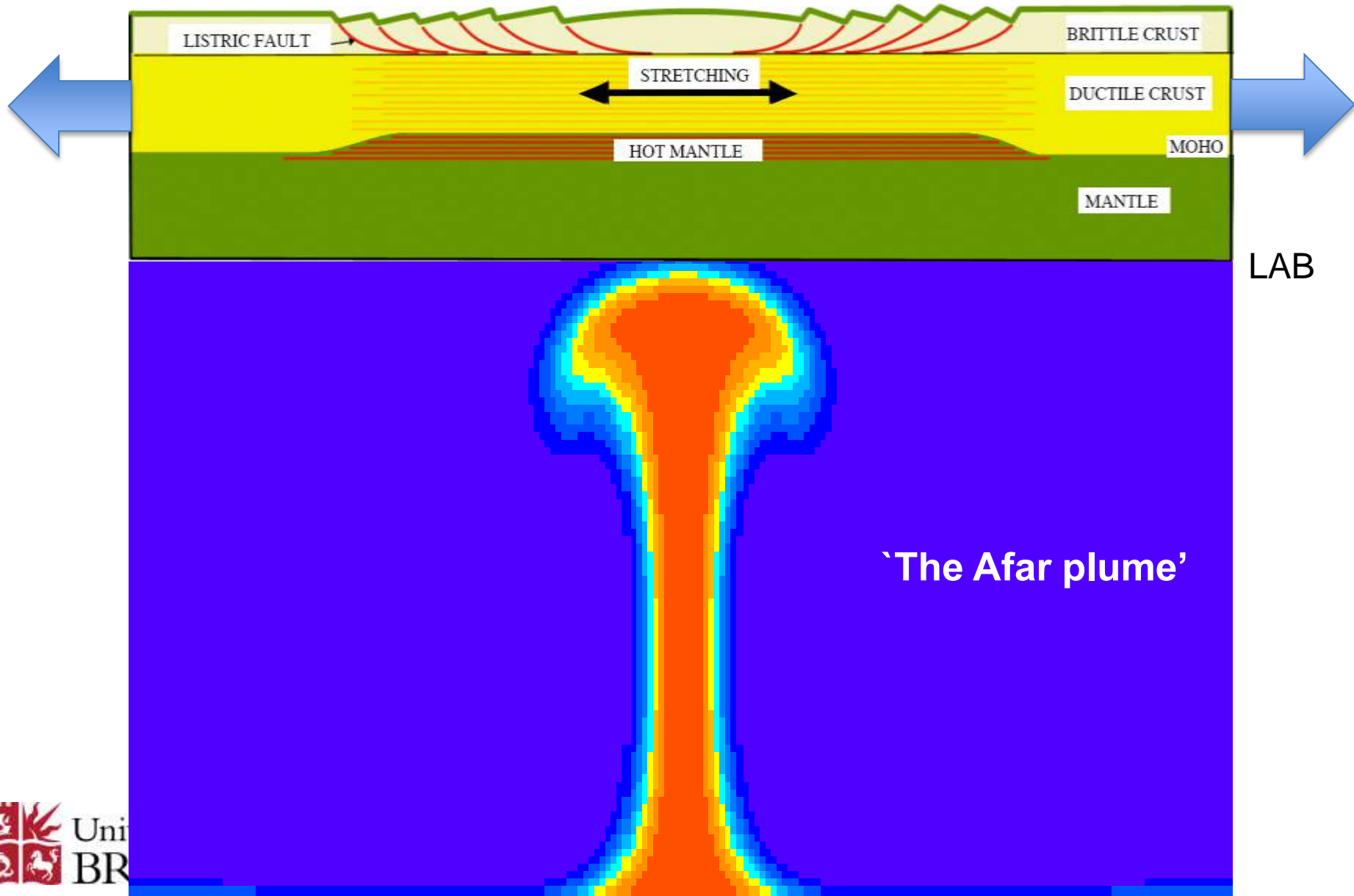
Thompson et al., 2013

Mantle transition zone



The old view

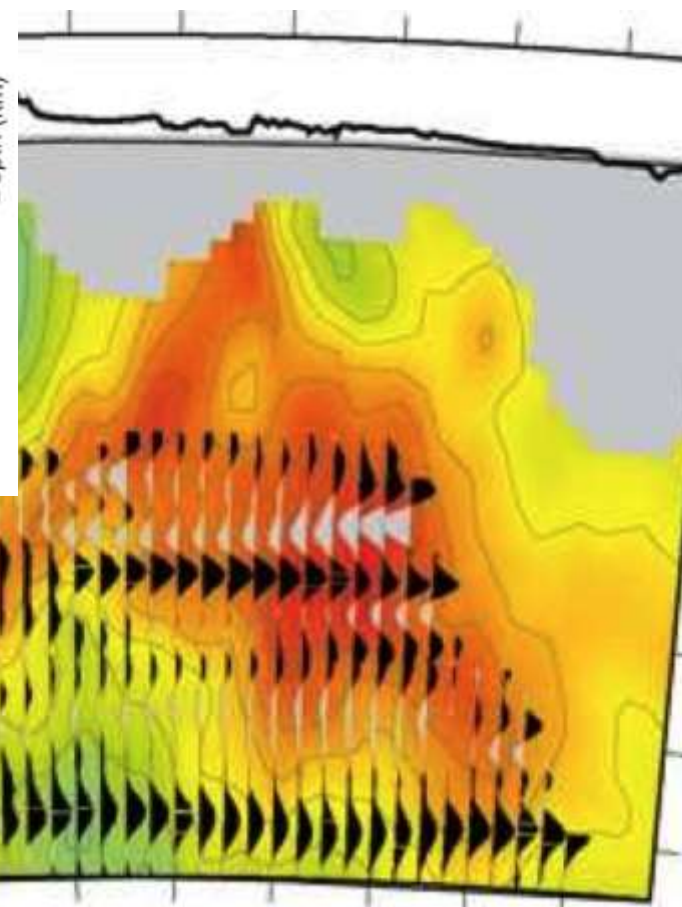
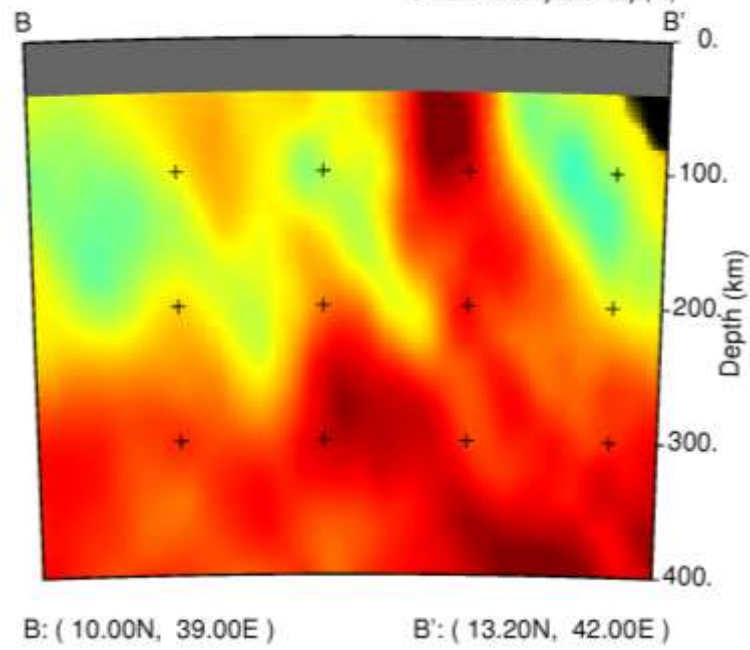
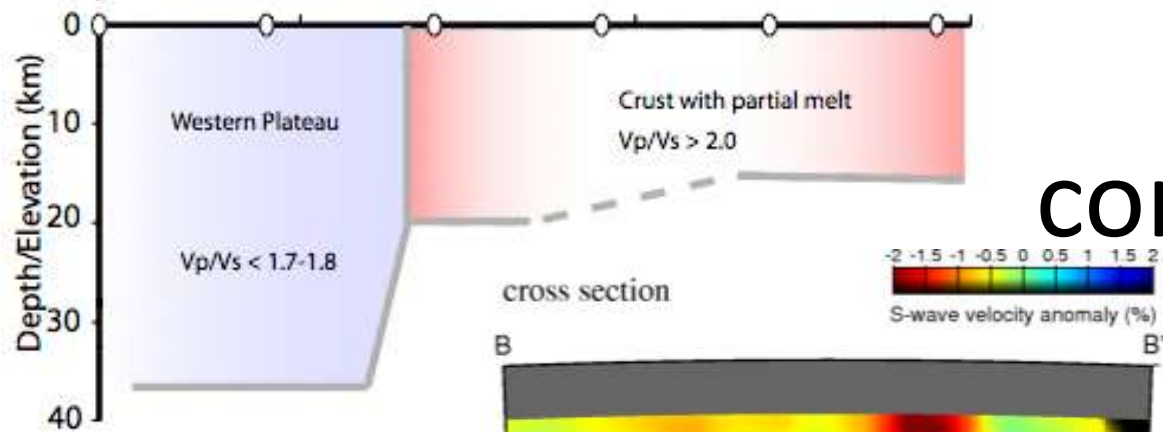
(a) Pure Shear



'The Afar plume'

LAB

New view of continental rifting



- Hammond et al., 2011
- Rychert et al., 2012
- Hammond et al., 2013
- Thompson et al., 2013

Conclusions

- Broad thermo-chemical upwelling from the CMB gives rise to more localised upwellings through the MTZ
- **Broad region of sub-lithospheric low velocities underlay region and give rise to dynamic topography (SW-NE orientation) – anisotropy reveals dynamic nature of this layer**
- Punctuated upwellings through the lithosphere (rift segmentation) mimic surface structure.
- **Eroded LAB beneath the rift – passive upwelling?**
- Observations provide ‘fingerprint’ of magma-assisted rifting; do not support simple mechanical stretching.

Issues

- Slowest traveltimes residuals on Earth and yet the mantle doesn't seem that hot:
 - LAB suggests ridge-like decompression melting
 - Mantle xenoliths (e.g., Rooney et al., 2011; Ferguson et al., 2013)
 - MTZ topography looks unremarkable, but may show fingerprint of water above 410km
- Is chemistry more important? Exotic melts (carbonates or sulfide melt)

