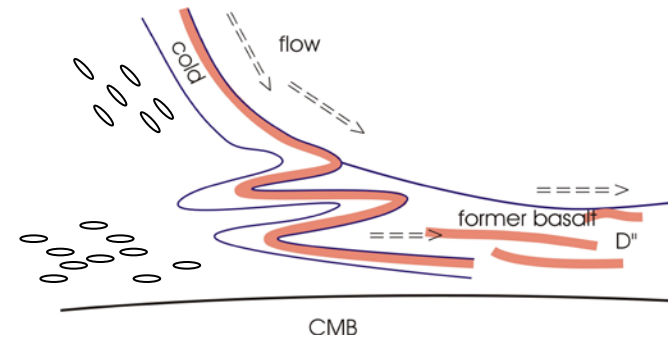
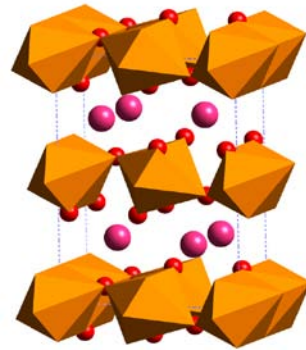
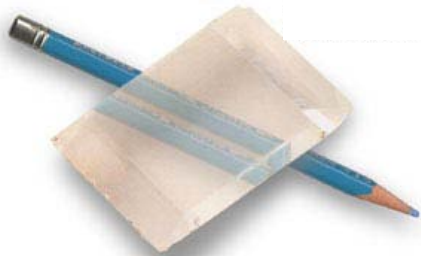


# Deformation in the lowermost mantle: linking seismic anisotropy, mineral physics and geodynamics



Michael Kendall  
University of Bristol

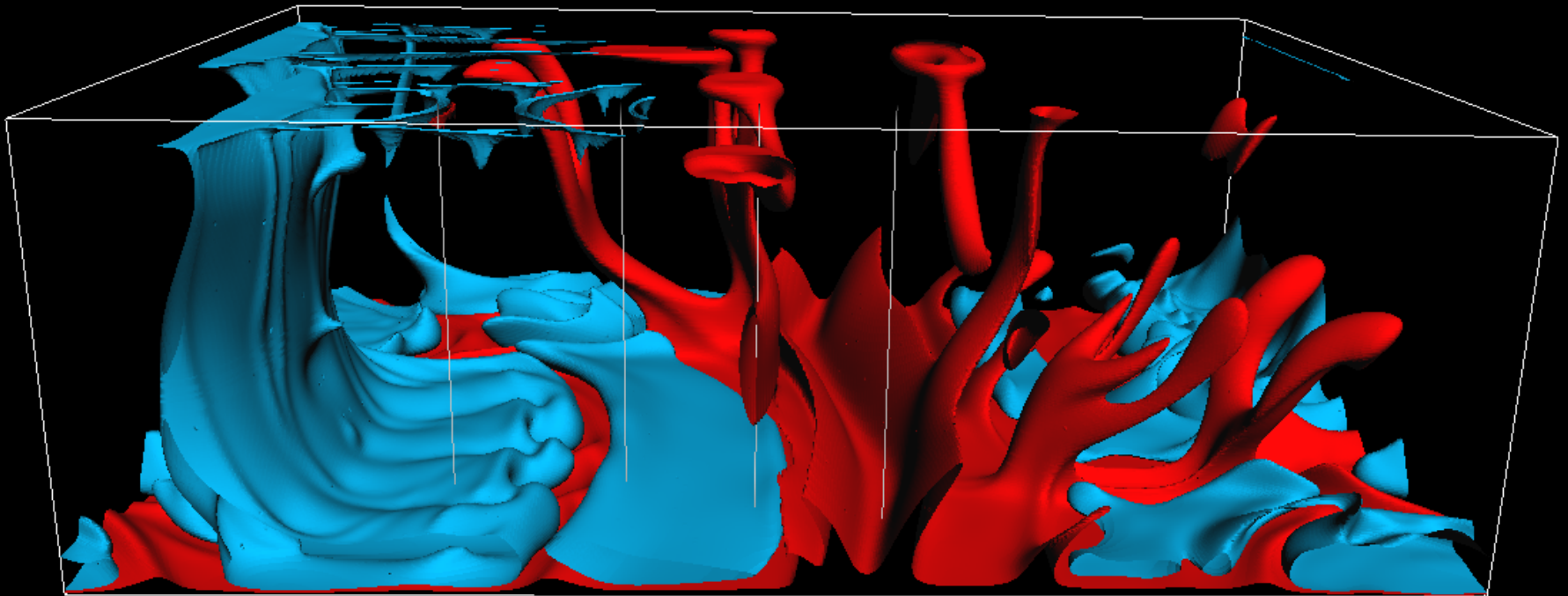
# Acknowledgements

- James Wookey - Bristol
- Andrew Walker - Bristol
- Andy Nowacki – Bristol
- Jack Walpole - Bristol
- Alessandro Forte – UQAM, Canada

COMITAC – ERC funding (PI - James Wookey)

# D'' and the dynamics of the core and mantle

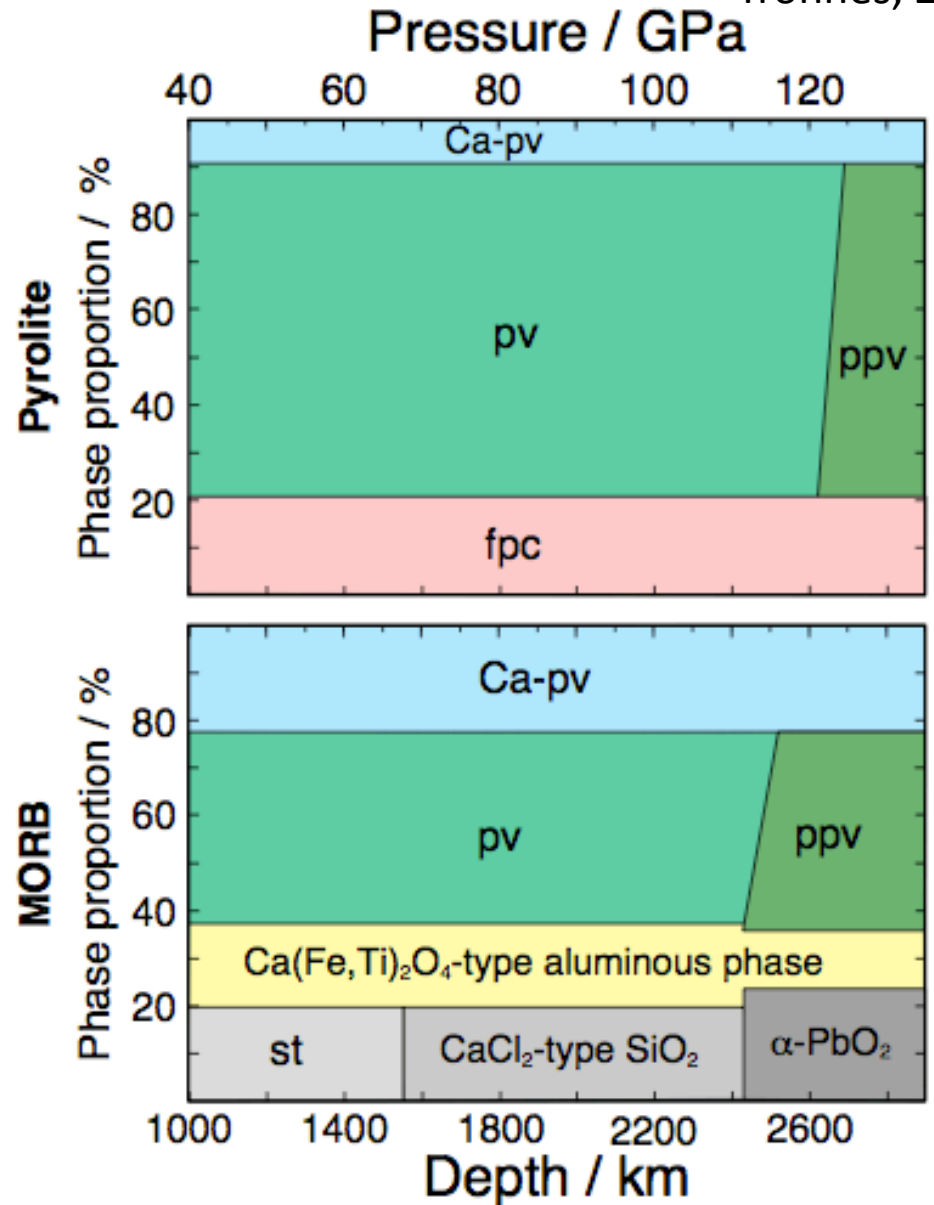
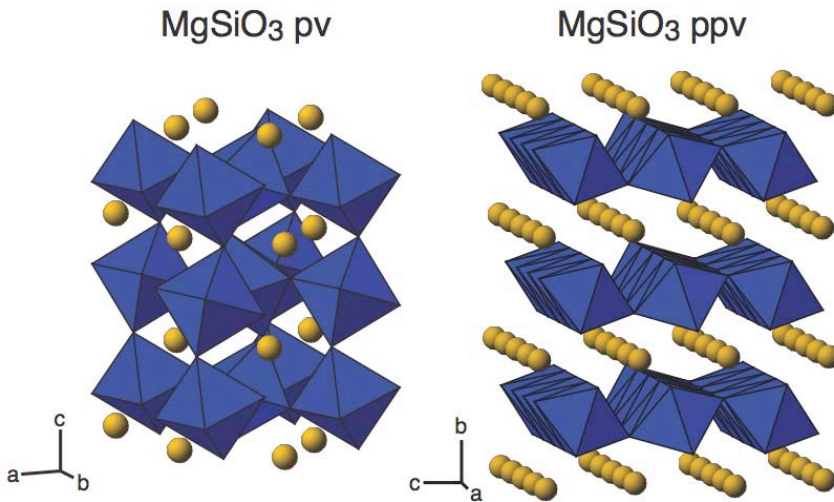
- Thermal boundary layer
- Site of plumes
- Birth of LLSVPs
- Resting place for slabs
- Boundary condition for core dynamics



# Mineralogy of the lower mantle

Tronnes, 2010

- Dominant mineral is perovskite
- Phase transition to post-perovskite



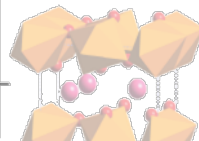
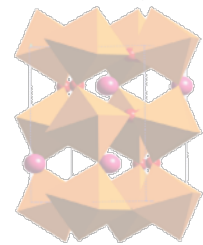
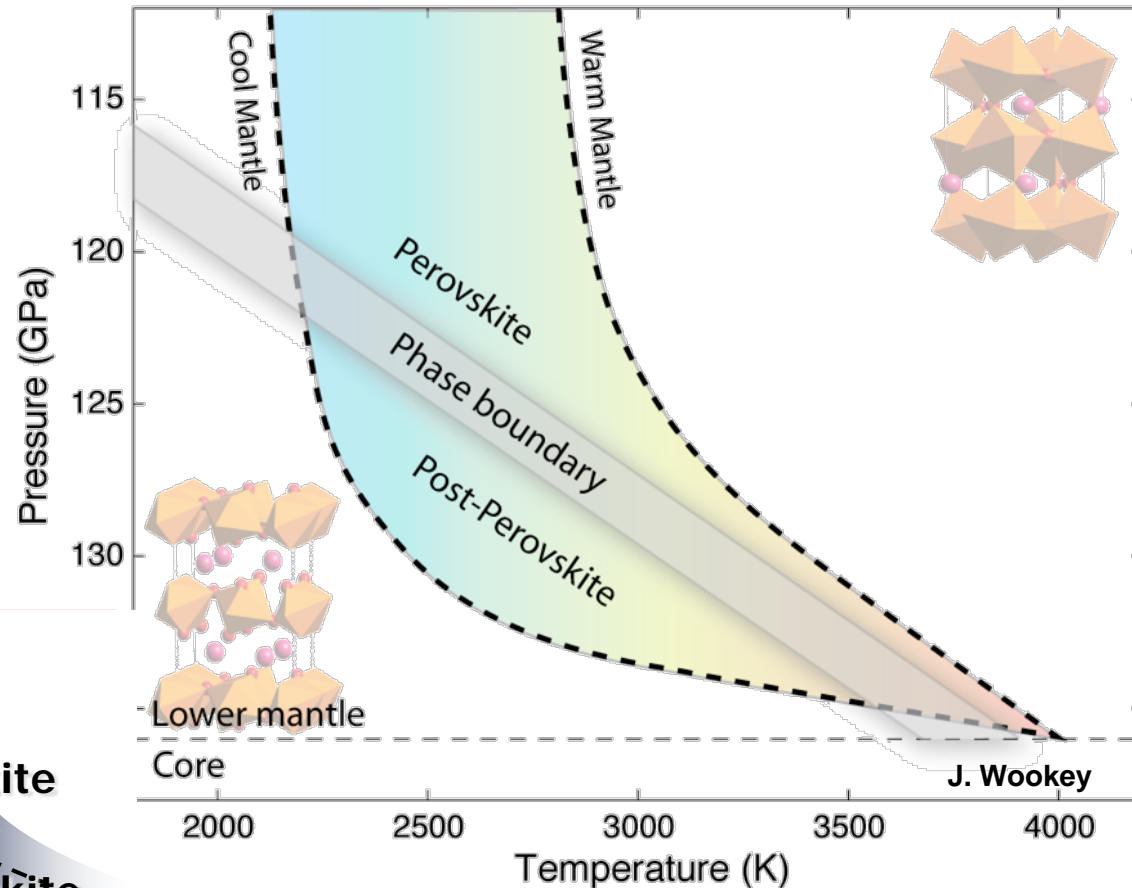
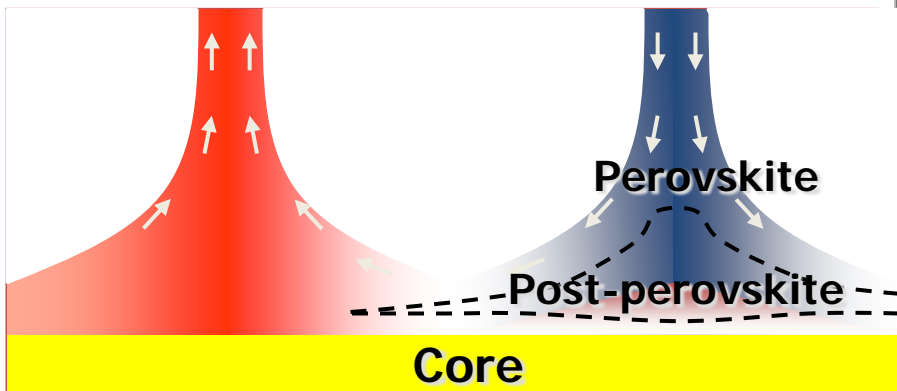
# Perovskite to Post-perovskite

## Ab initio modelling of PPV

(Oganov and Ono, 2004; Stackhouse et al, 2005; Wookey et al, 2005b; Tsuchya et al, 2004; 2005; Wentzcovich et al, 2006 ...)

- Simulate PPV at a range of near-CMB P-T
- Observe strong (exothermic) temperature dependence in phase-boundary depth

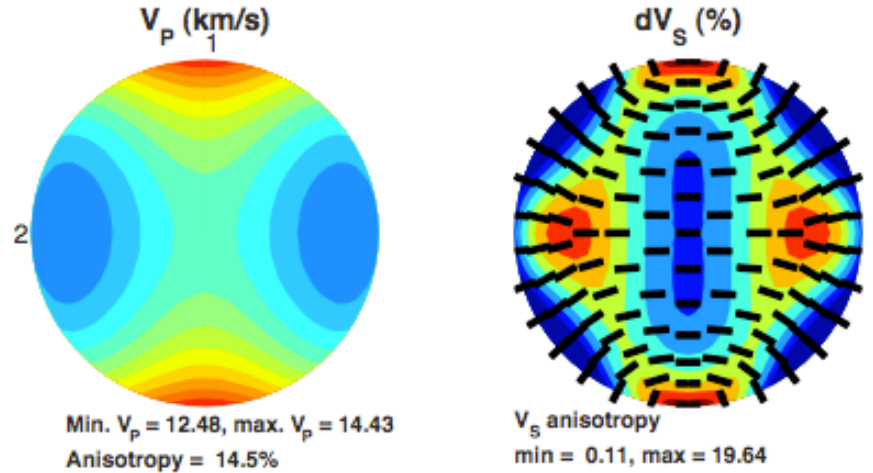
A lower mantle temperature probe.



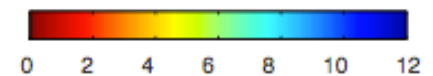
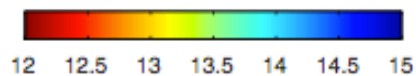
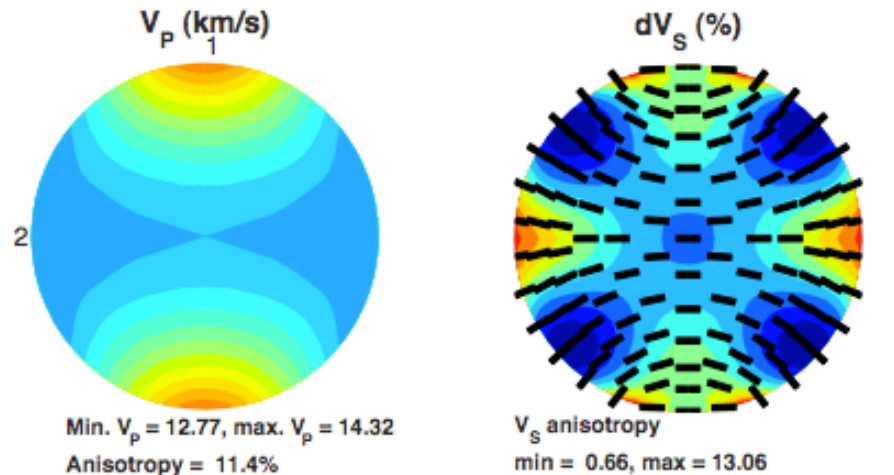
# The elasticity of the lower mantle

- Perovskite
- Orthorhombic symmetry
- Moderately anisotropic comparable magnitude and symmetry between authors.

MgSiO<sub>3</sub>, Wookey et al., 2005, P=126 GPa, T=2800 K



MgSiO<sub>3</sub>, Wentzcovitch et al., 2006, P=125 GPa, T=2500 K

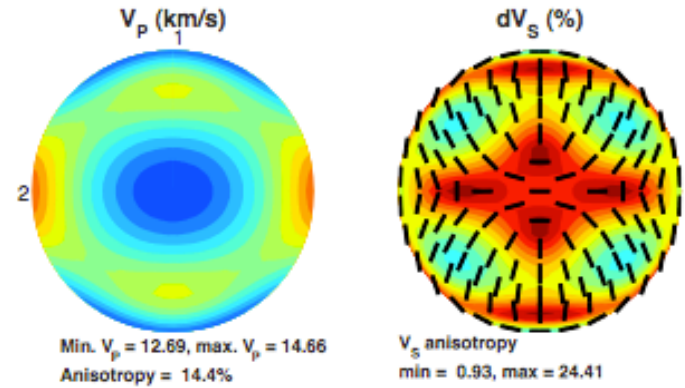




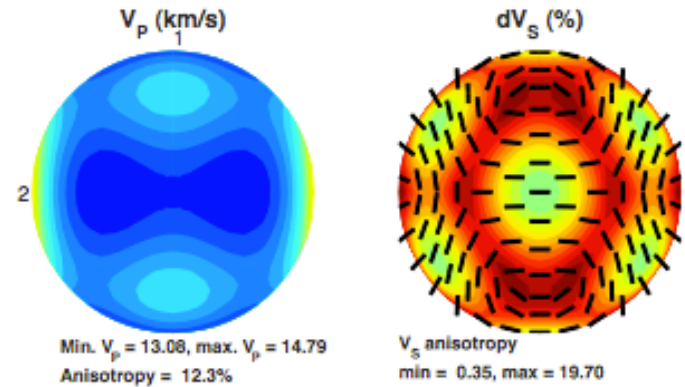
# The elasticity of the lower mantle

- Postperovskite
- Orthorhombic symmetry
- Moderate to high amounts of anisotropy; more disagreement between authors.

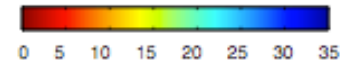
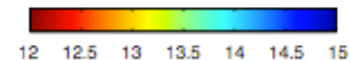
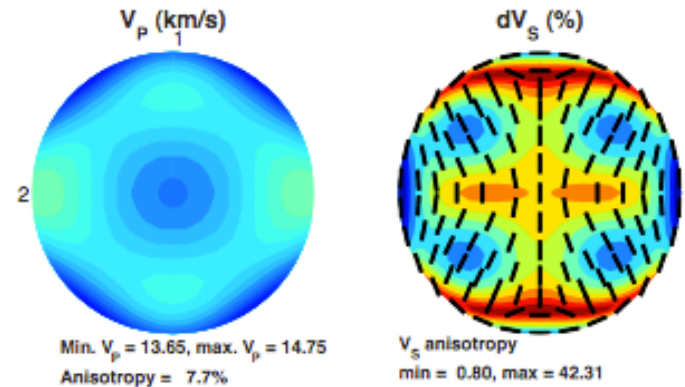
MgSiO<sub>3</sub>, Stackhouse et al., 2005, P=135 GPa, T=4000 K



MgSiO<sub>3</sub>, Wentzcovitch et al., 2006, P=140 GPa, T=4000 K



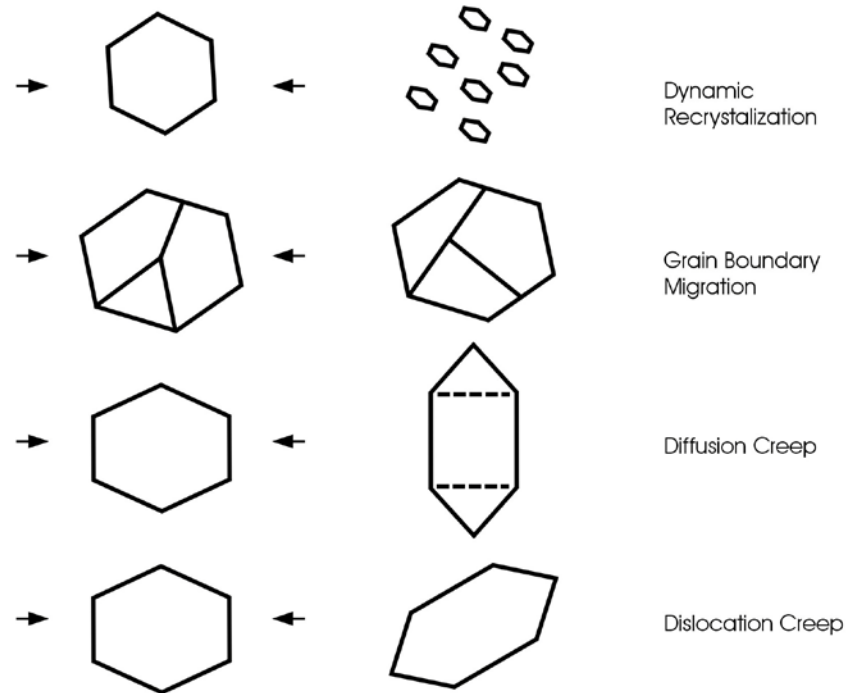
(Mg<sub>0.6</sub>Fe<sub>0.4</sub>)SiO<sub>3</sub>, Mao et al., 2010, P=140 GPa, T=2000 K



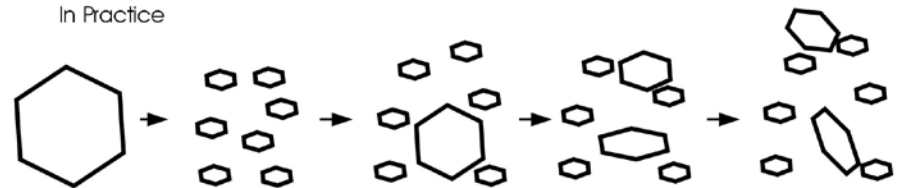
# Lattice preferred orientation

- Deformation accommodated by many mechanisms.
- Dislocation more effective than diffusion in generating anisotropy.
- Temperature, grain size, strain rate, strain history, pressure, fluids
- Polymineralic effects?
- Inherited textures?
- Need to know slip systems.

## DEFORMATION MECHANISMS



## In Practice

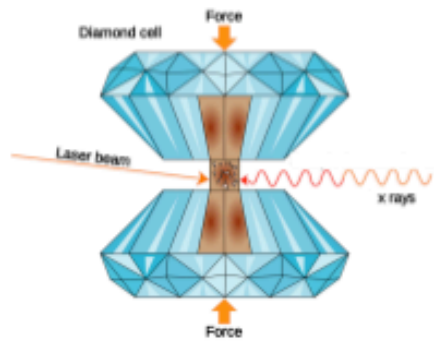


Controlling Factors:

temperature, grain size, stress-strain rate, pressure, fluids.

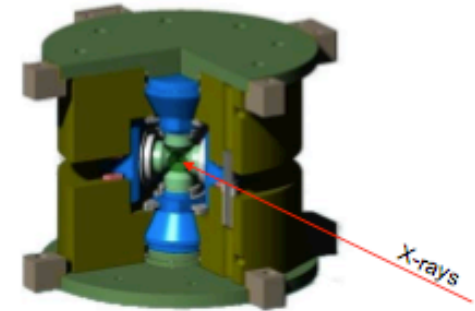
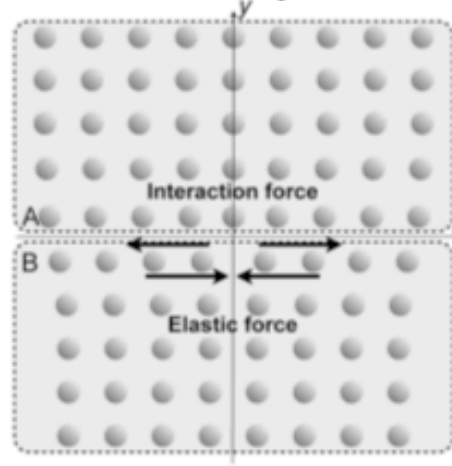


# Postperovskite slip systems



In situ deformation  
in the diamond anvil  
cell (to CMB  
pressure)

Peierls-Nabarro  
modelling using  
density functional  
theory



Large volume  
deformation  
experiments on  
analogues  
(elevated P & T)

$\text{MgSiO}_3$  (100)  
or (001)

(010)

-

$\text{CaIrO}_3$  (010)

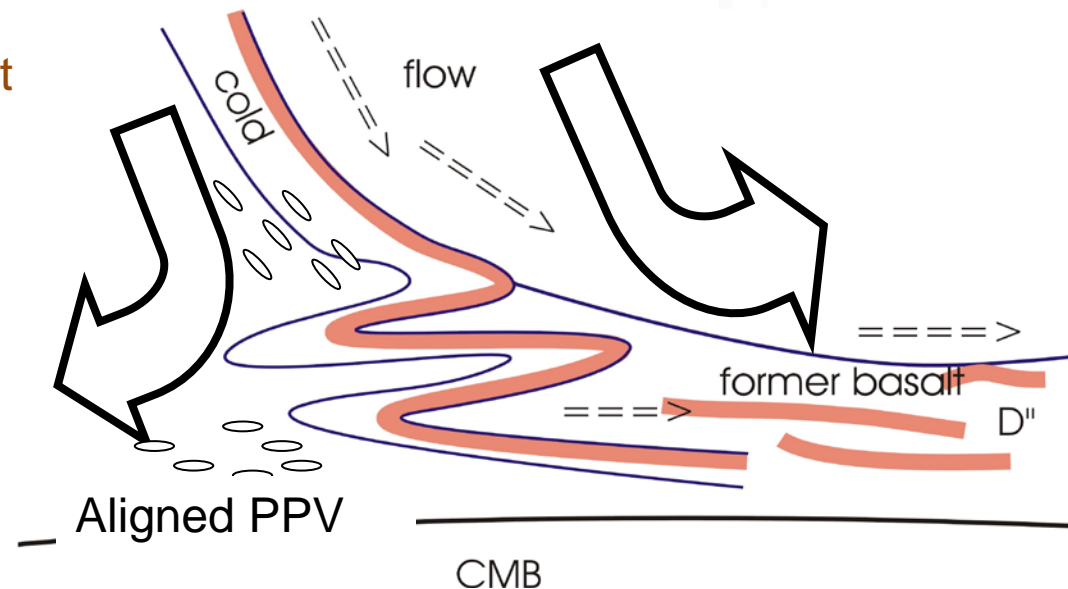
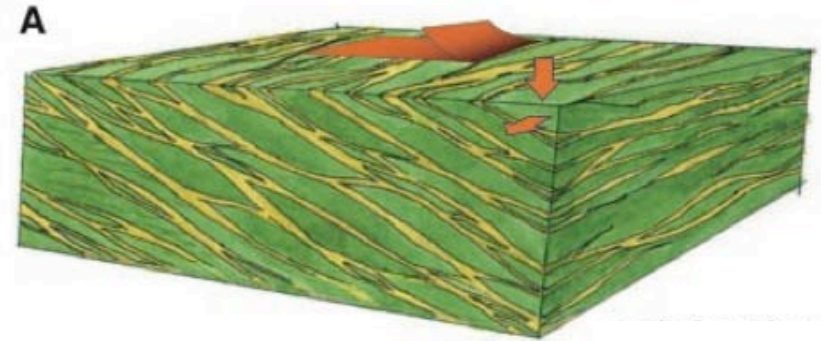
(010)

(010)

# Other anisotropy mechanisms

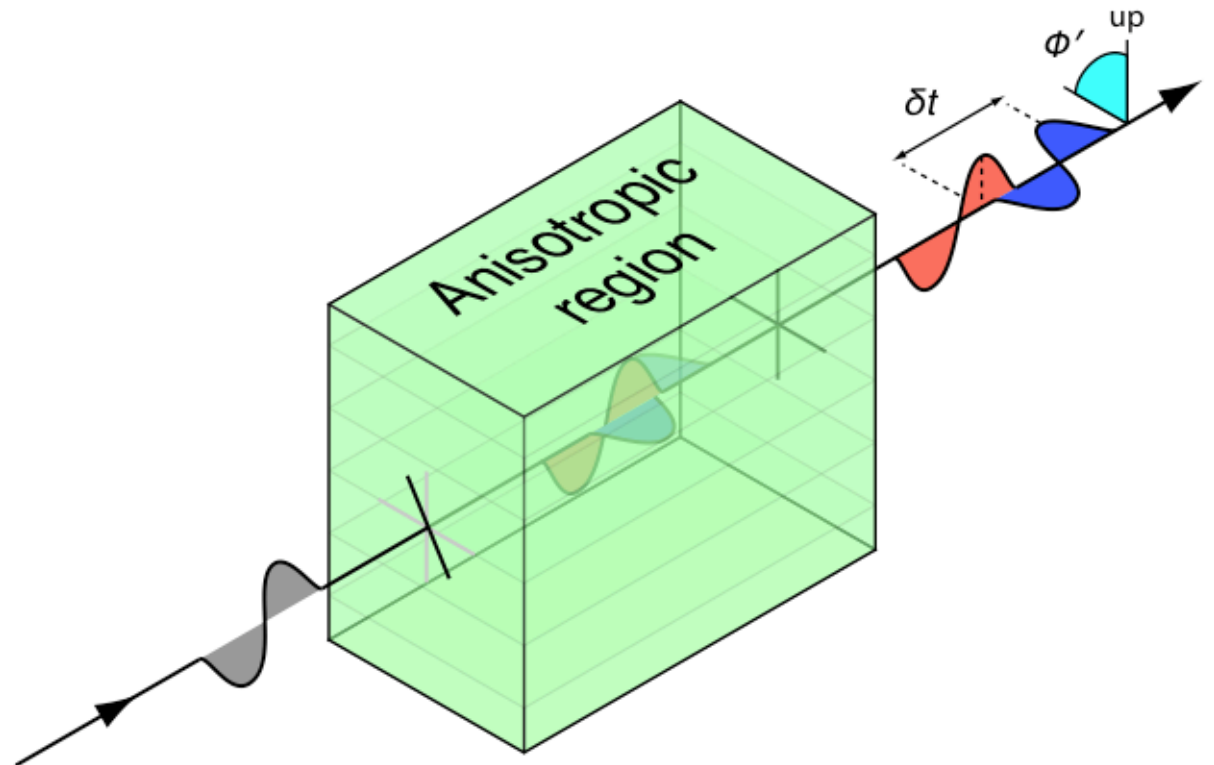
- **Shape-preferred orientation (SPO)**

- Layering of contrasting materials (Backus 1962)
- Alignment of inclusions (e.g., Kendall and Silver, 1996)
- Candidate inclusion materials include basaltic melt (Hirose et al, 1999), or intrusions of core material
- Combination of SPO and LPO (e.g., Holtzman and Kendall, 2010)

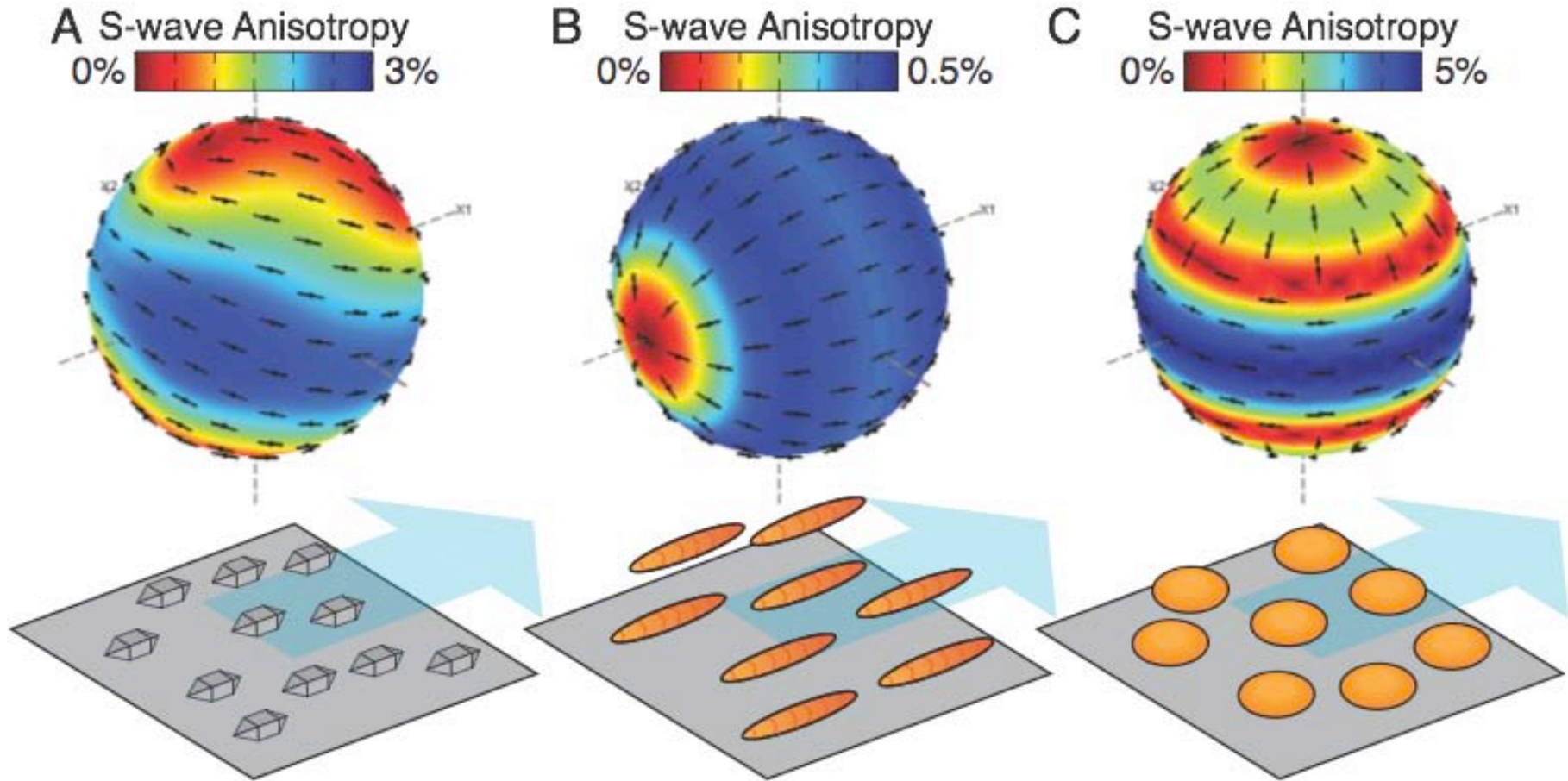


# A seismic probe of lower-mantle anisotropy

- Shear-wave splitting

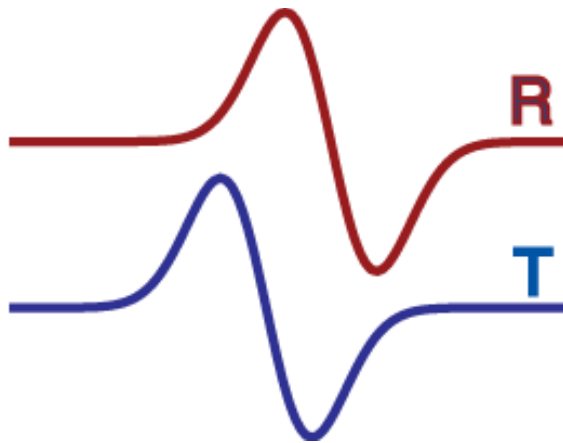
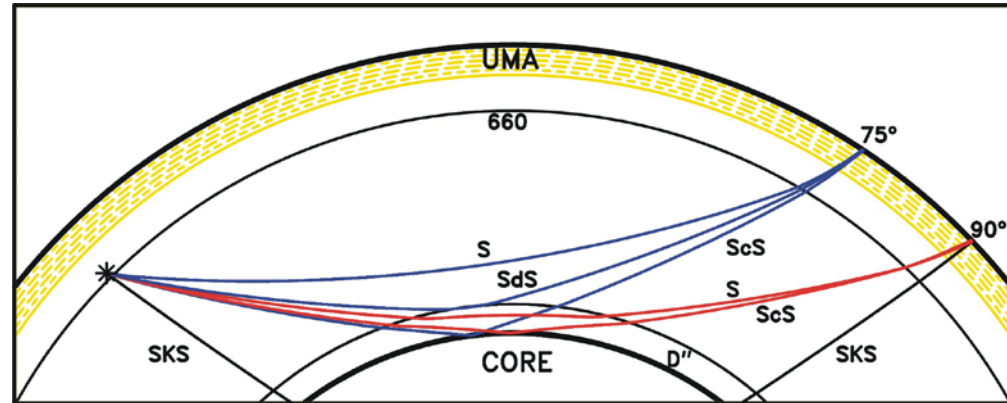
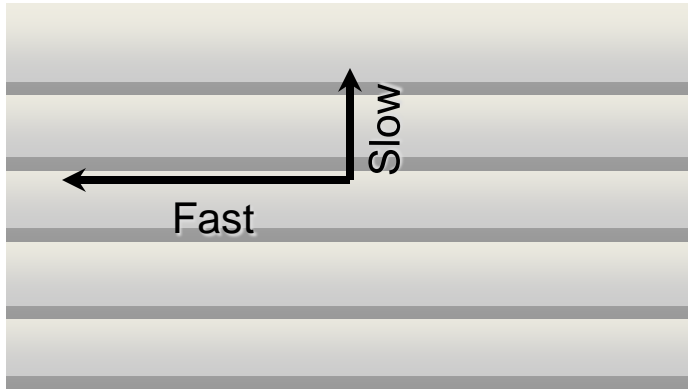


# Lower mantle anisotropy

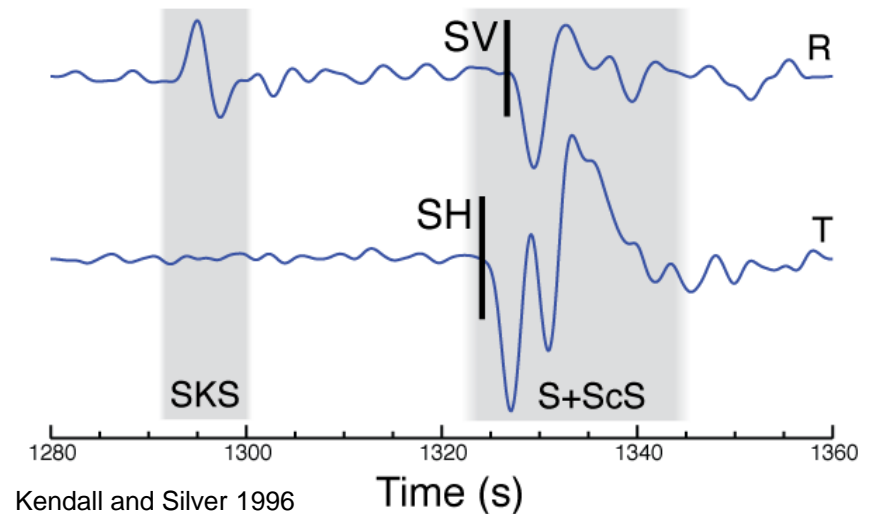


# VTI anisotropy

- Simple example is layering ...



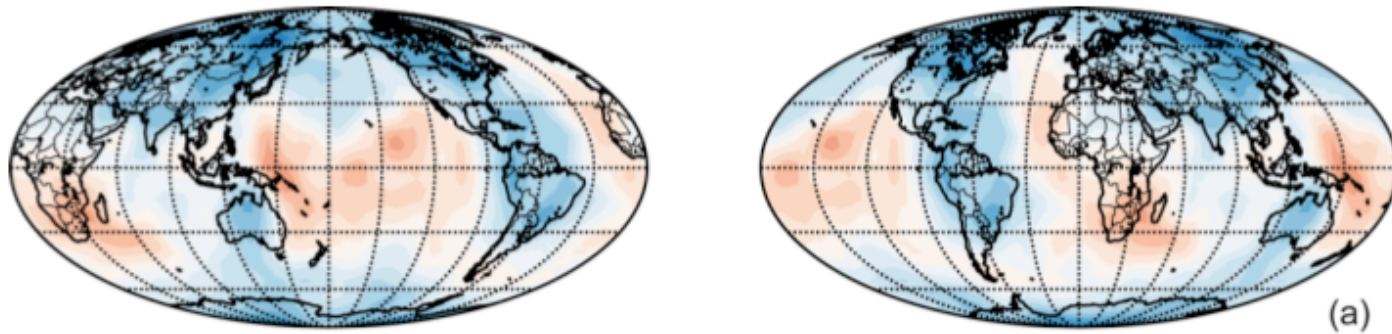
**B** CNSN Station FCC (Manitoba, Canada)  
Distance = 90°, corrected for UM anisotropy



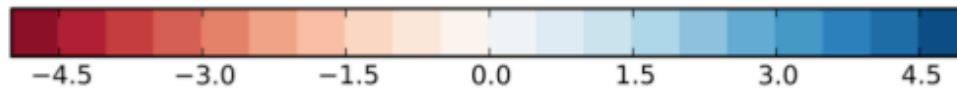


# Global map of VTI anisotropy: clear regional variations

Tomography 75 km above CMB

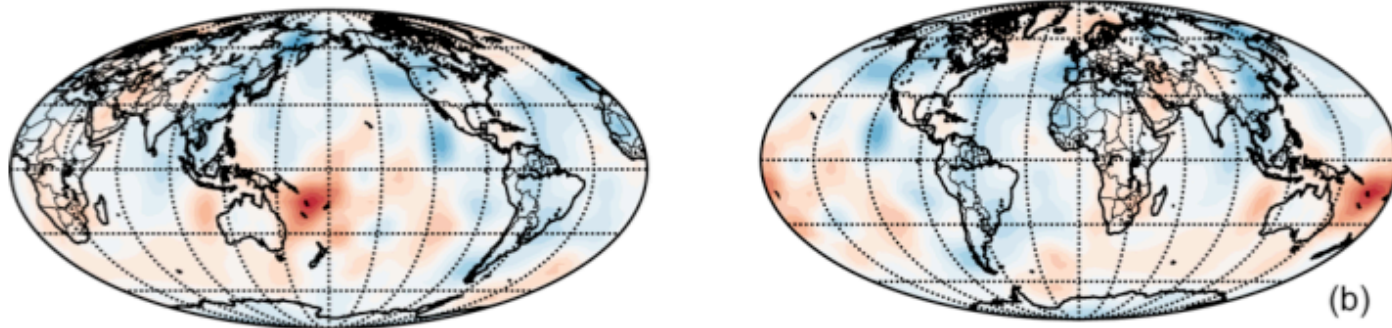


$V_{SH} < V_{SV}$



$V_{SH} > V_{SV}$

$d \ln(\xi) \%$



$V_{SH} < V_{SV}$



$V_{SH} > V_{SV}$

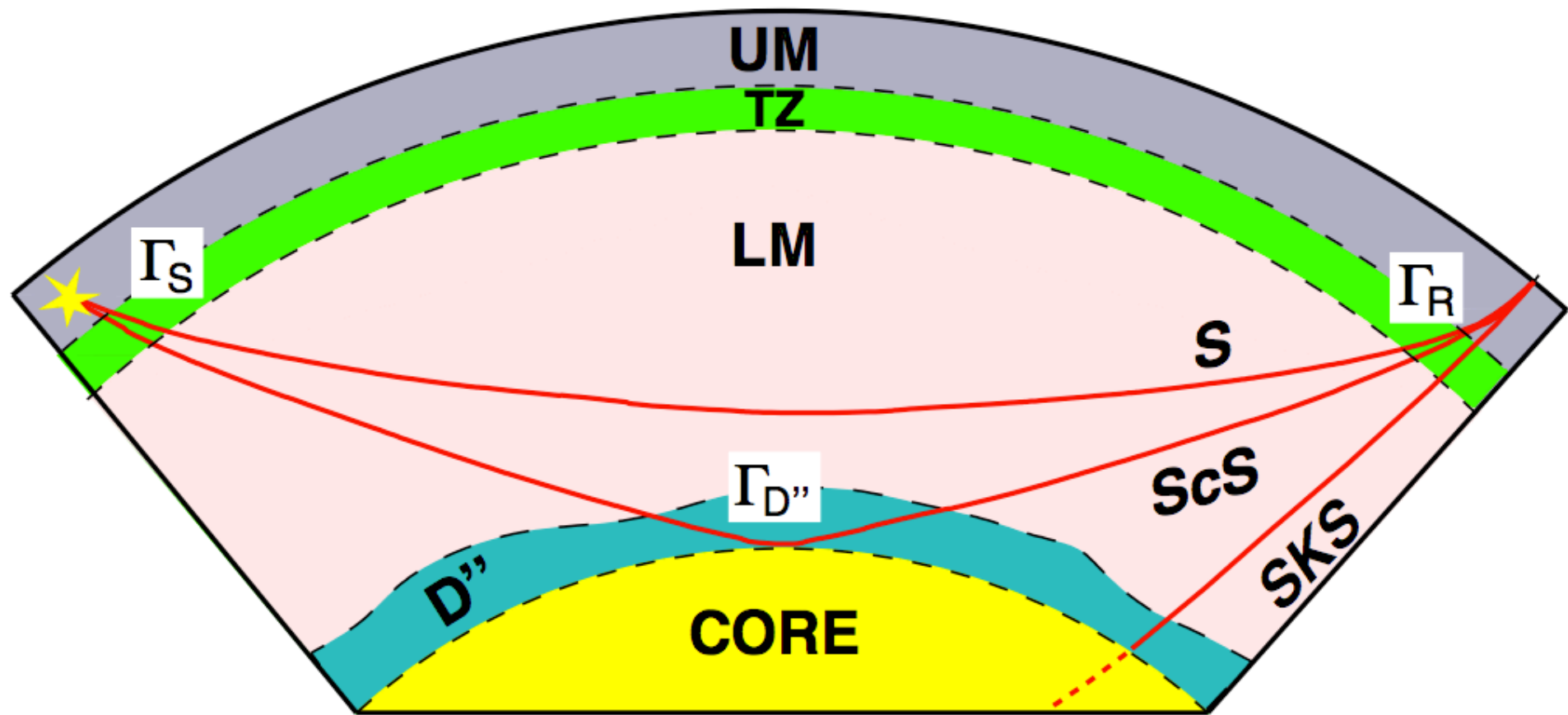
$d \ln(\xi) \%$

(a) Panning and Romanowicz 2006 *Geophys. J. Int.* 167:361-379

(b) Kustowski et al. 2008 *J. Geophys. Res.* 113: B06306



# Isolating more general forms of anisotropy (non-VTI)

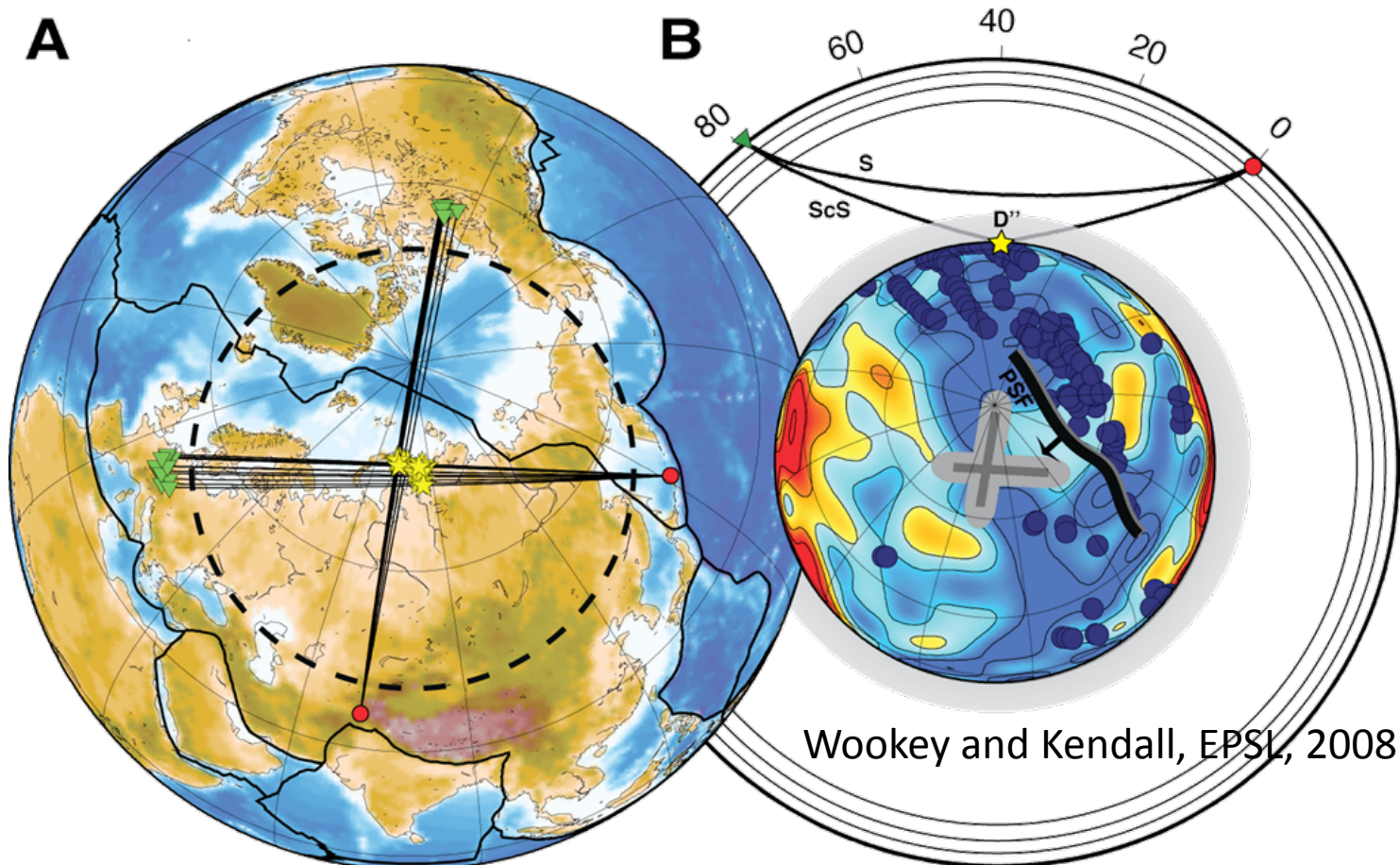


Source side shear-wave splitting analysis (e.g., Wookey et al., 2005; Wookey and Kendall, 2008; Nowacki et al., 2010; 2013)



# Crossing paths in the lowermost mantle:

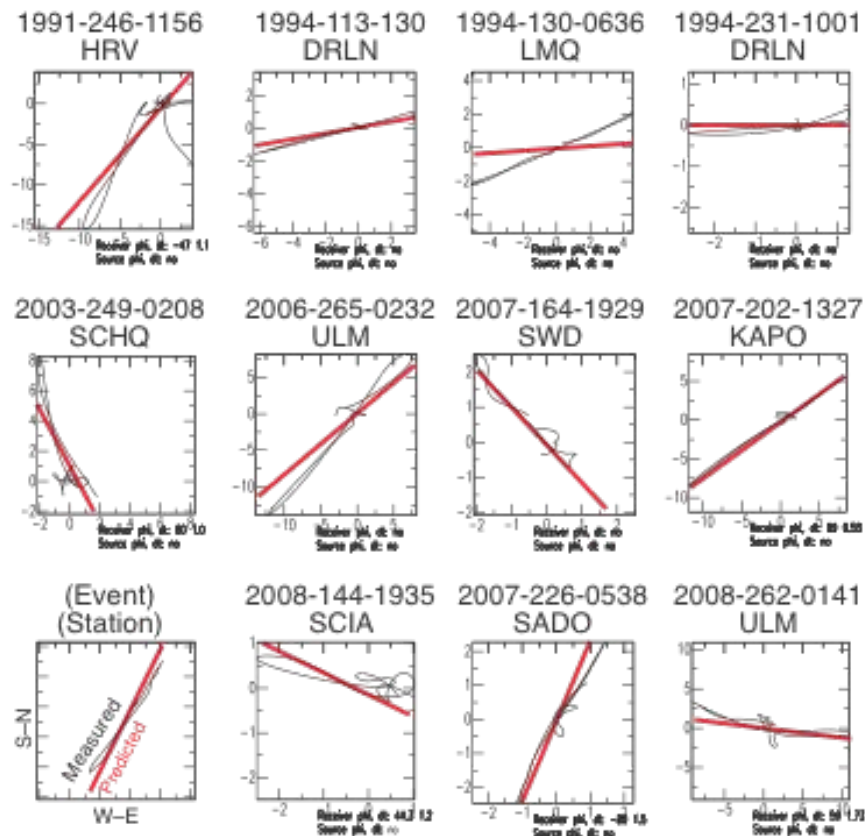
- Much better resolution of anisotropy (2 vectors of fast shear-wave polarisation to define the symmetry plane)
- Potential to distinguish TI from orthorhombic medium



# NWK methodology

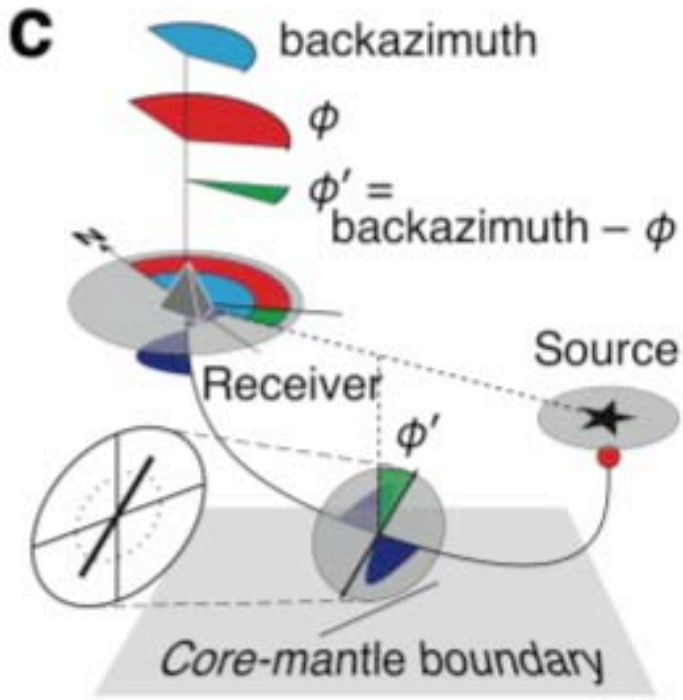
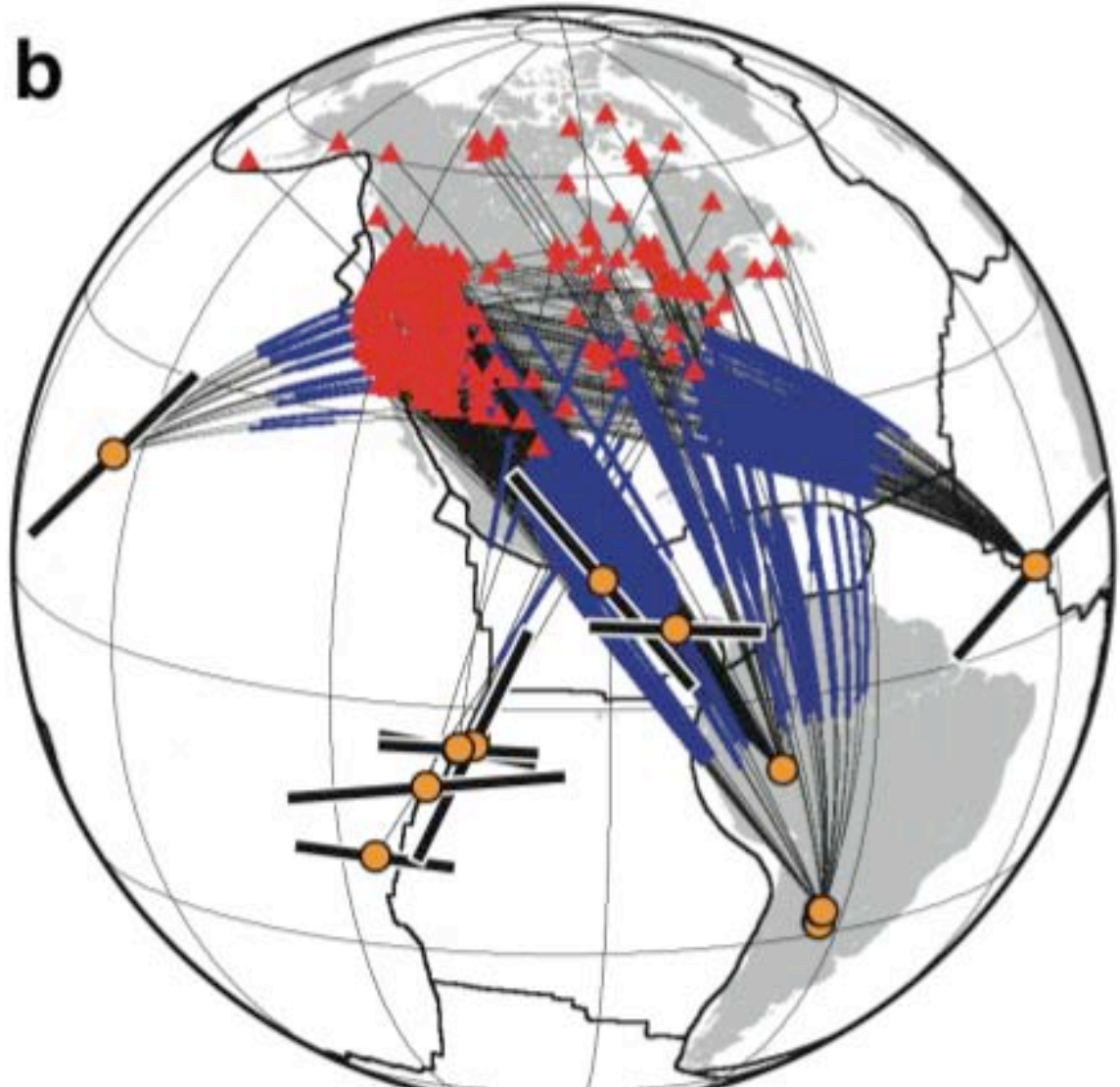
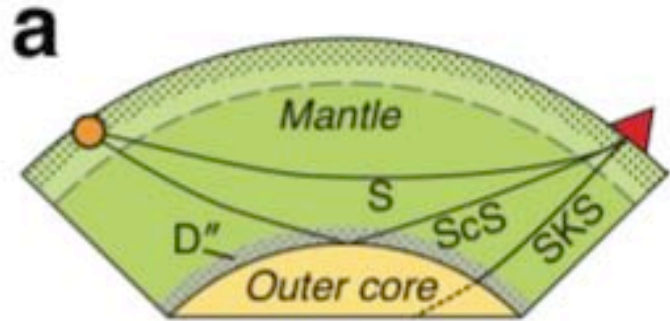
Rigorous data selection and quality control:

- Simple and well characterised receiver anisotropy.
- Use data from similar azimuths for both S and SKS.
- High-quality measurements, low error, clear signal and anisotropy.
- Inferred source polarisation must agree with CMT solution for the earthquake.



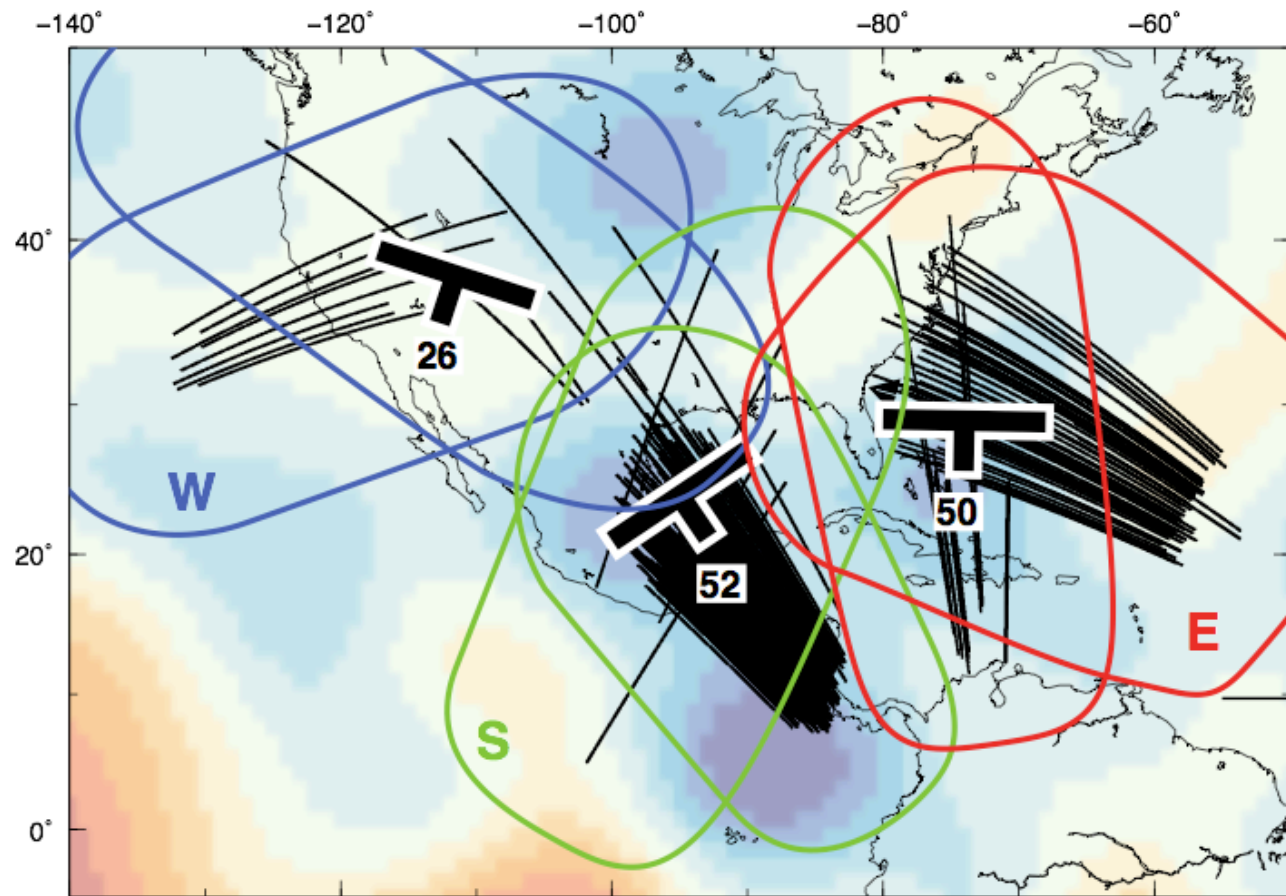
Nowacki, Wookey and Kendall, Nature 2010

# Anisotropy beneath the Americas



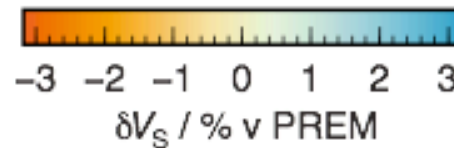
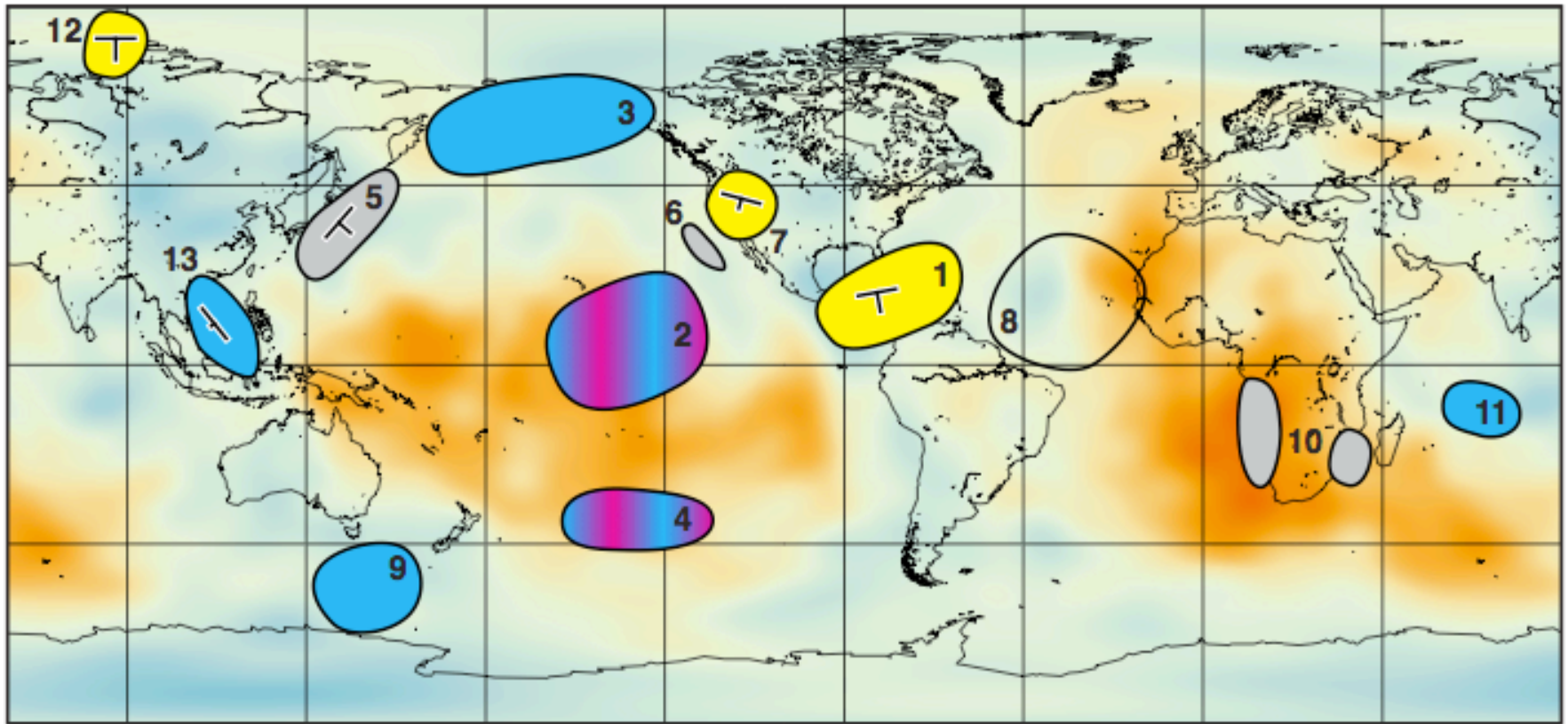


# Anisotropy beneath the Americas: best fitting TTI model



S20RTS  $\delta V_S / \% \text{ v PREM}$   
-1.5 -1.0 -0.5 0.0 0.5 1.0 1.5

# The current picture



Nowacki et al., J. Geodyn., 2011

  $V_{SH} > V_{SV}$

  $V_{SH} < V_{SV}$

  $V_{SH} > V_{SV}$   
 $V_{SH} < V_{SV}$   
 $V_{SH} \approx V_{SV}$

 TTI:  
dip  
60°

 TTI:  
dip  
40°

 TTI:  
dip  
20°

 See text

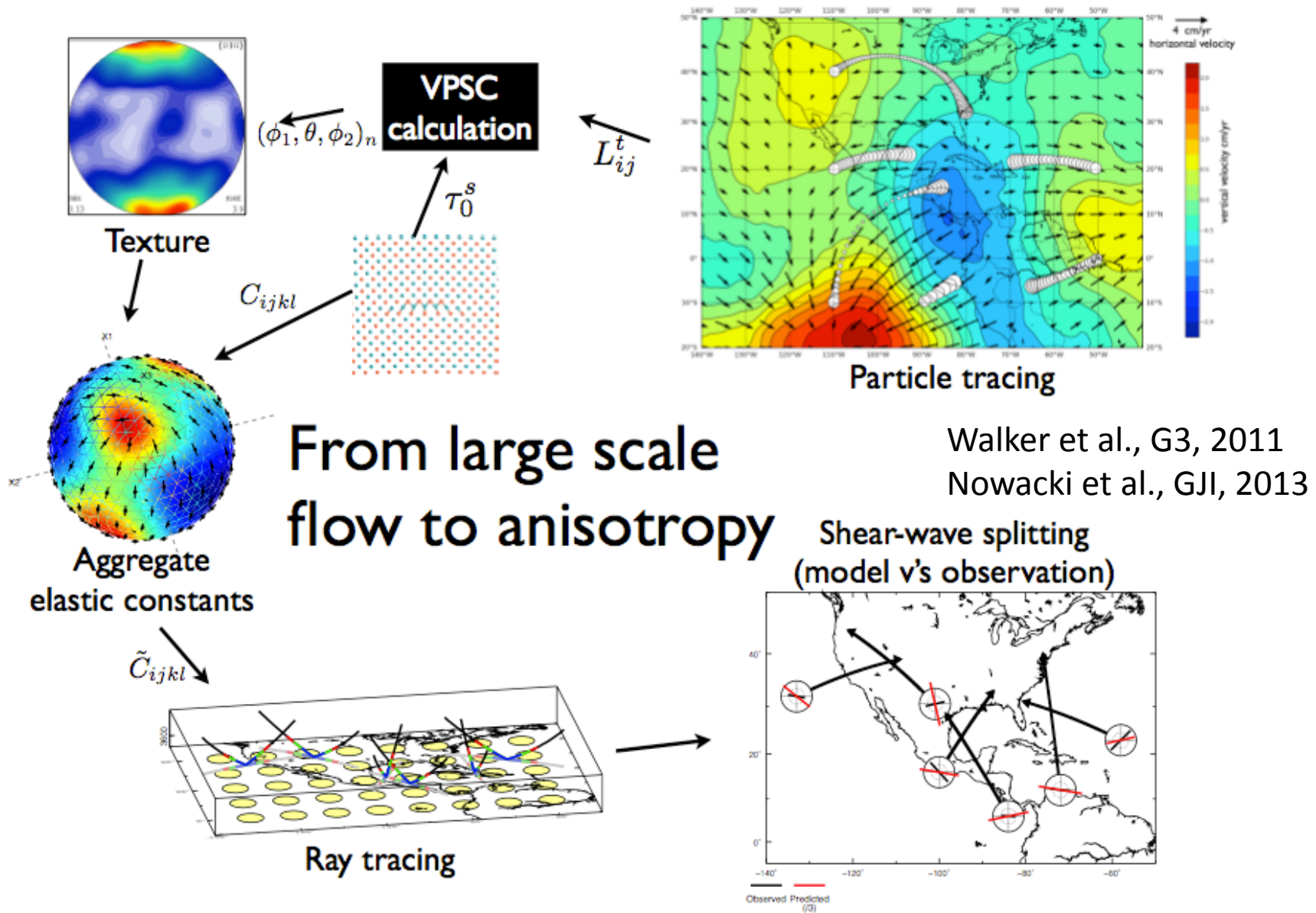
  $V_{SH} \approx V_{SV}$

## ... next steps

- We test ppv as a causative mechanism for D'' anisotropy.
- Use current best knowledge of ppv elasticity and slip systems
- Use state of the art models of mantle flow, which incorporate chemistry (e.g., Simmons et al., 2009)
- Ideally consider strain history in 3D
- Ideally impose no restriction on style of anisotropy

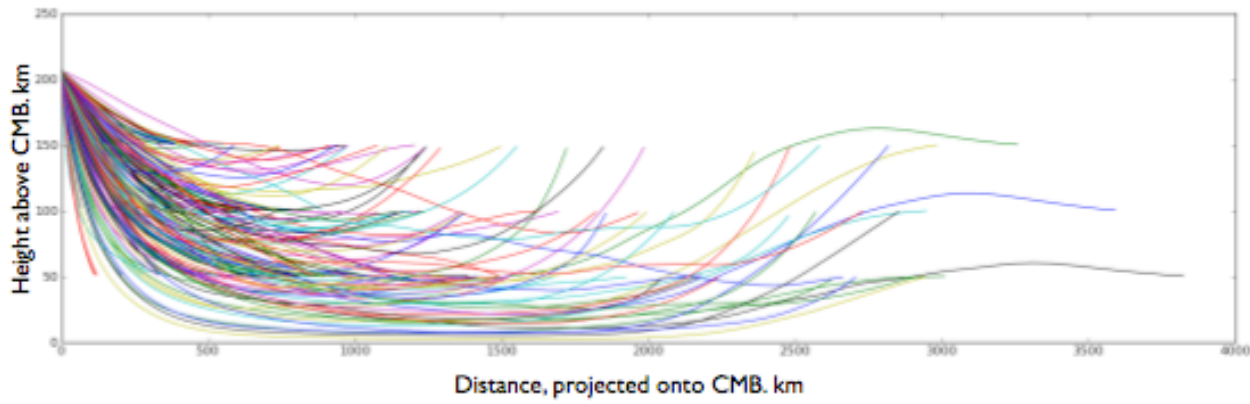
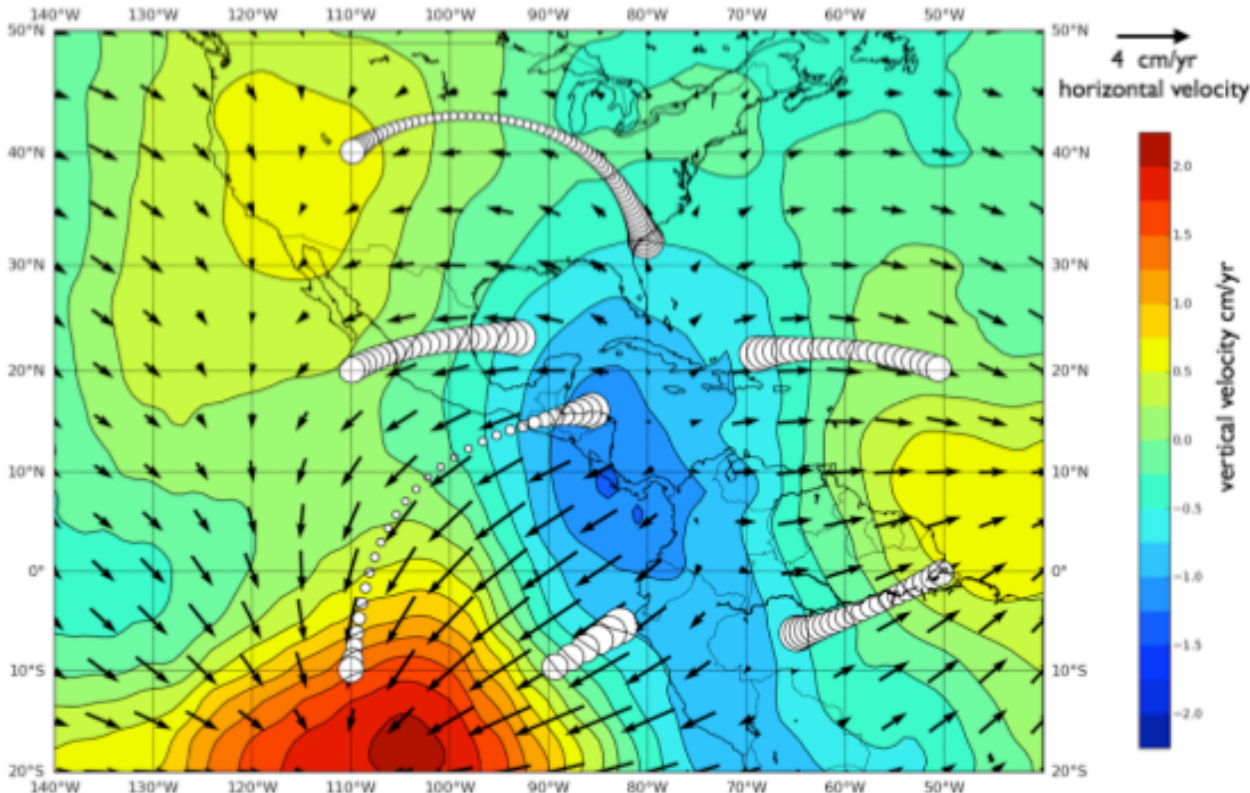
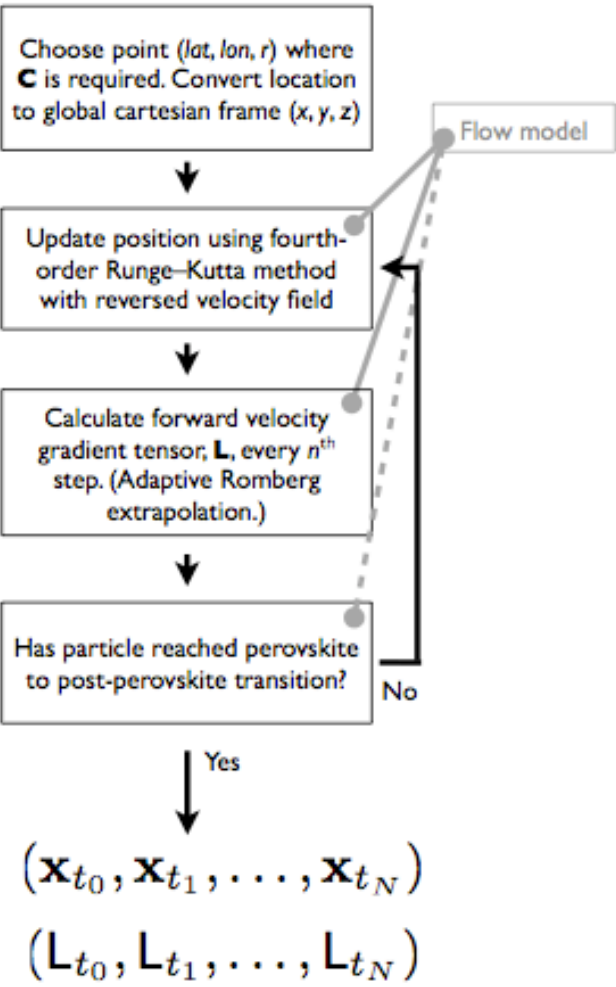
See: Walker et al., G3, 2011; Wenk et al., EPSL, 2011; Nowacki et al., GJI, 2013

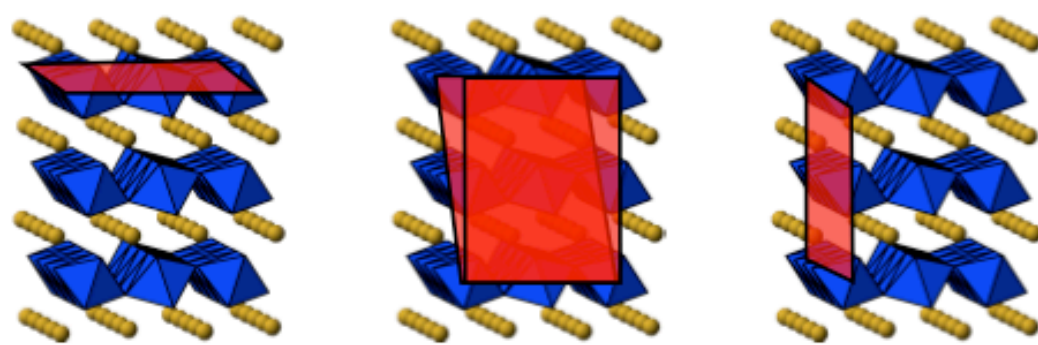
# Linking flow, mineralogy and seismic observations





# Flow model from joint inversion





Slip system	PN: (010) <sup>a</sup>	{110}/(100) <sup>b</sup>	(001) <sup>c</sup>
[100](001)	2.6	10	1
[010](001)	4.1	10	1
[001](010)	1.0	∞	∞
[001](100)	5.4	2	10
[010](100)	5.2	1	10
[001]{110}	2.9	4	10
<110>(001)	4.1	10	2
[100]{011}	6.8	∞	∞
[100](010)	4.7	∞	∞
<110>{110}	8.8	1	10

<sup>a</sup> Metsue *et al.* 2009 *PEPI* 74:165-173

<sup>b</sup> Merkel *et al.* 2007 *Science* 316:1729-1732

<sup>c</sup> This work (*c.f.* Miyagi *et al.* 2010 *Science* 329:1636-1638)

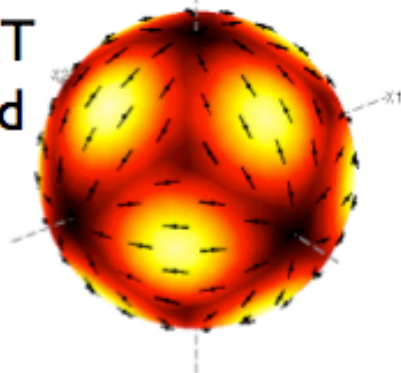


% S-wave anisotropy

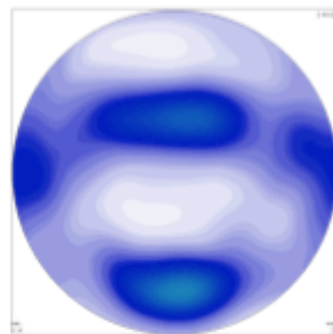


Single crystal elastic constants (from DFT at high pressure and temperature)

$$c(p, t)$$



Texture from VPSC calculation



$$g(\varphi_1, \Phi, \varphi_2)_1 \dots g(\varphi_1, \Phi, \varphi_2)_n$$

Voigt-Reuss-Hill average

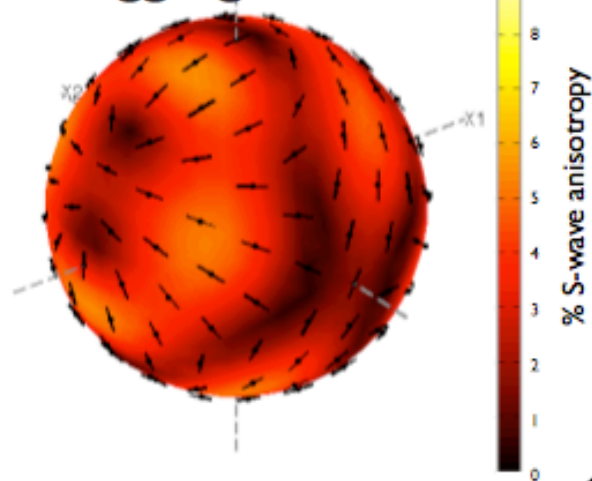
$$s(p, t) = c^{-1}(p, t)$$

$$C_{ijkl} = \frac{1}{500} \sum_{n=1}^{500} g_{i\alpha}^n g_{j\beta}^n g_{k\gamma}^n g_{l\delta}^n c_{\alpha\beta\gamma\delta}$$

$$S_{ijkl} = \frac{1}{500} \sum_{n=1}^{500} g_{i\alpha}^n g_{j\beta}^n g_{k\gamma}^n g_{l\delta}^n s_{\alpha\beta\gamma\delta}$$

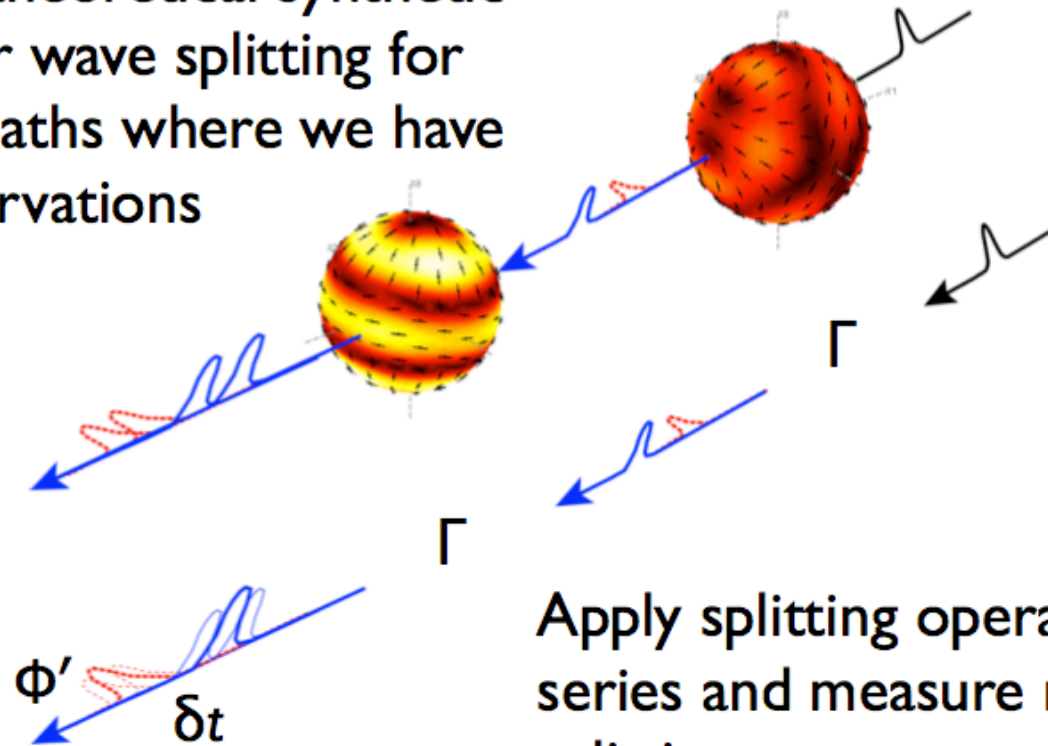
$$\tilde{C}_{ijkl} = \frac{C_{ijkl} + S_{ijkl}^{-1}}{2}$$

Elastic constants of textured aggregate



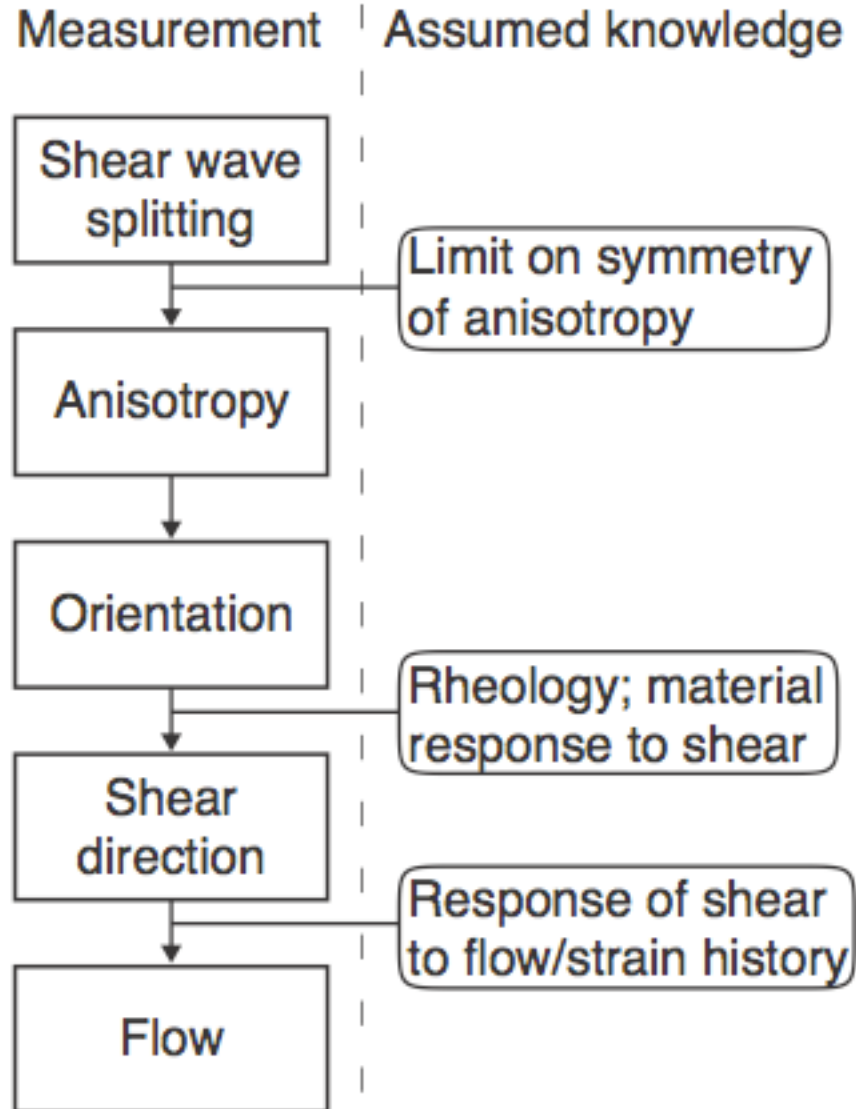
# Generating Synthetic Seismograms

Ray-theoretical synthetic shear wave splitting for ray paths where we have observations



Apply splitting operators in series and measure resultant splitting parameters.  
NB:  $\delta t$  smaller than period of wave

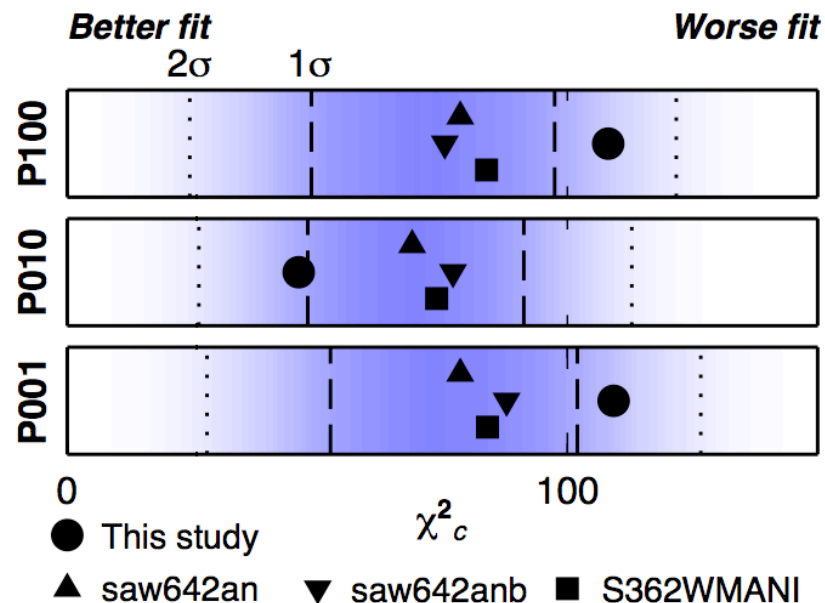
# From observation to flow



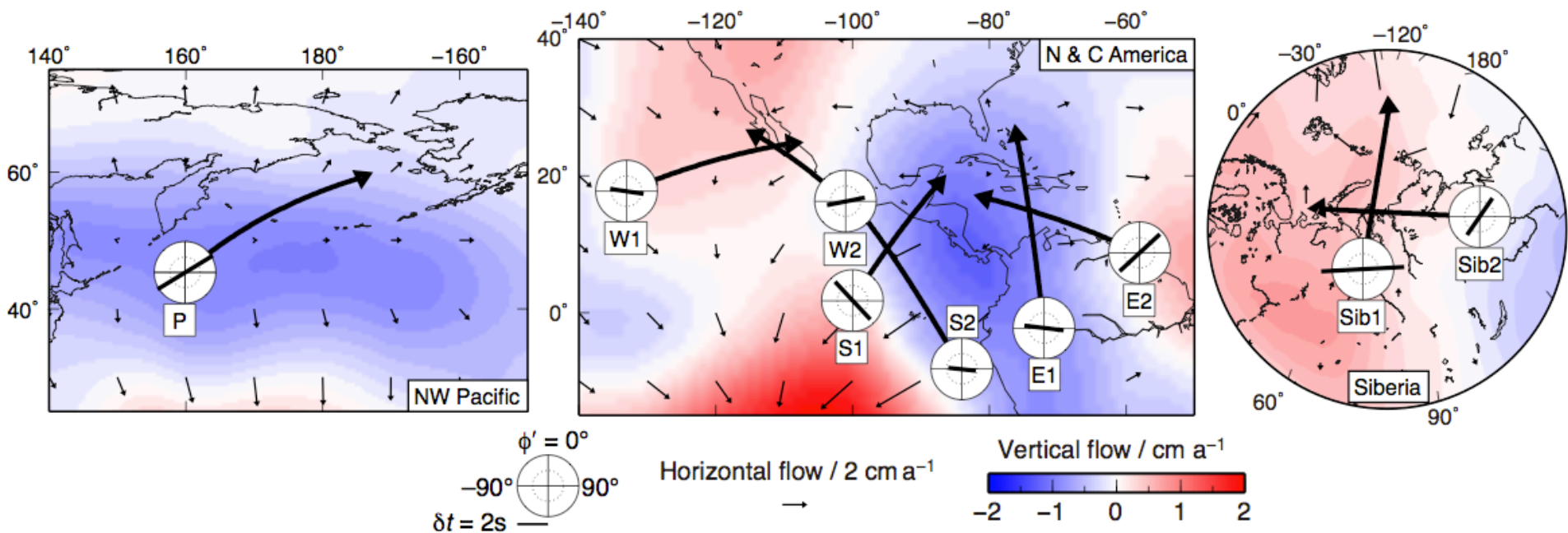


# Which slip system best fits the data? Ans = P010

- flow 200 km above the CMB; predicted by TX2008 (Simmons et al., 2009)
- VTI models: Panning et al. (2004; 2010); Kustowski et al. (2008)
- General form: Wookey et al. (2005; 2008); Nowacki et al. (2012)



Nowacki et al., GJI, 2013



# Conclusions

- In general, the lowermost mantle is anisotropic, but clear regional variations are evident.
- Anisotropy is the fingerprint of D'' mineralogy and style of deformation.
- Compared observations with numerical predictions from linked flow and LPO modelling.
- Three different plasticity cases for dislocation creep in ppv are considered, with that favouring slip on (010) matching best (at least in regions of subduction).



# Future Directions

- More observations – global dataset
- Full waveform modelling – how important are finite-frequency effects?
- To what degree do dynamic recrystallisation, work hardening, climb and the presence of other phases change the texture?
- What are the effects of inherited petrofabric textures (e.g, pv to ppv)?
- Effects of melt
- Is LPO in post-perovskite important everywhere in D'' - do other anisotropy generating processes play a role?

# Towards better data coverage

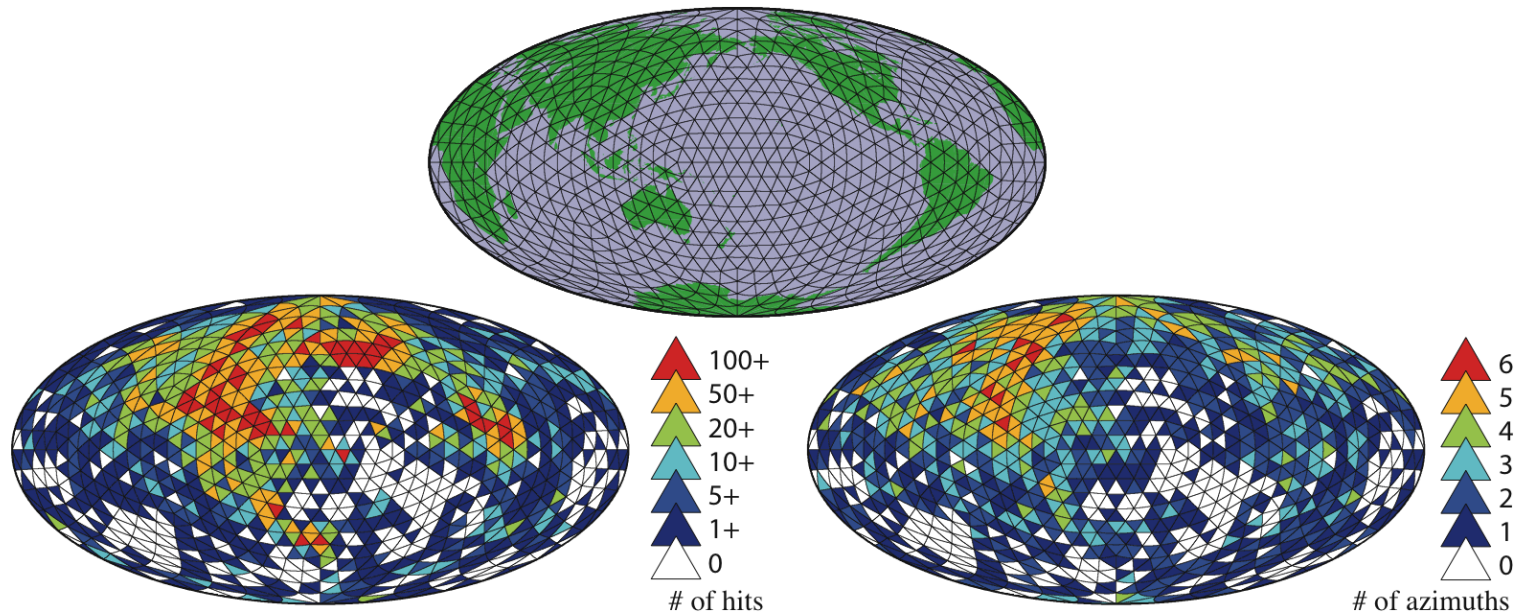


Image courtesy of Jack Walpole