



Great Earthquakes: Observations and modeling



COLLÈGE
DE FRANCE
— 1530 —

Rupture Complexity of Great Earthquakes and its Effects on Seismic Radiation

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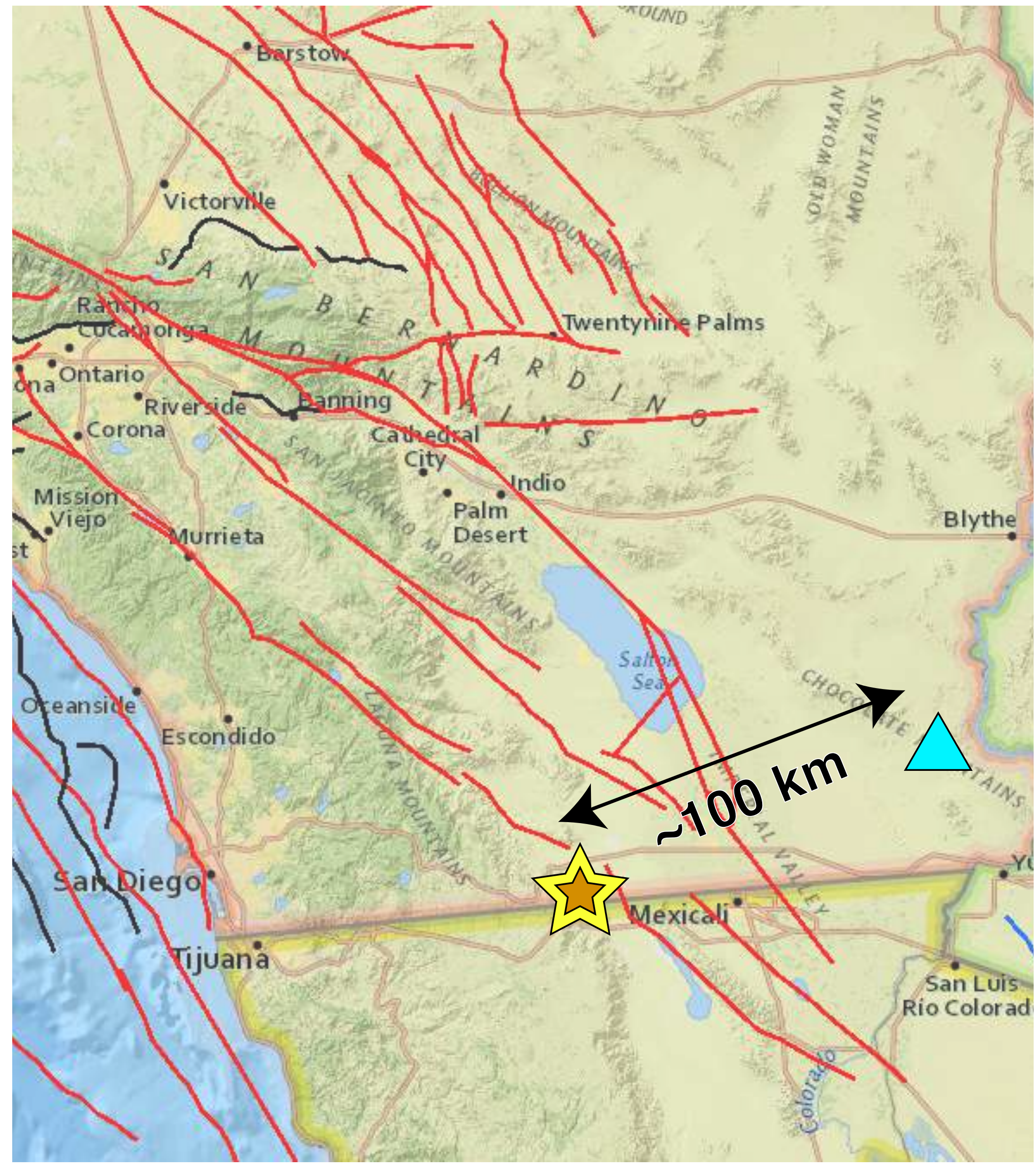
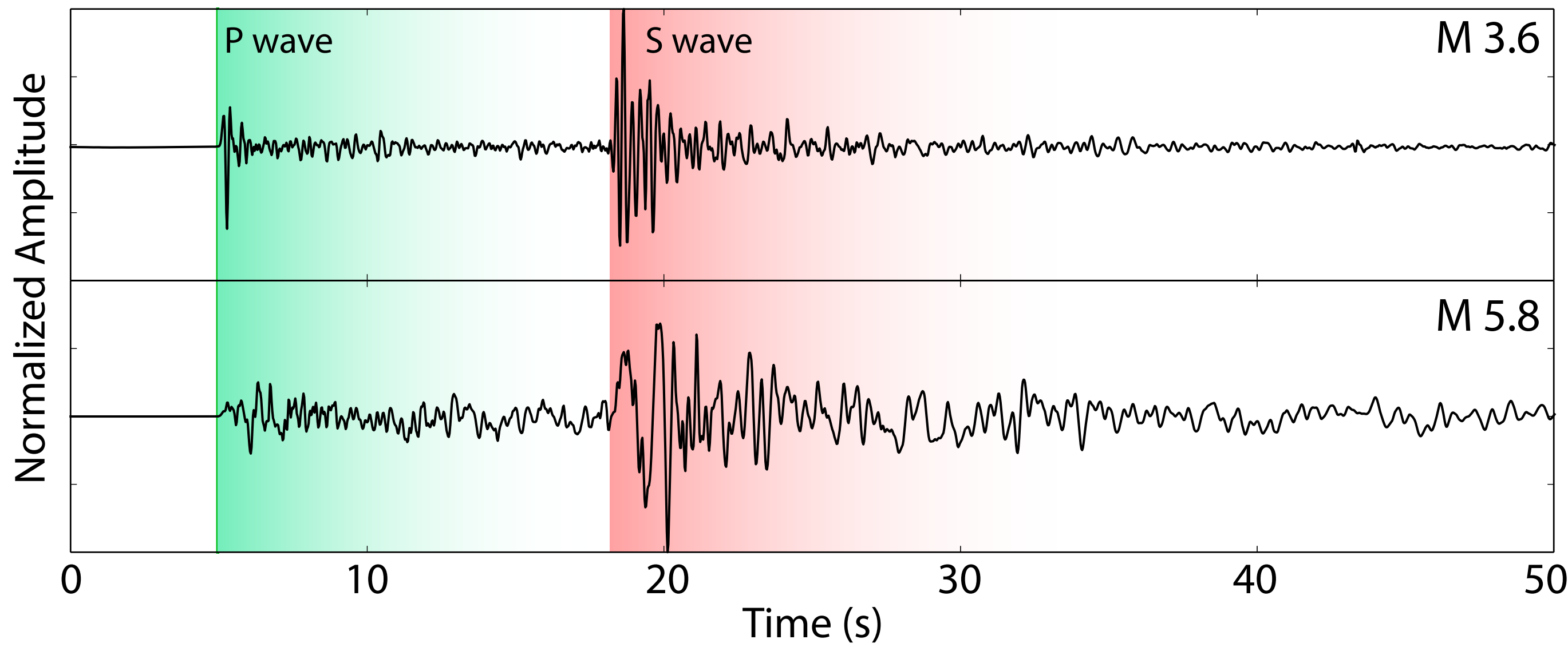
With contributions from:

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Laurent Bollinger (CEA, France)
Hiroe Miyake (ERI, Japan), Naoki Uchida (Tohoku University, Japan),
Viviana Dionicio (Servicio Geológico Colombiano)

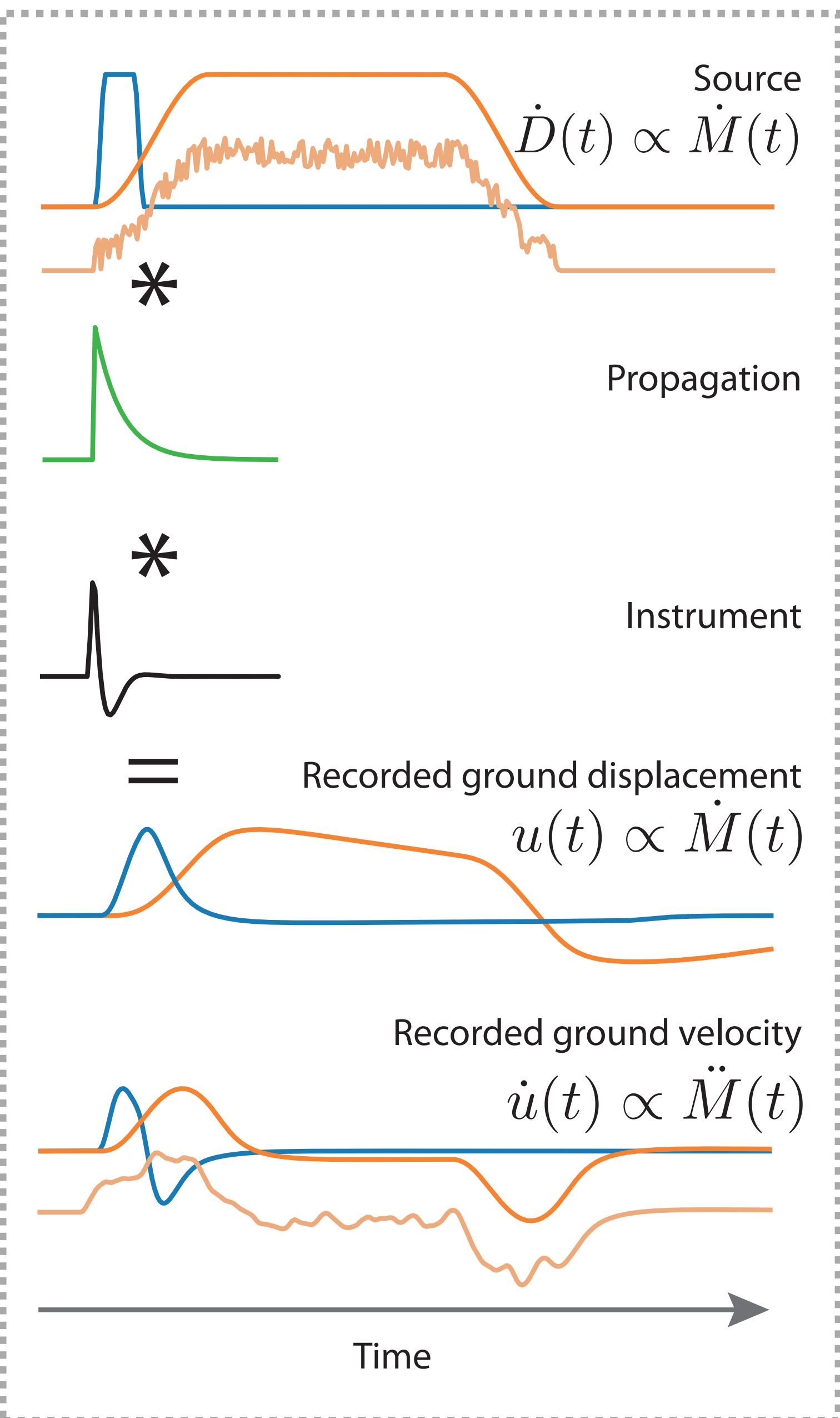
Collège de France, November 30th 2017

Seismic radiation and earthquake size

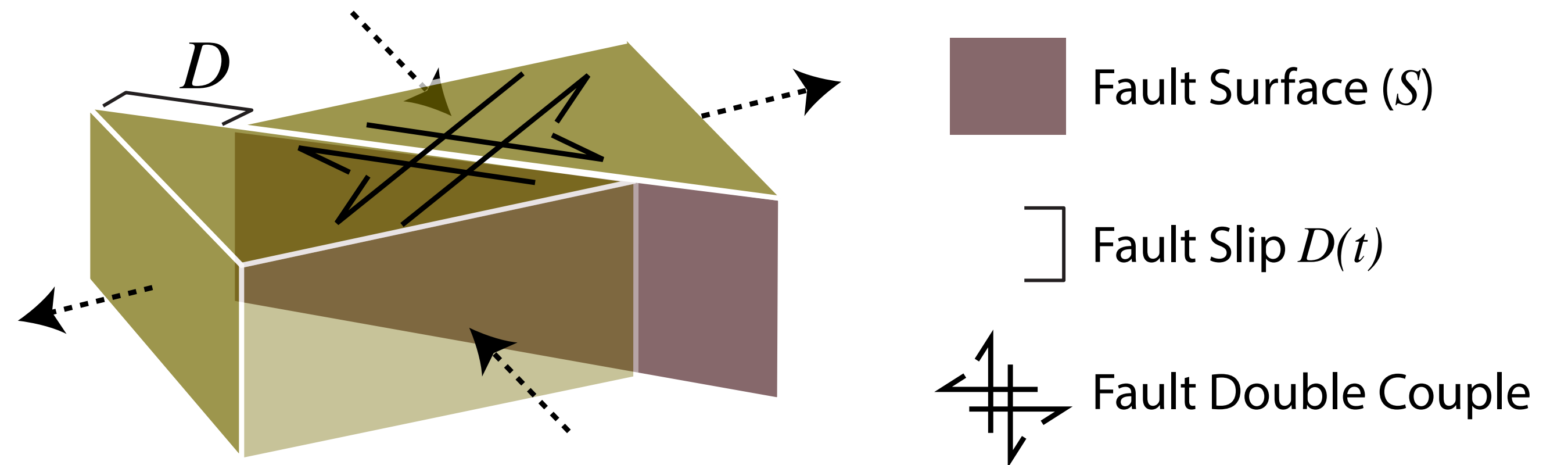
Since the early days of Seismology (1930-1940), it has been well known that **the greater the size of an earthquake, the more efficiently longer-period waves are generated.**



Far field radiation from a dislocation source



Far field ground displacement is proportional to slip (moment) rate on the fault



Seismic moment (Aki, 1966)

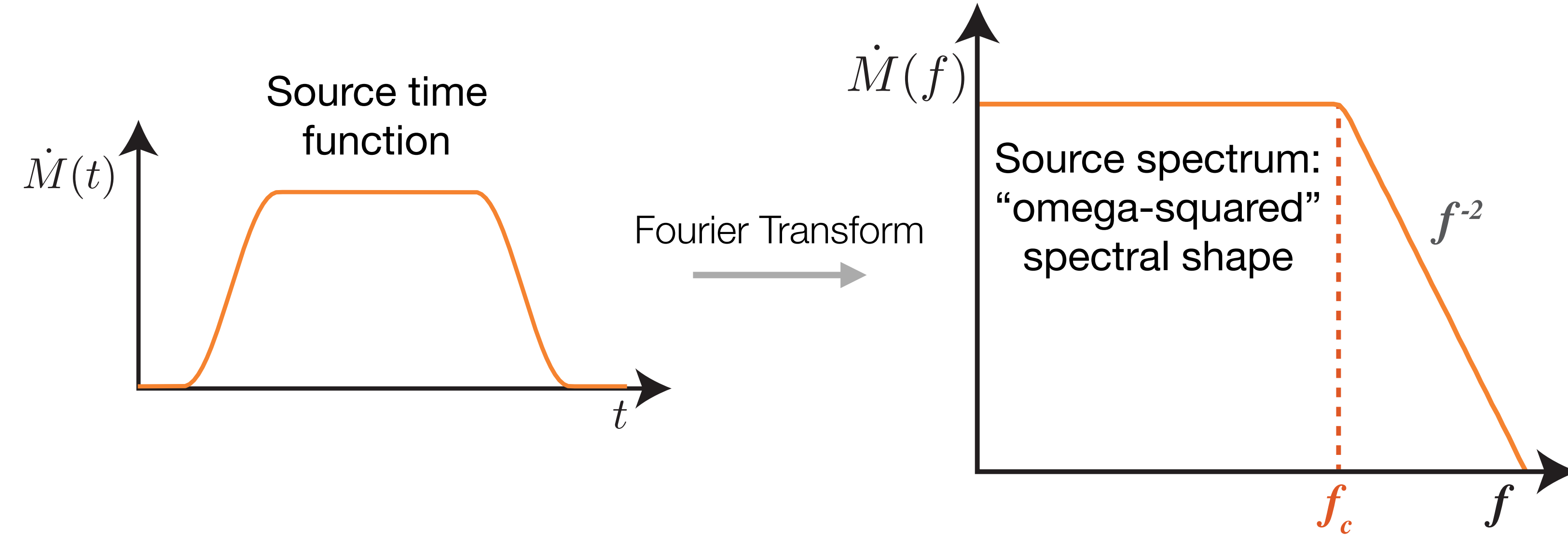
$$M_0 = \mu S D(t \rightarrow \infty) \quad (\mu : \text{rigidity})$$

e.g.,
 Knopoff and Gilbert (1960)
 Haskell (1964, 1966)
 Aki (1967)

Scaling law of seismic spectrum (Aki, 1967)

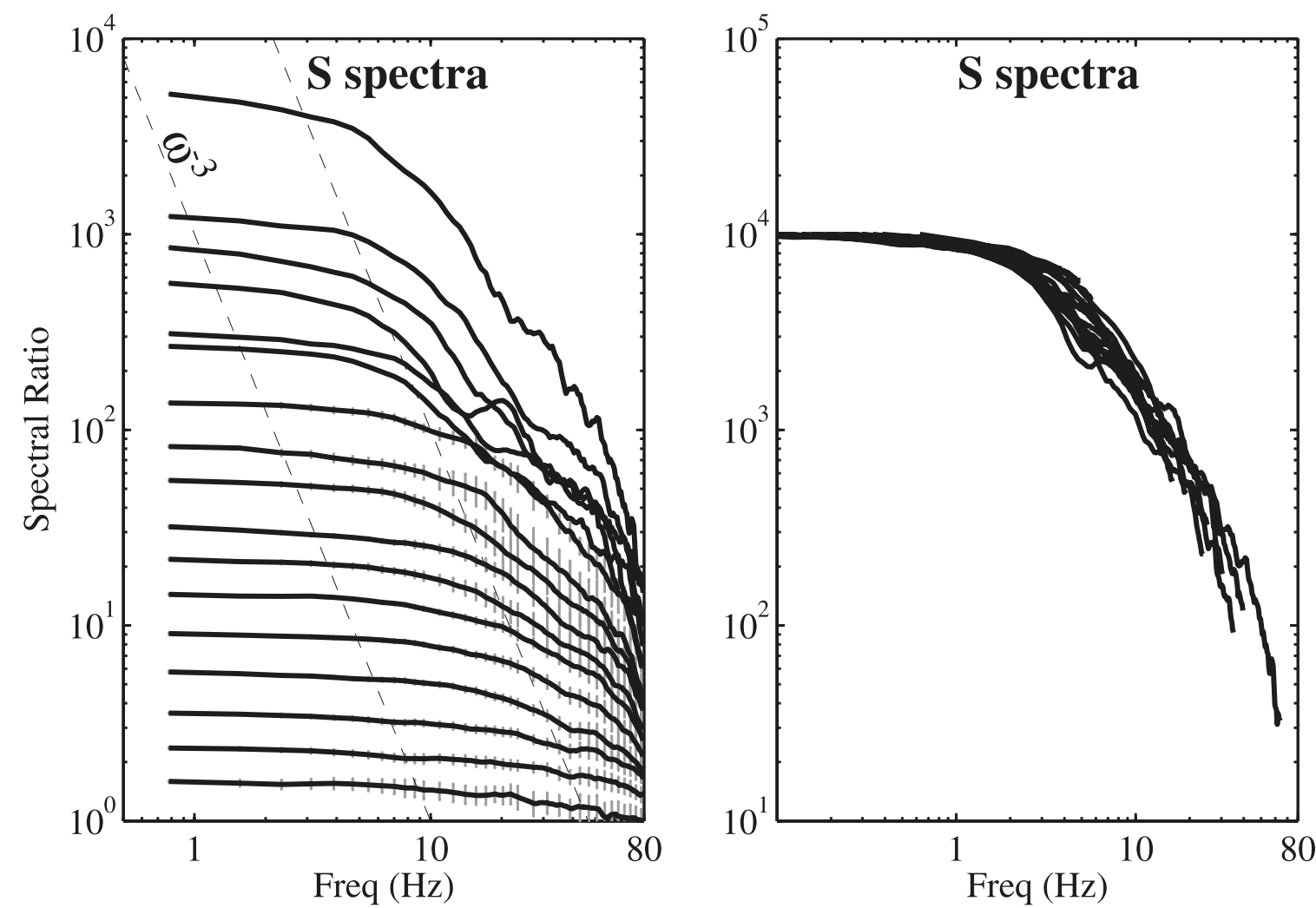
1. The seismic spectrum has a characteristic “omega-squared” shape:

$$\omega := 2\pi f$$

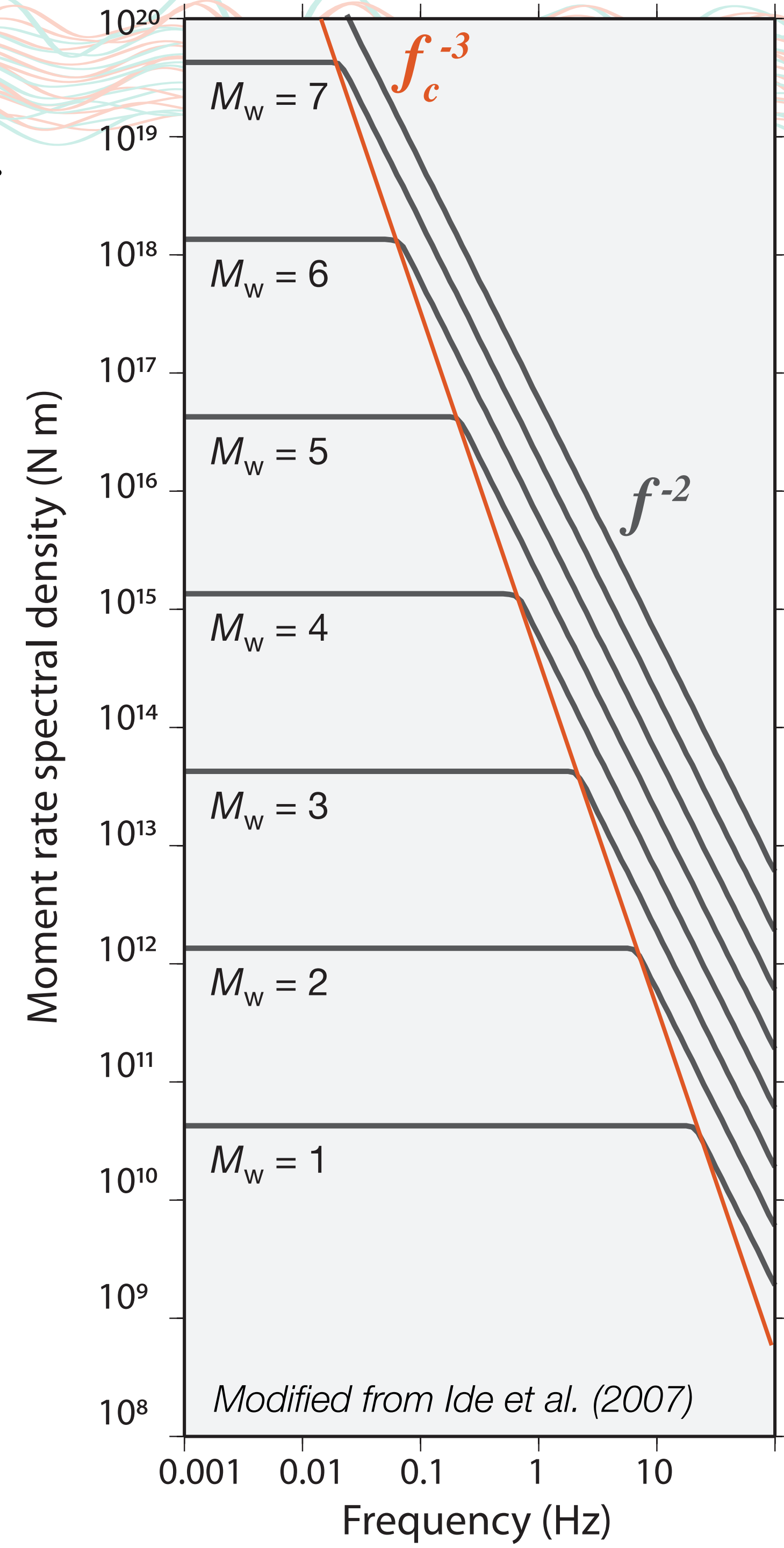


2. The seismic spectrum is **self-similar**:

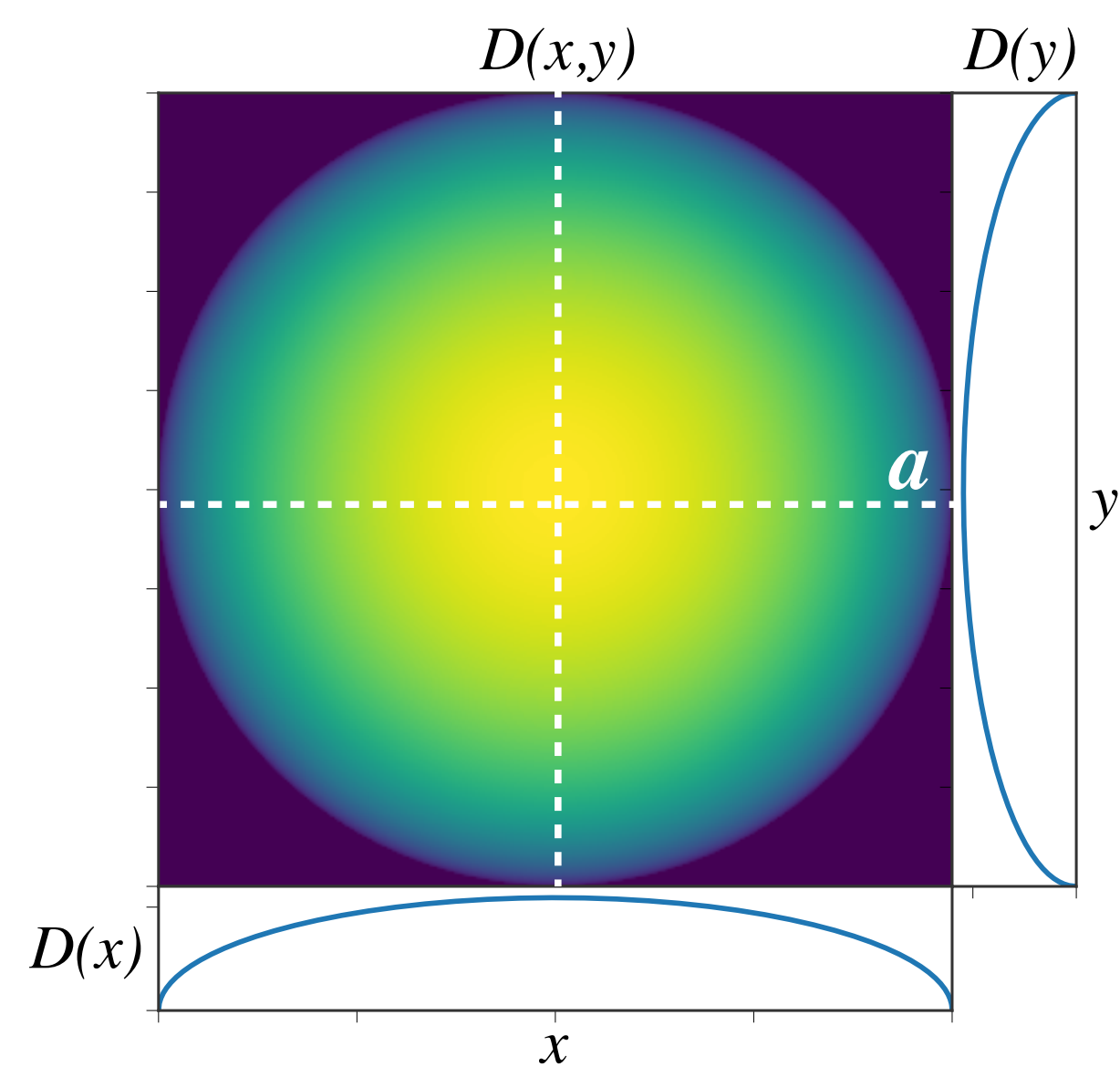
$$M_0 \propto f_c^{-3}$$



e.g., Prieto et al, 2004



Source spectrum and rupture complexity

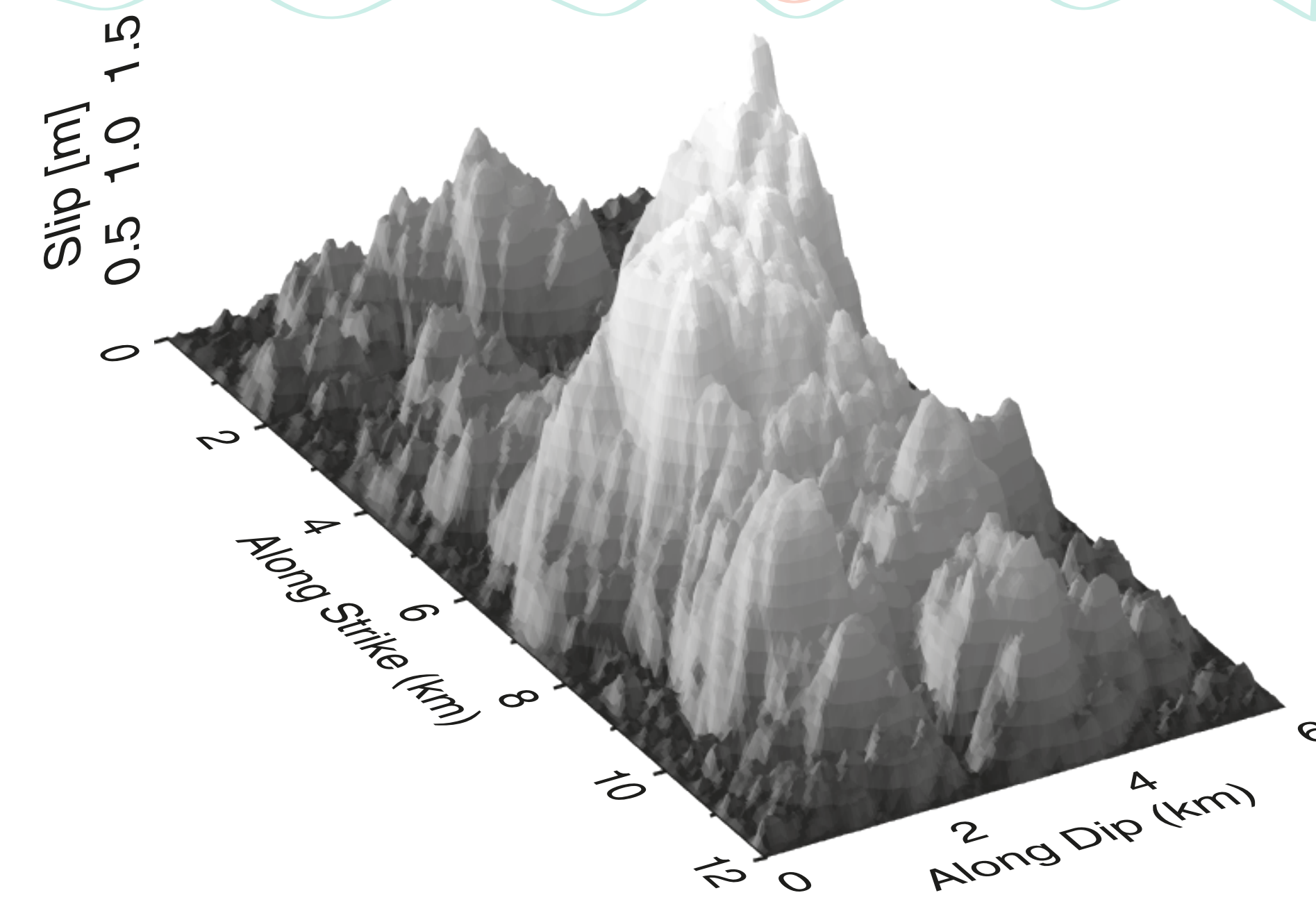
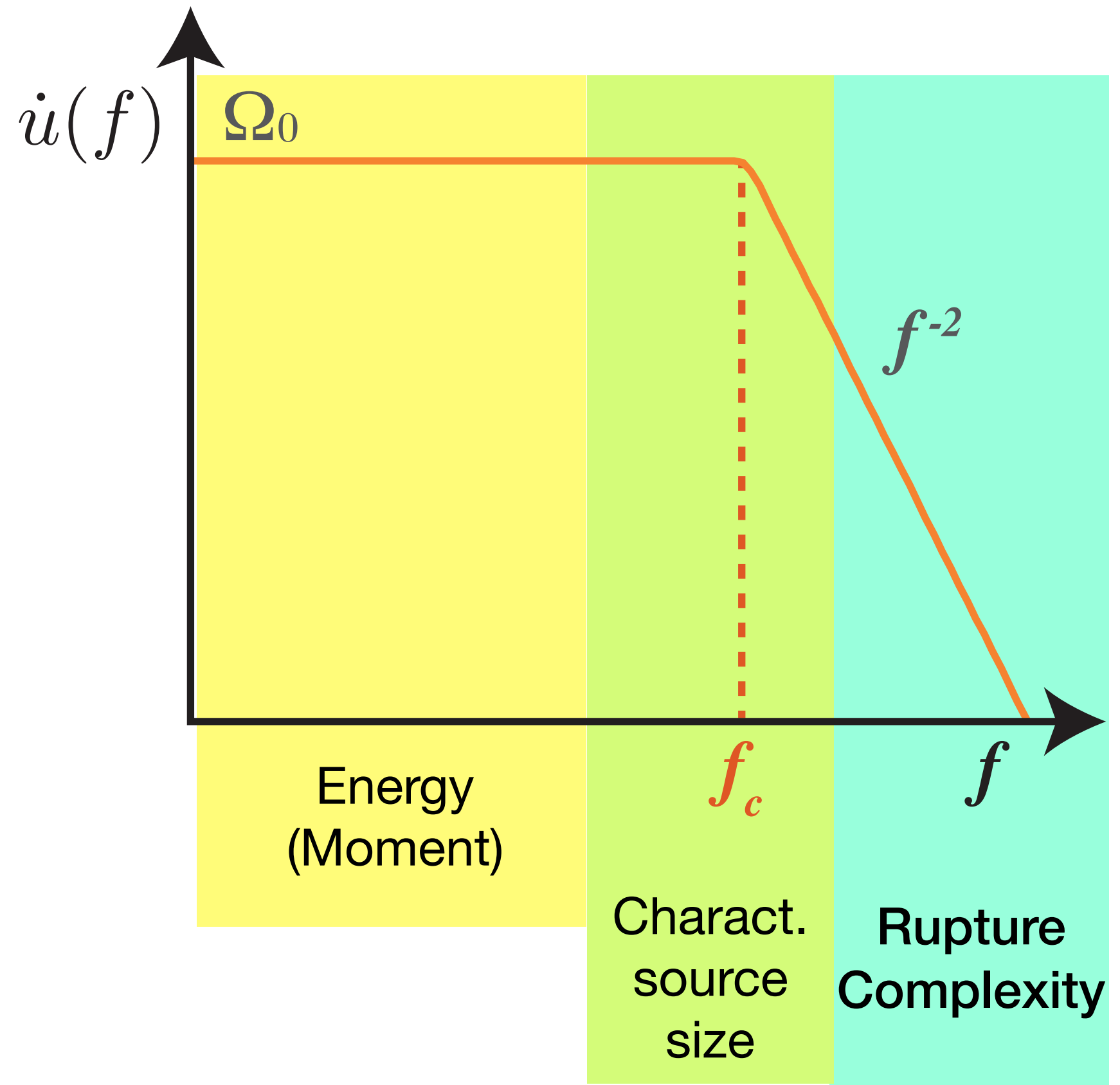


Circular crack model
(Brune, 1970; Madariaga, 1976)

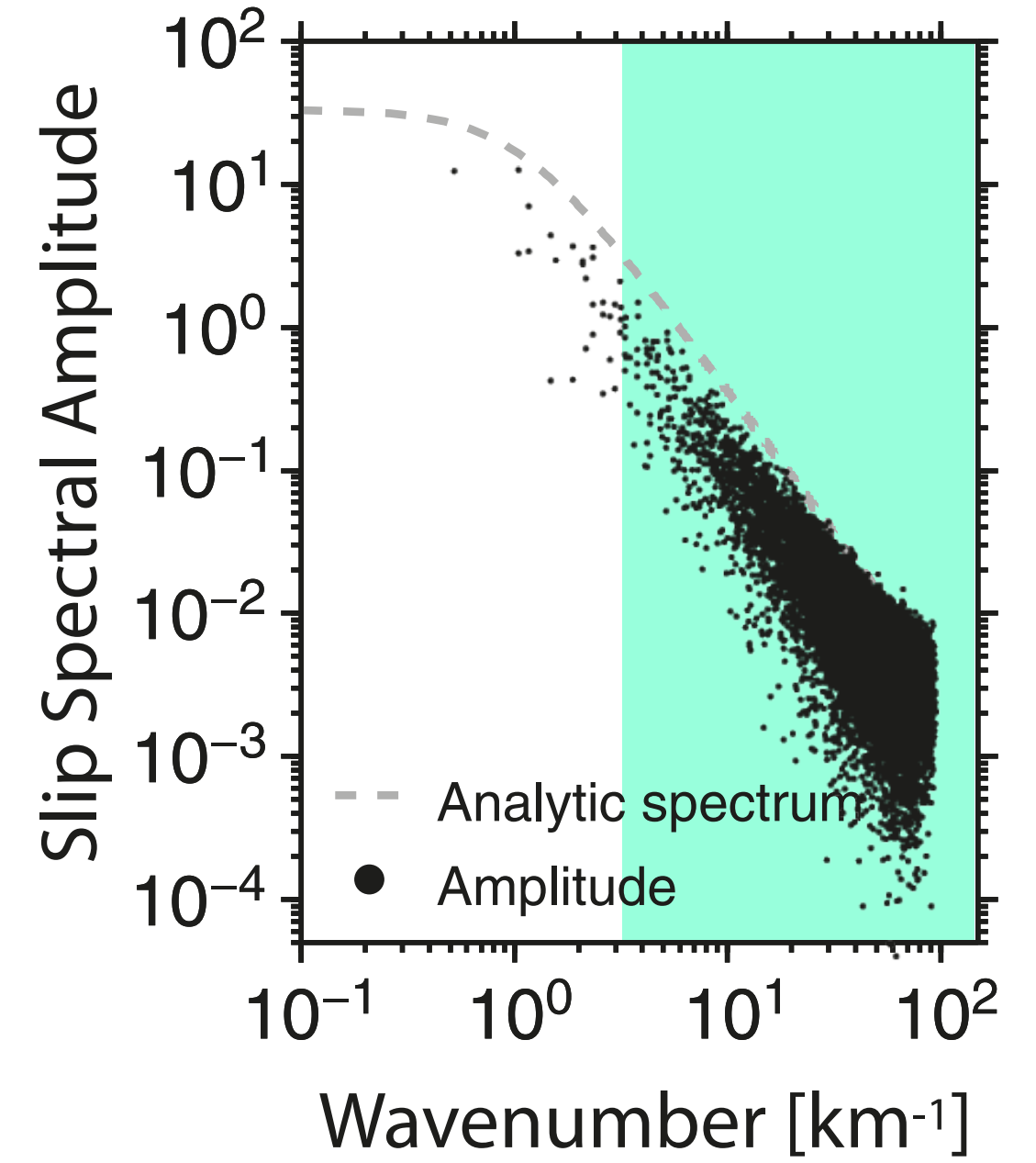
$$u(f) \propto \frac{\Omega_0}{1 + \left(\frac{f}{f_c}\right)^2}$$

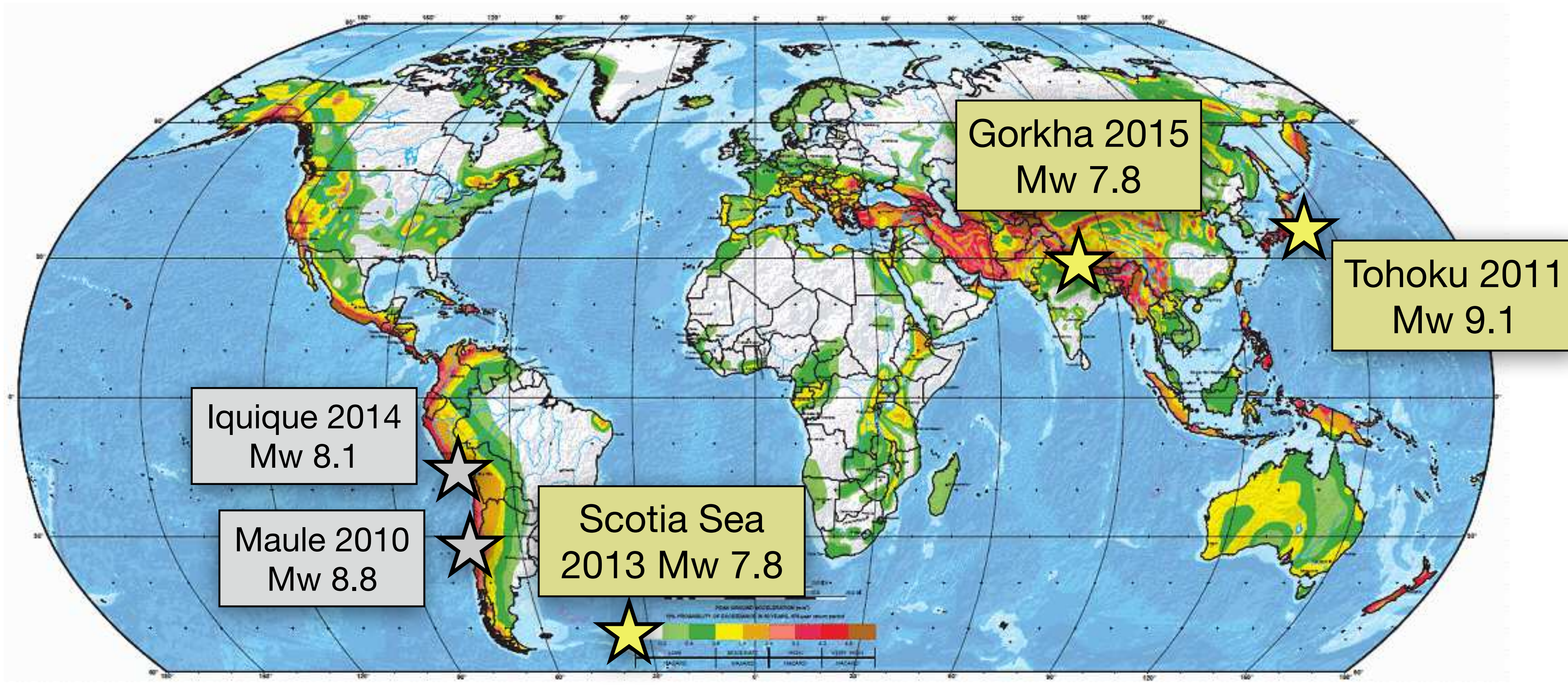
$$f_c = 0.3724 \frac{\beta}{a}$$

β = S-wave speed
 a = source radius

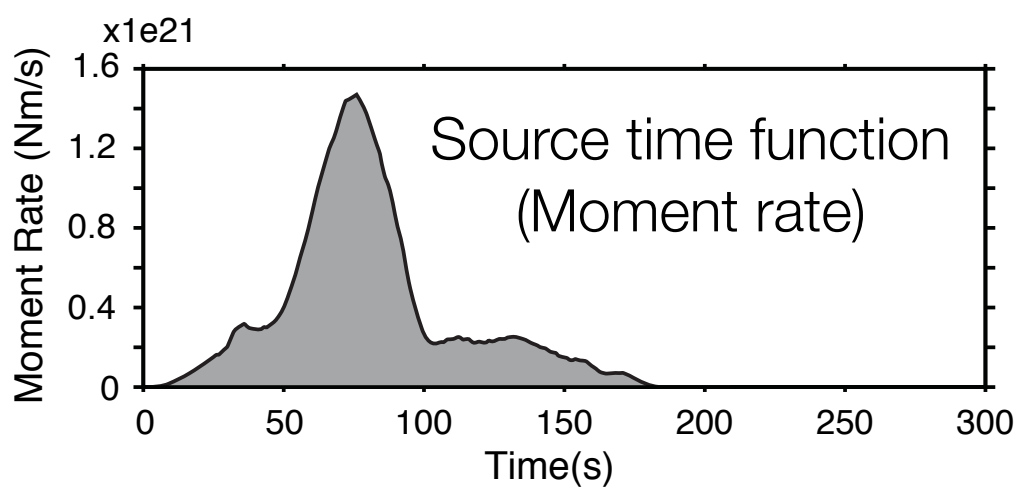
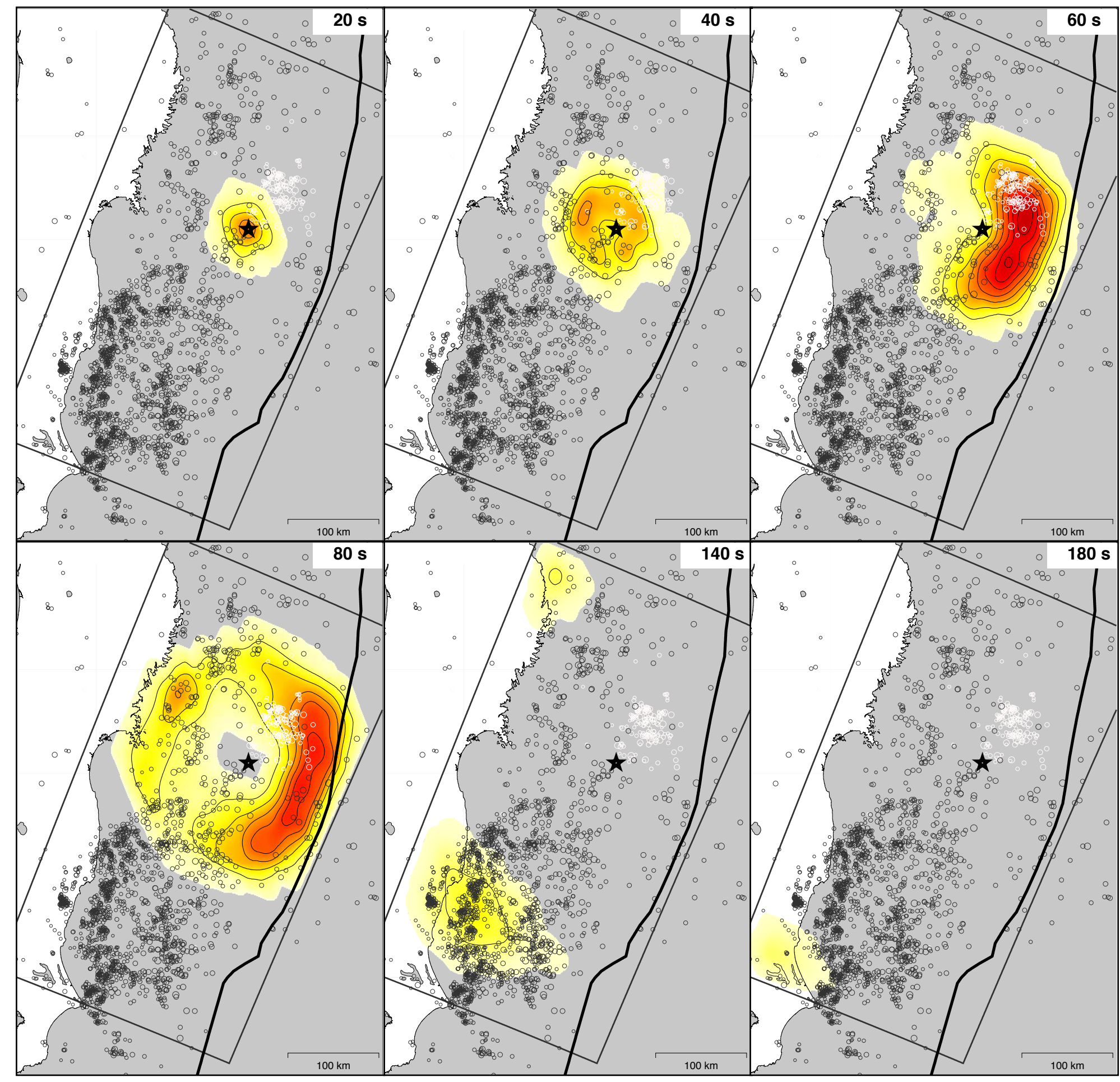
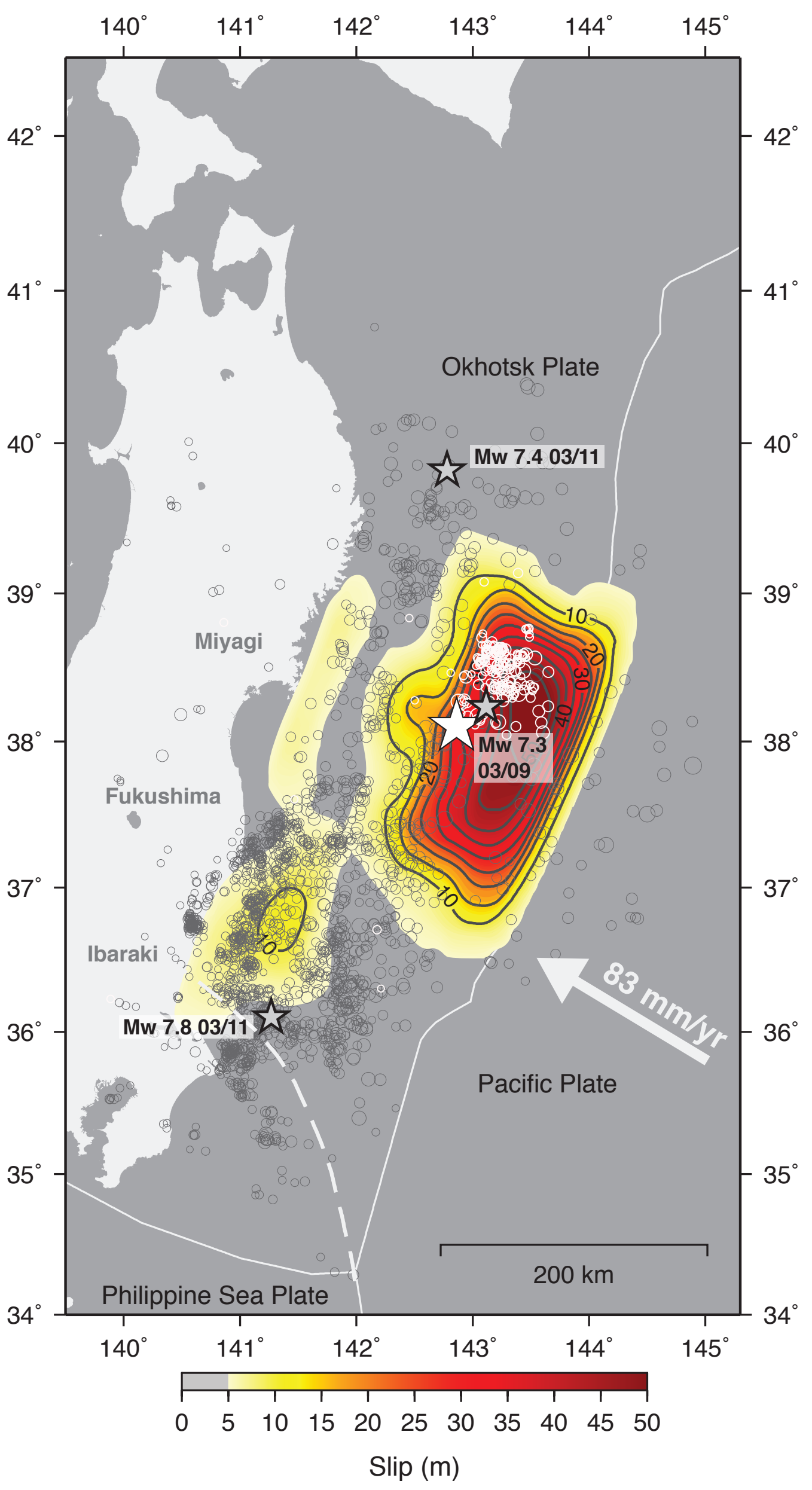


“k-squared” models:
self-similar slip on the fault
at high wavenumbers k .
(Herrero and Bernard, 1994;
Causse et al., 2010;
Ruiz et al., 2011;)



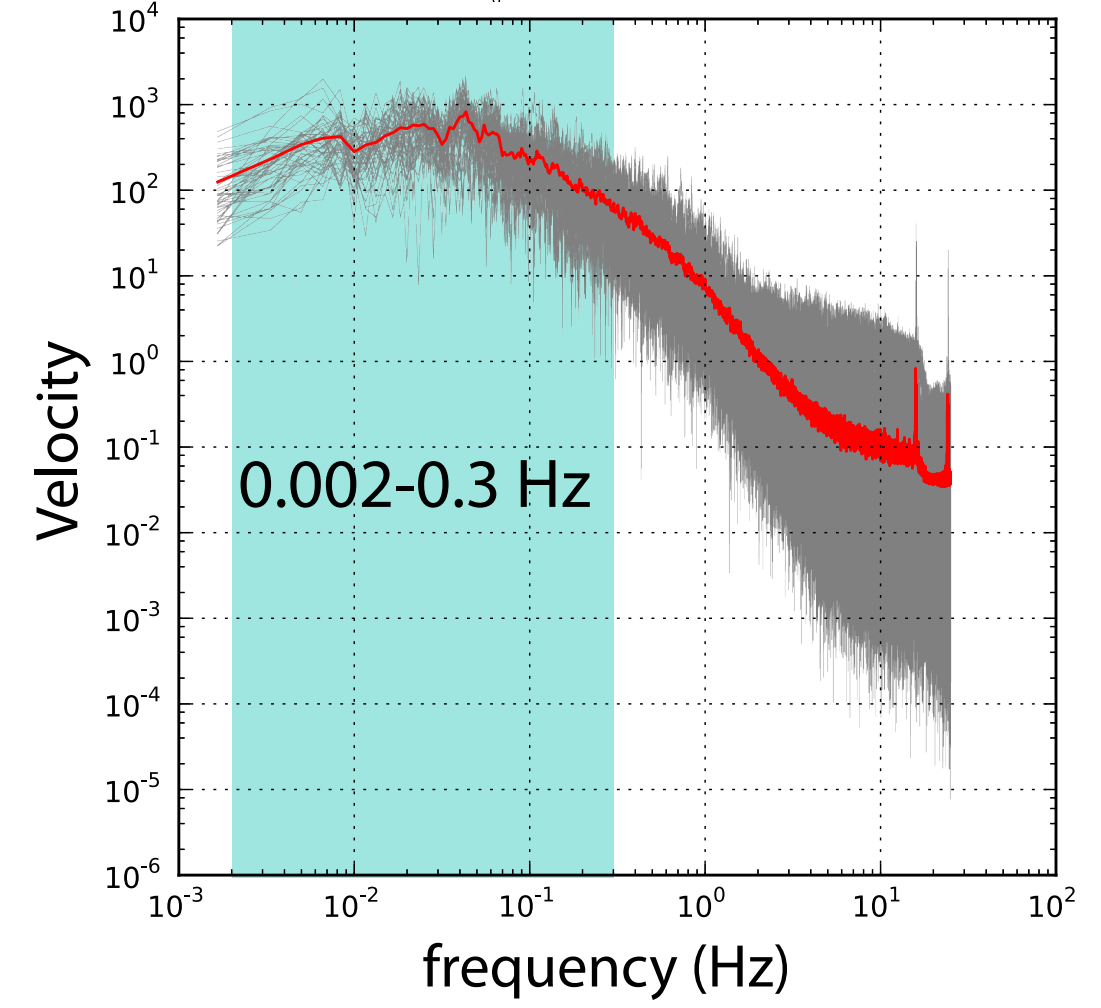
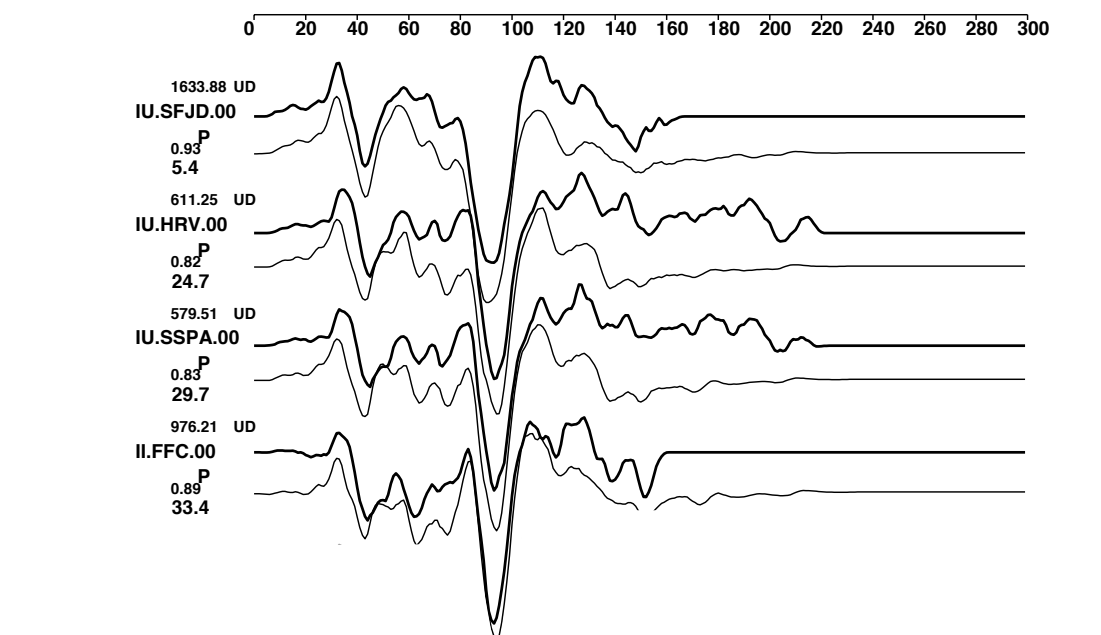


The 2011 Tohoku earthquake (Mw 9.1): slip from kinematic modeling

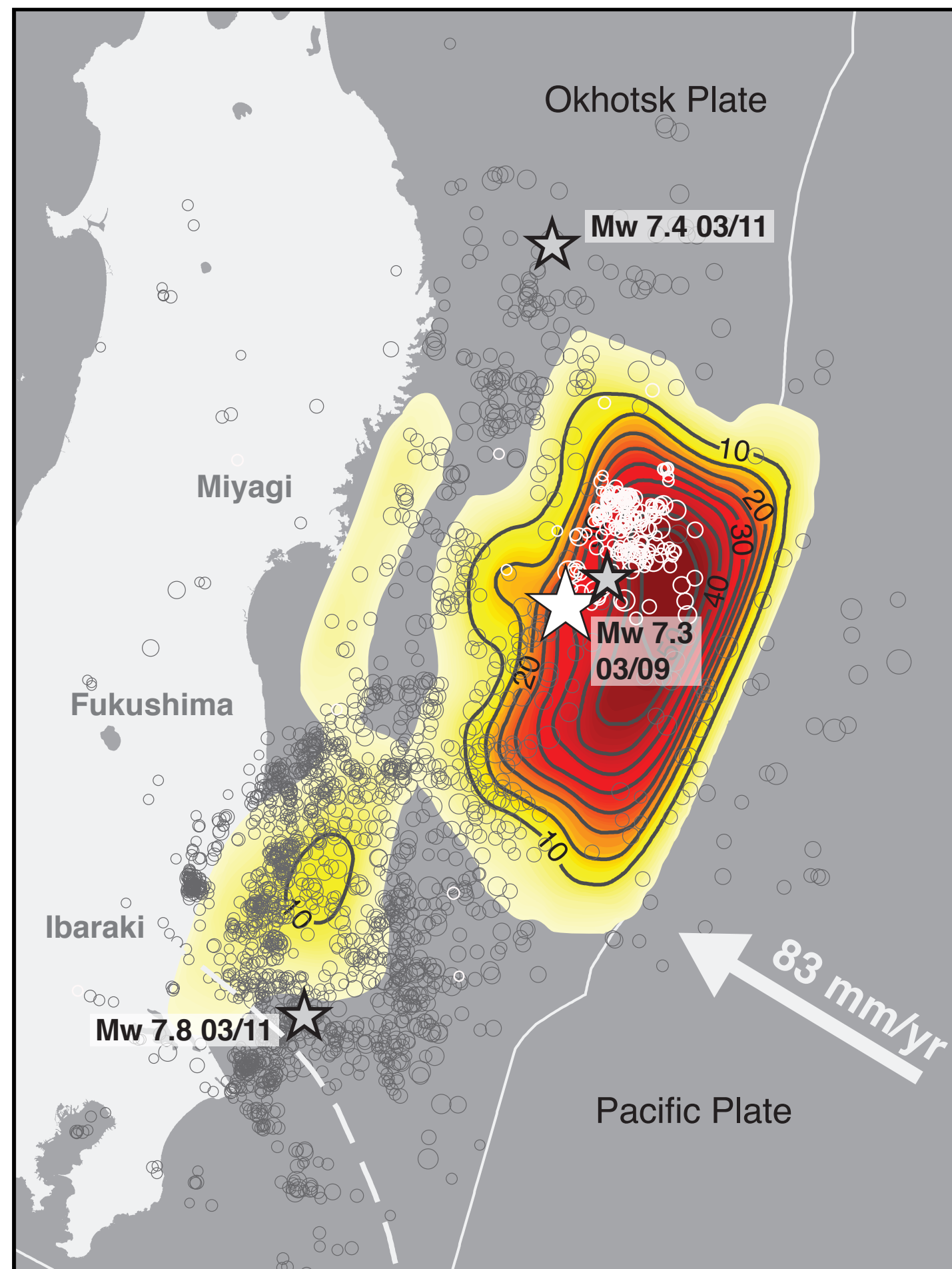


Satriano et al., EPSL (2014)

Data

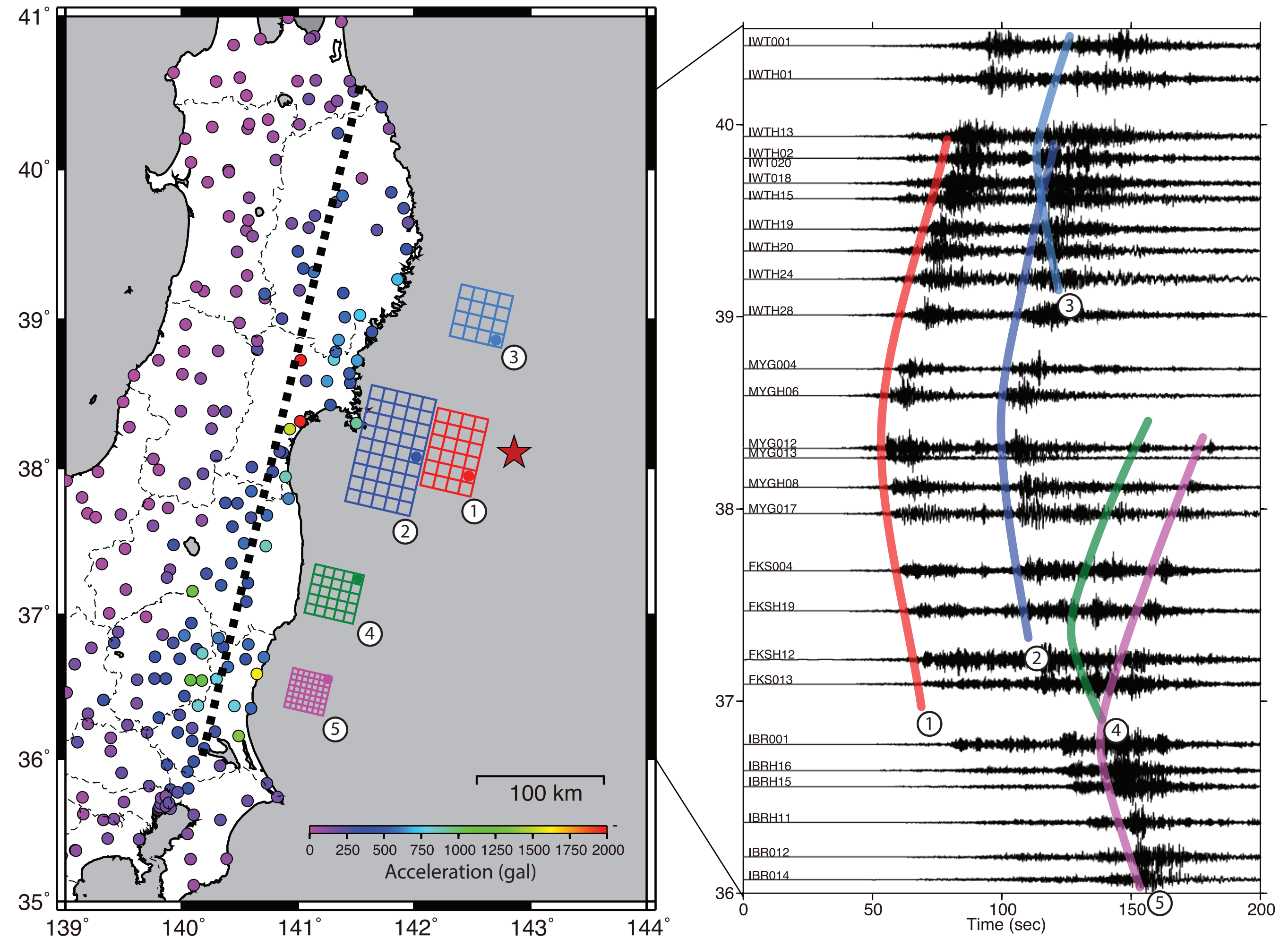


Tohoku: a complex rupture when seen through strong motion



Fault slip imaged by teleseismic data
A compact rupture, close to the trench.

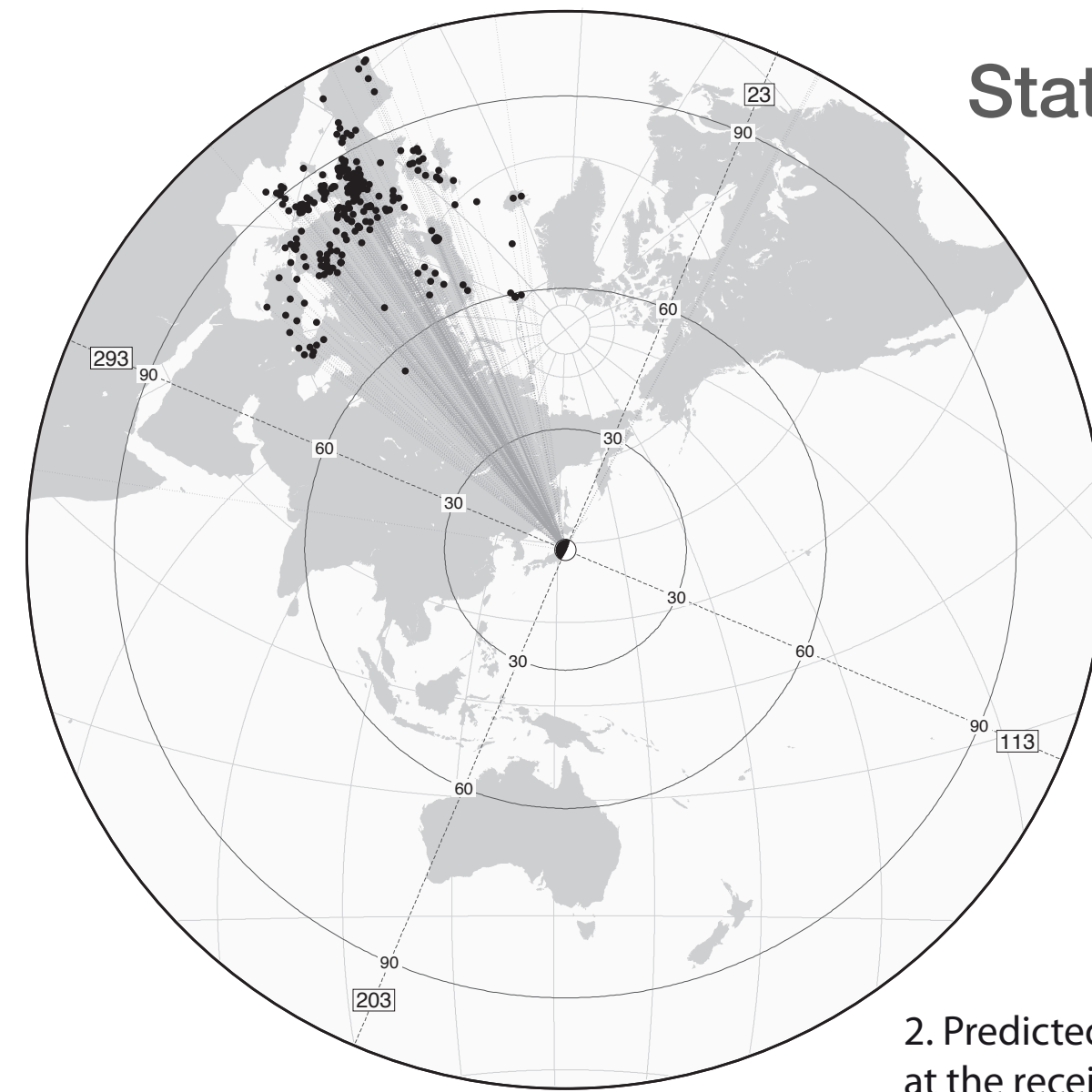
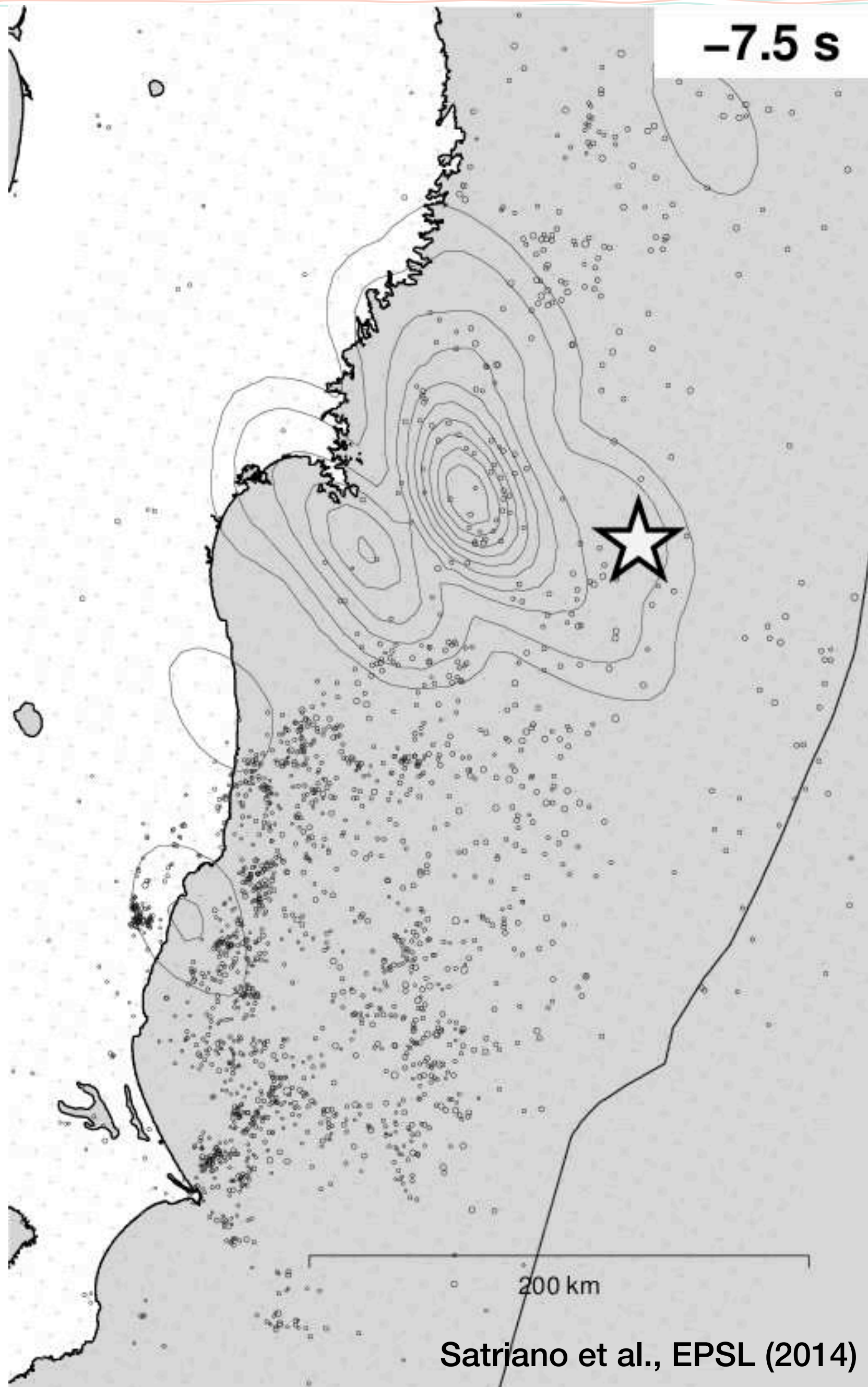
Satriano et al., EPSL (2014)



Ground acceleration at regional scale
A complex rupture: at least 5 sub-events, close to the coast.

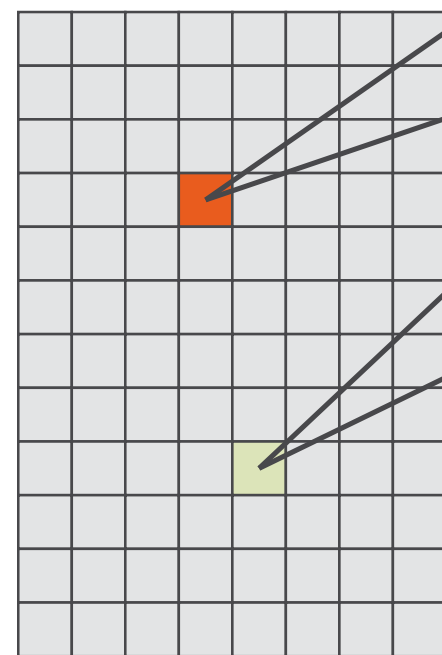
Lee et al., 2011 (redrawn)

Teleseismic imaging of high-frequency radiation: back projection

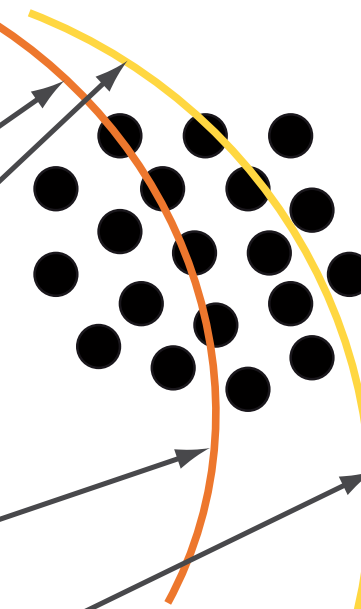


Method (Ishii et al., 2005)

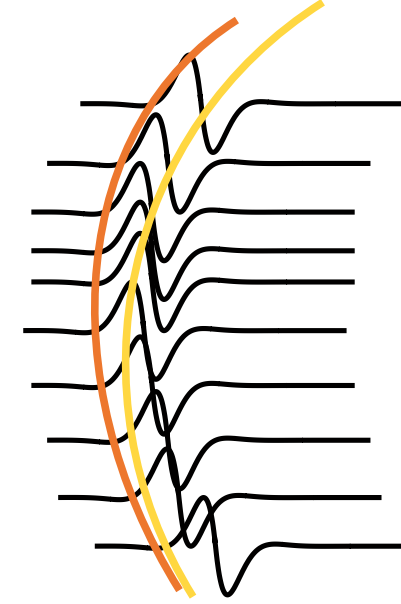
1. Grid of possible source locations



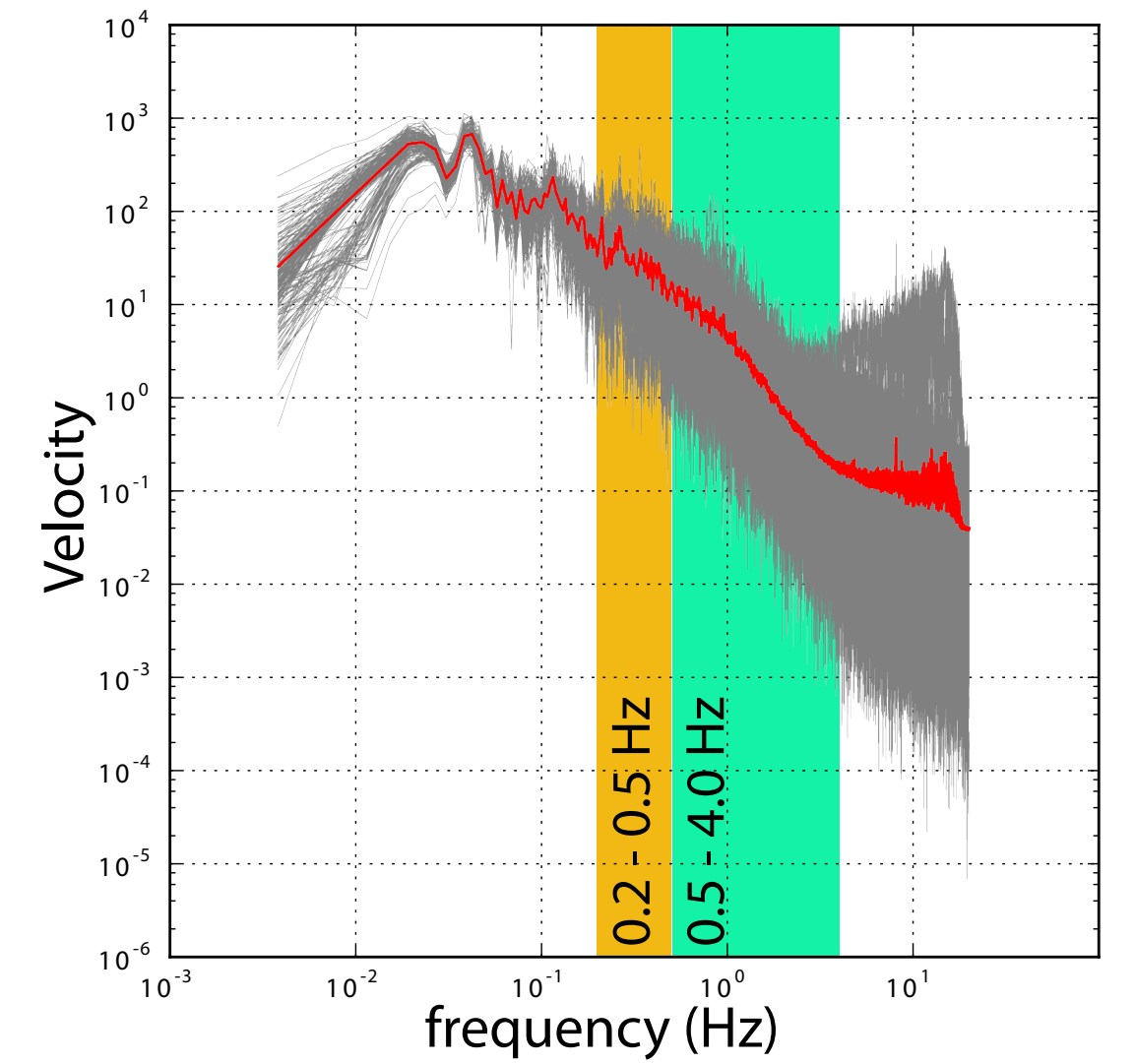
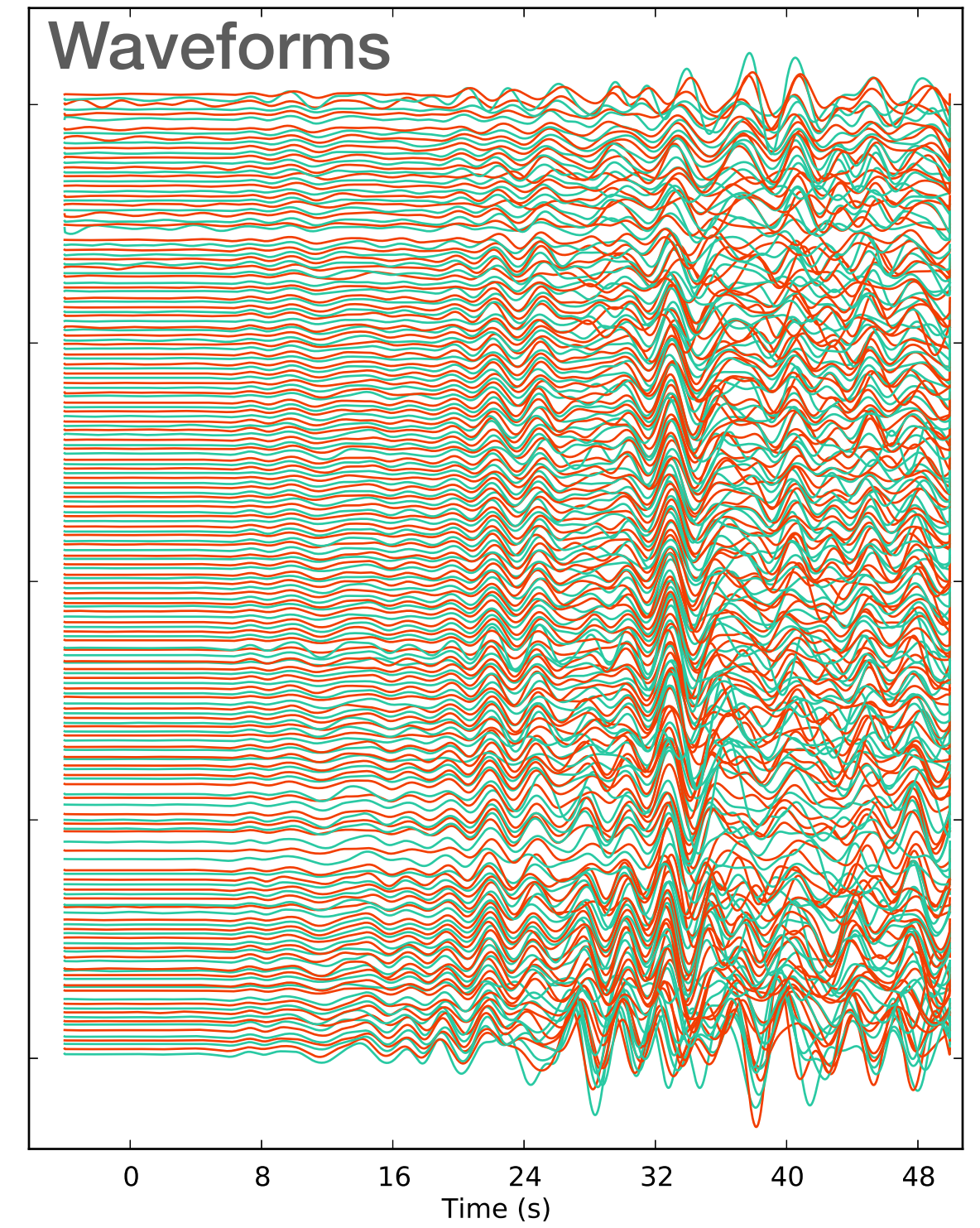
2. Predicted P-wave travel time at the receiver network



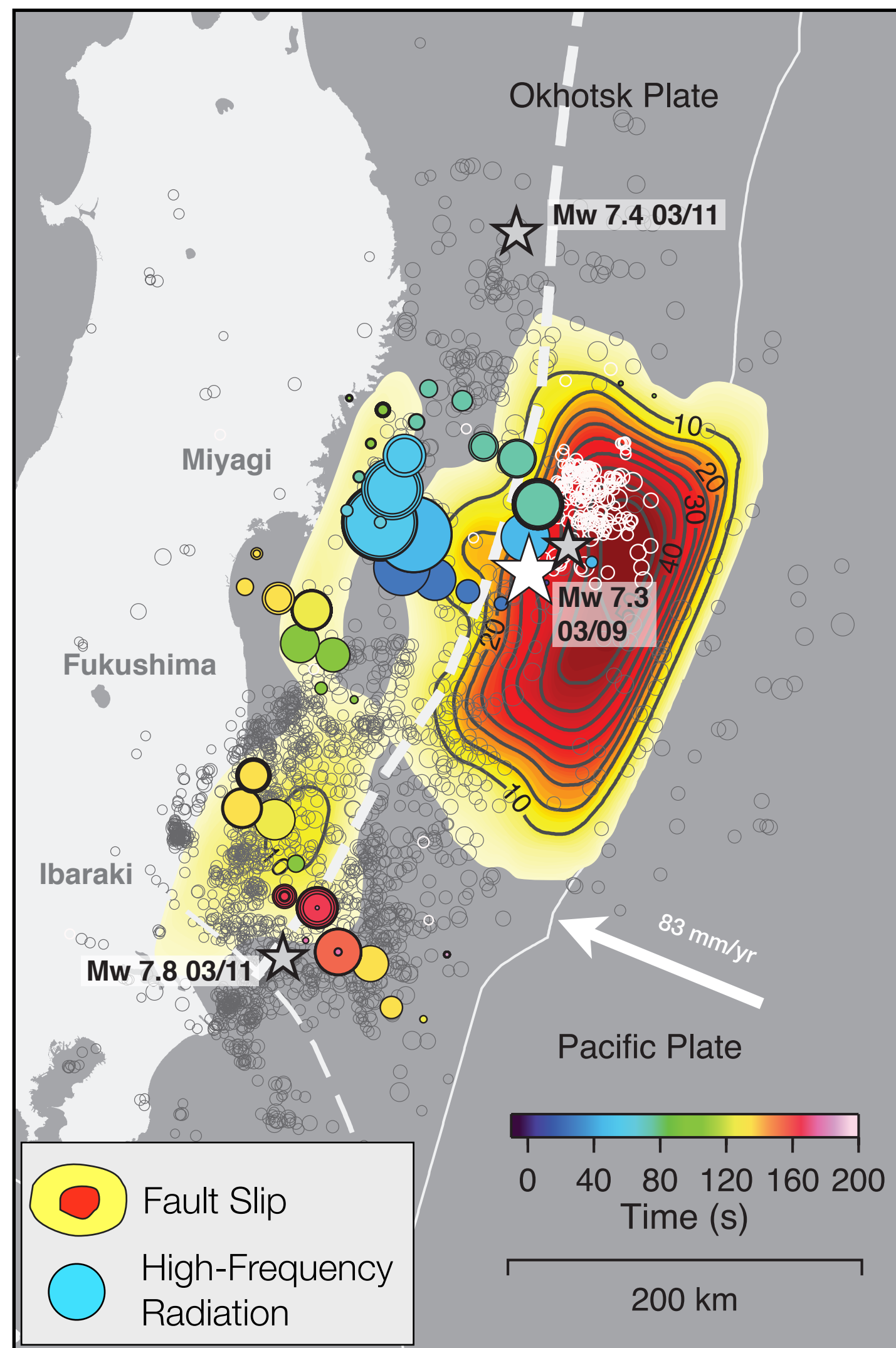
3. Stacking along the predicted travel-time curves



4. The value of the stack is assigned to the corresponding grid point

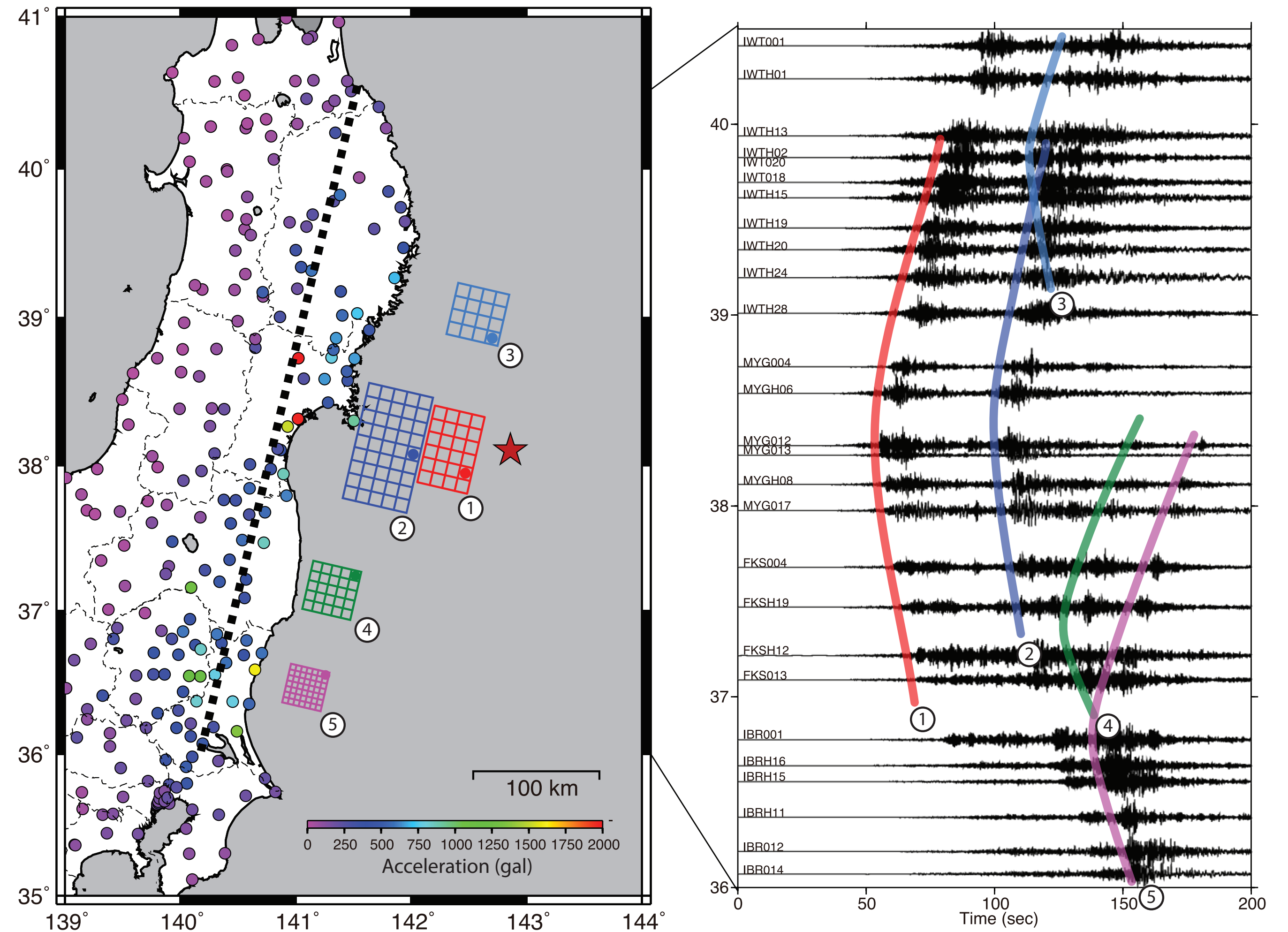


High-frequency sources and strong ground motion



Fault slip and high-frequency sources imaged by teleseismic data

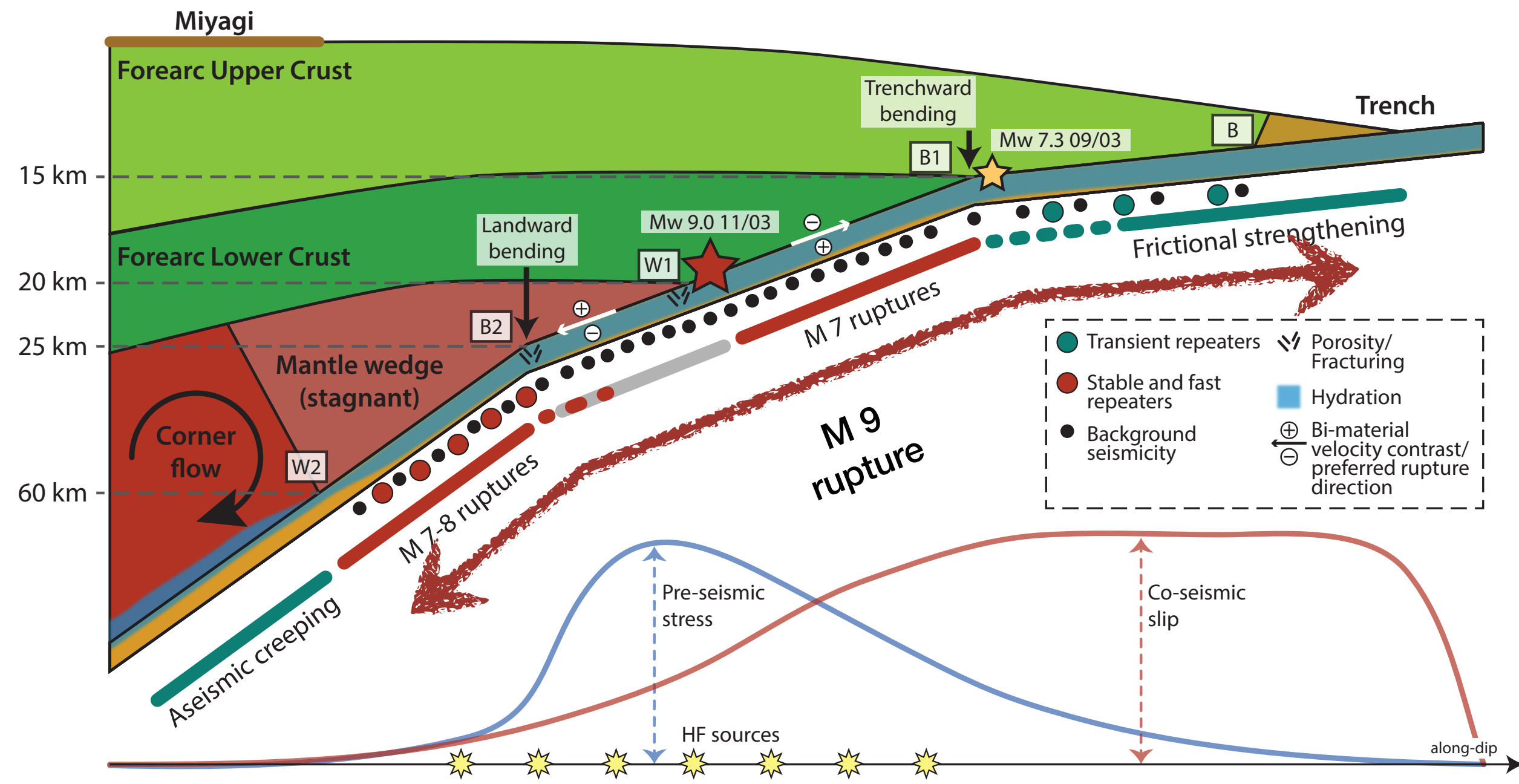
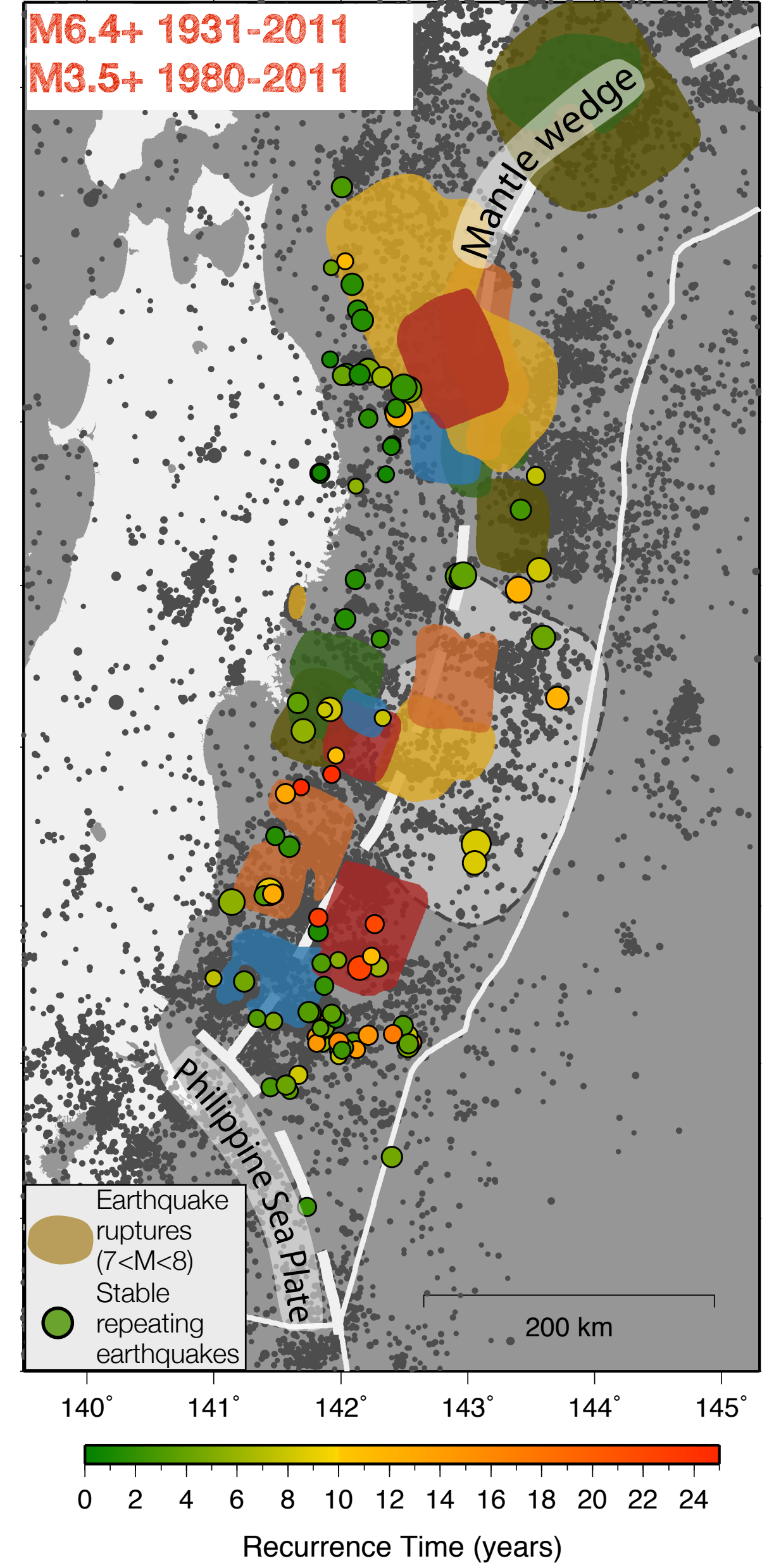
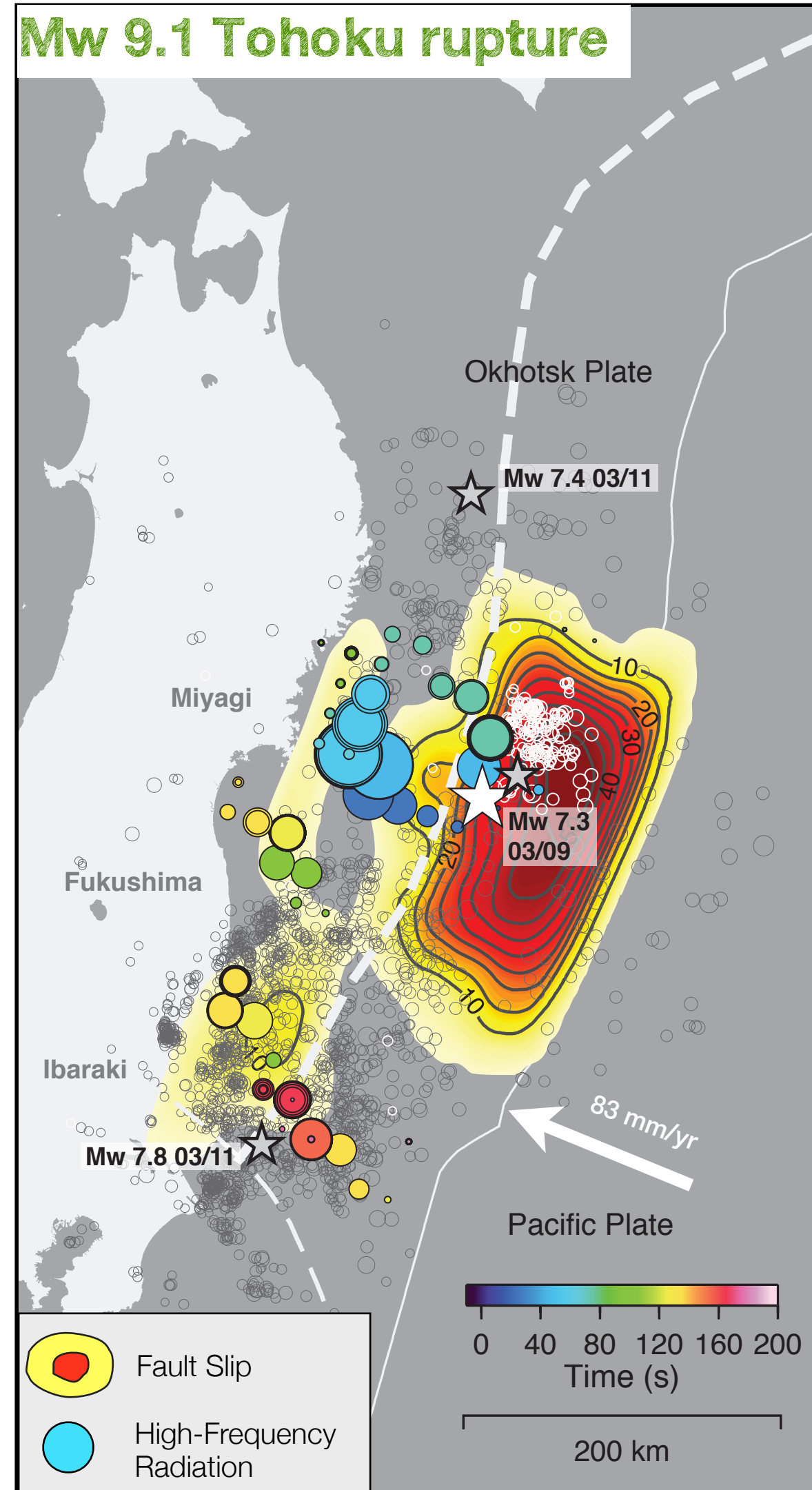
Satriano et al., EPSL (2014)



Ground acceleration at regional scale

Lee et al., 2011 (redrawn)

Northwest Japan: Structural heterogeneity and seismic asperities

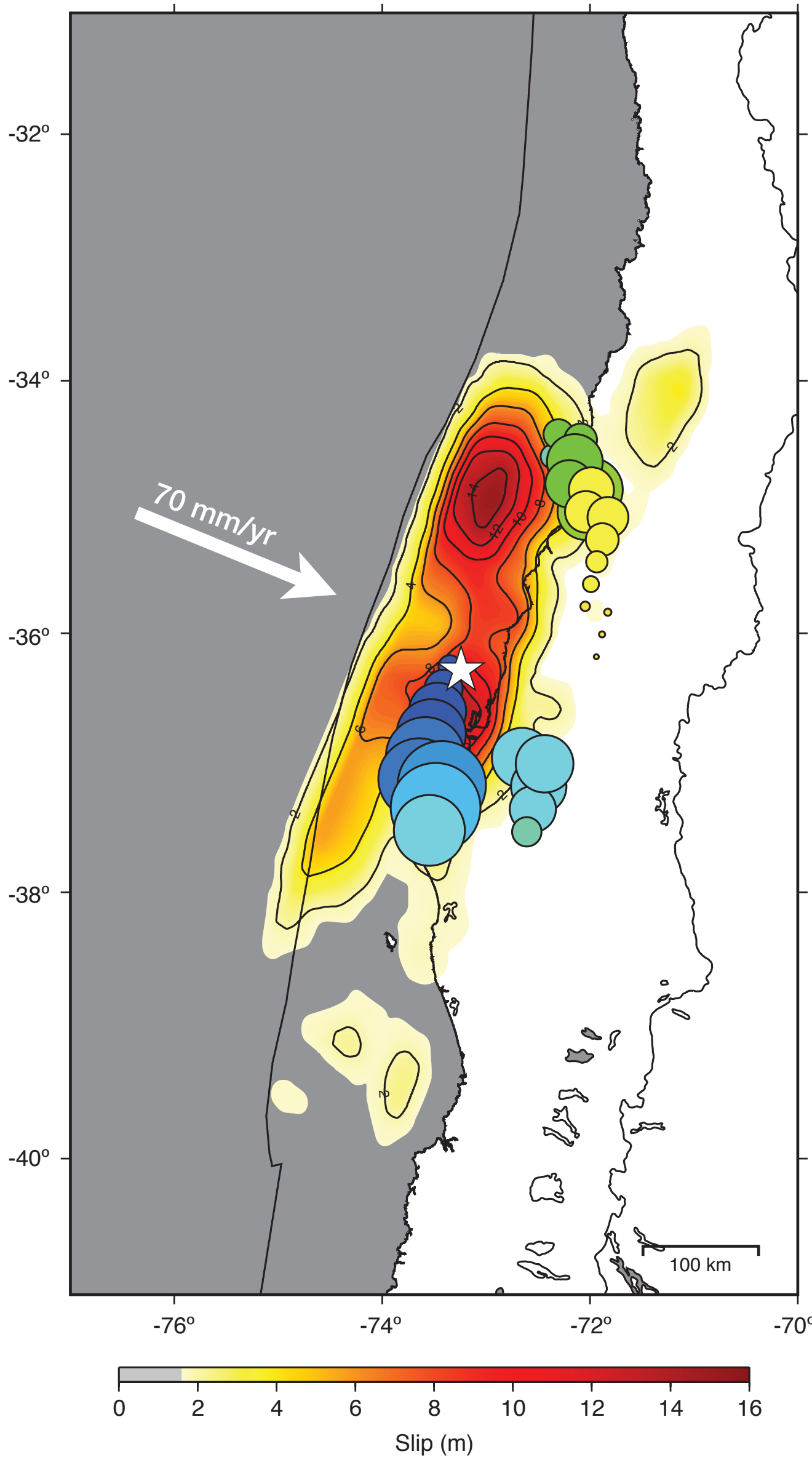


Background seismicity and large earthquake ruptures (M 7-8) inform us on the mechanical state of the plate interface.

This information is useful to forecast the rupture and radiation properties of future large earthquakes.

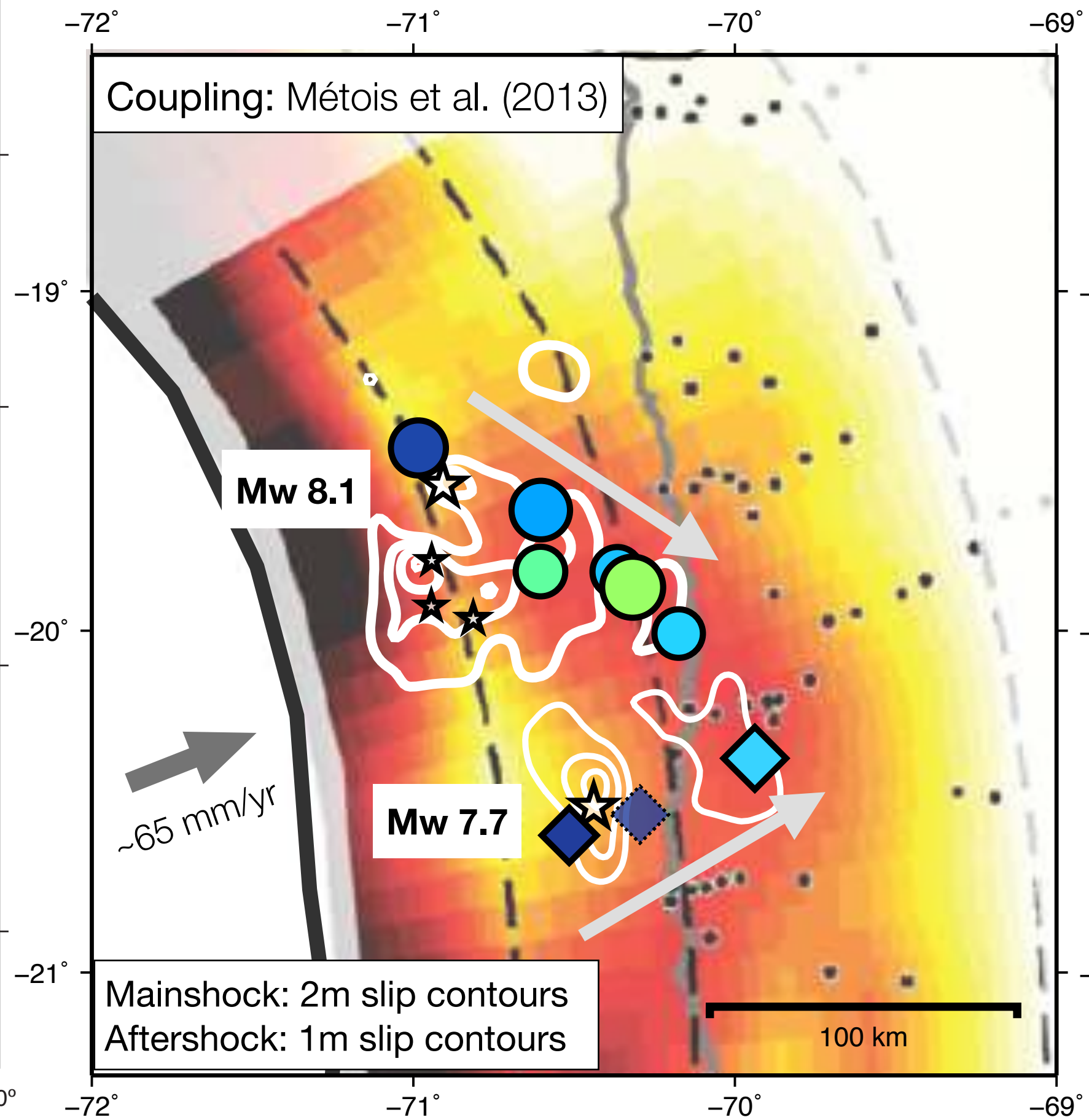
Other examples: downdip HF emission for large subduction earthquakes

Maule (Chile) 2010 Mw 8.8



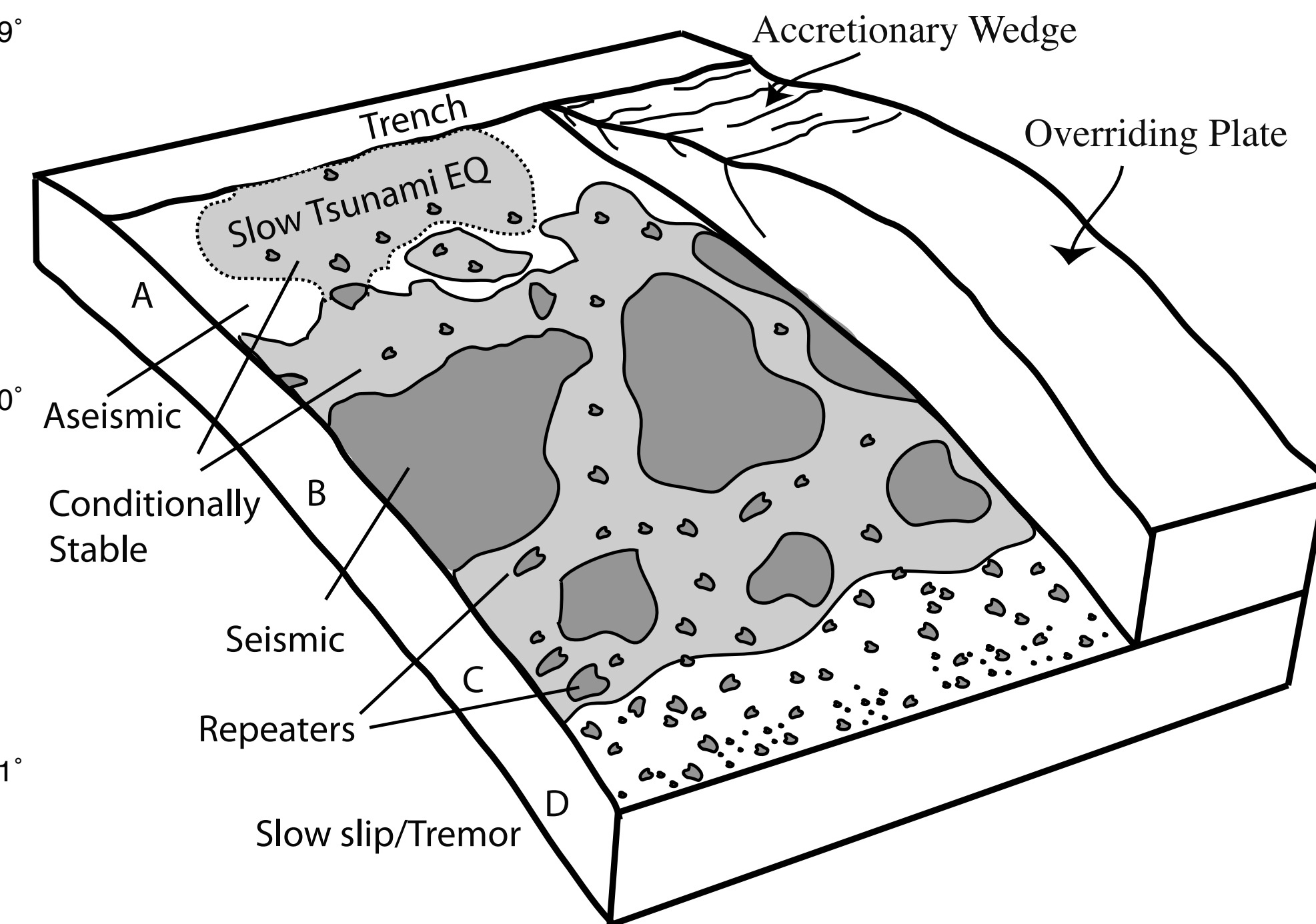
Satriano et al., AGU (2010)

Iquique (Chile) 2014 Mw 8.1

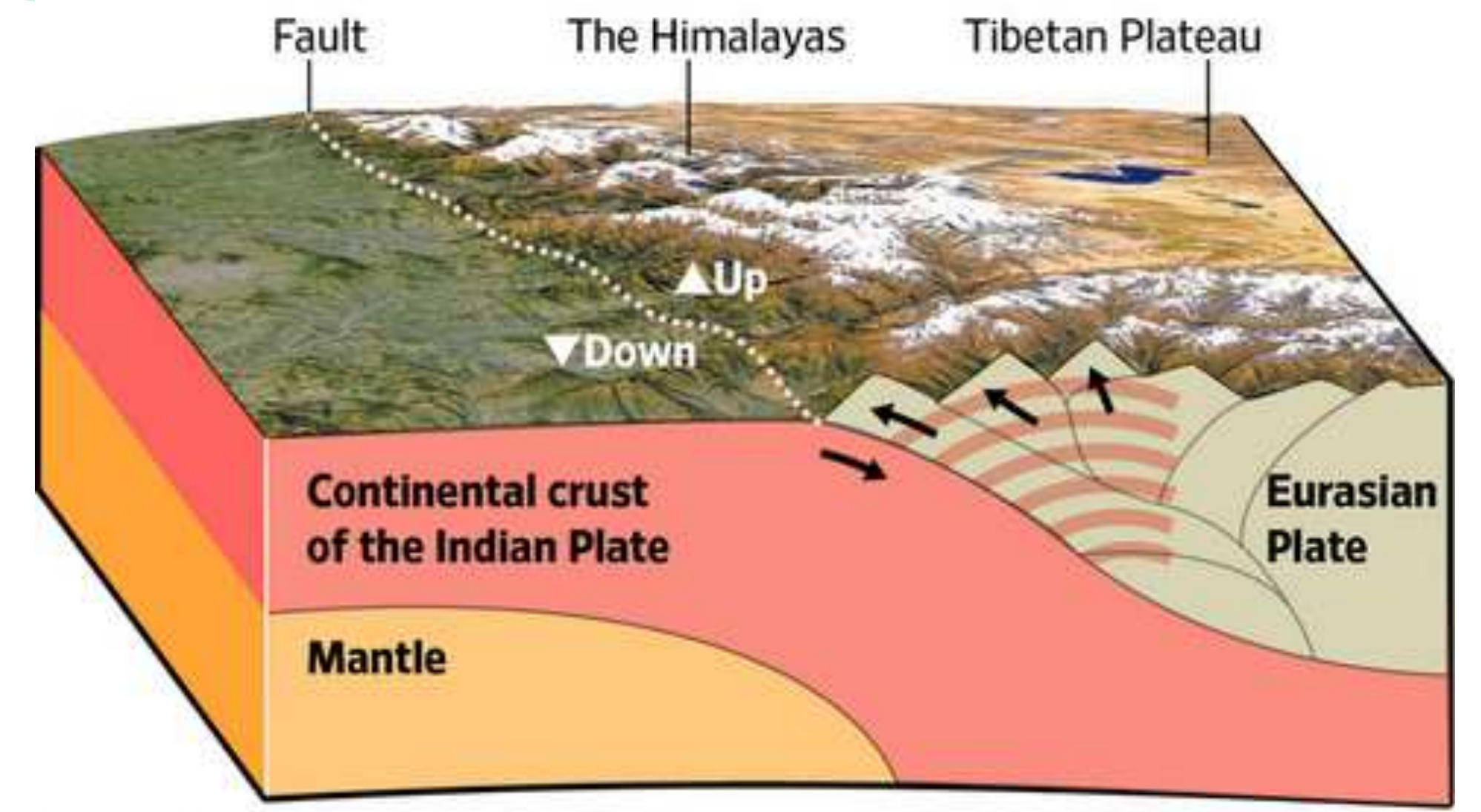
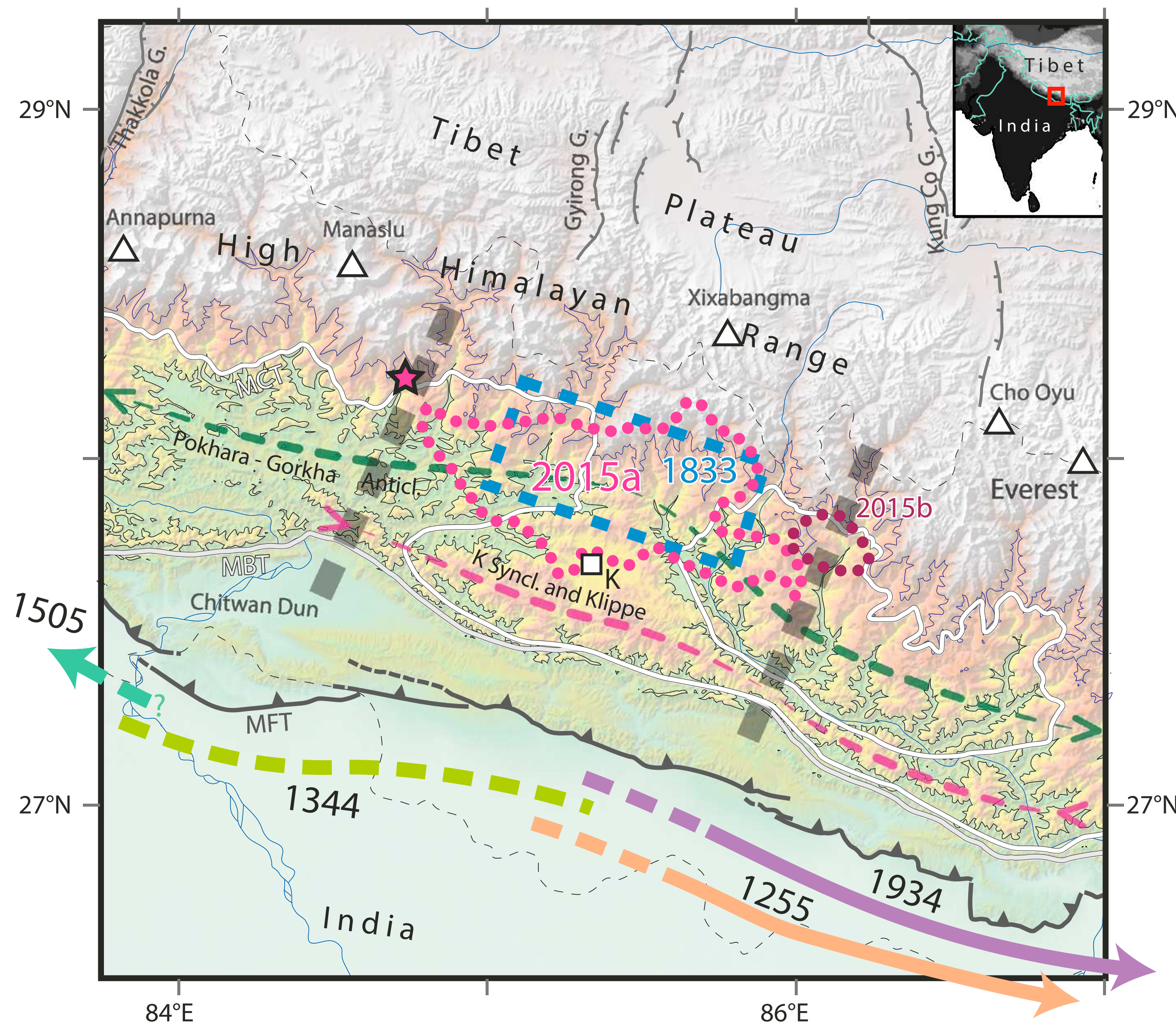


Satriano et al., AGU (2014)

Lay et al., JGR (2012)



The 2015 Mw 7.8 Gorkha earthquake

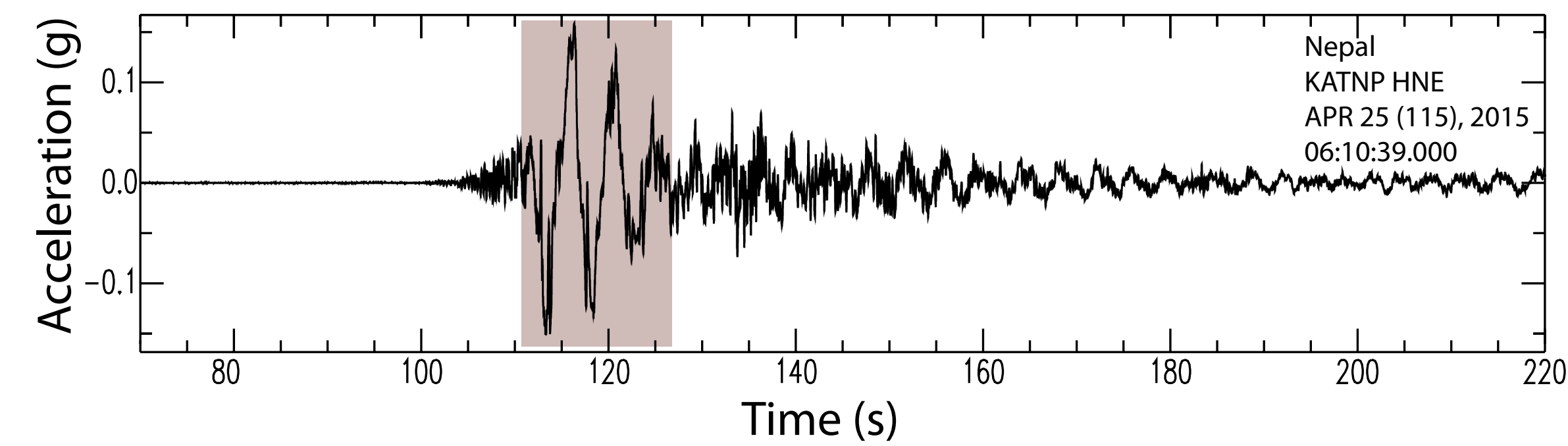


Death toll of ~9000 people :

- About 1700 in Kathmandu and its valley
- 600-1000 by landslides

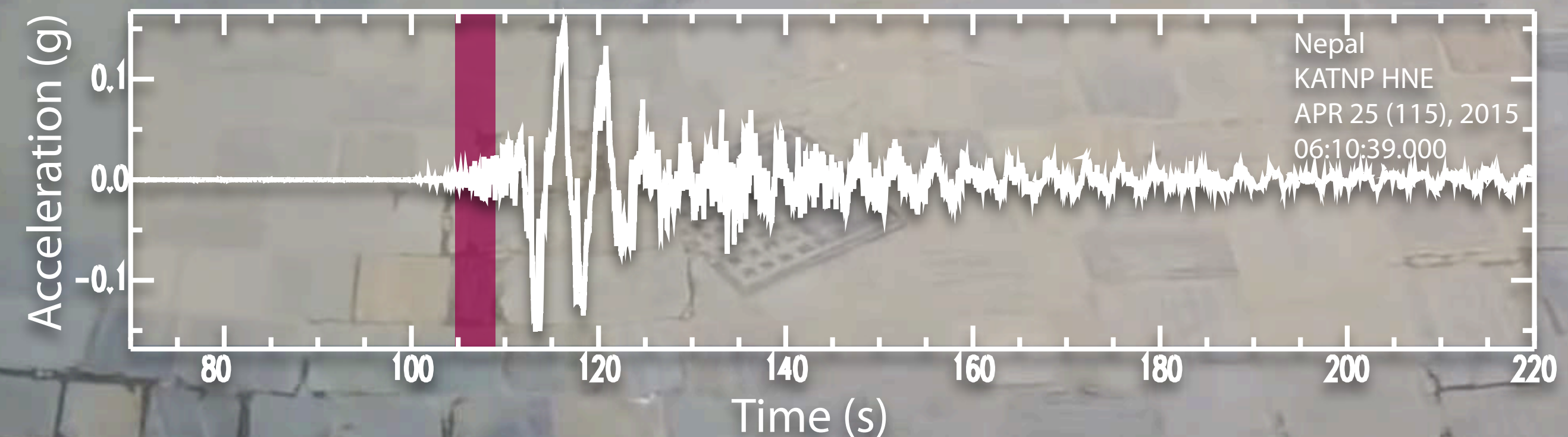
Maximum acceleration in Kathmandu (PGA) ~0.15 g:

- lower than expected
- in agreement with large damages but not full destruction



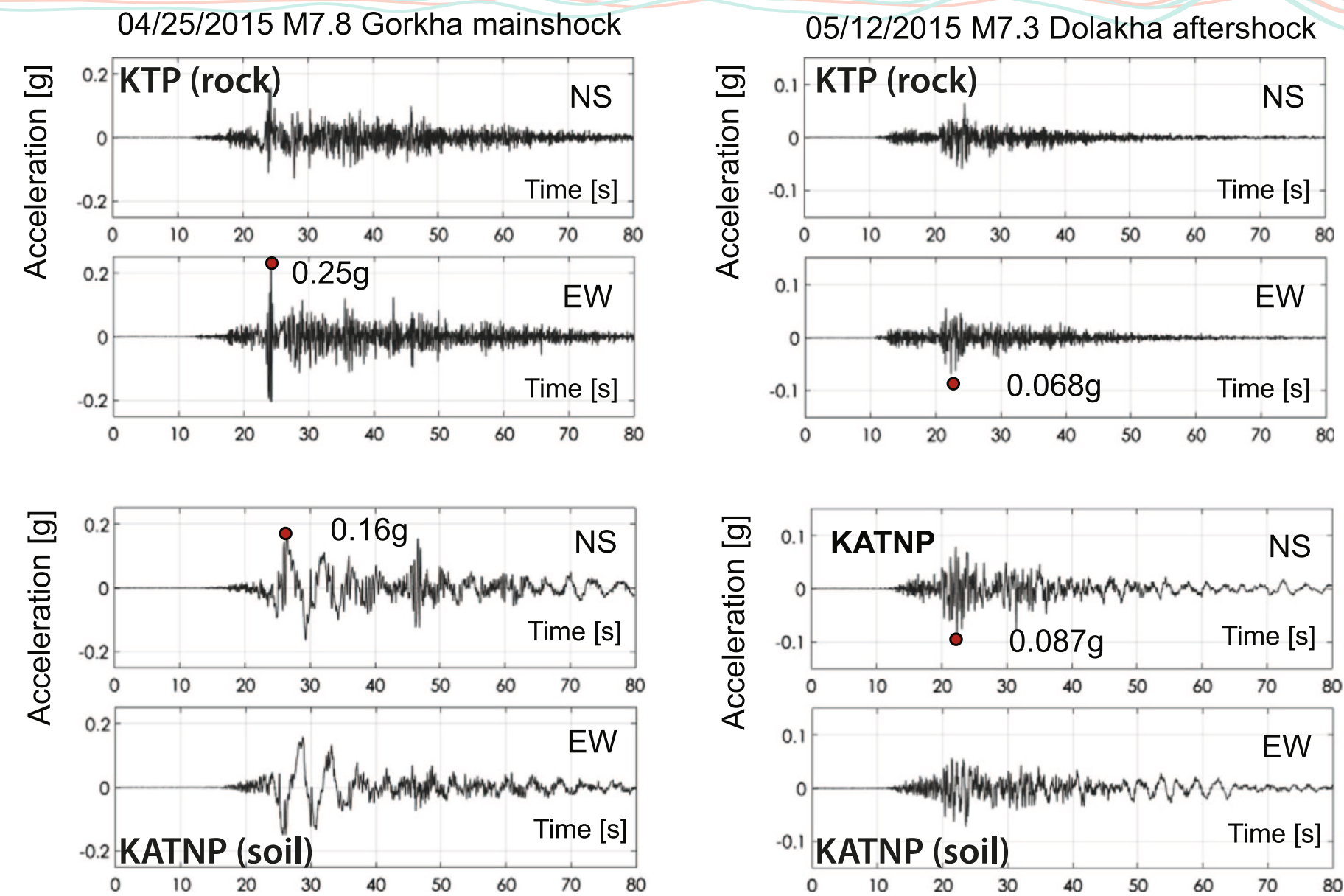
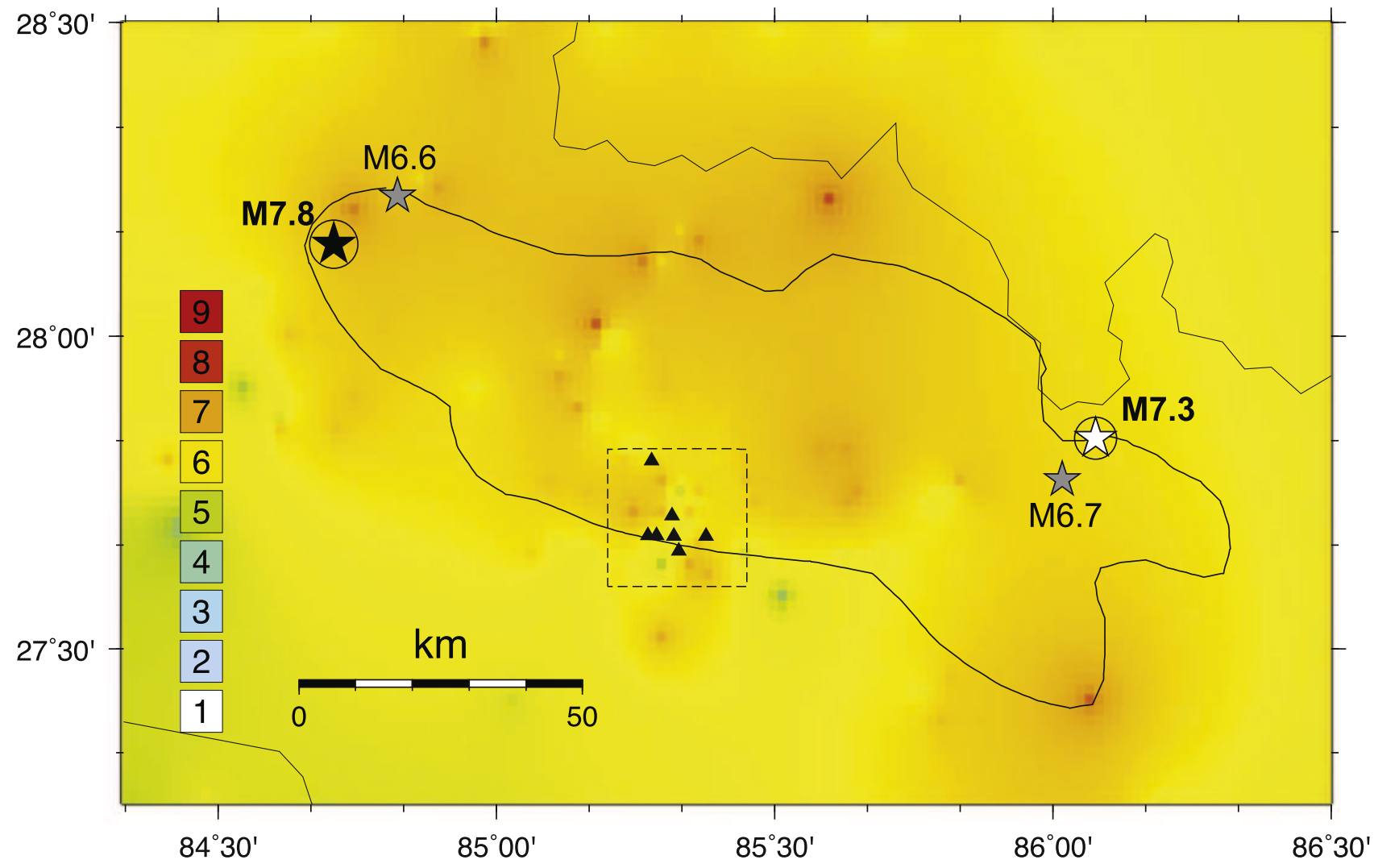
Grandin, Vallée, Satriano et al., GRL (2015)

04-25-2015 Sat 11:56:28



DABALI

Gorkha earthquake: intensity and ground acceleration above the rupture



Modified from
Rajaure et al., Tectonophysics (2016)

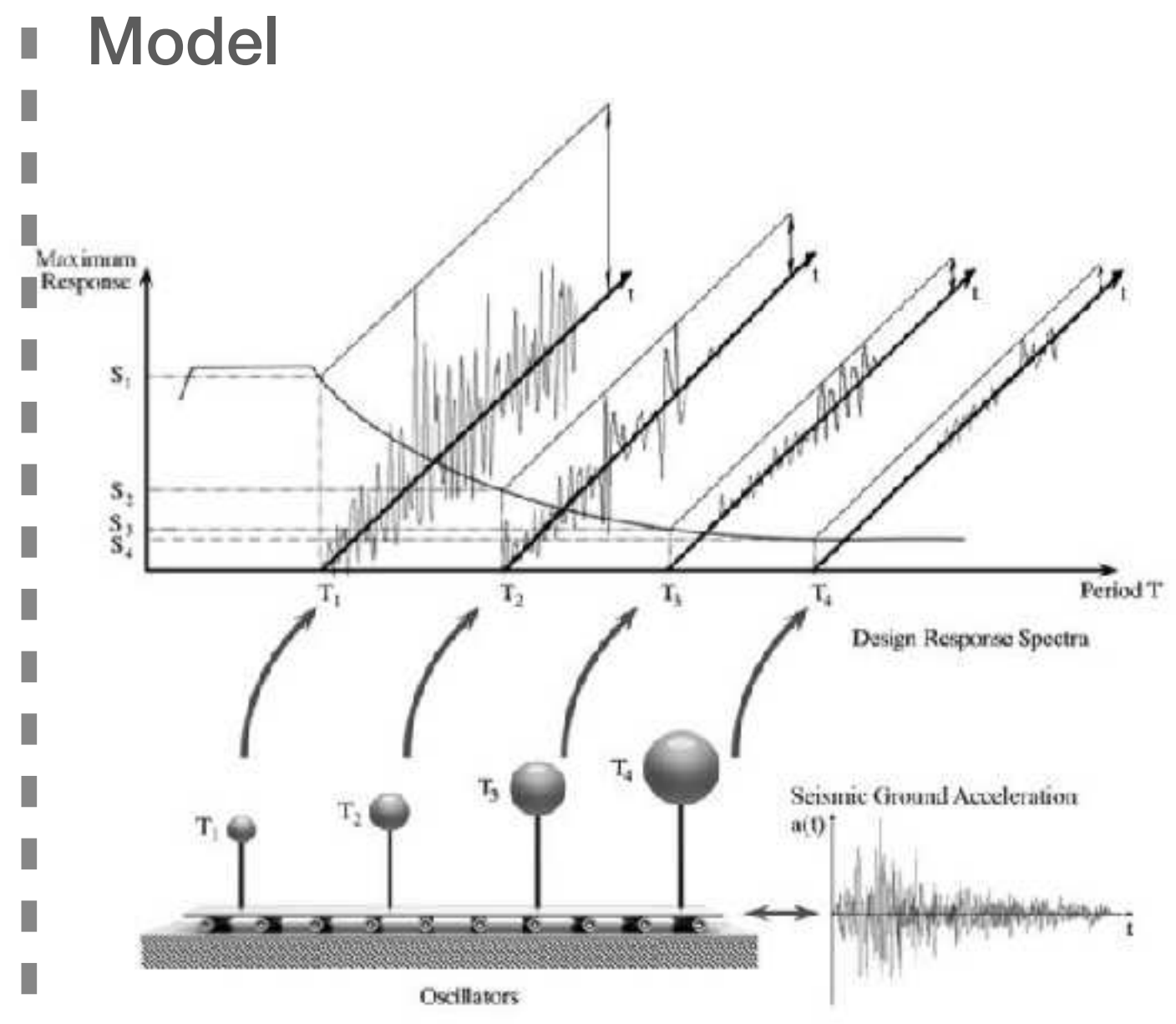
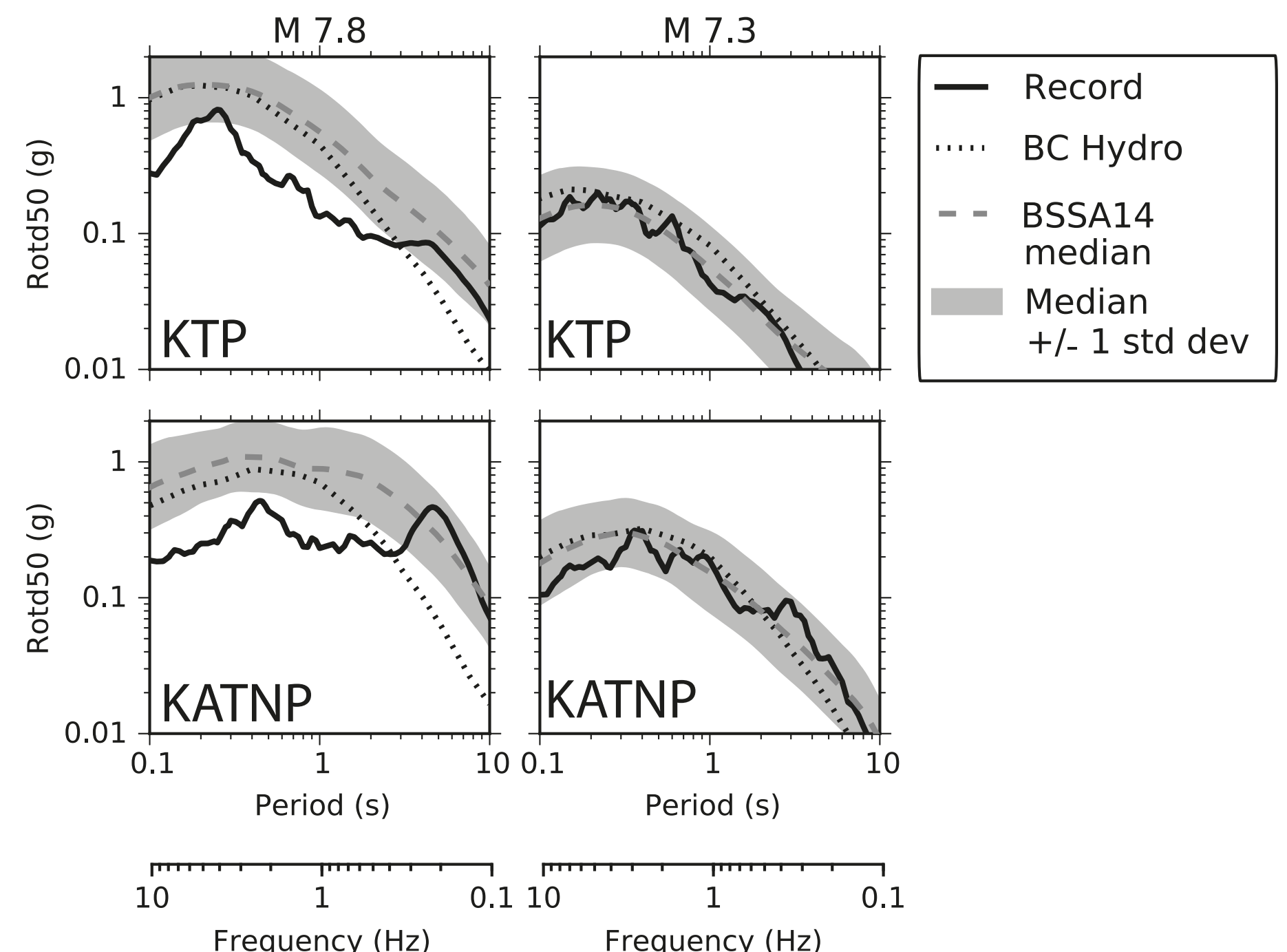
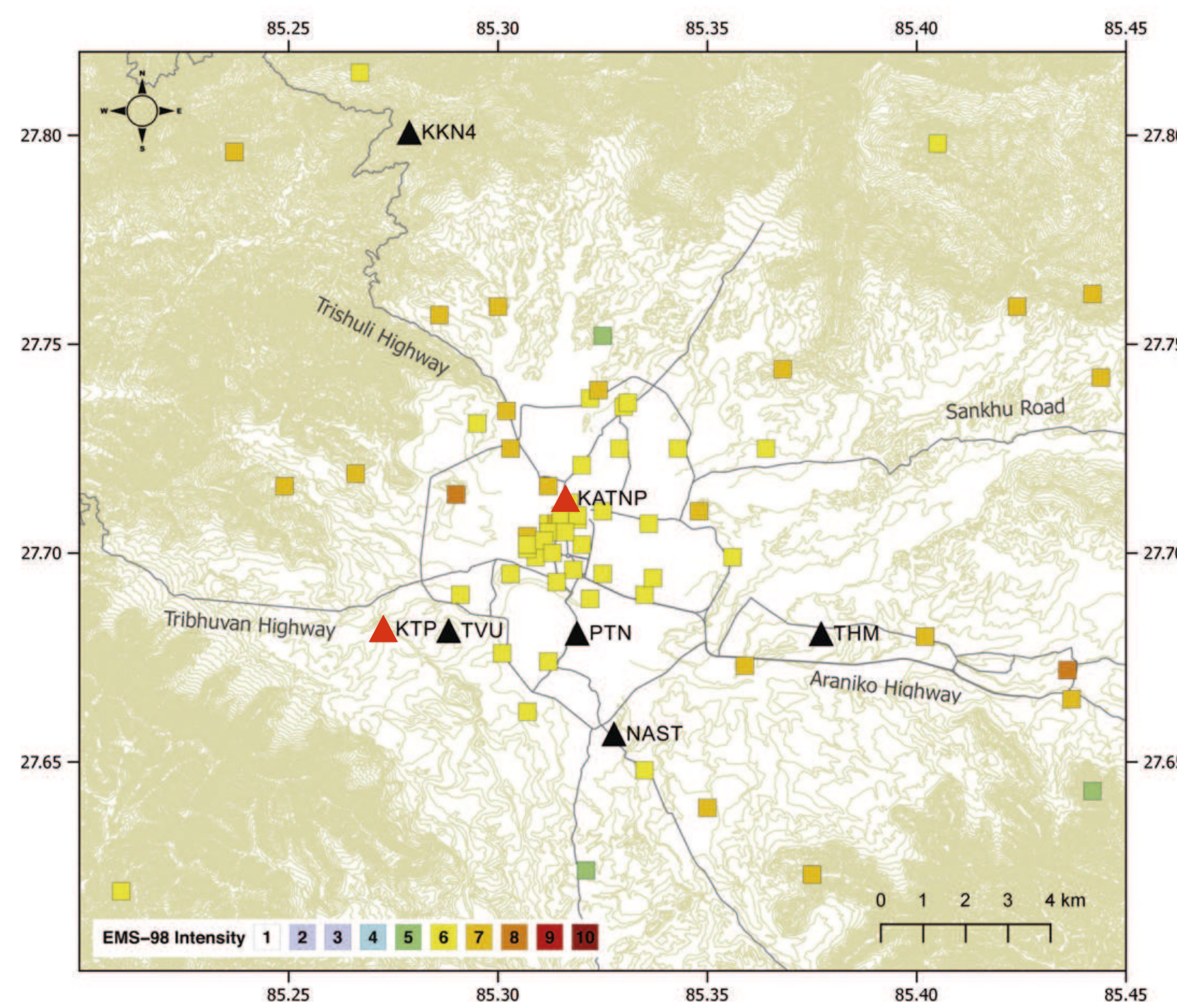
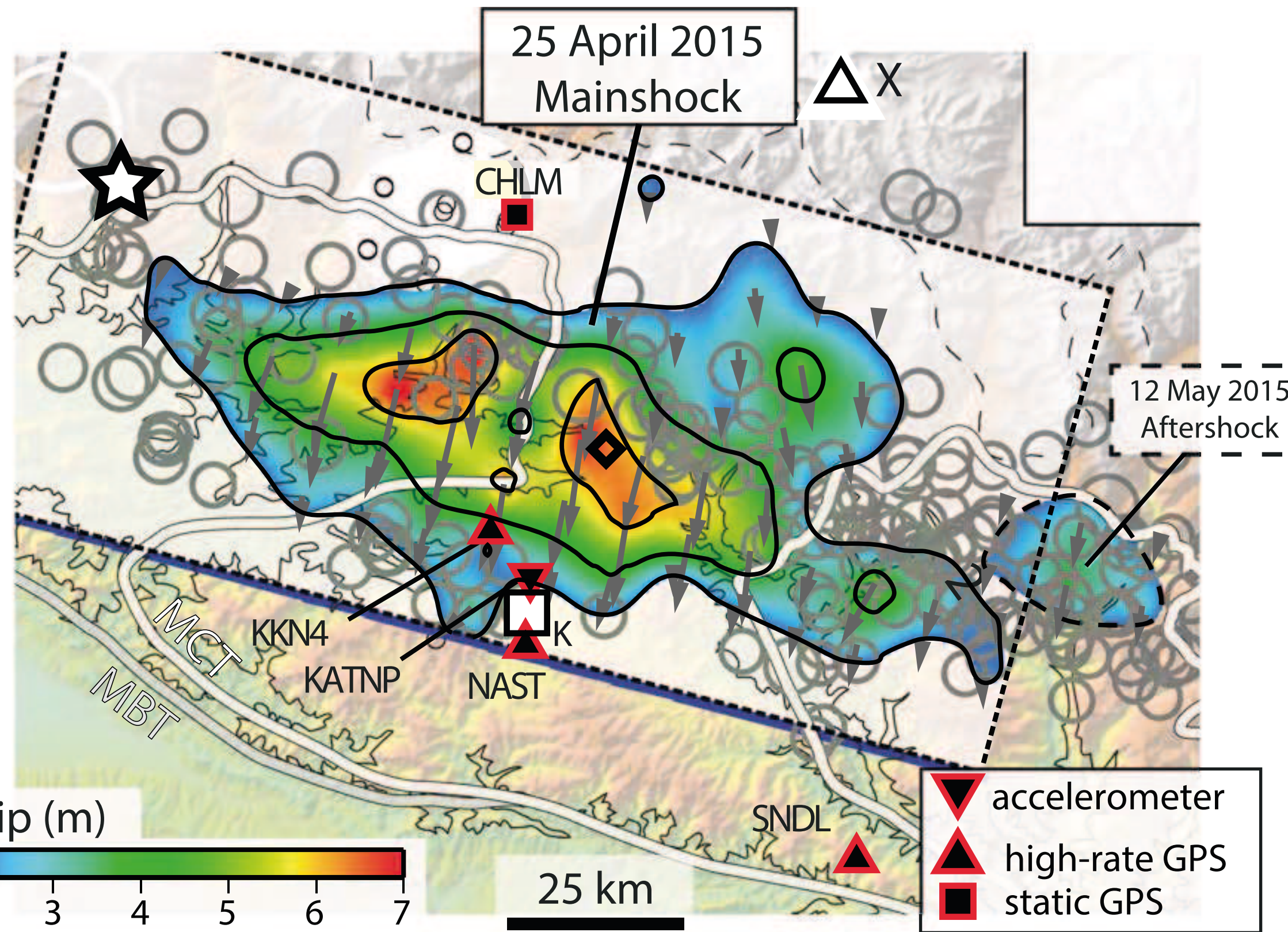
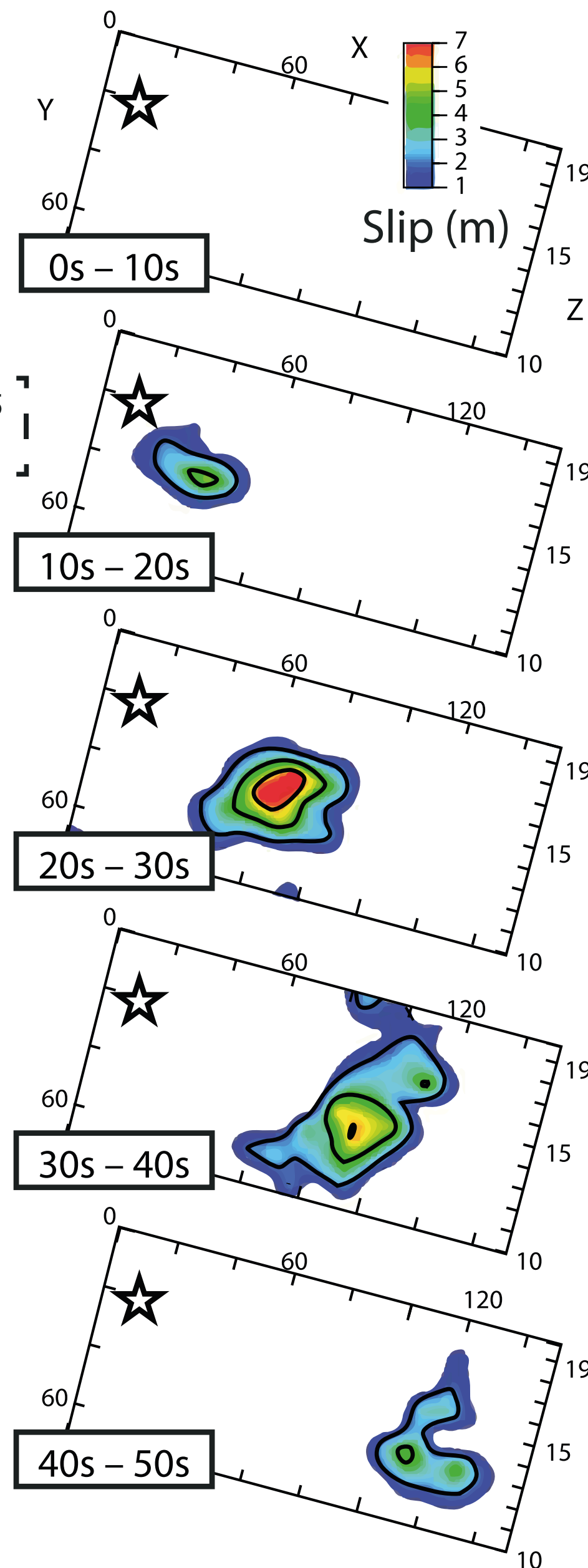
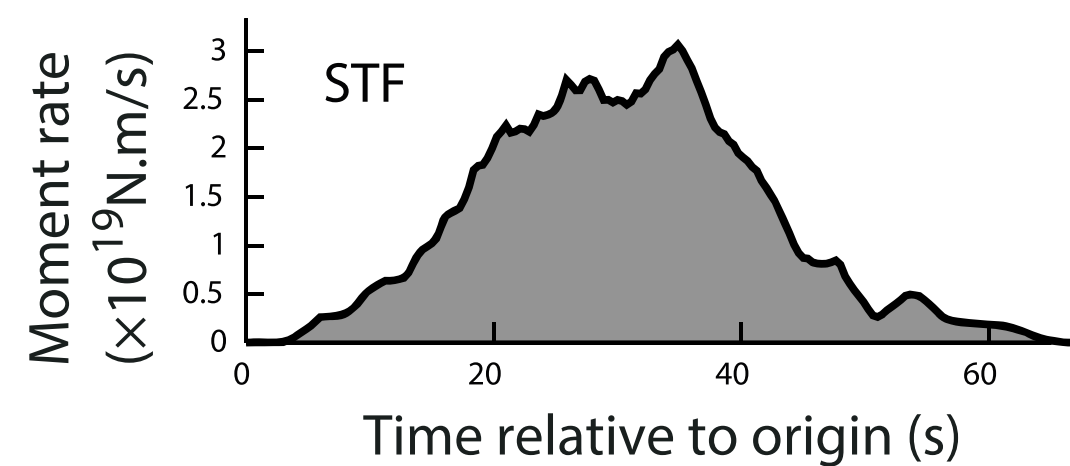


Figure 1. Concept of earthquake response spectrum.
e.g., Vrochidou et al. (2014)

Gorkha earthquake: kinematic slip model

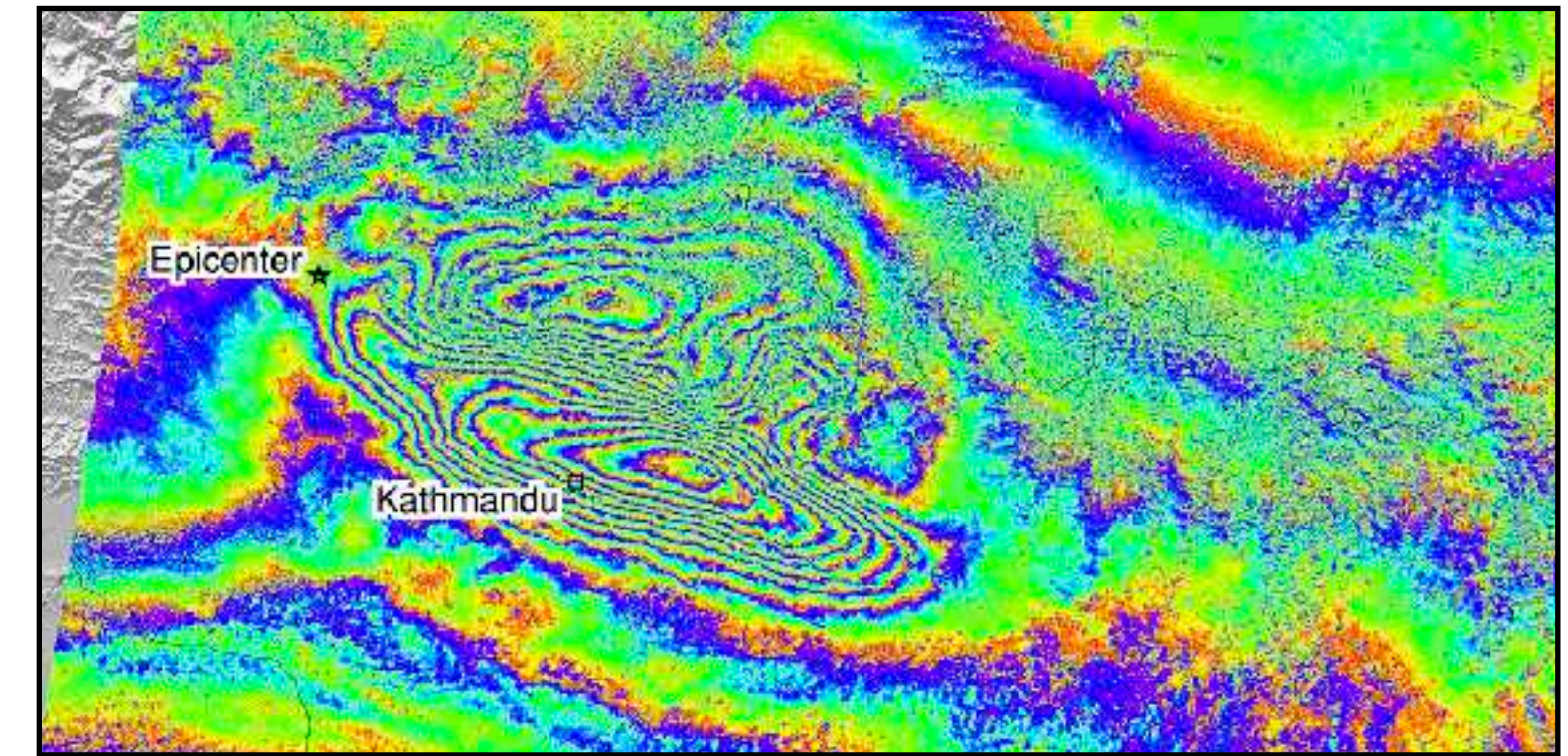


Grandin, Vallée, Satriano et al., GRL (2015)

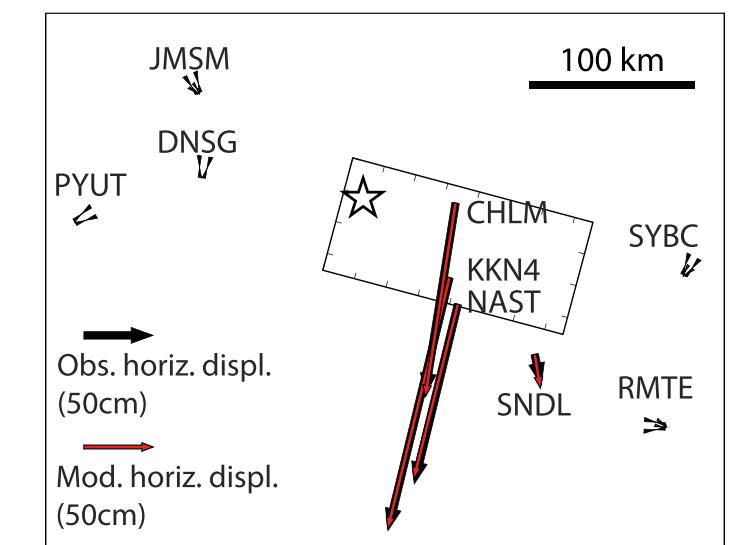


Data

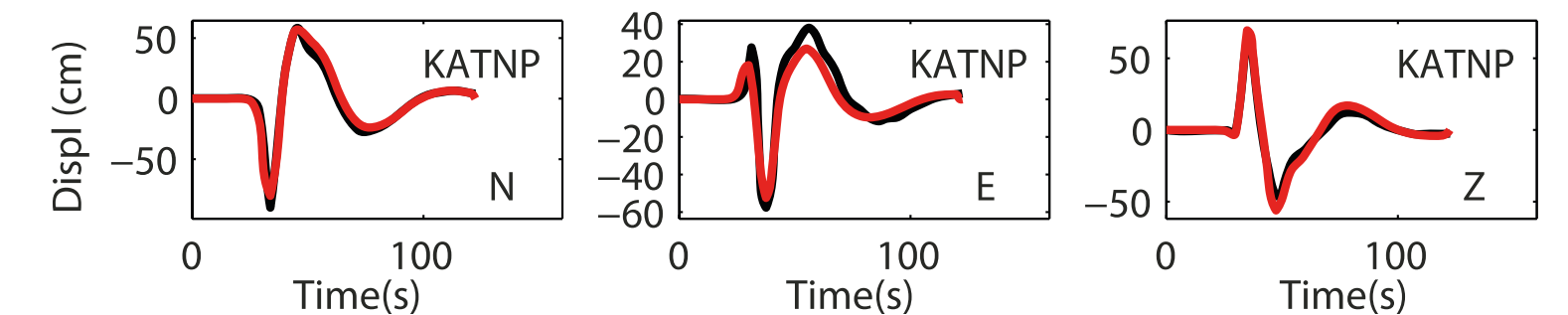
1. Interferometric synthetic aperture radar



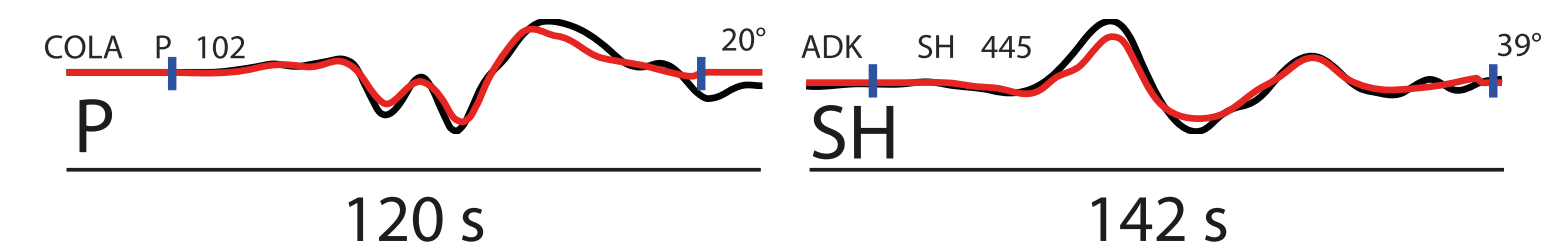
2. Static GPS displacement



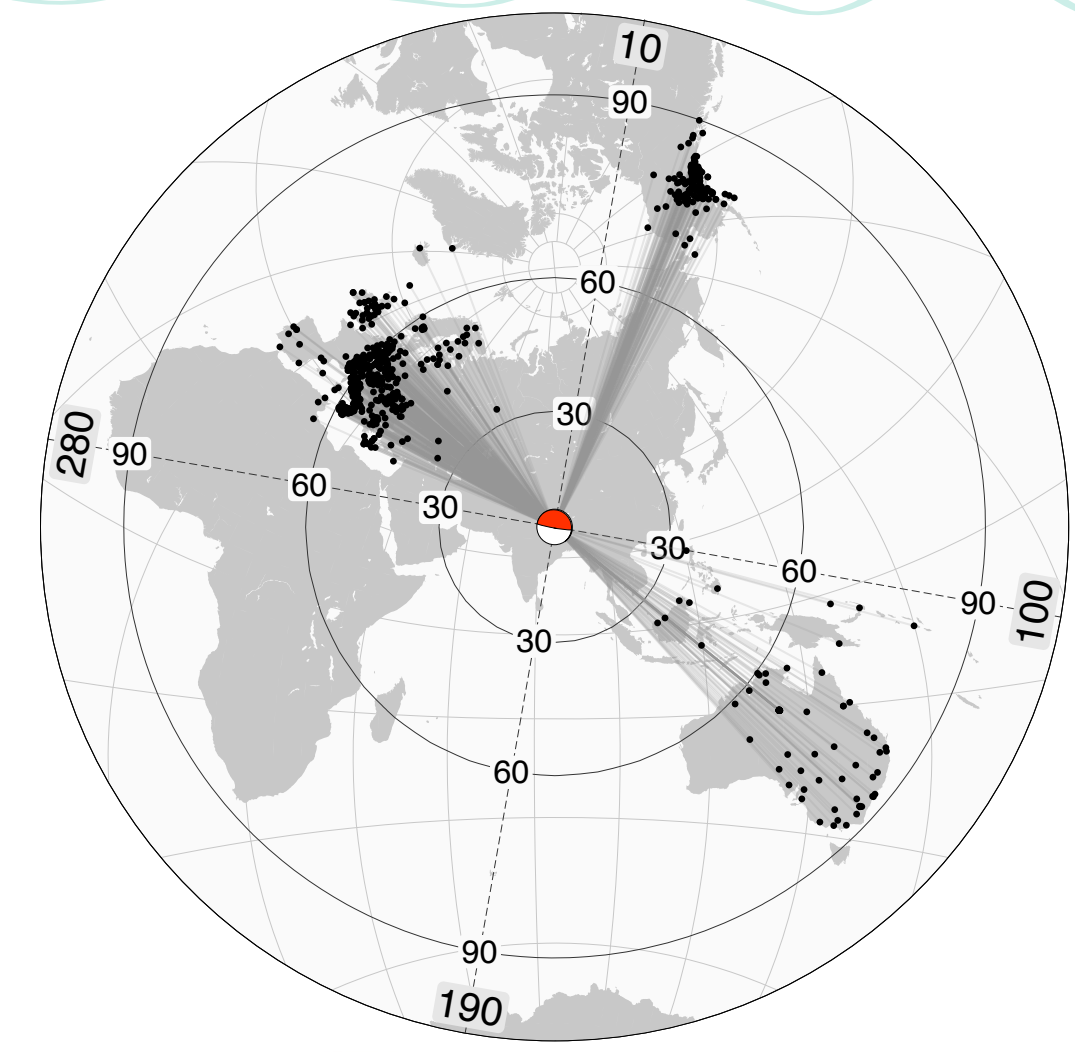
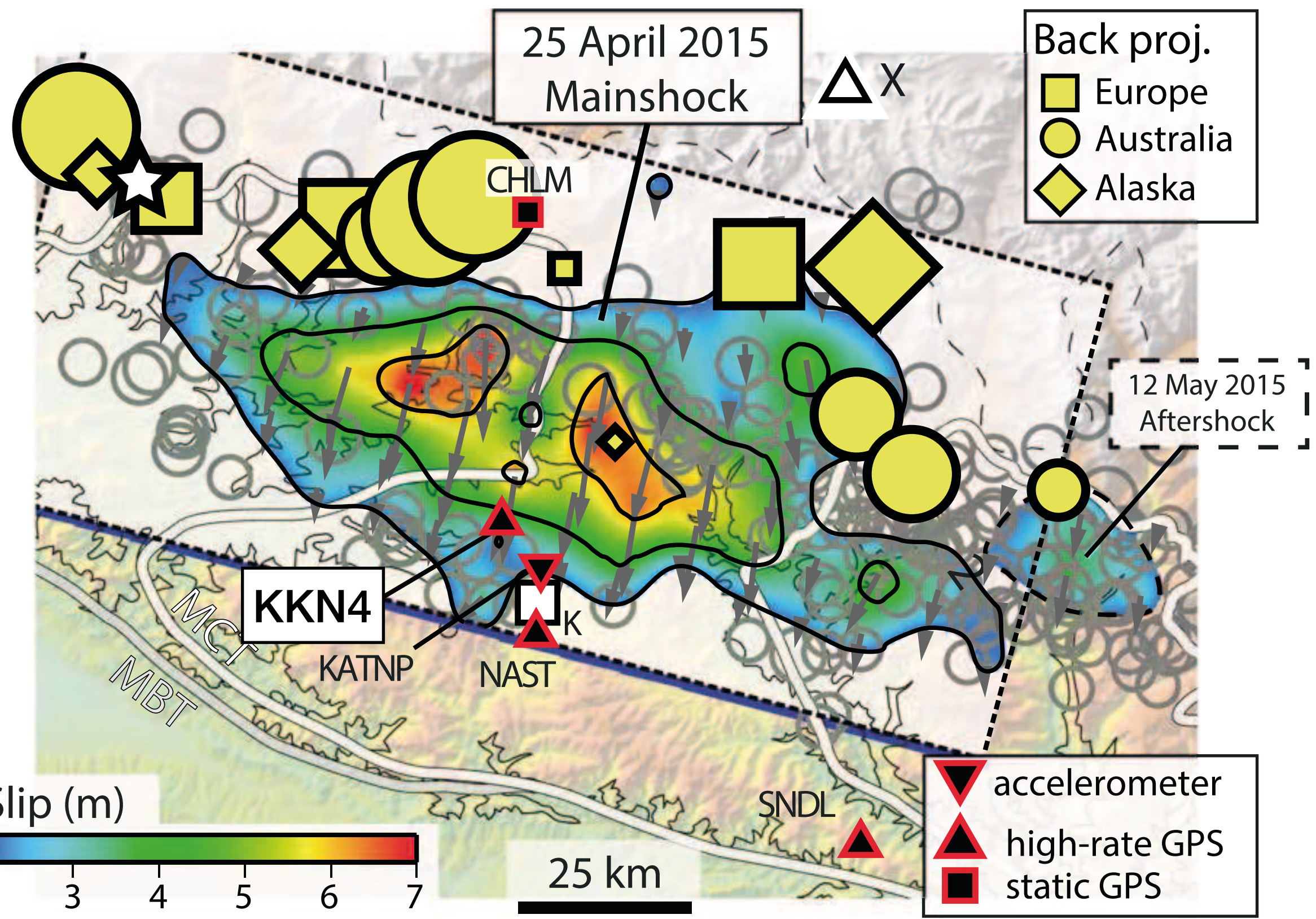
2. Local ground motion (up to 0.1 Hz)



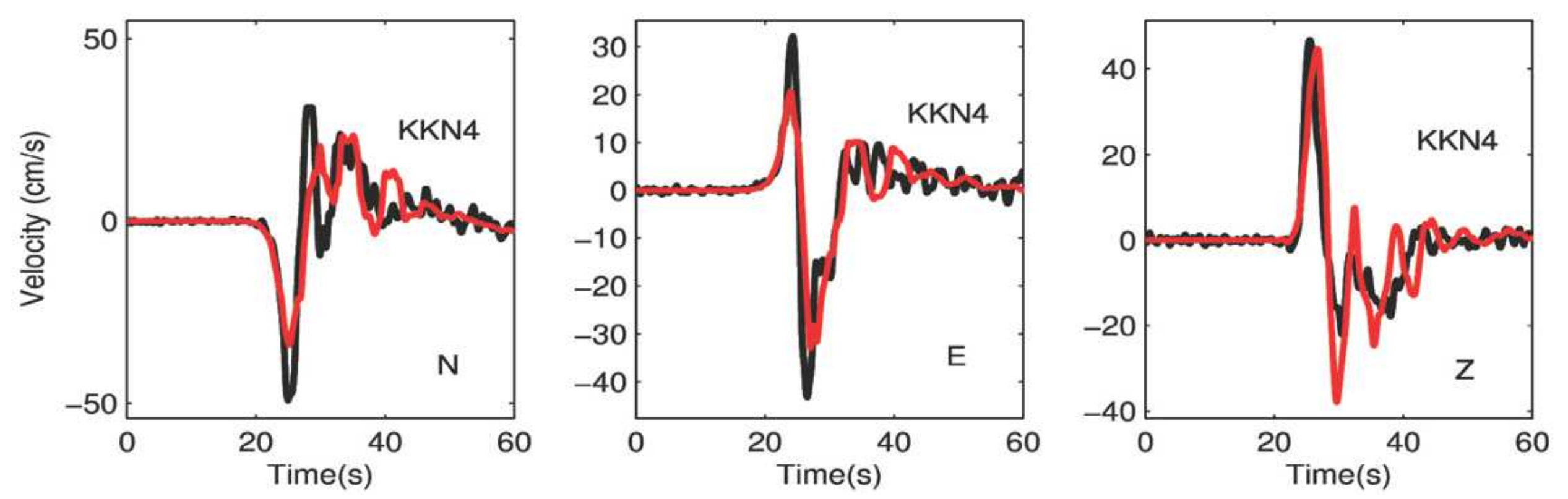
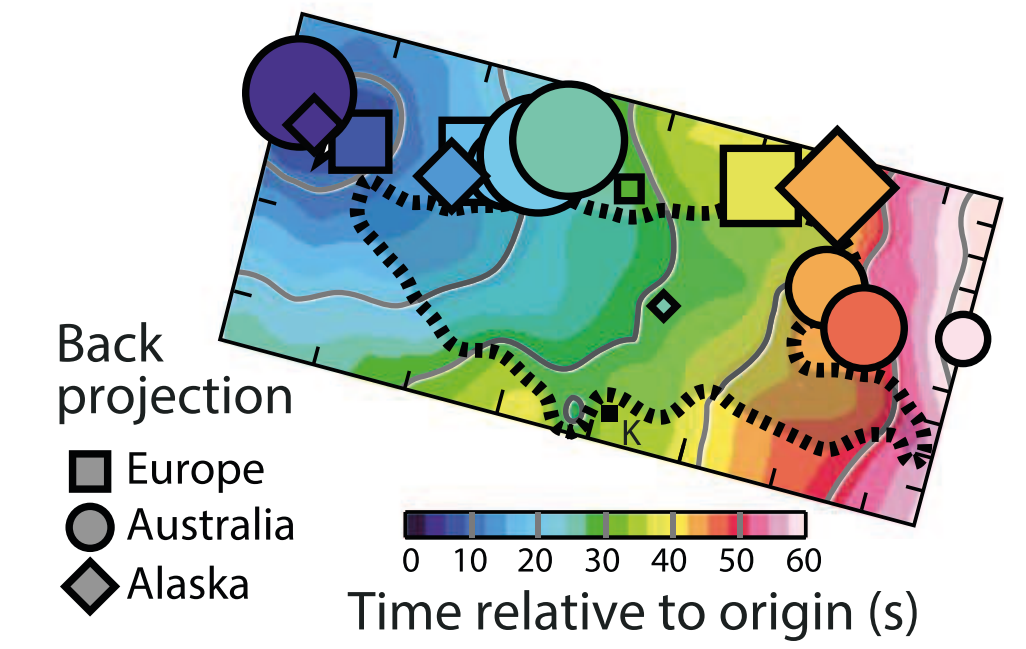
3. Teleseismic waveforms (up to 0.125 Hz)



LF earthquake in Kathmandu, HF at the base of the rupture

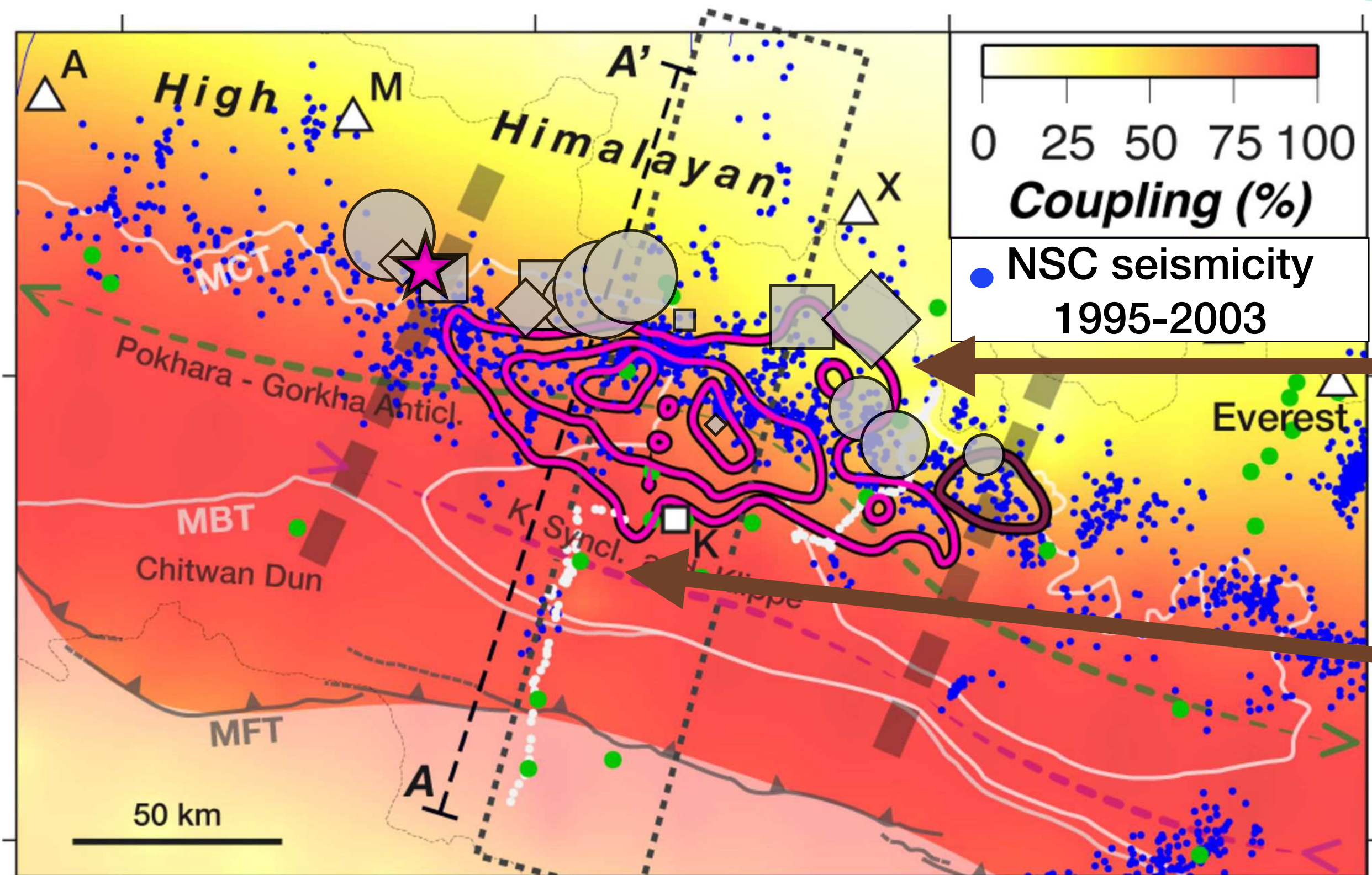


High-Frequency peaks (1-4 Hz) from back projection are at the down dip edge of the rupture



The kinematic model (inversion of data up to 0.1 Hz), well explains near-field data at frequencies up to 1 Hz

Rupture properties and fault zone heterogeneity



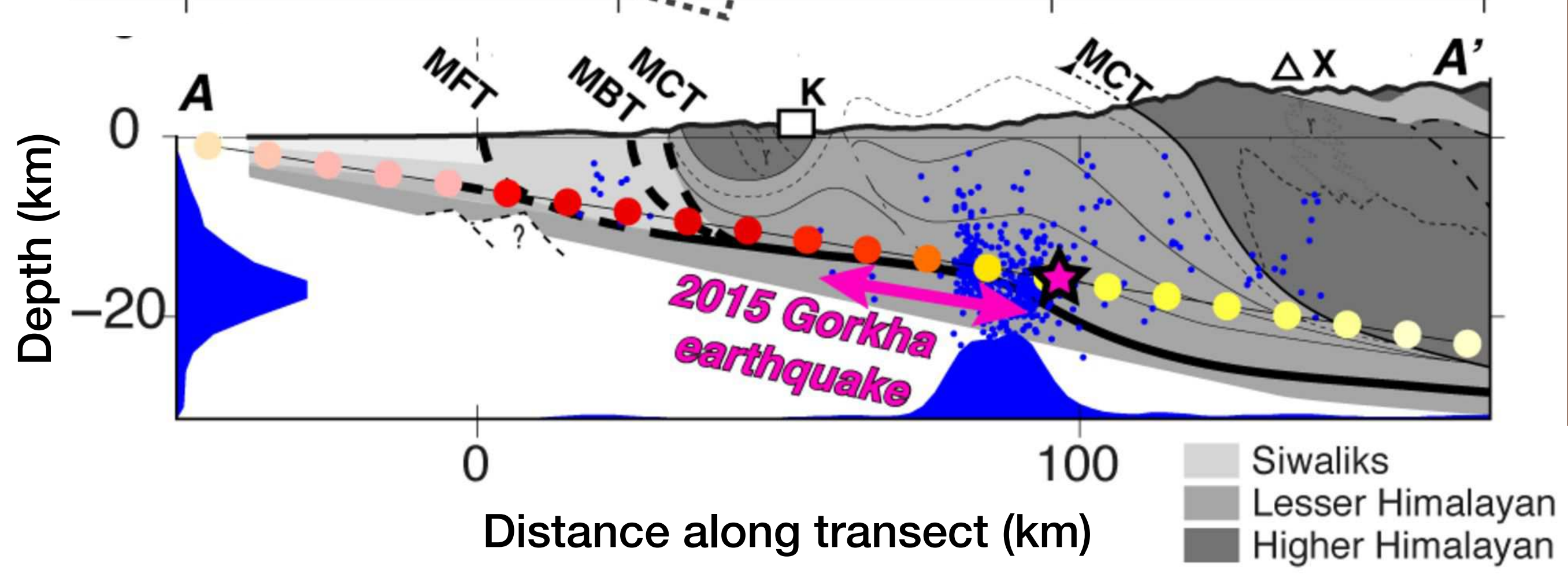
Rupture abruptly stops down-dip, with high-frequency radiation, where the plate interface strongly decouples.

Almost all the seismicity before 2015 Gorkha earthquake is in this zone: stress concentration?

Rupture gradually slows down up-dip, with no high-frequency radiation.

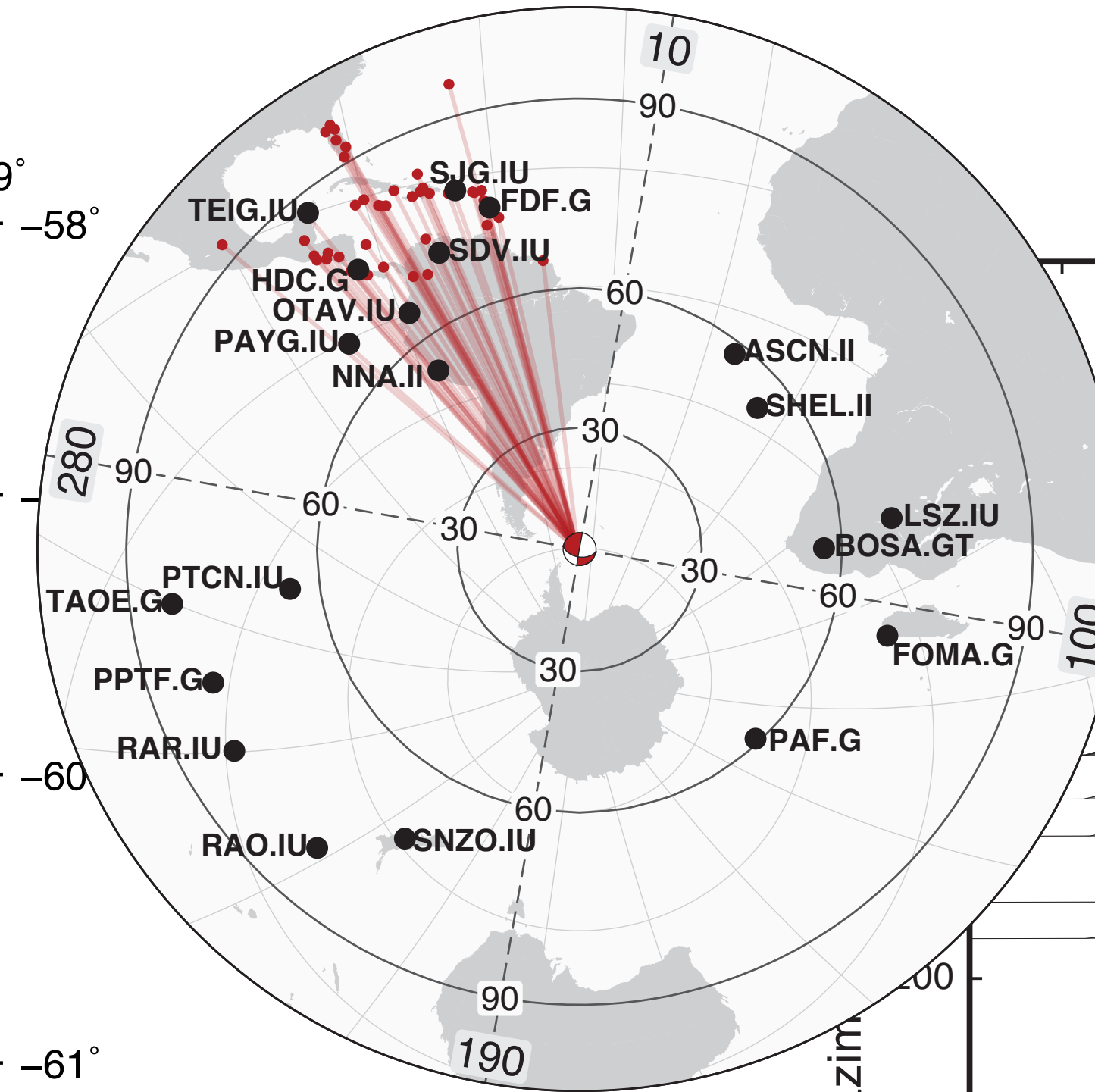
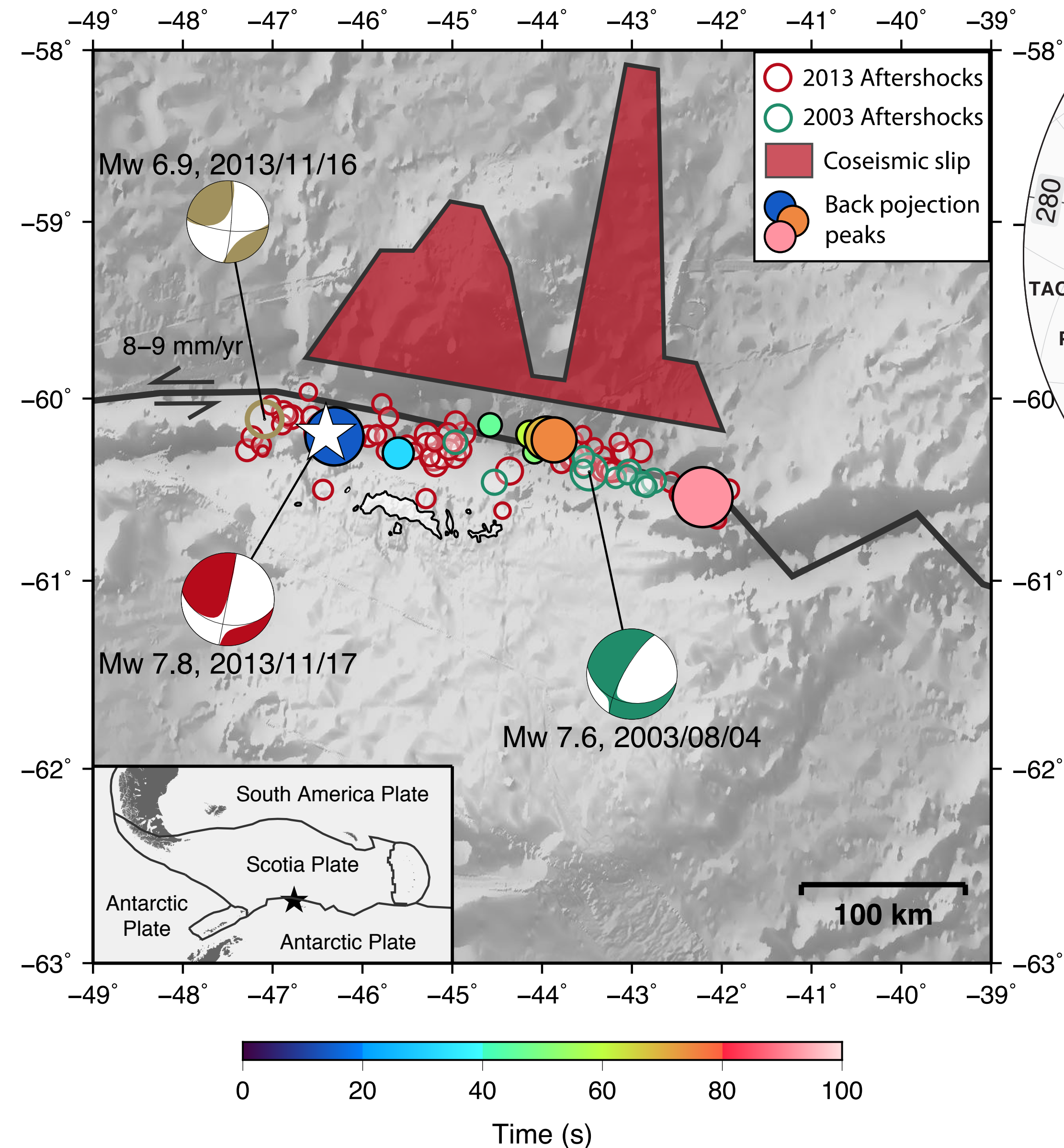
No seismicity in this zone before 2015: fully locked and/or strengthening friction?

Larger ruptures (M 8+) could reach the surface but only with low-frequency radiation?

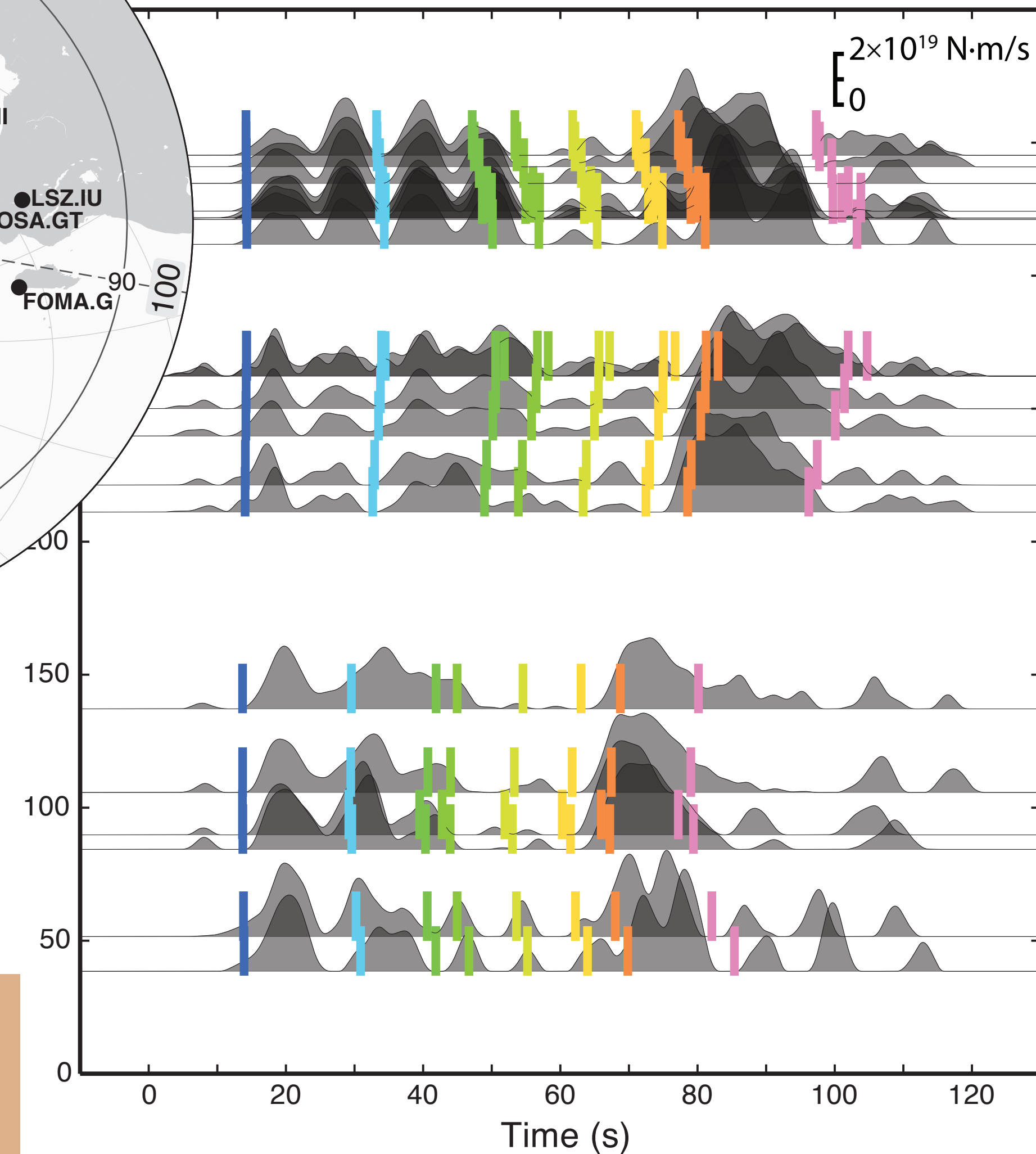


The 2013 Mw 7.8 Scotia Sea earthquake: fault segmentation and HF

Coseismic slip and back projection



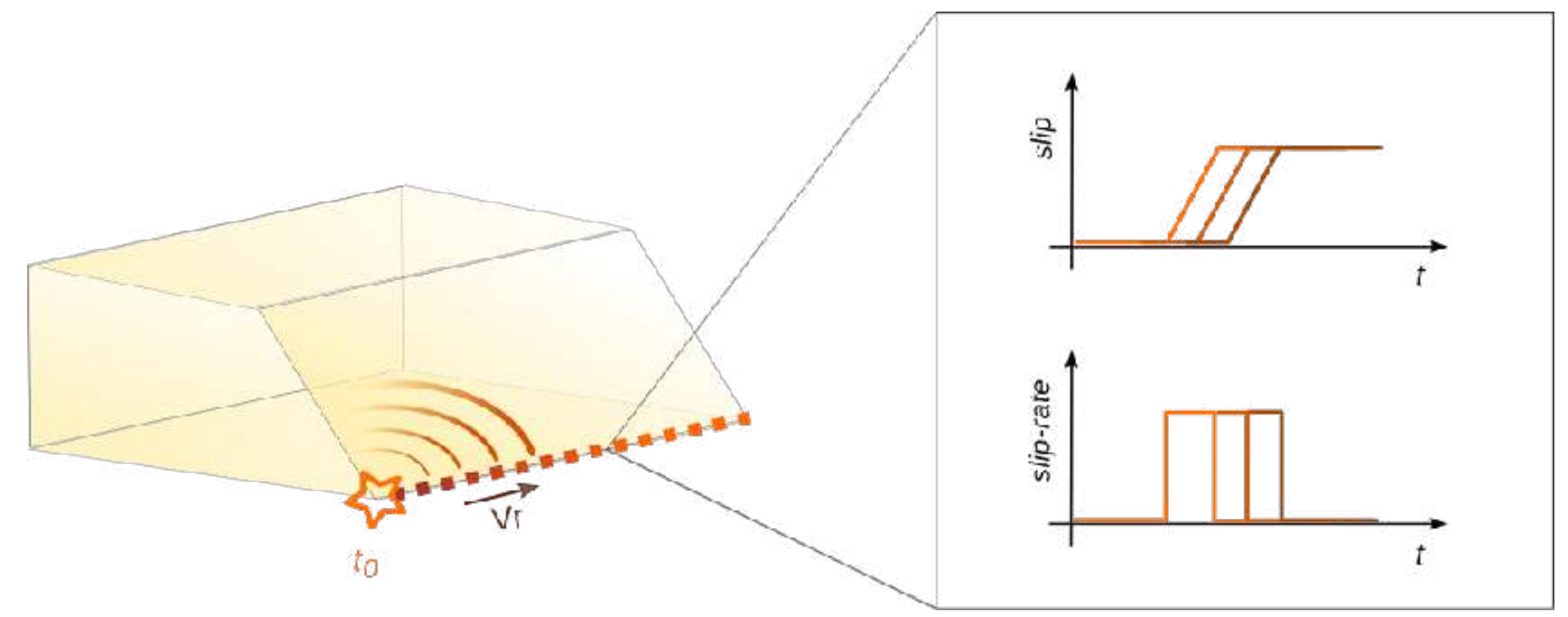
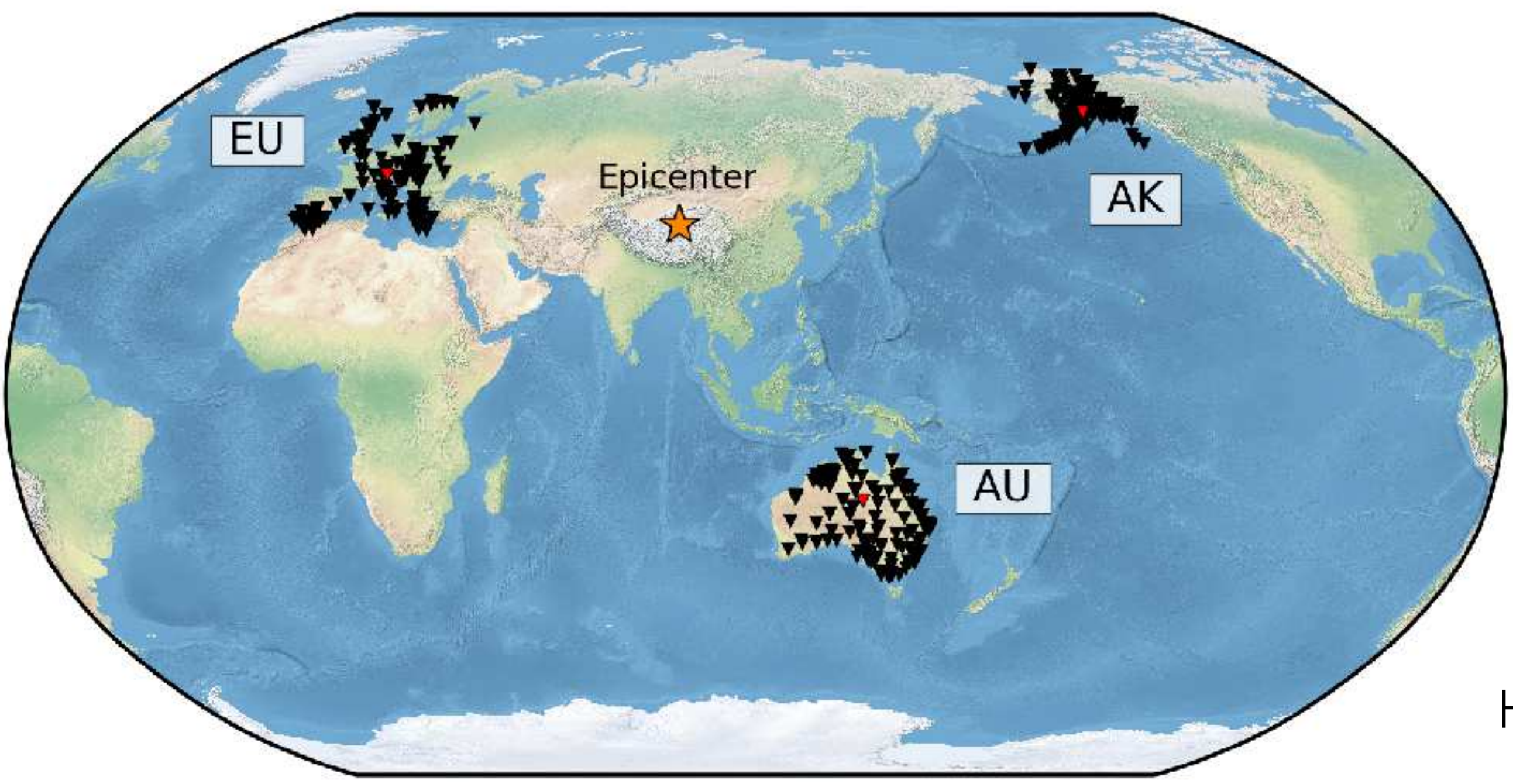
Apparent source time functions + BP peak times (0.5 - 1.0 Hz)



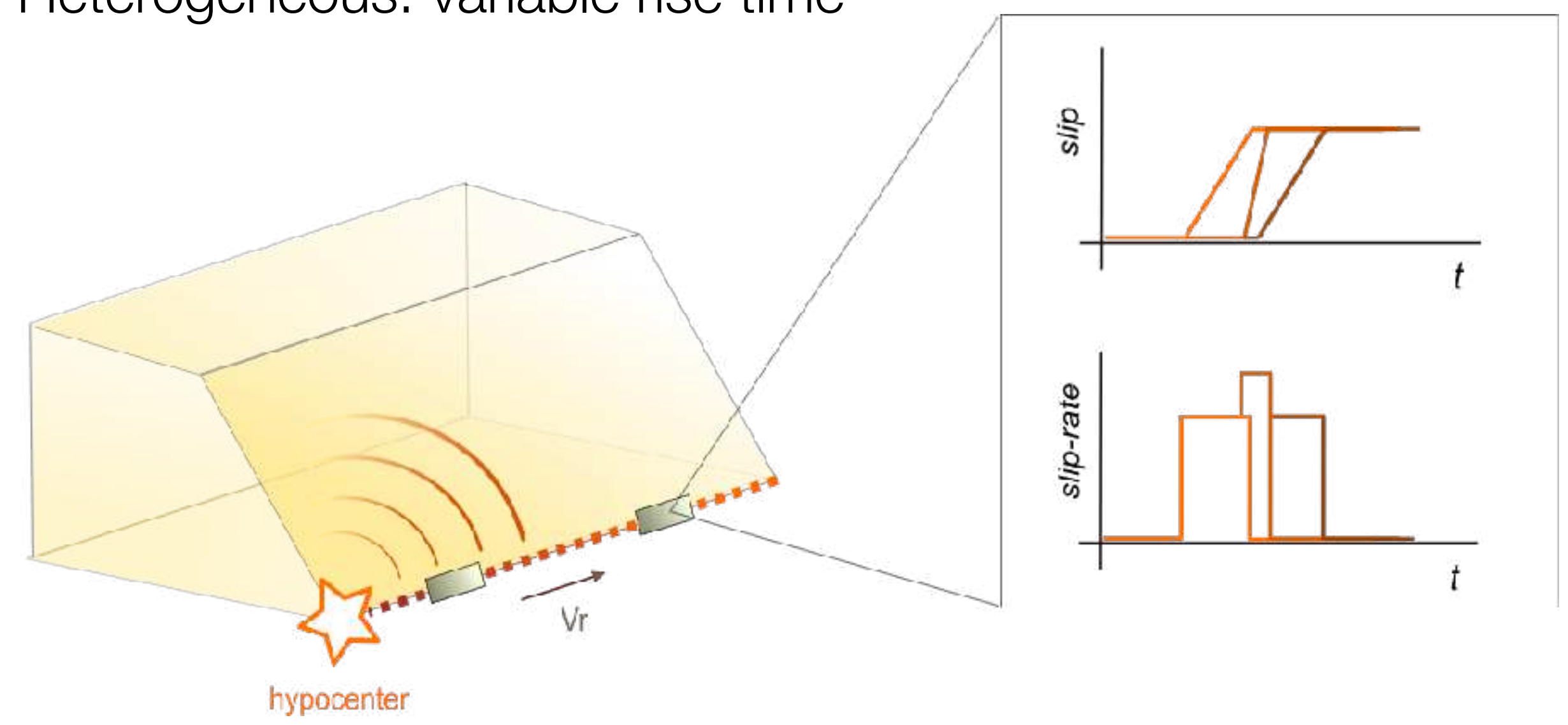
High-frequency radiation associated with slip acceleration

Fault heterogeneity and high-frequency radiation: synthetic tests

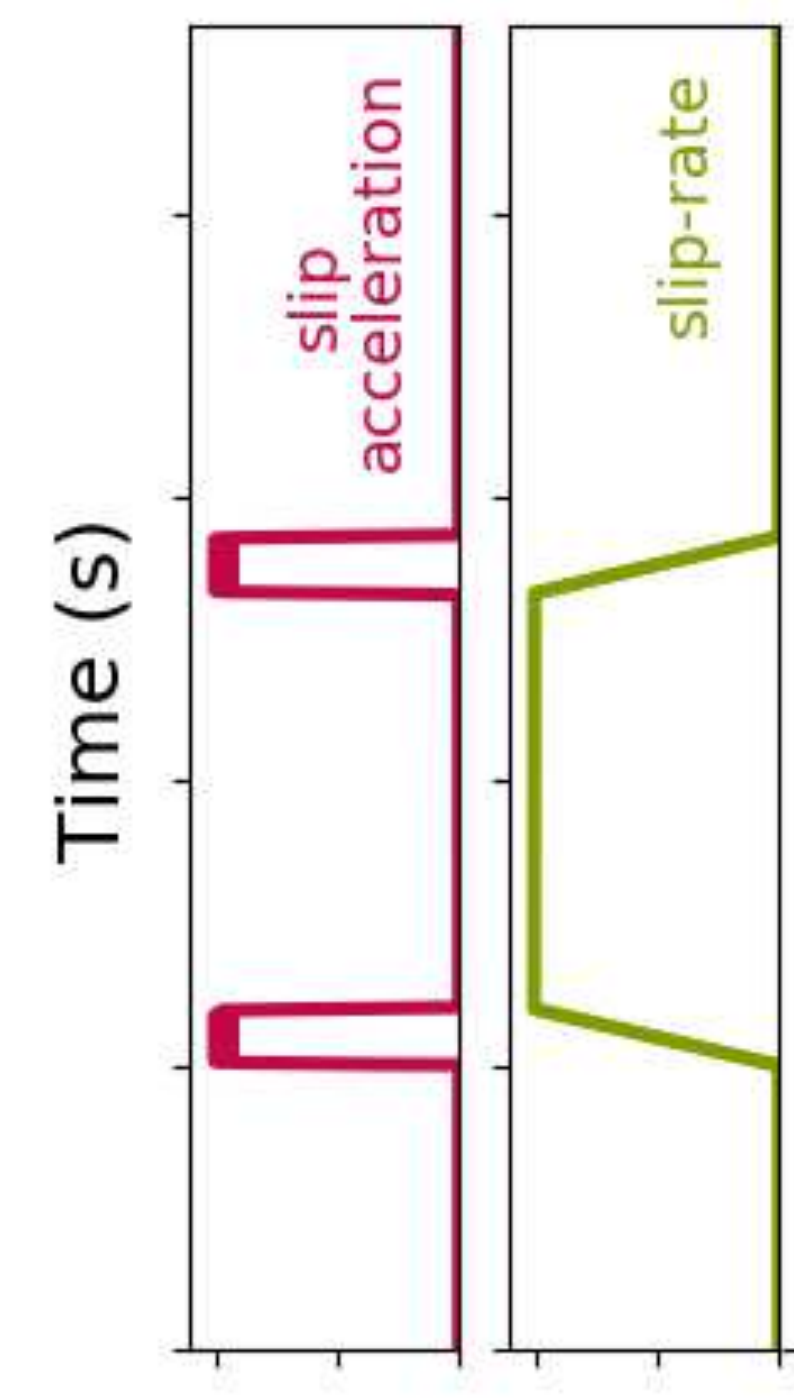
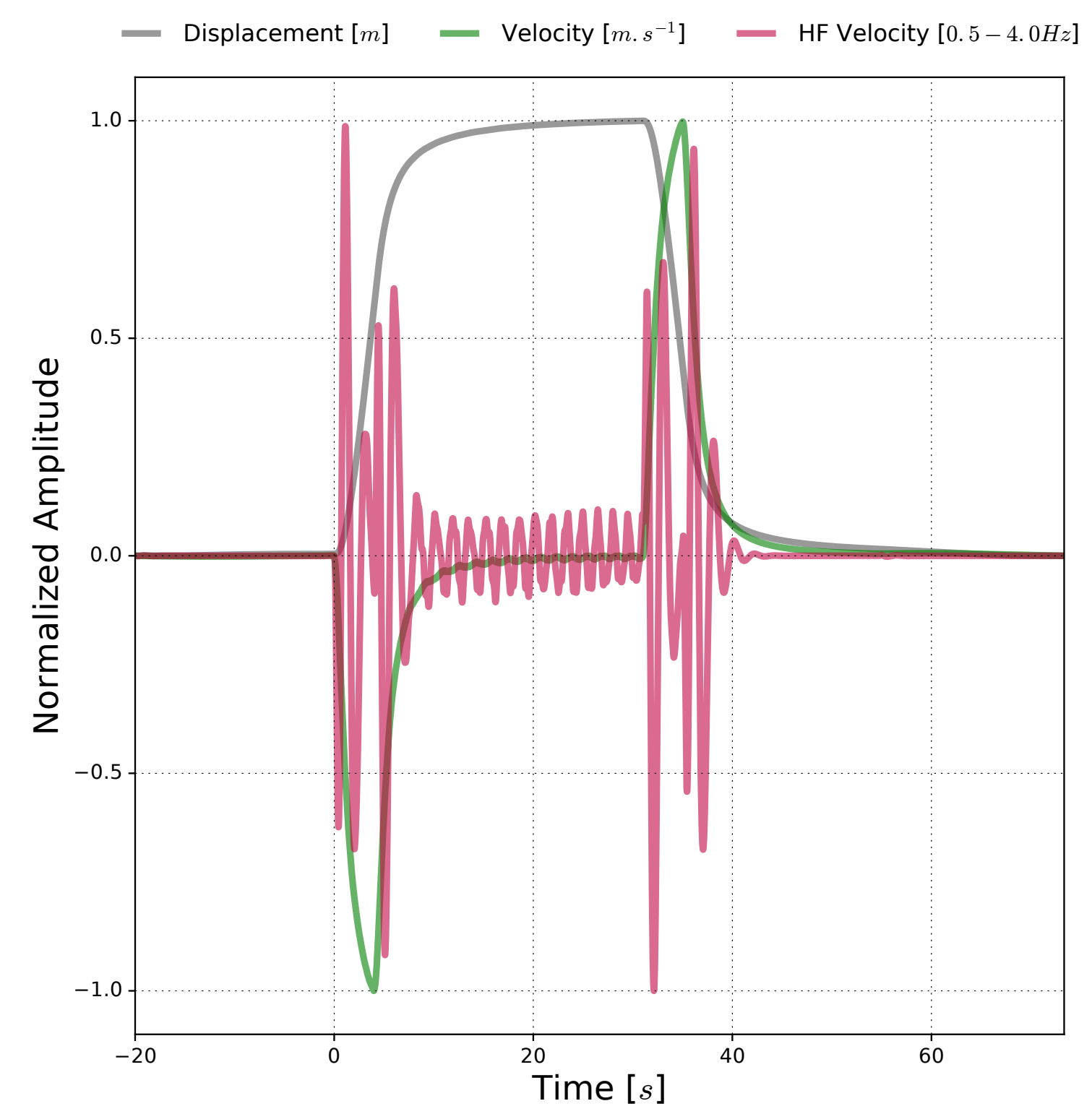
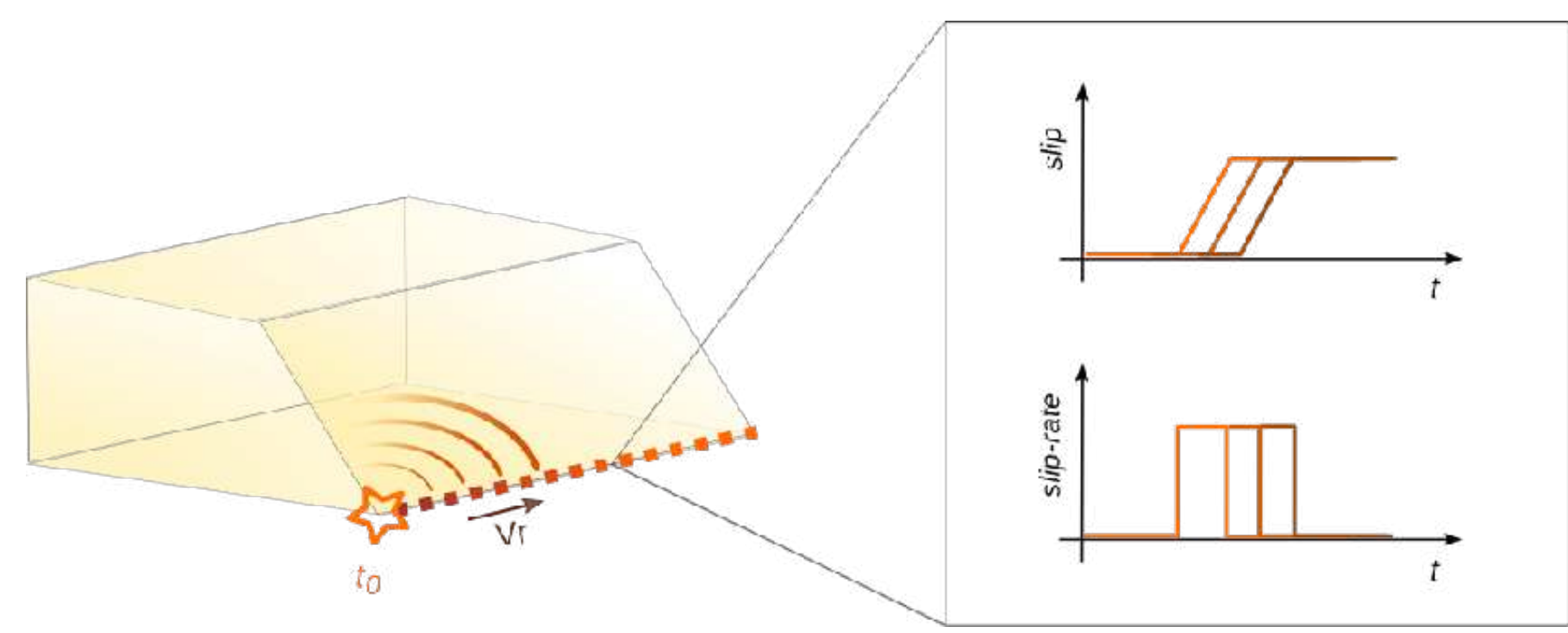
Homogenous: "Haskell" fault



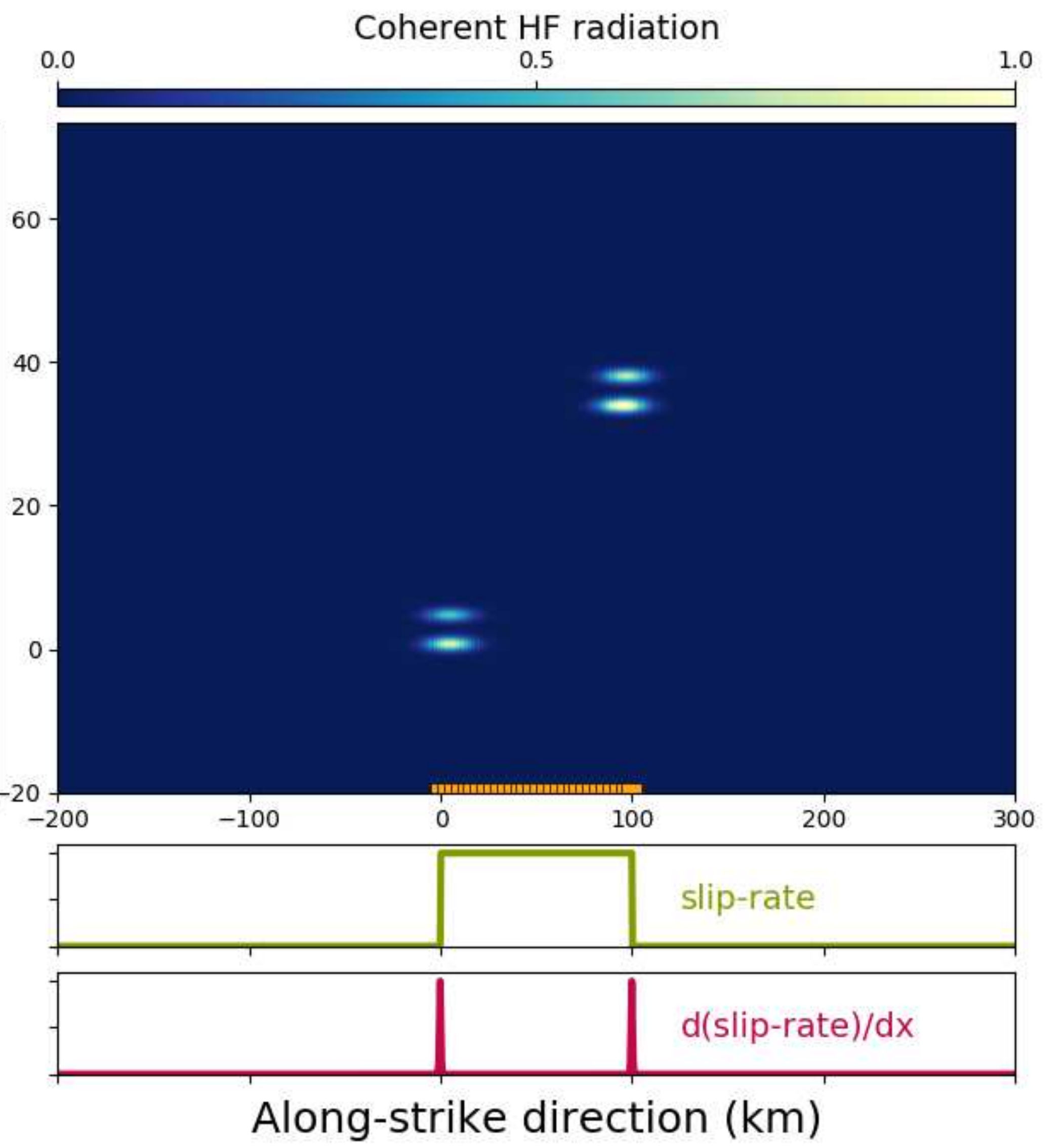
Heterogeneous: variable rise time



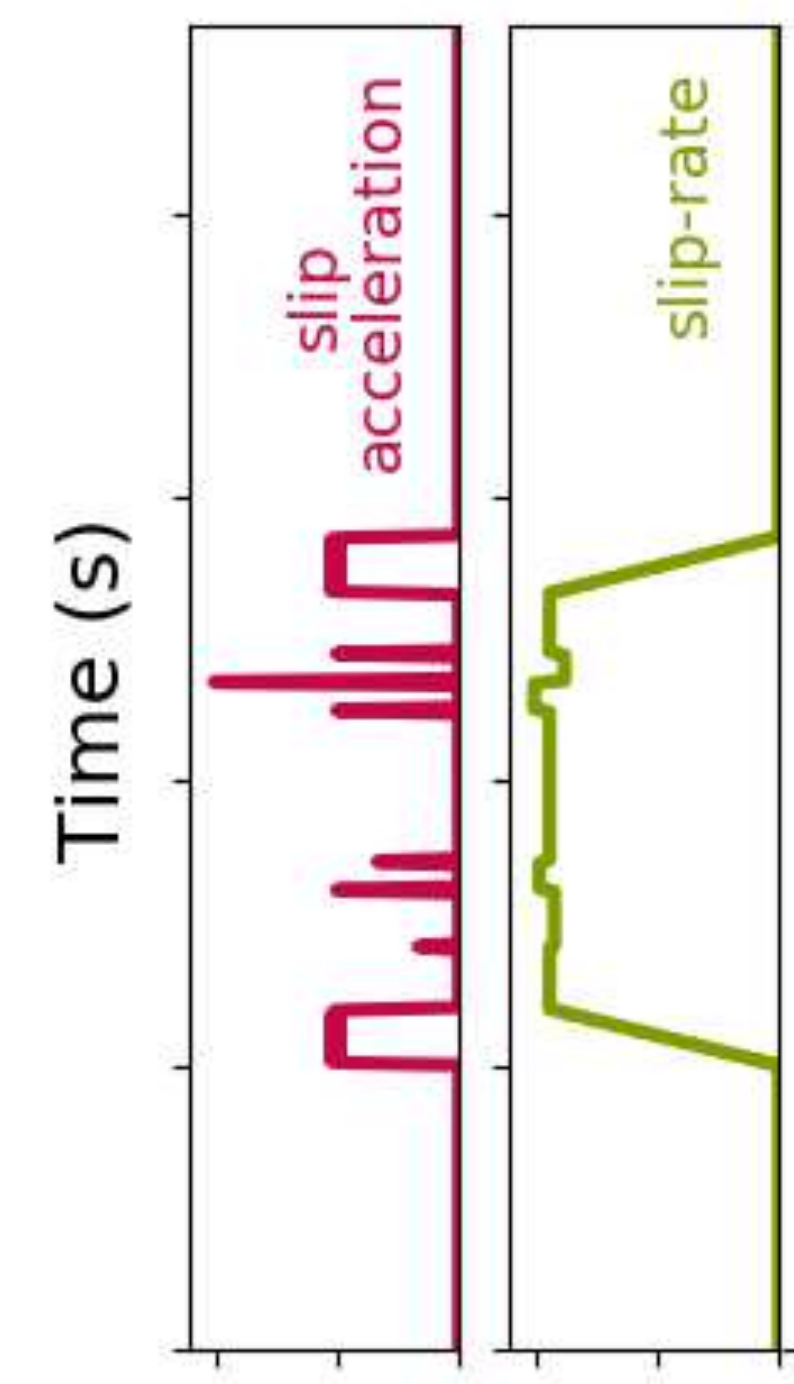
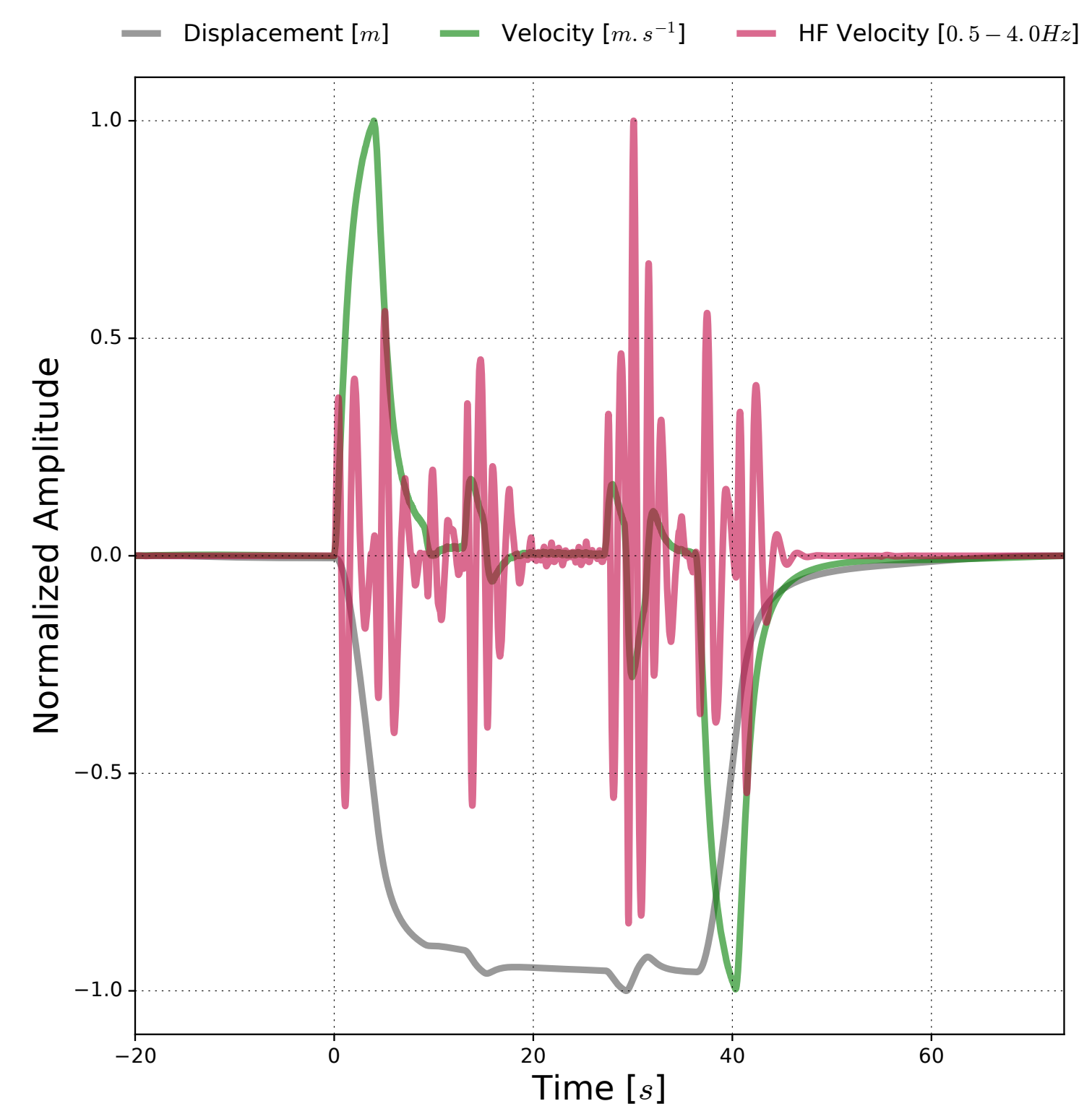
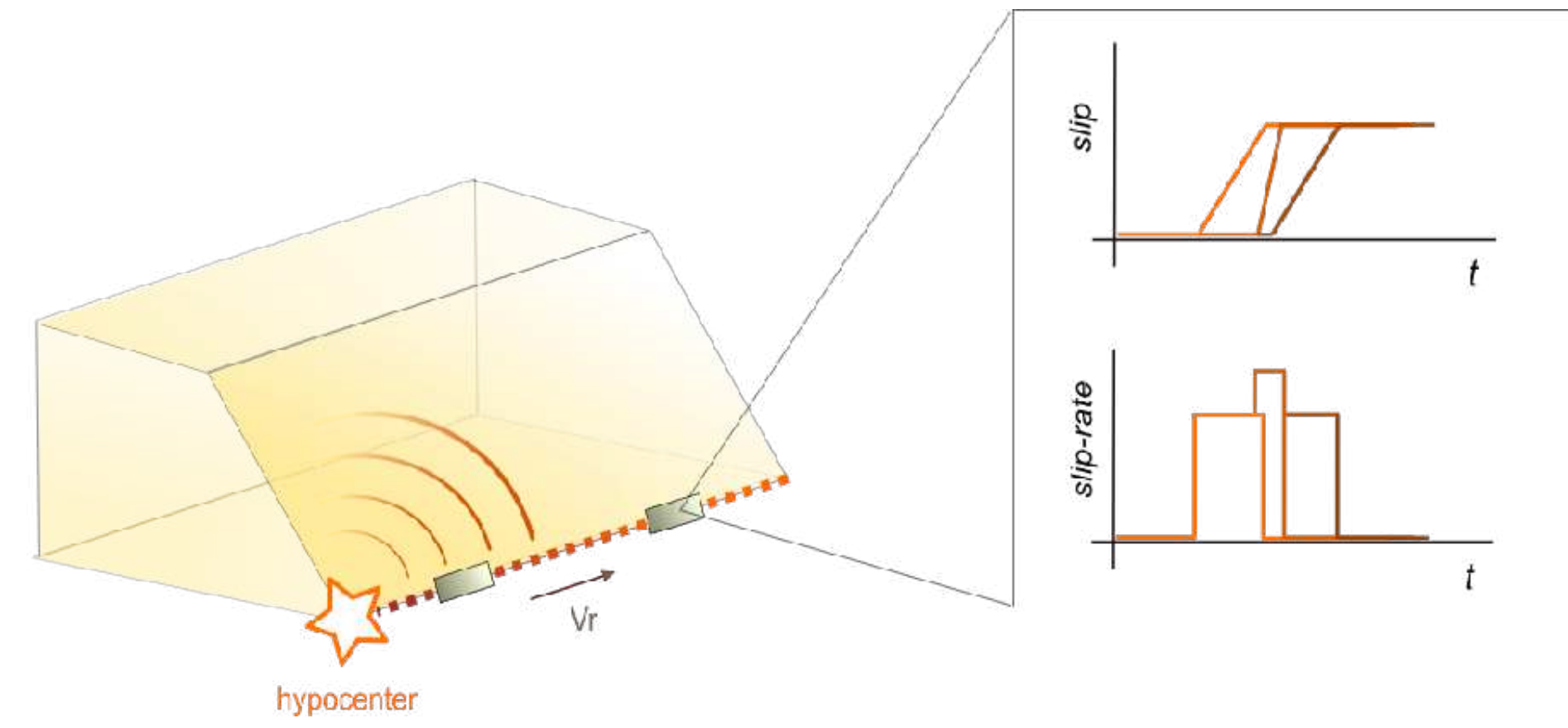
Synthetic test: homogenous model



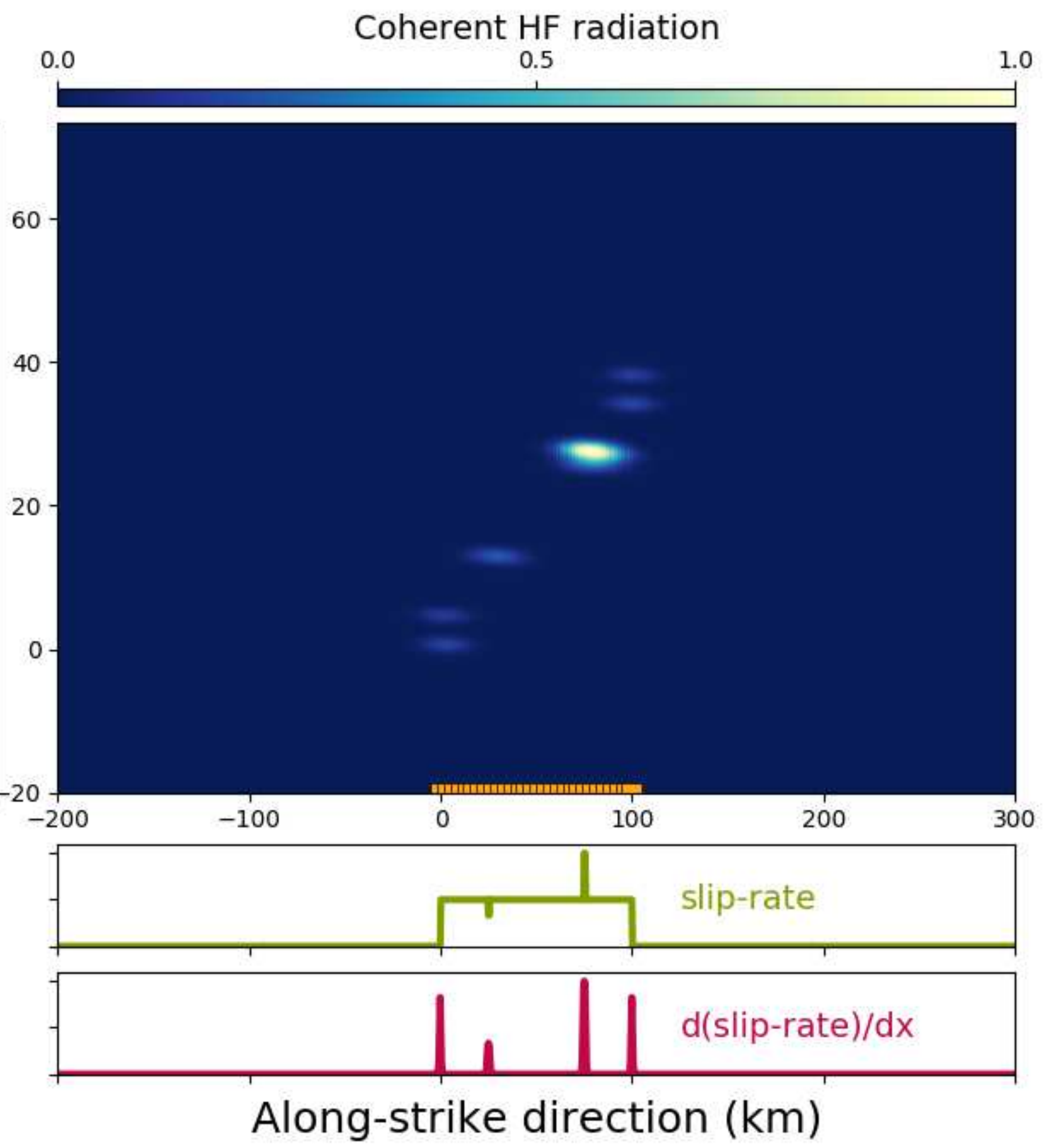
AK



Synthetic test: variable rise time



EU



High-frequency radiation linked to space-time variability of slip rate on the fault

Conclusions

Variability of fault-zone mechanical properties (e.g., coupling, asperities, geometry) controls the variability of earthquake radiation (high-frequency - low-frequency).

Study of seismicity and deformation before large earthquakes is necessary to characterize those properties and define future large earthquake scenarios.

New imaging techniques like back projection can help in illuminating the complex nature of earthquake rupture. Still need to better understand the link between high-frequency images and rupture process.

