

Université
Paris 7-Denis Diderot



Oceanic and continental plates as seen by seismic anisotropy

Jean-Paul Montagner^{1,2,3}

Gael Burgos, Satish Maurya

Eric Beucler, Yann Capdeville, Ravi Kumar, Eleonore Stutzmann

1- Laboratoire Sismologie, I.P.G., Paris, France

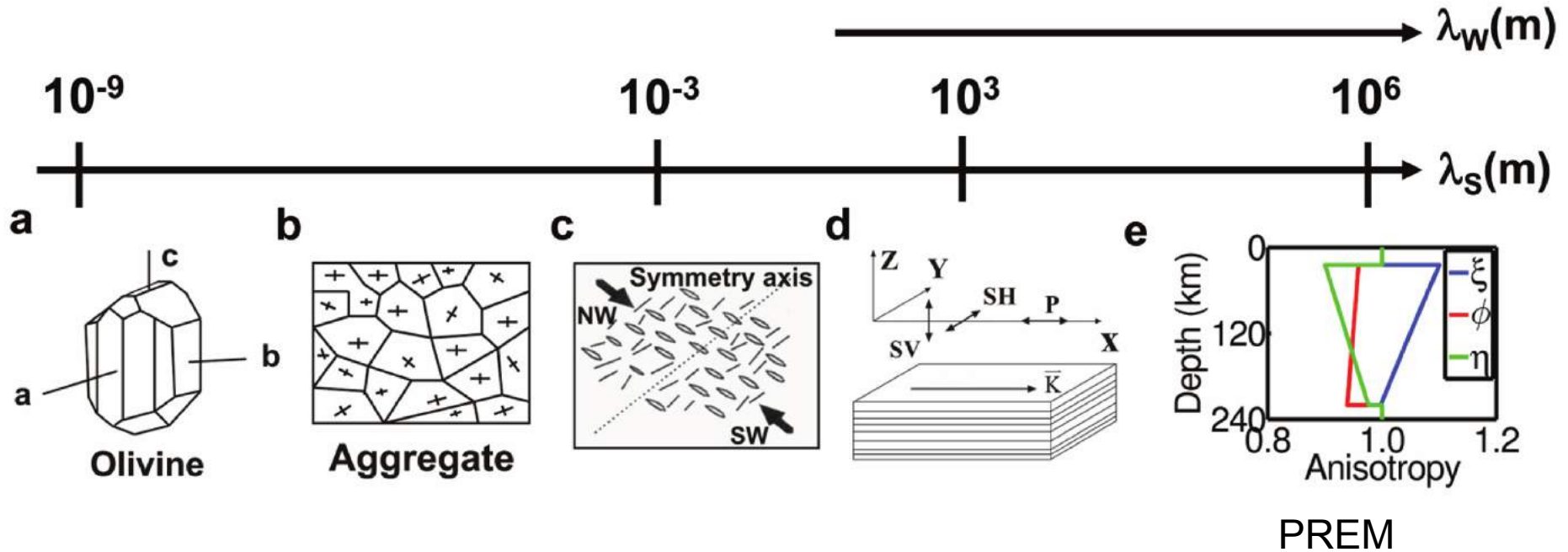
2- Université Paris-Diderot, USPC

2- Institut Universitaire de France

OUTLINE

- What is a plate? Role of continents?
Surface manifestation of mantle convection,
role of mantle plumes, slabs?
- Seismic Anisotropy: many processes, different interpretations
- 3D- anisotropic structure of the Earth
Lithosphere- Asthenosphere Boundary (LAB)
Oceans (Burgos et al., 2014)
Continents: Indian Continent (Maurya et al., 2016)

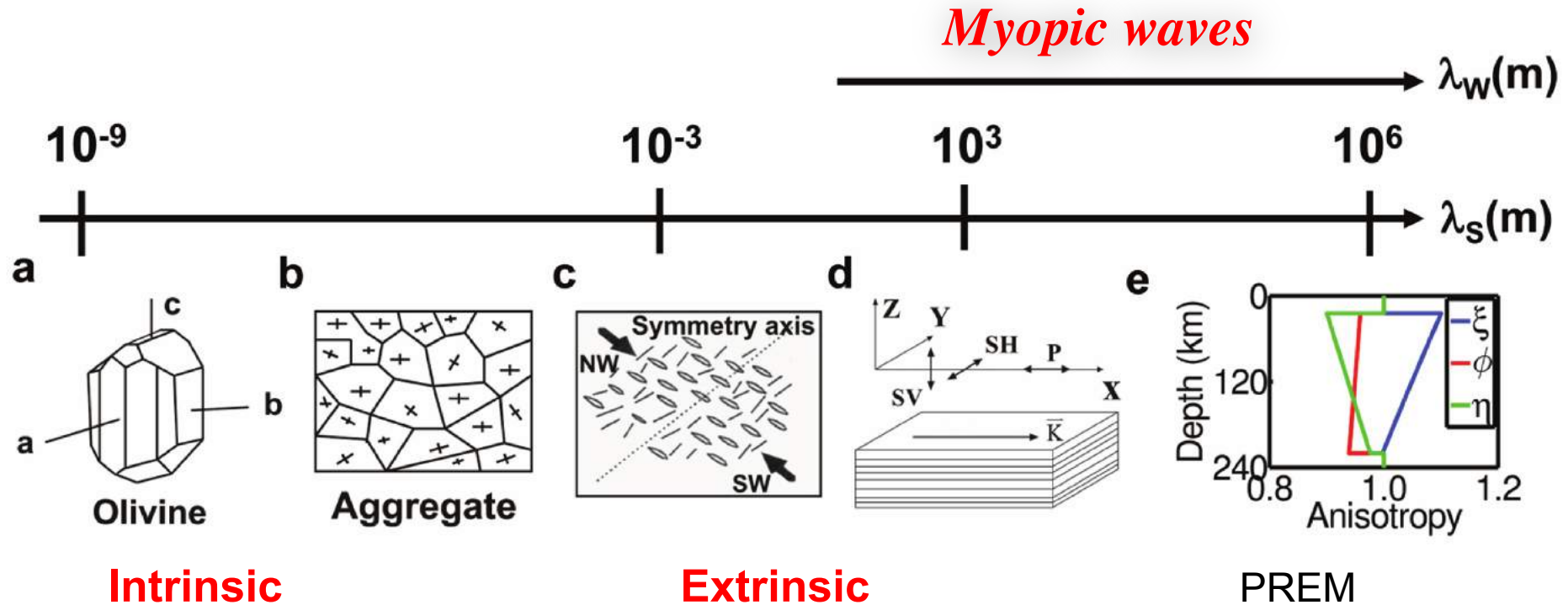
Seismic Anisotropy at all scales



PREM: radial anisotropy: up to 10%

λ_W seismic wavelength
 λ_S spatial scale

Seismic Anisotropy at all scales

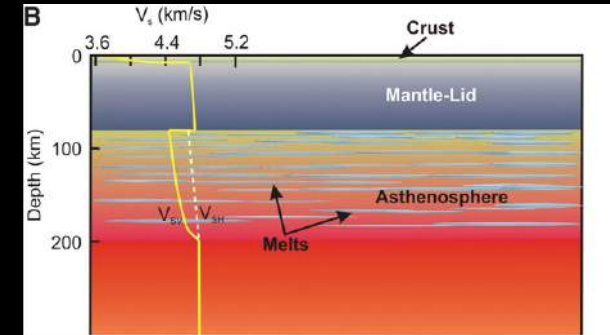
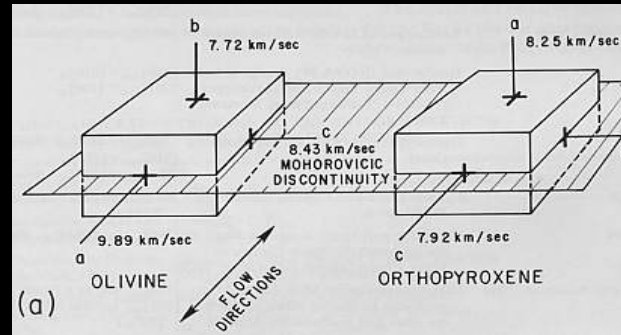
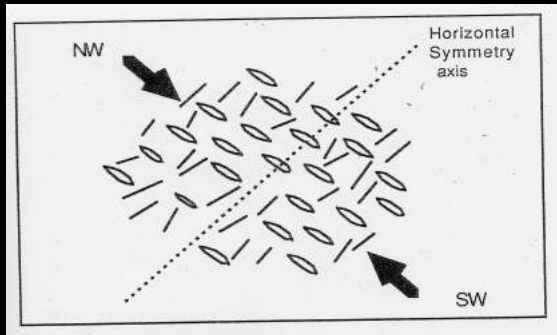


Observed (apparent) anisotropy
Intrinsic versus Extrinsic anisotropy

$$\alpha = p\alpha^{\text{int}} + (1-p)\alpha^{\text{ext}}$$

Different processes in different layers

-S.P.O. (stress) -L.P.O.(strain) Fine Layering



- ***Mineralogy, Water and fluid content, layering***
- ***Present day tectonics, geodynamic processes***
- ***Mantle Convection***
- ***Past processes (frozen anisotropy)***

***Separation of the different kinds of anisotropy in different layers => Different interpretations
(Stratification of anisotropy in the crust & mantle)***

Effect of anisotropy on the phase of surface waves

Effect on eigenfrequency ω_k for multiplet $k=\{n,l,m\}$ (Rayleigh's principle)

$$\frac{\delta\omega_k}{\omega_k} = \frac{\int_{\Omega} \varepsilon_{ij}^* \delta C_{ijkl} \varepsilon_{kl} d\Omega}{\int_{\Omega} \rho_0 u_r^* u_r d\Omega} = \frac{\delta V}{V} \Big|_k$$

ε strain tensor, u displacement, δC_{ijkl} elastic tensor perturbation,

V phase velocity (V_R Rayleigh; V_L Love)

Phase velocity perturbation $\delta V(T, \theta, \phi, \Psi)$ at point $r(\theta, \phi)$

(Smith & Dahlen, 1973; Montagner & Nataf, 1986)

Ψ Azimuth (angle between North and wave vector)

Azimuthal Terms: $0-\psi, 2-\psi, 4-\psi$

$$\delta V(T, \theta, \phi, \Psi) / V = \alpha_0(T, \theta, \phi) + \alpha_1(T, \theta, \phi) \cos 2\Psi + \alpha_2(T, \theta, \phi) \sin 2\Psi + \alpha_3(T, \theta, \phi) \cos 4\Psi + \alpha_4(T, \theta, \phi) \sin 4\Psi$$

• *Cijkl 21 elastic moduli*

• $\alpha_0 = 0$ - ψ term: 5 parameters *A, C, F, L, N (PREM)*

VTI Model (transverse isotropy with vertical symmetry axis)

• *Best resolved parameters from surface waves (among 13 parameters when including azimuthal anisotropy 2ψ -, 4ψ - terms)*

$L = \rho V_{SV}^2$ *Isotropic part of V_{SV} or $\mu_{iso} \approx (2L+N)/3$*

$\xi = N/L = (V_{SH}/V_{SV})^2$ *Radial Anisotropy*

G, Ψ_G *Azimuthal Anisotropy of V_{SV} , also related to SKS splitting*

Geodynamic
Interpretation: LPO

Convective cell: anisotropic parameters

Tomographies of:

-S- Velocity

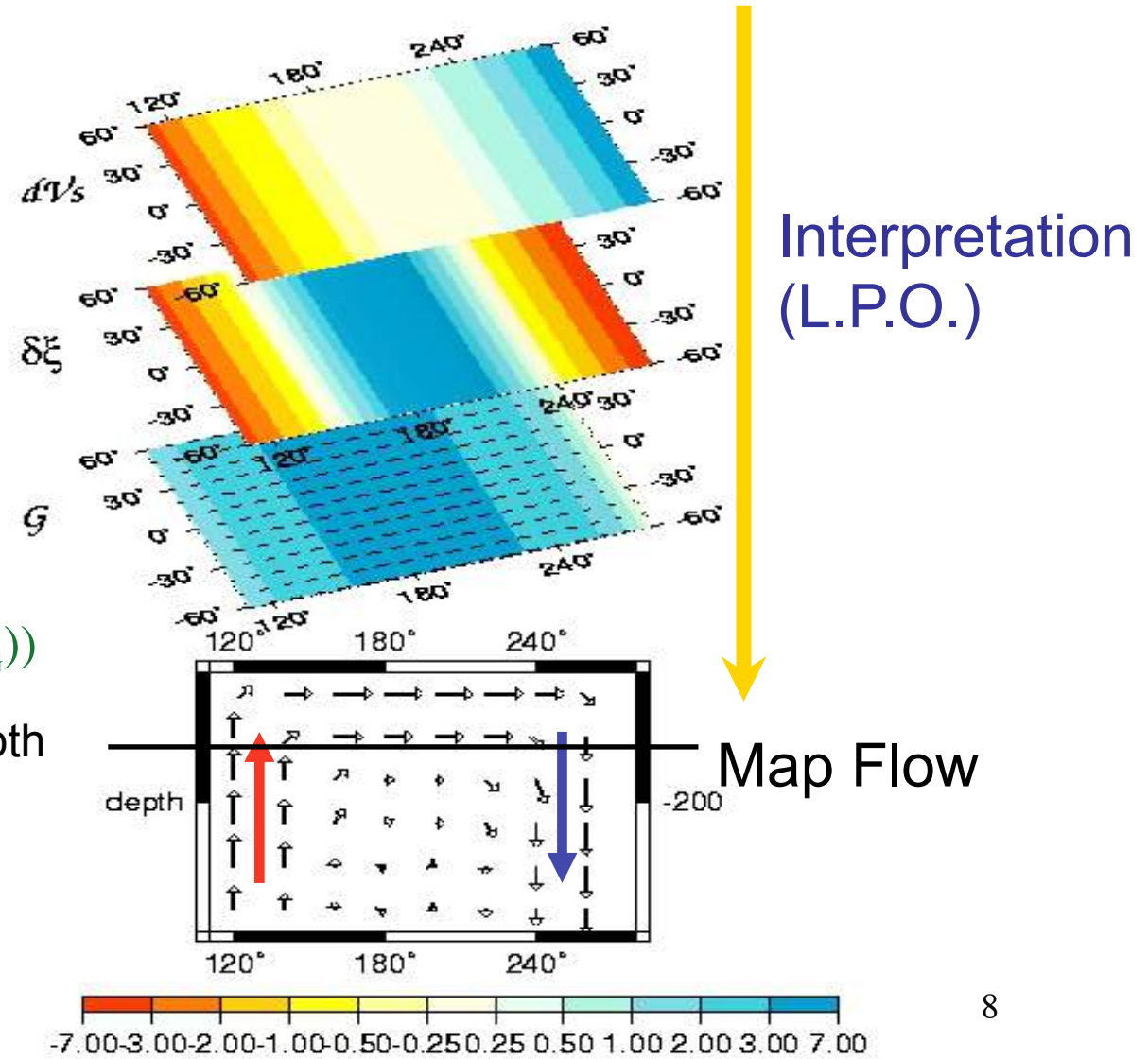
-Radial Anisotropy

$$\delta\xi = (V_{SH}^2 - V_{SV}^2) / V_{SV}^2$$

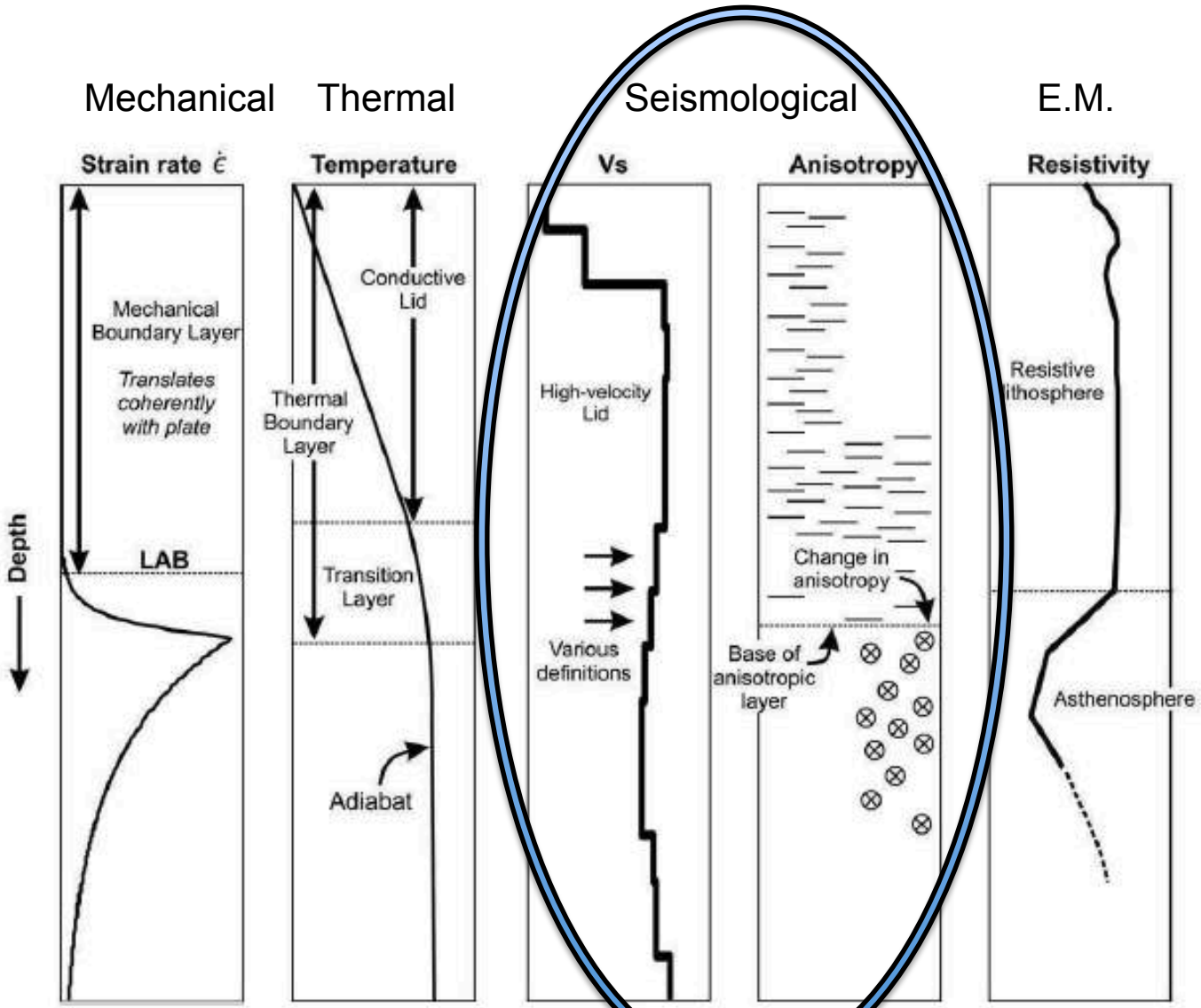
-Azimuthal Anisotropy

$$V_{SV} \approx V_{SV0} + \frac{1}{2} G \cos(2(\Psi - \Psi_G))$$

At a given depth



L.A.B.: Lithosphere-Asthenosphere Boundary (many different approaches and definitions)

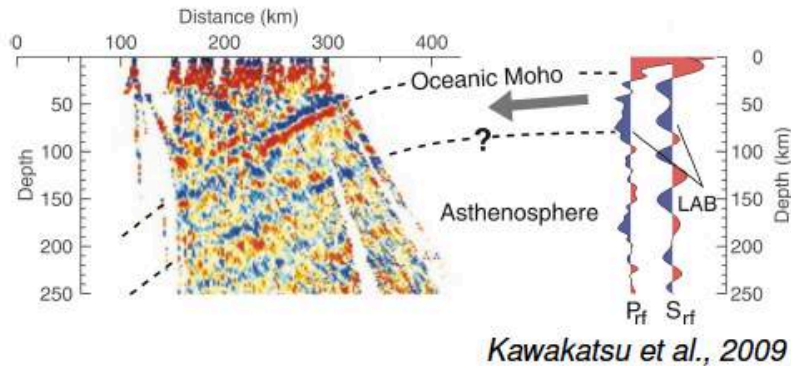


02 Dec. 2016

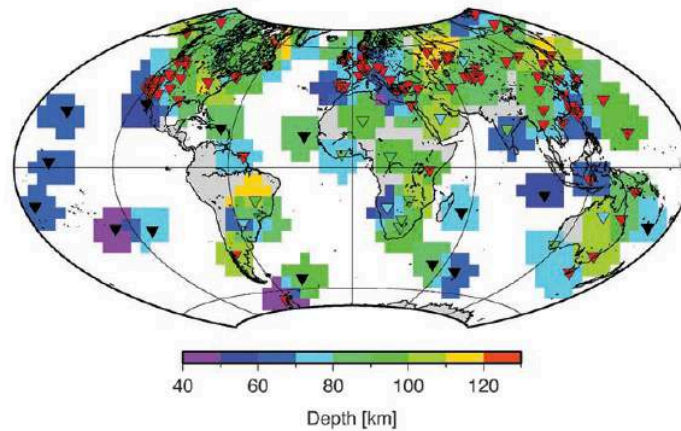
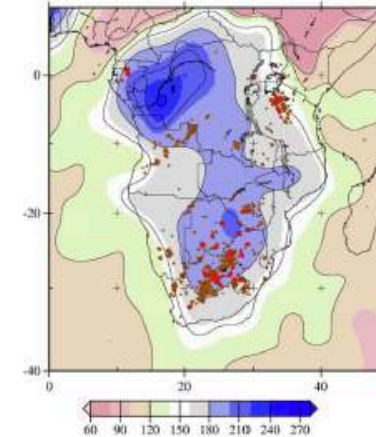
Eaton et al., 2008

LAB : from seismic data

Receiver functions

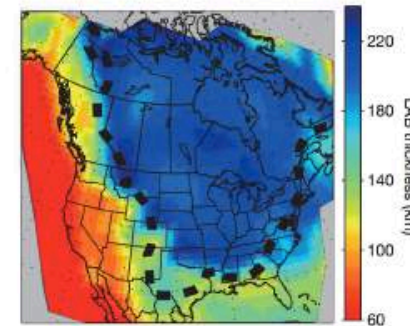


Surface waves



Fishwick, 2010

Rychert & Shearer, 2009

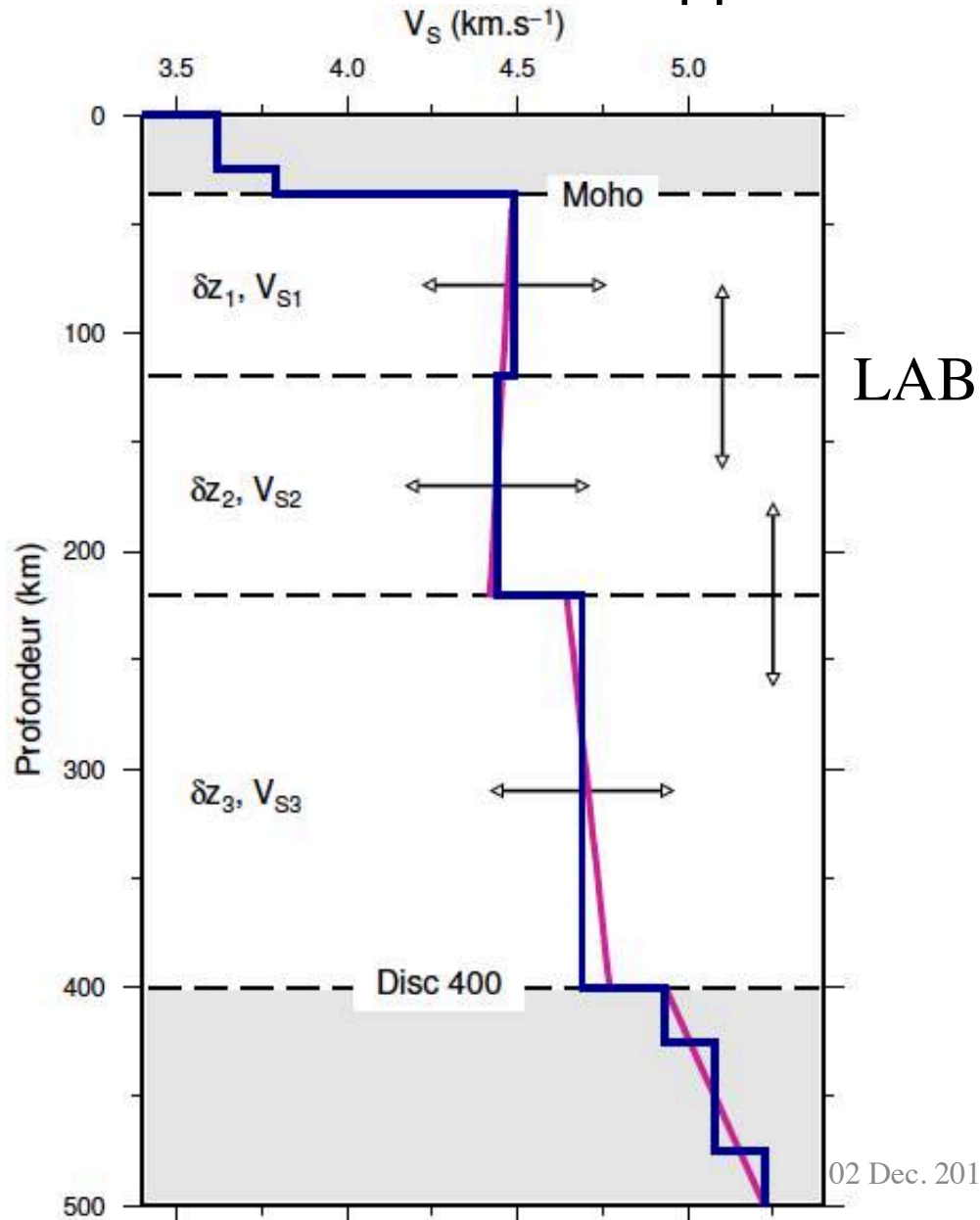


+SKS Yuan & Romanowicz, 2010

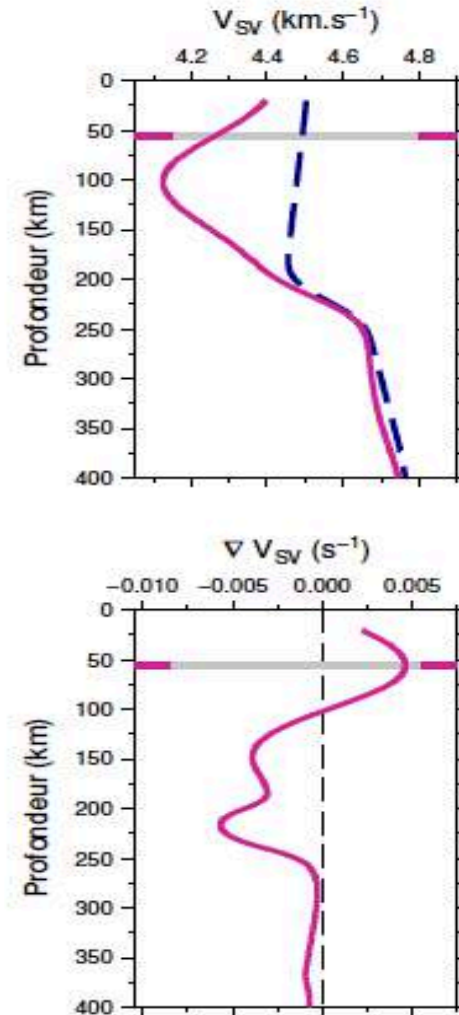
Much discrepancy between different estimates:
Global surface wave tomographies give 200-250km depth for continental roots
Mid-lithospheric Discontinuity (Yuan & Romanowicz, 2010)

From Surface wave dispersion

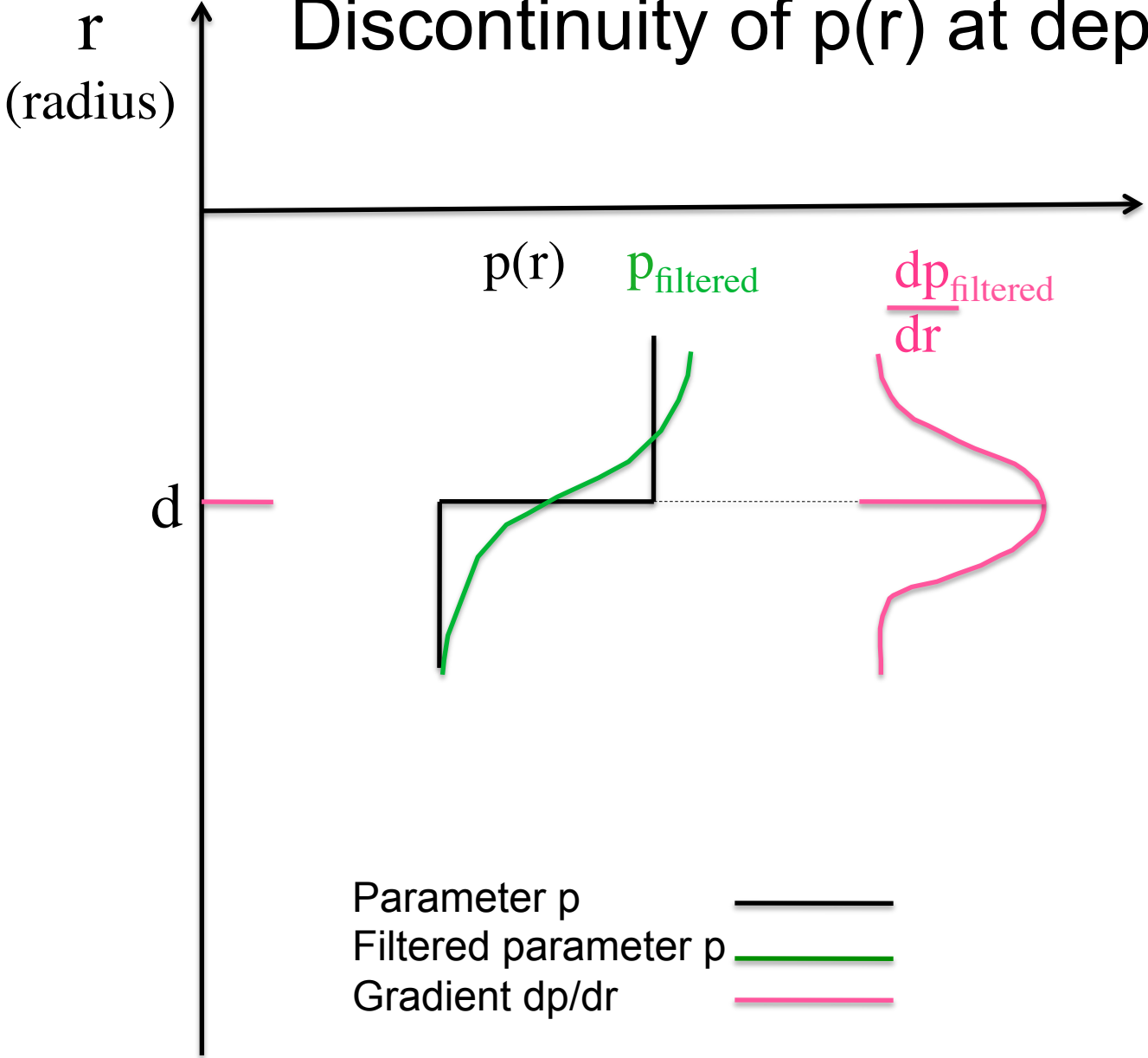
Statistical Monte-Carlo Approach First order Perturbation theory



Proxy from parameter V_{SV}



Discontinuity of $p(r)$ at depth d



Other proxies from parameters obtained from anisotropy tomographic models

Well resolved parameters:

V_{SV} S-wave velocity

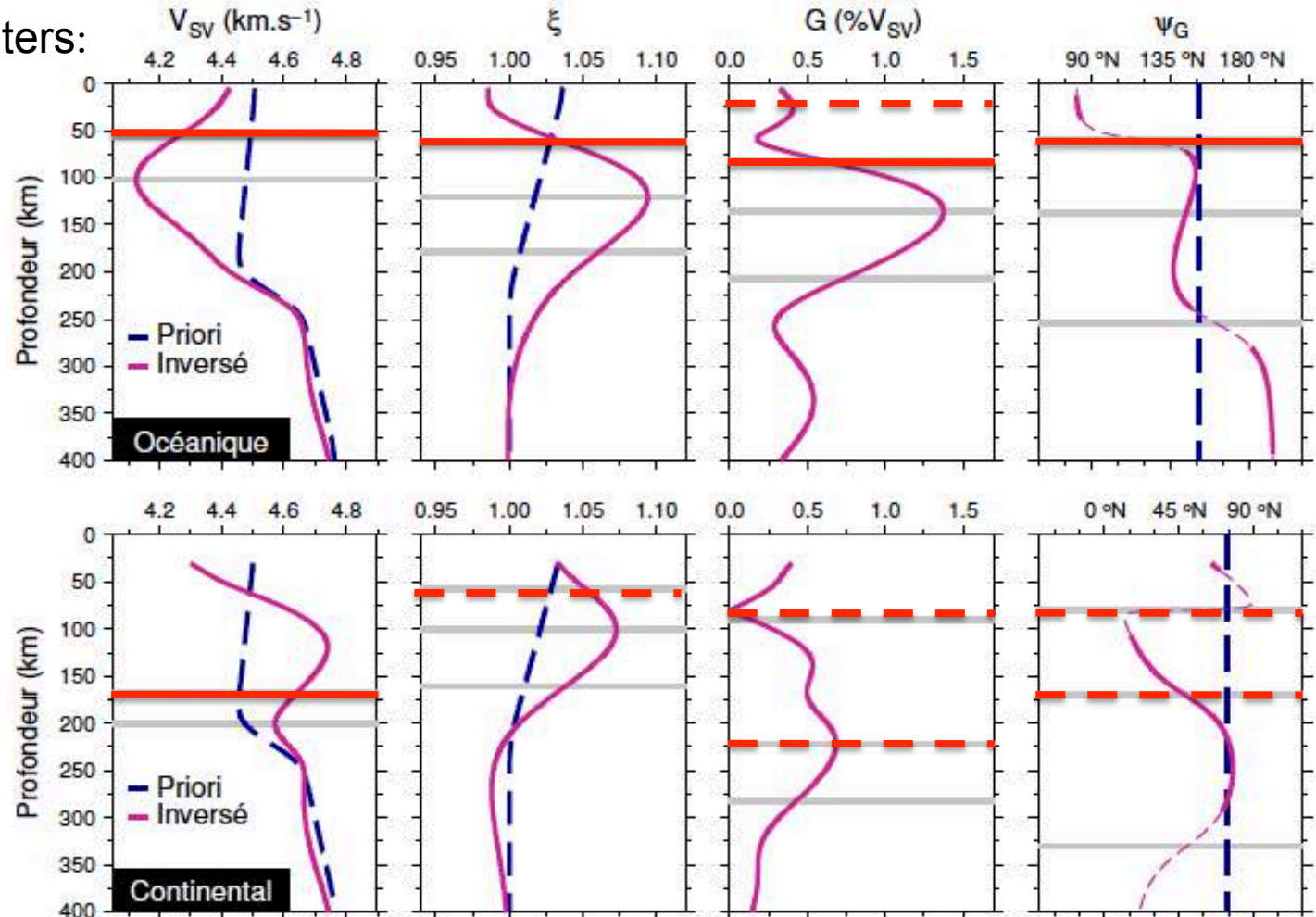
ξ , radial anisotropy

G, Ψ_G S-wave

azimuthal anisotropy

Oceanic profile

$\lambda=35^\circ, \phi=-35^\circ$



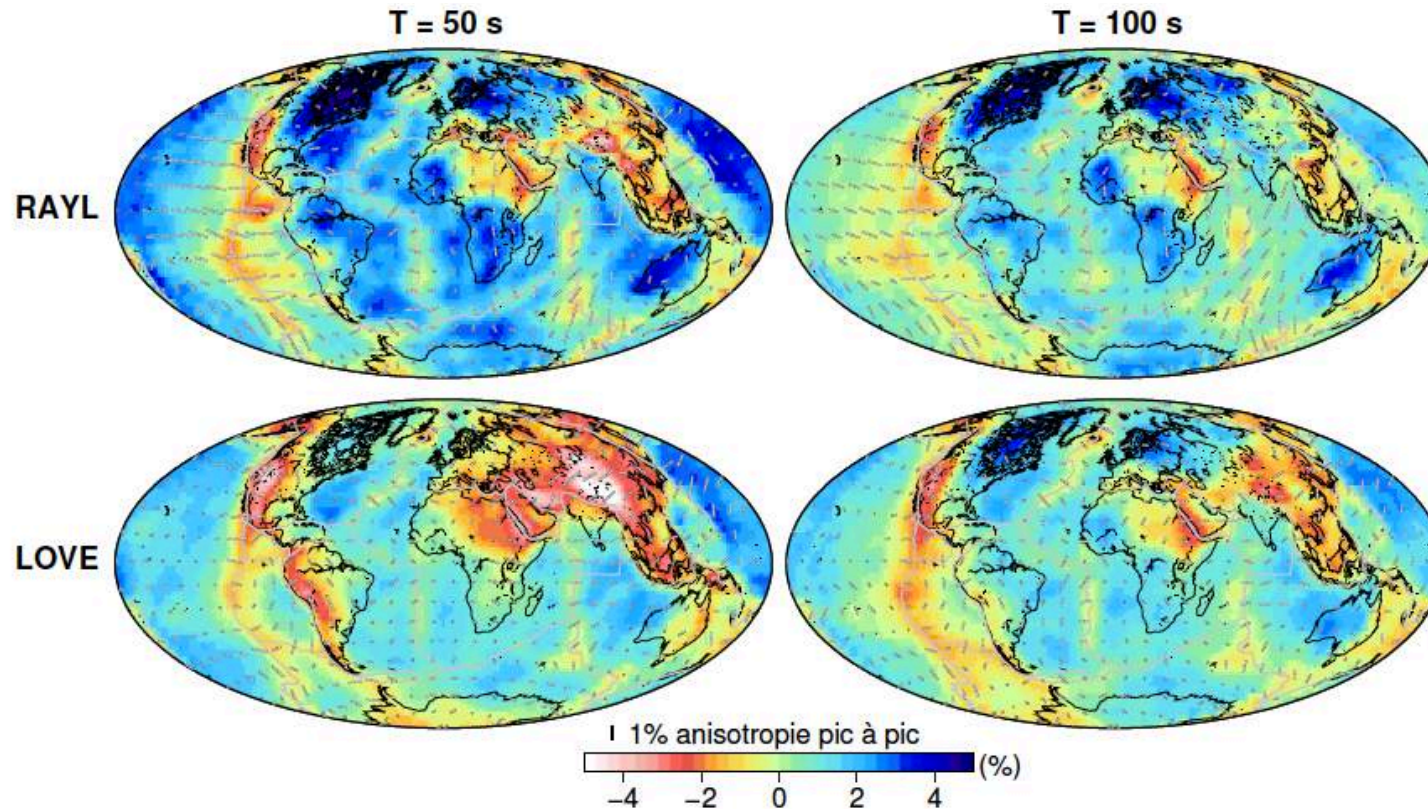
Continental profile

$\lambda=63^\circ, \phi=-96^\circ$

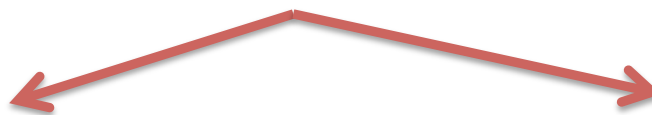
Collection of surface wave dispersion data (IPGP, Harvard, Utrecht, Boulder)

First step: Regionalization => local dispersion velocity $V(T, \theta, \phi, \psi)$

Rayleigh phase velocity and azimuthal anisotropy



Second step: Inversion at depth

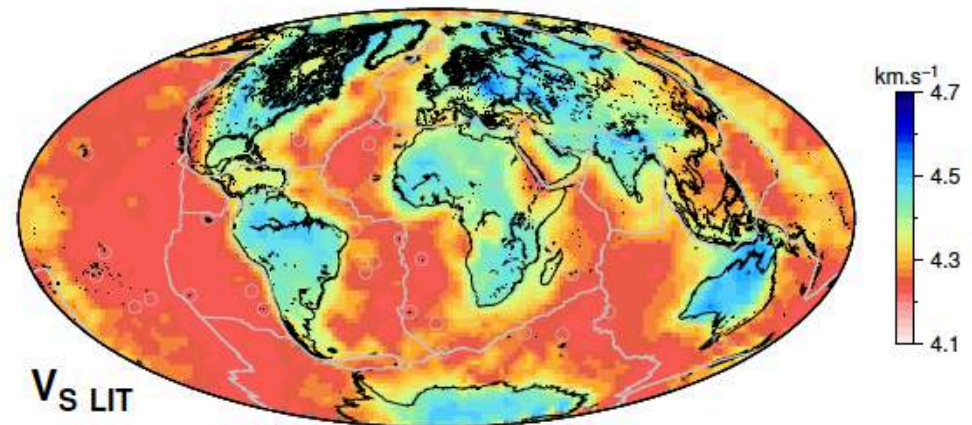
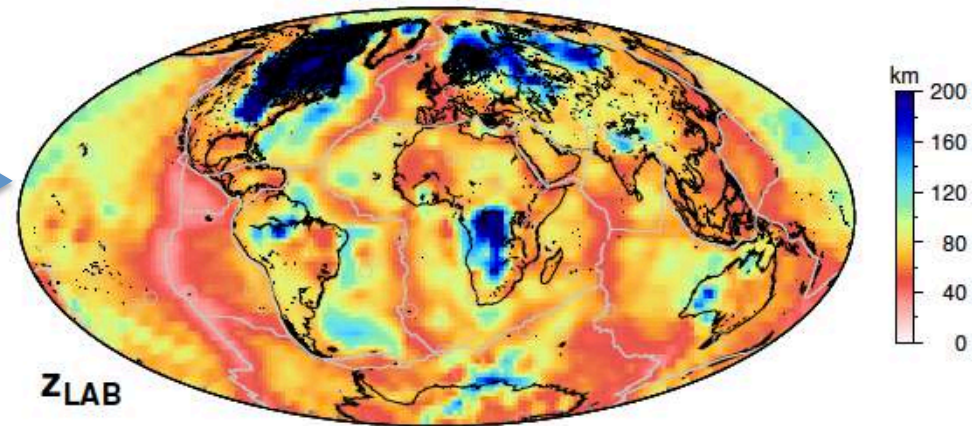
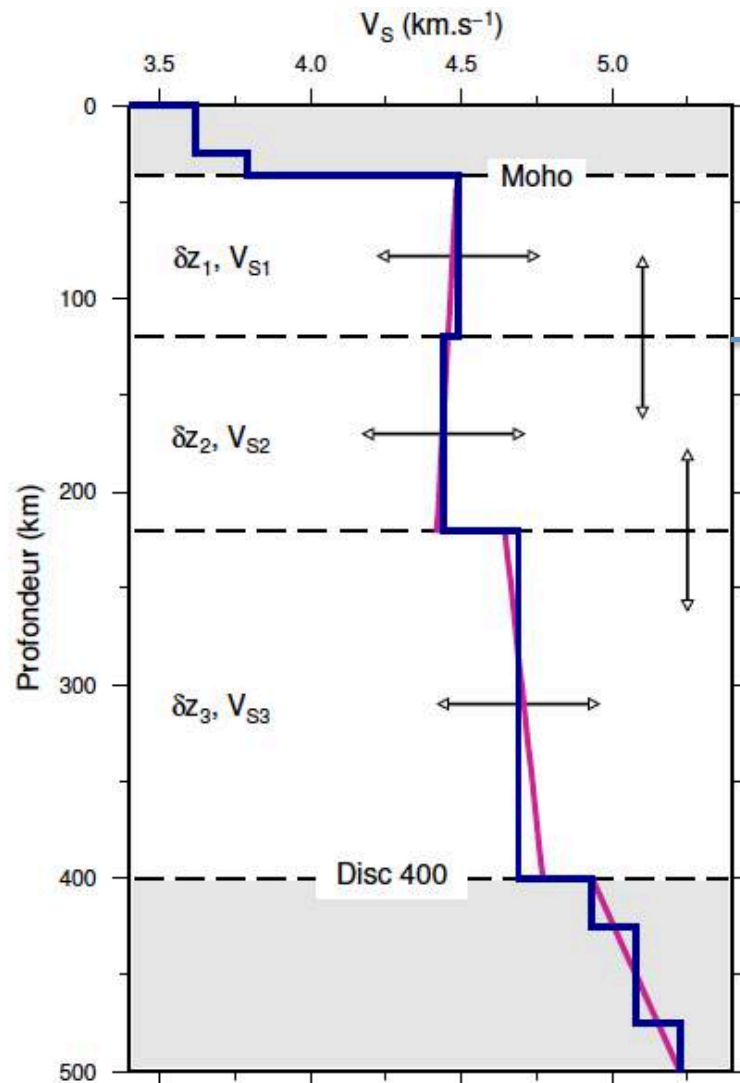


Statistical Monte-Carlo Inversion

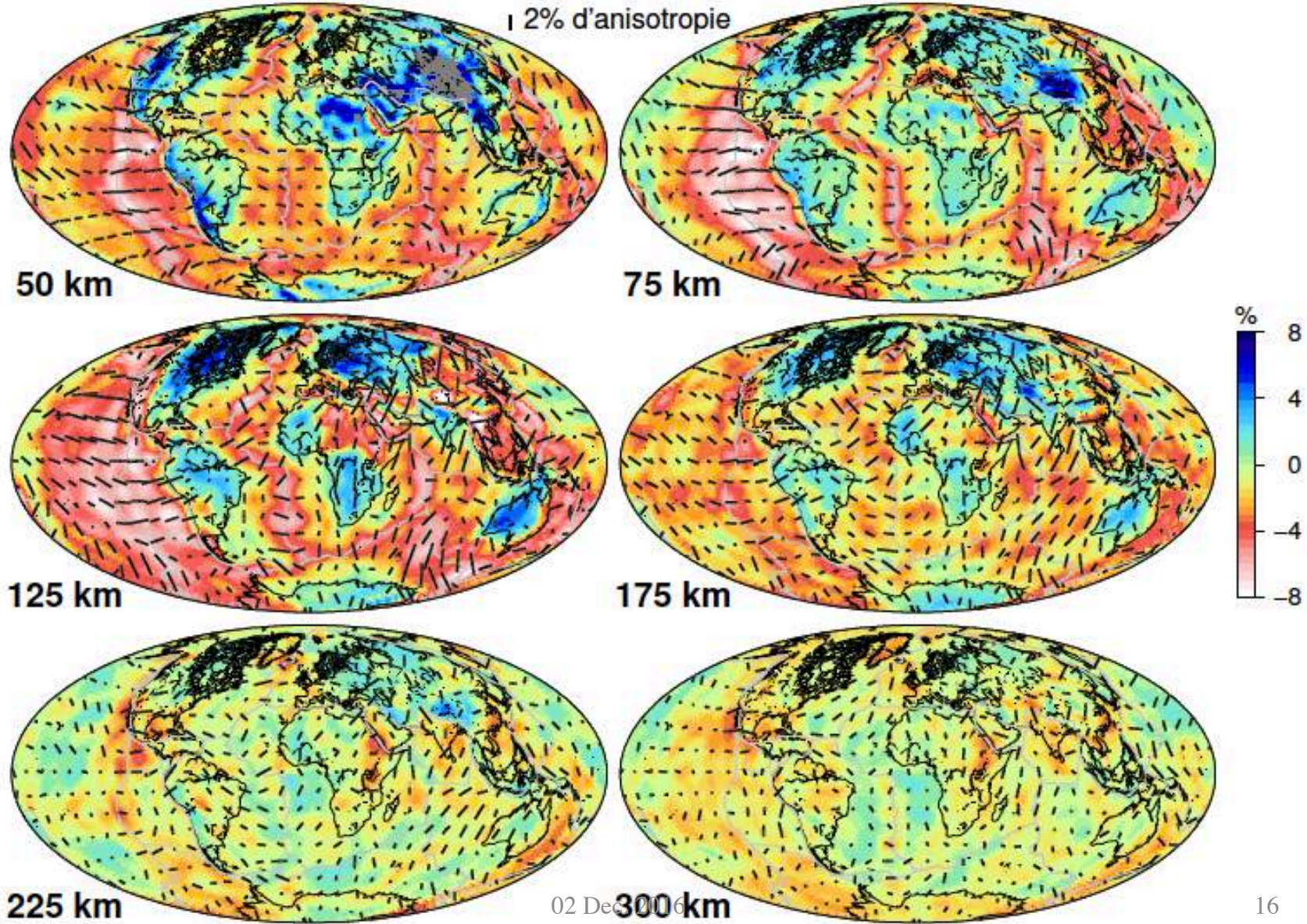
First order Perturbation

LAB: Statistical M.C. Isotropic Inversion

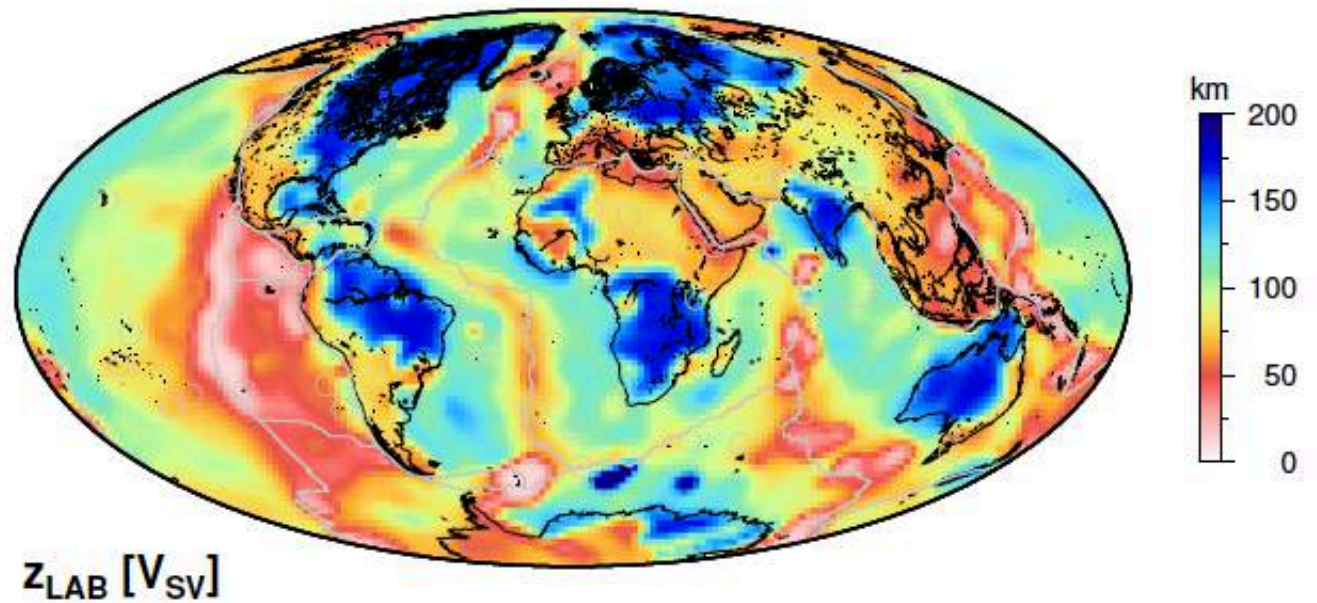
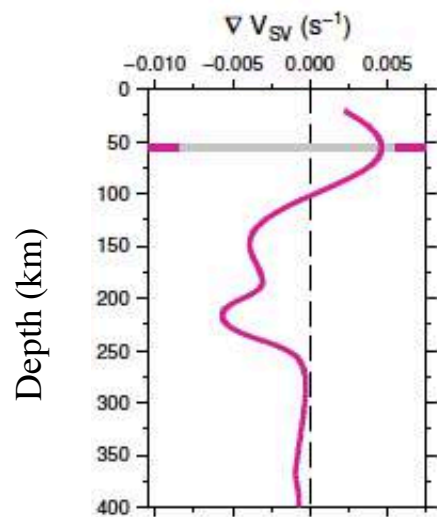
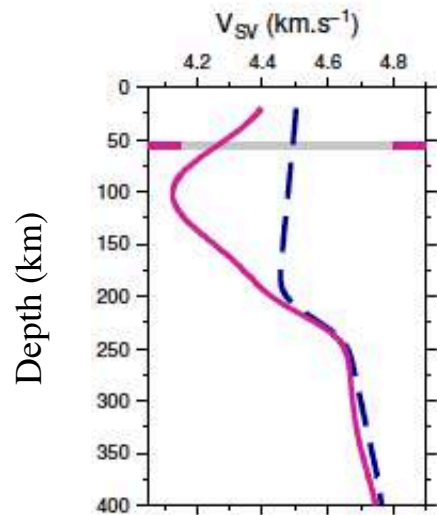
Data: C_R, C_L, U_R, U_L [30-300s], Parameters: $3V_s, 2\delta z$



First order perturbation Theory => depth distribution of V_{sv} , G (and ξ)

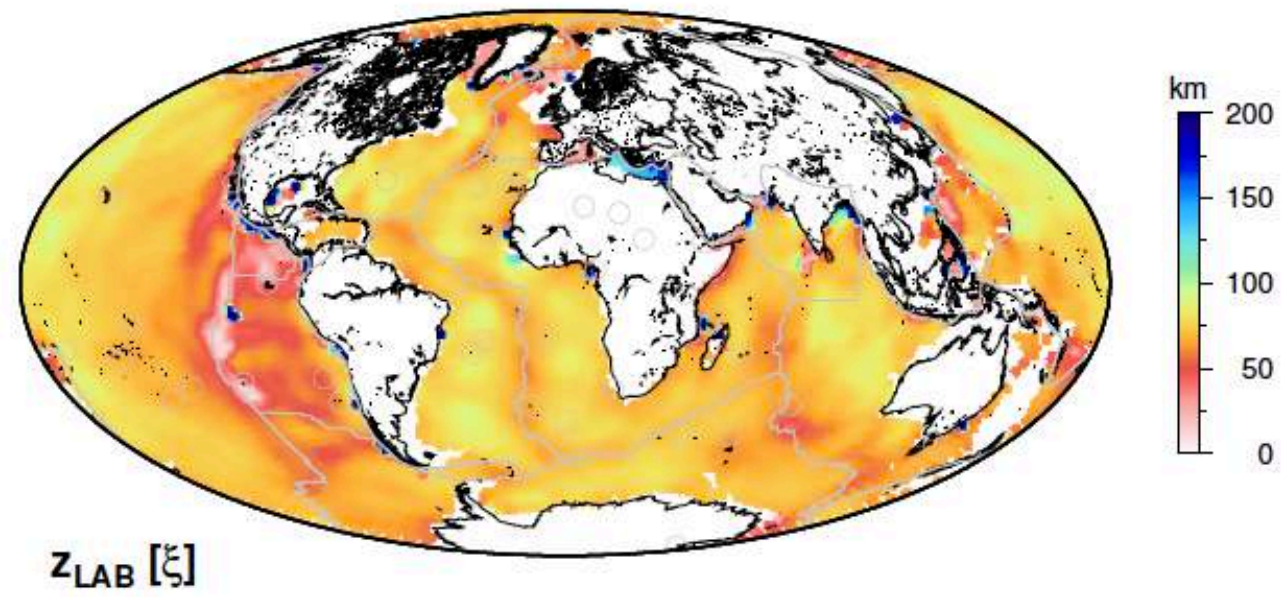
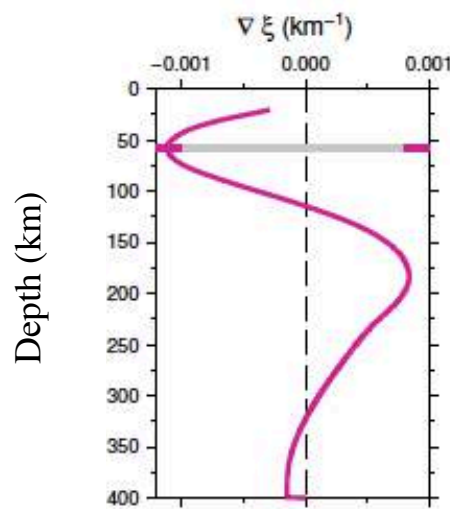
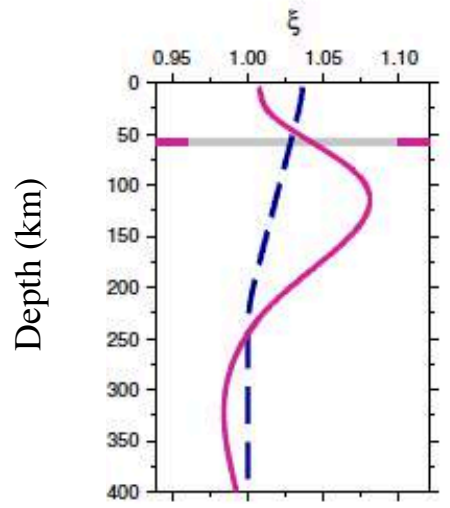


LAB from the gradient of VSV parameter

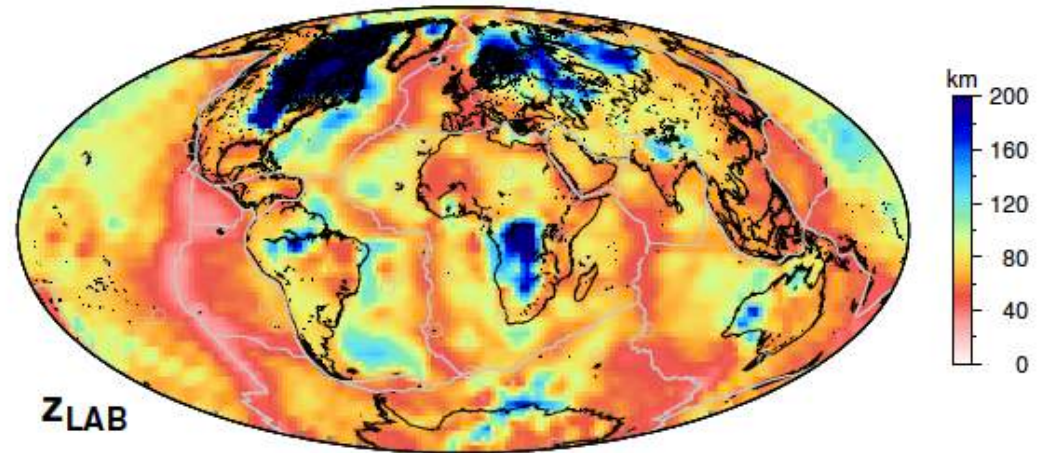


LAB from the gradient of ξ parameter (only oceans)

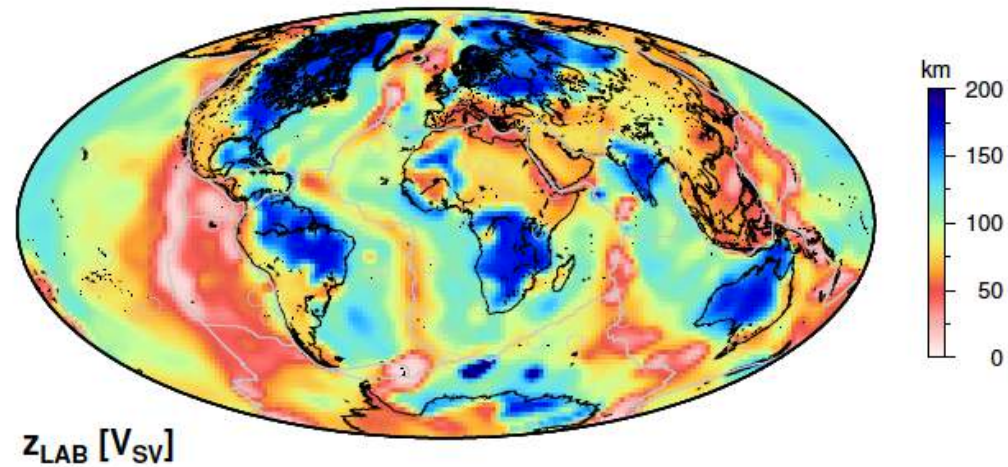
Radial anisotropy $\xi = (V_{SH}/V_{SV})^2$



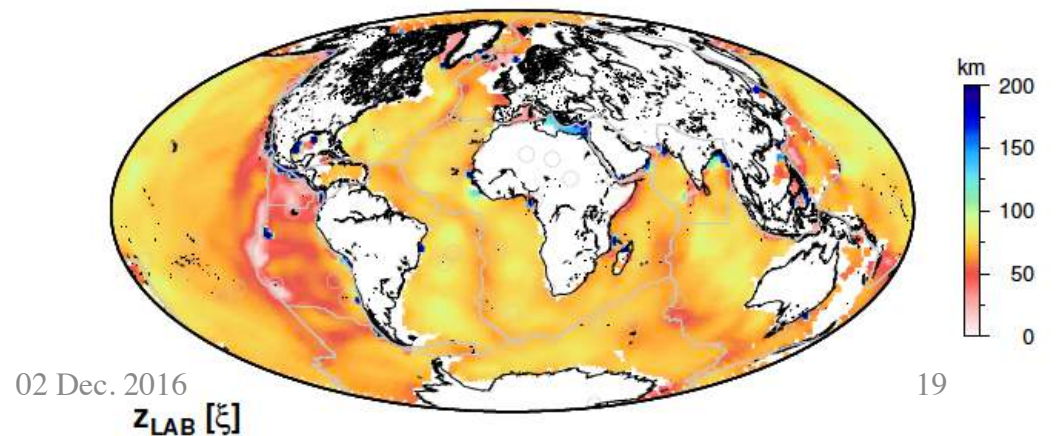
Statistical MC Isotropic
Inversion



V_{sv} proxy (1st order
Perturbation Theory)

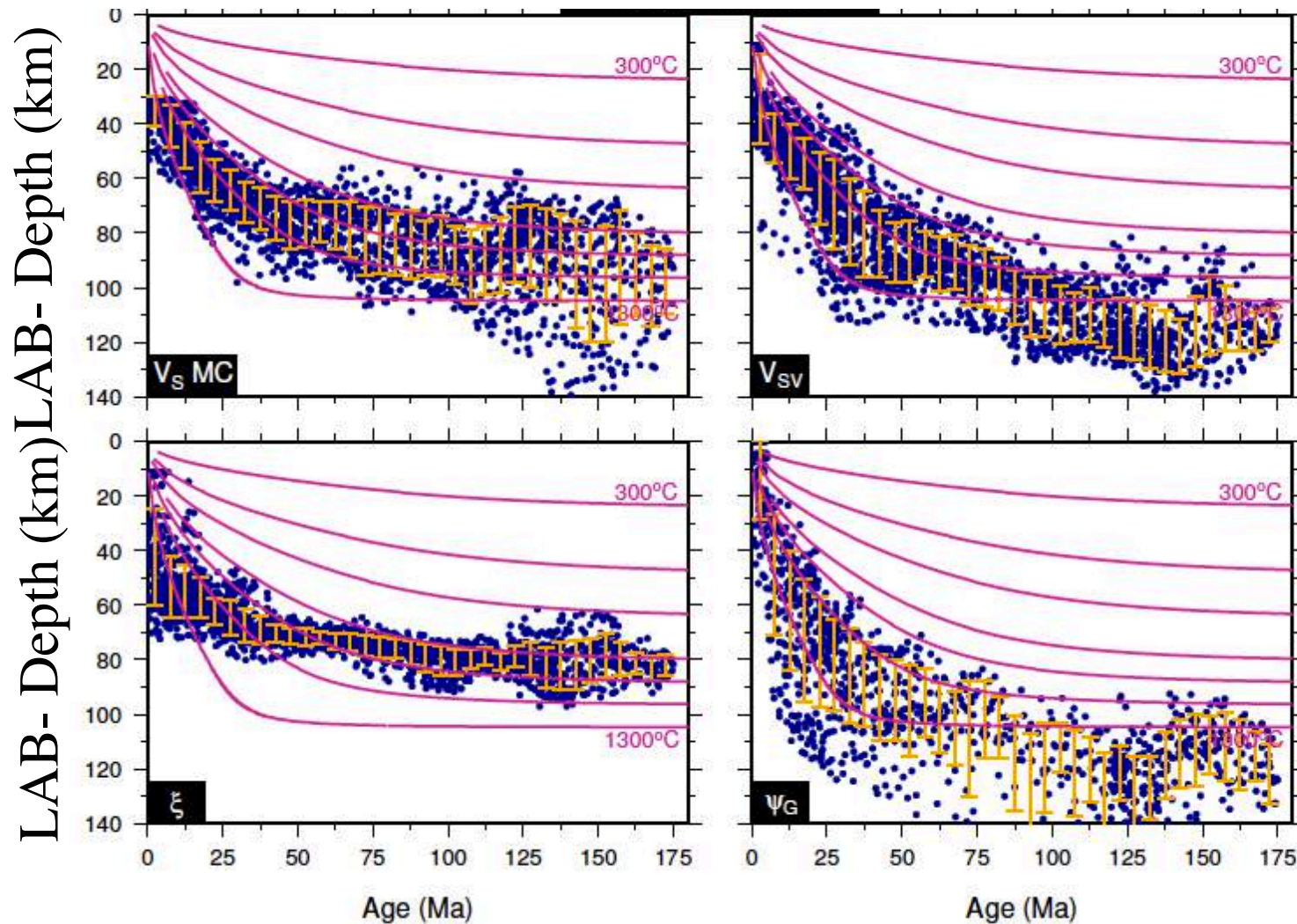


ξ proxy (1st order
Perturbation Theory)

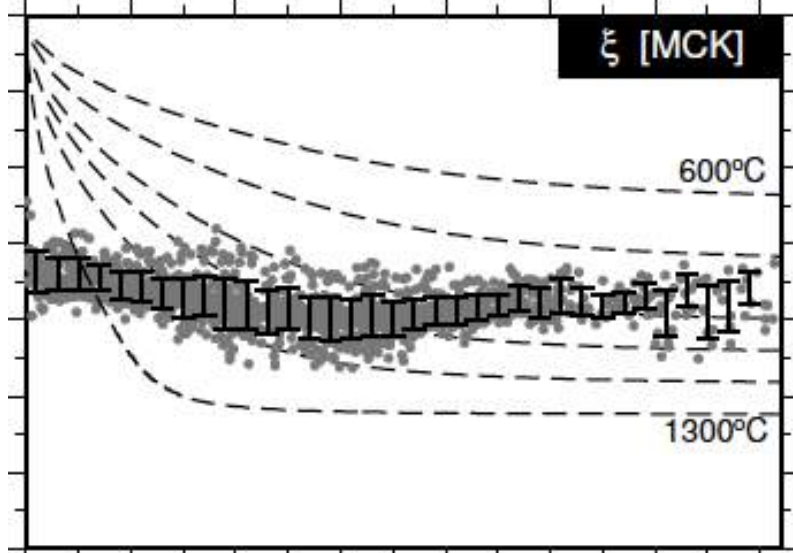
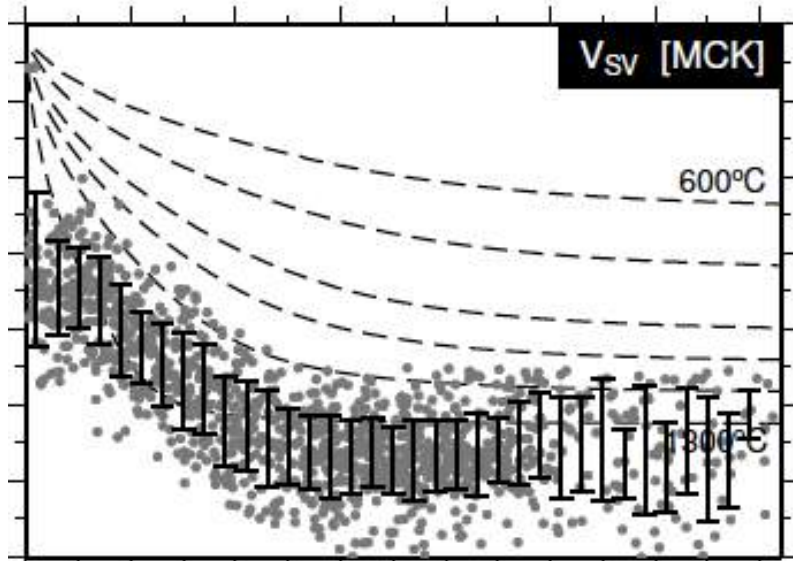


Age Variation of LAB depth in oceanic regions

Compared with Plate model



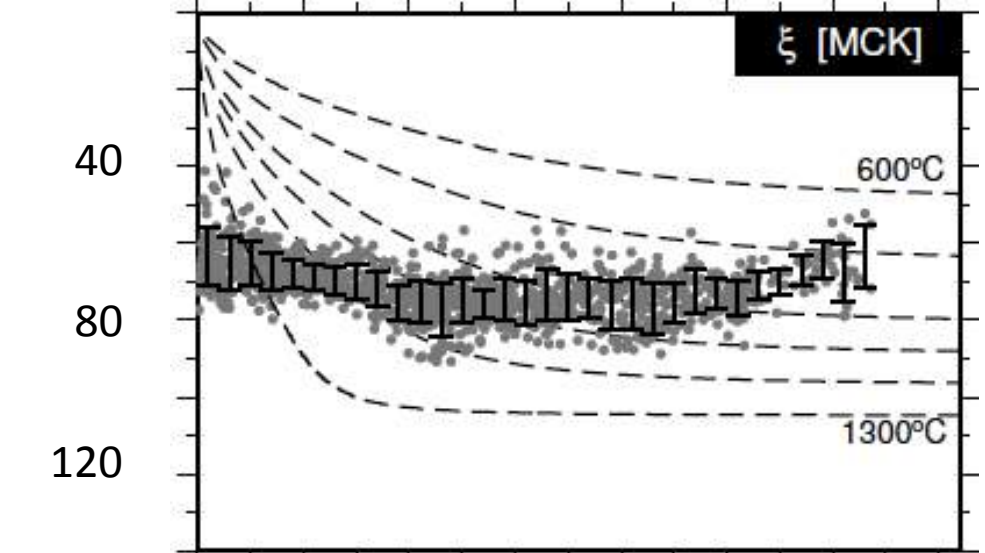
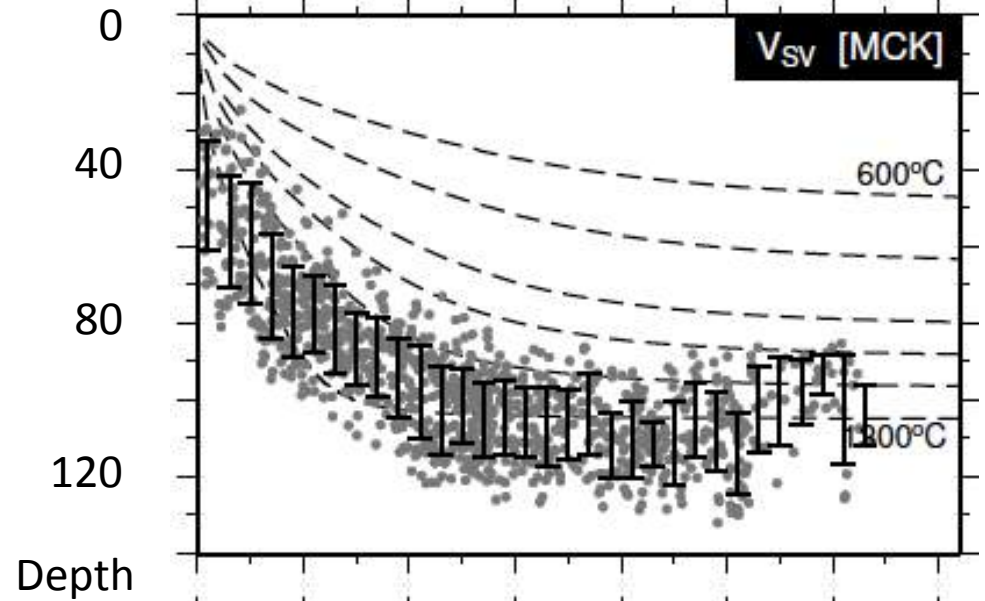
Atlantic Ocean



Age (My)

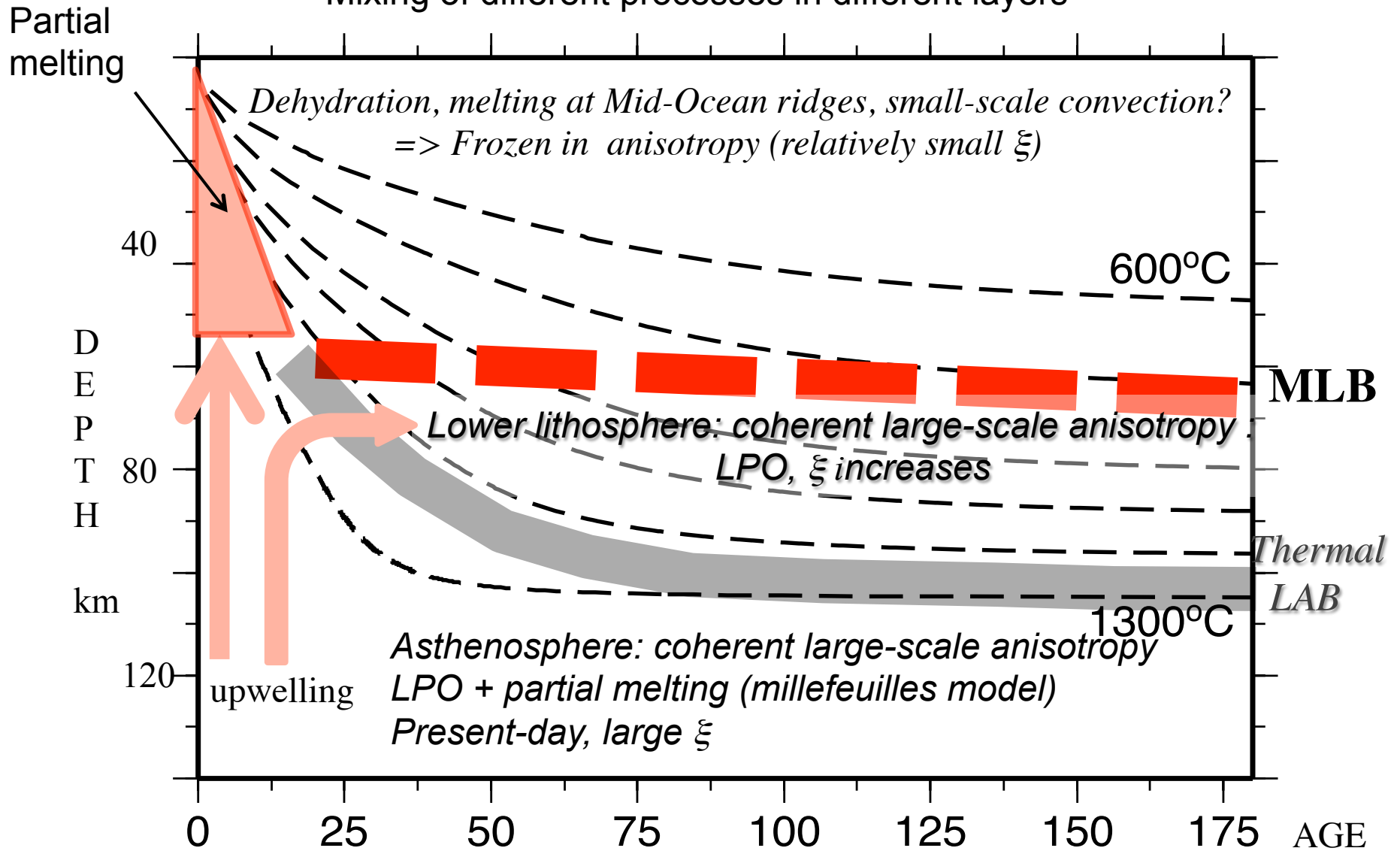
02 Dec. 2016

Indian Ocean



Age (My)

Mixing of different processes in different layers



MLB: Mid-Lithospheric Boundary

New Discontinuity within the lithosphere

-LAB topography derived from surface wave data on a global scale

-The ocean lithosphere not so simple!

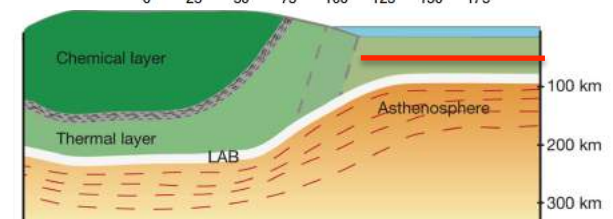
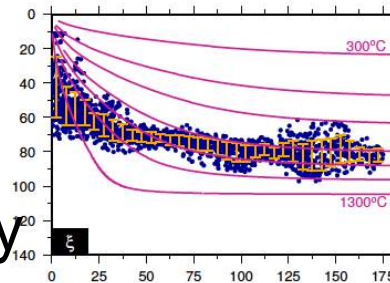
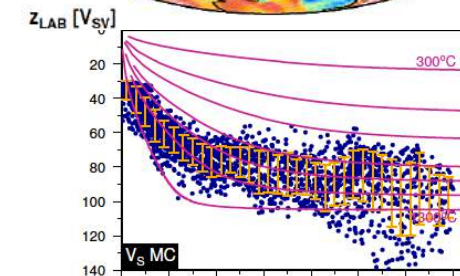
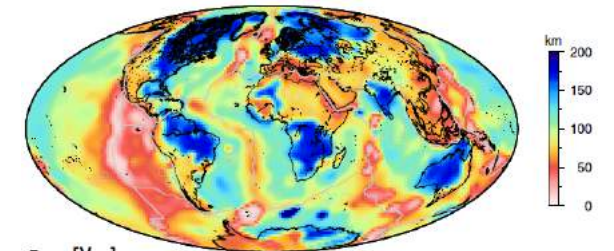
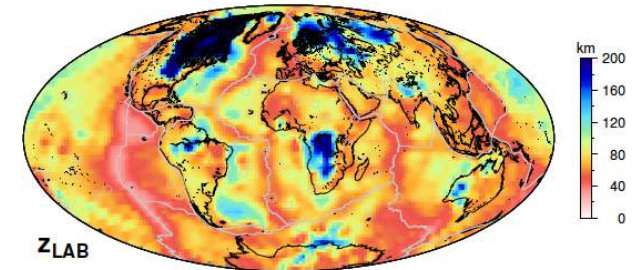
- For oceans, the model of formation of lithosphere must be revisited in view of results from radial and azimuthal anisotropies.

-Existence of a strong gradient of ξ between 60-80km (plate; related to dehydration boundary layer?)

Mid-Lithospheric Boundary

CONTINENTS?

02 Dec. 2016



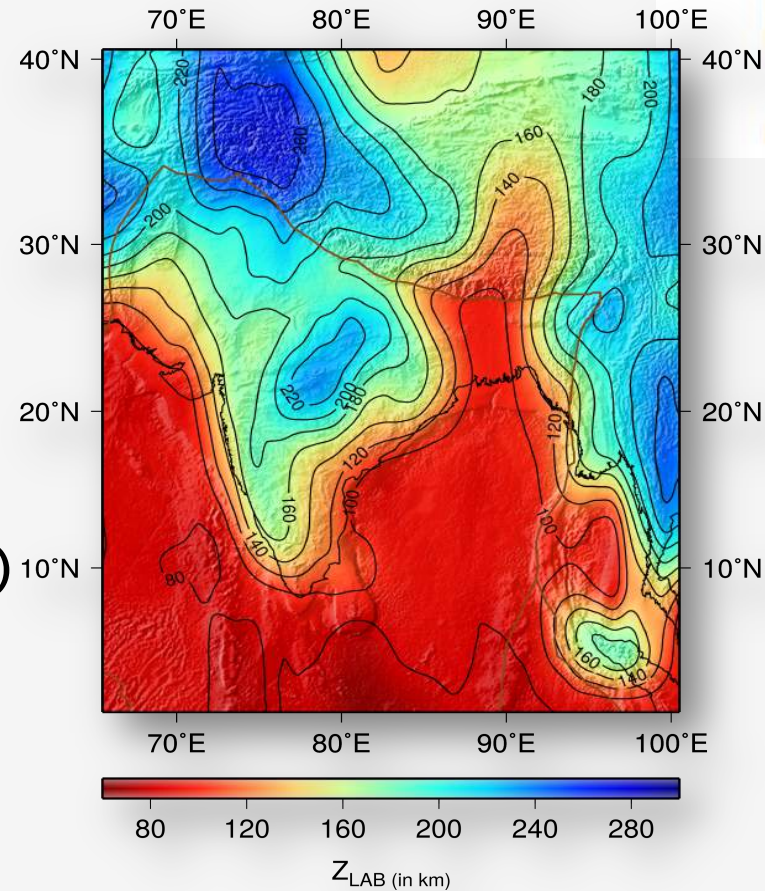
Yuan and Romanowicz, 2010

Deep Structure of the Indian Continent

1: IPG-Paris, France (J.-P. Montagner)

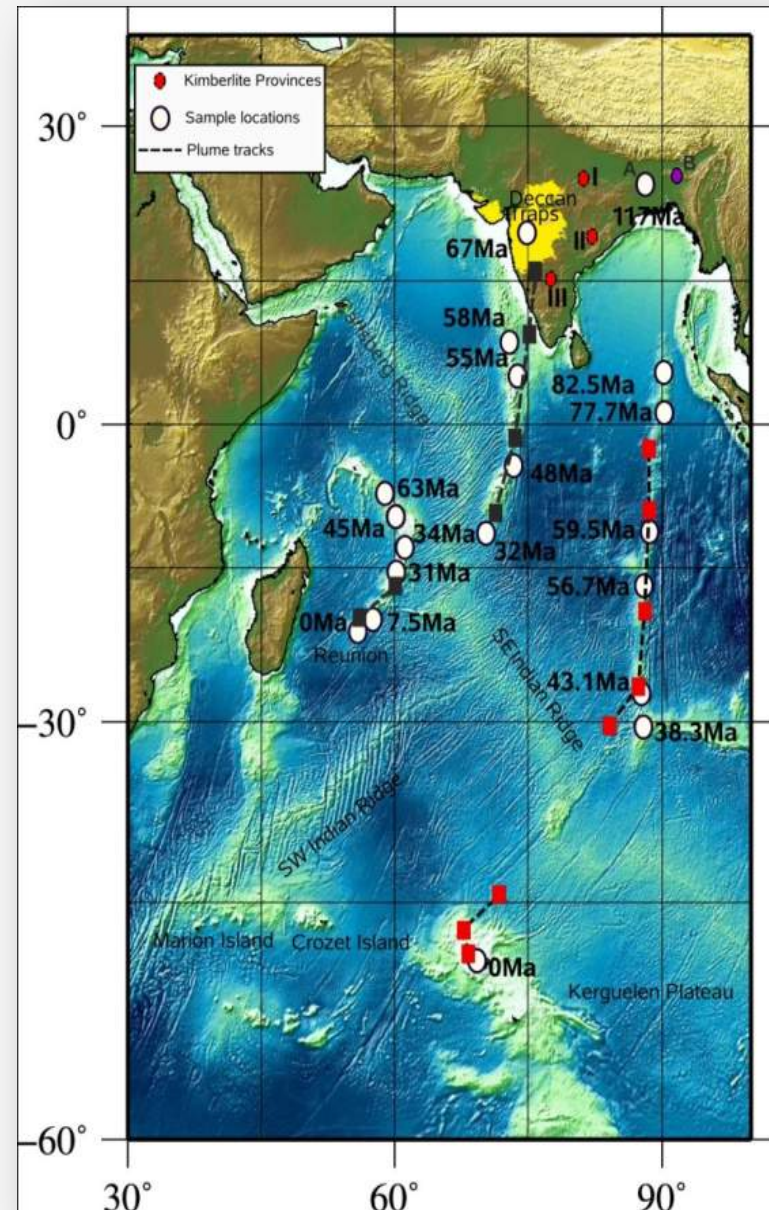


2: NGRI-Hyderabad, India (Ravi Kumar)



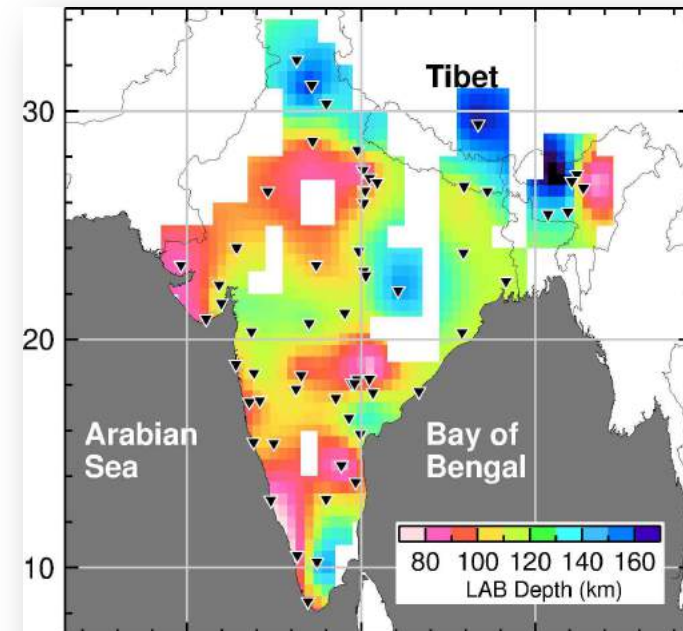
Motivation and Scientific Challenges

- Indian continent is **unique** in many respects.
- Indian plate moved at exceptionally high speeds of **18-20 cm/yr** after its breakup from Gondwanaland ~65 Myr. Ago (Paleomagnetic data).
- Ravaged by hotspots and experienced large scale magmatism.
- Five cratons. Indian continental lithosphere might be underlain by an anomalously hot sub-lithospheric mantle
- Interaction with plumes (*Réunion, Marion, Crozet and Kerguelen*)

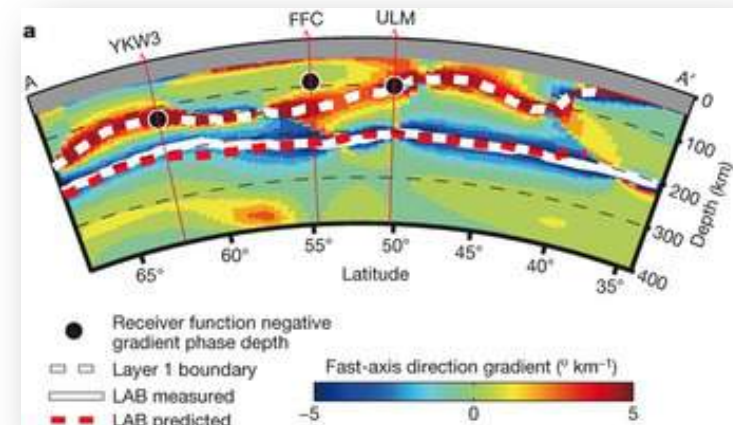


Scientific Challenges – Debate on Indian LAB

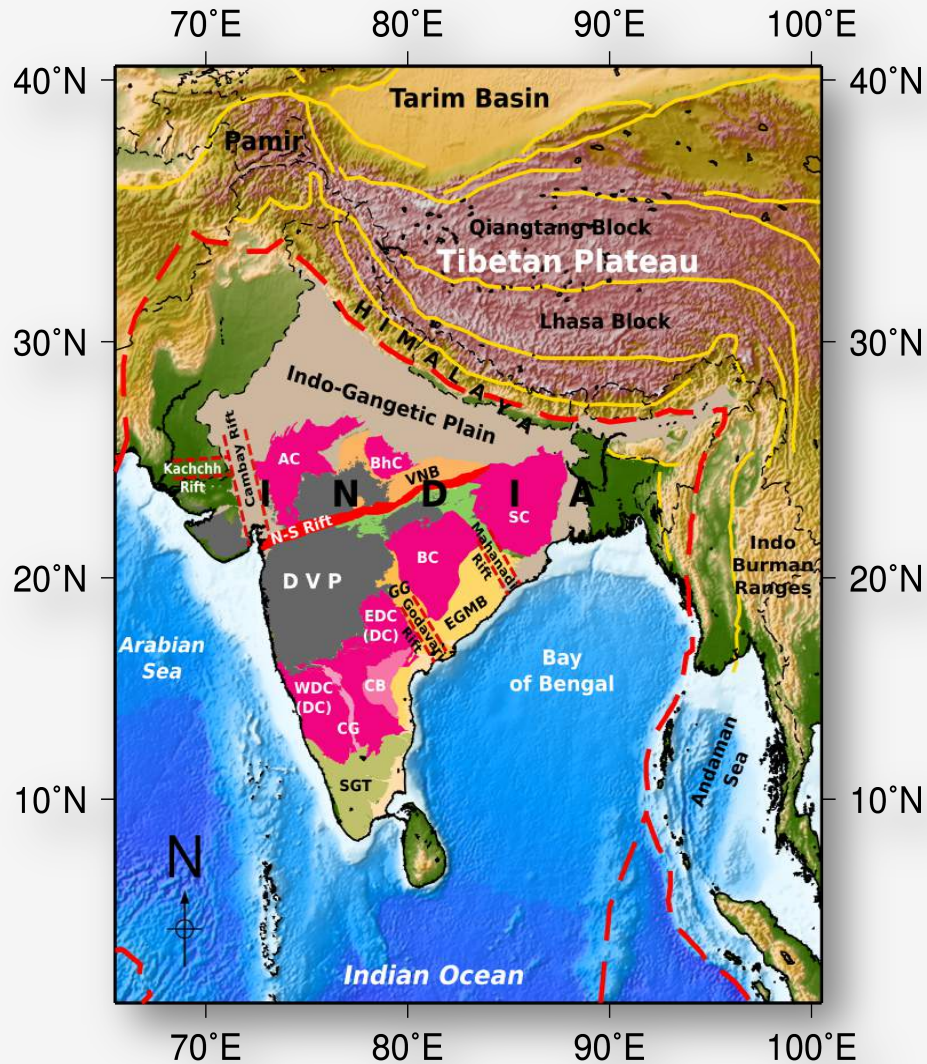
- **Super mobility** due to a thin seismic lithosphere (~100km) [Kumar et al., 2007 & Negi et al., 1986] - in total disagreement with common consensus on cratons.
- Pressure-Temperature studies on mantle xenoliths reveal a thick lithosphere that is ~**230 km** at ~1100 Ma. underneath **Dharwar Craton** [Griffin et al., 2008].
- ~**175 km** at ~ 65 Ma underneath **Bastar craton** [Babu et al. 2009].
- Evidence for postcollisional flexuring of the Indian plate with a wavelength of ~1000 km



Layering in the lithosphere in NA
[Yuan and Romanowicz, 2010]



Study Area

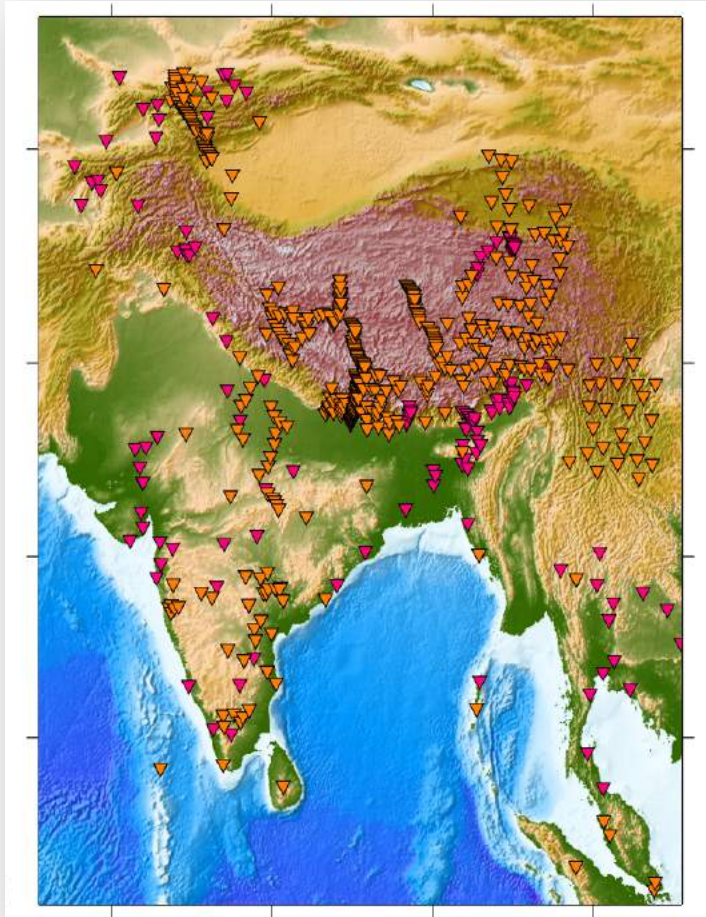


- | Precambrian | Phanerozoic |
|---------------------------------|--------------------------------|
| South Granulite Terrian (SGT) | Deccan Volcanic Province (DVP) |
| Cratons (AC, BC, BhC, DC, SC) | Gondwana Rocks (GR) |
| Closepet Granite (CG) | Indo-Gangetic Plains (IGP) |
| Cuddaph Basin (CB) | Alluvium |
| Eastern Ghat Mobile Belt (EGMB) | |
| Godavari Graben (GG) | |
| Vindhyan Basin (VNB) | |

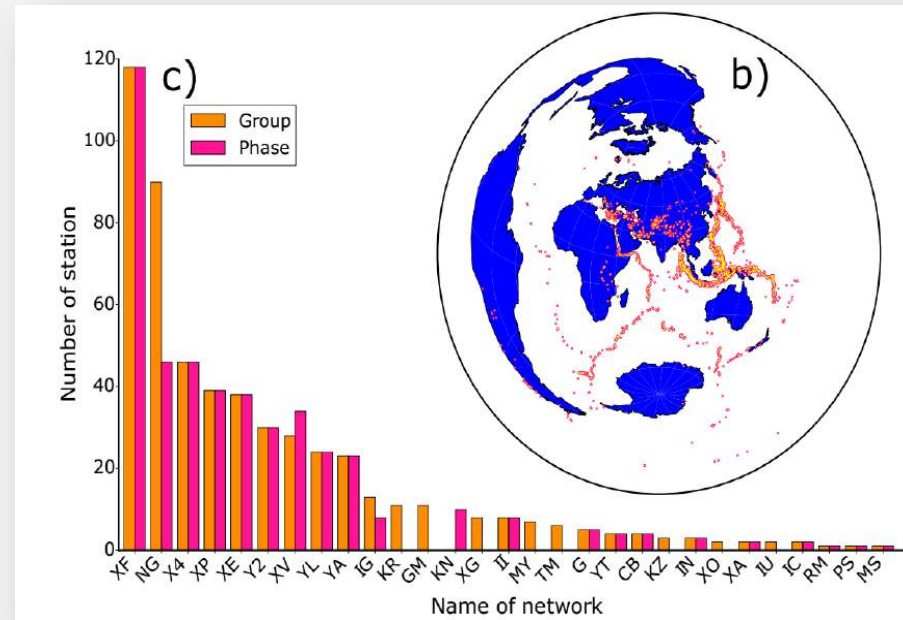
AC: Aravalli craton
 BC: Bastar craton
 BhC: Bundhelkhand craton
 DC: Dharwar craton
 SC: Singhbhum craton

Data

Stations



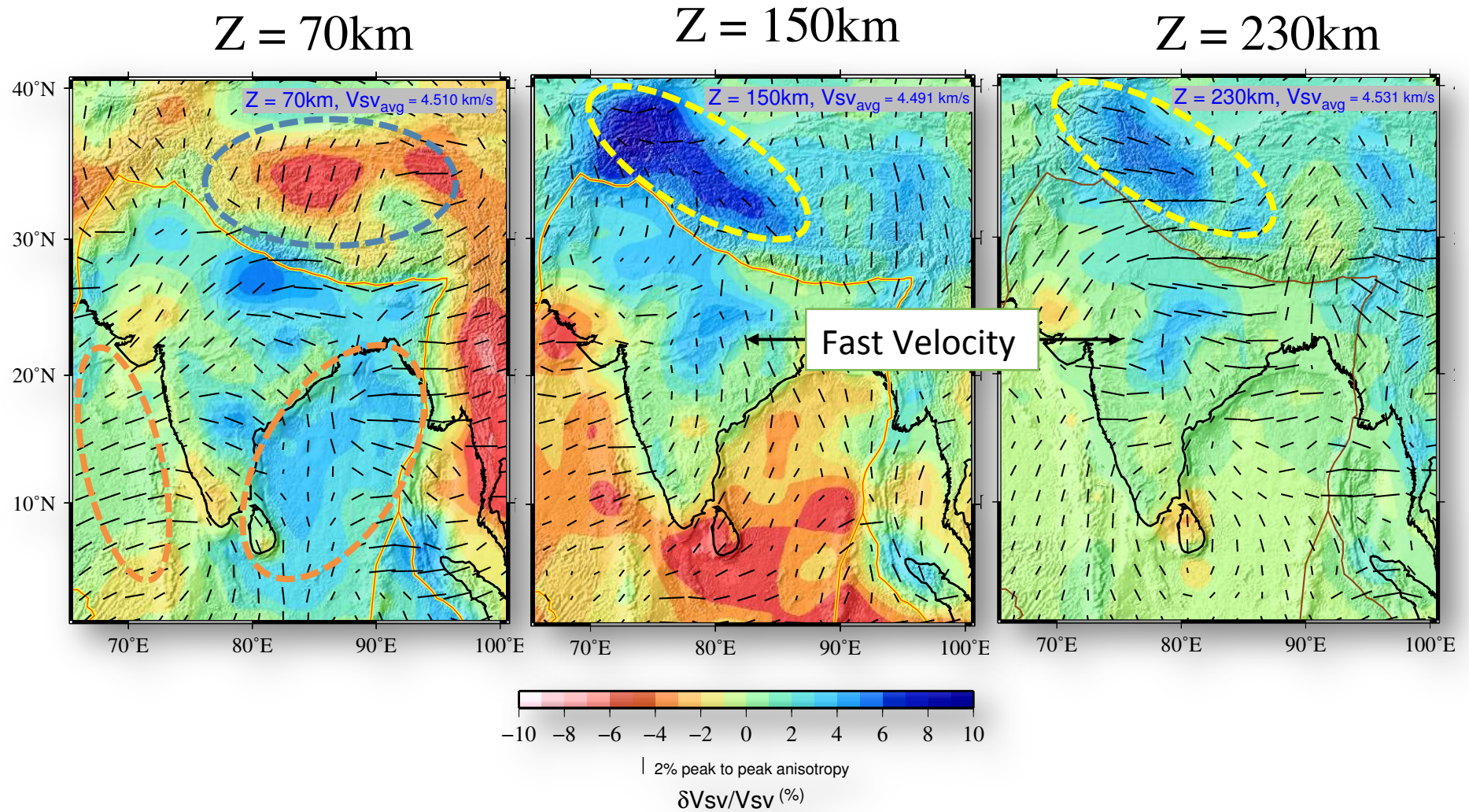
Earthquakes



- 29 Seismic broadband Networks (**global and regional**)
- Over 550 seismic stations
- Earthquakes of magnitude >5.5
- Surface wave data in the period range of 10-400S.

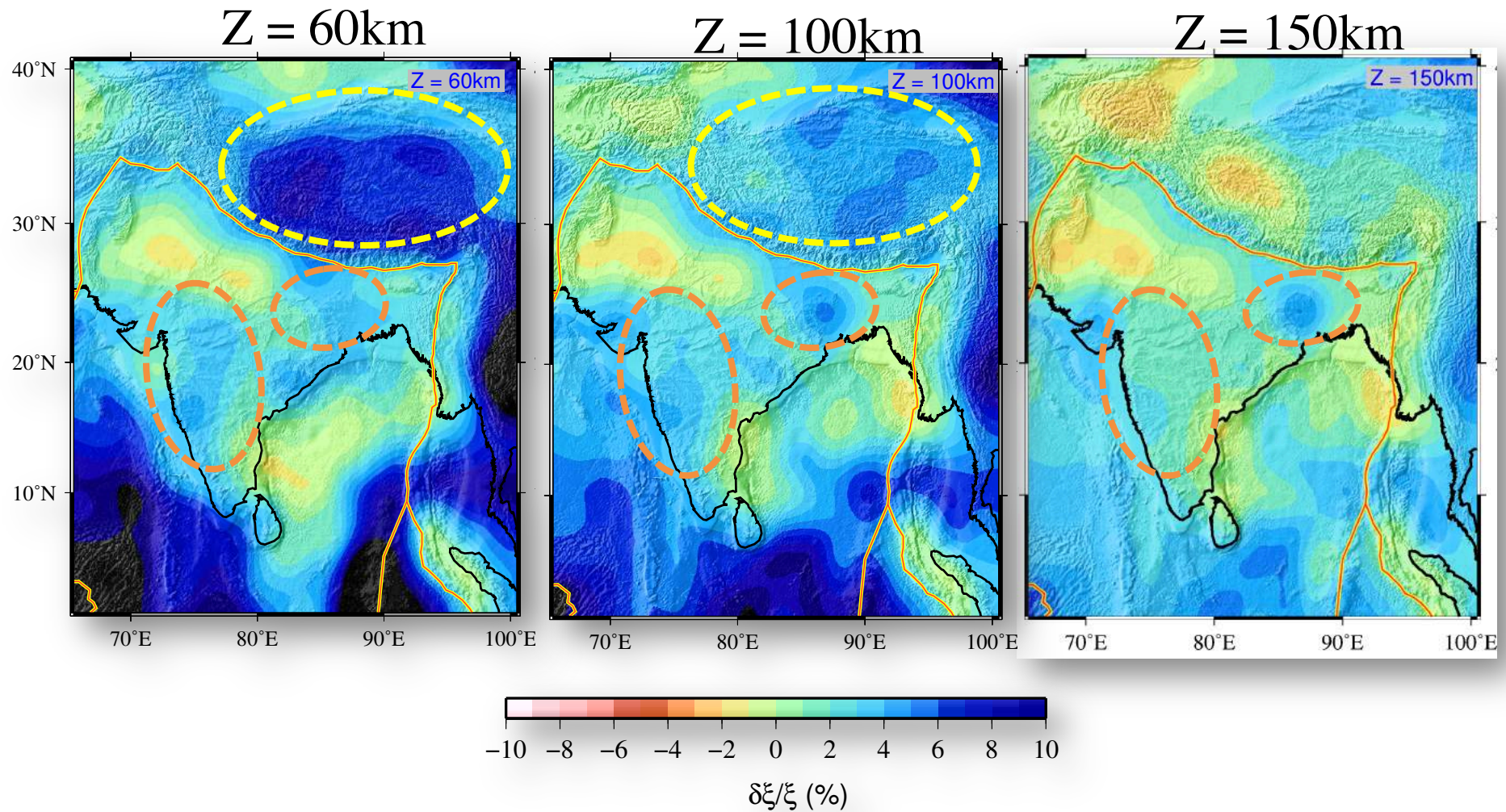
3-D tomography model of the Indian continent

Velocity and Azimuthal Anisotropy

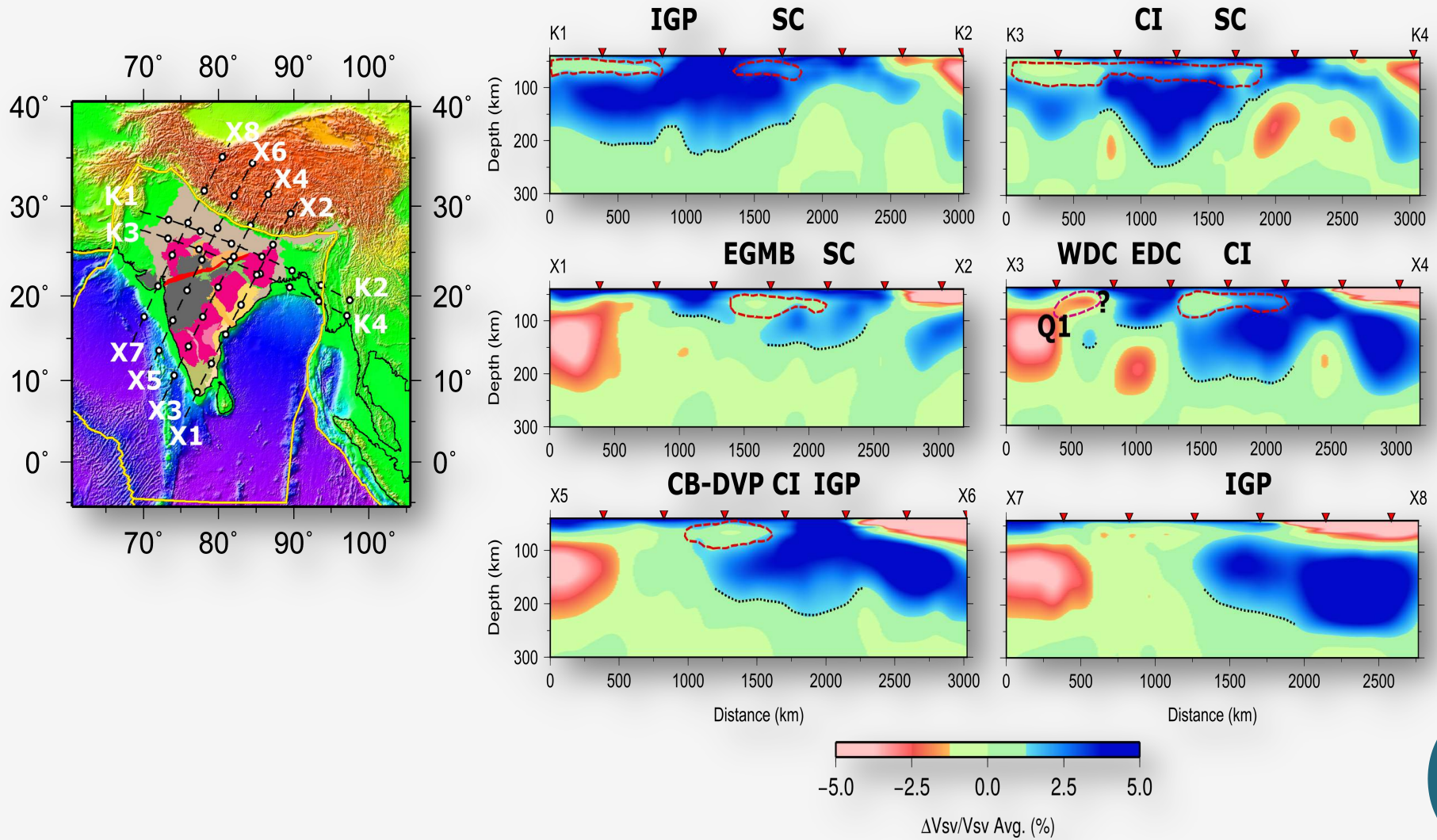


3-D tomography model of the Indian continent

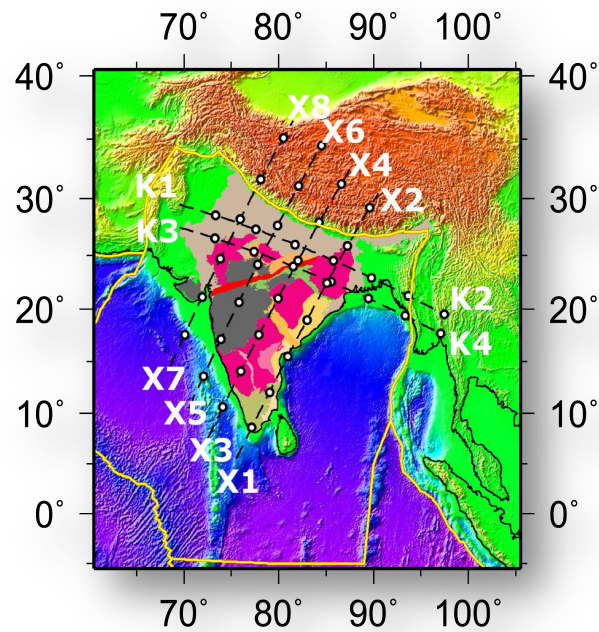
Radial Anisotropy ξ



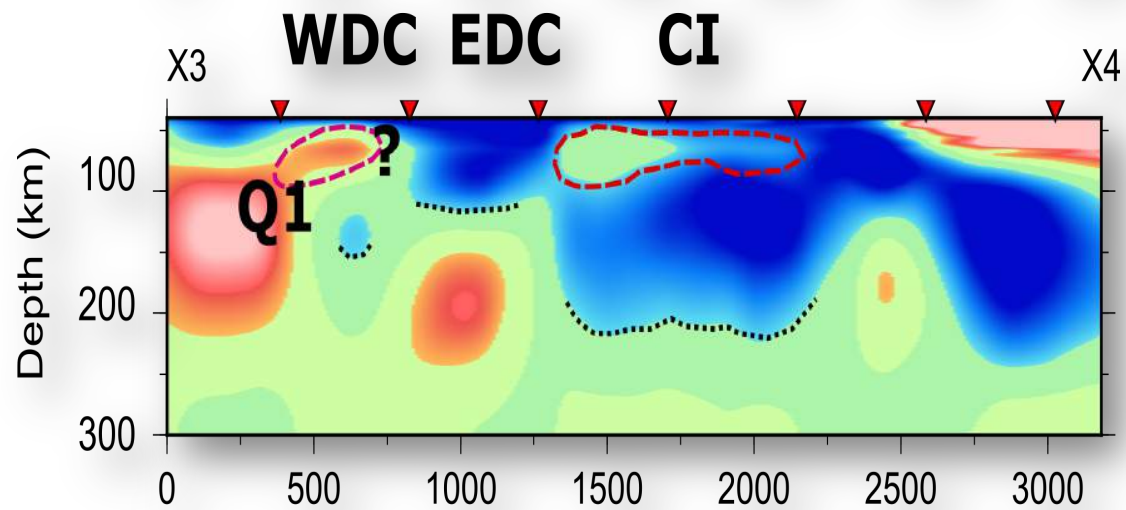
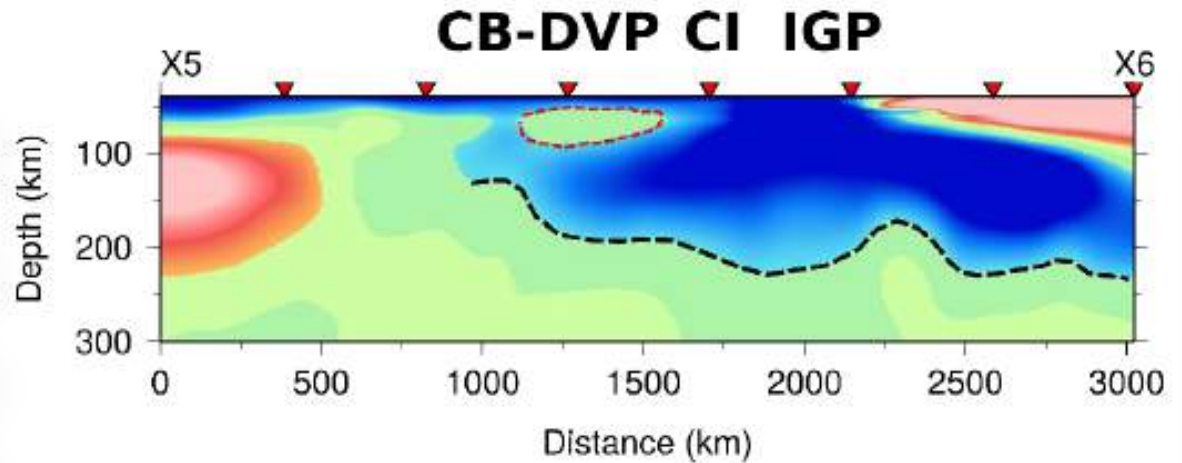
3D-Perturbation model



3D-Perturbation model



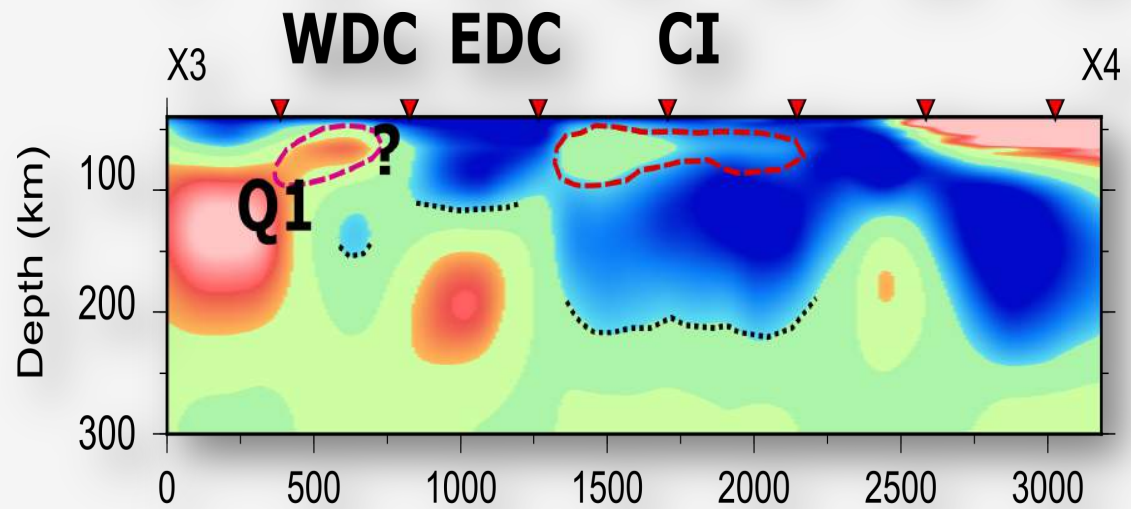
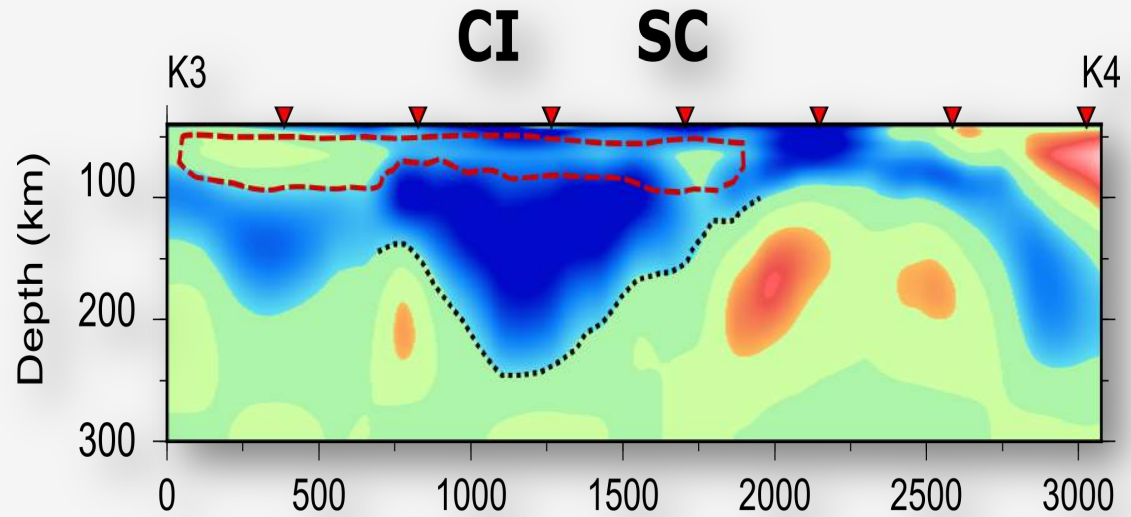
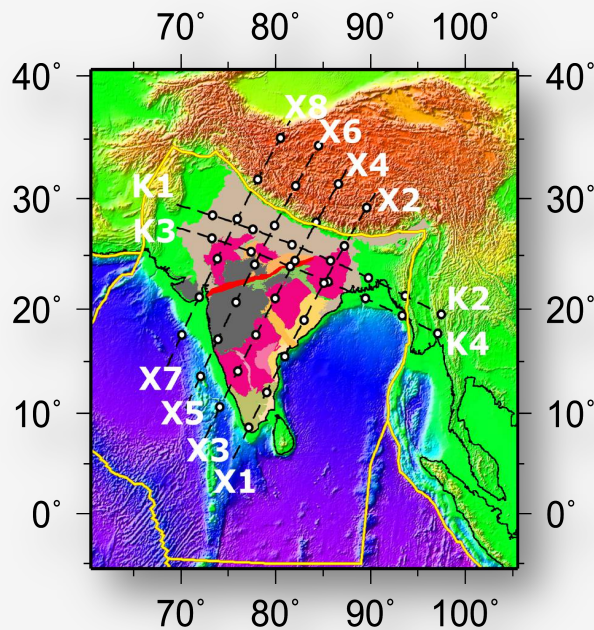
DVP: Low velocity zone
Remnant of hotspot birth



MLB: Mid-lithospheric low velocity zone

3D-Perturbation model

Indian Keel



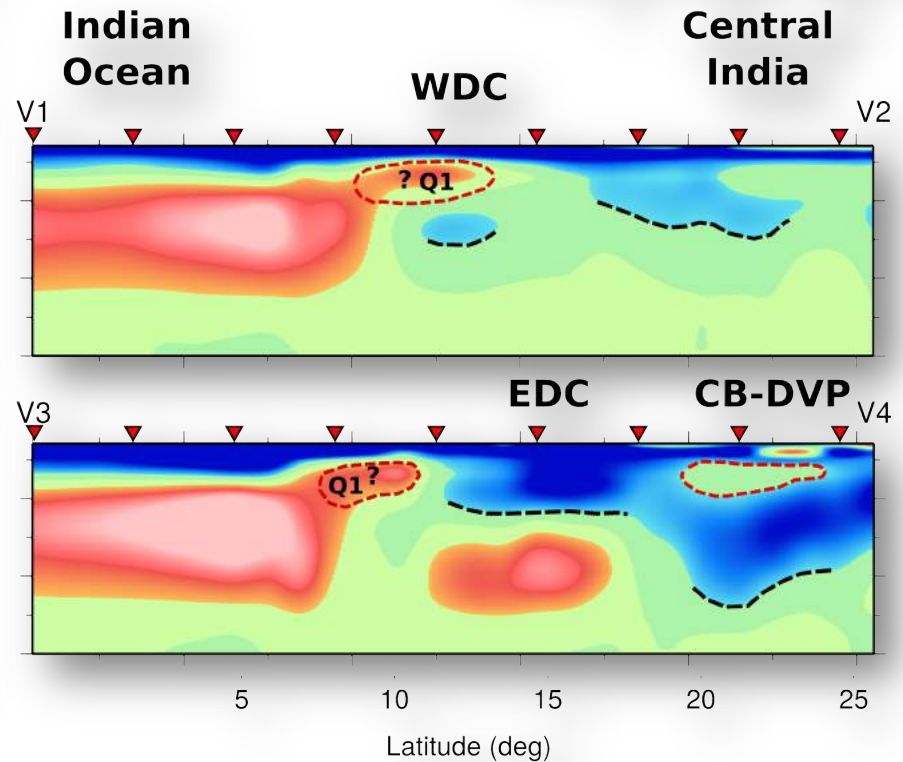
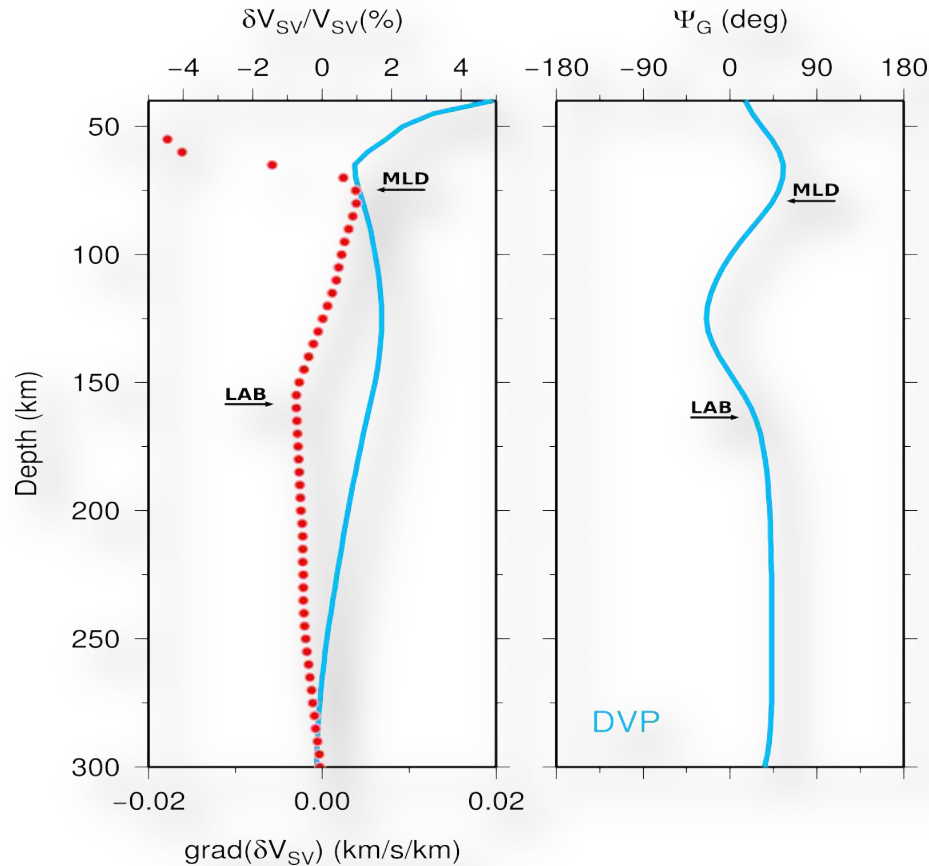
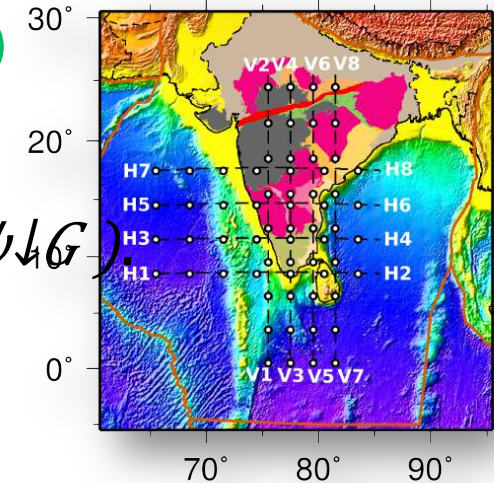
02 Dec. 2016

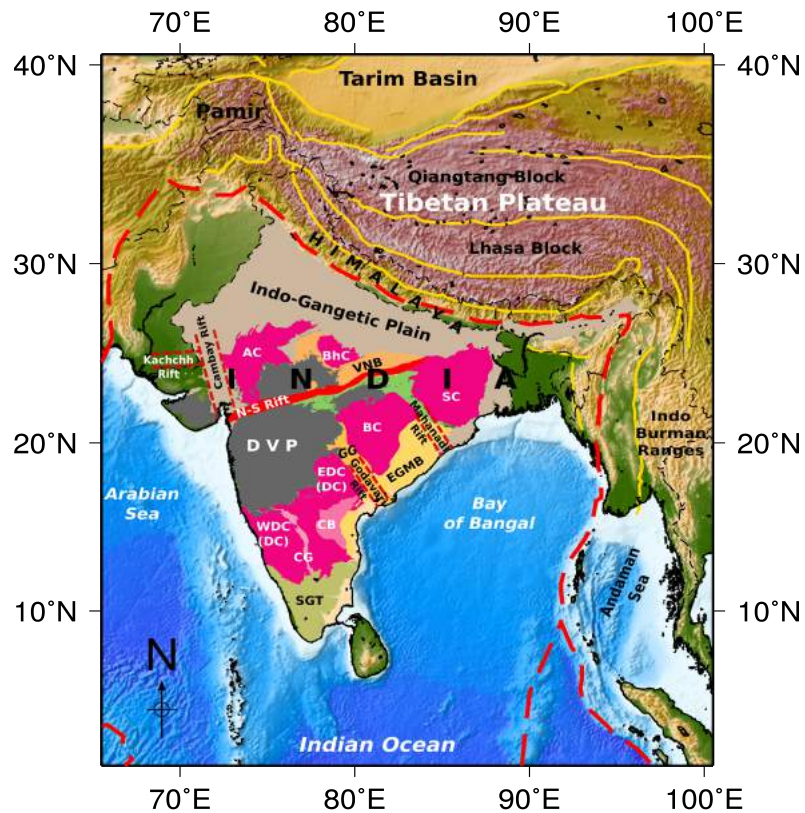
MLB: Mid-lithospheric low velocity zone

Discussion: Central India (CI)

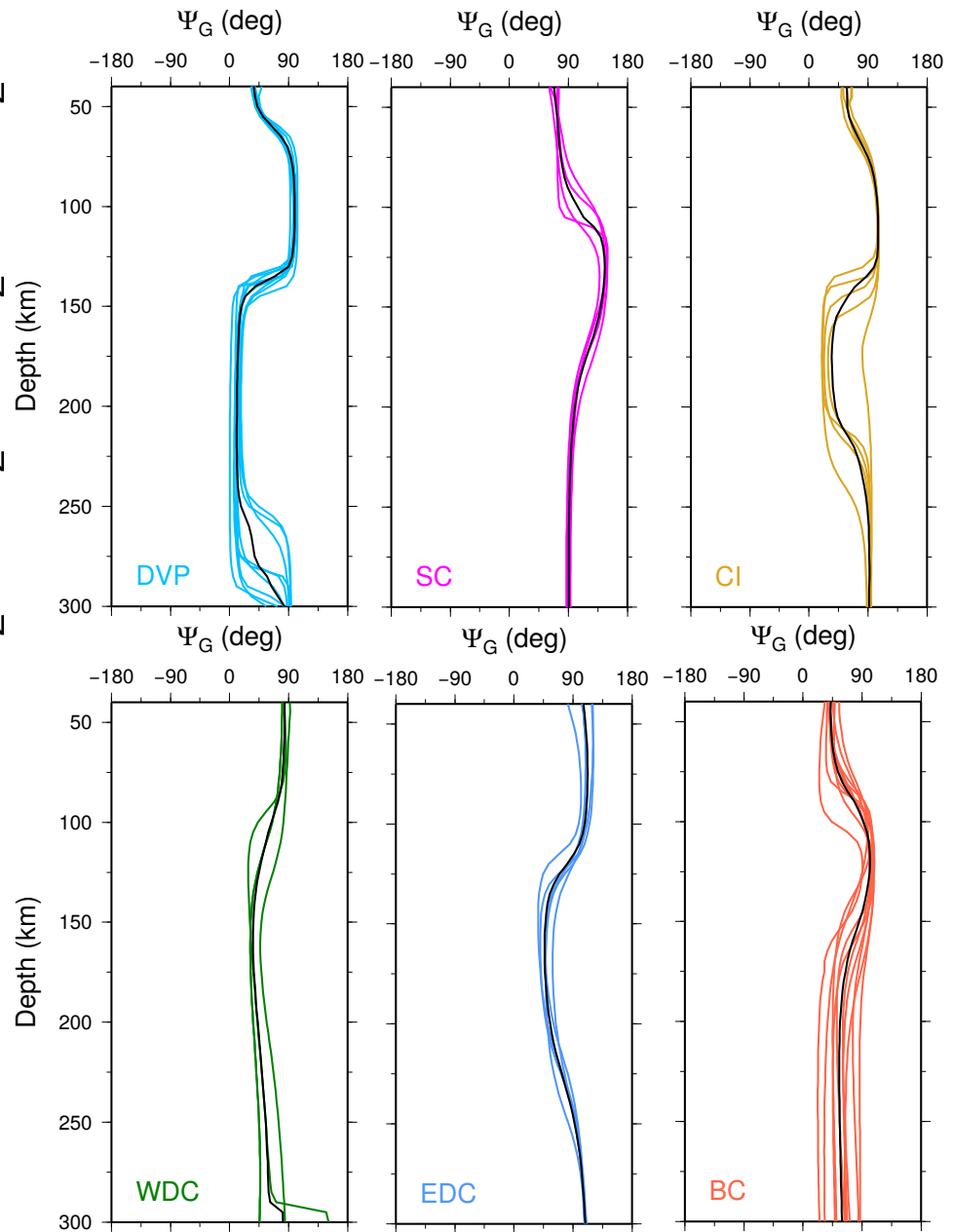
- Small velocity drop beneath DVP.
- Good agreement in both parameters ($V \downarrow SV, \psi \downarrow G$)
- Slow velocity present in the lithosphere.

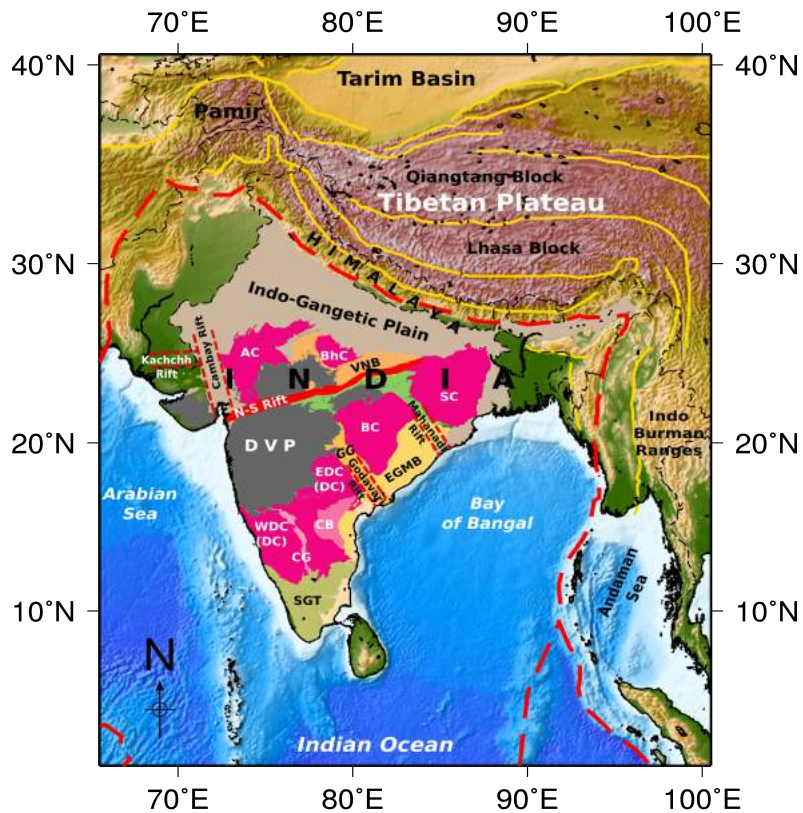
Deccan Volcanic Province (DVP)



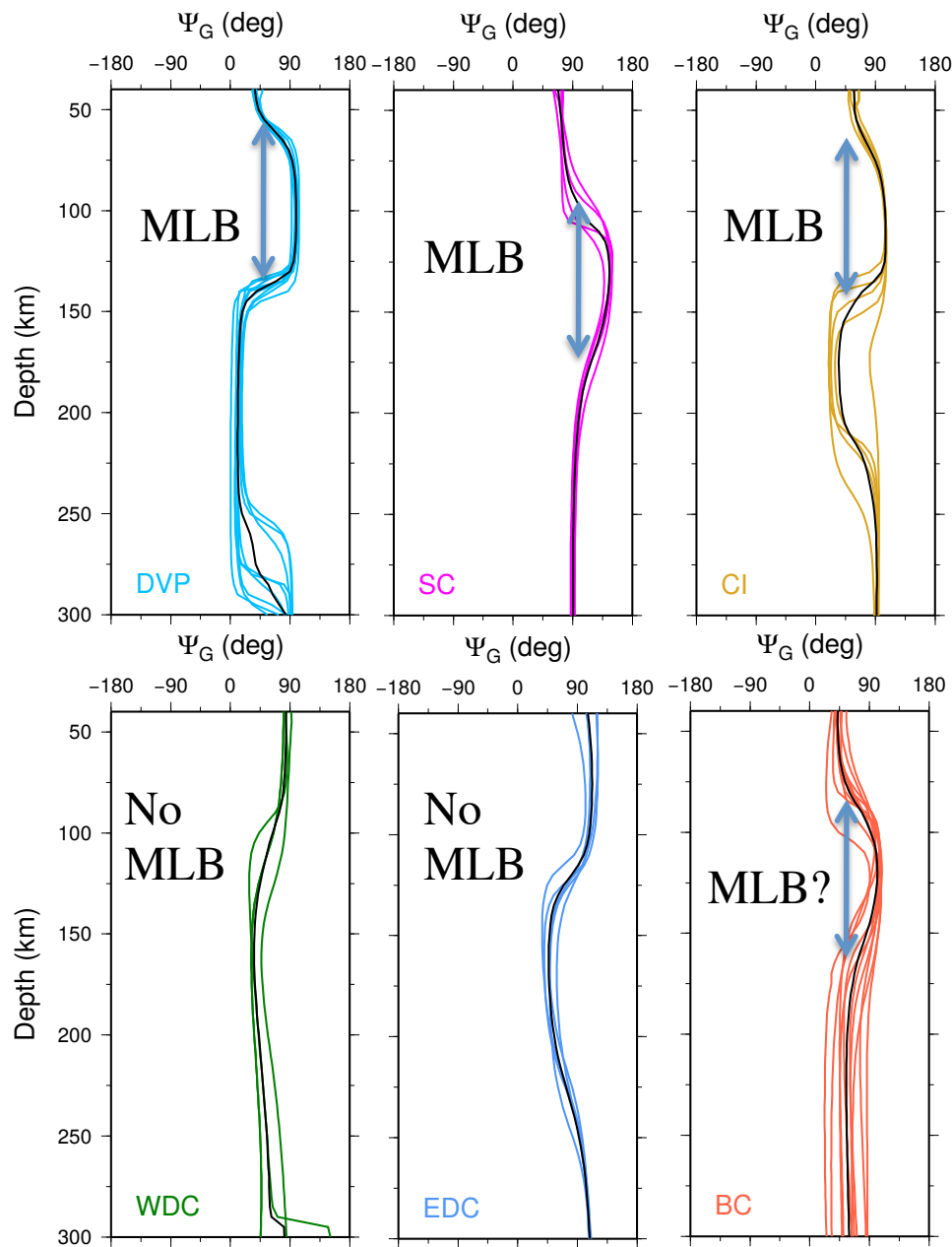


- | Precambrian | | Phanerozoic | |
|--|---|---|---|
| ■ South Granulite Terrian (SGT) | ■ Deccan Volcanic Province (DVP) | ■ Gondwana Rocks (GR) | ■ Indo-Gangetic Plains (IGP) |
| ■ Cratons (AC,BC,BhC,DC,SC) | ■ Gondwana Rocks (GR) | ■ Indo-Gangetic Plains (IGP) | ■ Alluvium |
| ■ Closepet Granite(CG) | ■ Indo-Gangetic Plains (IGP) | ■ Alluvium | |
| ■ Cuddaph Basin (CB) | | | |
| ■ Eastern Ghat Mobile Belt (EGMB) | | | |
| ■ Godavari Graben (GG) | | | |
| ■ Vindhyan Basin (VNB) | | | |

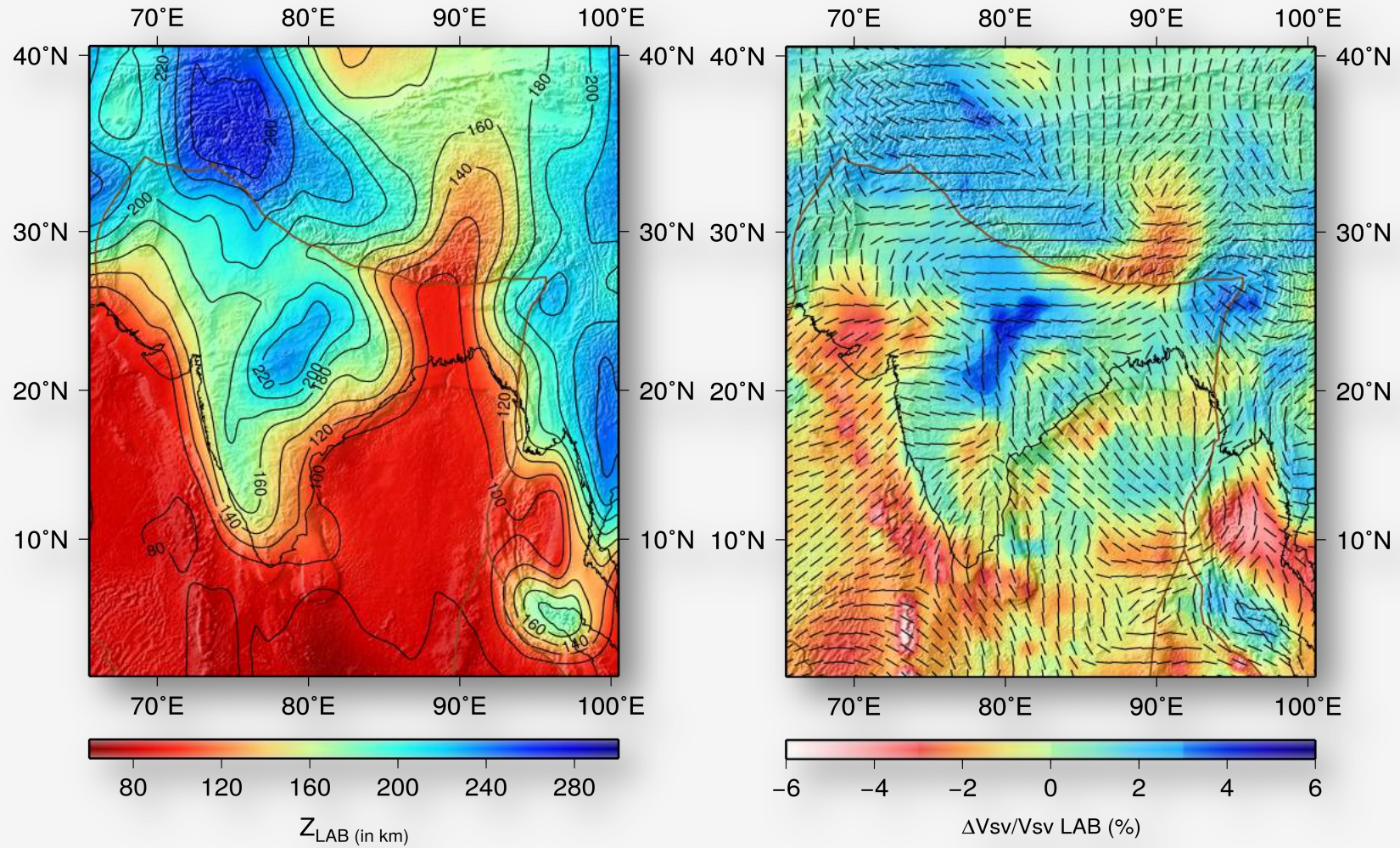




MLB = ML-LVZ



Indian Plate LAB (Lithosphere-Asthenosphere Boundary)

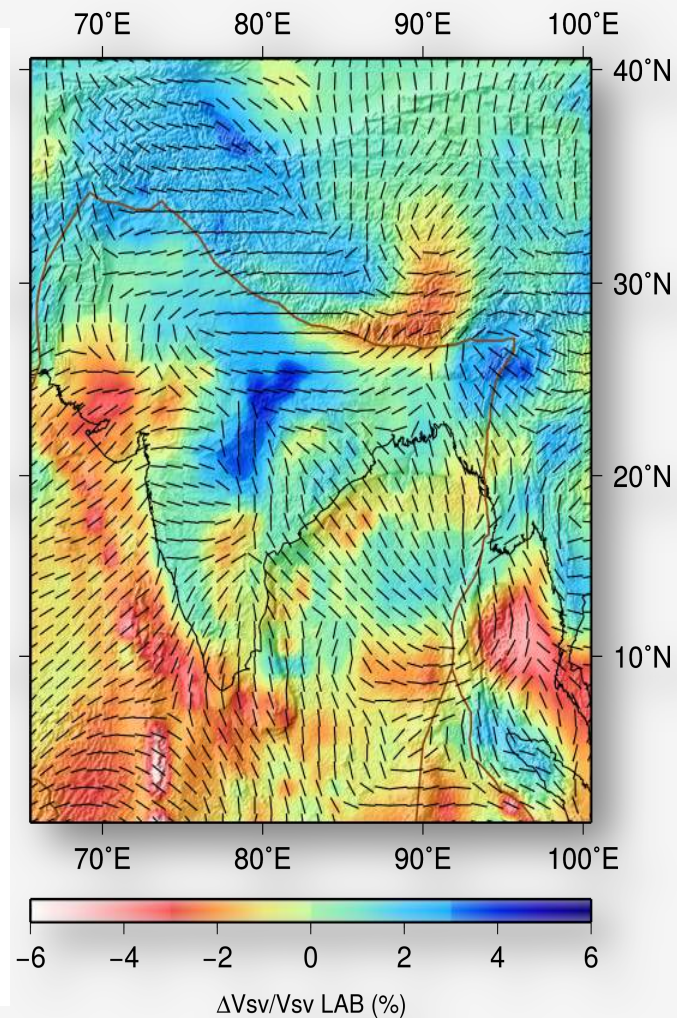


02 Dec. 2016

Indian Plate LAB: Keel

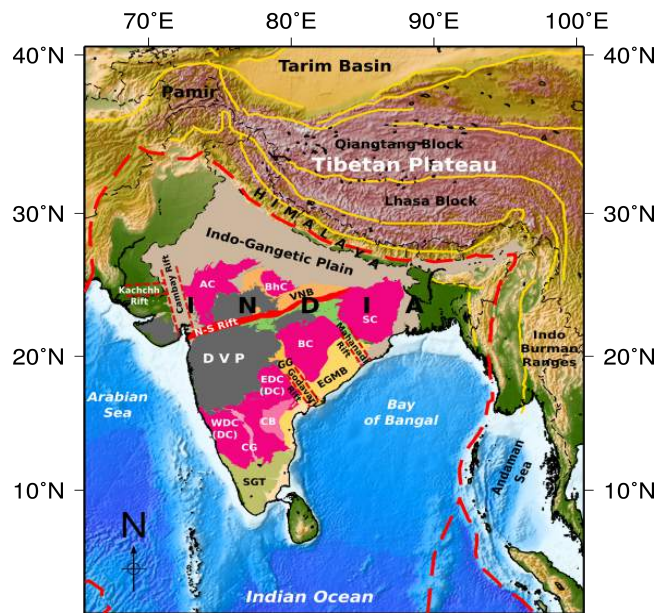


←LAB (in km)

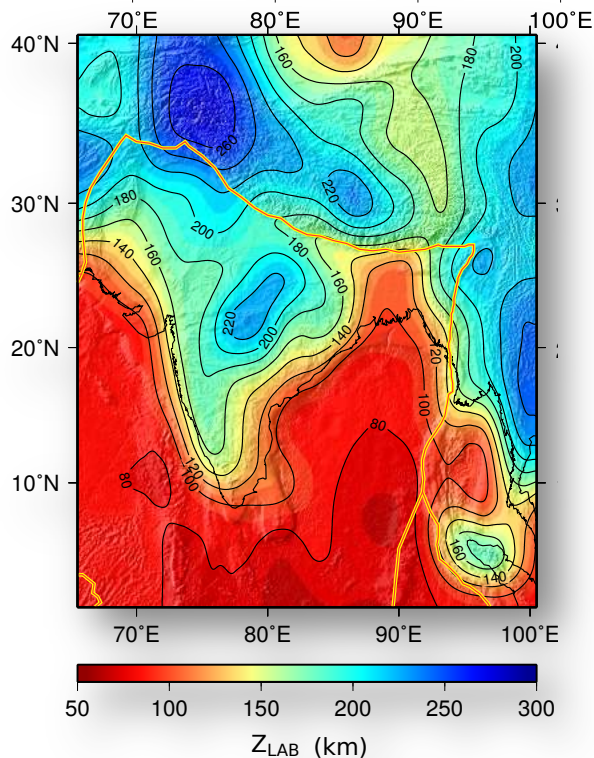


Geodynamic Role?
Plume influence?

02 Dec. 2016



- Large variability of craton thicknesses
- MLB (ML-LVZ): low velocity zone
- MLB: Change in azimuthal anisotropy
- MLB not present in all blocks



- DVP (Deccan Volcanic Province)
MLB: memory of La Réunion
Hotspot birth
- Indian Keel: geodynamic role?
- Continents much more complicated than oceans

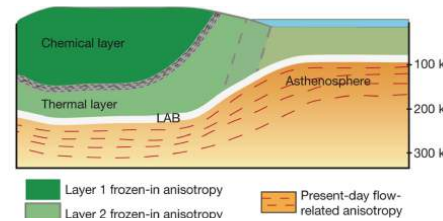
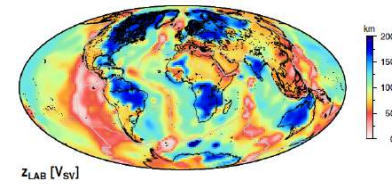
Conclusions

- Seismic Anisotropy can be mapped in different depth ranges
- Interpretation of seismic anisotropy is non-unique (intrinsic L.P.O. versus extrinsic anisotropy)
- Imaging of geological objects such as LAB (Lithosphere-Asthenosphere Boundary)
- New findings from anisotropy :

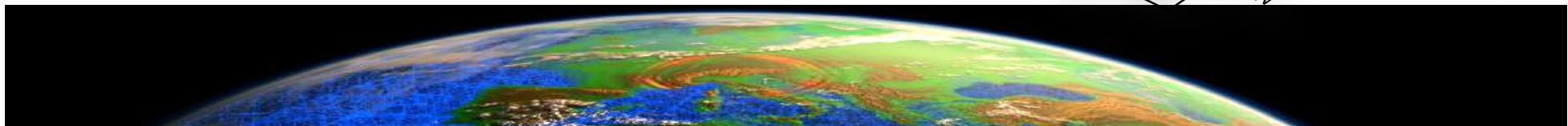
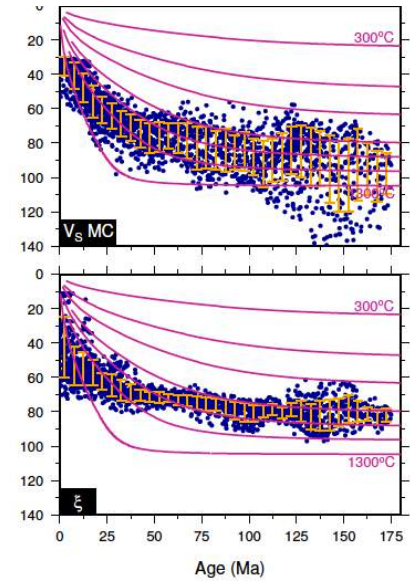
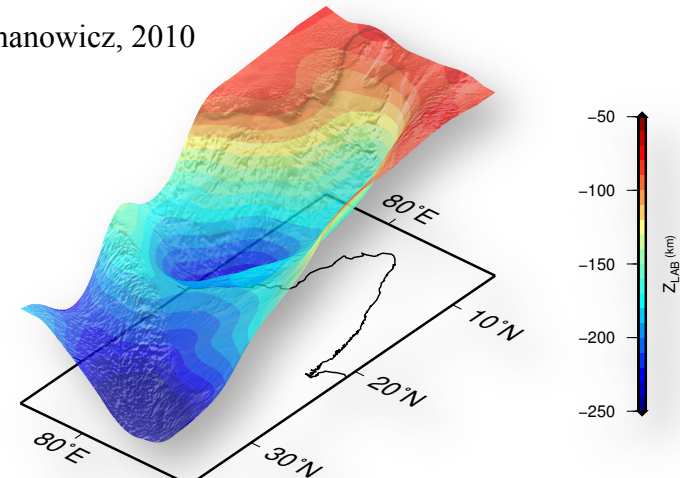
Oceans: MLB

Continents: variable thickness, Mid-LVZ
Keel

-

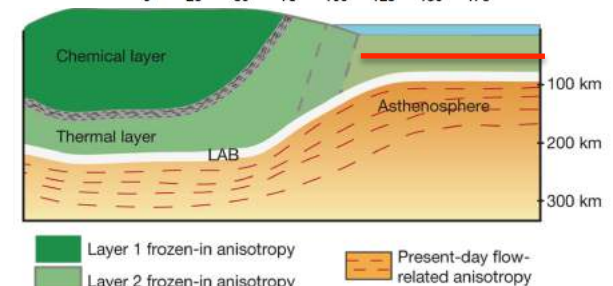
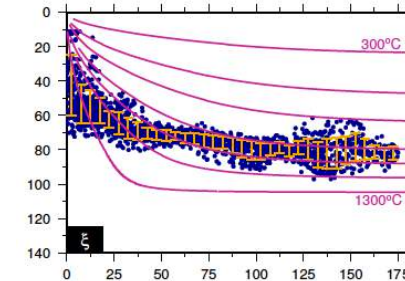
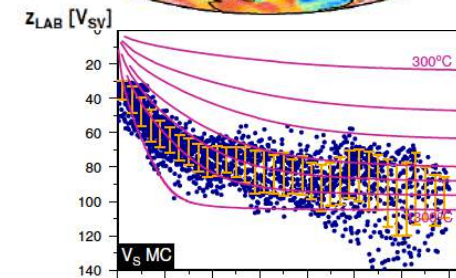
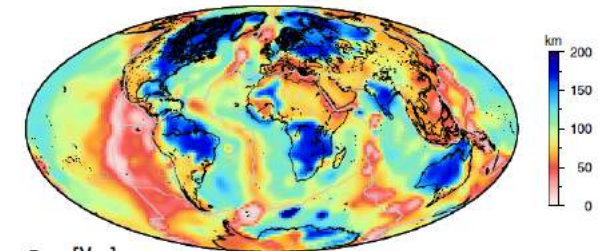
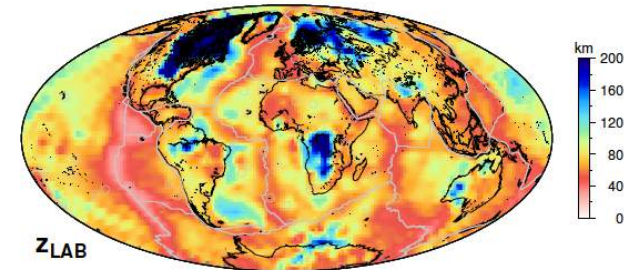


Yuan and Romanowicz, 2010



Oceanic and Continental plates

- LAB topography derived from surface wave data on a global scale (for oceans) and regional scale (India)
- For oceans: The ocean lithosphere not so simple! Stratification, existence of a Mid-Lithospheric Discontinuity (MLD: strong gradient of ξ between 60-80km)
- For continents: Large variability of craton thickness Stratification: MLB, low velocity zone (not present in all blocks). Relationship with MLD?
- The model of formation of lithosphere must be revisited in view of results from radial and azimuthal anisotropies in oceans and continents.



Yuan and Romanowicz, 2010

3-D Indian LAB model

