

The InSight Mission Exploring the Interior of Mars



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Jet Propulsion Laboratory

November 20, 2014



InSight

JPL



cnes



DLR

ETH



Imperial College
London



CAB

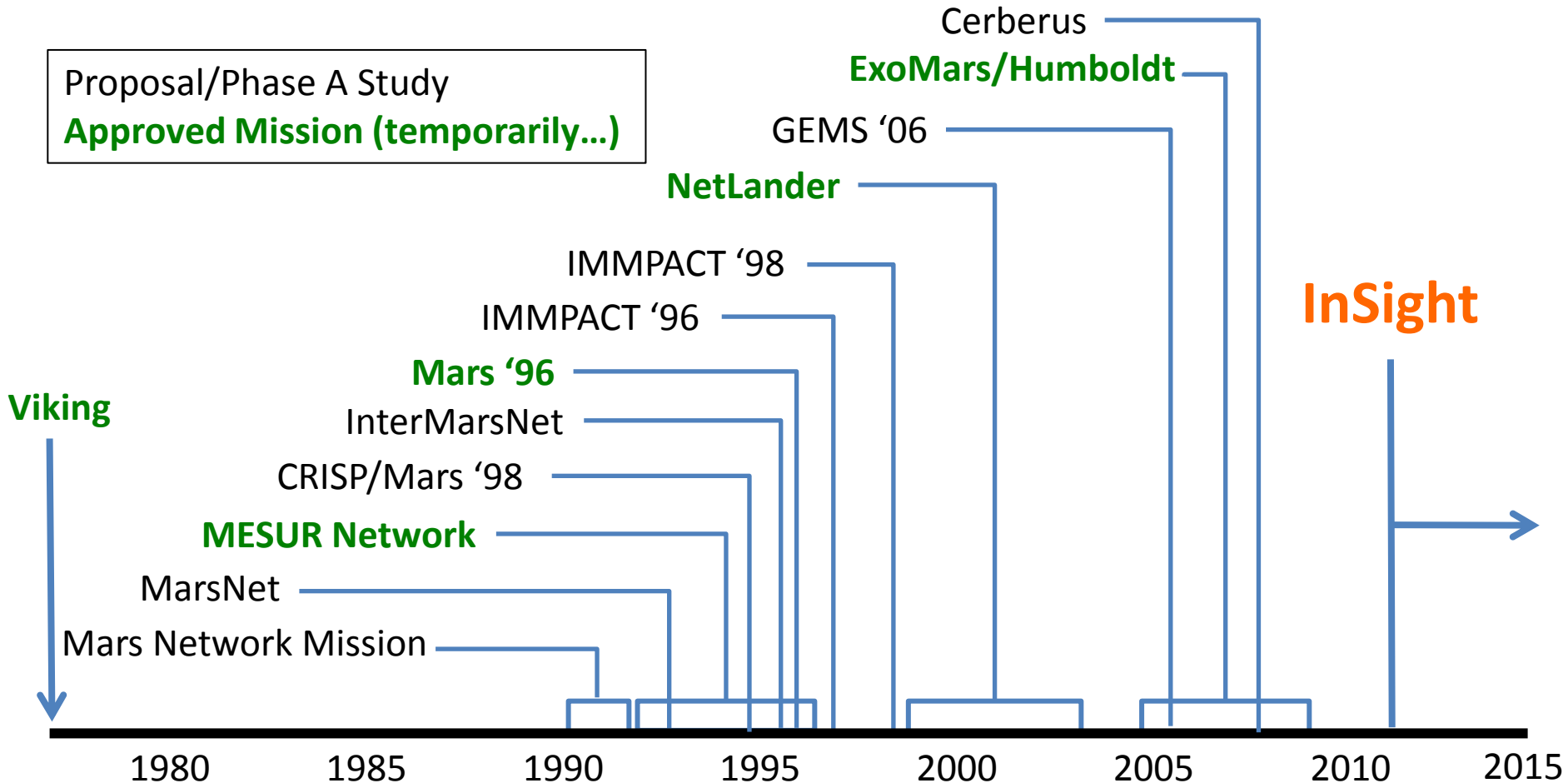


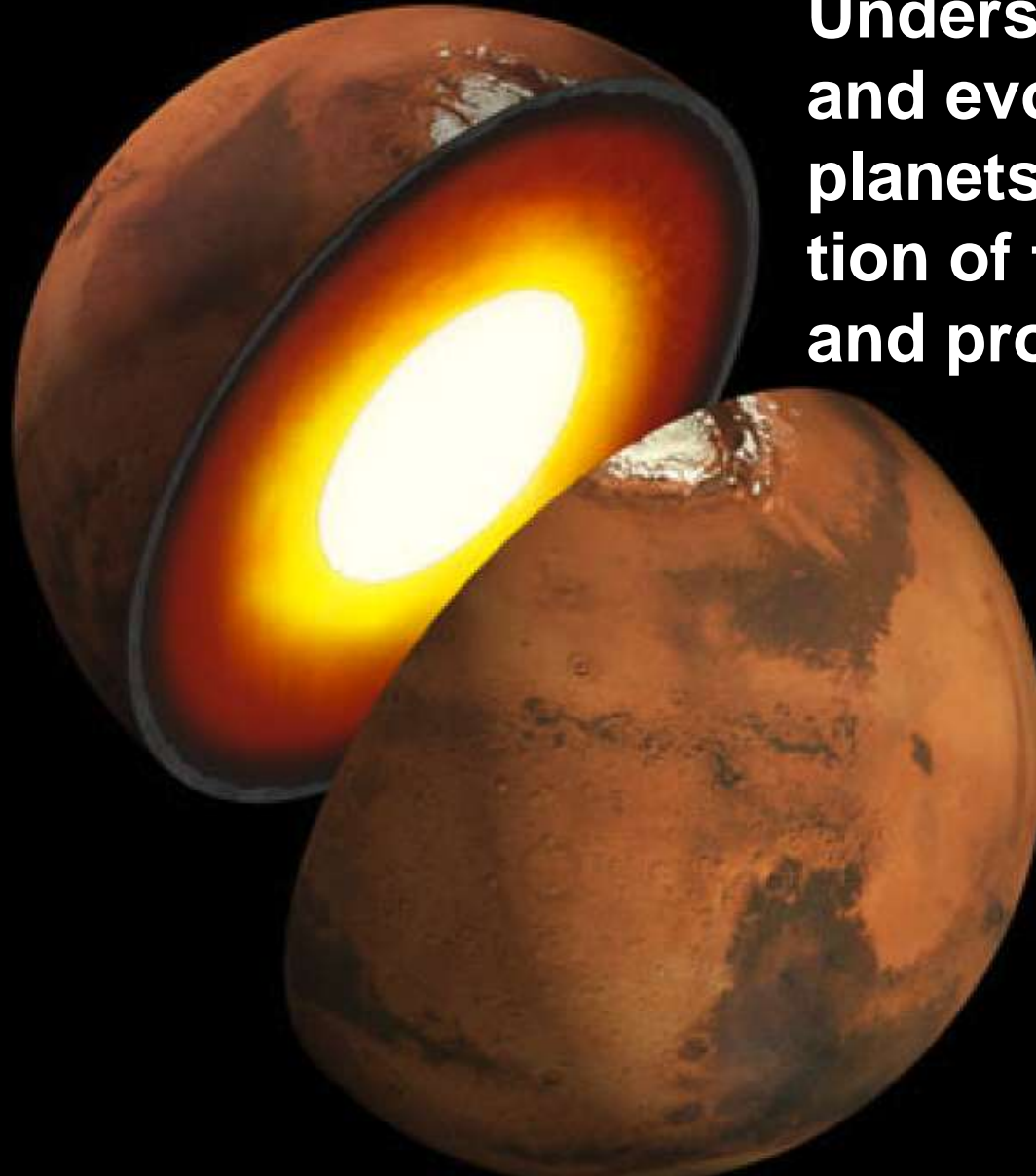
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- At the dawn of the age of planetary exploration, seismology was considered a key technique for understanding a planet and its interior.
 - The first instruments sent to the surface of another planet were seismometers.
 - Rangers 3–5; 1962
 - The two highest scientific priorities of the Apollo program were sample return and seismology.
 - Apollos 11, 12, 14, 15, 16; 1969–1977
 - The first landers sent to Mars carried seismometers.
 - Vikings 1, 2; 1975–1977

- Over the 35 years since Viking and Apollo, despite many proposals and several mission starts, there have been no further seismic investigations of the interior of any planet... until now!



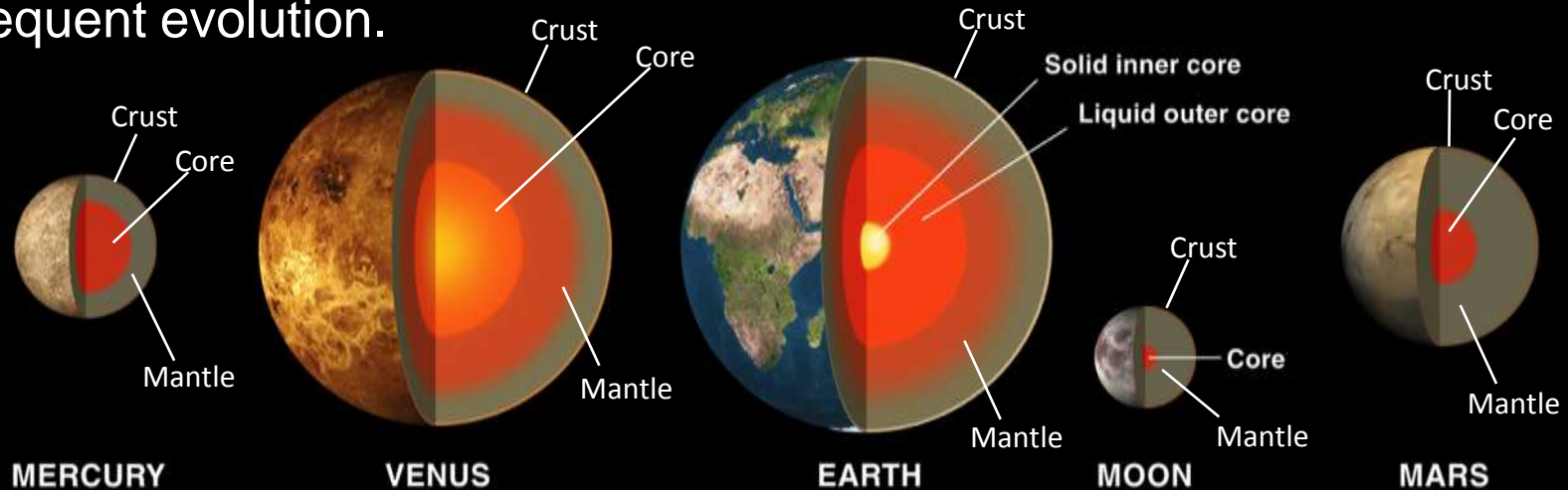


Understand the formation and evolution of terrestrial planets through investigation of the interior structure and processes of Mars.

- **Seismology**
- **Geodesy**
- **Heat Flow**

Mars is Key to Understanding Early Formation of Terrestrial Planets, Including Rocky Exoplanets

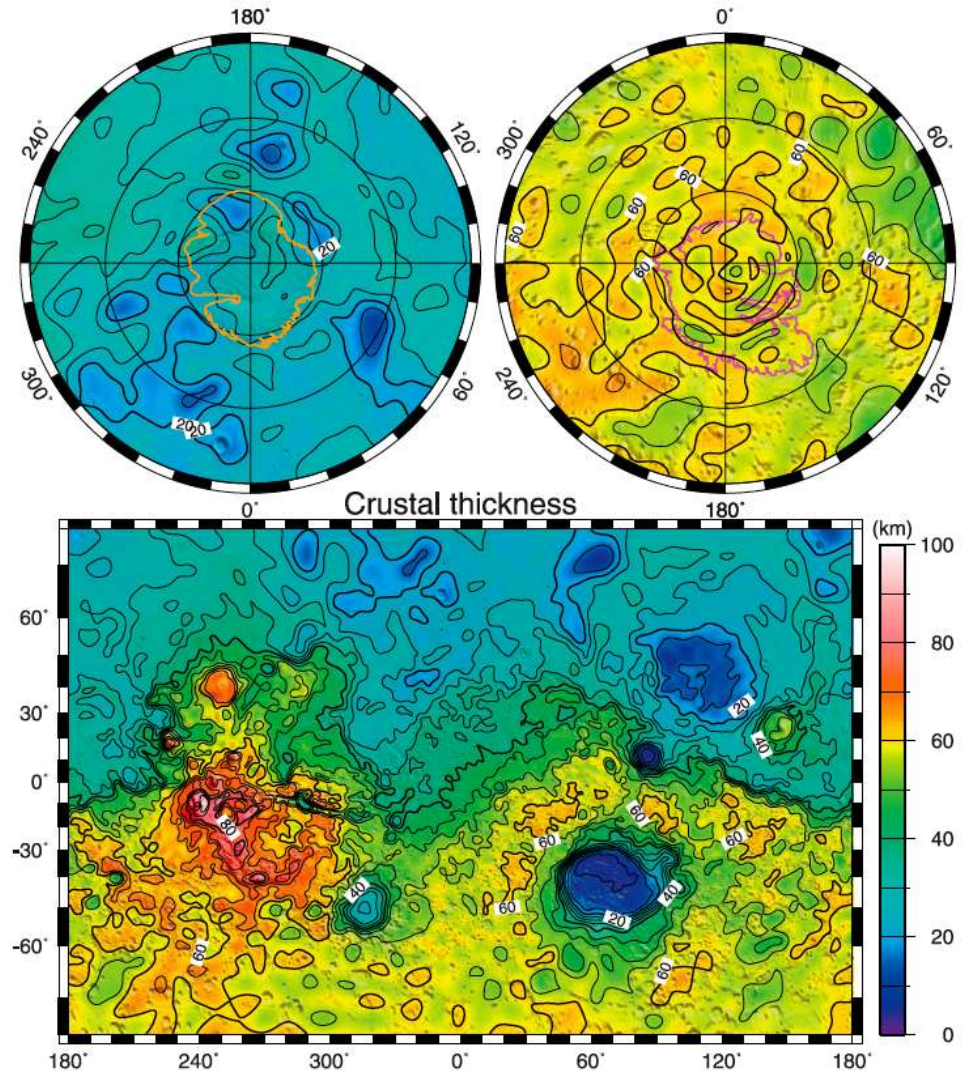
Terrestrial planets all share a common structural framework (crust, mantle, core), which is developed very shortly after formation and which determines subsequent evolution.



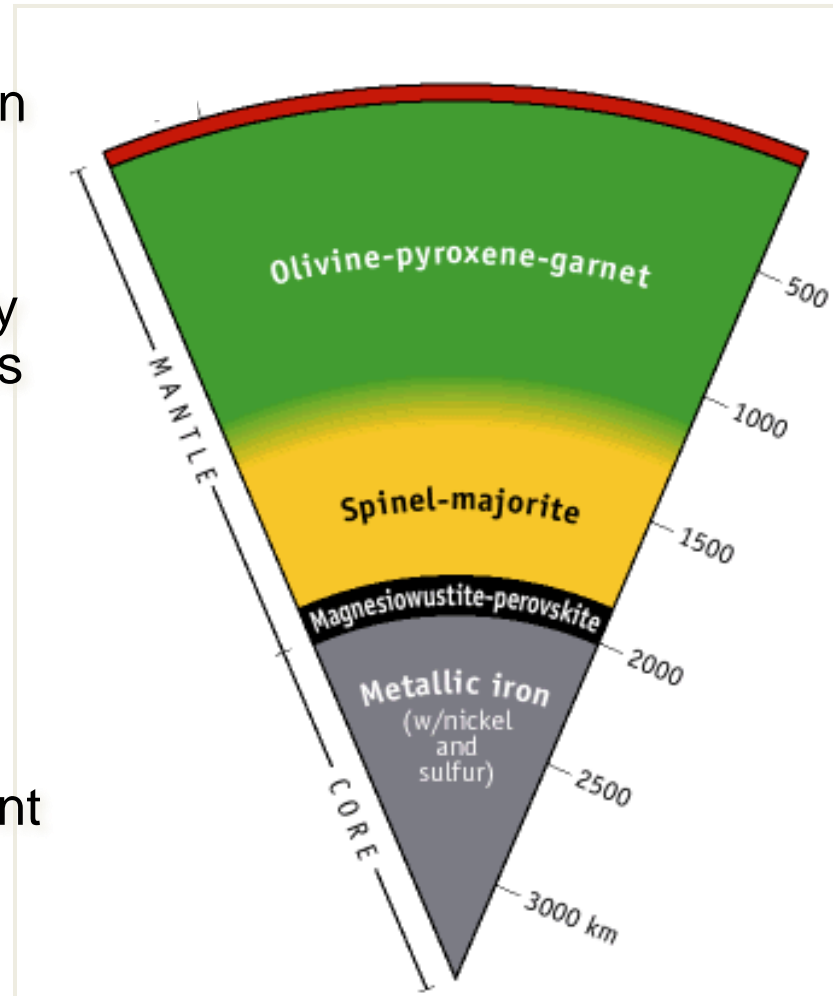
Mars is uniquely well-suited to study the common processes that shape all rocky planets and govern their basic habitability.

- There is strong evidence that its basic crust and mantle structure have survived little changed from the first few hundred Myr of formation.
 - Ancient crustal surface age from crater and QCD density
 - SNC evidence for persistent distinct mantle reservoirs, Ar isotope ratios
- Its surface is much more accessible than Mercury, Venus.
- Our knowledge of its geology, chemistry, climate history provides a rich scientific context for using interior information to increase our understanding of the solar system.

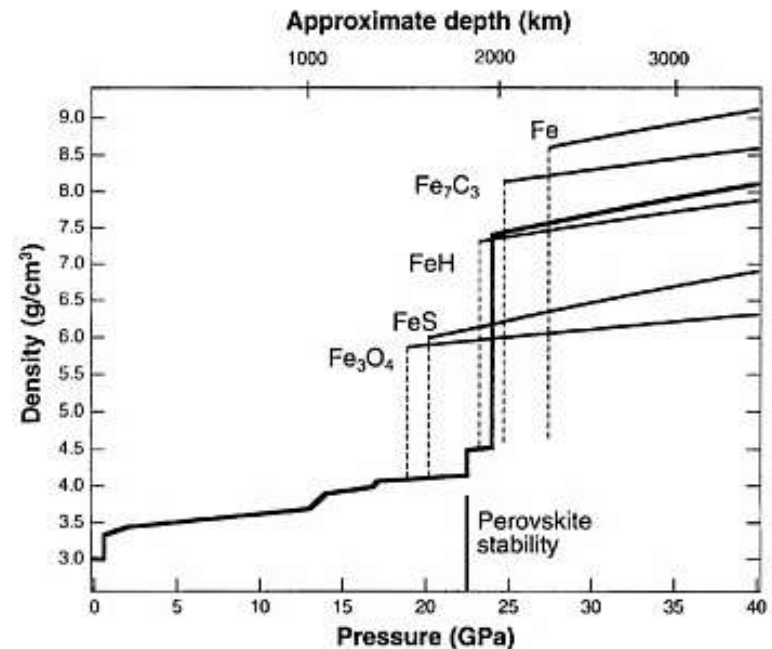
- From orbital measurements we have detailed information on variations in crustal thickness (assumes uniform density).
- But we don't know the volume of the crust to within a factor of 2.
- Does Mars have a layered crust? Is there a primary crust beneath the secondary veneer of basalt?
- Is the crust a result of primary differentiation or of late-stage overturn?
- These questions and more can be addressed by a simple thickness measurement.



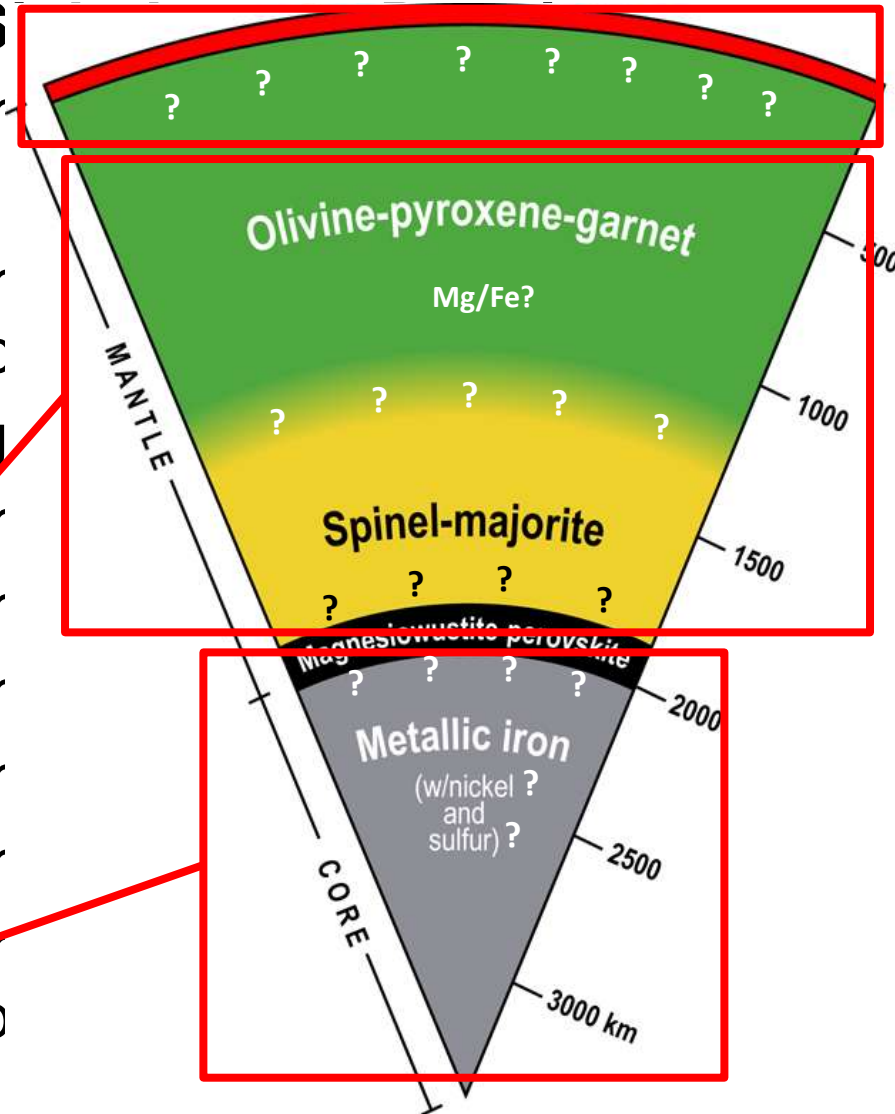
- What is the actual mantle density, which can be related to composition (e.g., Mg#, mineralogy, volatile content)?
- To what degree is it compositionally stratified? What are the implications for mantle convection?
- Are there polymorphic phase transitions?
- What is the thermal state of the mantle?
- All can be addressed to some extent by basic seismic observations.



- Radius is 1600 ± 300 km, density is uncertain to $\pm 15\%$
- Composed primarily of iron, but what are the lighter alloying elements?
- At least the outer part appears to be liquid; is there a solid inner core?
- How do these parameters relate to the initiation and shut down of the dynamo?
- Does the core radius preclude a lower mantle perovskite transition?
- Without radius and shear measurements we are stuck with a family of possible core structures, each with significantly different implications for Mars' origin and history.



- Crust:** Its **thickness** and vertical structure (**layering** of different compositions) reflects the depth and crystallization processes of the magma ocean and the early post-differentiation evolution of the planet (plate tectonics vs. crustal overturn vs. immobile crust vs. ...).
- Mantle:** Its behavior (e.g., convection, partial melt generation) determines the manifestation of the thermal history on a planet's surface; depends directly on its **thermal structure** and **stratification**.
- Core:** Its **size** and composition (**density**) reflect conditions of accretion and early differentiation; its **state** (liquid vs. solid) reflects its composition and the thermal history of the planet.

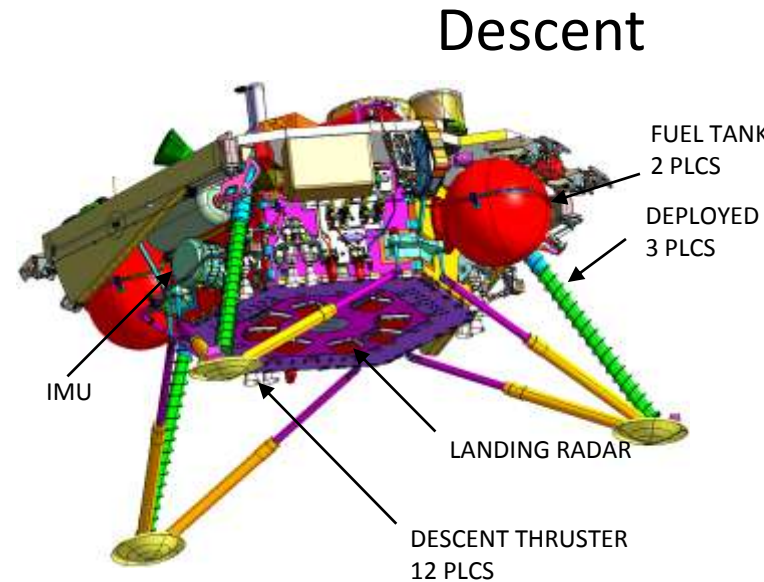
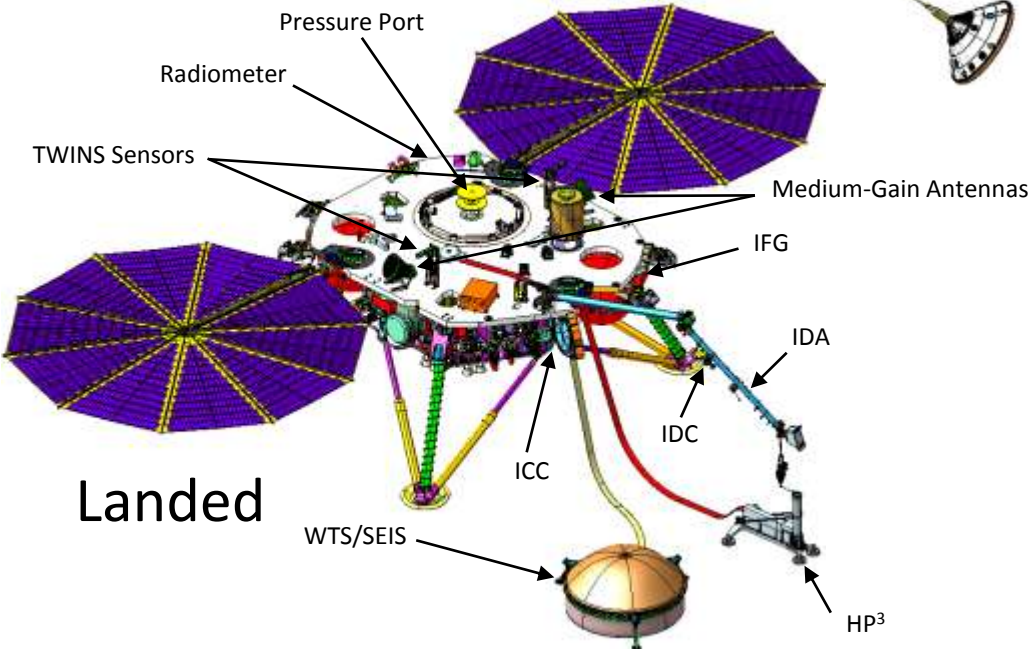
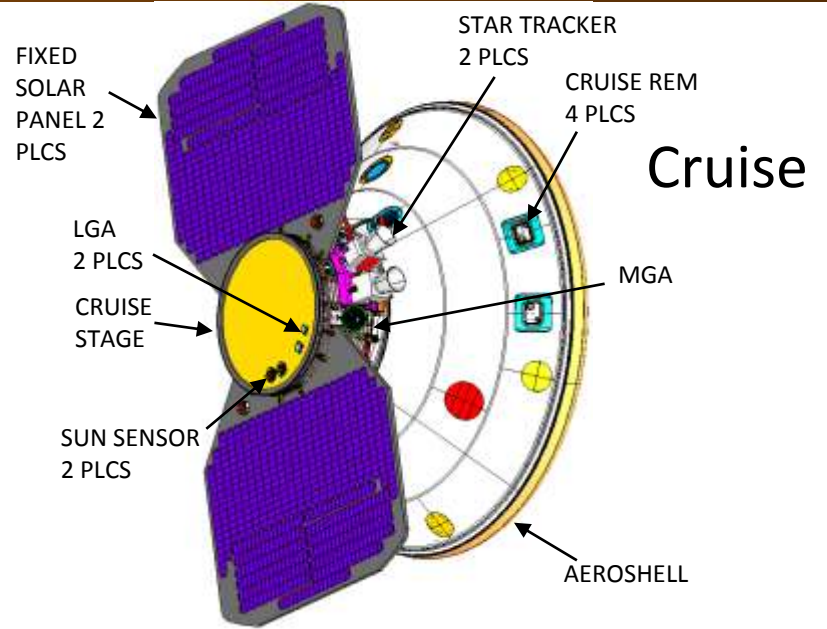
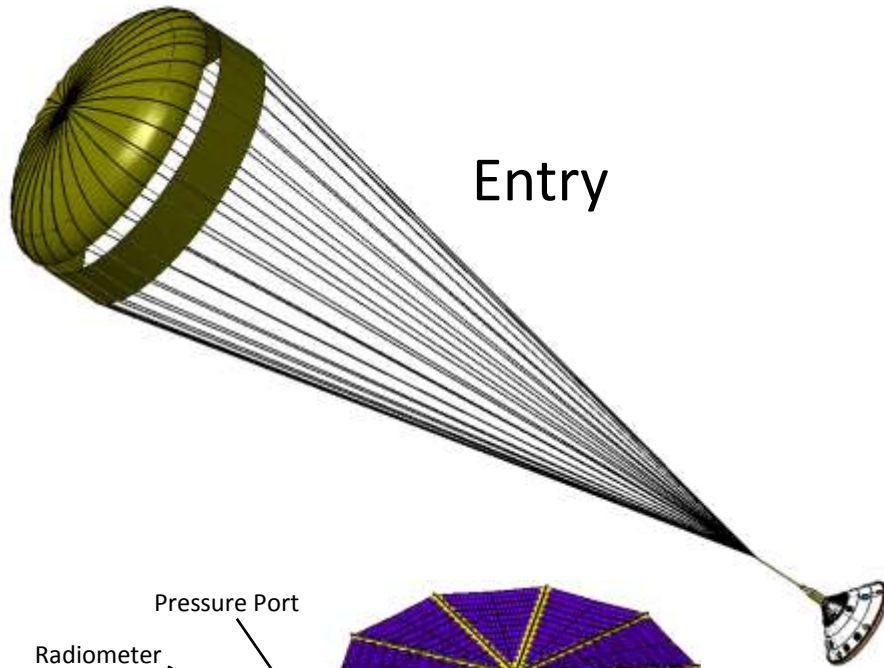


What Do We Know About the Interior of Mars?

Measurement	Current Uncertainty	InSight Capability	Improvement
Crustal thickness	65±35 km (inferred)	±5 km	7X
Crustal layering	no information	resolve 5-km layers	New
Mantle velocity	8±1 km/s (inferred)	±0.13 km/s	7.5X
Core liquid or solid	“likely” liquid	positive determination	New
Core radius	1700±300 km	±75 km	4X
Core density	6.4±1.0 gm/cc	±0.3 gm/cc	3X
Heat flow	30±25 mW/m ² (inferred)	±3 mW/m ²	8X
Seismic activity	factor of 100 (inferred)	factor of 10	10X
Seismic distribution	no information	locations ≤10 deg.	New
Meteorite impact rate	factor of 6	factor of 2	3X

- InSight will fly a near-copy of the successful Phoenix lander
- Launch: **March 4-24, 2016** from **Vandenberg AFB, California**
- Fast, type-1 trajectory, 6.5-mo. cruise to Mars
- Landing: September 28, 2016
- 67-sol deployment phase
- Two years (one Mars year) science operations on the surface; repetitive operations
- Nominal end-of-mission: October 6, 2018







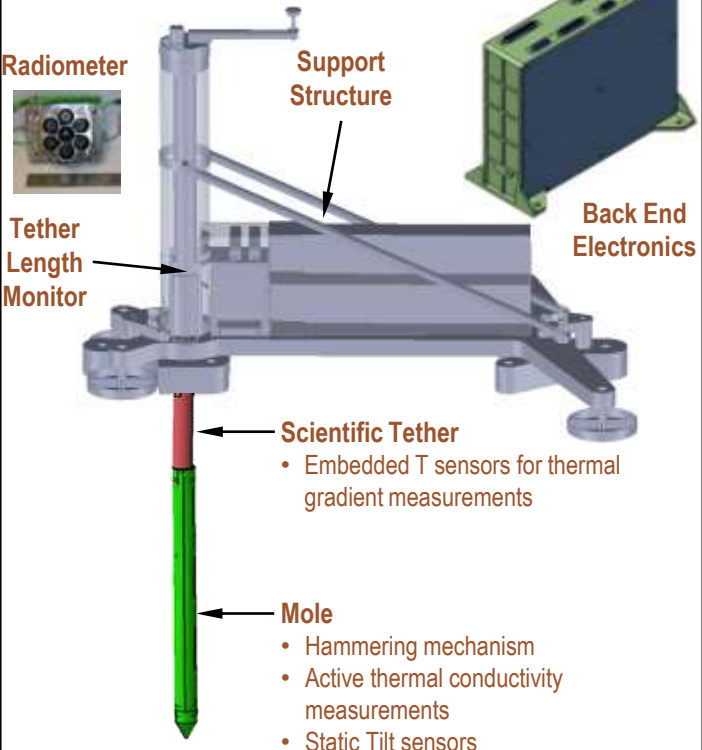
RISE (S/C Telecom)

Rotation and Interior Structure Experiment

Small Deep Space Transponder

HP³ (DLR)

Heat Flow and Physical Properties Probe



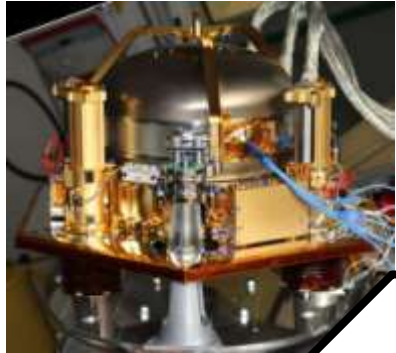
Wind and Thermal Shield (WTS)

SEIS (CNES)

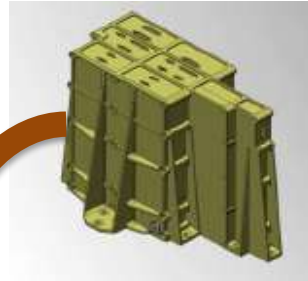
(also IPGP, ETH/SSA, MPS/DLR, IC/Oxford/UKSA, JPL/NASA)
Seismic Experiment for Interior Structure



Thermal Enclosure (RWEB)



Sensor Head: VBB, SP, LVL



Electronics (Ebox)

Tether (THR)

IDS (JPL)

Instrument Deployment System



IDA – Instrument Deployment Arm

IDC – Instrument Deployment Camera

ICC – Instrument Context Camera



IFG (UCLA) – InSight FluxGate



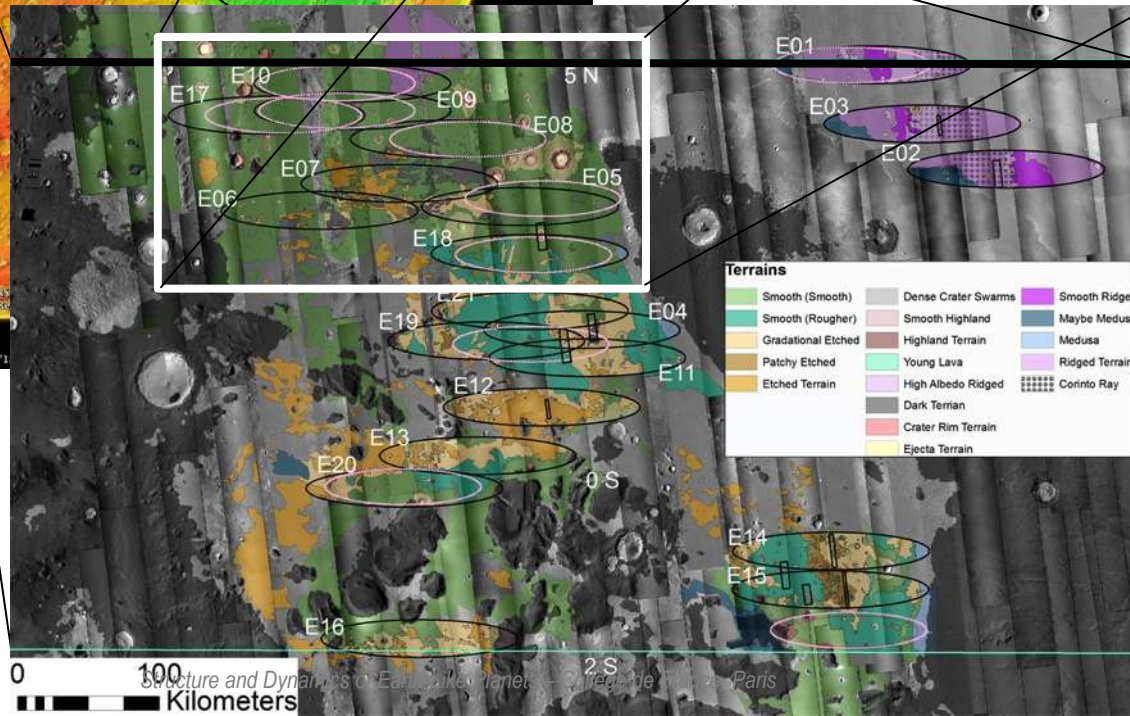
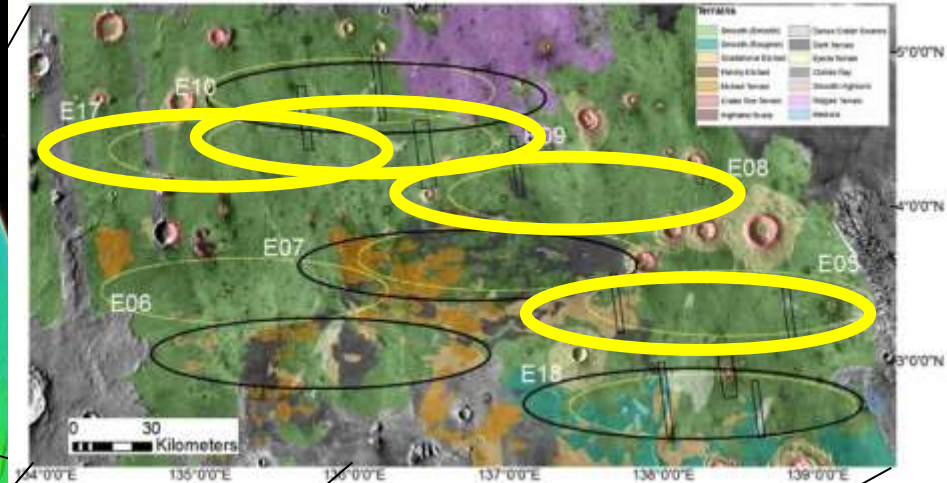
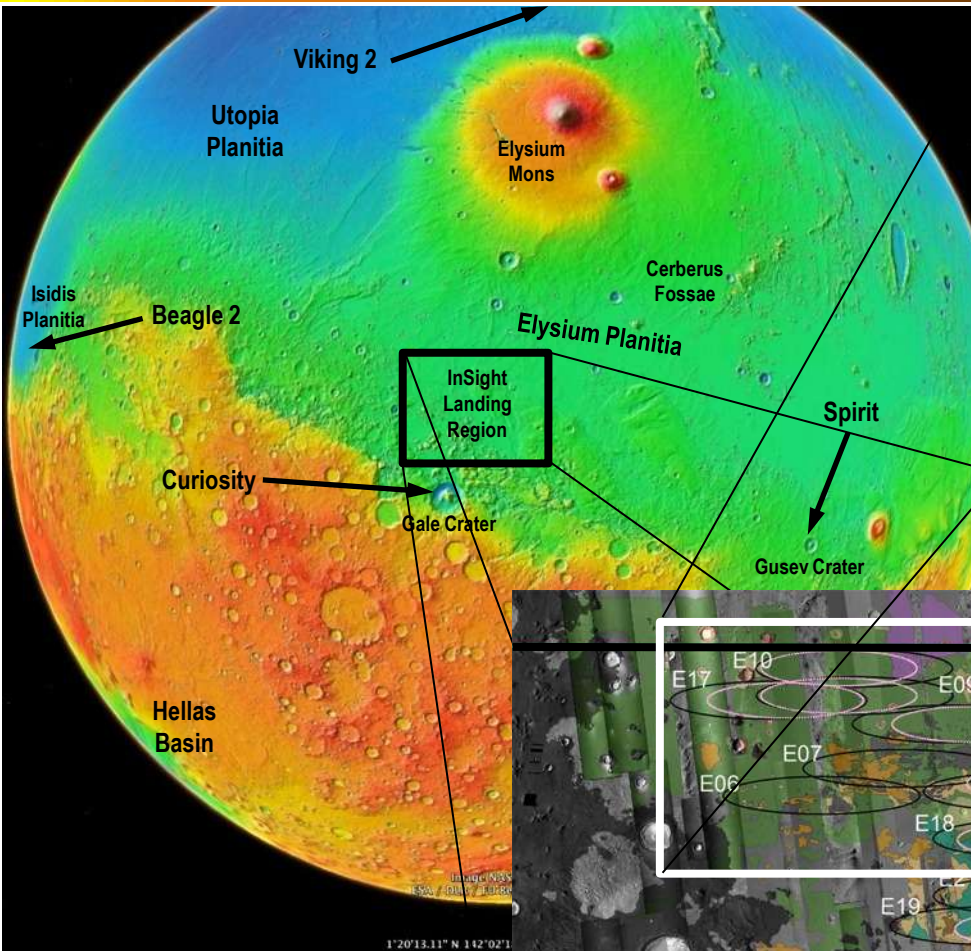
Pressure Sensor



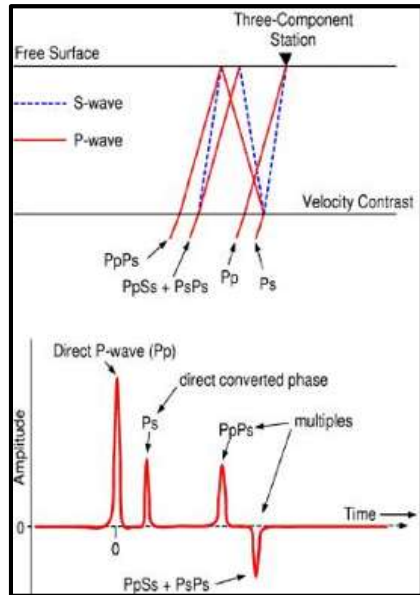
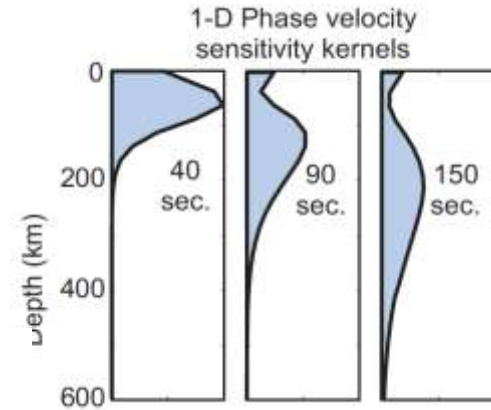
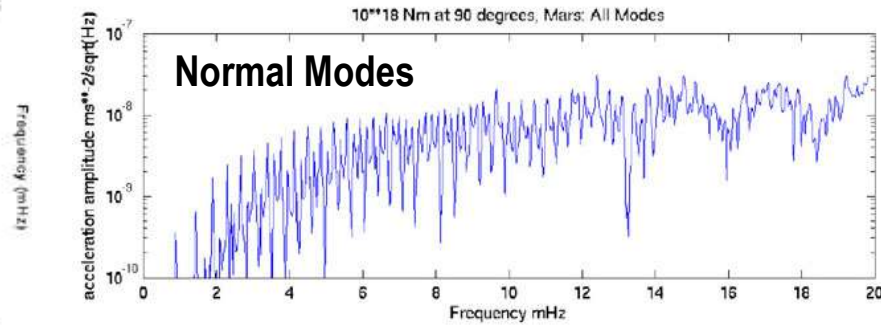
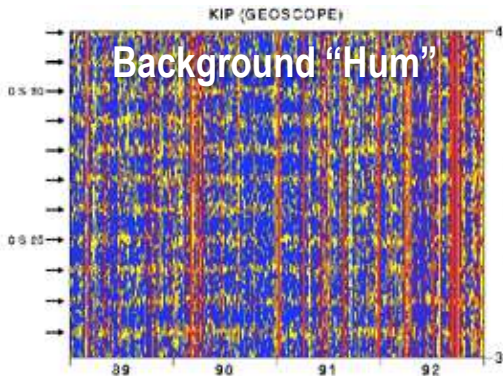
TWINS (CAB) – Temp. and Wind for InSight

APSS (JPL)

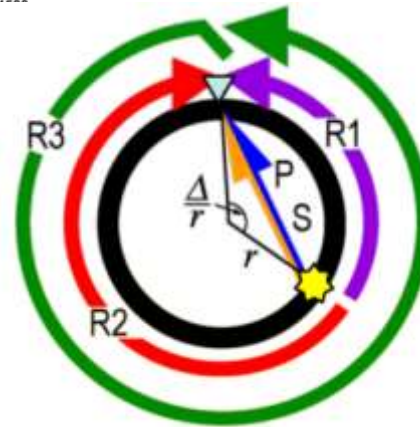
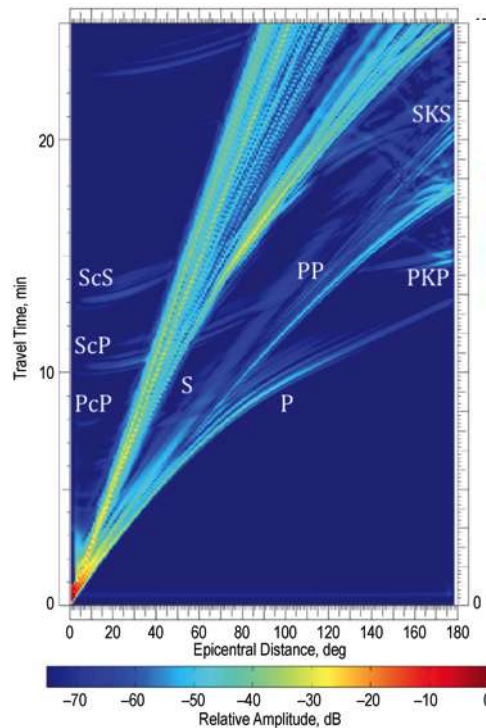
Auxiliary Payload Sensor Suite



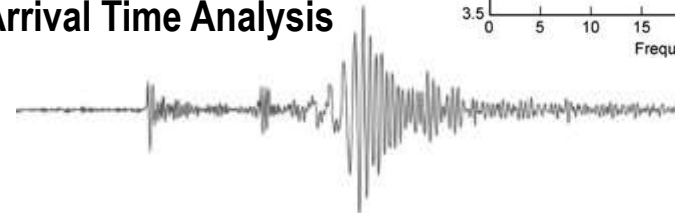
Ellipse	Lon.	Lat.
E05	138.2	3.3
E08	137.3	4.1
E09	136.0	4.4
E17	135.1	4.3



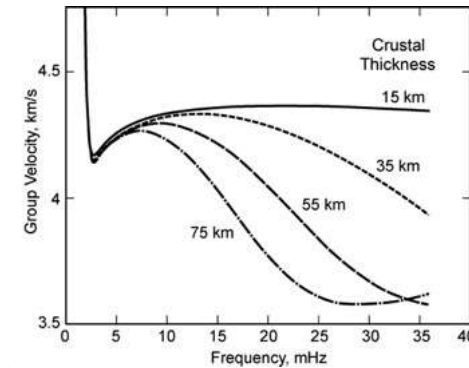
Receiver Function



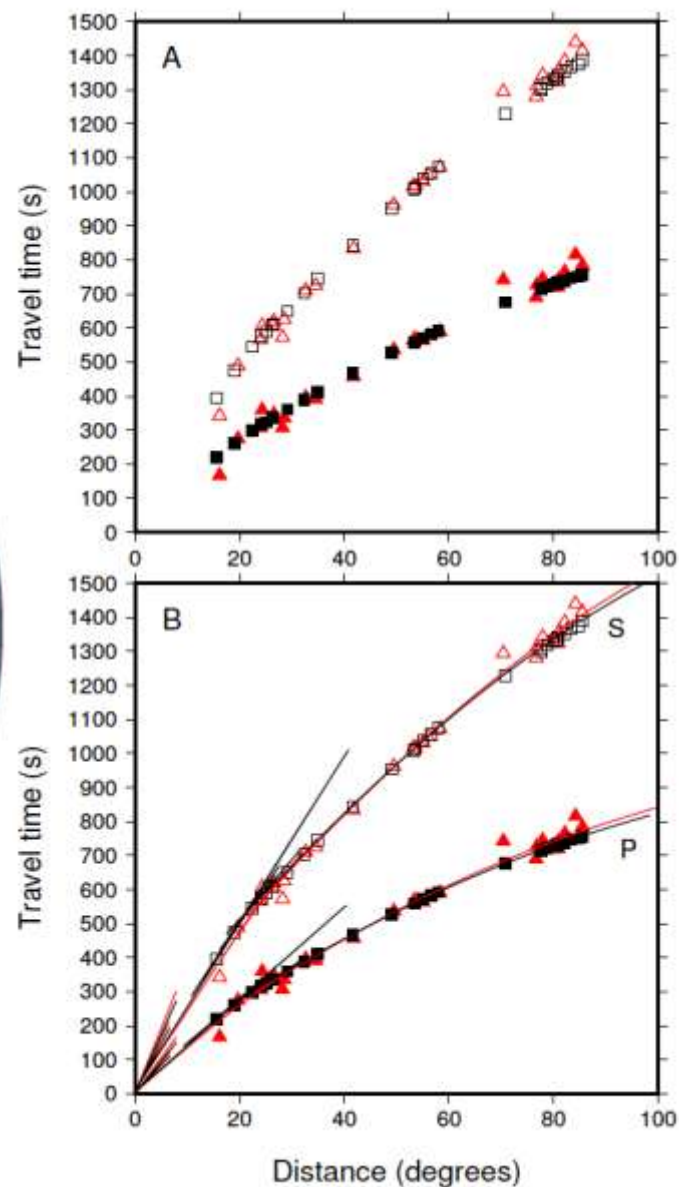
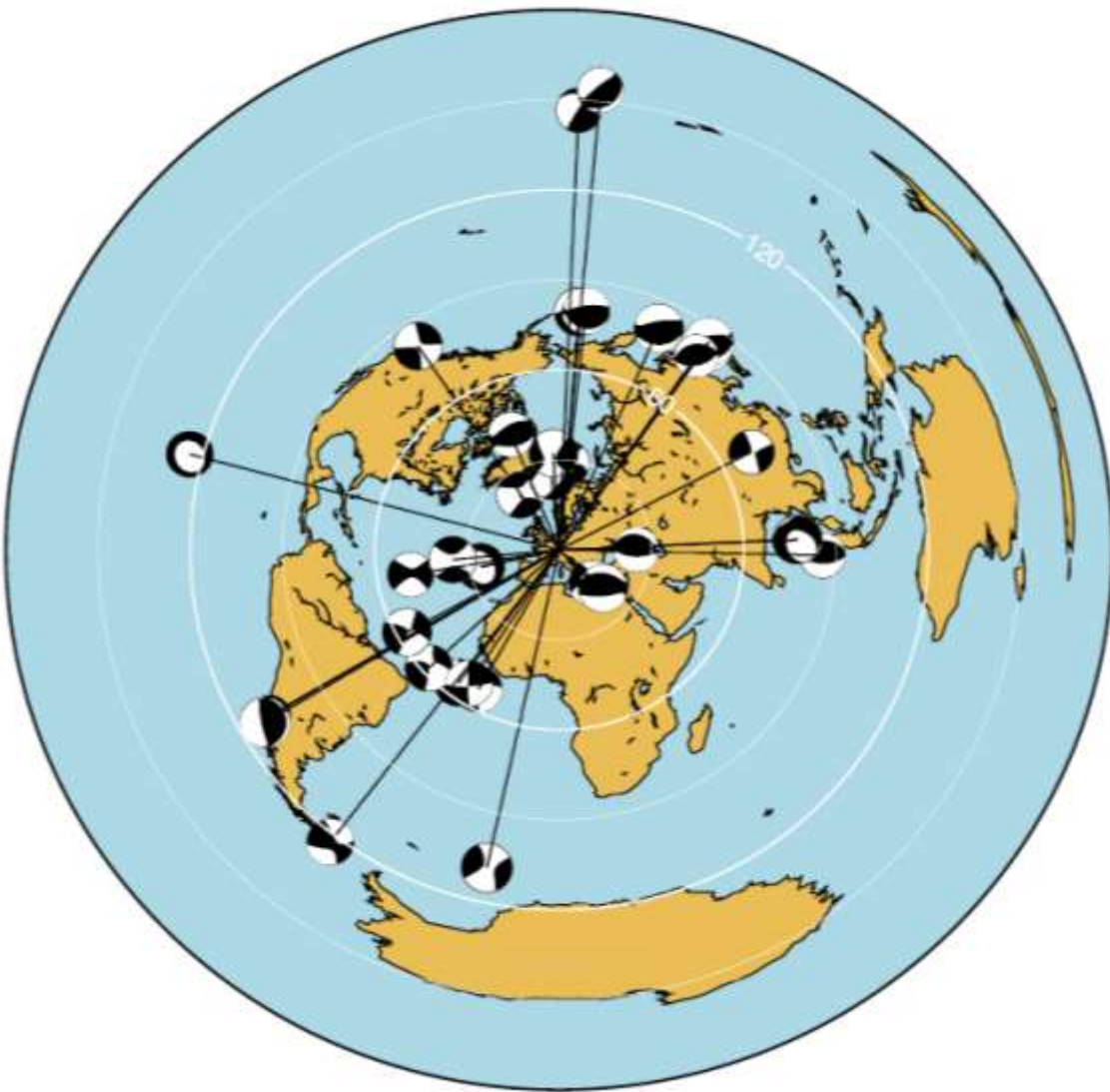
Arrival Time Analysis

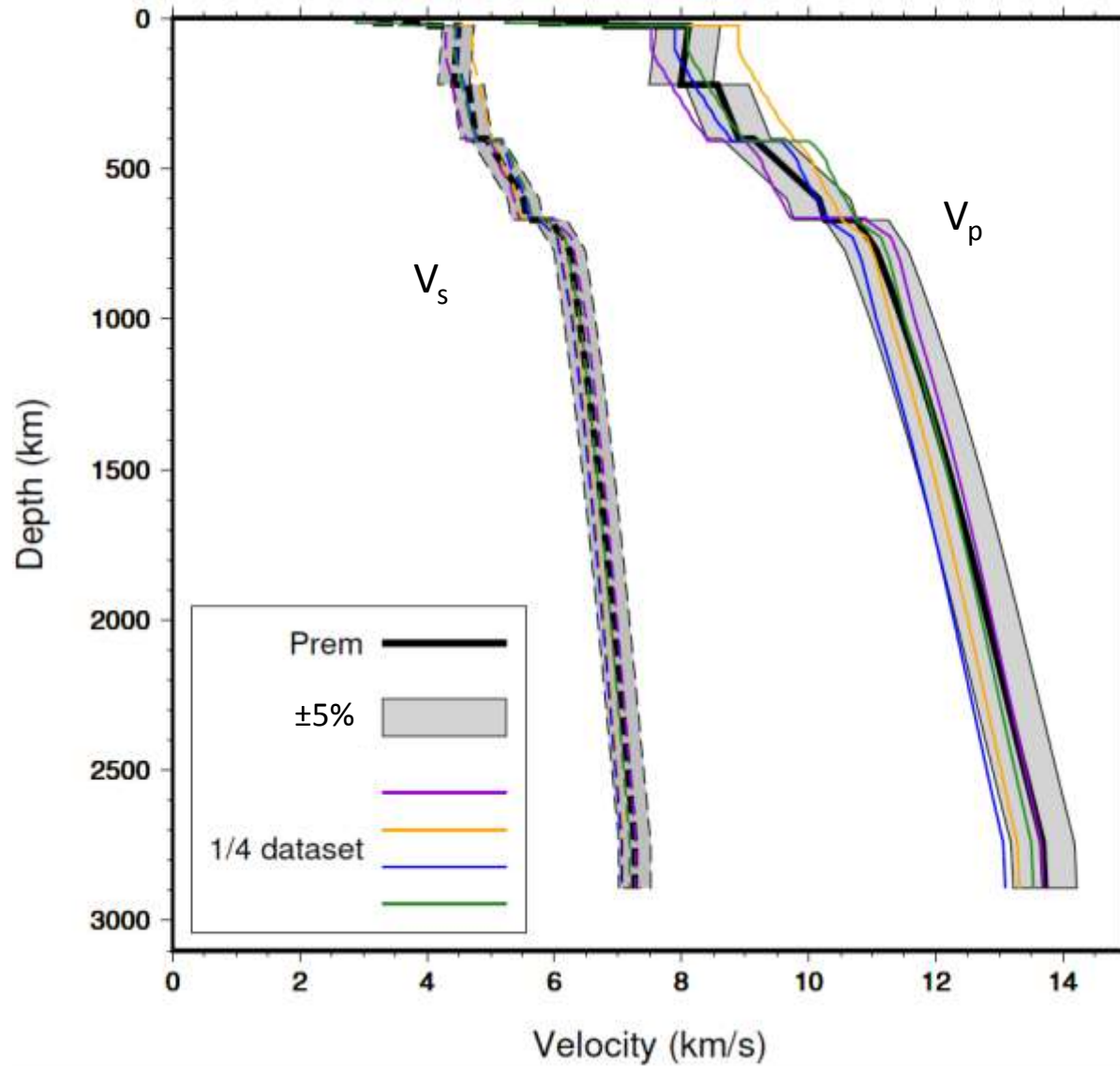


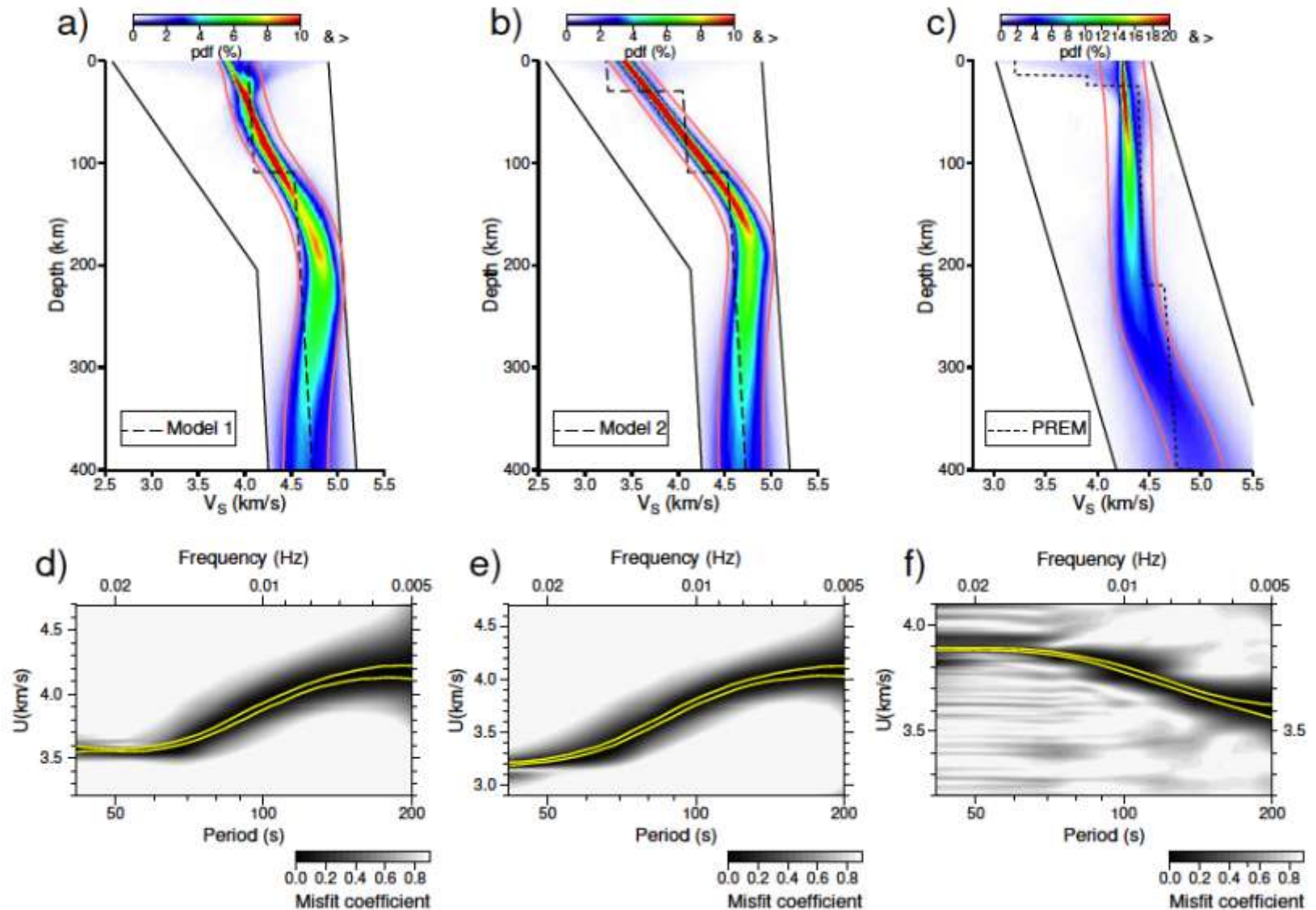
Surface Wave Dispersion



M. Panning et al., *Icarus*, in press

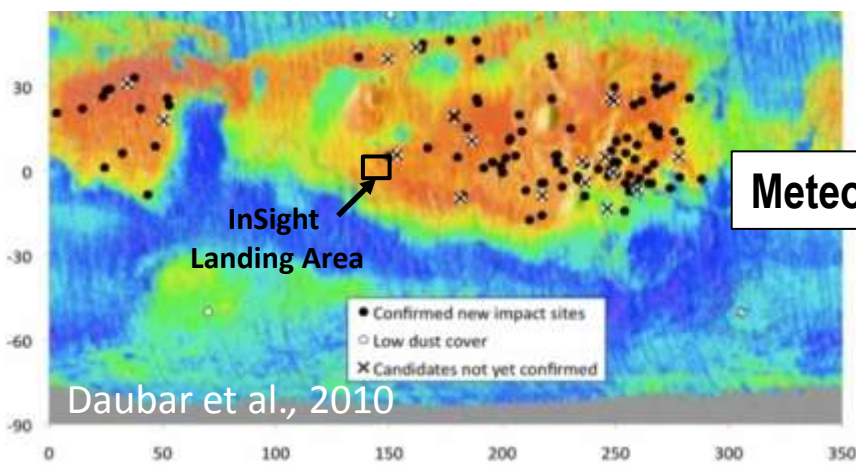
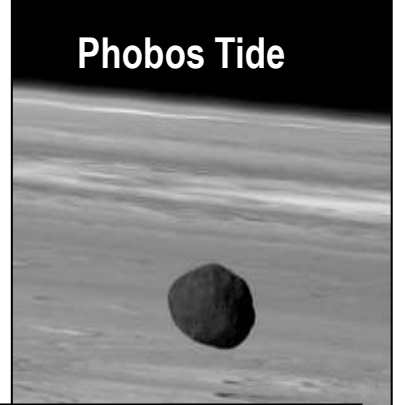
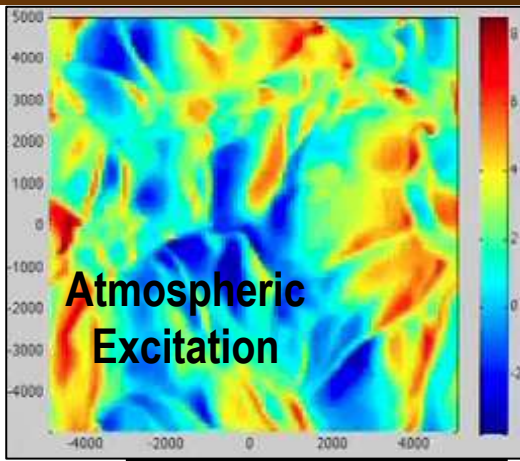
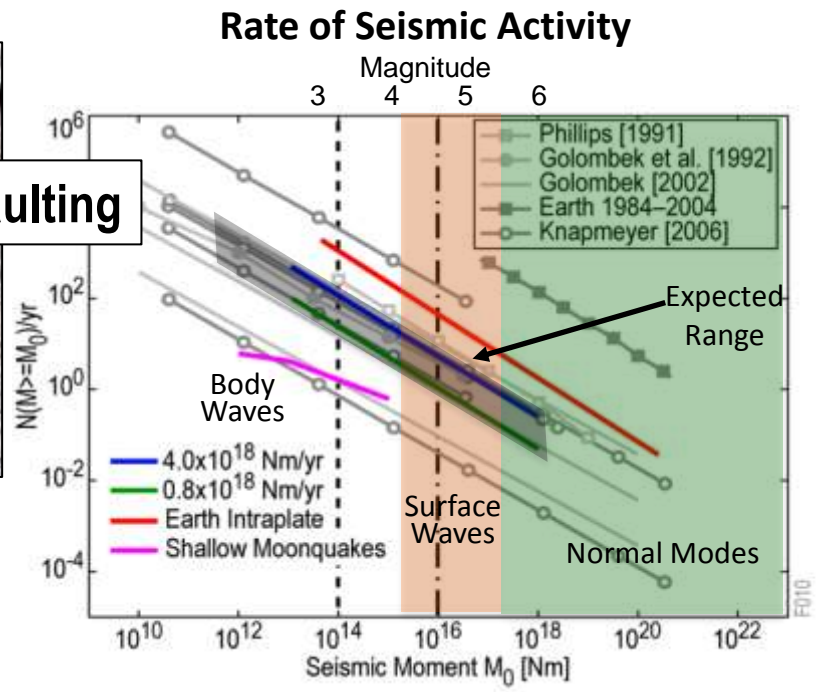








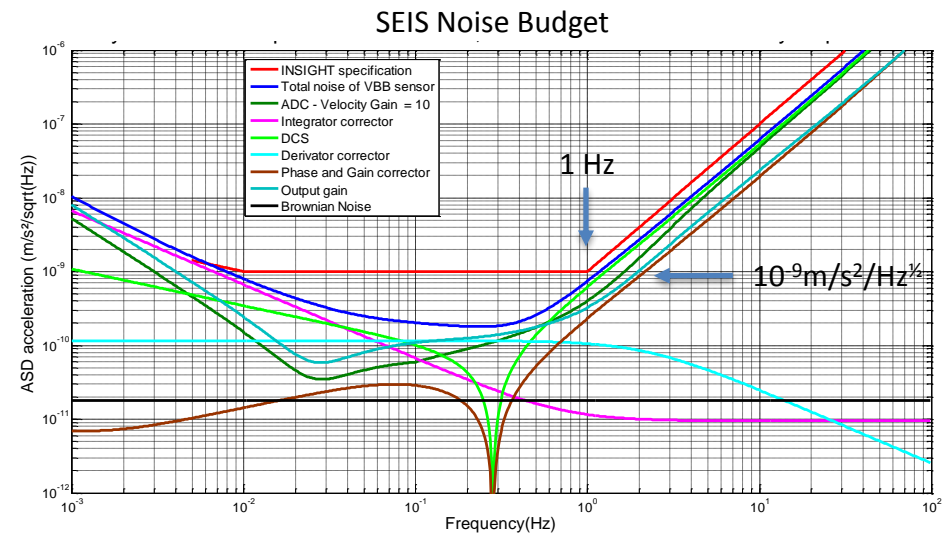
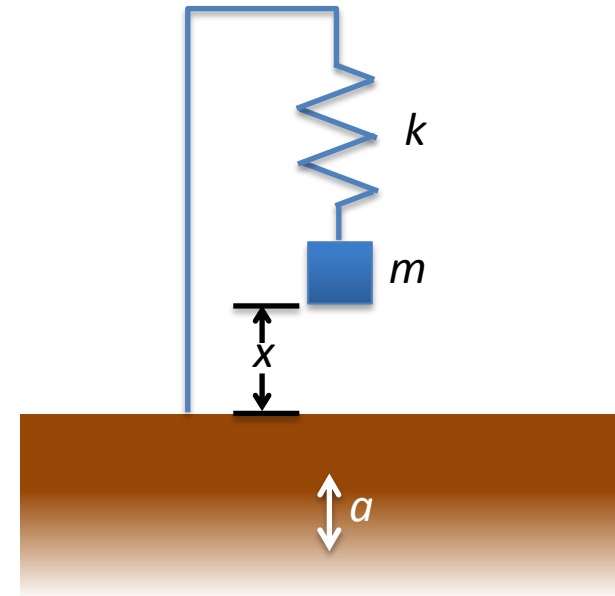
Faulting



- Acceleration noise req. over 1 Hz: $\leq 10^{-9} \text{ m/s}^2/\text{Hz}^{1/2}$
 - For oscillatory motion, $x = a/\omega^2 = a/4\pi^2 f^2$

⇒ SEIS is sensitive to displacements of $\sim 2.5 \times 10^{-11} \text{ m}$

This is about half the Bohr radius of a hydrogen atom

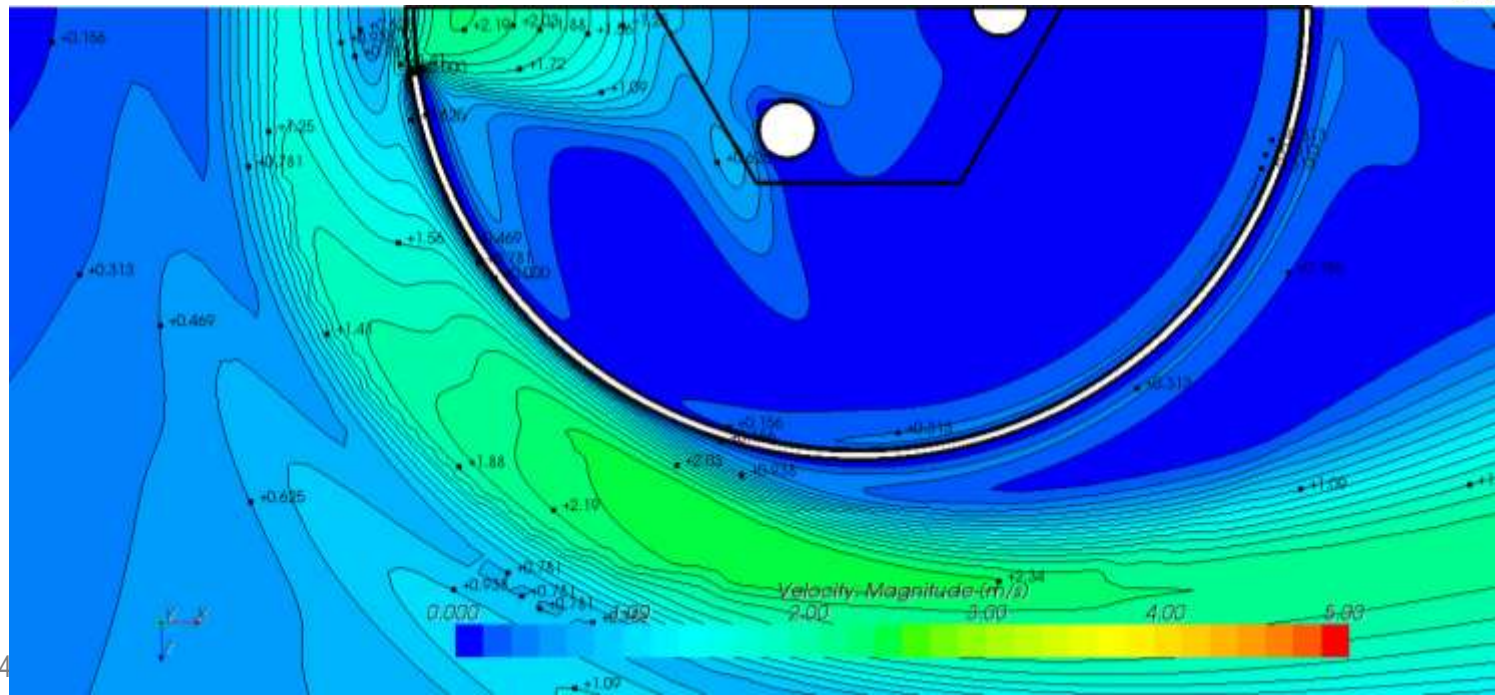
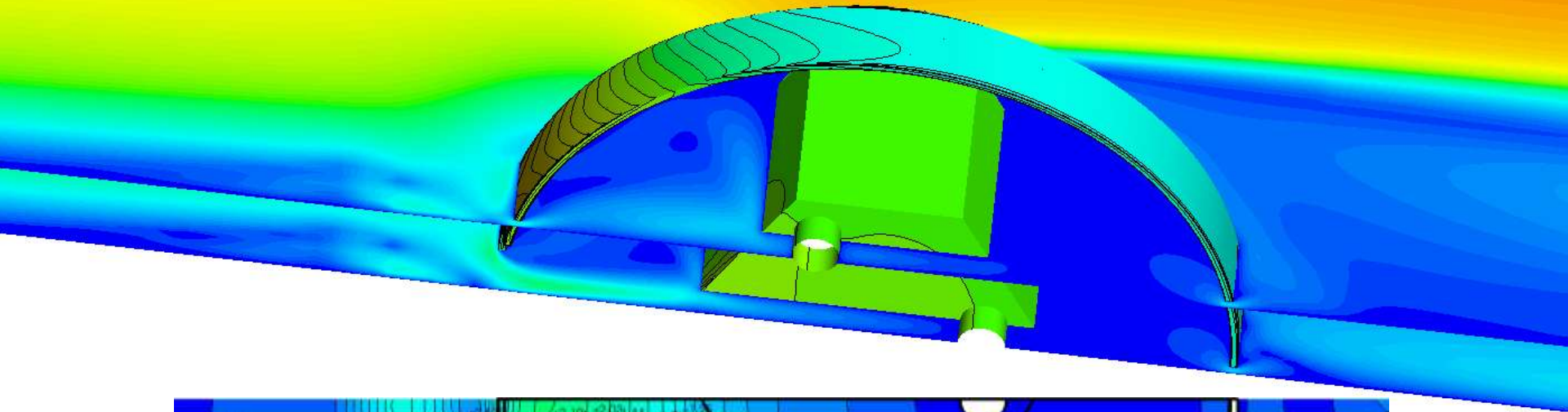


- **Environmental monitoring**

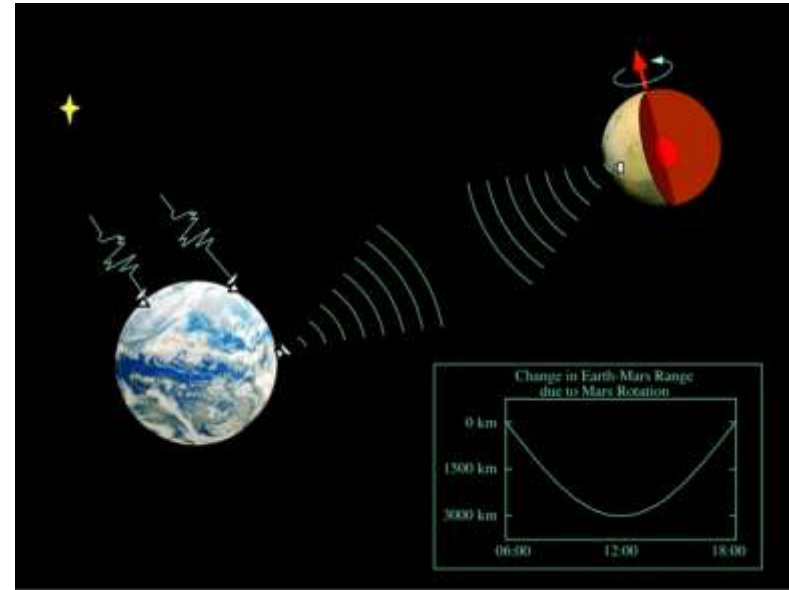
- A seismometer measures almost everything in its vicinity better than it measures seismic waves.
 - Could sense the flapping wings of a butterfly on top of the instrument!
- InSight will carry a sensor package to characterize the atmospheric and electromagnetic noise environment for SEIS
 - **Pressure** – 10 mPa barometer, DC-10 Hz
 - **TWINS** (Temperature and Wind for INSight) – Wind speed and direction, air temperature, twice per second
 - **Radiometer** – Ground temperature from IR emission
 - **IFG** (Insight FluxGate) – Magnetic field to 0.1 nT, DC-10 Hz



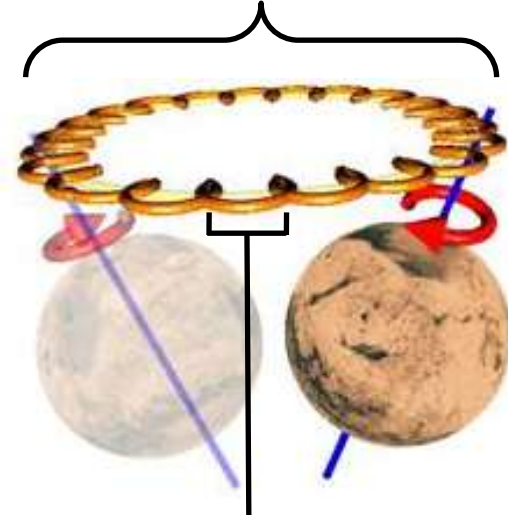
“Stealth” instruments



- First measured constraint on Mars core size came from combining radio Doppler measurements from Viking and Mars Pathfinder, which determined spin axis directions 20 years apart
 - Difference of spin axis direction gives precession rate and hence planet's moment of inertia
 - Constrains mean mantle density, core radius and density, although only 2 parameters are independent.
- InSight will provide another snapshot of the axis 20 years later still, providing stronger constraints on precession.
- In addition, nutation amplitudes can be determined after 2 years of tracking with sub-decameter precision
 - Free core nutation constrains core MOI directly, allowing separation of radius and density.



Precession ($T=165,000$ yr)



Nutation ($T \leq 1$ Mars yr)

Self-Hammering Mole, Thermal Sensors in 5-m Tether



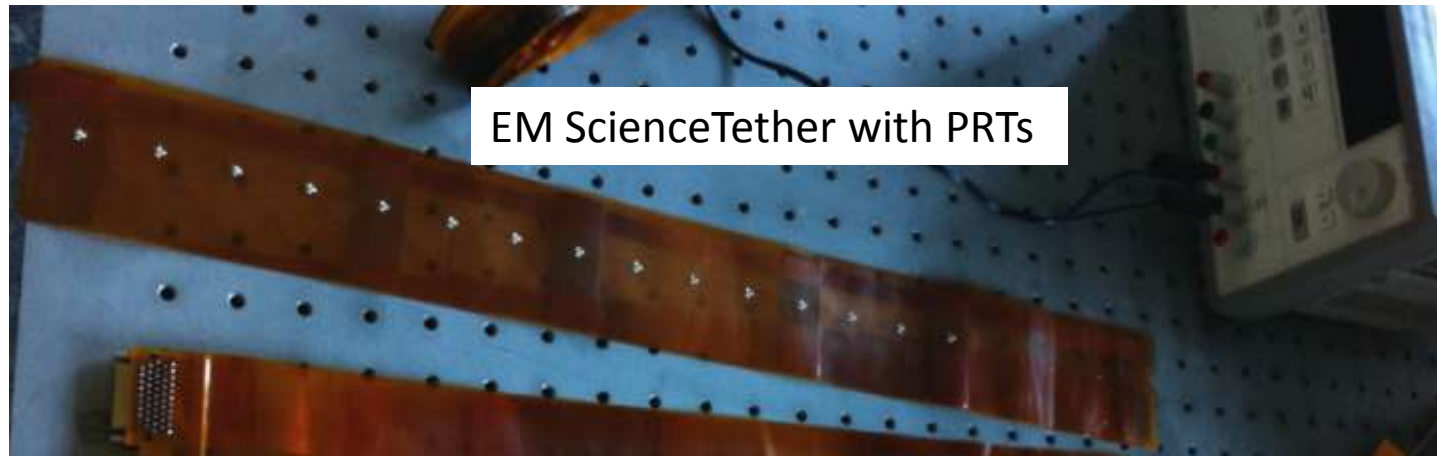
STATIL

Motor

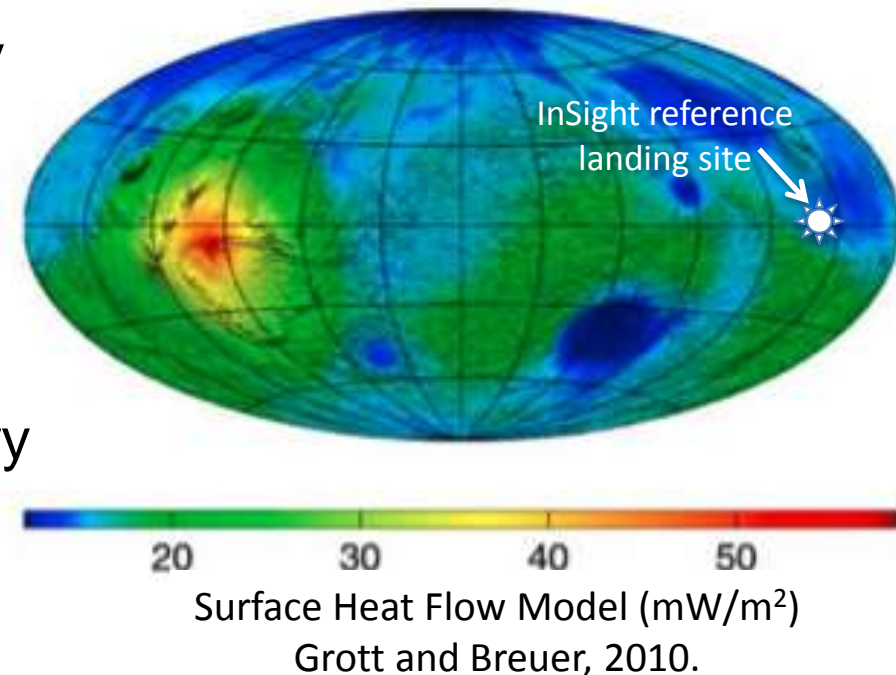
Hammer Mechanism



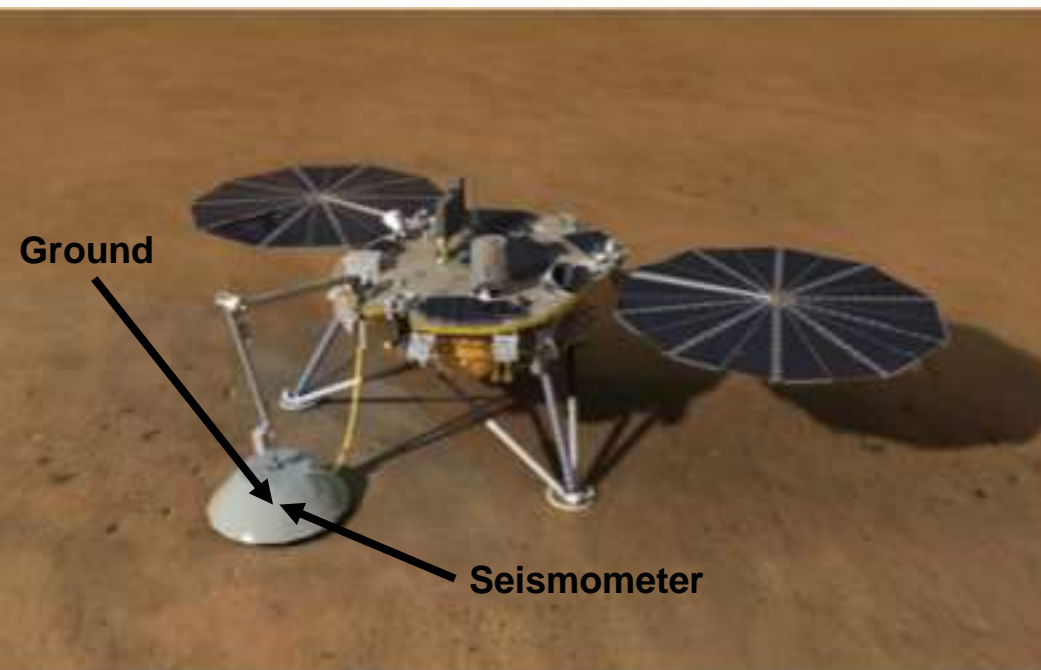
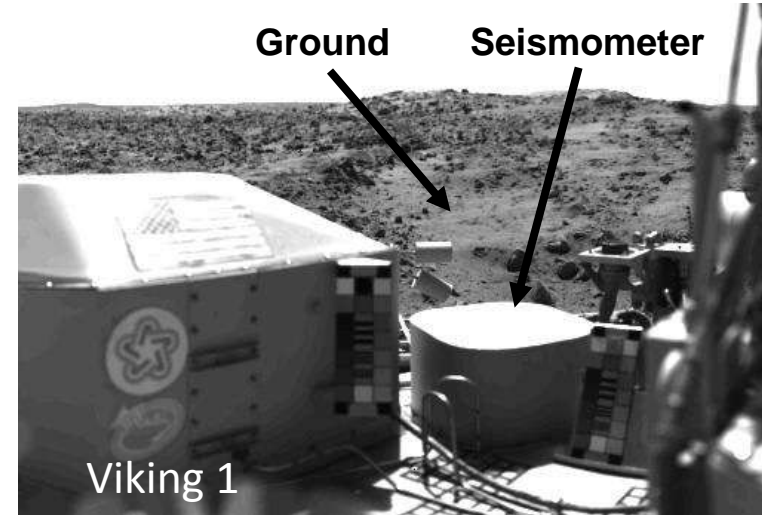
TEM-A: foils within Mole outer hull)



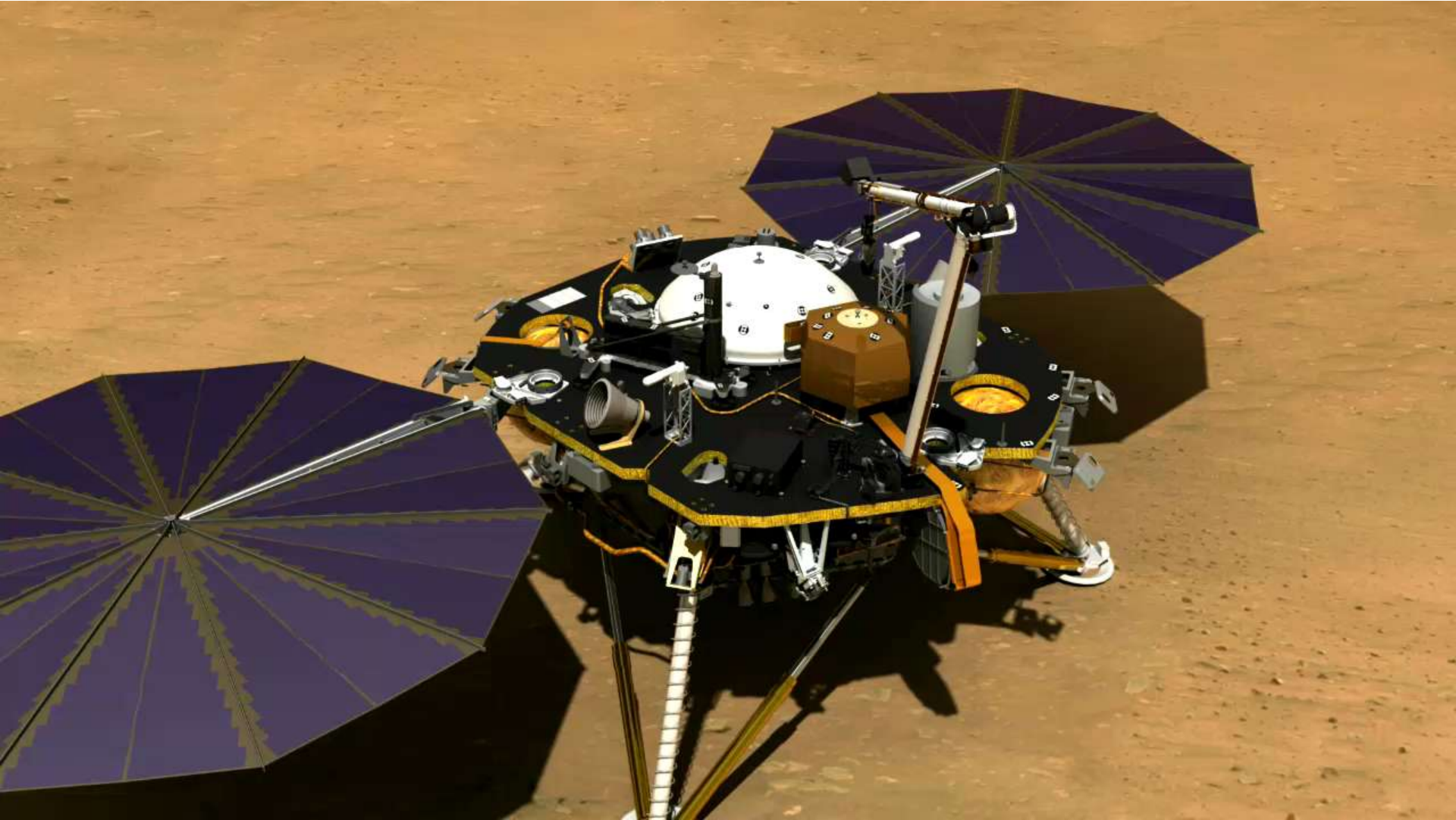
- HP³ (Heat Flow and Physical Properties Probe) has a self-penetrating “mole” that burrows down 3 to 5 meters below the surface.
 - It trails a tether containing precise temperature sensors every 35 cm to measure the temperature gradient of the subsurface.
 - It also contains a heater to supply a heat pulse for an active determination of thermal conductivity every 50 cm.
- Together, gradient and conductivity yield the rate of heat flowing from the interior.
- Present-day heat flow at a given location provides a critical boundary condition on models of planetary thermal history.



- The quality of a seismic station is directly related to the quality of its installation.
 - Installation couples the instrument to the ground and isolates it from the rest of the environment.
- But after landing, the instruments are still ~1 m from the ground...

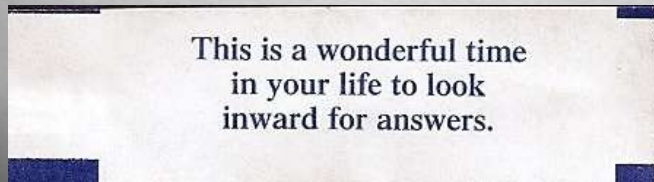


- InSight takes advantage of the payload mass capability of the Phoenix lander to fly an unprecedented deployment system.
- It will place the seismometer on the surface and cover it with an effective wind and thermal shield
- This will allow the seismometer sensitivity to reach the micro-seismic noise level of the planet





“Look deep into nature, and then you will understand everything better.” – Albert Einstein



– Chinese Fortune Cookie

Gusev Crater