

Early mantle differentiation and preservation of primordial reservoirs

The case of the atmosphere/mantle system

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Outline

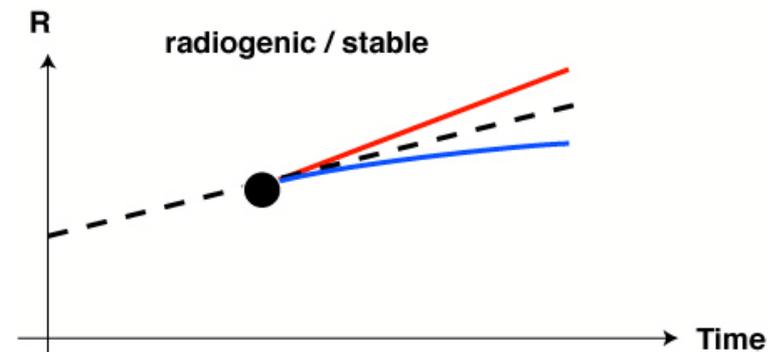
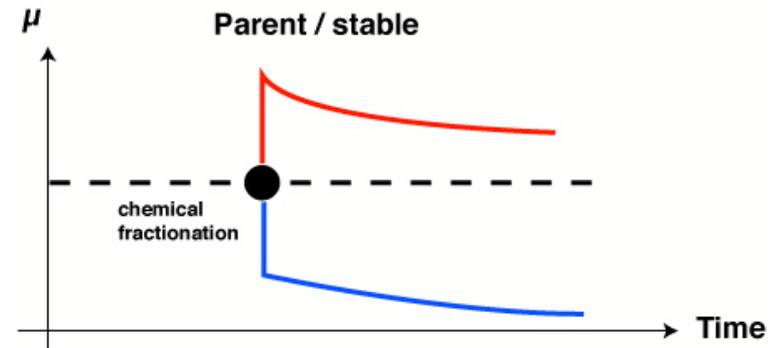
- Studying the Mantle – Atmosphere system
- When occurred mantle degassing ?
- A “primitive” undegassed reservoir in the deep mantle ?
- Is it preserved for 4.4Ga? If so, where it is ?

Isotope geochemistry made simple

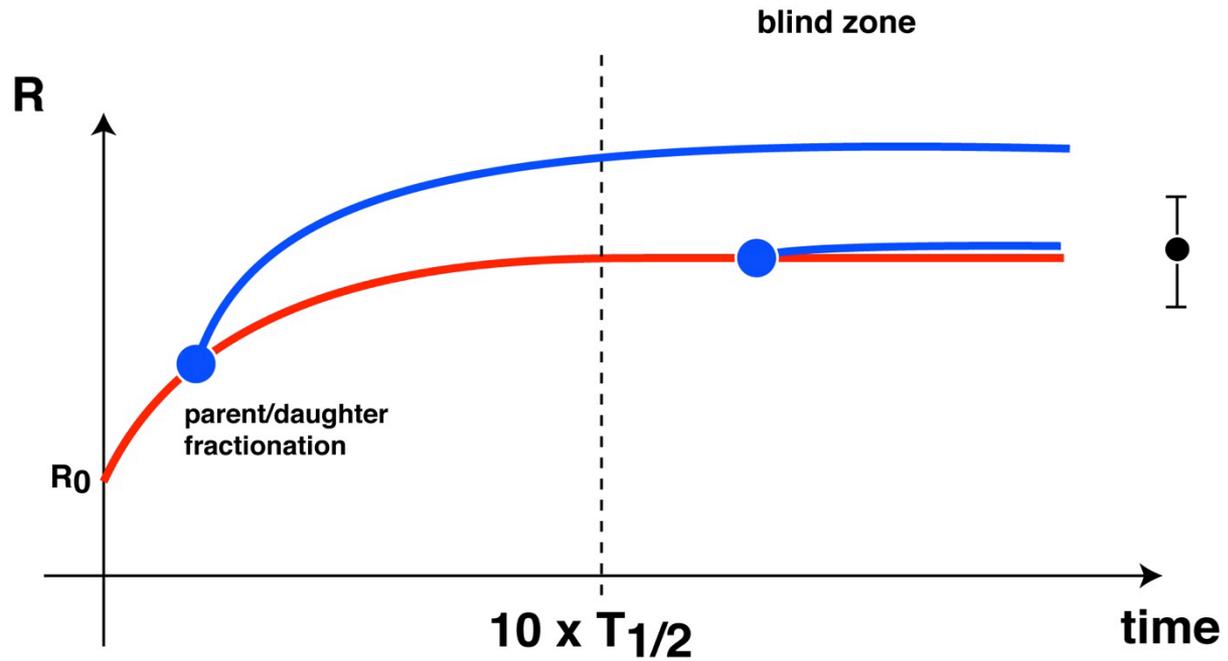
- A radiogenic isotopic ratio reflects a chemical fractionation and time
- A chemical fractionation is the translation to chemistry of a geological event (core formation, melting, crystallization, degassing, ...)
- “stable” isotopic ratios can fractionate during some chemical or physical processes but this will be negligible for radiogenic isotopes.

Radiogenic isotopic ratios (1): long period

- F=daughter, S=stable, P=parent
- $R=F/S$ (isotopic ratio)
- $\mu=P/S$ (chemical ratio)
- $R(t)=R_0+\mu(t)[e^{\lambda t}-1]$

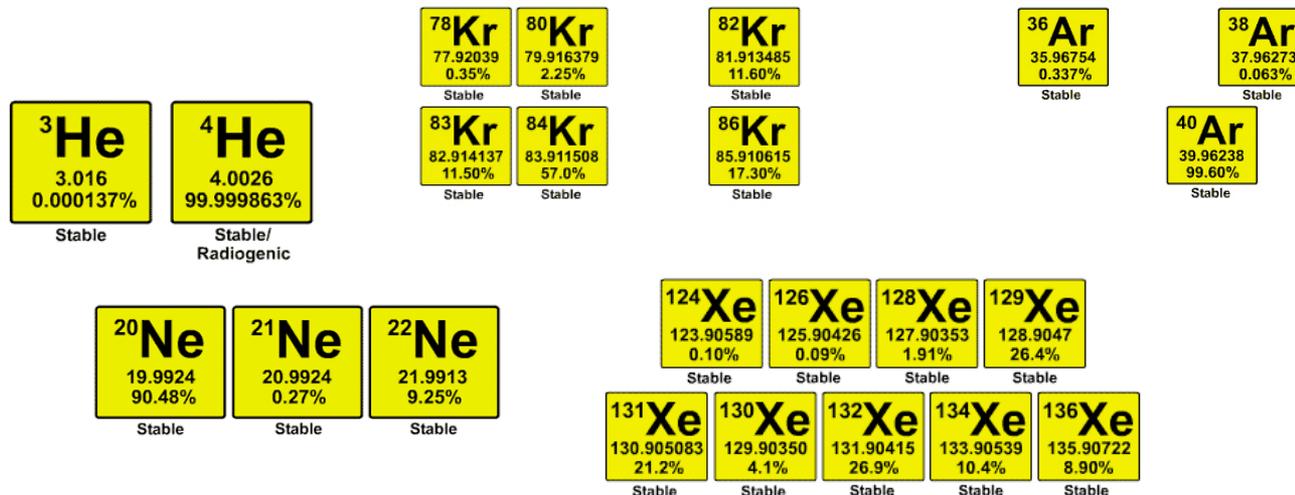


Radiogenic isotopic ratios (2): short period



Noble gas geochemistry: a tool to constrain mantle/Atmosphere evolution

- A single family – chemically inert
- Radiogenic isotopes are used to constrain time-evolution of the mantle/atmosphere system
- Non-radiogenic isotopes can be used as “stable” isotopes for physical processes during accretion

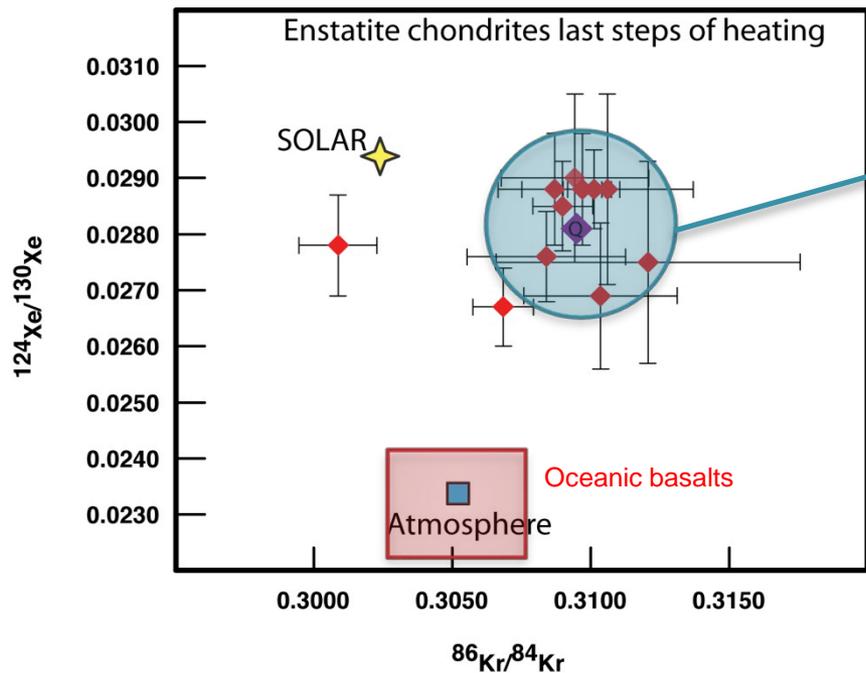


Noble gases of special interest today

- Helium: ^4He is radiogenic (U, Th), ^3He is primordial. Important: helium can leave the atmosphere: interesting for flux determination
- Neon: three isotopes, ^{21}Ne is “radiogenic”, ^{20}Ne , ^{22}Ne are stable
- Argon: ^{40}Ar is radiogenic (^{40}K). Important: $^{40}\text{Ar}_0 \sim 0$
- Xenon: couple ^{129}I - ^{129}Xe ($T_{1/2} := 17\text{Ma}$)

On the difficulty to have the composition of the primordial mantle

- Earth is **not** chondritic for noble gases
- It is unknown why
- Origin of the chondritic compositions is unknown



Carbonaceous Component
of unknown origin
Degassed at high T, acid
resistant

↓
Either not present either
lost during Earth accretion

Some evidences that (a part of) atmosphere is degassed from the mantle

- - ^{40}Ar is entirely of radiogenic origin
 - ^{40}Ar is 1% in atmosphere
 - So mantle was degassed to produce the atmosphere
- - Non-radiogenic isotopic ratios of the xenon are similar in the mantle and in the atmosphere, and this is a unique signature in the solar system
- - ^3He is detected above ridges or OIB, so the mantle is still degassing primordial isotopes

Evidences for a less degassed reservoir in deep Earth

(1) Argon budget (Allègre et al., 1995)

CC: [U]= 15 ± 3 ppb

EC: [U]= 8 ± 3 ppb

OC: [U]= 14 ± 4 ppb



CC: [K]=290ppm in silicate Earth

EC: [K]=152ppm in silicate Earth

OC: [K]=270ppm in silicate Earth



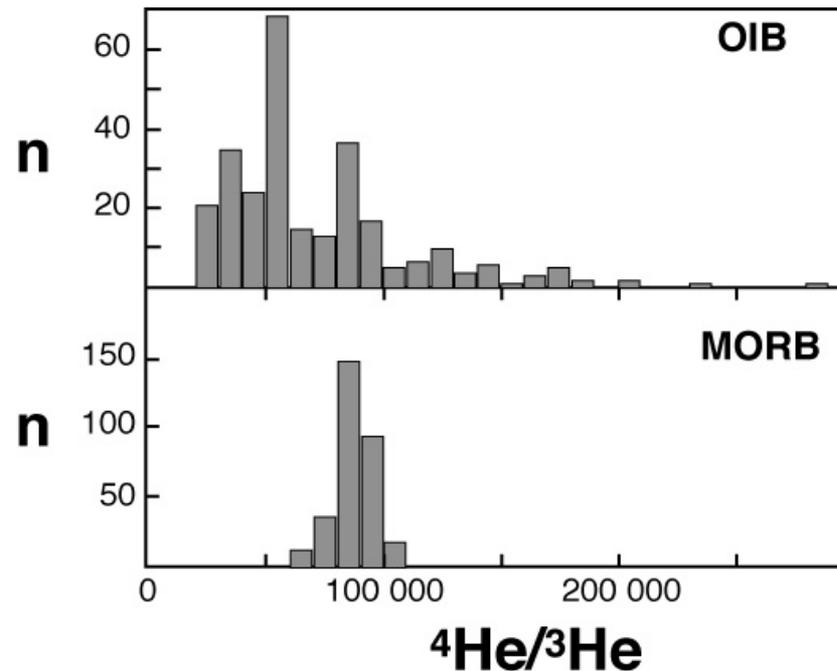
For CC: 57% of the ^{40}Ar is in the lower mantle; 5% in UM)

For EC: 20% of the ^{40}Ar is in the lower mantle; 10% in UM)

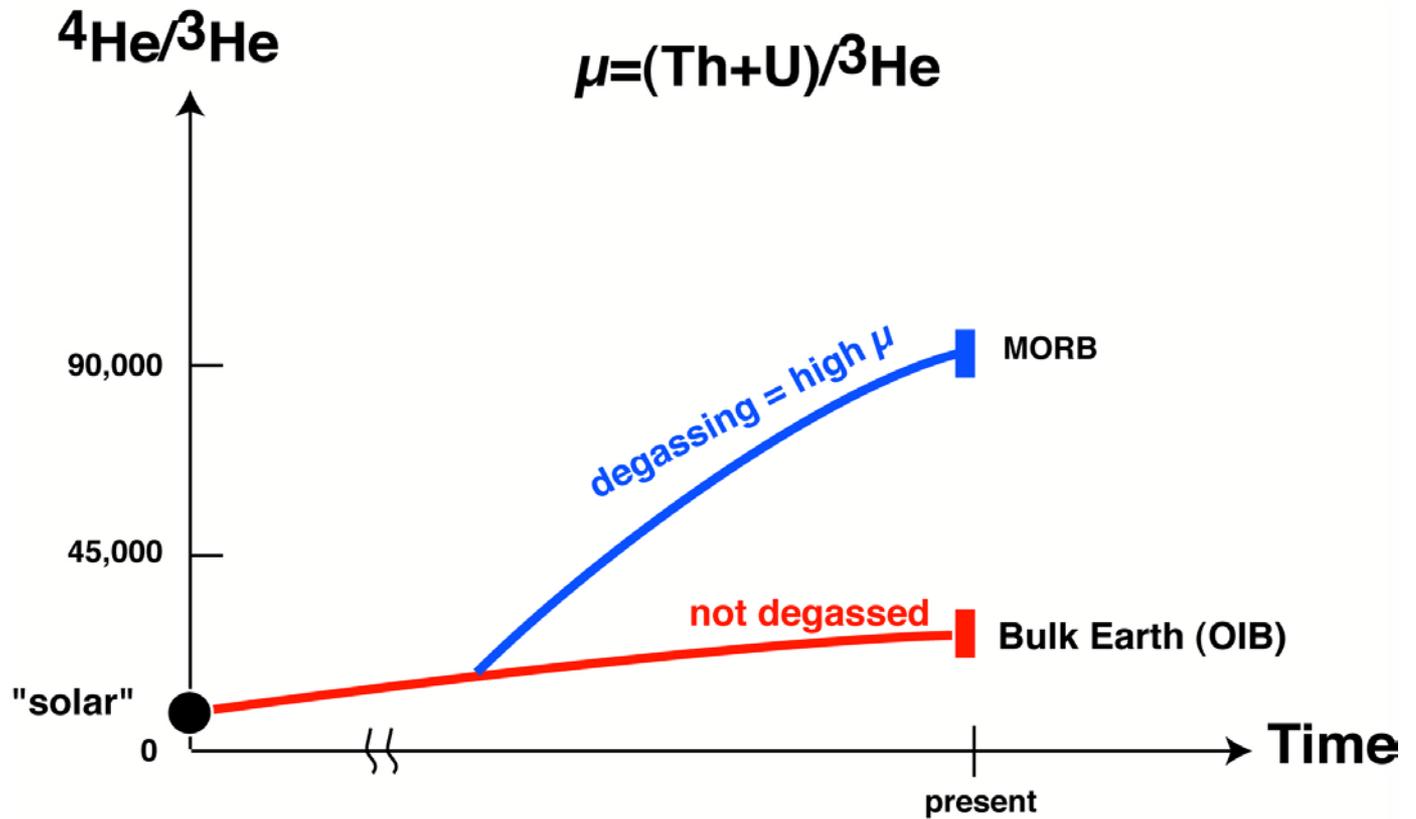
Mantle is not fully degassed of its radiogenic argon

Helium isotopic ratios as evidence for an undegassed mantle

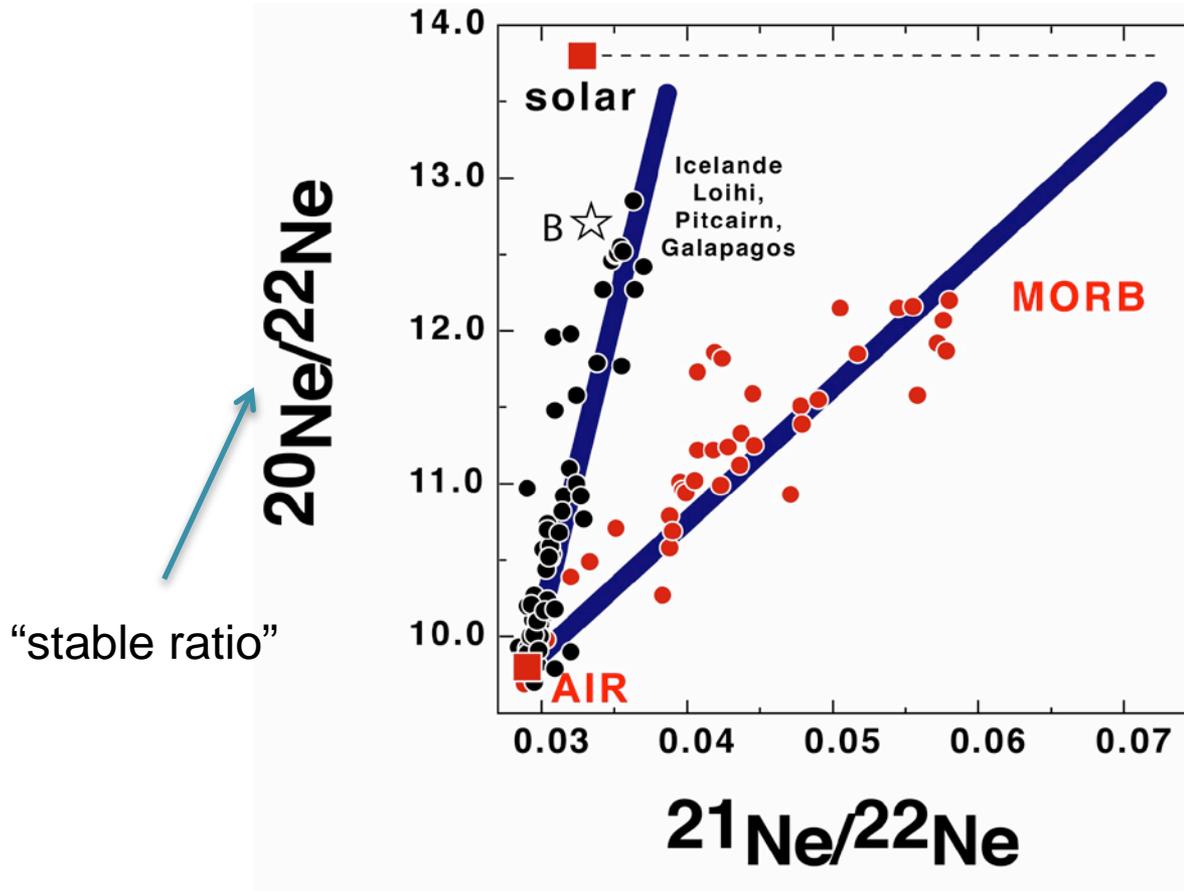
- ^4He is radiogenic (U, Th), ^3He is primordial
- MORB source is relatively homogeneous in respect with helium
- Most OIBs have lower $^4\text{He}/^3\text{He}$ ratios (lower $^3\text{He}/^4\text{He}$) than MORB



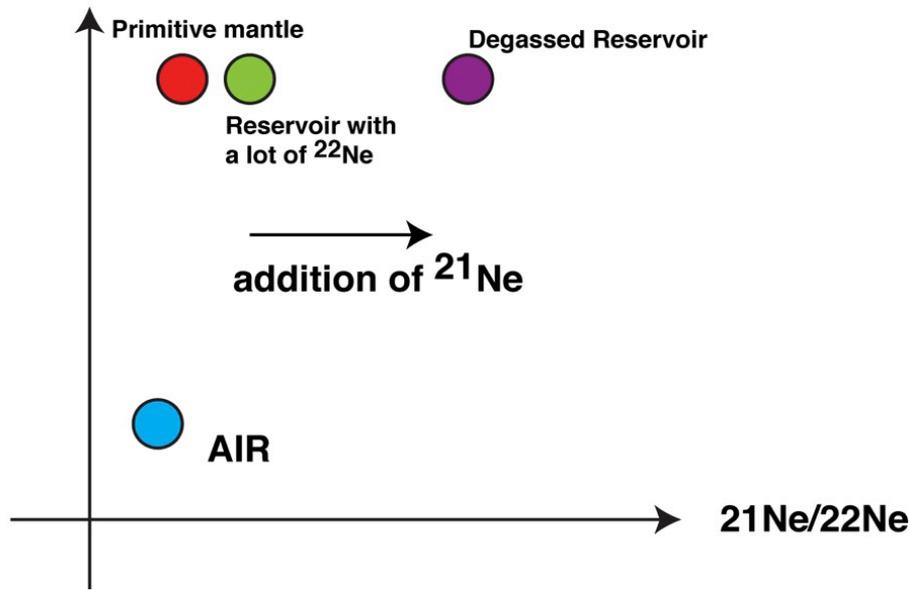
$$\left(\frac{{}^4\text{He}}{{}^3\text{He}} \right) \approx \left(\frac{{}^4\text{He}}{{}^3\text{He}} \right)_0 + \left(\frac{U + \text{Th}}{{}^3\text{He}} \right) \cdot f(t)$$



Neon isotopes



20Ne/22Ne

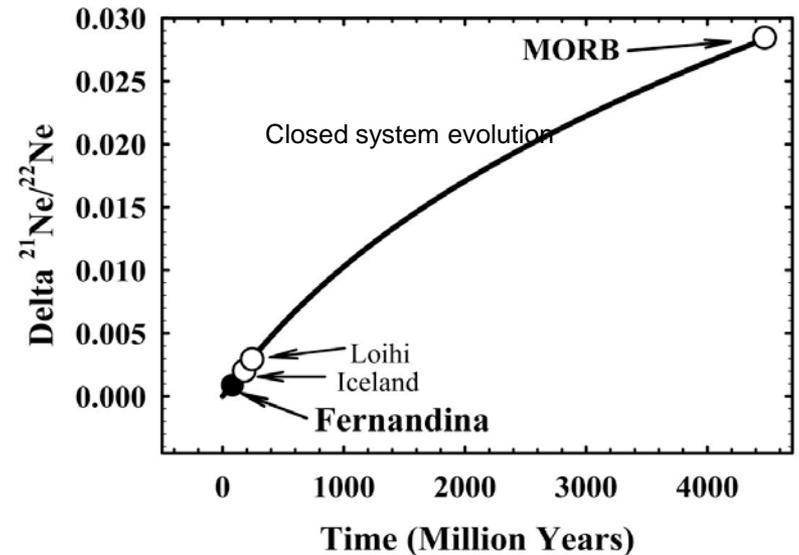
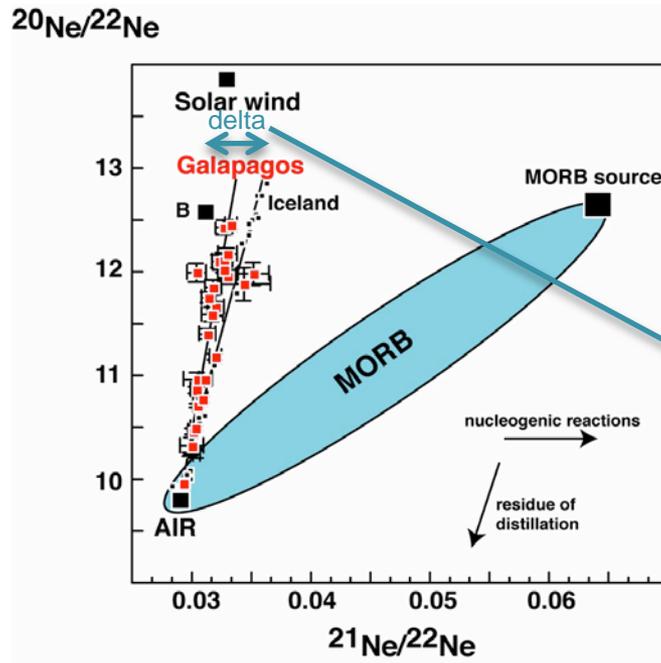


Month salary x2

Month salary x1.0000000000000000



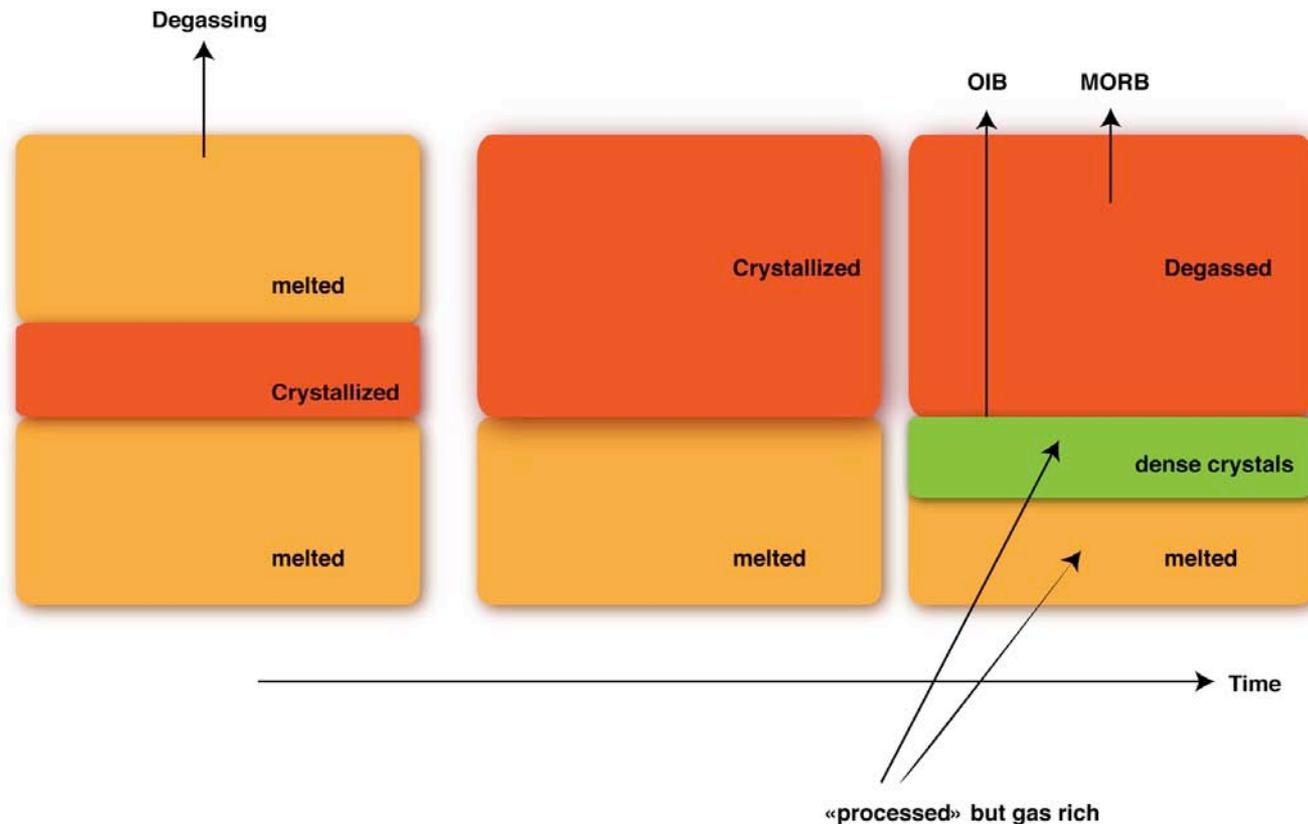
Neon suggests either it is primordial or can be a 4.4Ga depletion of a MORB-type source



Kurz et al. (2009)

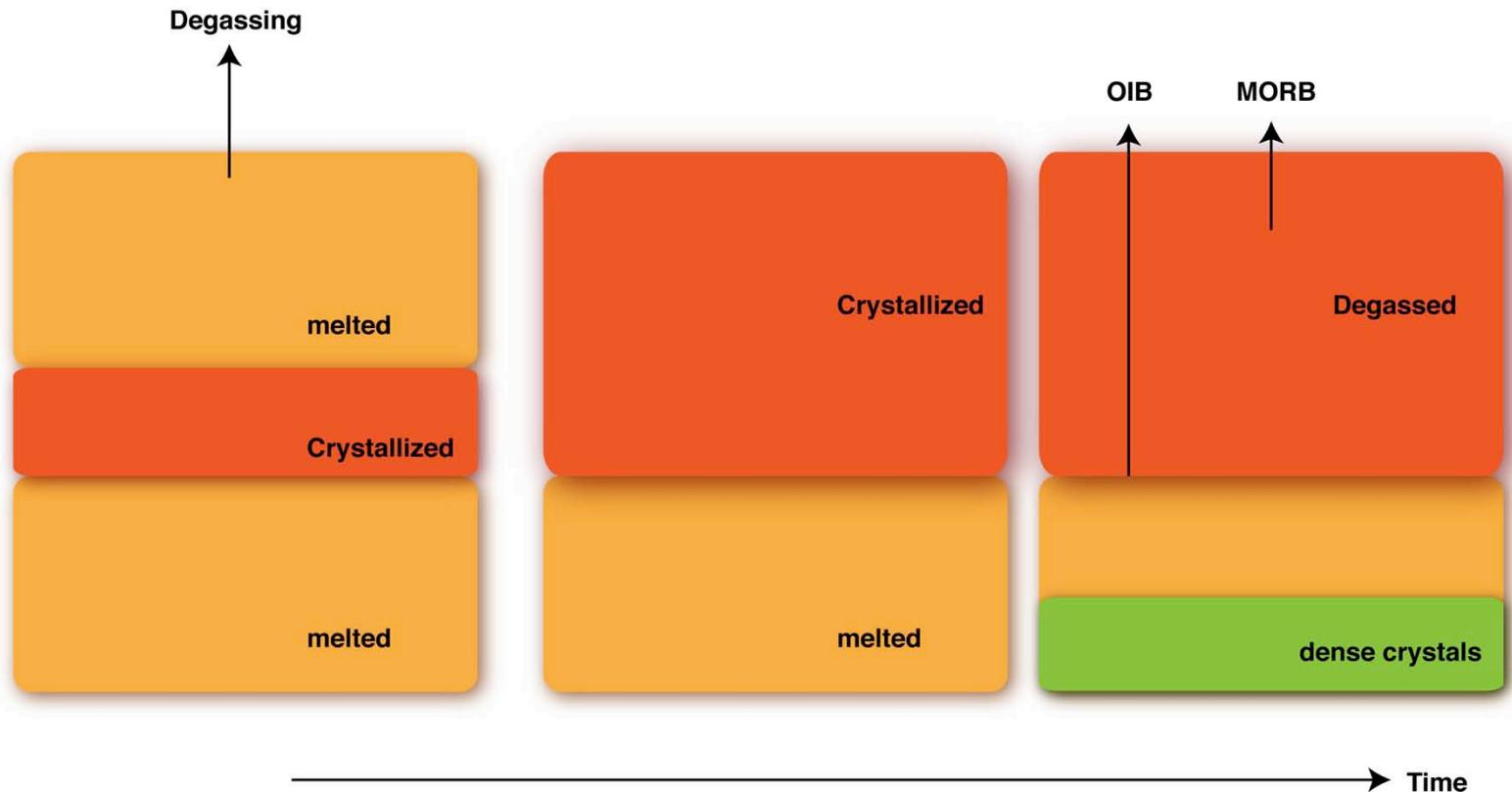
If depletion, it has to be older than 4.4Ga

- He and Ne isotopic compositions require a less degassed reservoir deep in the mantle
- He and Ne suggest that this reservoir is ancient and gas-rich. It does not mean it is primitive in the sense that it was never melted or crystallized, or never saw any recycled component in it.
- **The important process is degassing** (requires to be at surface)



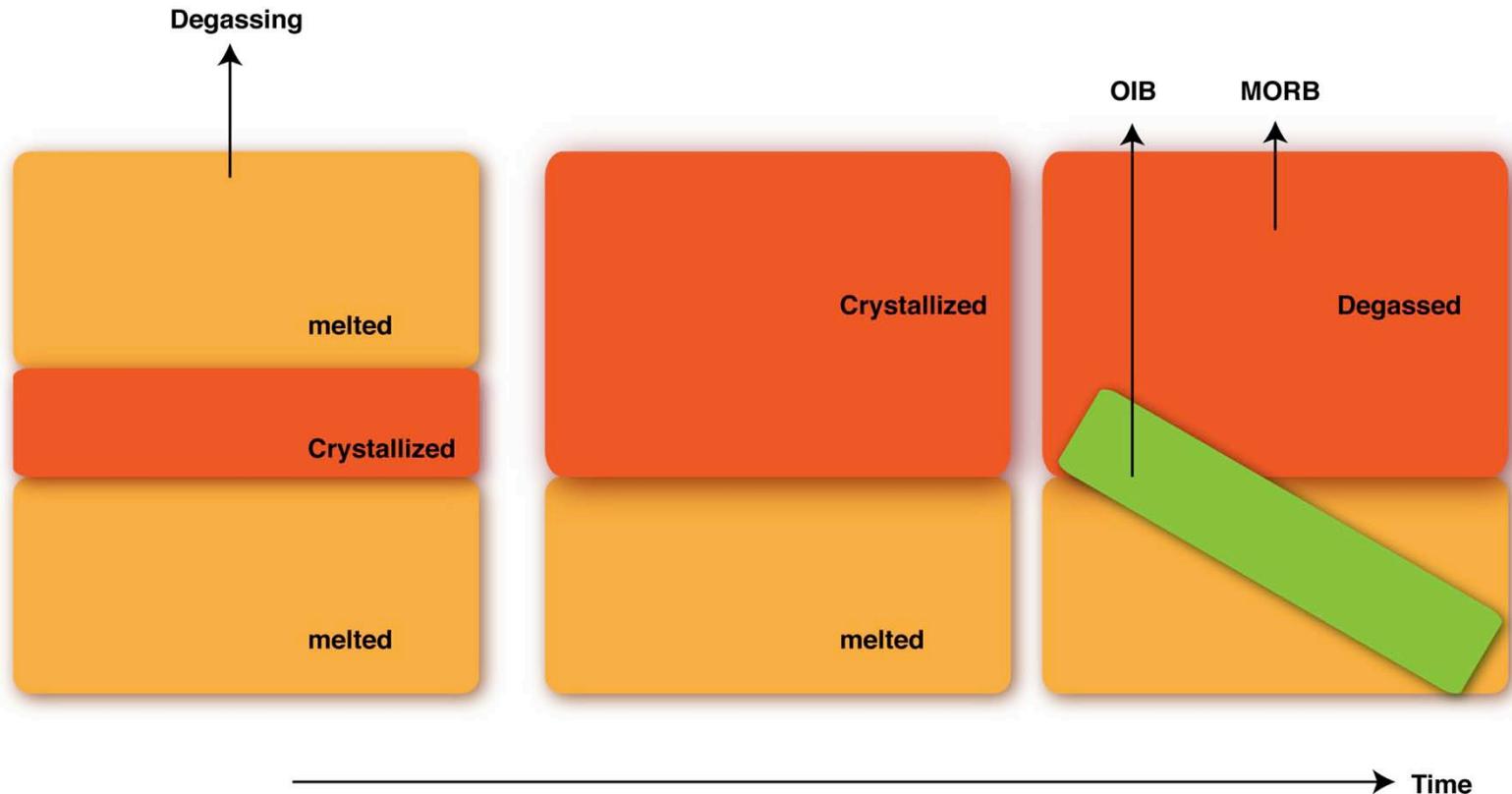
Inspired from Labrosse et al. (but with nicer colors)

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Inspired from Labrosse et al. (yesterday talk)

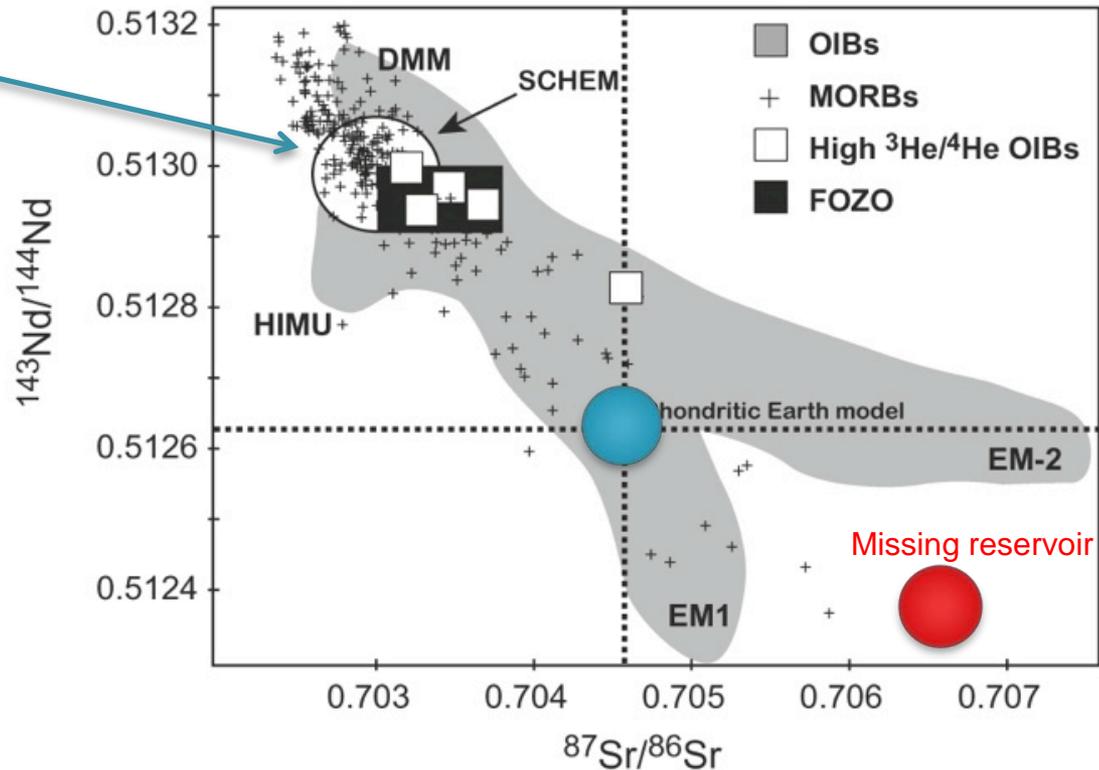
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Inspired from Labrosse et al. (next year talk)

Indeed, the high $^3\text{He}/^4\text{He}$ mantle is not chondritic for Nd

Reservoir deduced from ^{142}Nd excess

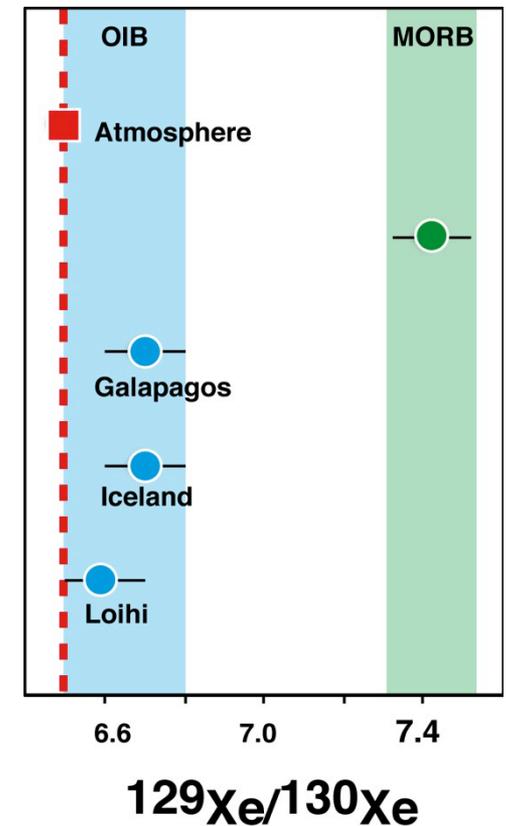
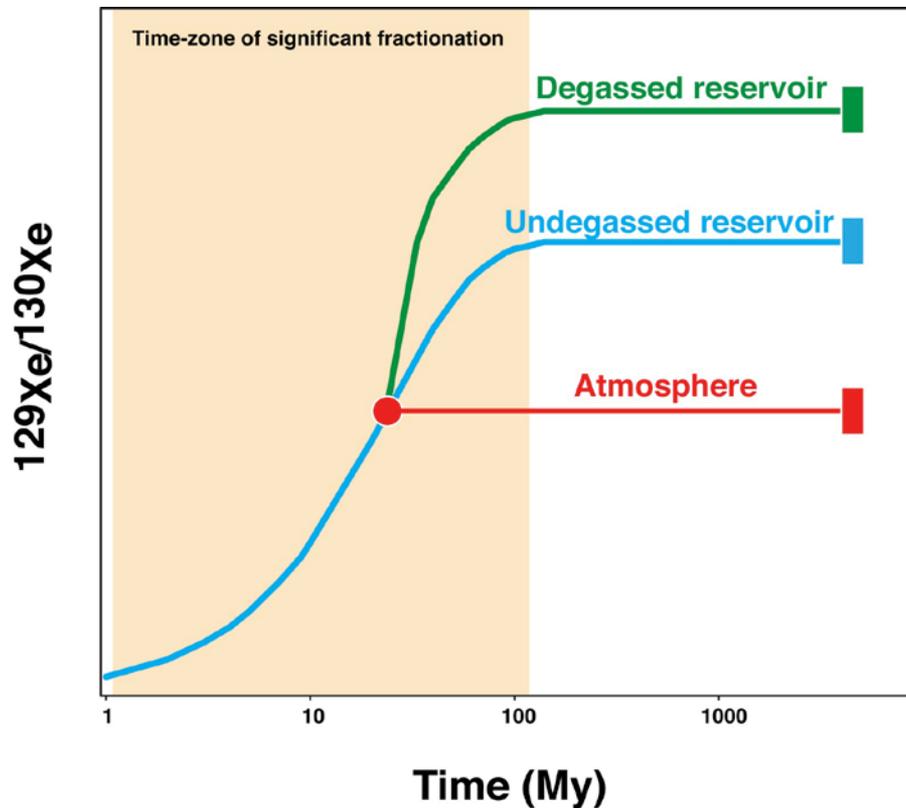


Caro and Bourdon (2010)

Either Earth is not chondritic for Nd, or there is some missing reservoir
Same process to explain non chondritic Xe isotopes ?

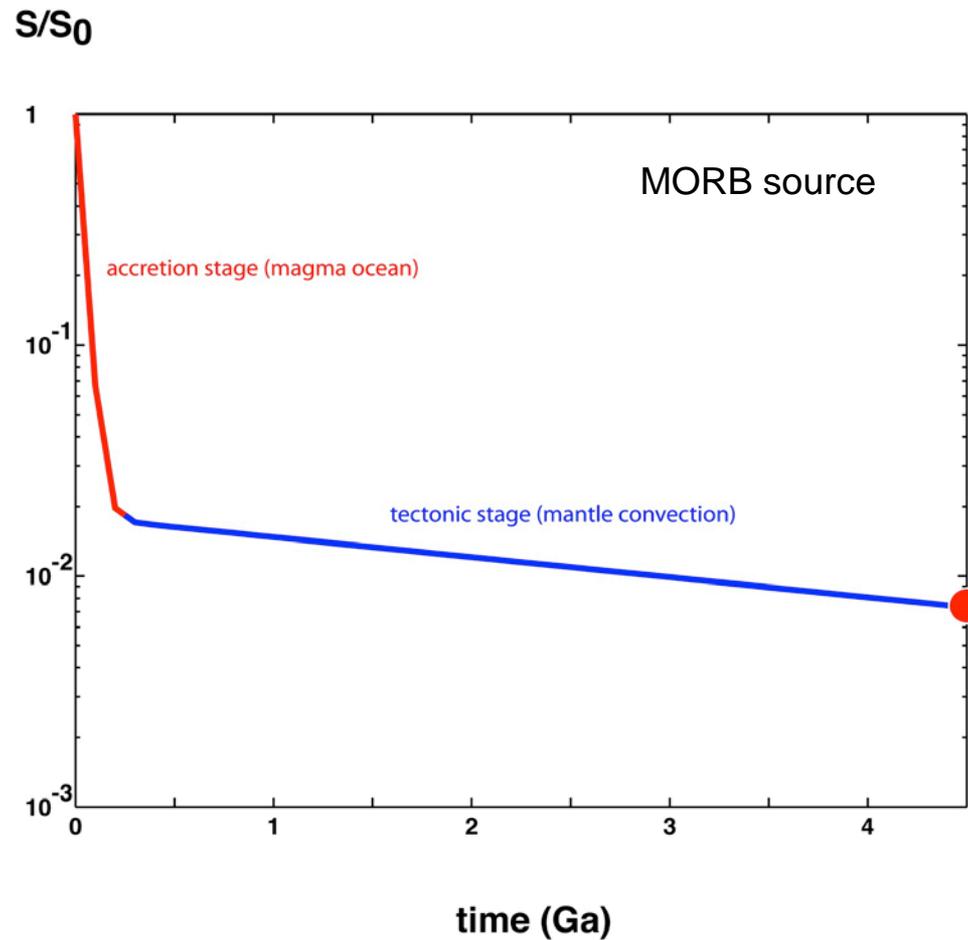
Evidence for an isolation of the OIB source since 4.4Ga

- Use of short life radioactive elements such as ^{129}I (17My) or ^{244}Pu (80My)



First order models of mantle degassing

- Use of K-Ar, ^3He flux, I-Xe



Conclusions 1

- Evidences that atmosphere is degassed from the mantle (massively during the first 100My, and then slower)
- MORB source is degassed
- A deep reservoir, sampled by oceanic island basalts, is rich in primordial He, Ne and in radiogenic Ar.
- It may not be primitive in the sense of lithophile elements, but it was not degassed.

Conclusions 2

- $^{129}\text{Xe}/^{130}\text{Xe}$ ratios suggest that MORB and OIB sources are separated since more than 4.4Ga
- Exact volume of the undegassed reservoir is unknown because primordial abundances are unknown since Earth is not chondritic for noble gases
- It can be the lower mantle (layered mantle)
- It can be located in piles at the bottom of the mantle (cf Labrosse et al. or Coltice et al.)
- Constraints on the magma ocean: cannot have a vigorous convection if whole mantle is melted
- Core is unlikely because noble gases are not siderophile elements.