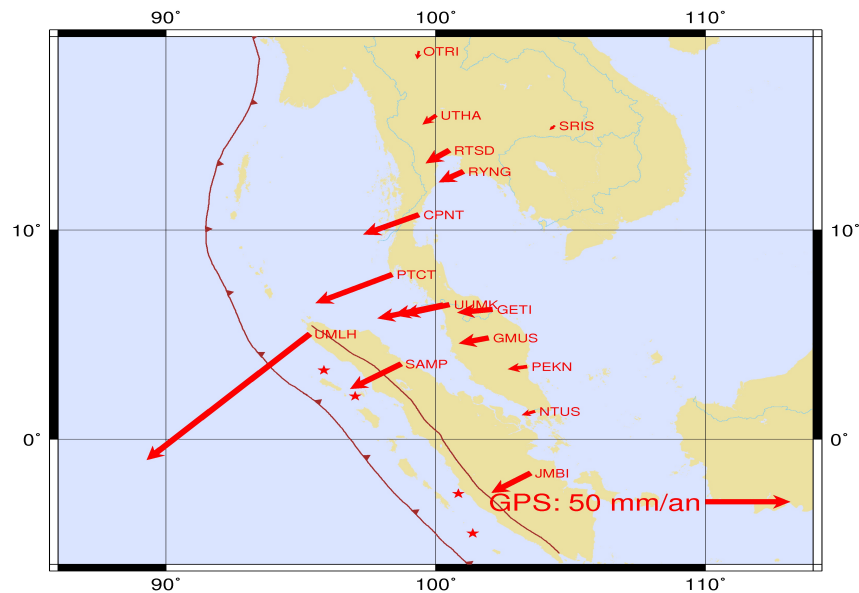


**Deformations associated with megathrust earthquakes:  
rheology of the asthenosphere and of the subduction interface**

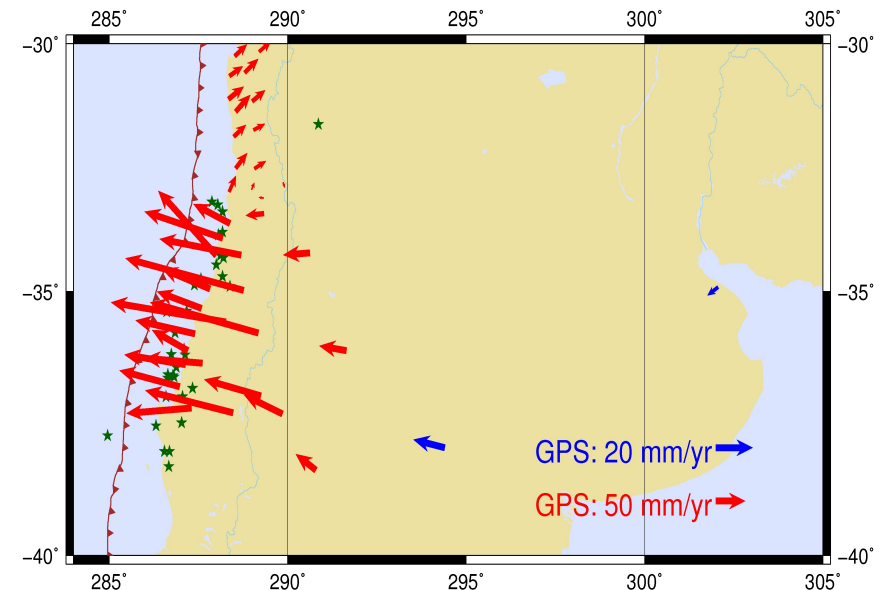
**L. Fleitout, J.D. Garaud, O. Trubienko, E. Klein, and numerous 'GPS' colleagues (C. Vigny...)**

**Collège de France 1-12-2016**

## Huge velocity perturbations are observed more than 1000km away from the earthquake



Velocity perturbations 4yrs after  
Aceh (dec 2004)

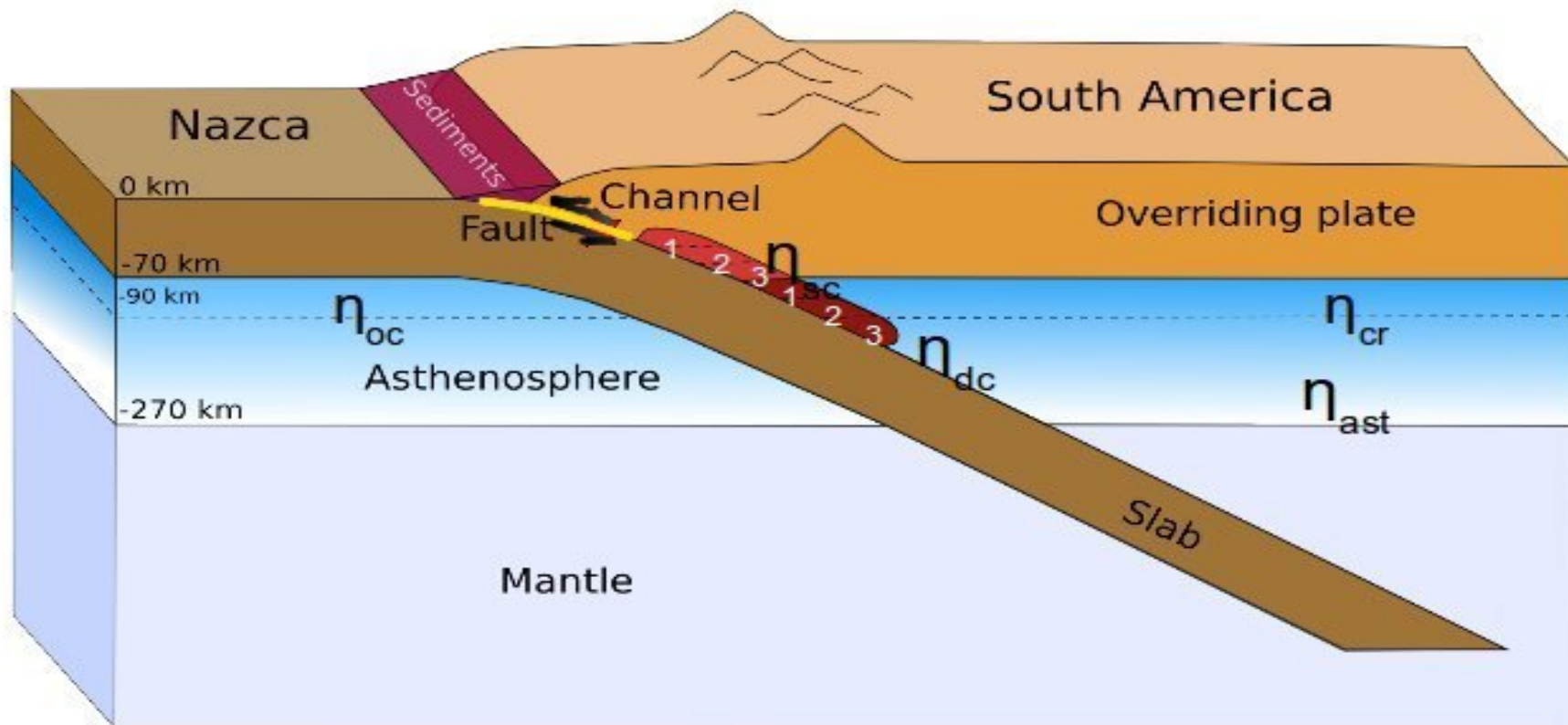


1 yr after Maule (febr 2010)

The three megaeearthquakes (Aceh 2004, Maule 2010, Tohoku 2011) are the first earthquakes of magnitude around 9 since modern techniques to monitor deformation are available. They are associated with far-field, long-lasting velocity perturbations

## The response of the Earth to megaequakes: A large-scale 'experiment'

- To test the Earth's mechanical properties (asthenosphere, subduction channel)
- To explore the 'time dependence' of stress/strains



# Some of the questions that we can tackle with postseismic deformations after megathrust earthquakes

Is there a low-viscosity asthenosphere?

What rheology (Newtonian, non-linear creep), What thickness?

What occurs below the brittle part of the subduction interface?

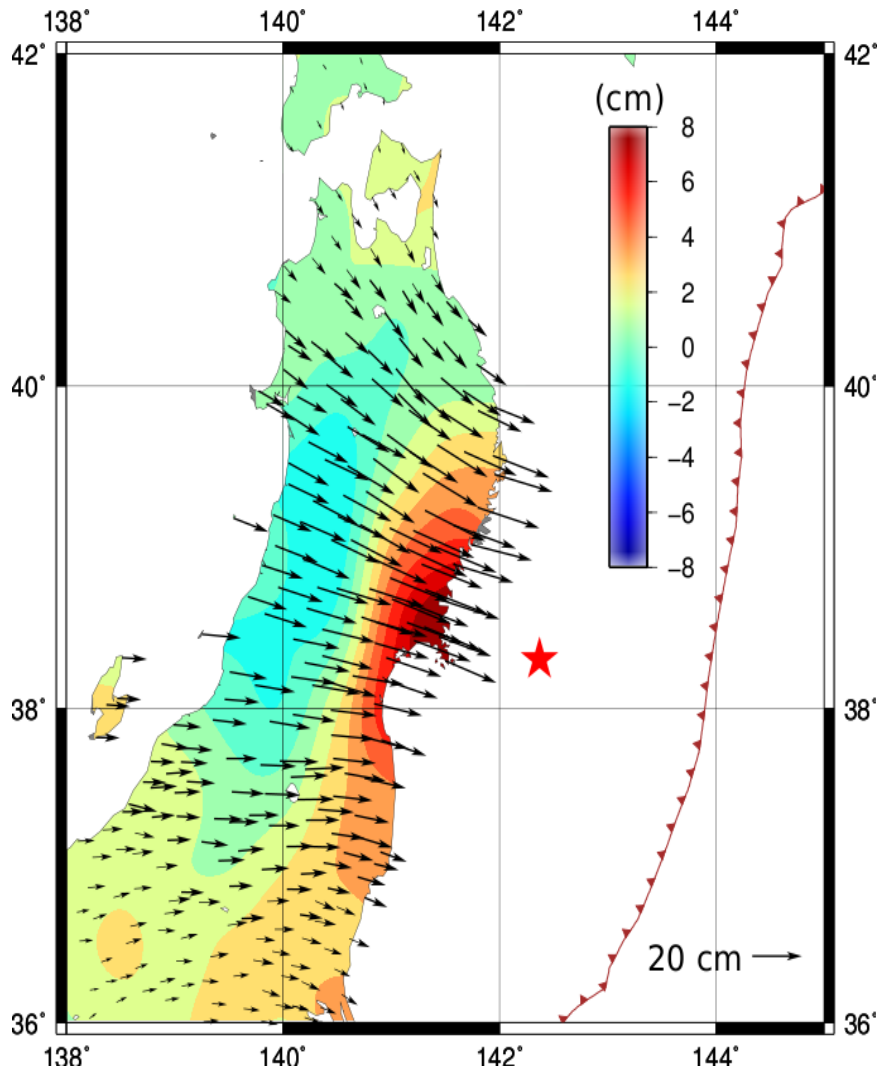
Localization of the deformation in low-viscosity ductile shear zones?

This is not a new problem (Elsasser, Pollitz and coauthors...)

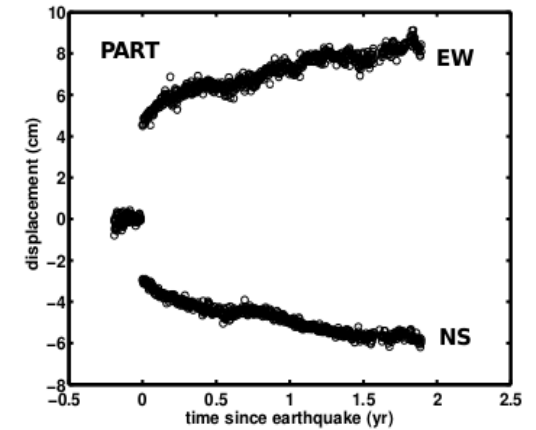
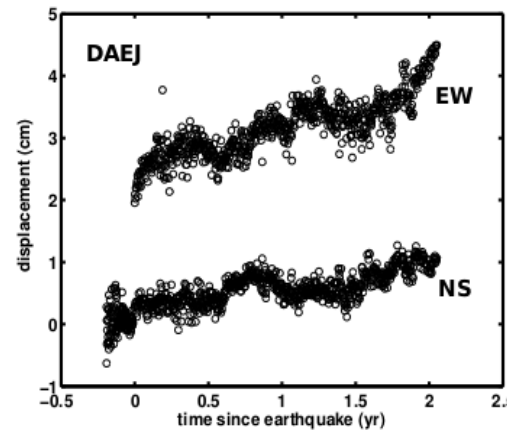
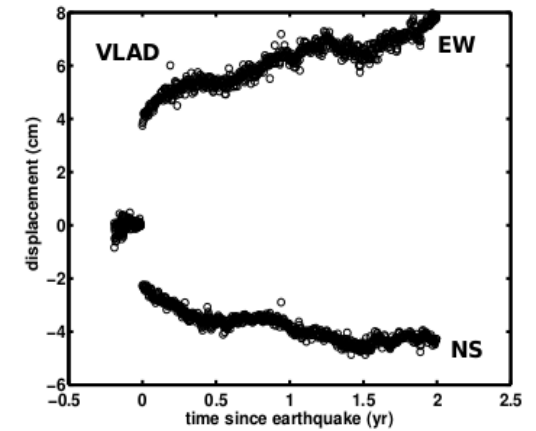
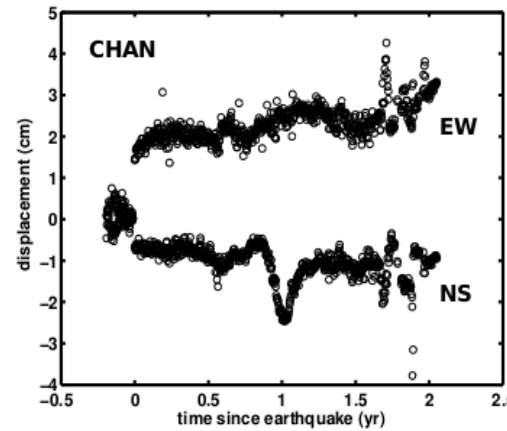
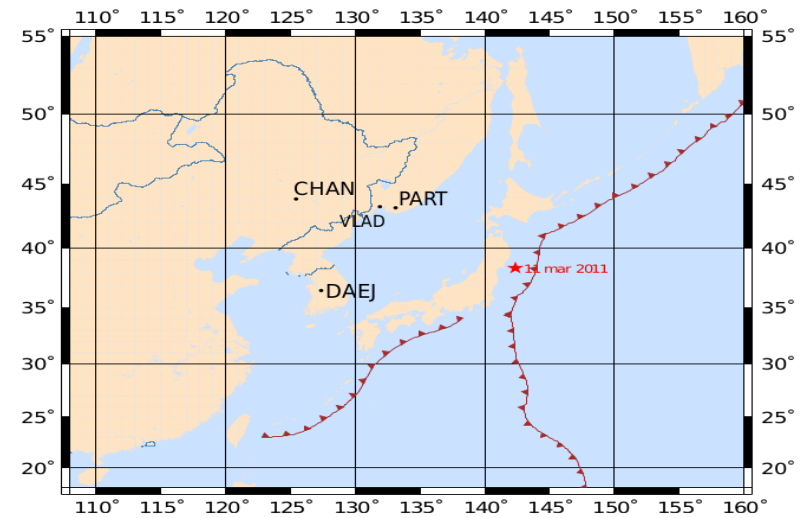
But we have now precise data and the numerical ability to treat the problem in 3D

Note: The rheology obtained here will be the rheology for a time-scale of a few years. Long term rheology?

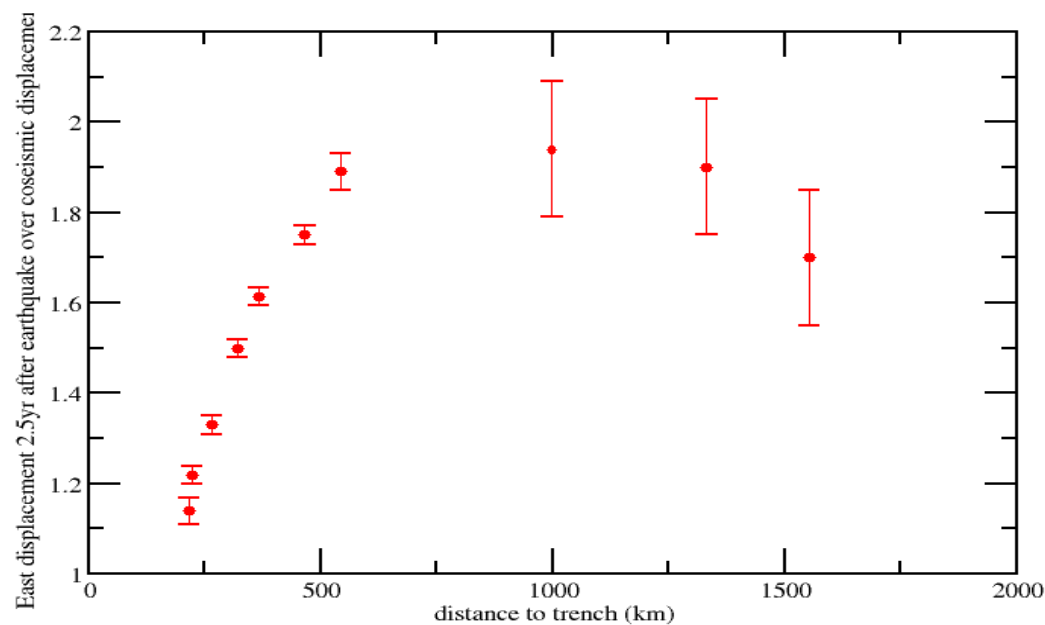
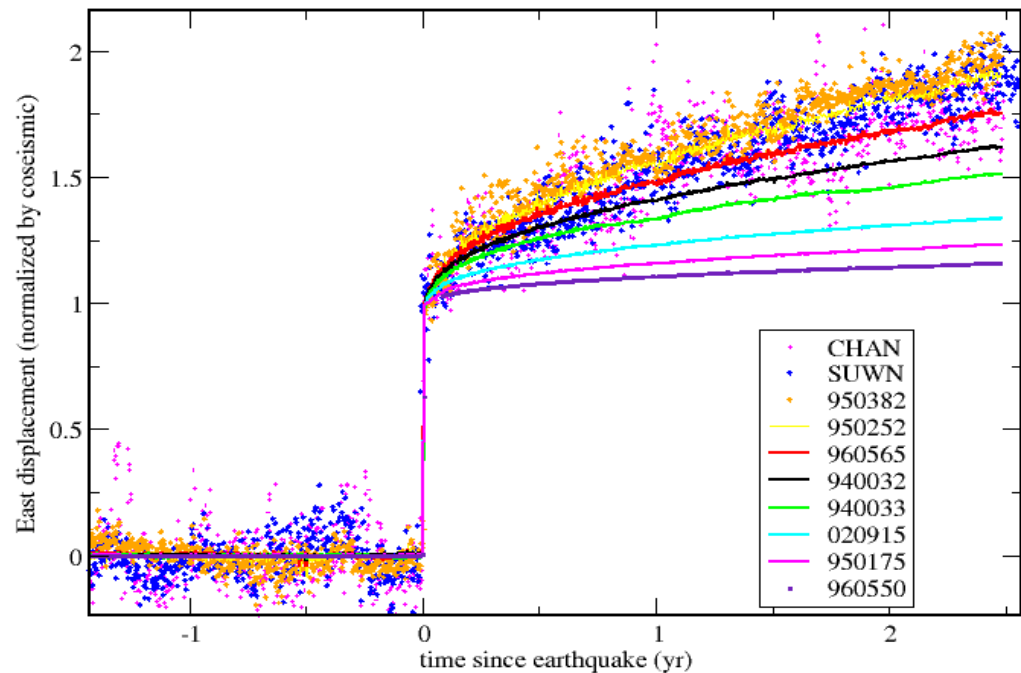
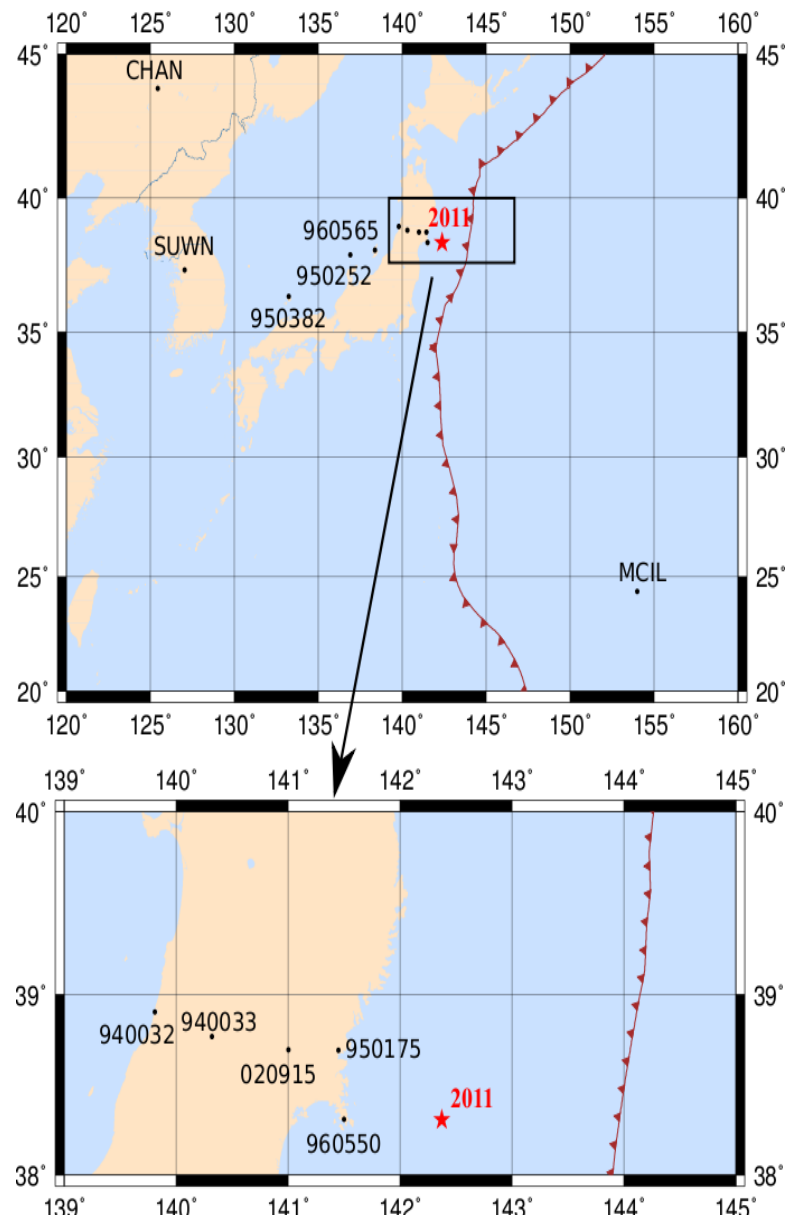
# GPS velocities and displacements after Tohoku (corrected for preseismic)



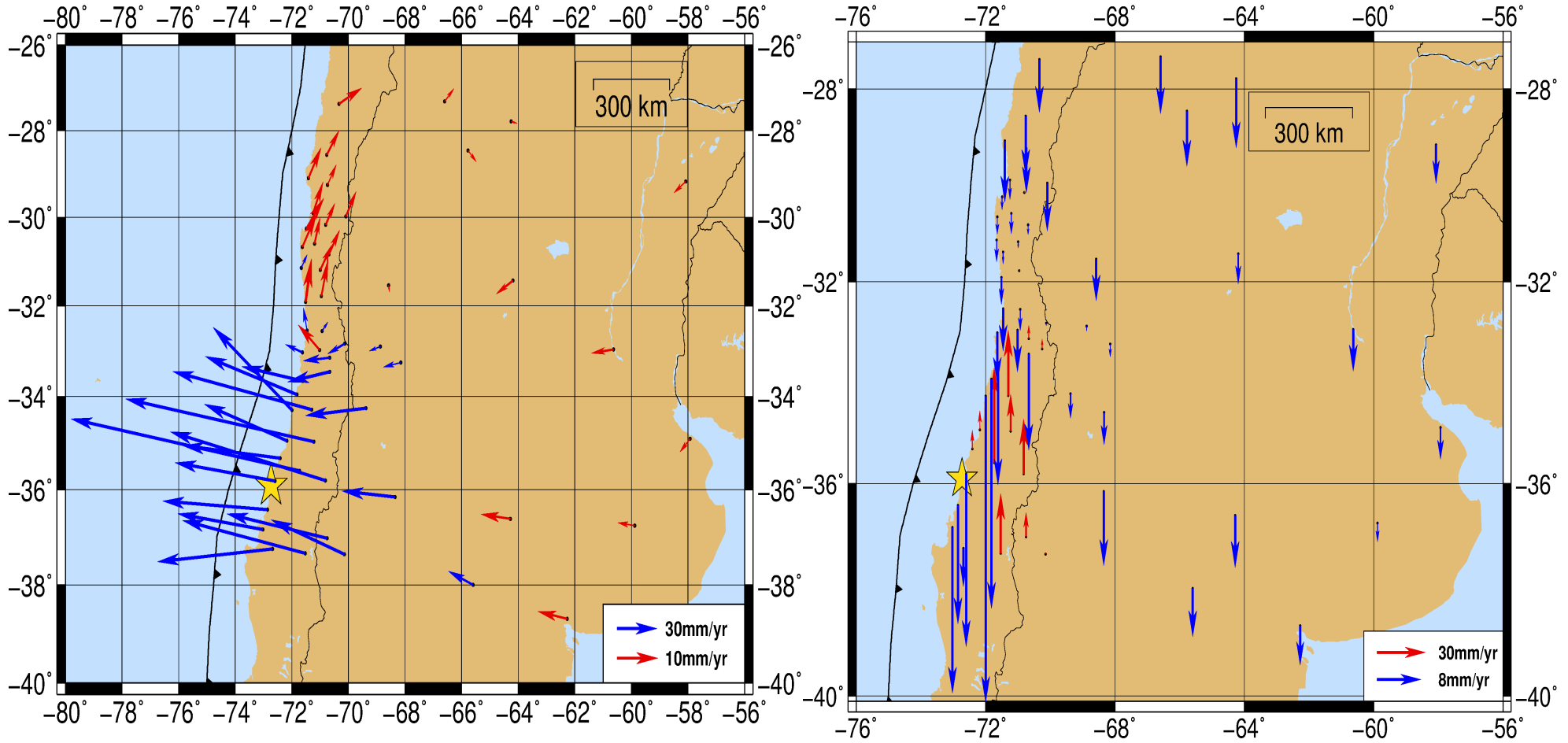
Horizontal velocities in 2012



# Non-dimensionalized horizontal displacements function of distance

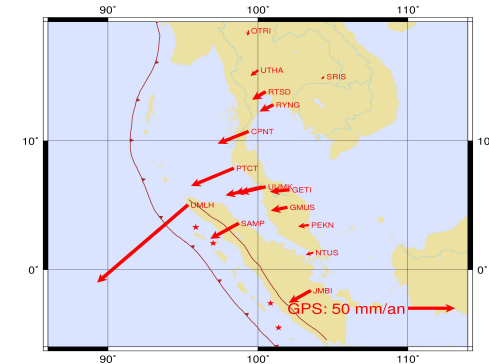
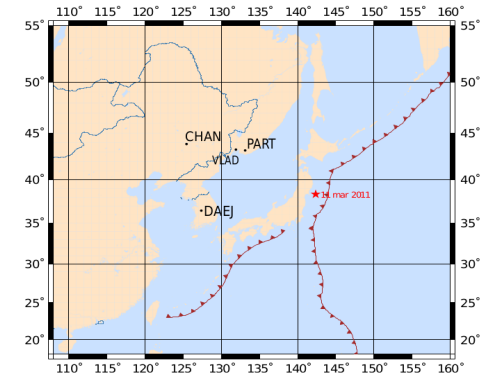
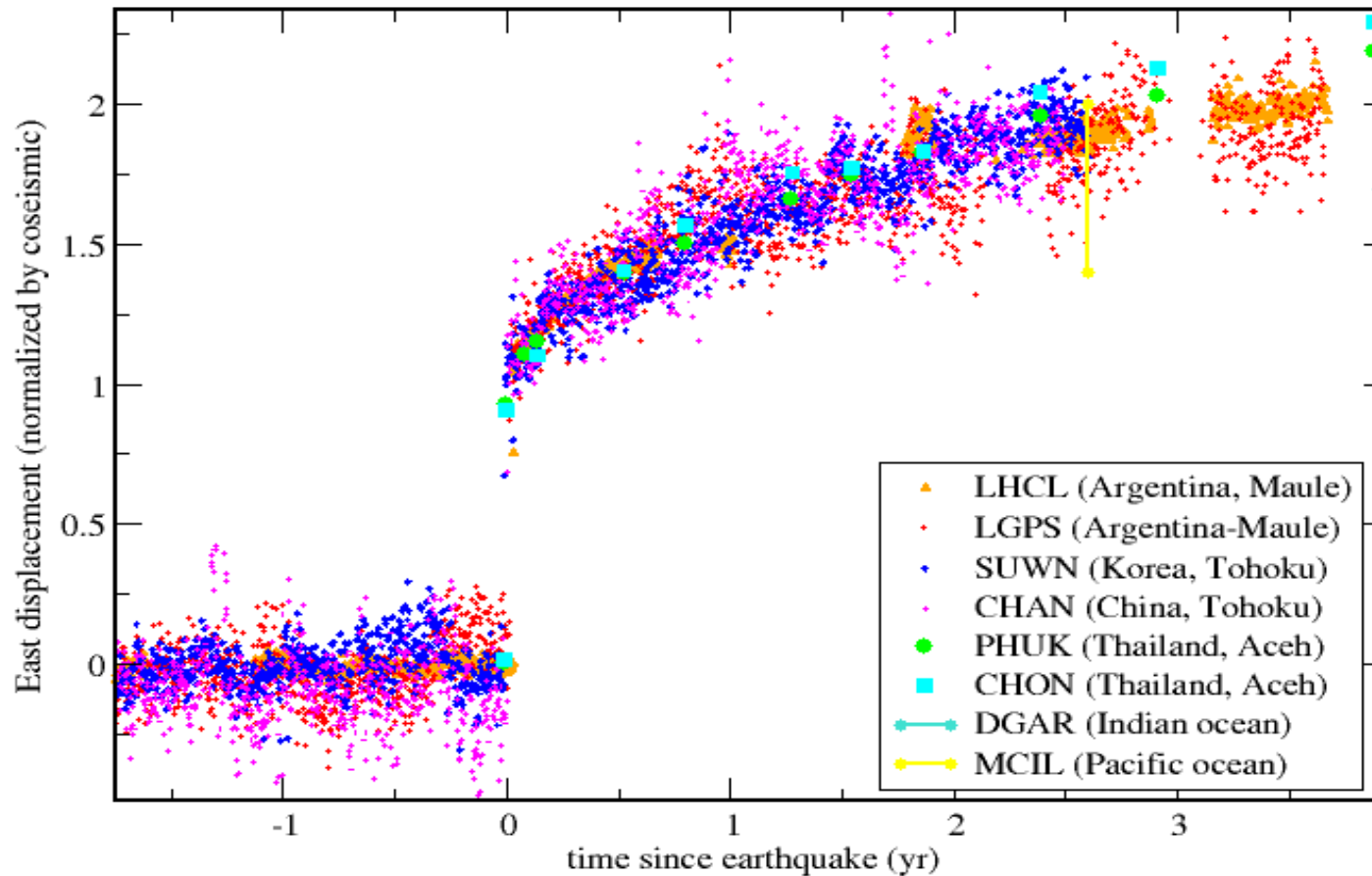


# AFTER MAULE



In 2012 (velocity difference between after and before Maule)  
Notice the vertical uplift on the Andes

# After the three giant earthquakes at distances between ( and 1000km)

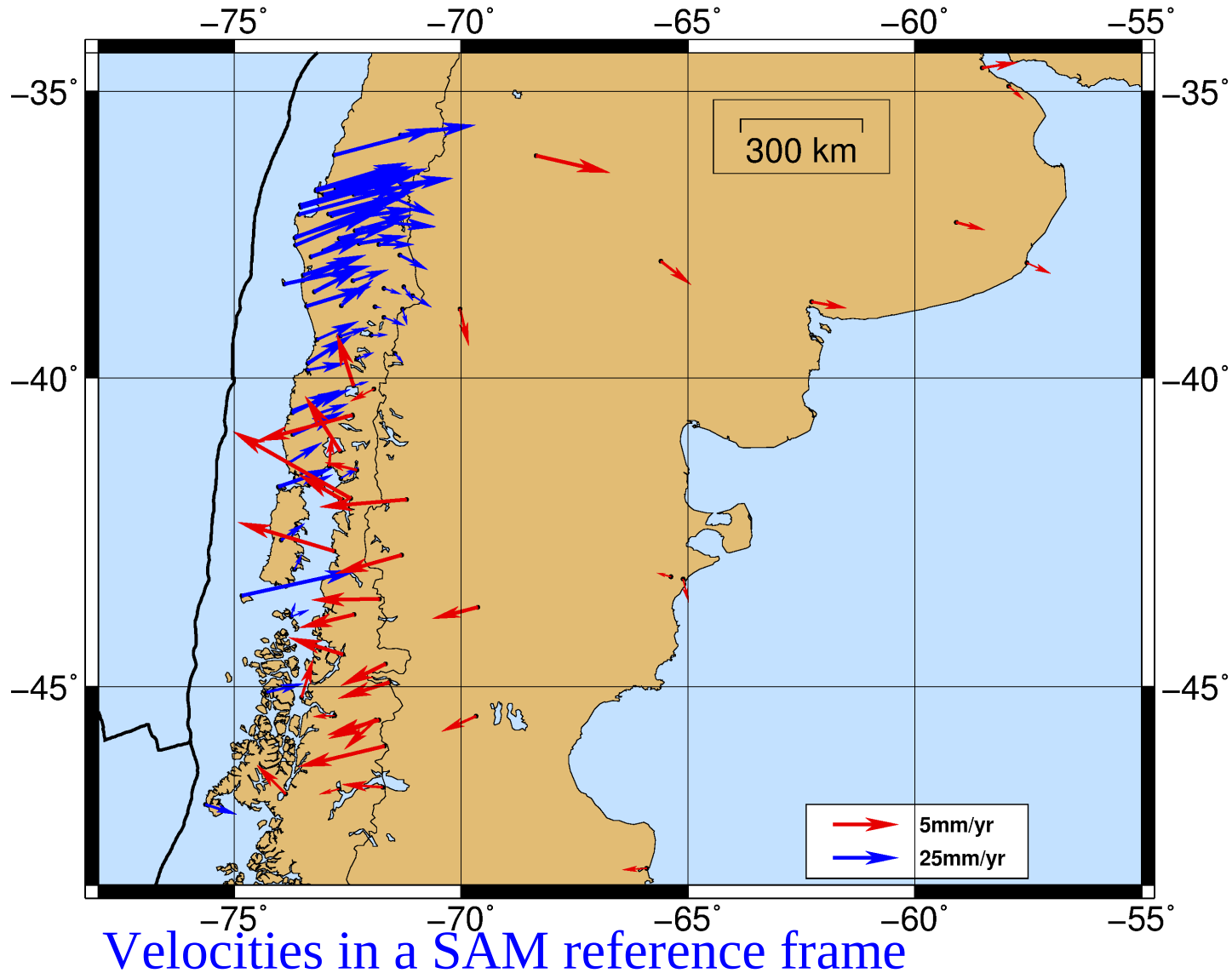


The postseismic signal non dimensionalised by coseismic is similar for the three earthquakes



# The postseismic phase continues for several decenies

After Valdivia (before Maule)



See Klotz et al. EPSL 2001

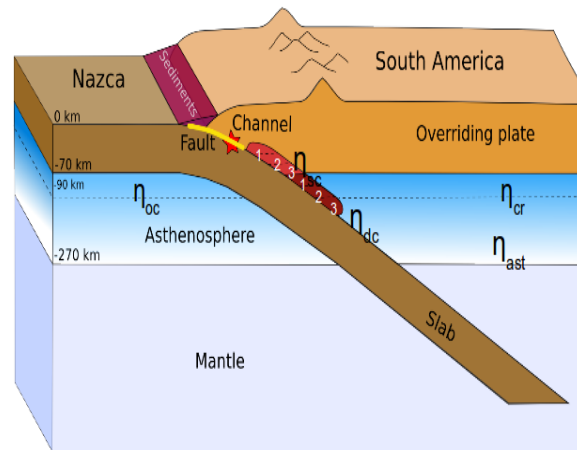
Hu et al. JGR 2004

## Postseismic velocities after a subduction earthquake:

- Large perturbation of horizontal velocities. 'Bell shape' of the postseismic over coseismic velocity function of distance curve
- Similar non-dimensional curve for the far-field stations in various zones of the world
- Subsidence in the far-field, uplift on the oceanward side of the volcanic arc
- The perturbation of the velocities persists for at least several decennies

**Finite element mechanical models are used to understand the origin of those postseismic velocities**

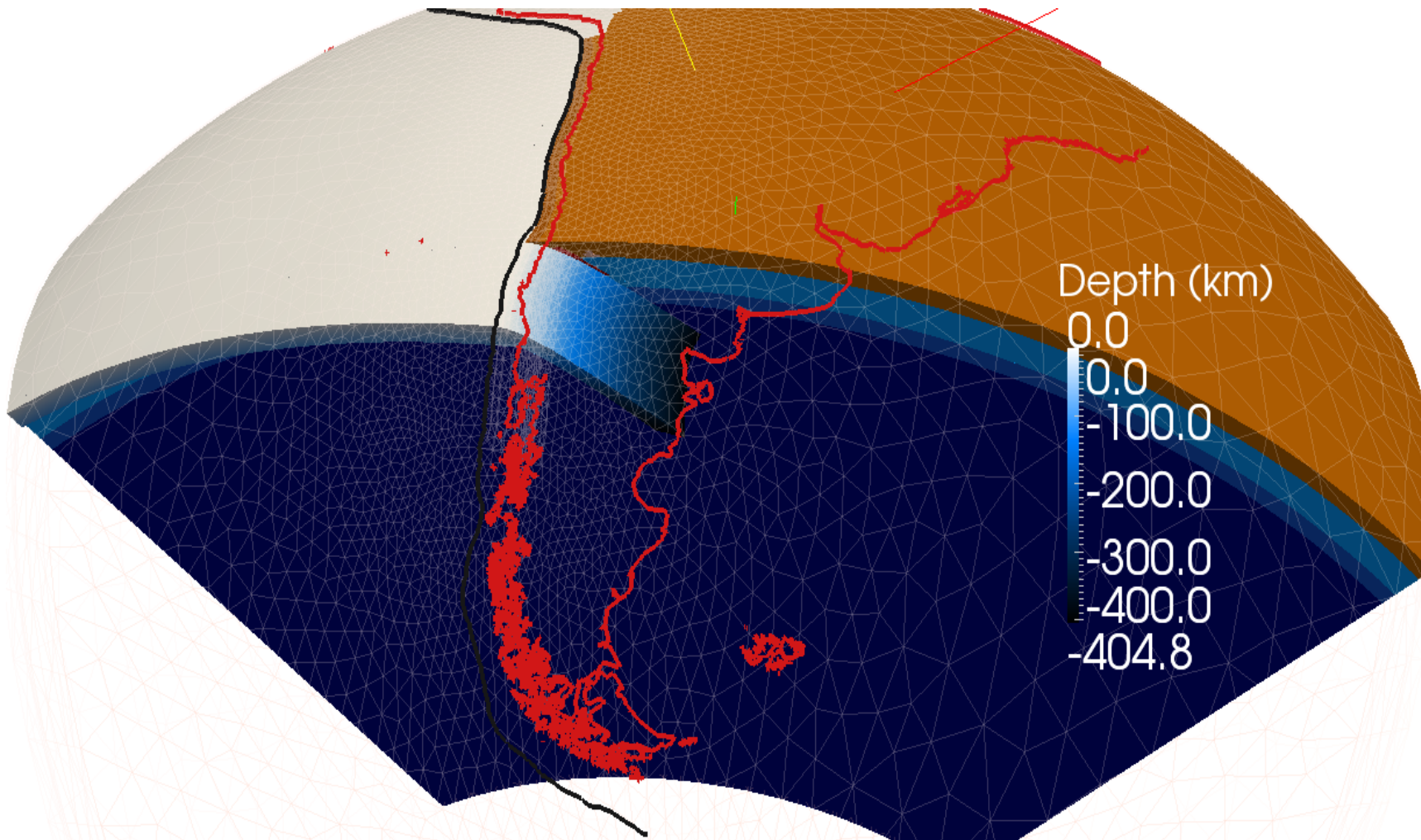
# Mesh for Maule earthquake



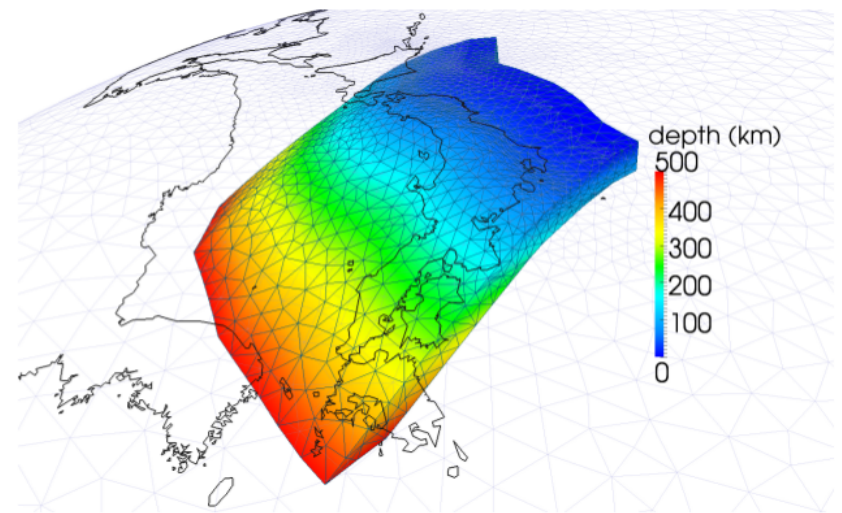
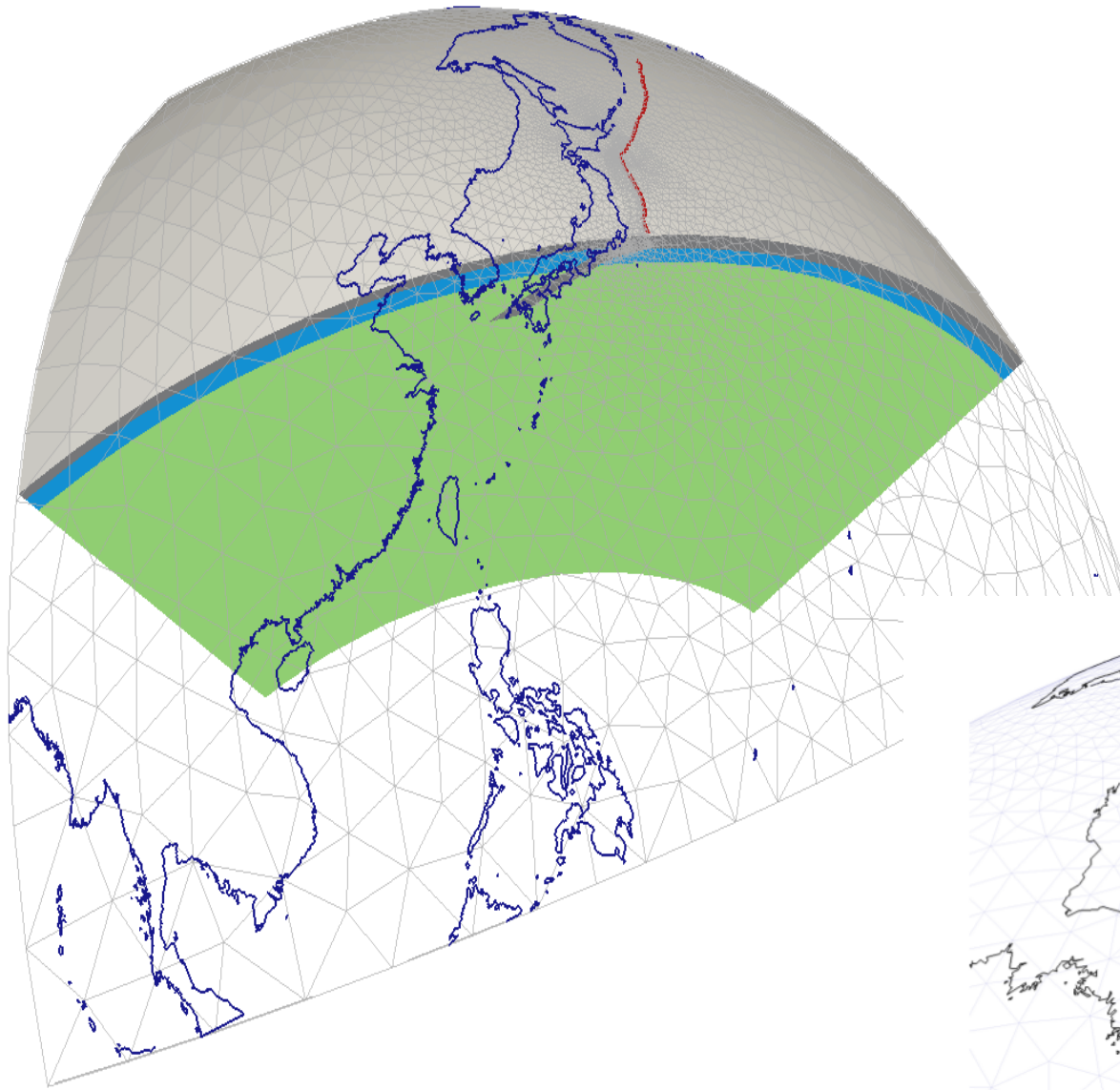
		$\eta_1$
<b>SC1</b>	shallow channel 1	1,85E+018
<b>SC2</b>	shallow channel 2	7,50E+017
<b>SC3</b>	shallow channel 3	5,01E+017
<b>DC1</b>	deep channel 1	1,86E+017
<b>DC2</b>	deep channel 2	5,72E+017
<b>DC3</b>	deep channel 3	5,72E+017
<b>CR</b>	crust (70-90 km)	2,00E+018
<b>OC</b>	ocean (70-90 km)	9,70E+018
<b>Asthenosphere</b>	90-200 km	3,00E+018
	200-270 km	8,40E+018

$$\eta_2/\eta_1 = \text{cst} = 6,5$$

$$\mu_2/\mu_1 = \text{cst} = 3$$

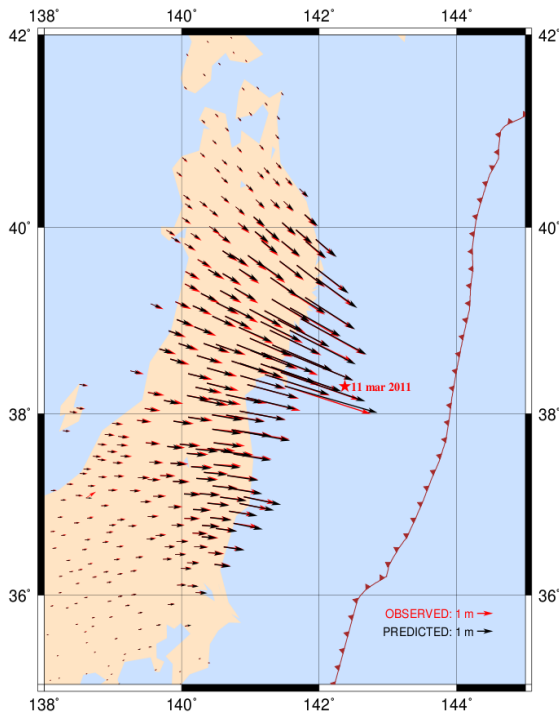
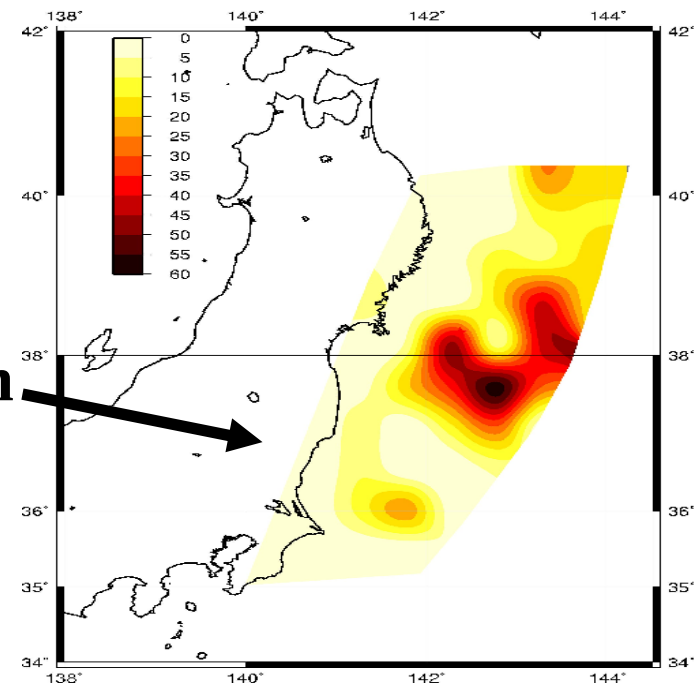


# Finite element mesh for Japan computations with Zset-Zebulon

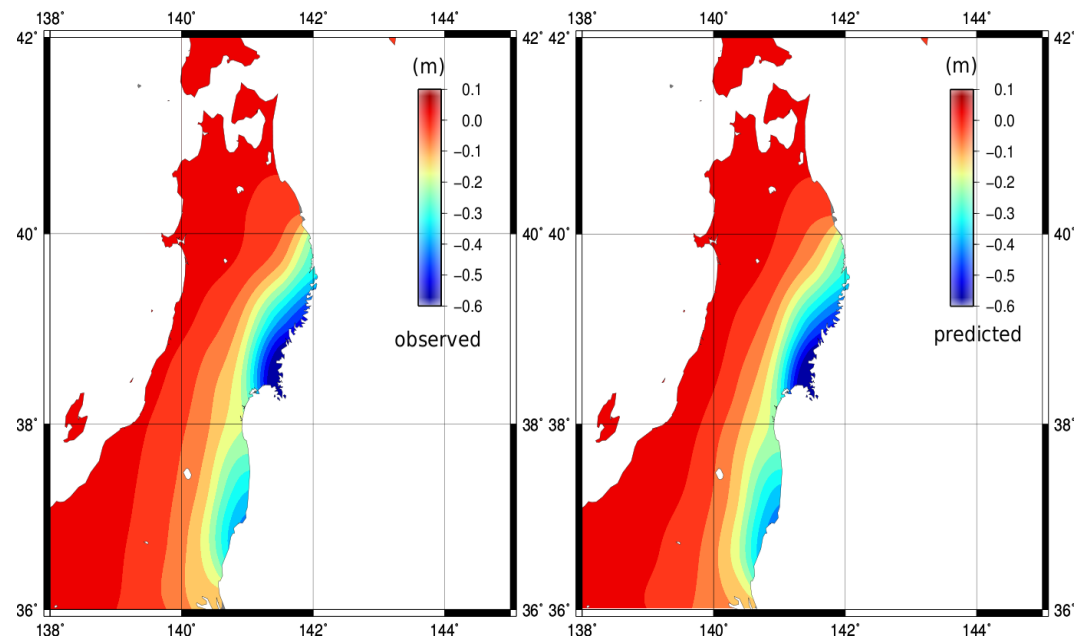


# Coseismic (here example of Japan)

**Predicted coseismic slip on the interface**



**horizontal**



**vertical**

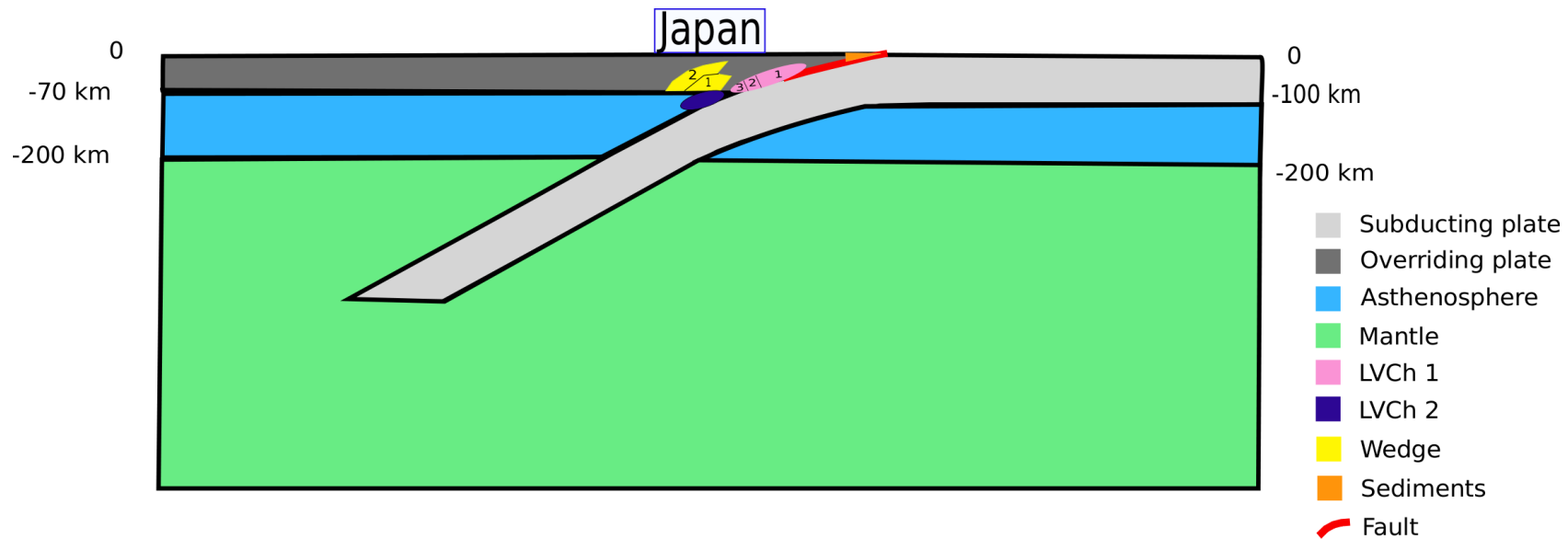
The coseismic deformation induces stresses in the mantle, and on the plate interface.

**What induces postseismic deformation?**

- Slip on the fault plane at shallow depths?
- Relaxation in the asthenosphere?
- Relaxation in a low viscosity channel (LVC)?

Then if there is viscoelastic deformation, what is the appropriate rheology?

# Viscosities obtained from inversion (Japan)



## viscosities:

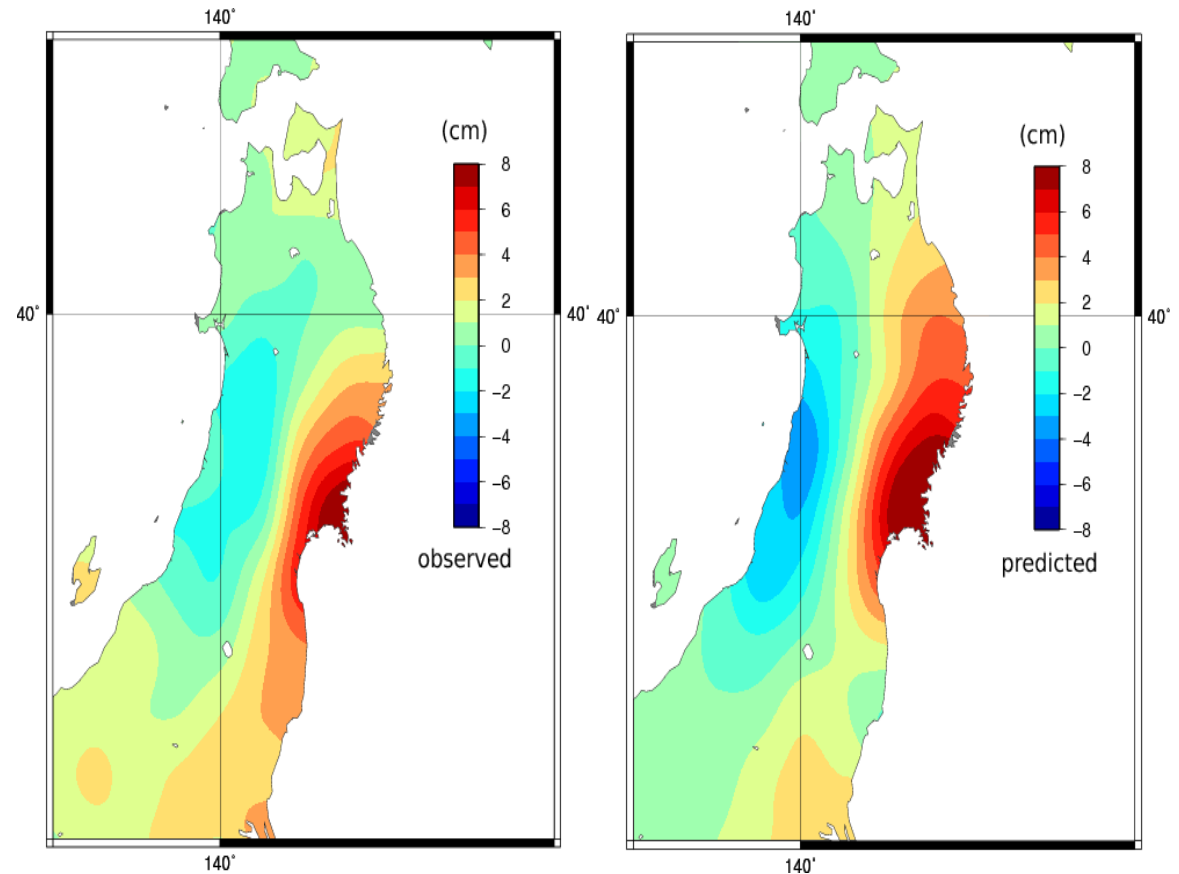
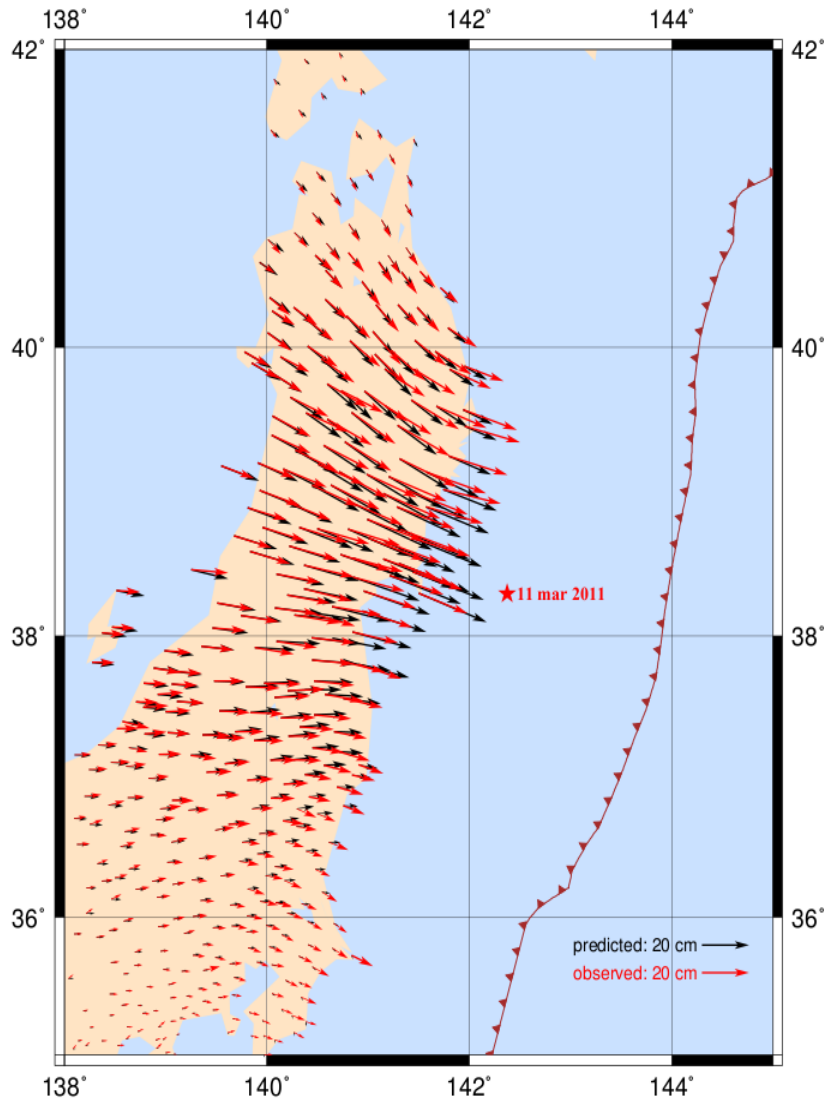
Asthenosphere (70-200km):  $1.6 \cdot 10^{18}$  Pas (a short-term viscosity?)

LVCh: from  $10^{17}$  Pas to  $6 \cdot 10^{17}$  Pas

LVW1 :  $4 \cdot 10^{17}$  Pas? (poorly constrained)

No additional postseismic slip on the interface required

# Comparison between observed and predicted postseismic velocities in Japan (jan to dec 2012)

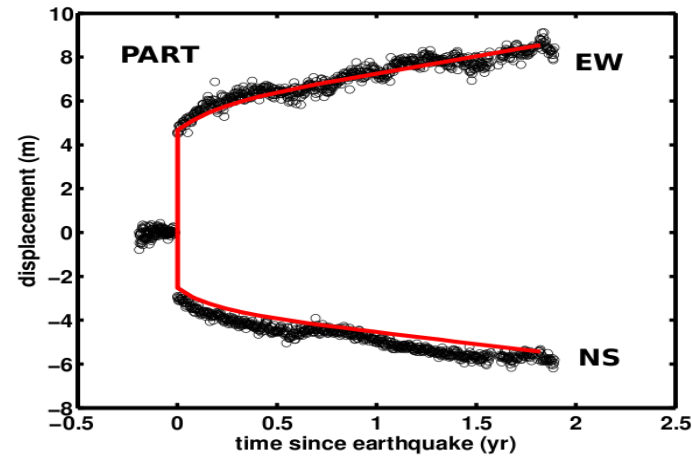
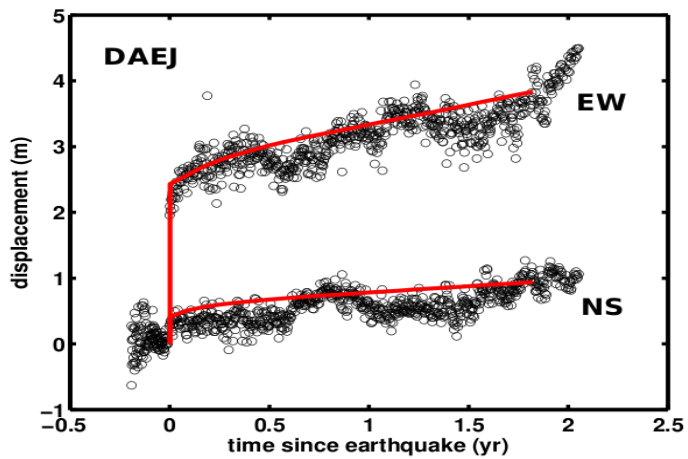
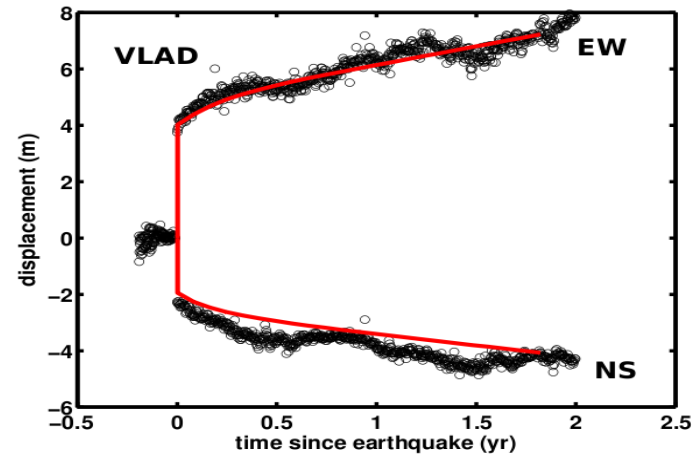
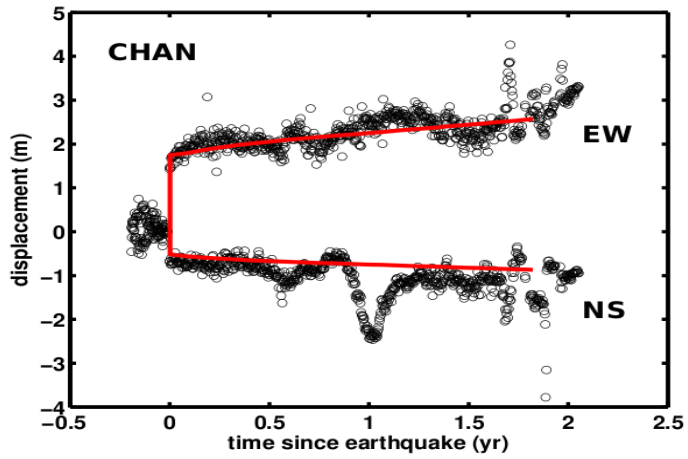
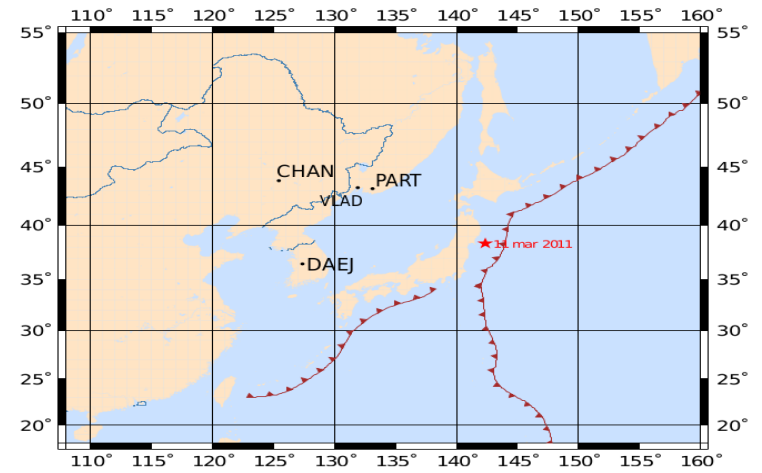


**Horizontal velocities**

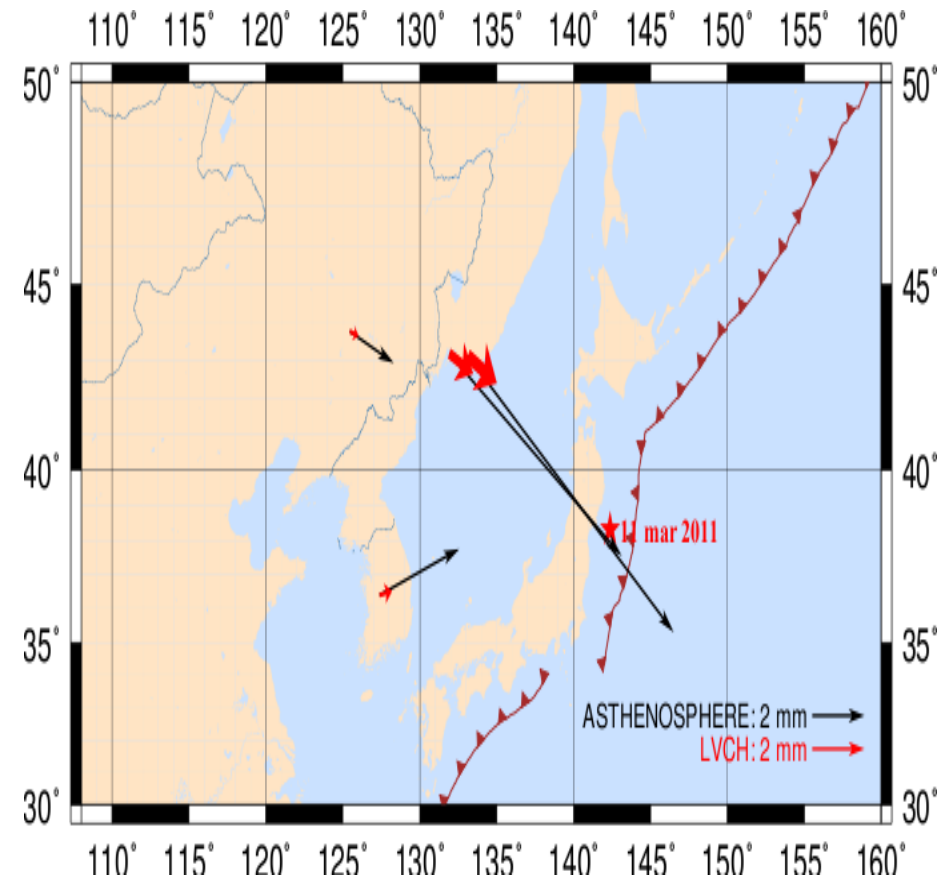
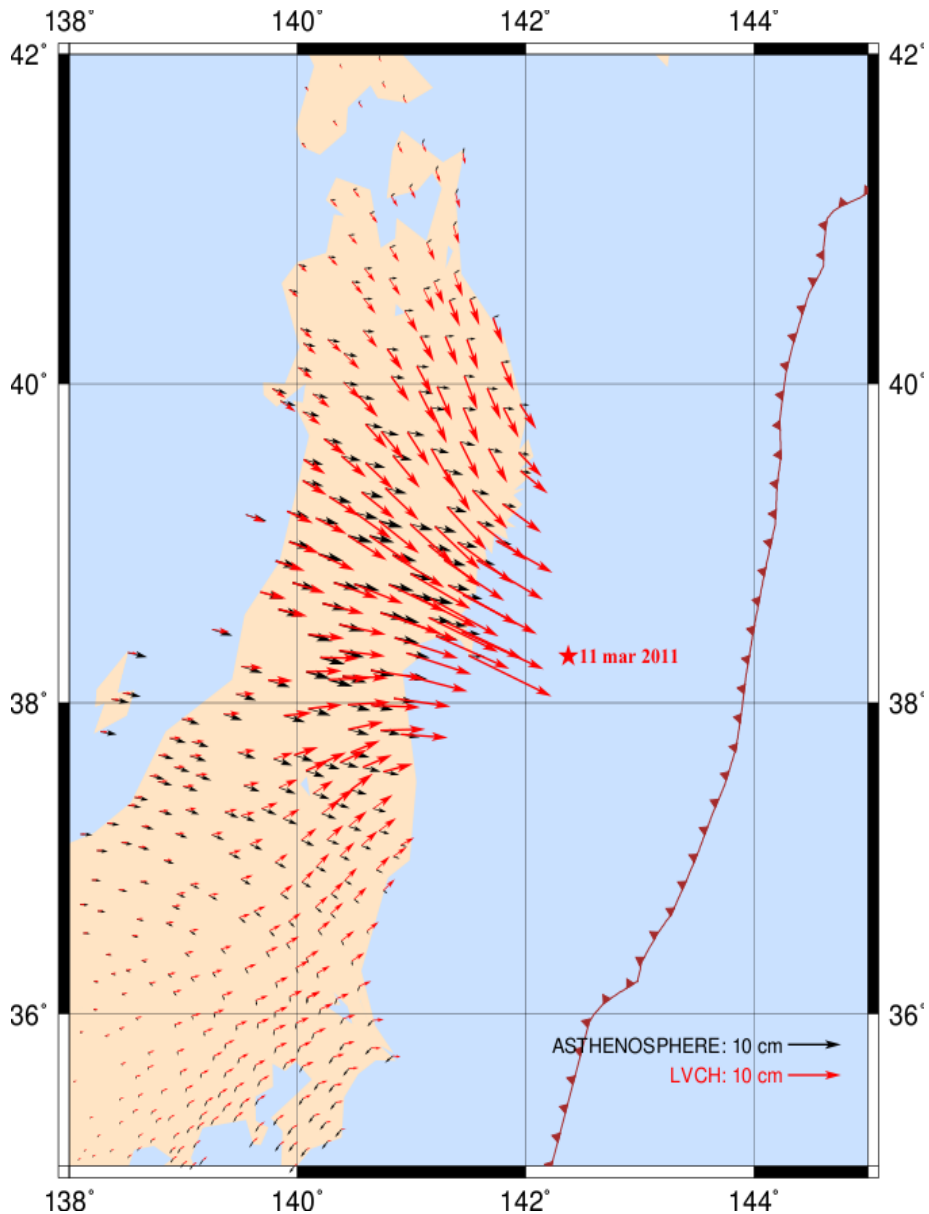
**Vertical velocities**



# Fit to far-field stations time series



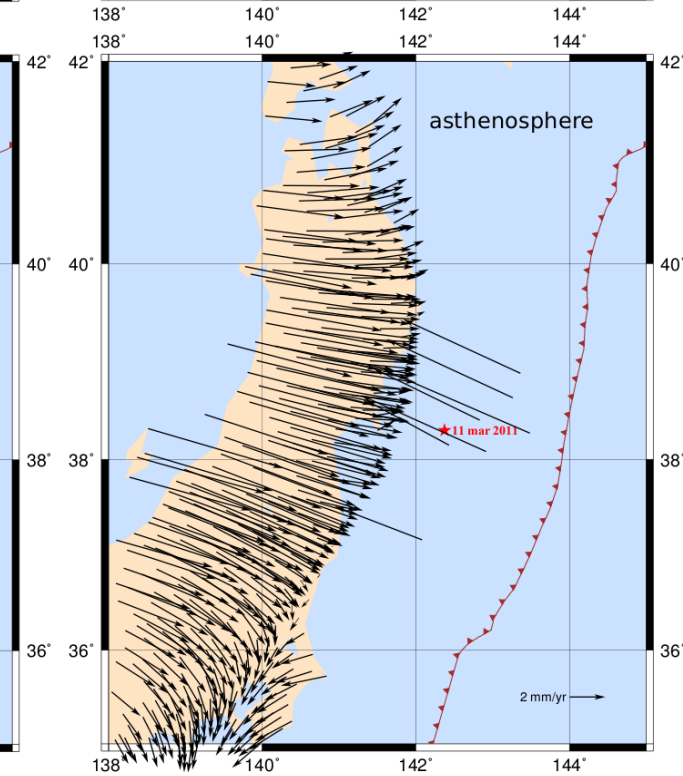
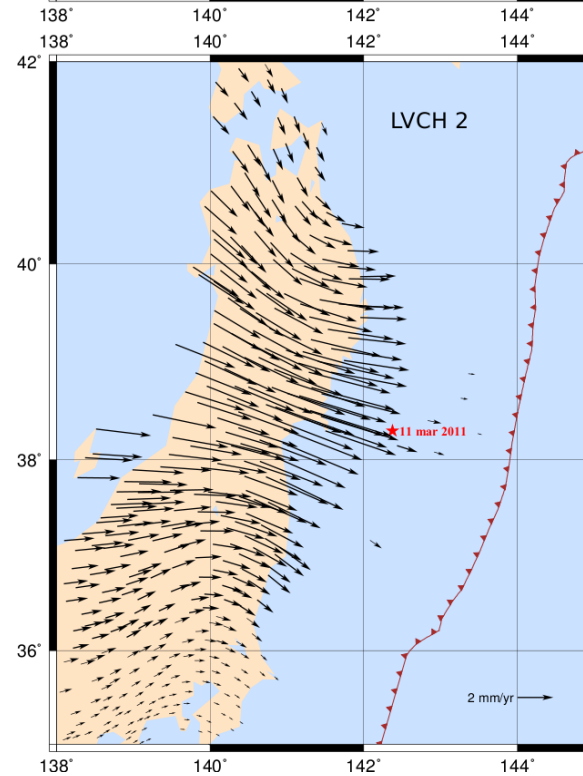
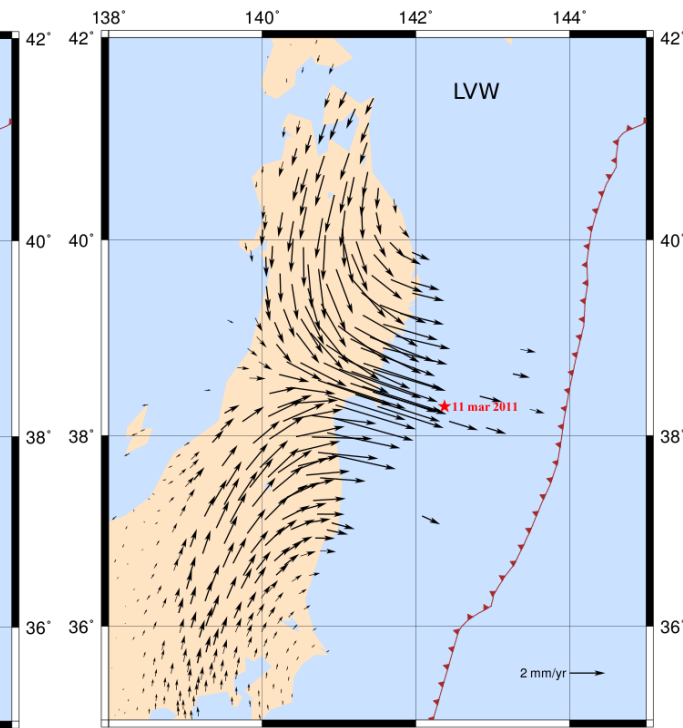
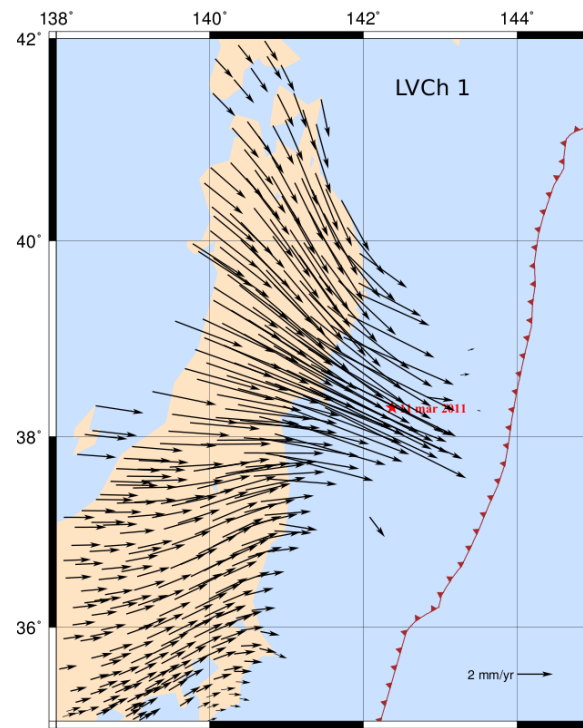
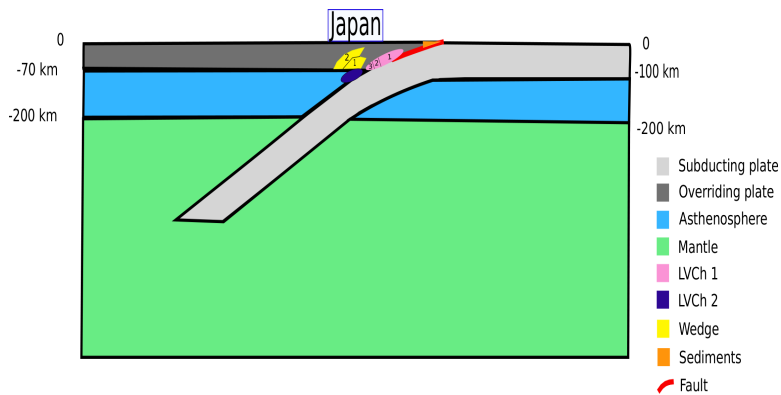
# Both relaxation in the asthenosphere and relaxation in LVW or LVCh are necessary



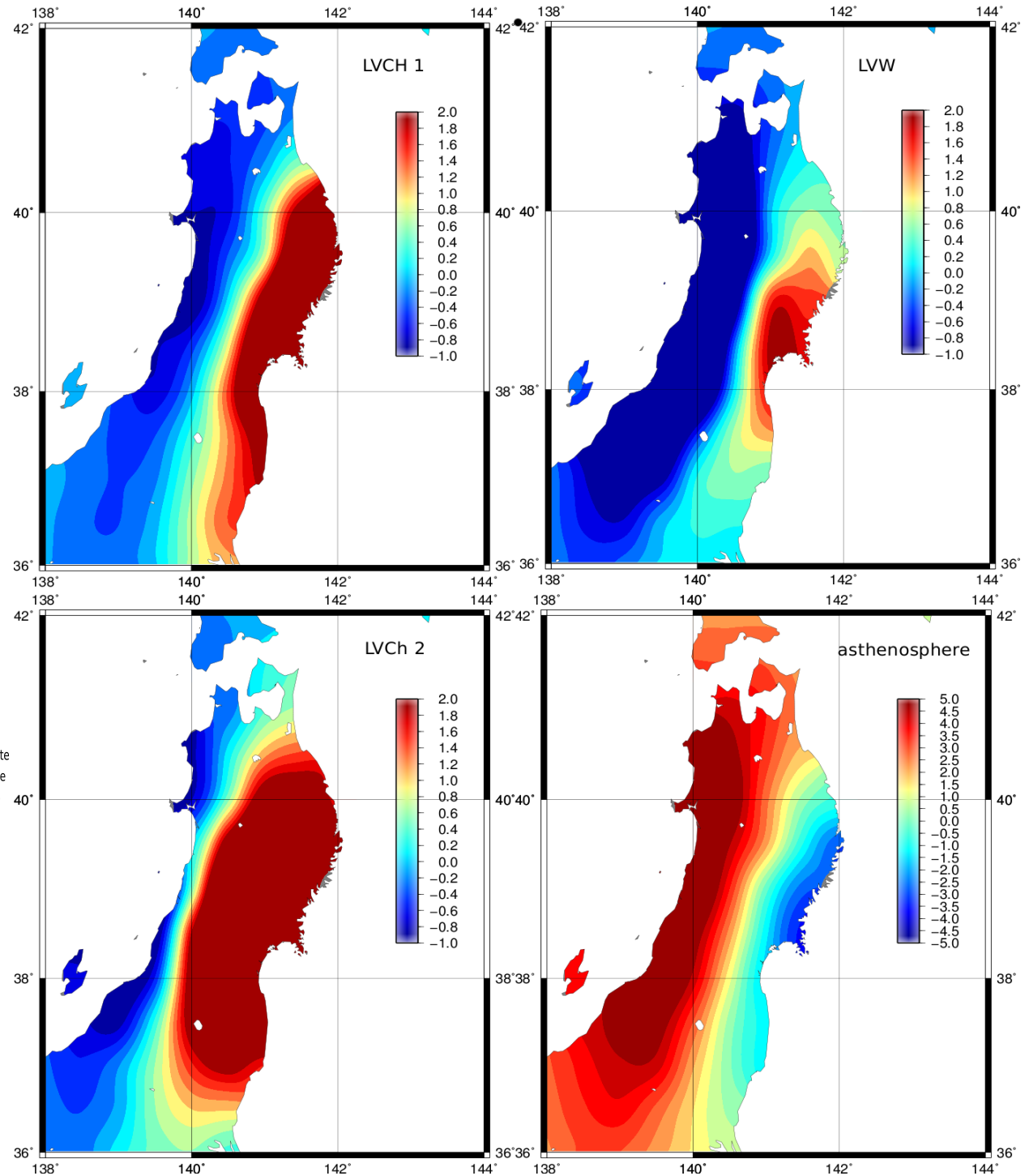
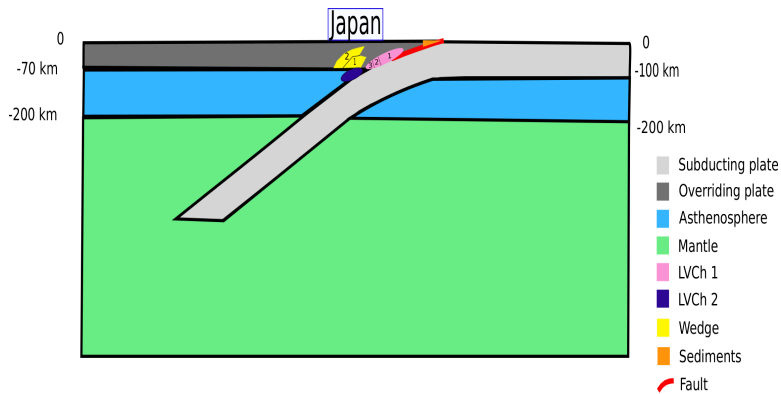
**A model with only relaxation in the asthenosphere which fits the far-field velocities (black) induces negligible near-field velocities**

**A model with only relaxation in the LVCh which fits the Japanese velocities (red) induces negligible far-field velocities.**

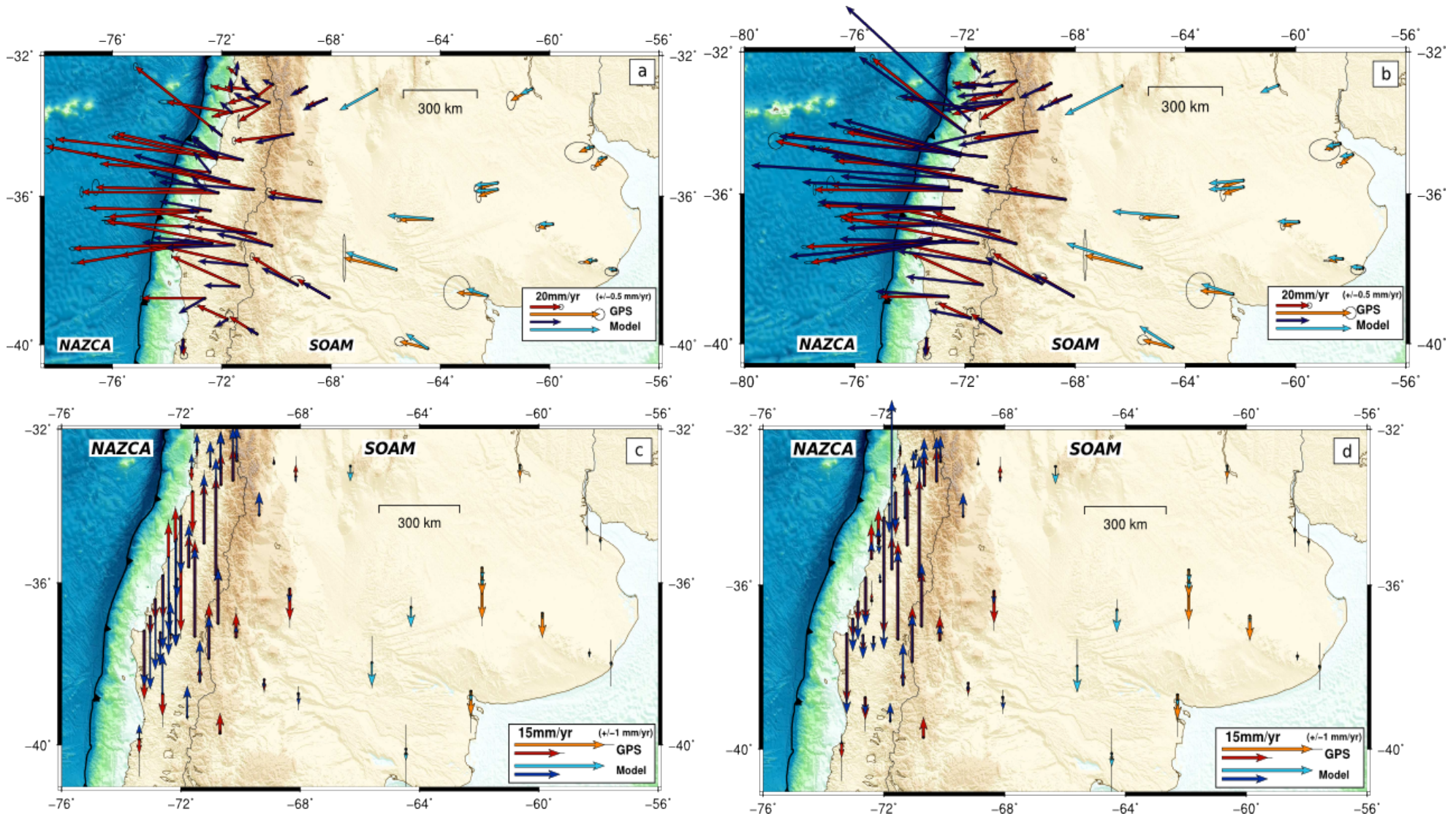
# Impact of the viscosity in various zones on horizontal velocities ('partial derivatives')



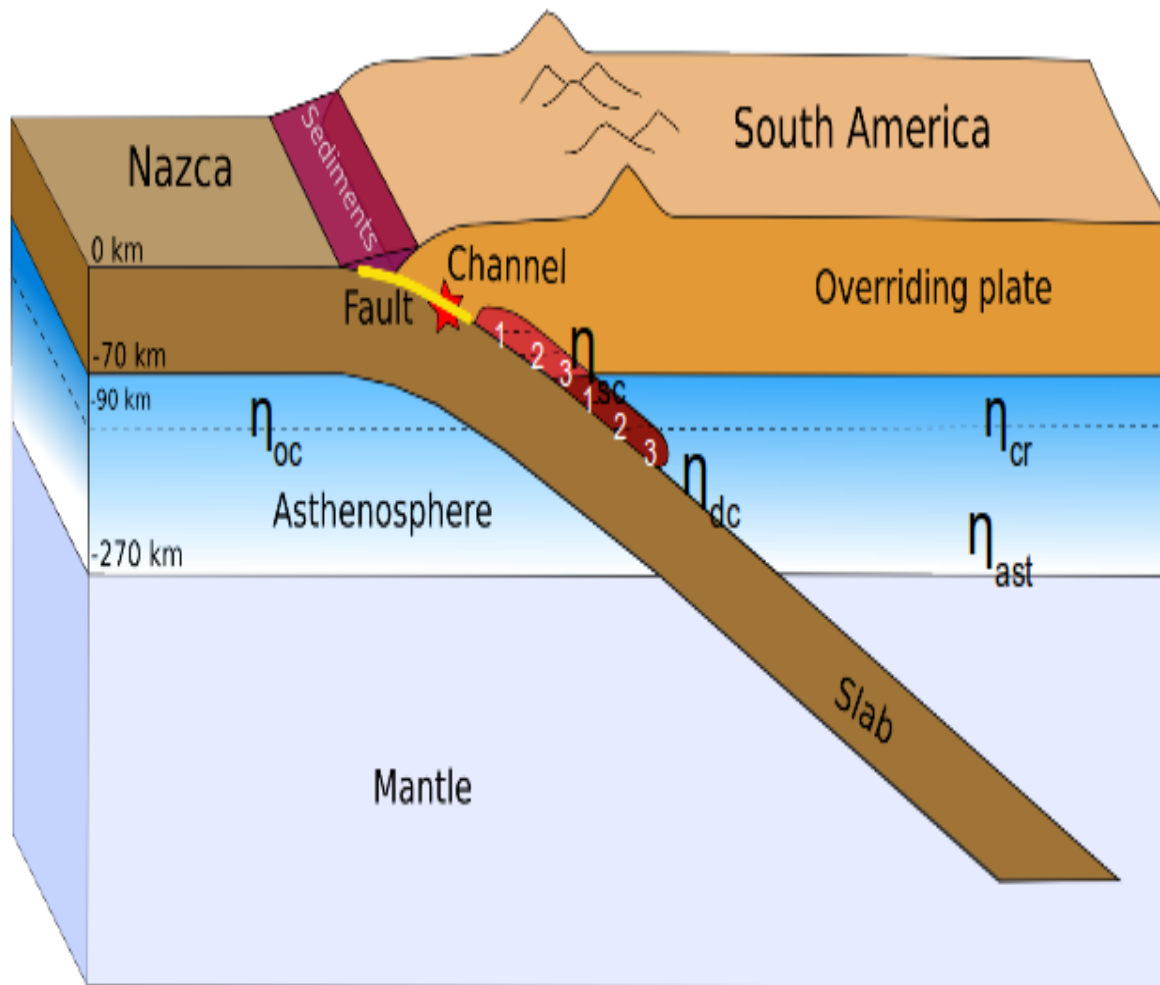
# Impact of the viscosity in various areas on vertical velocities ('partial derivatives')



# In Chile, additional slip on the interface is required



*Modeled vs measured mean velocities over the 2nd year after the earthquake :*  
a ) horizontal vel. in the case of a pure viscoelastic relaxation model  
b ) vertical vel. in the case of a pure viscoelastic relaxation model  
c ) horizontal vel. in the case of a viscoelastic relaxation + afterslip model  
d ) vertical vel. in the case of a viscoelastic relaxation + afterslip model



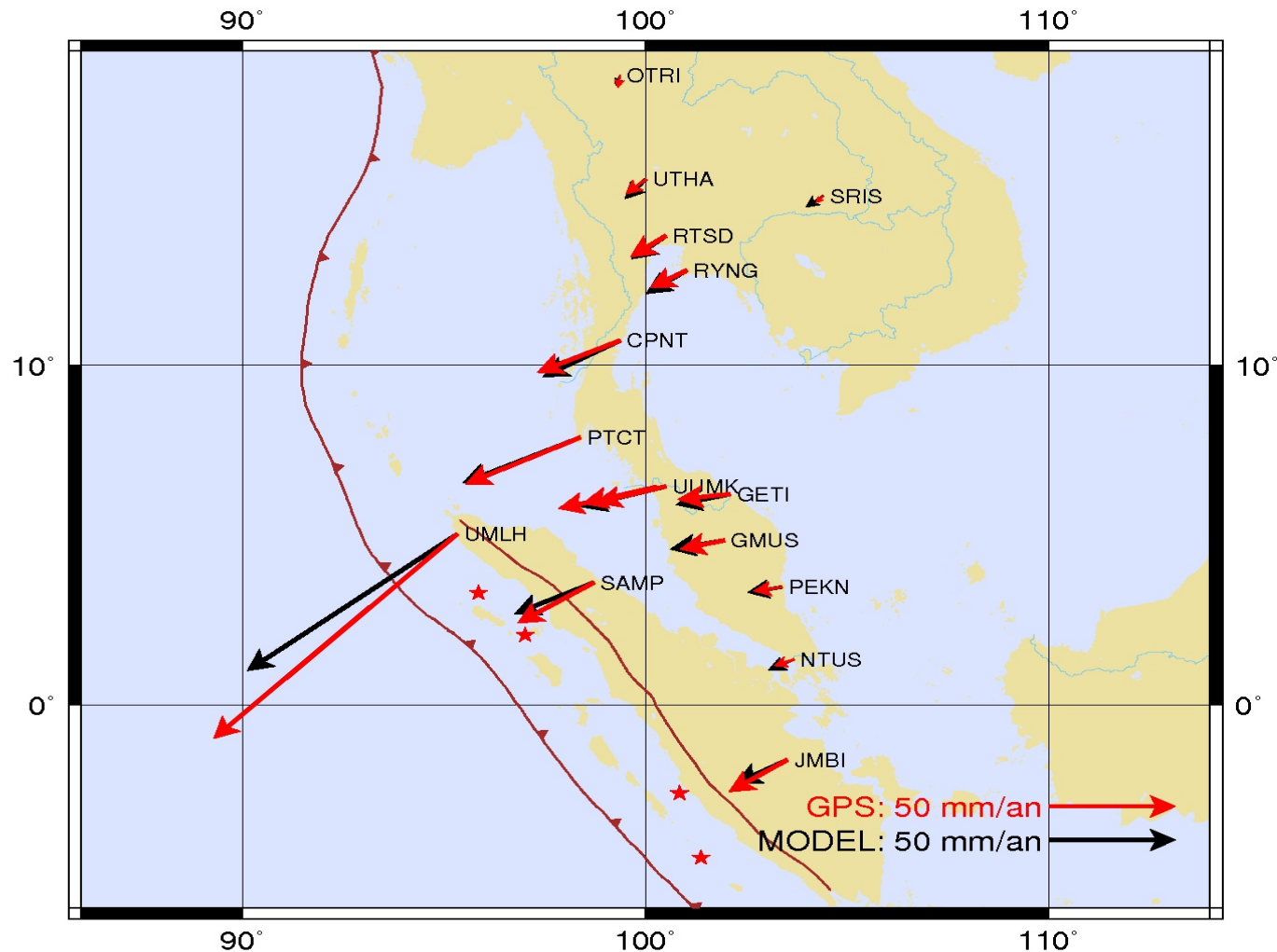
		$\eta_1$
<b>SC1</b>	shallow channel 1	1,85E+018
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$$\eta_2/\eta_1 = \text{cst} = 6,5$$

$$\mu_2/\mu_1 = \text{cst} = 3$$

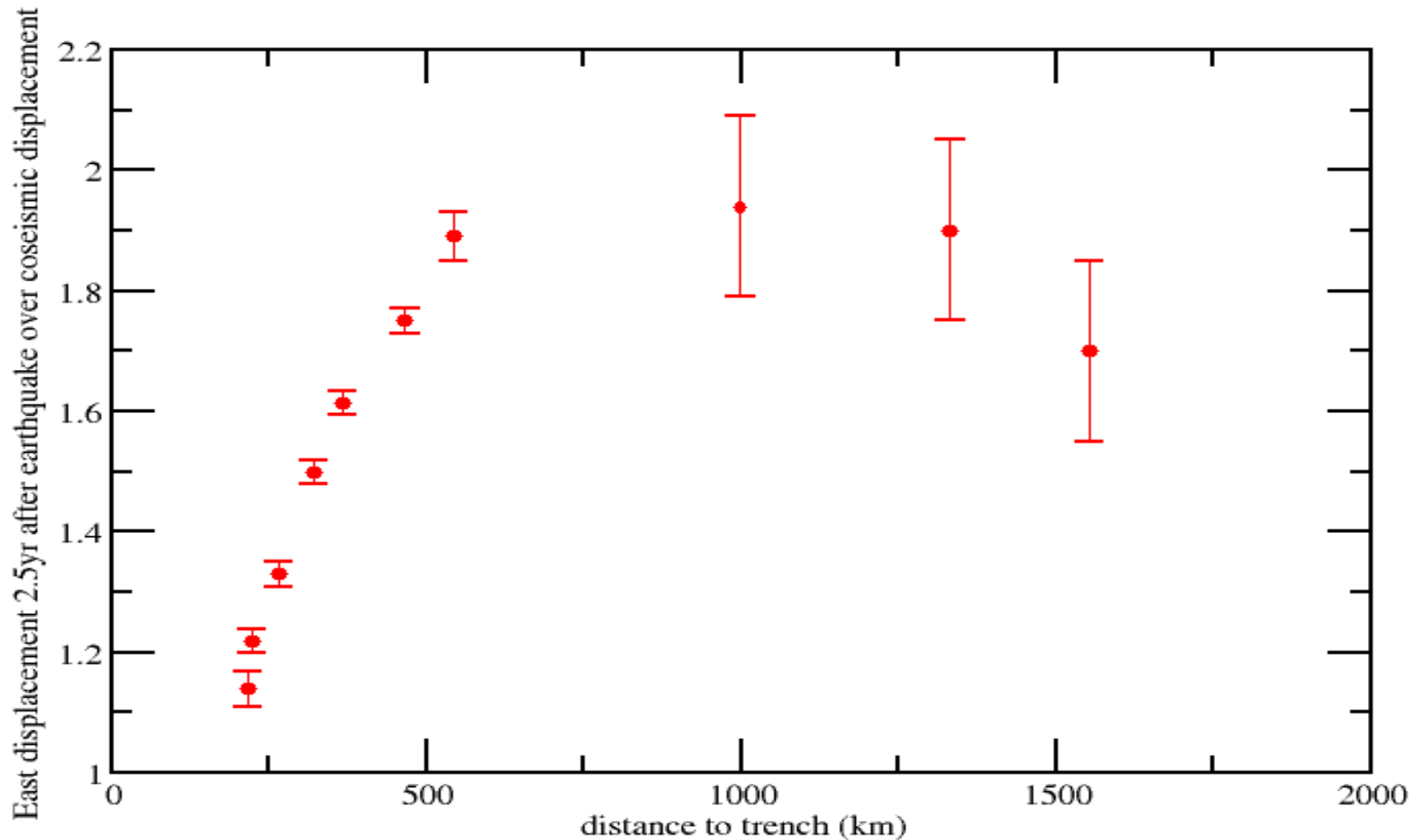
# Far-field horizontal postseismic velocities after 2004 Sumatra earthquake

2008 velocities- predicted vs observations.  $\mu = 3 \cdot 10^{18}$  Pas



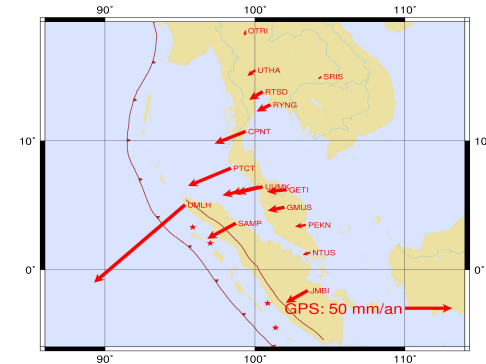
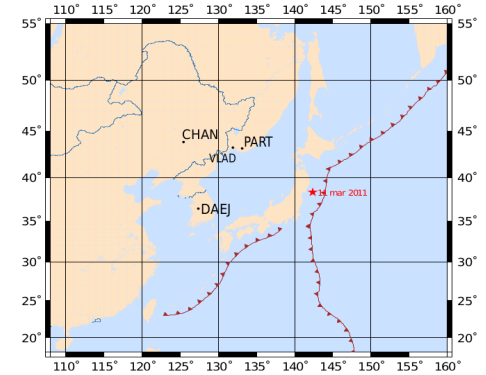
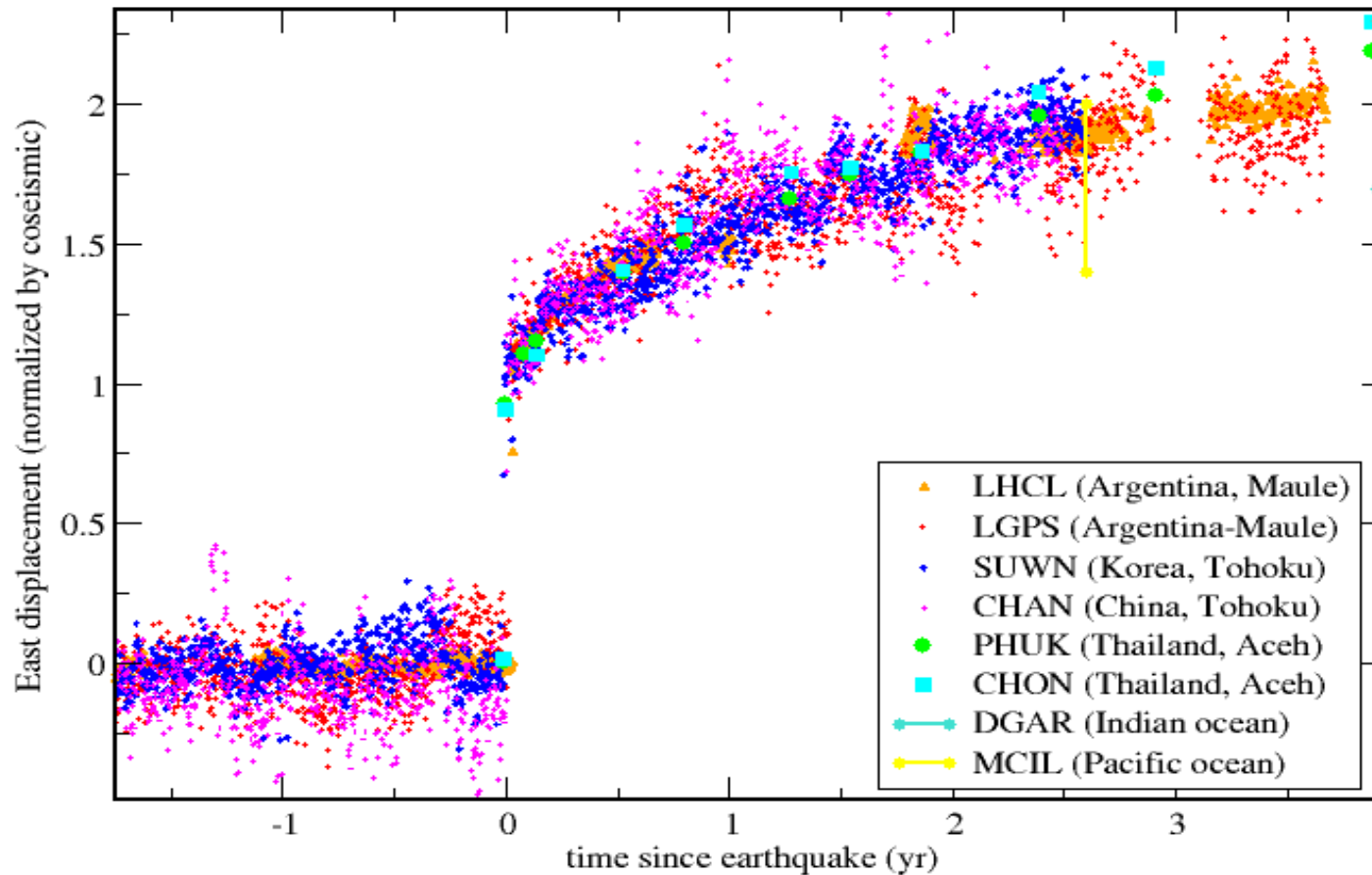
# An asthenosphere of finite thickness:

This curve keeps on increasing in case of low viscosity in the whole upper mantle





# After the three giant earthquakes at distances between ( and 1000km)



The postseismic signal non dimensionalised by coseismic is similar for the three earthquakes **but the deviatoric stresses differ by a factor 5!!!**

## Summary of the results from models of postseismic deformation after large subduction earthquakes:

**Relaxation in the asthenosphere:** Necessary for explaining far-field horizontal and vertical velocities. Viscosity around  $2 \text{ to } 4 \times 10^{18}$  Pas (likely a long-term 'transient'). Asthenosphere 150-200km thick;

**Relaxation in a 'low viscosity channel':** necessary for explaining middle-field velocities (uplift).

**Slip on the interface at shallow depths:** Important for Aceh and Maule. Less for Tohoku

## What we have learnt from the deformations following megathrust earthquakes:

Is there a low-viscosity asthenosphere?  
What rheology (Newtonian, non-linear creep), what thickness

Asthenosphere (200km thick?)  
with a viscosity around  $3 \cdot 10^{18}$  Pa.s. Transient creep but not power law rheology

What occurs below the brittle part of the subduction interface? Localization of the deformation in low-viscosity ductile shear zones?

A low viscosity channel is clearly identified: serpentine; hydrated material?

Other tectonic implications:  
What are the ingredients of the stress/strain-field?

Stresses and strains vary in time: Only close to megathrust earthquakes?

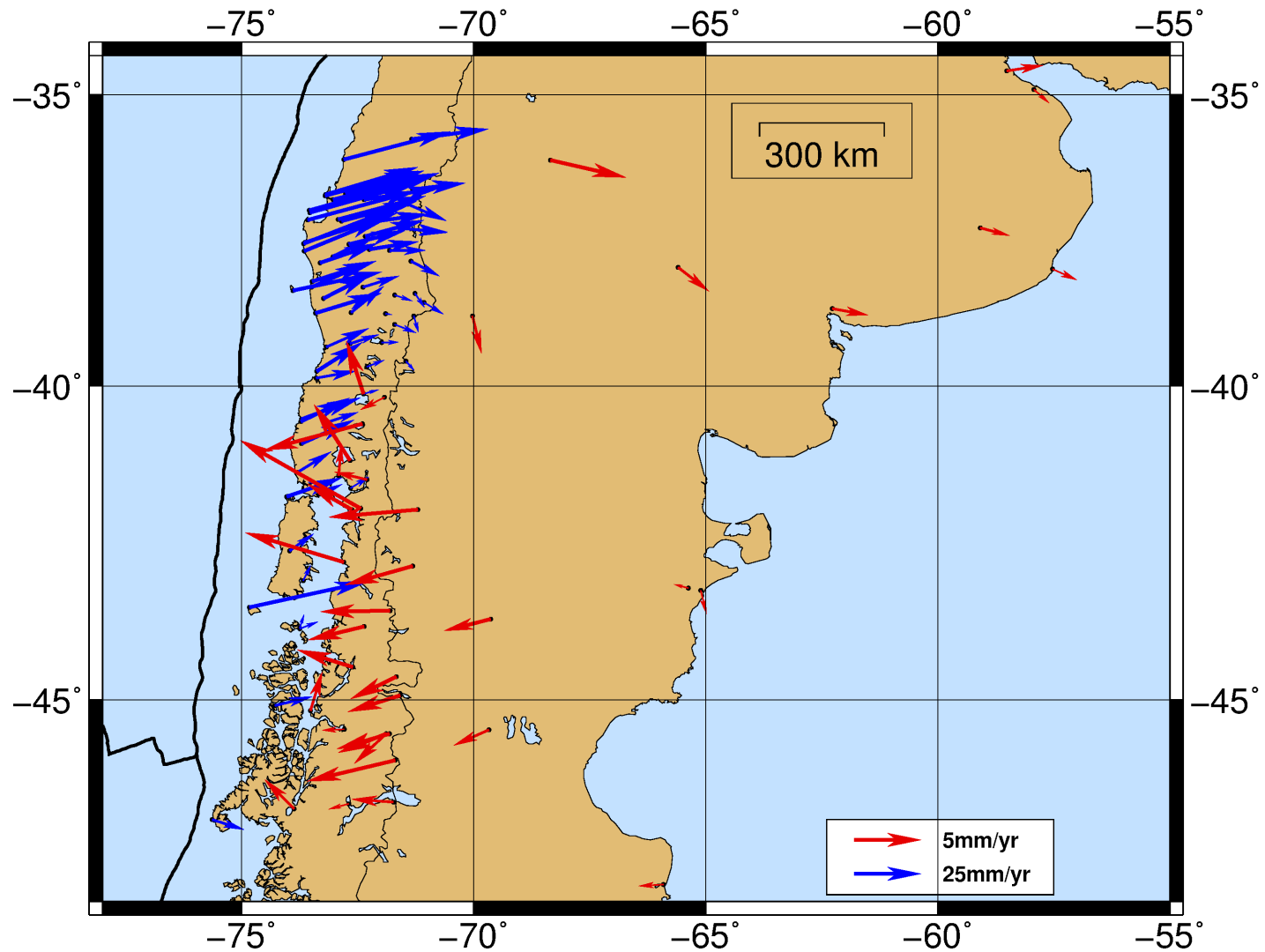
Implications concerning the stress evolution along the slab interface and within the overriding and subducting plates

What occurs on longer time-scales?

## **The seismic cycle**

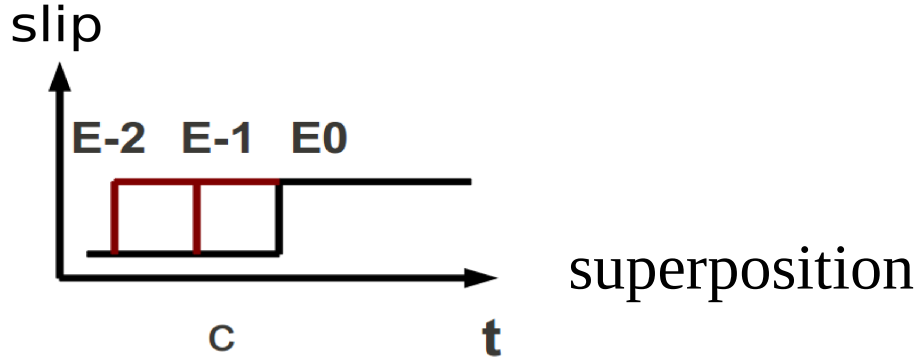
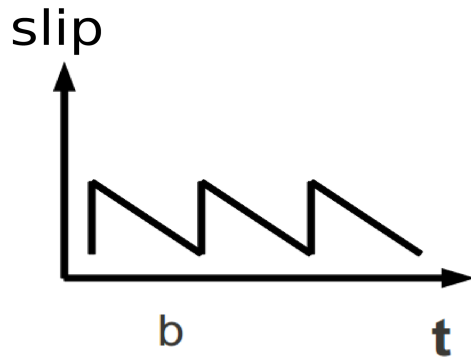
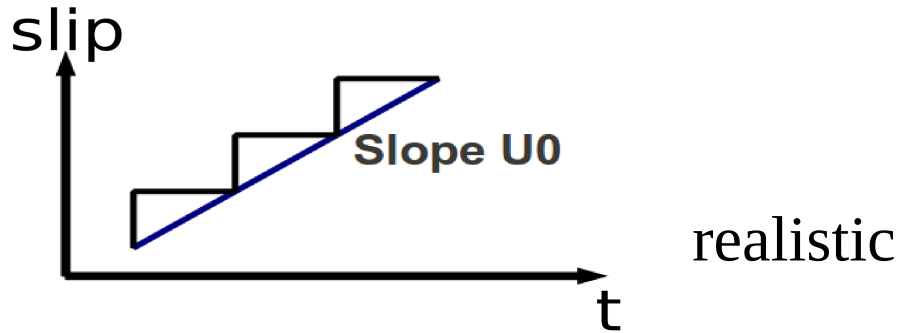
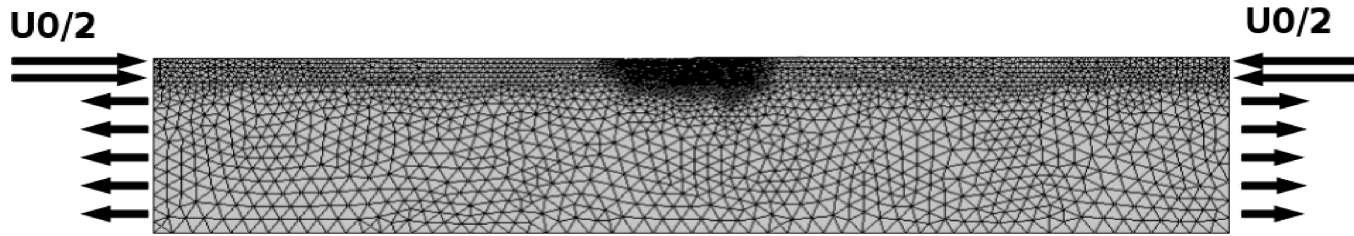
Viscoelastic versus elastic backslip...  
Observations and models

40yrs after Valdivia, before Maule (referential of North-East South America)



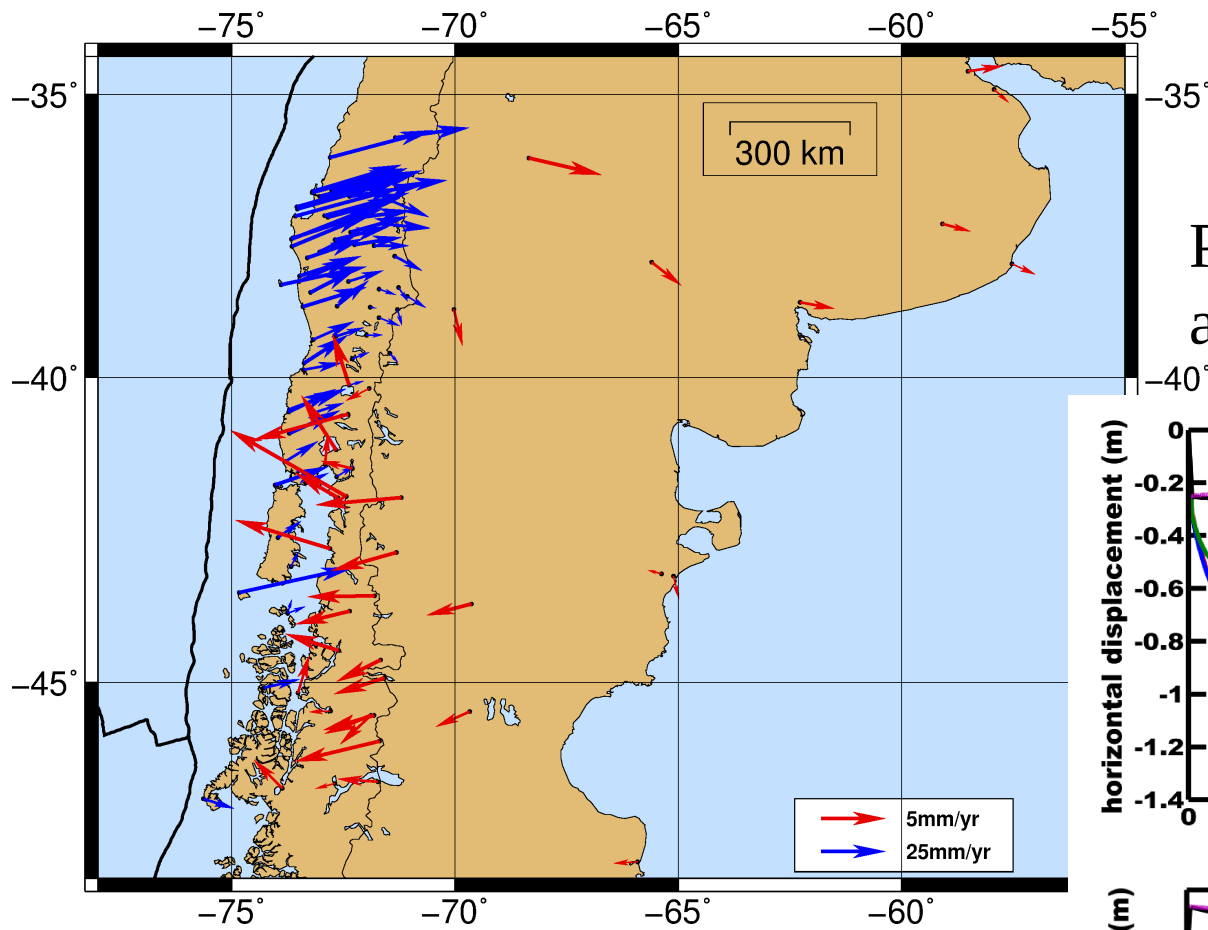
Notice eastward velocities far inland in front of Maule

# Back-slip

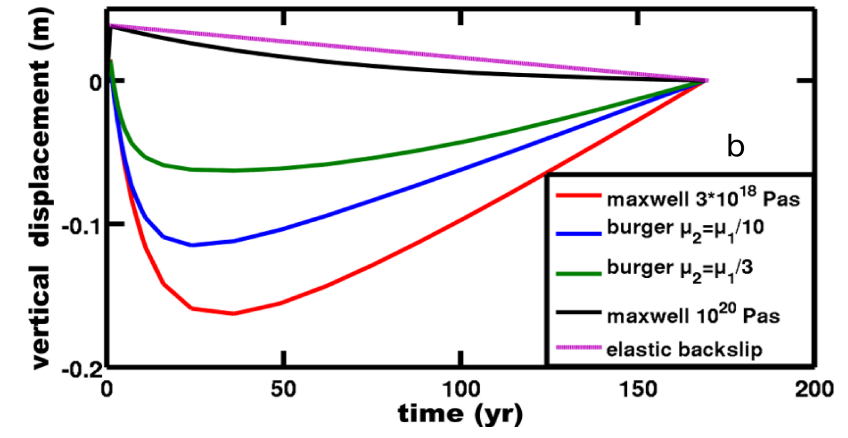
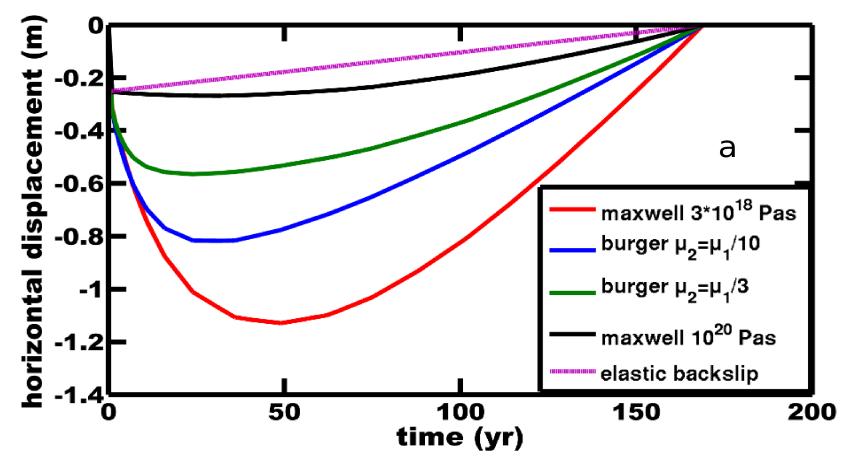


There are several equivalent methods to model the seismic cycle  
Elastic backslip implicitly assumes a large viscosity ( $>10^{21}$  Pas)  
But the asthenosphere has a low viscosity....

# Observed 40yrs after Valdivia, before Maule (referential of North-East South America)

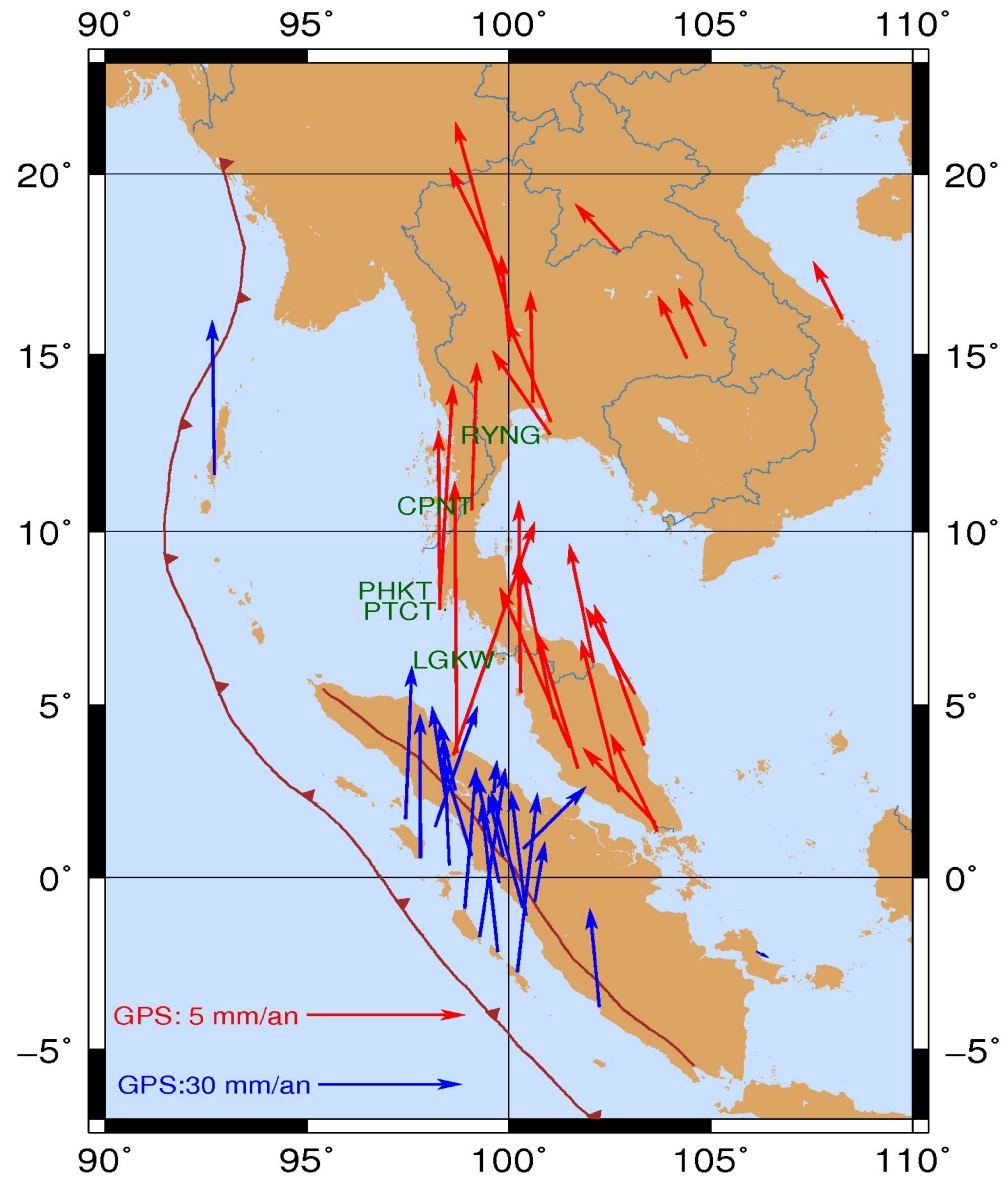


Predicted displacements at a distance of 700km



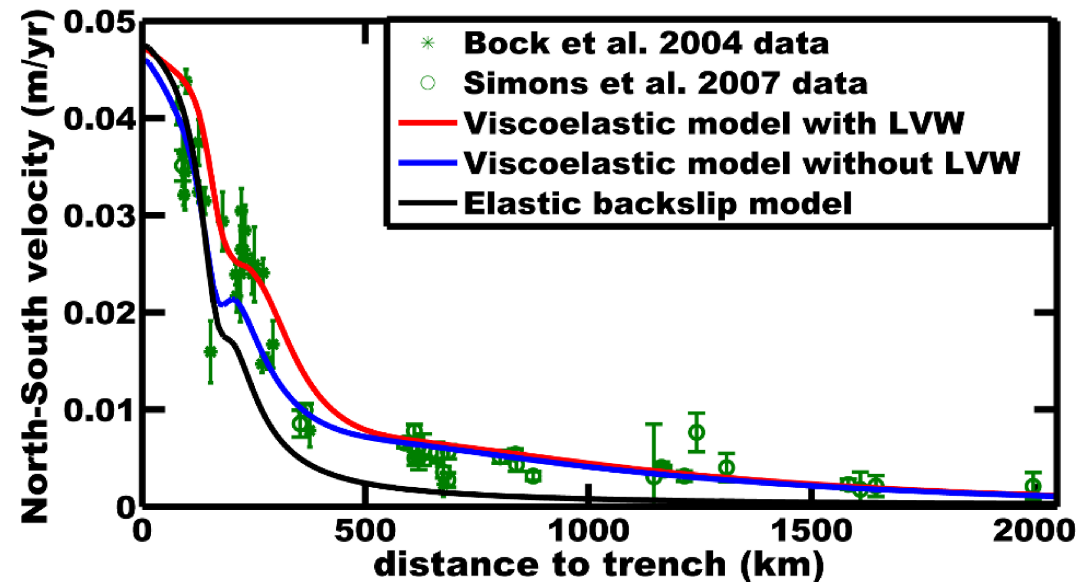
The velocity and strains are almost never equal to the long-term average

# Are the present-day strains 'Geologic' strains or effect of the seismic cycle?



Observed velocities with respect to south China before 2004

Is there a Sunda plate distinct from South-China?

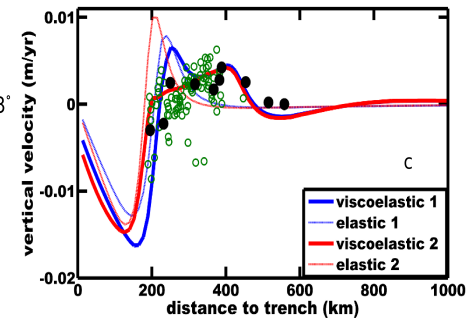
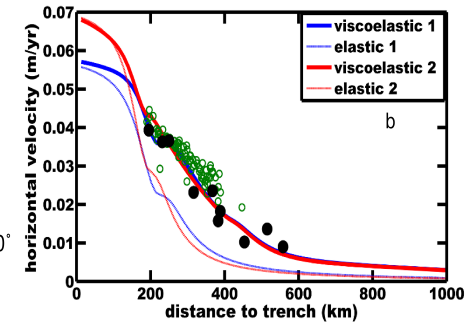




What is the long-term convergence? velocity between the Okhotsk and Amour plates?



East West



Westward velocities in Japan before Tohoku

- Before Tohoku Earthquake, Western Japan moves westward /Amour plate
- After Tohoku, Western Japan moves eastward/Amour plate

Deformations within continental Asia?