



# Deep thermal structure of the cratonic lithosphere

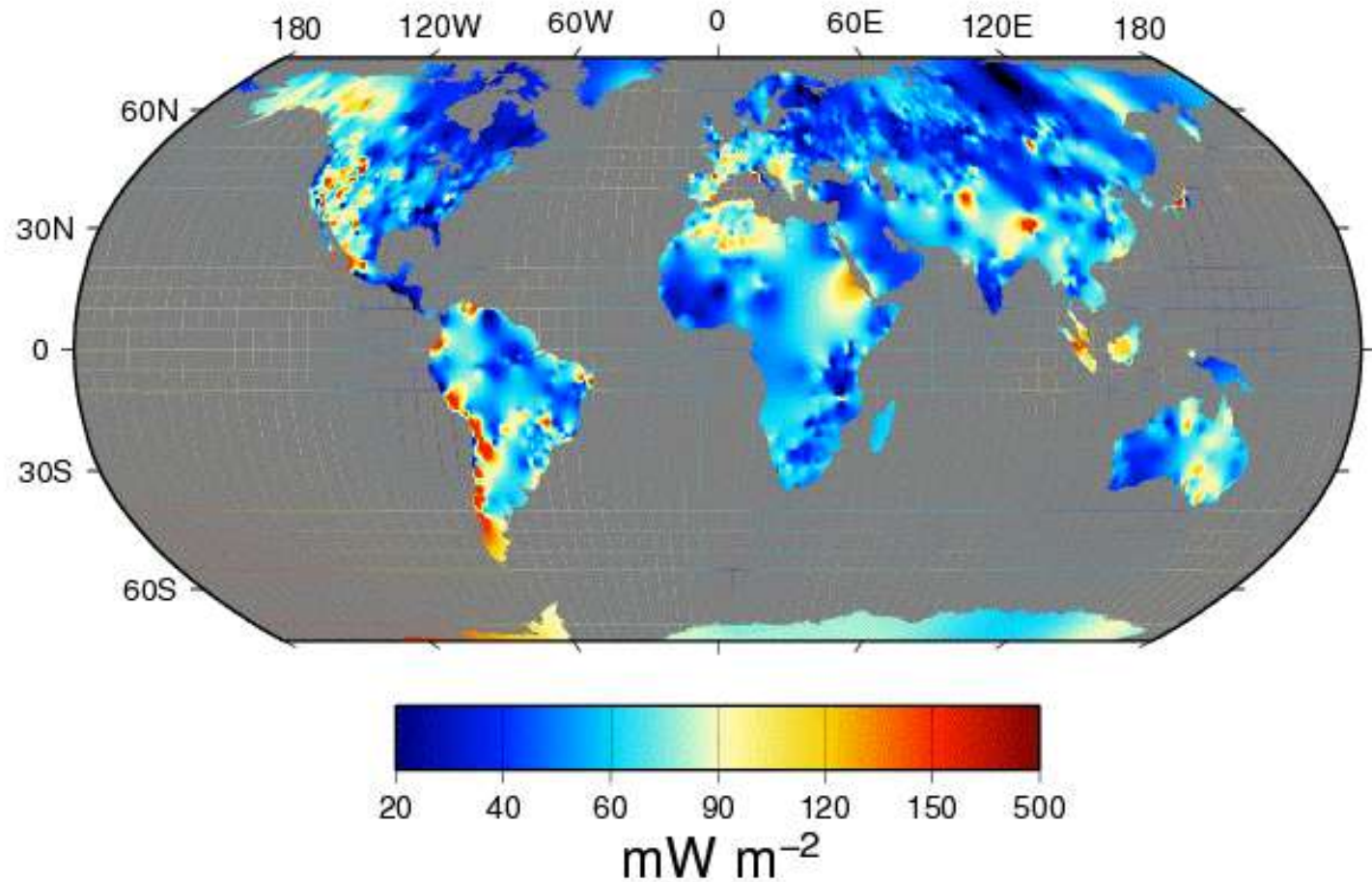
Jean-Claude Mareschal, with a little help from my friends, Claude Jaupart, Genevieve Savard, and many others...

# Deep thermal structure of lithosphere?

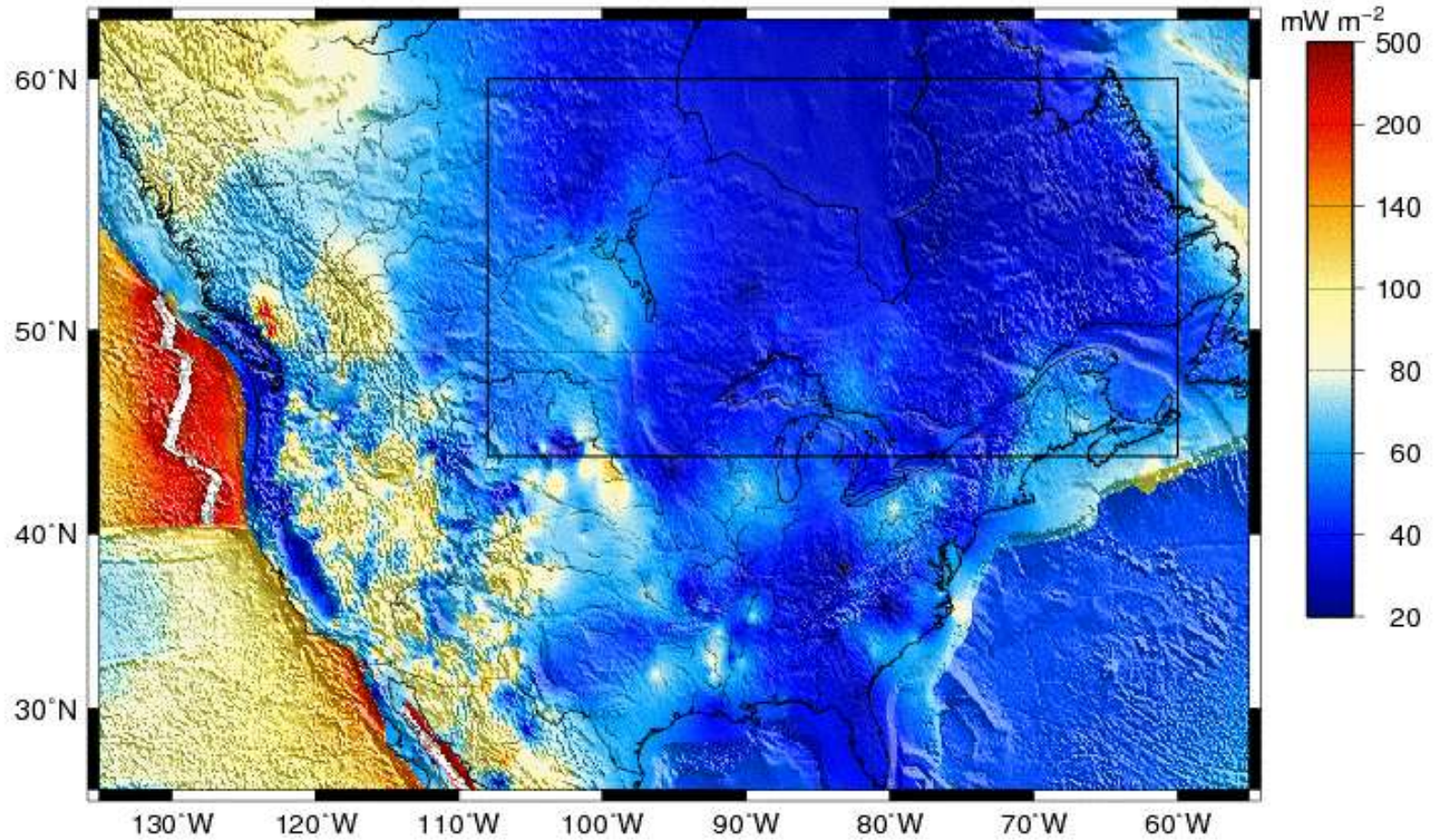
- What determines mantle temperature?
  - Surface heat flux
  - Vertical distribution of HPEs
- Controls on temperature models?
  - Xenoliths
  - Shear wave velocity profiles

# Continental heat flux

## Cold cratons: EOS?

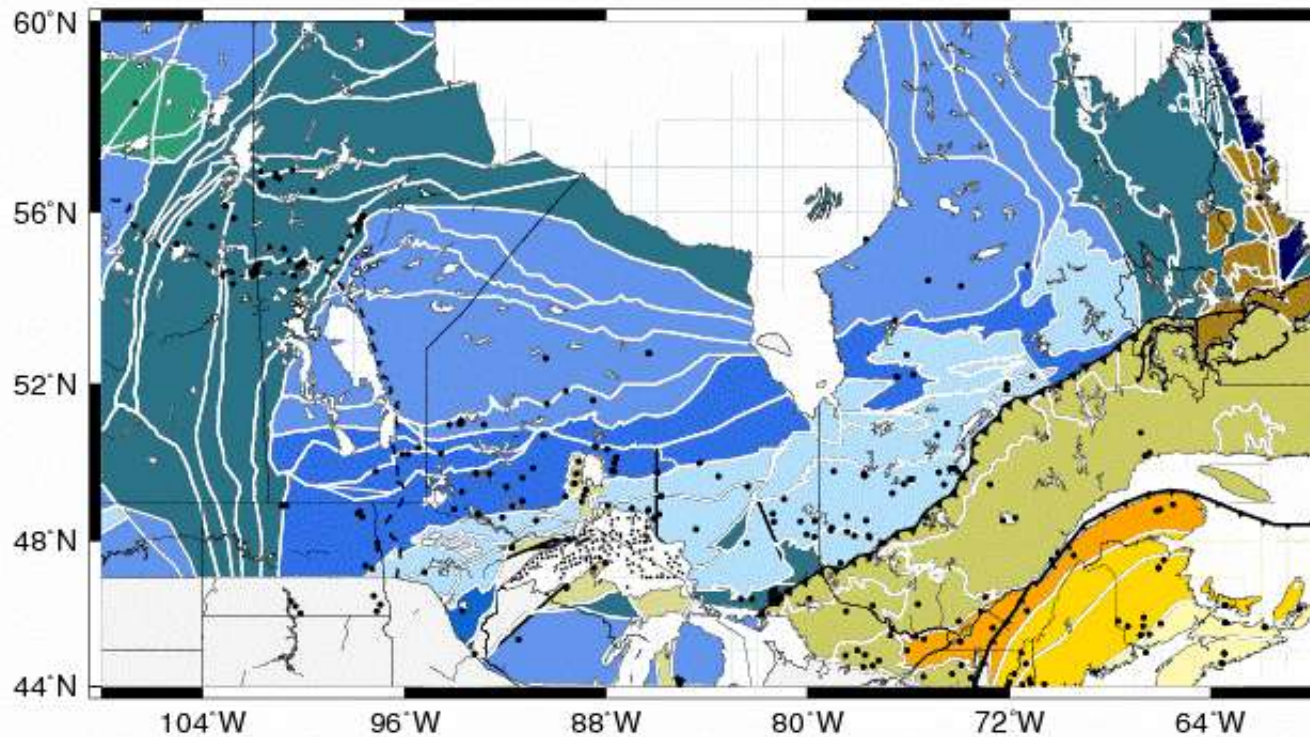


# Heat flow



# Canadian Shield

## Basement Age



# How do we calculate temperature?

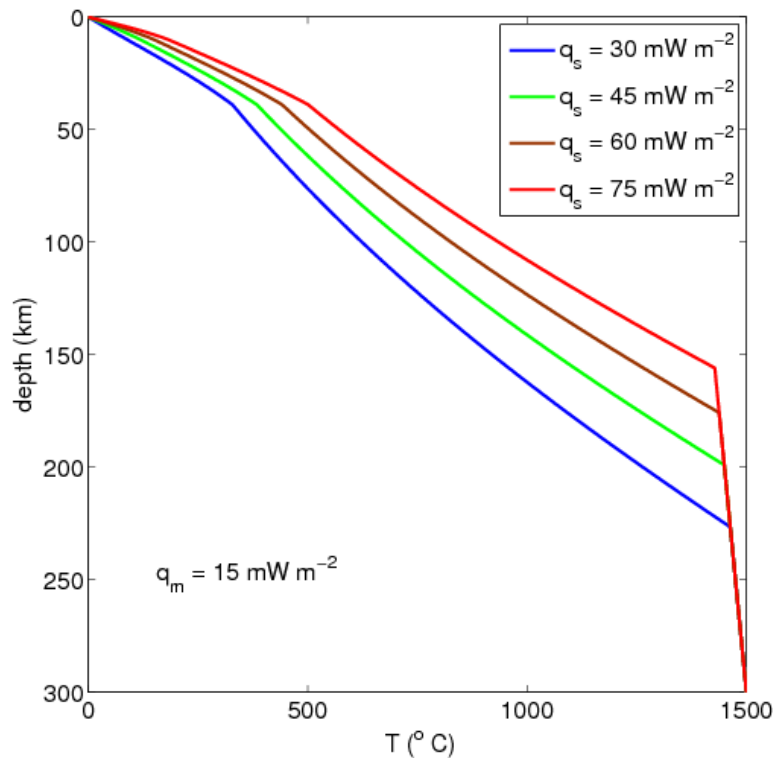
## Steady-state 1-D equation for $\langle T(z) \rangle$

- $Q(z) = -\lambda(T) \frac{dT}{dz}$
- $\frac{dQ}{dz} = H(z)$
- $\frac{d}{dz} \lambda(T) \frac{dT}{dz} = -H(z)$
- $T(z=0) = T_0$
- $Q(z=0) = Q_0$
- $\lambda(T)$  thermal conductivity
- Heat production  $H(z)$ ?

## Controlling parameters for temperature profile

- $Q_0$
- $Q_m = Q_0 - \int_0^{z_m} H dz$
- $H_m$  ? small?
- $H_c$  in crust characterized by differentiation index, DI
- $DI = \frac{z_m \langle H_0 \rangle}{Q_0 - Q_m}$

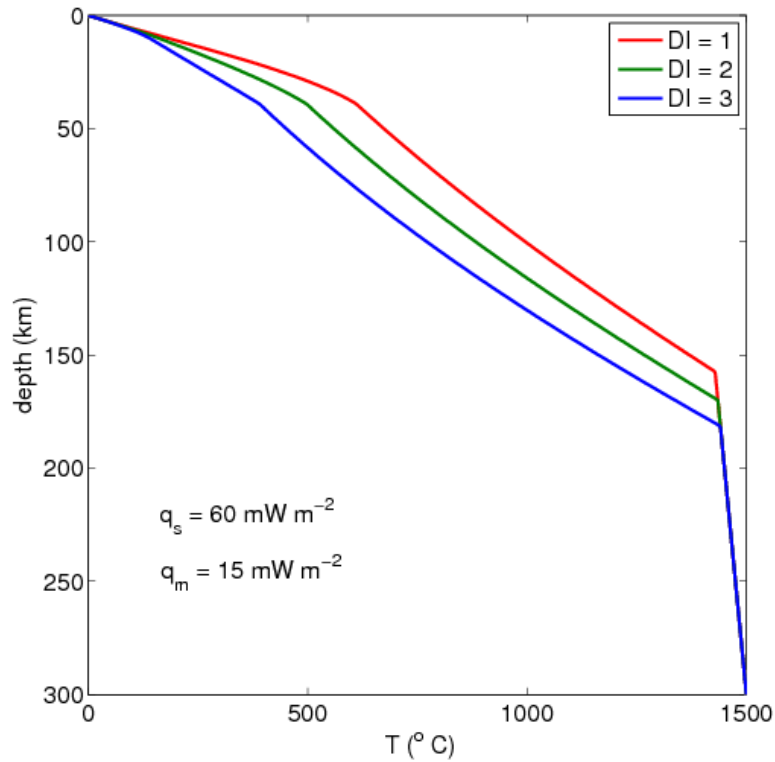
# Surface heat flux?



Crustal HP uniform,

- In shields,  
 $20 < Q_0 < 80 mW m^{-2}$
- In Canadian Shield  
 $26 < \langle Q_0 \rangle < 60 mW m^{-2}$

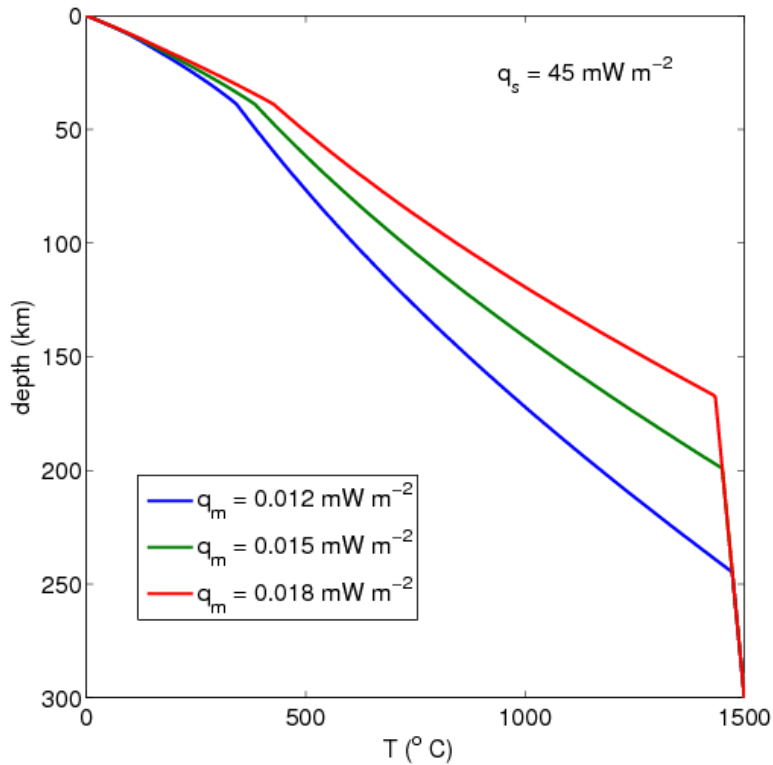
# Differentiation index



- In Canadian Shield
  - $DI = 0.7$  (Abitibi Subprovince)
  - $DI = 2.2$  (Slave Province)



# Moho heat flux

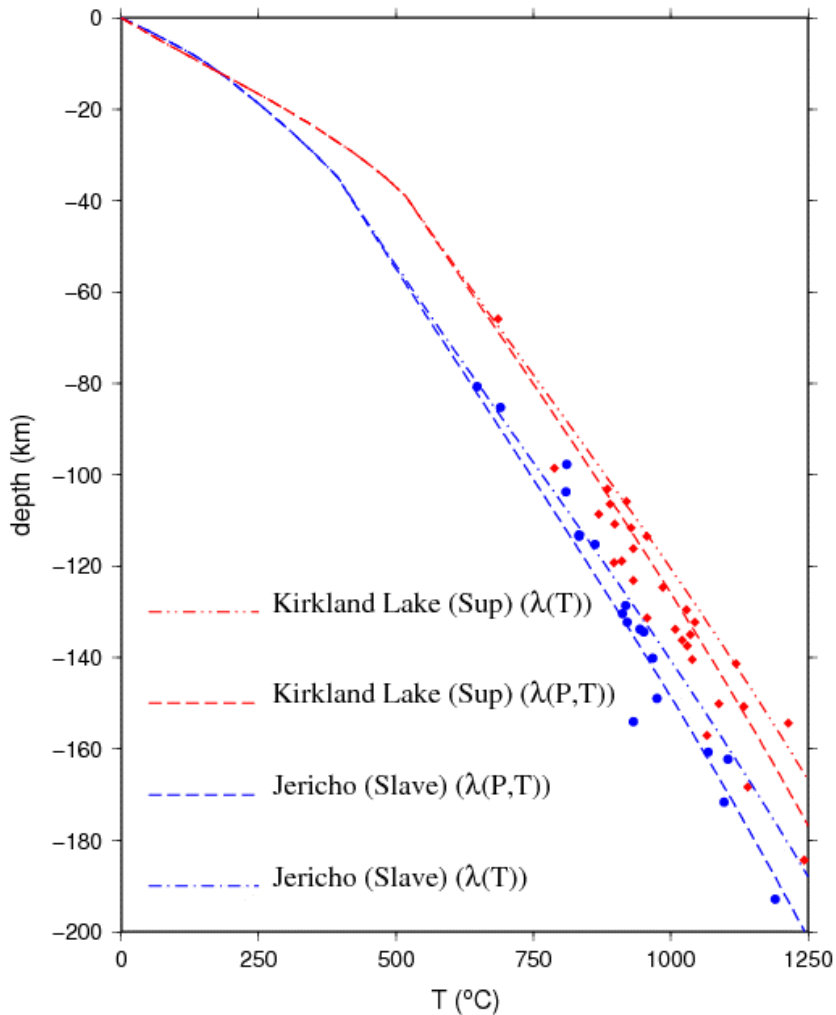


- Largest differences in mantle temperatures come from small differences in  $Q_m$
- Difficult to resolve  $Q_m$  to  $< 3 \text{ mW m}^{-2}$  from surface heat flux only

# Mantle heat flux variations in stable continents

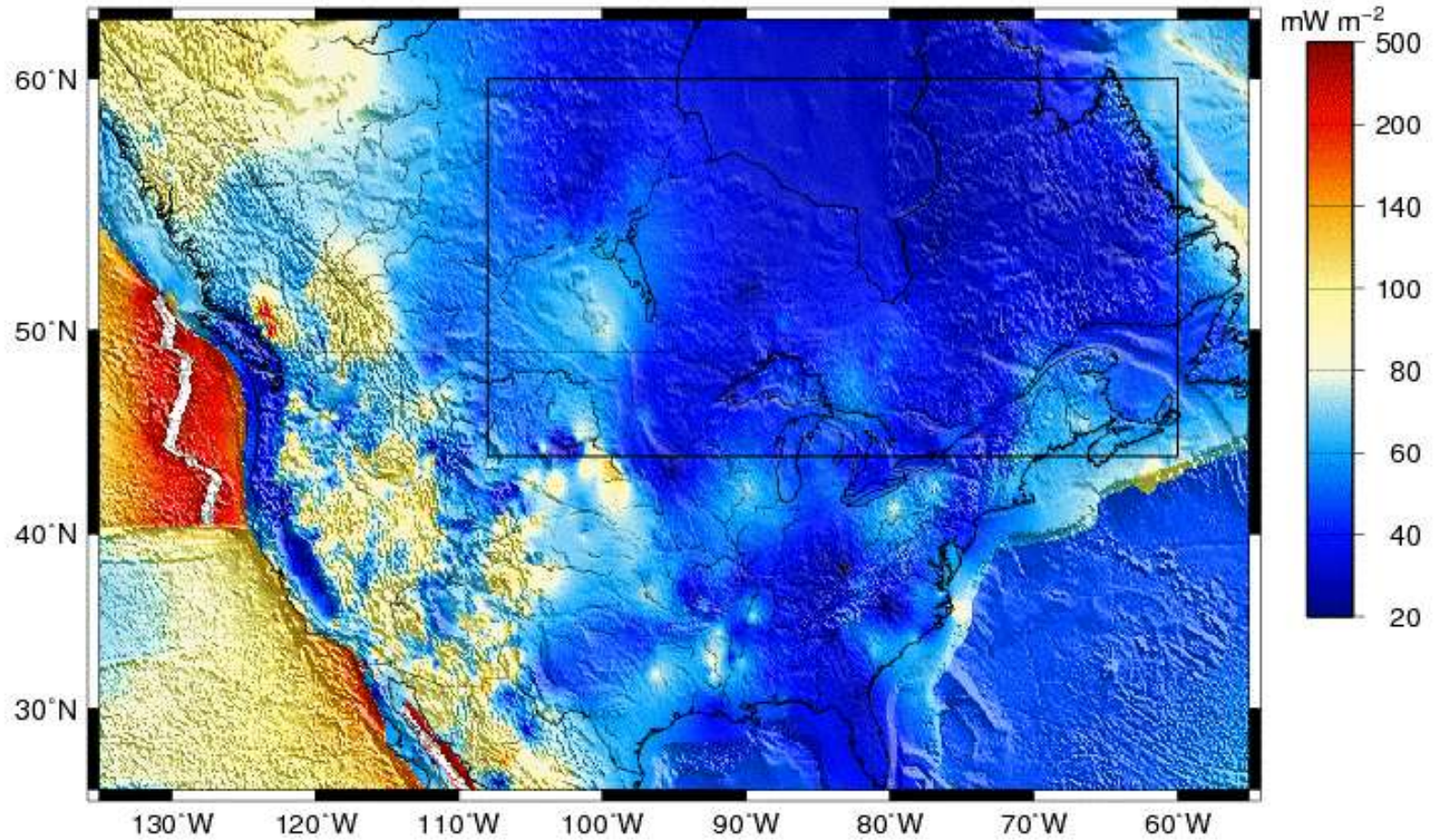
- $Q_m < 22 \text{ mWm}^{-2}$  (lowest heat flux measured)
- Best estimates  $13 < Q_m < 15 \text{ mWm}^{-2}$ 
  - Kapuskasing & other crustal sections
  - Grenville province
  - Gravity, crustal thickness, and heat flux data inversion
- $Q_m$  can not vary by more than  $\pm 3 \text{ mWm}^{-2}$
- Variability small but very significant

# Geothermobarometry on xenoliths from Slave and Superior Provinces

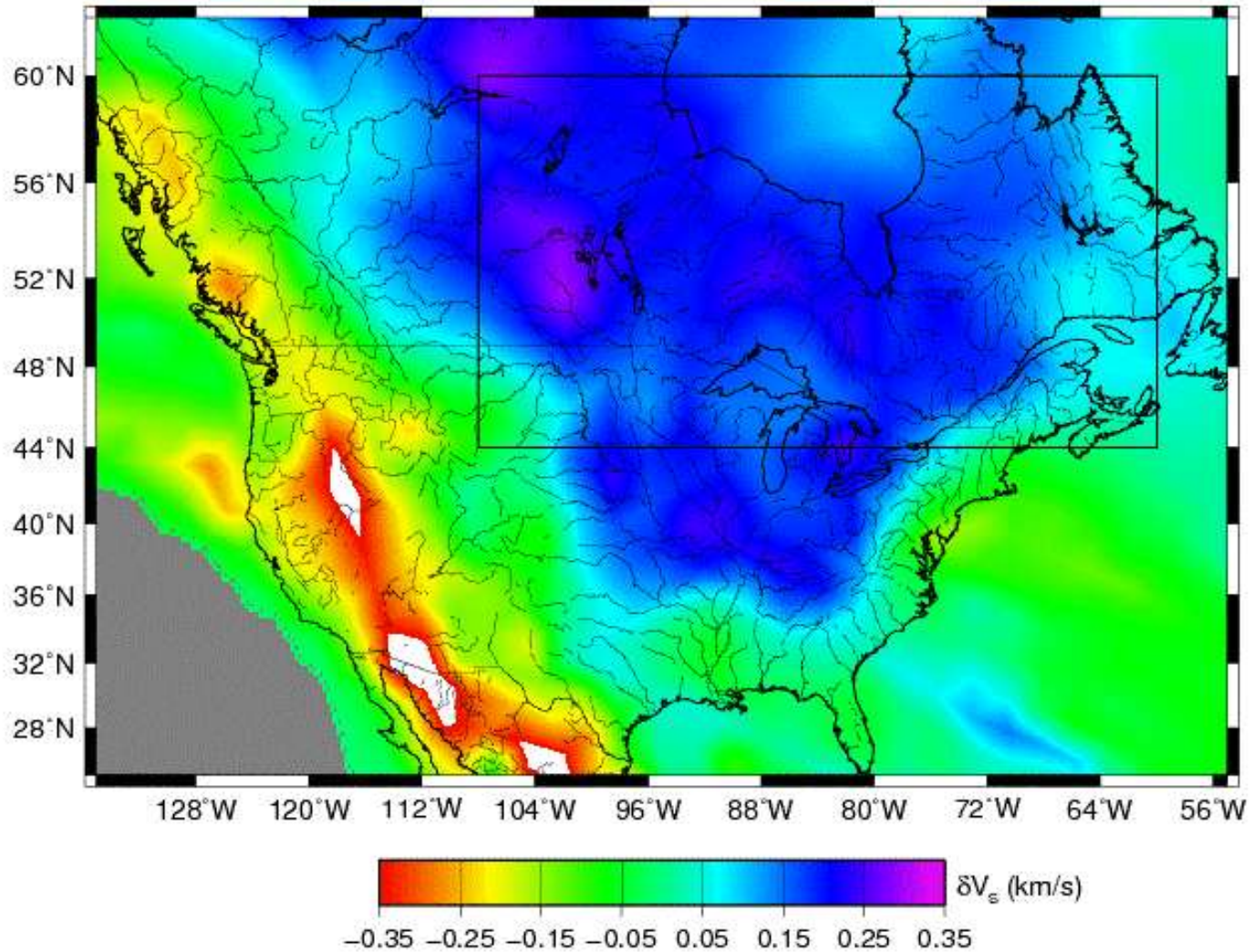


- Mantle xenoliths in kimberlites
- Phase equilibrium relationships  $\Rightarrow$  (P T) conditions at time of eruption
- Mantle temperature profile and heat flux
- $Q_0$  higher in Slave, but mantle temperature higher in Superior

# Heat flow

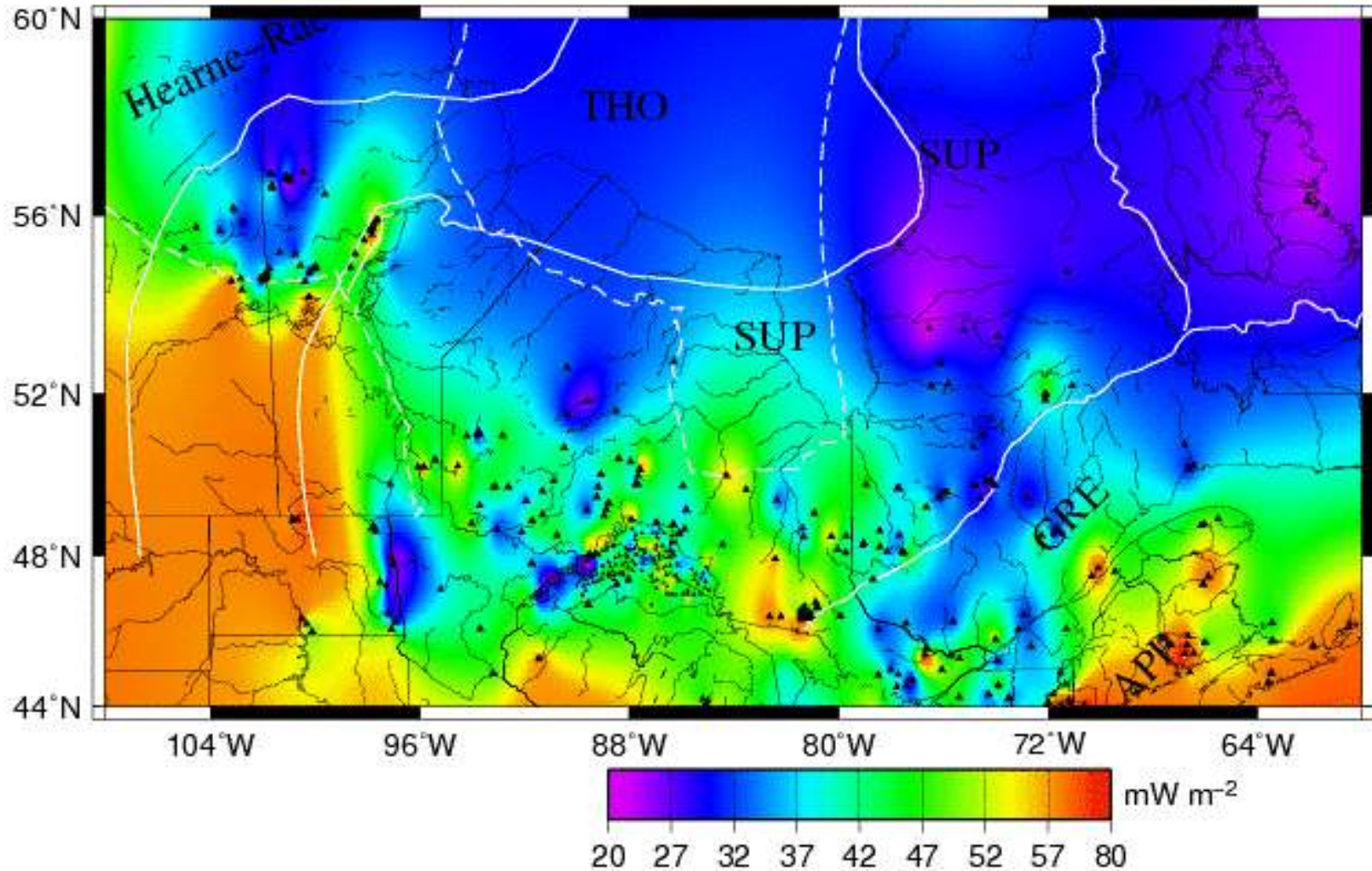


# $\Delta V_s$ 150km (NA07)

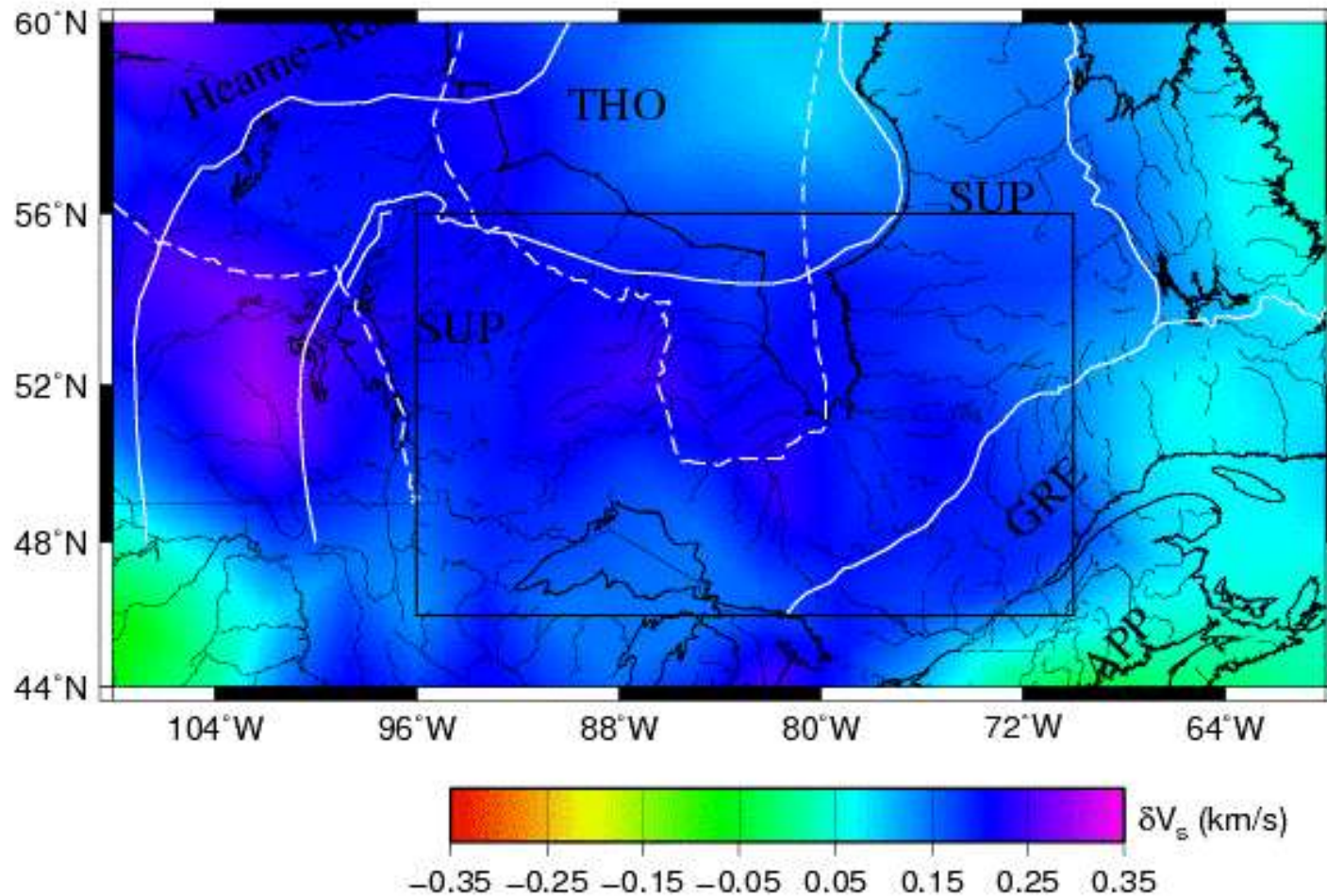


Beddle, H. and Van der Lee, S., 2009, JGR, 114, B7, B07308 .

# Surface heat flux in SECS

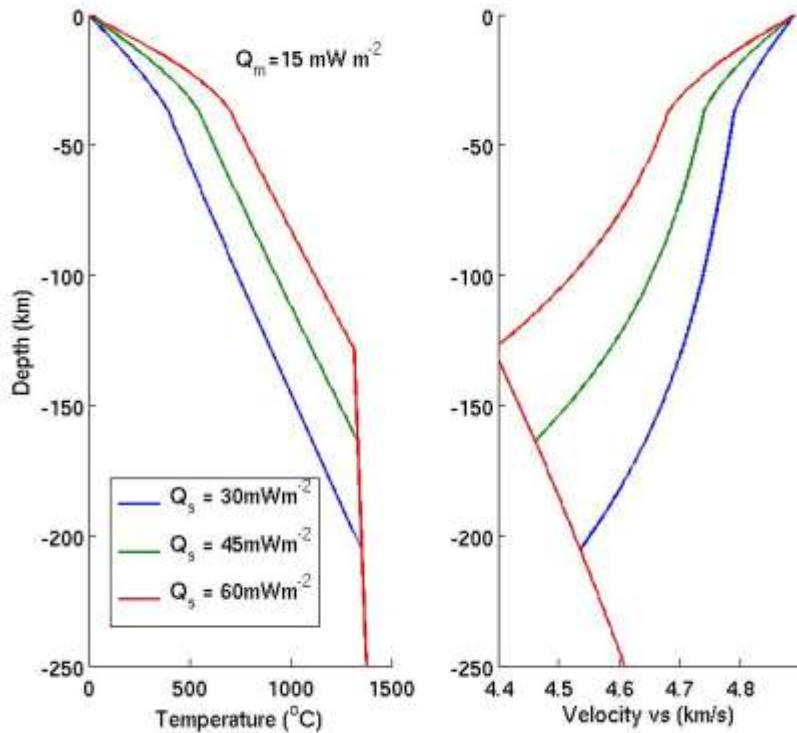


# Shear wave velocity anomaly (150km)

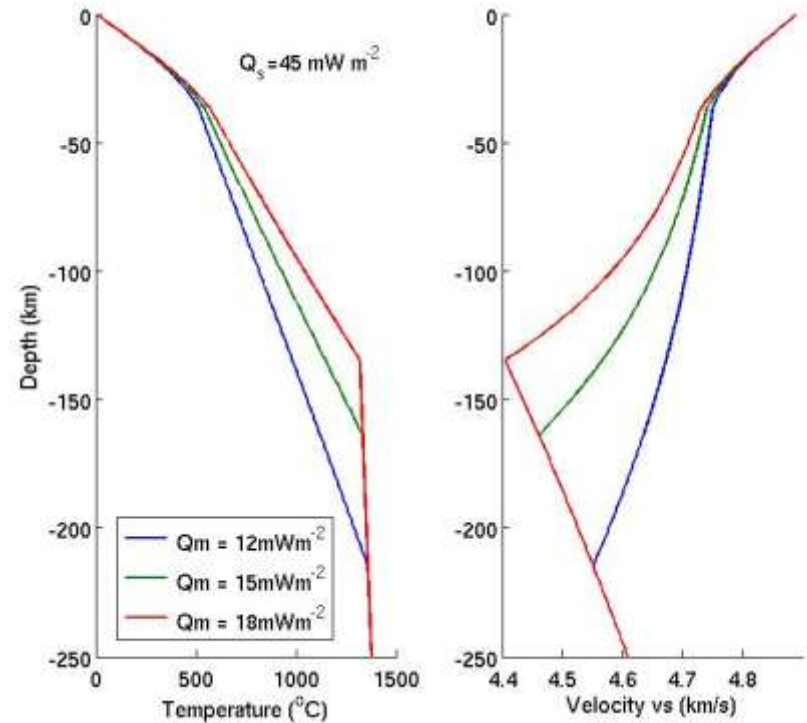


# Temperature and velocity profiles

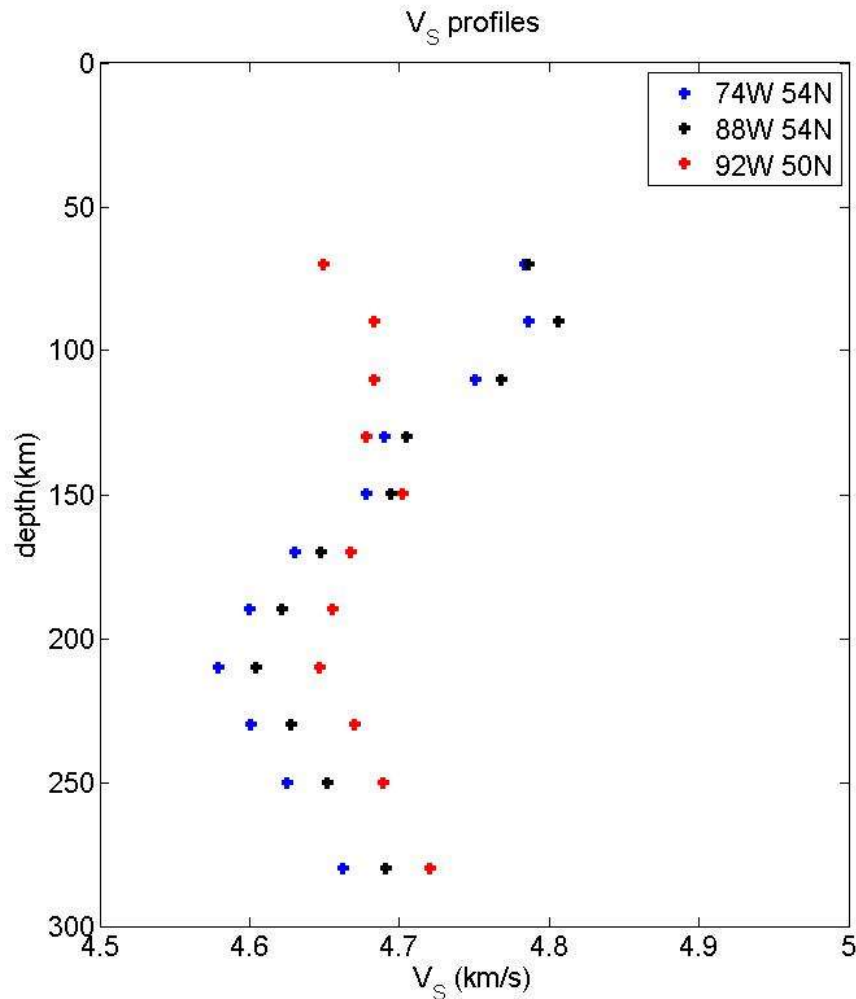
Large change in  $Q_s$   
No change in  $Q_m$



Small change in  $Q_m$   
No change in  $Q_s$





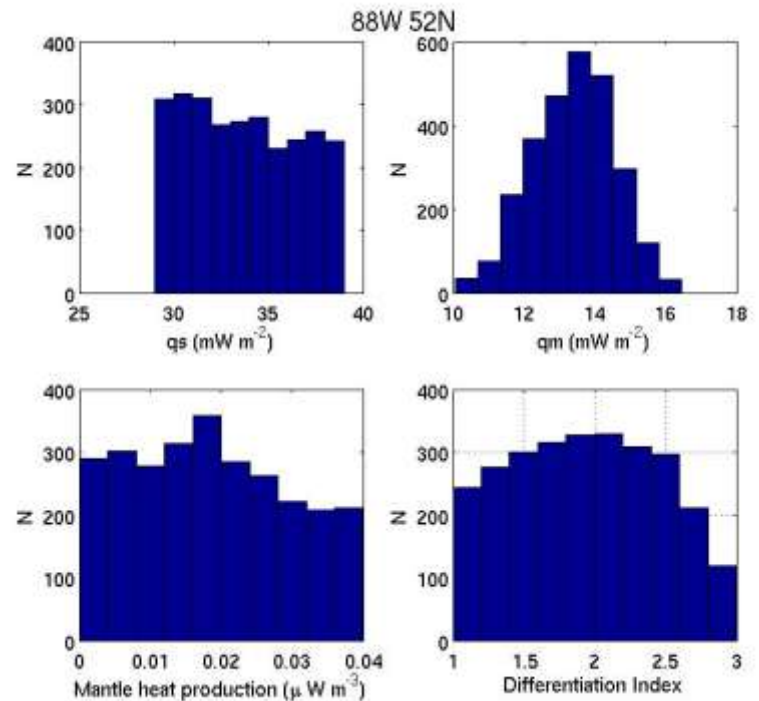
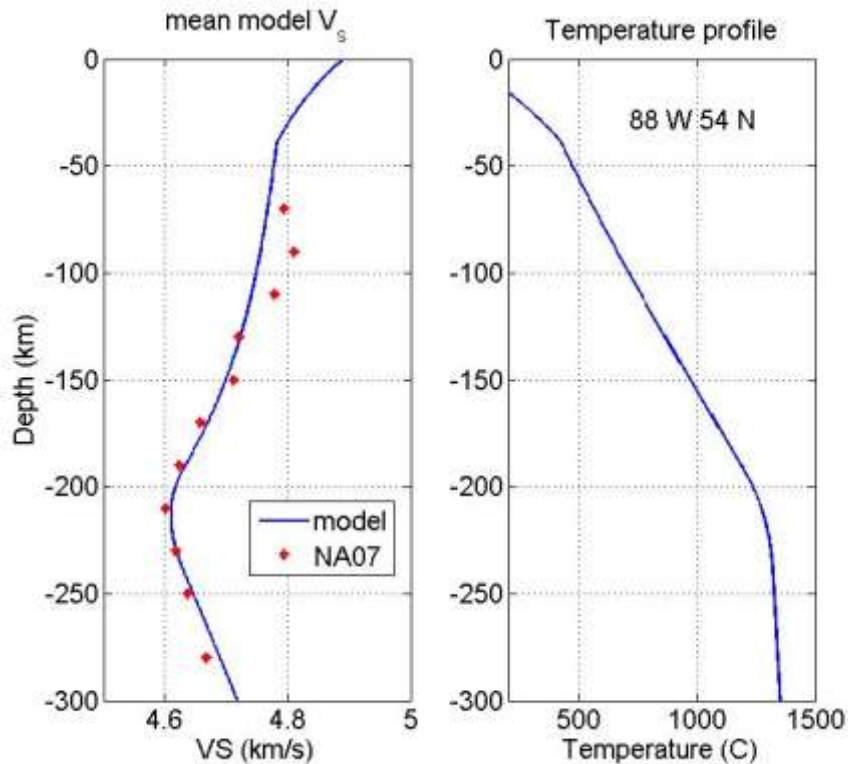


- Variations due to temperature
- Lithospheric temperatures depend on:
  - $Q_s$
  - Vertical distribution of HPEs ( $Q_m \leq$  total crustal HP and  $DI =$  crustal differentiation, mantle HP small)

# Monte Carlo inversion of vertical velocity profiles and surface heat flux

- Free parameters for calculating temperature profiles
  - $Q_0$  (varies +/- 4 mWm<sup>-2</sup> around average value in cell)
  - $Q_m$
  - $\langle hp \rangle = (Q_m - Q_s) / z_m$
  - $DI = hp(\text{surface}) / \langle hp \rangle$
  - Mantle heat production ( $<0.04 \mu\text{Wm}^{-3}$ )
- Crustal thickness constrained by seismic refraction or receiver functions.
- Different compositions (McDonough & Rudnick, Griffin Archon, ... )

# Example best fitting model and parameters for 1 cell

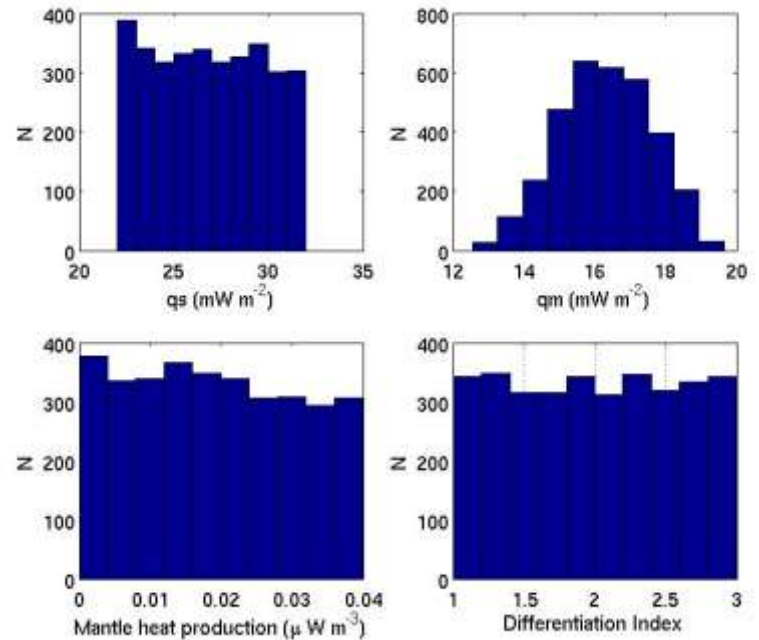
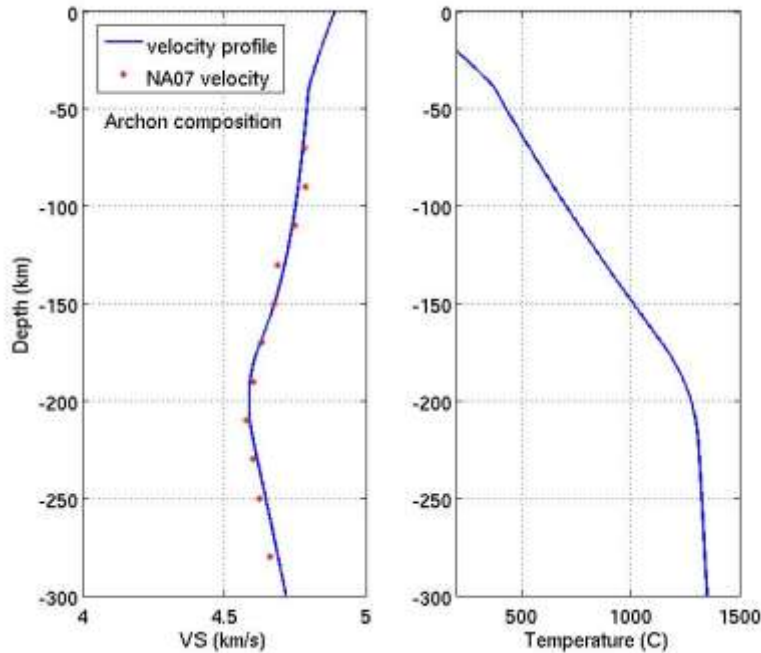


Composition Griffin Archon

Measured  $Q_s = 34 \text{ mWm}^{-2}$

Location West of James Bay

# Example best fitting model and parameters

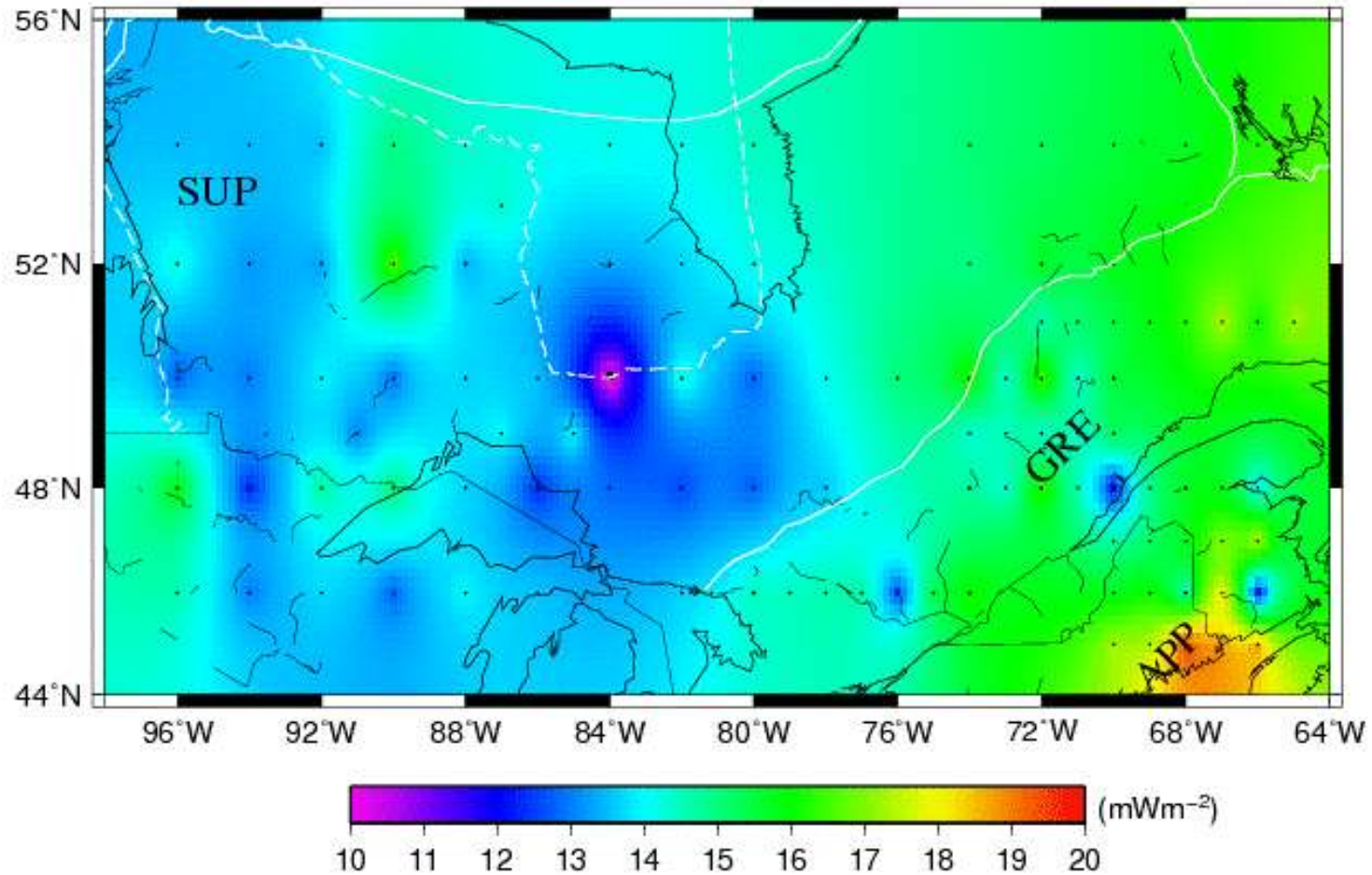


74 W 54 N  
Location East of James Bay

Mean surface heat flux 28 mWm<sup>-2</sup>

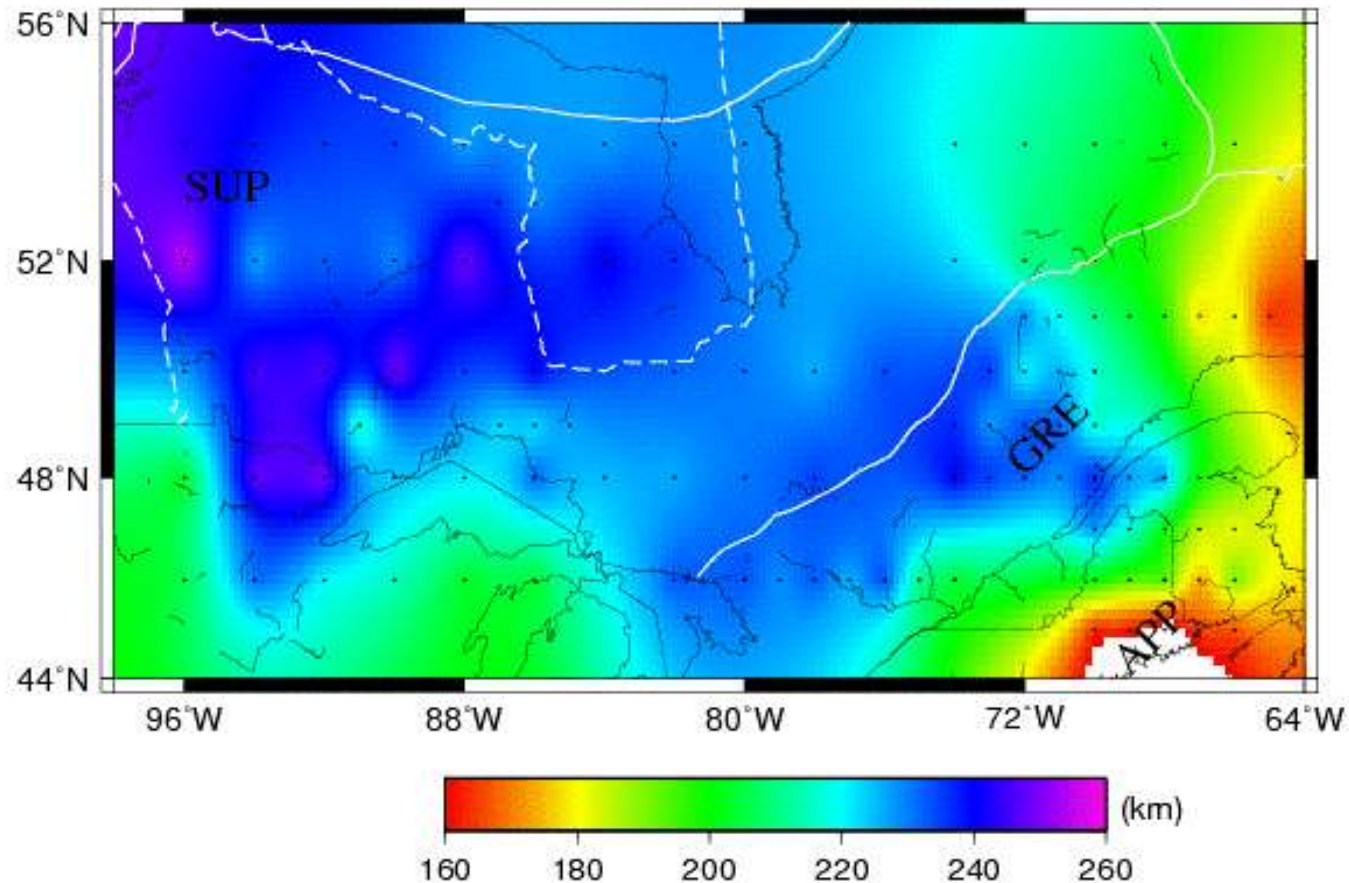
# Moho heat flux from inversion

$$\mu = 14.5 \quad \sigma = 1.3 \text{ mWm}^{-2}$$



# Lithospheric thickness from inversion

$$\mu = 225 \quad \sigma = 15 \text{ km}$$



# Summary

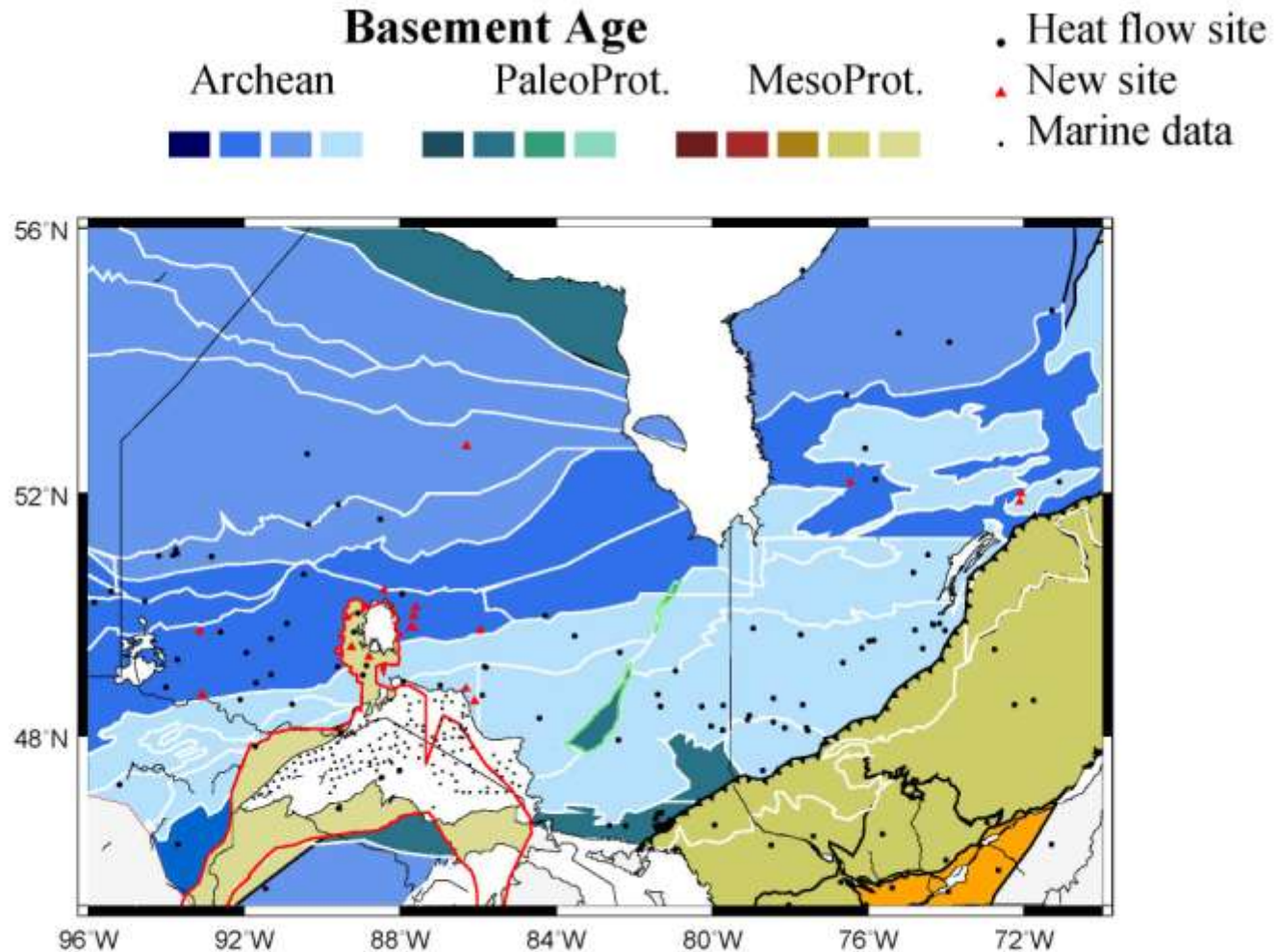
	mean Qm mW m <sup>-2</sup>	std Qm mW m <sup>-2</sup>	mean LT (km)	std LT (km)
SUPERIOR	14.7	1.3	225	16
GRENVILLE	14.9	1.3	220	19

# Conclusions

- Combining surface heat flux and shear wave velocity profiles works
- Moho heat flux and lithospheric thickness very strongly constrained
- Central value for heat flux at Moho  $Q_m$   $15\text{mWm}^{-2}$  in southeastern Canadian Shield. Consistent with previous estimates.
- Variations in  $Q_m$  small ( $\pm 2\text{mW m}^{-2}$ ) in southeastern Canadian Shield.
- Lithospheric thickness  $220 \pm 25$  km
- Lowest  $Q_m$  and thickest lithosphere do not always coincide with lowest surface heat flux.
- Is it EOS?
  - THO?
  - East West James Bay?

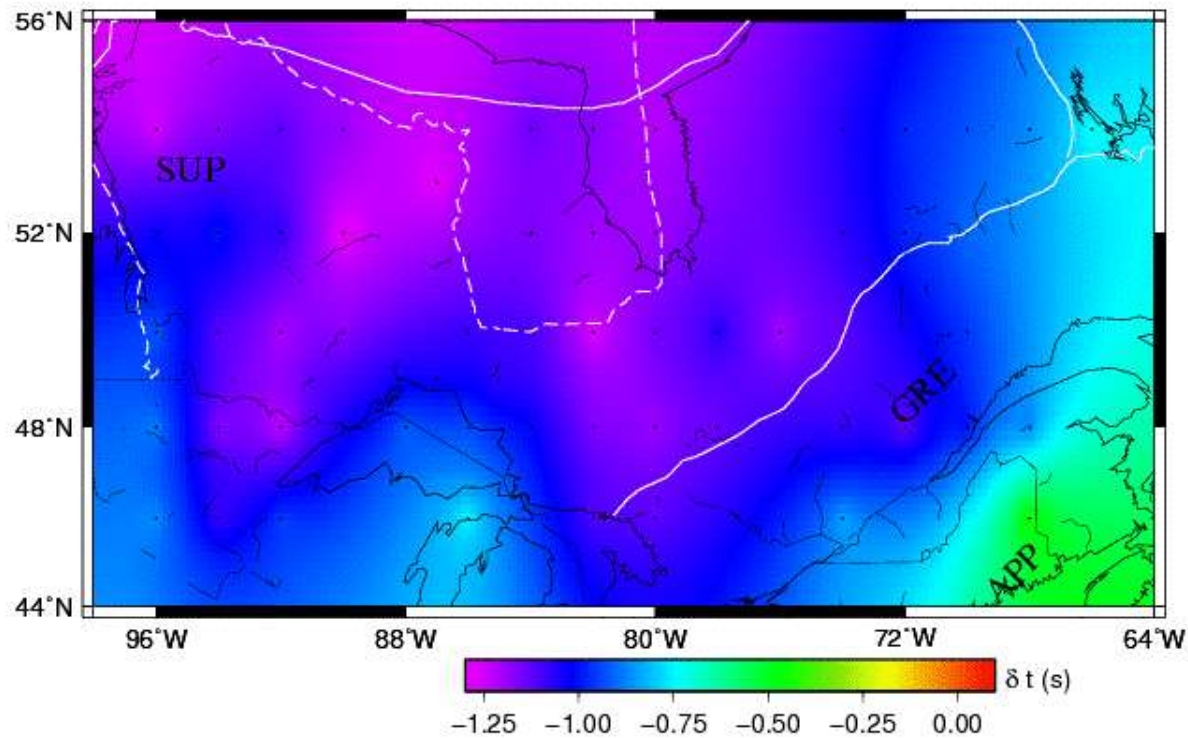


# Southwestern Superior Province final assembly ~2.65 Ga

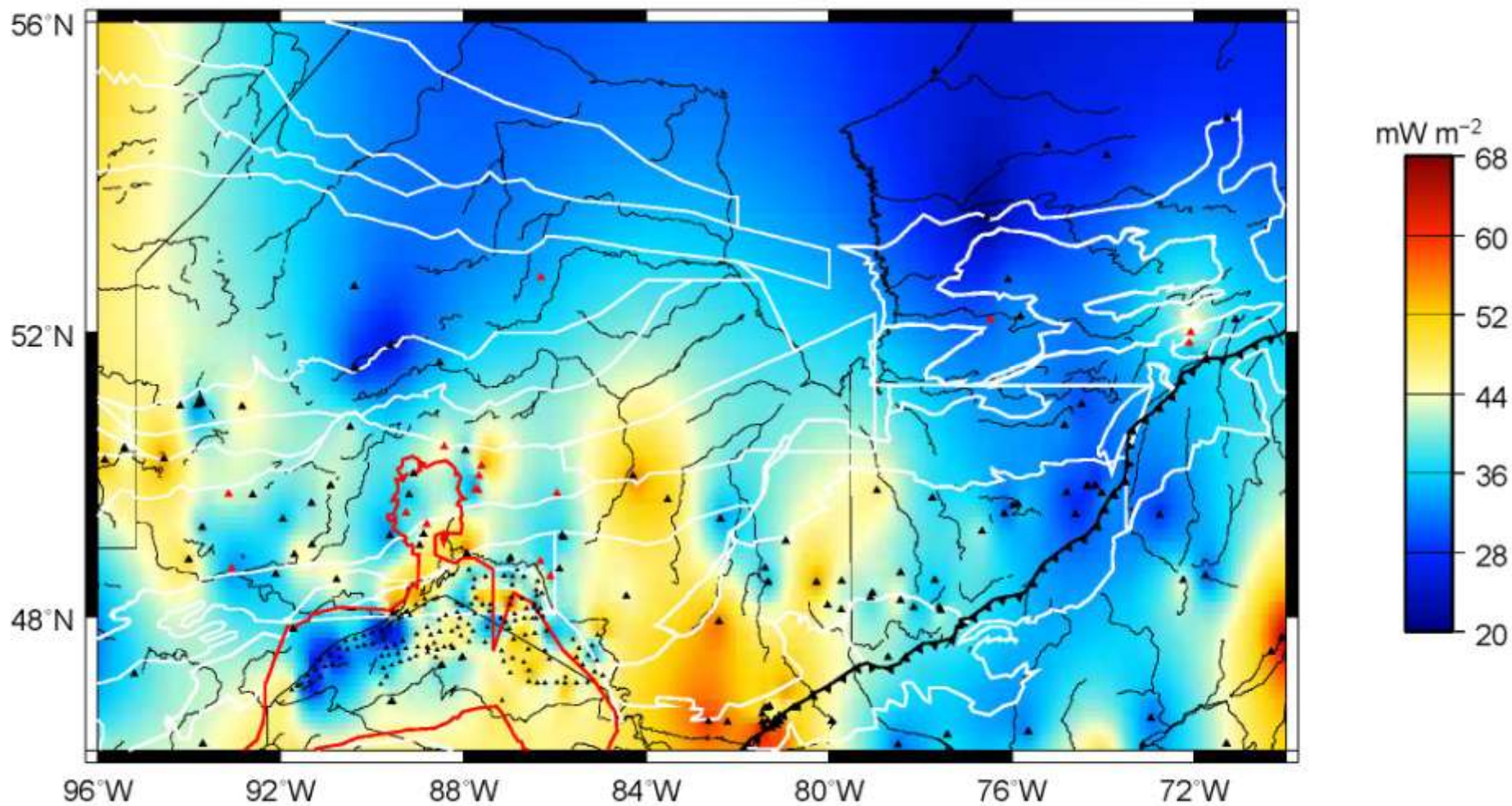


# Travel time delays (60-260km) from inversion

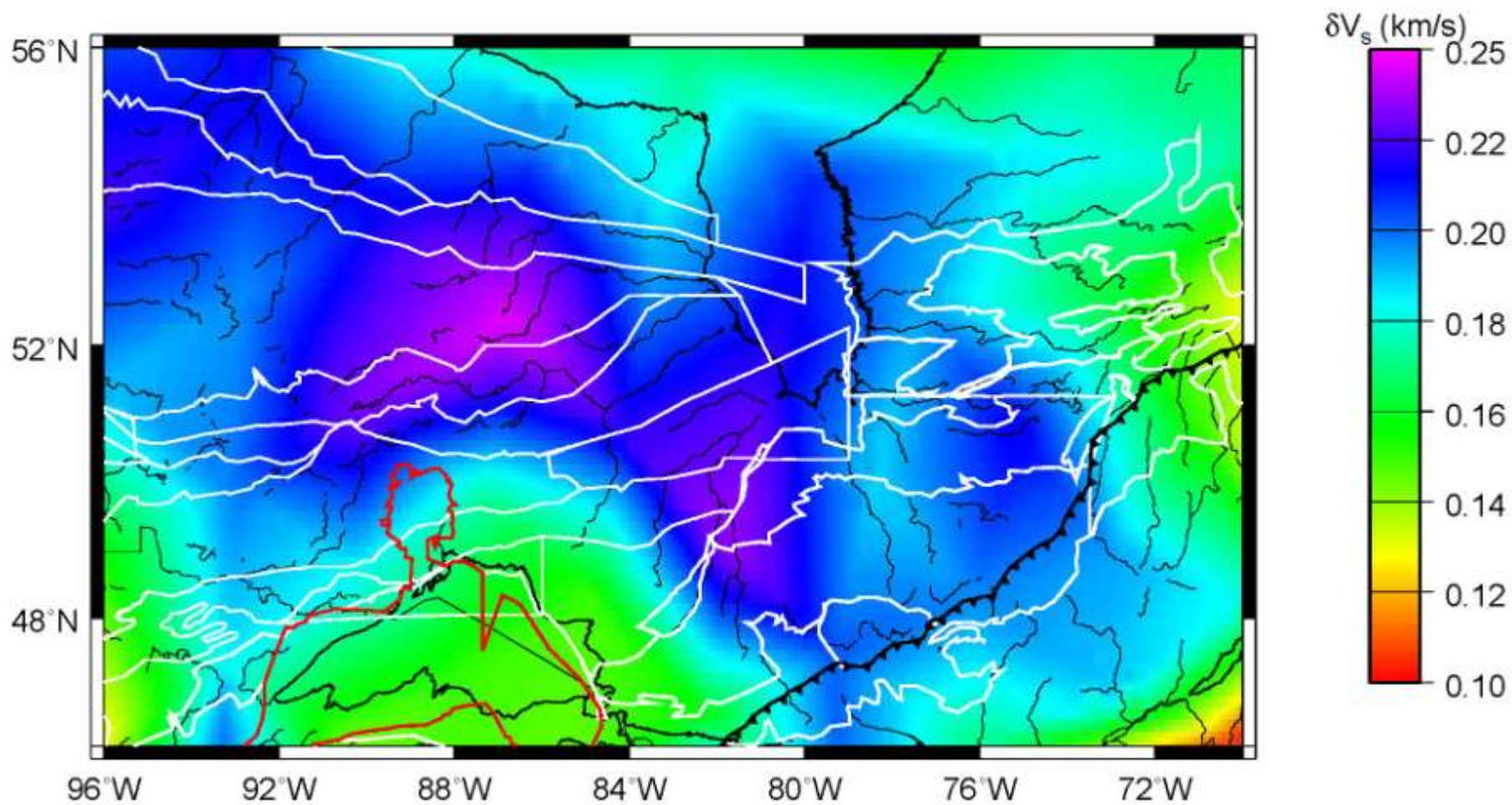
Travel time delay (Archon model)



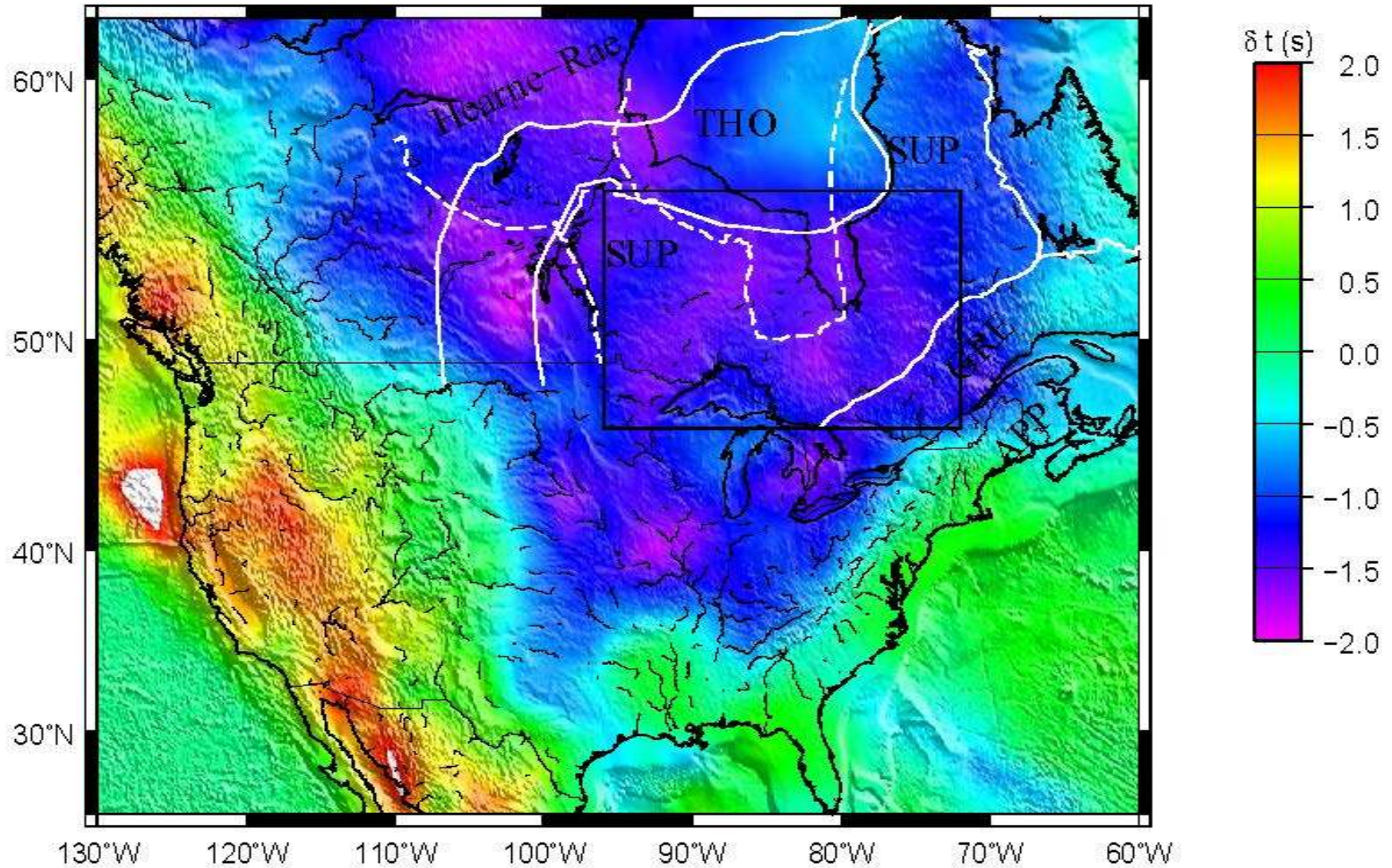
# Heat flux



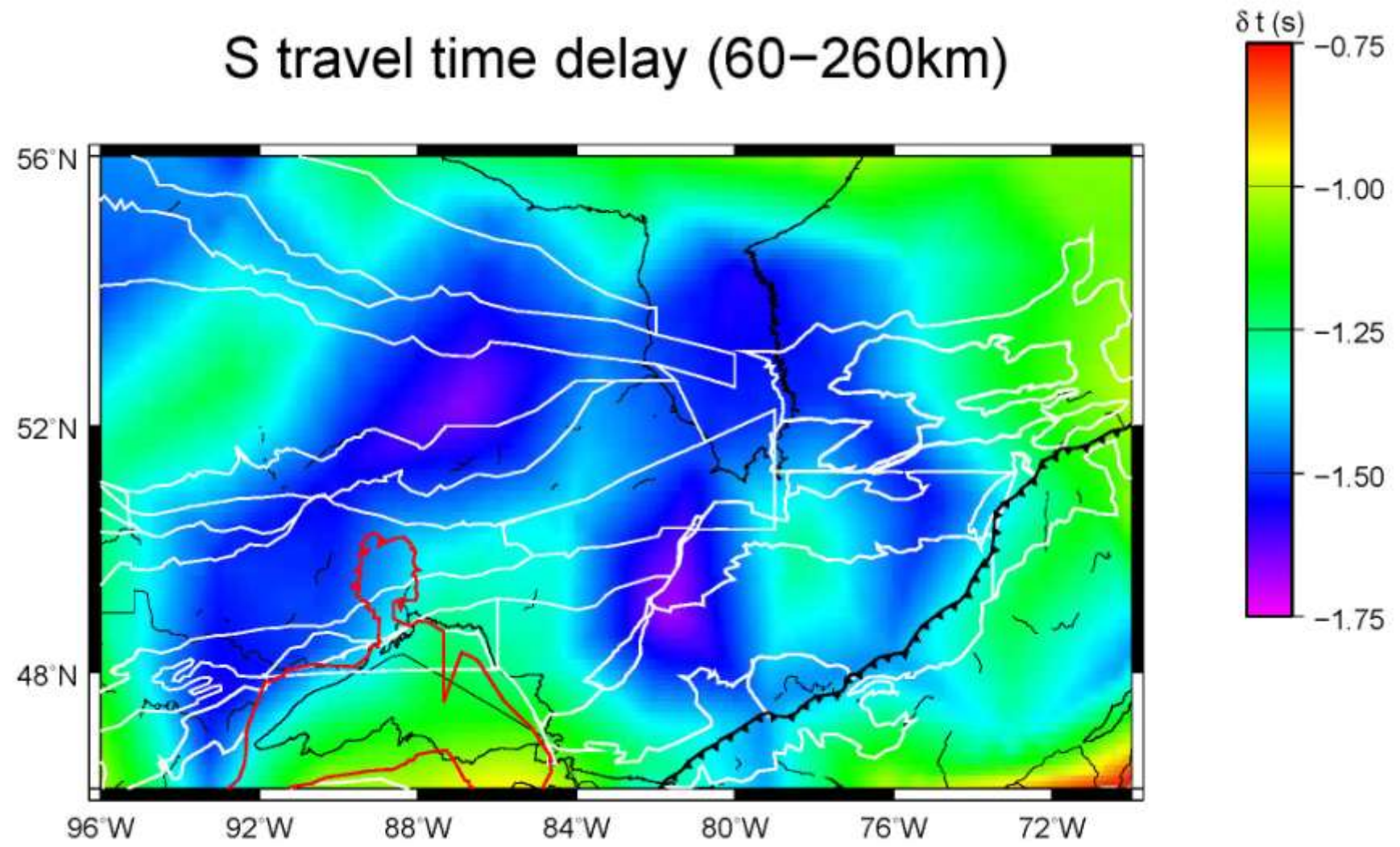
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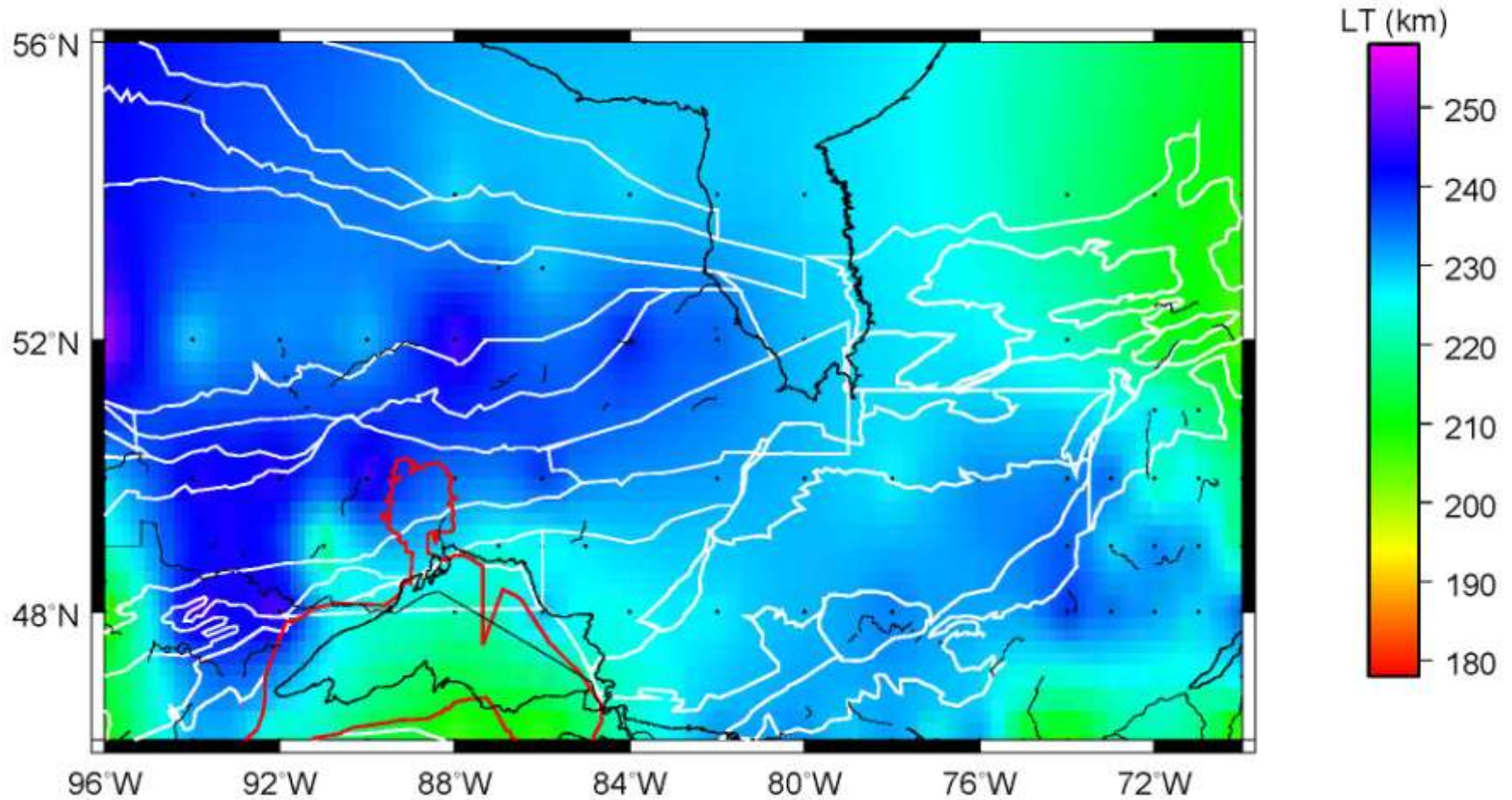
# S travel time delay (60–260km)



# S travel time delay (60–260km)



# Lithospheric thickness



Not constrained south of 48N

## Travel time delay (Archon model)

