

GyPSuM: A Joint Seismic-geodynamic Tomographic Image of the Mantle with Minimal Compositional Heterogeneity



LLNL-PRES-598995

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This work performed under the auspices of the U.S. Department of Energy by Lawrence Livermore National Laboratory under Contract DE-AC52-07NA27344.



- ❑ Motivation and our previous work
- ❑ GyPSuM Earth model development
- ❑ Observations from GyPSuM
- ❑ Summary



- ❑ Combining multiple data types yields more robust seismic images
 - Reduce non-uniqueness
 - Predict heterogeneity where certain seismic constraints are weak or non-existent

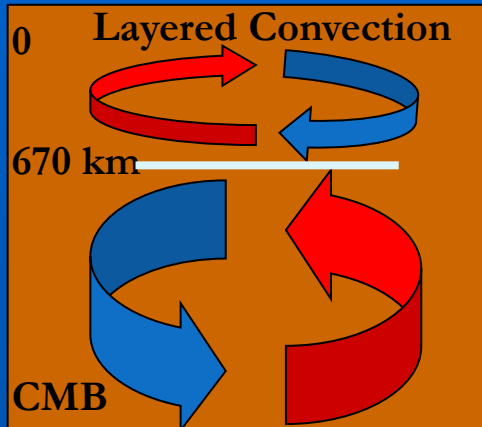
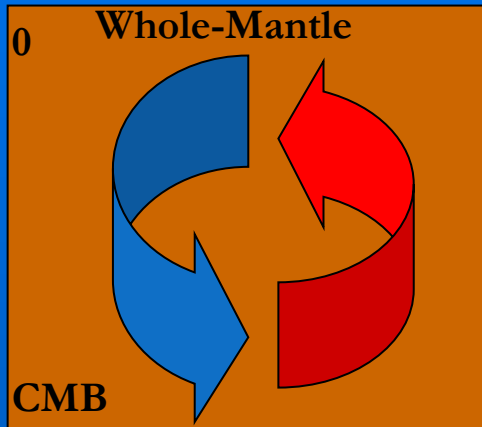
- ❑ Need detailed density structure for flow modeling
 - Scaling a model derived with only seismic data is inadequate
 - Density heterogeneity should be consistent with geodynamic observations
 - Solve for density directly and simultaneously

- ❑ Evaluate the relative behavior of mantle properties

Joint seismic-geodynamic mantle imaging



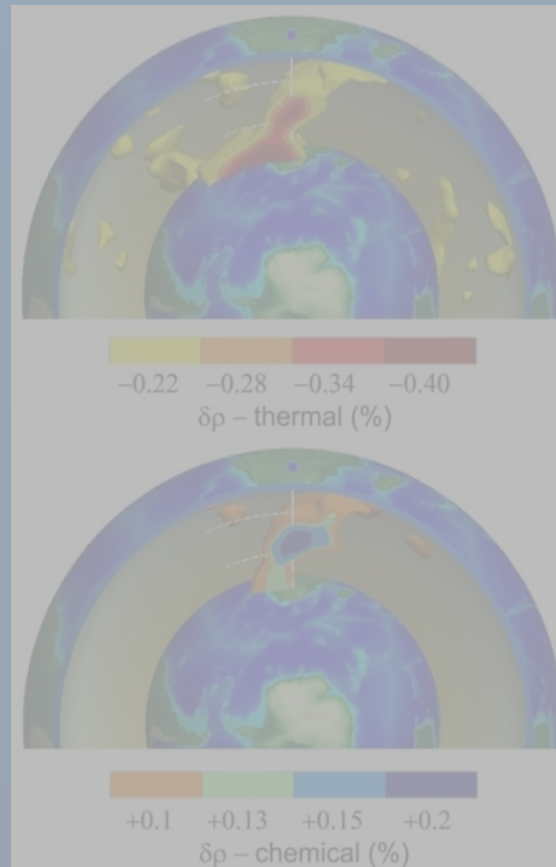
Simmons, Forte & Grand [2006]



Tested mantle flow hypotheses

- Effects the forward system
- Whole-mantle style prevailed

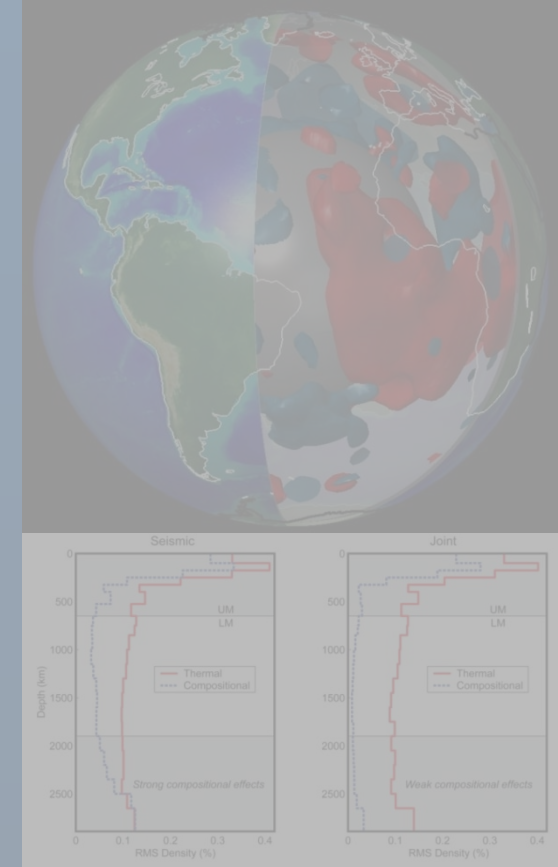
Simmons, Forte & Grand [2007]



Imaged mantle density

- Detailed image of the intrinsically dense material in the African superplume

Simmons, Forte & Grand [2009]



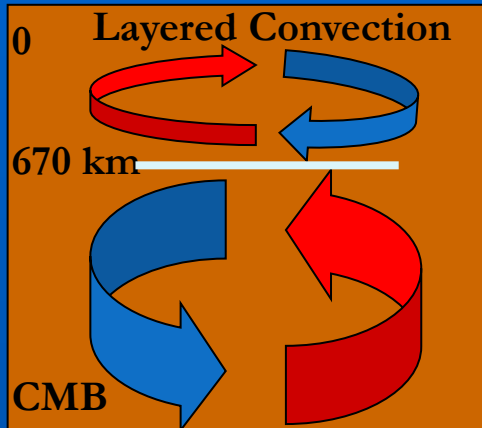
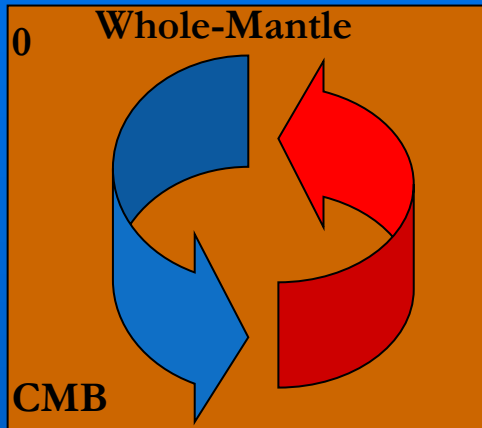
Tested the thermal hypothesis

- Evaluated the relative effects of temperature and composition
- Minimize the apparent role of composition

Joint seismic-geodynamic mantle imaging



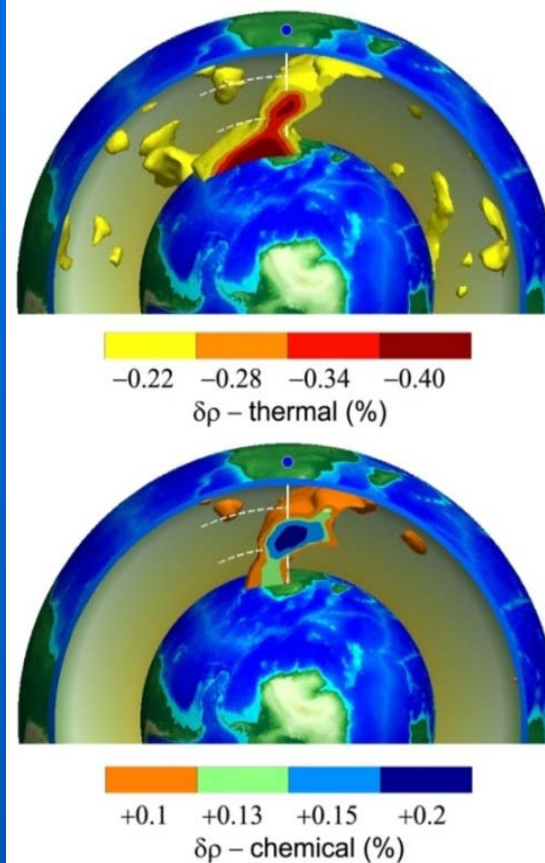
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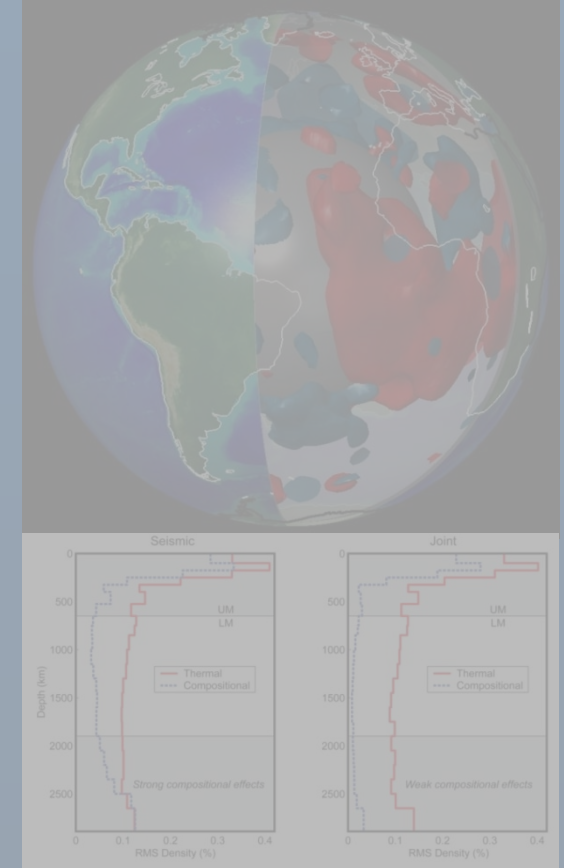
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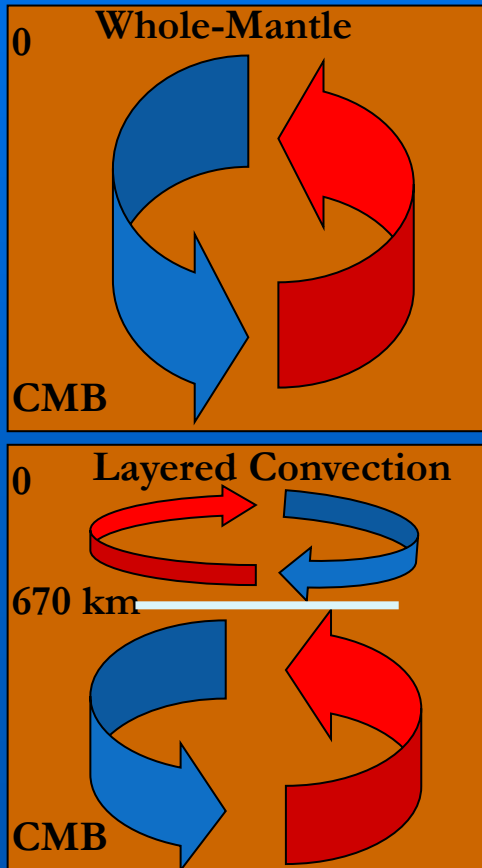


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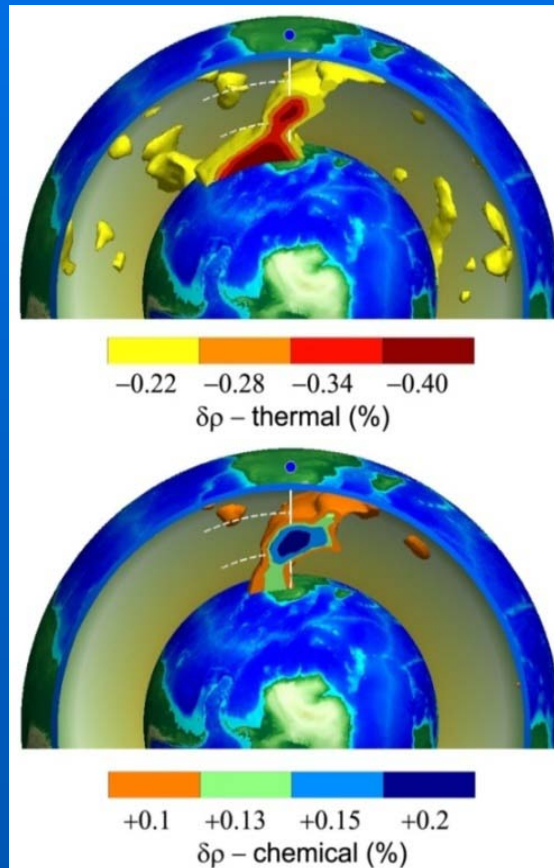
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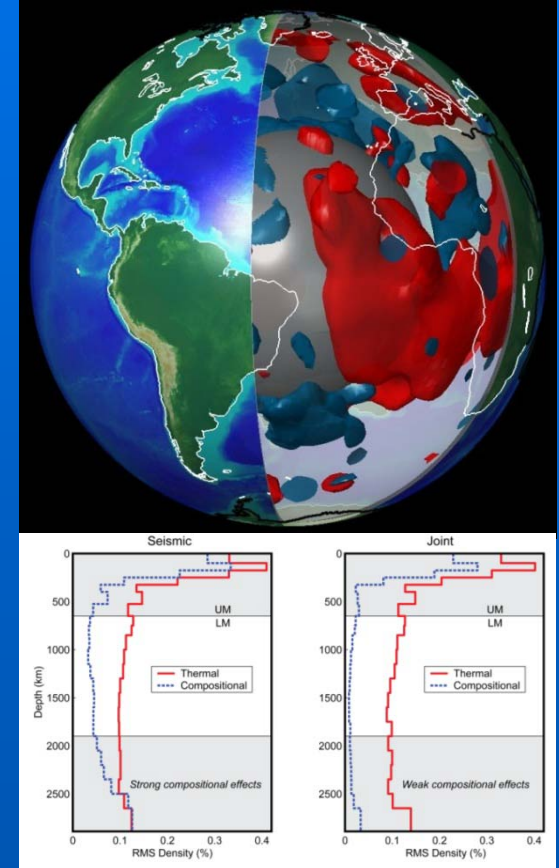
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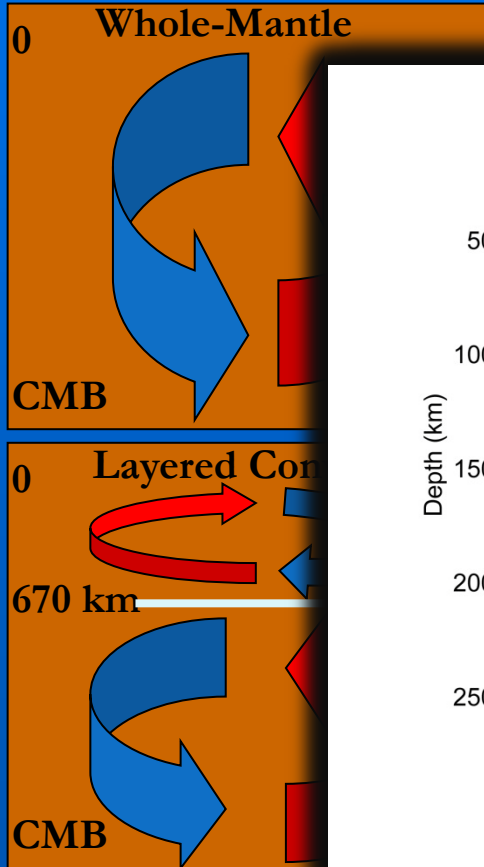
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Joint seismic-geodynamic mantle imaging



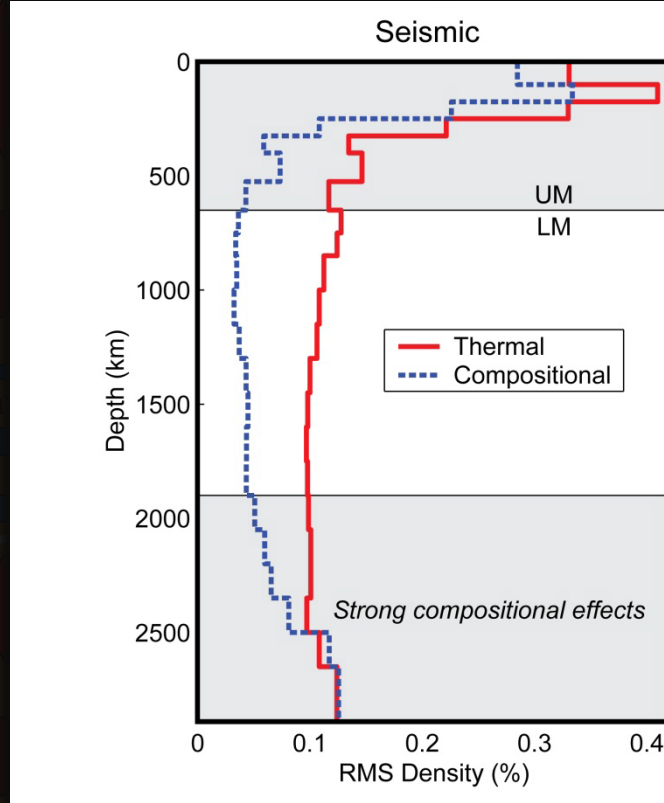
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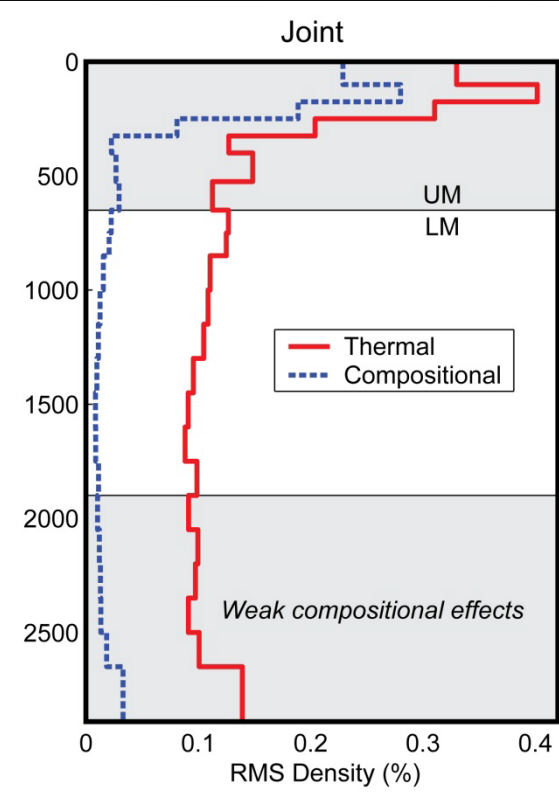
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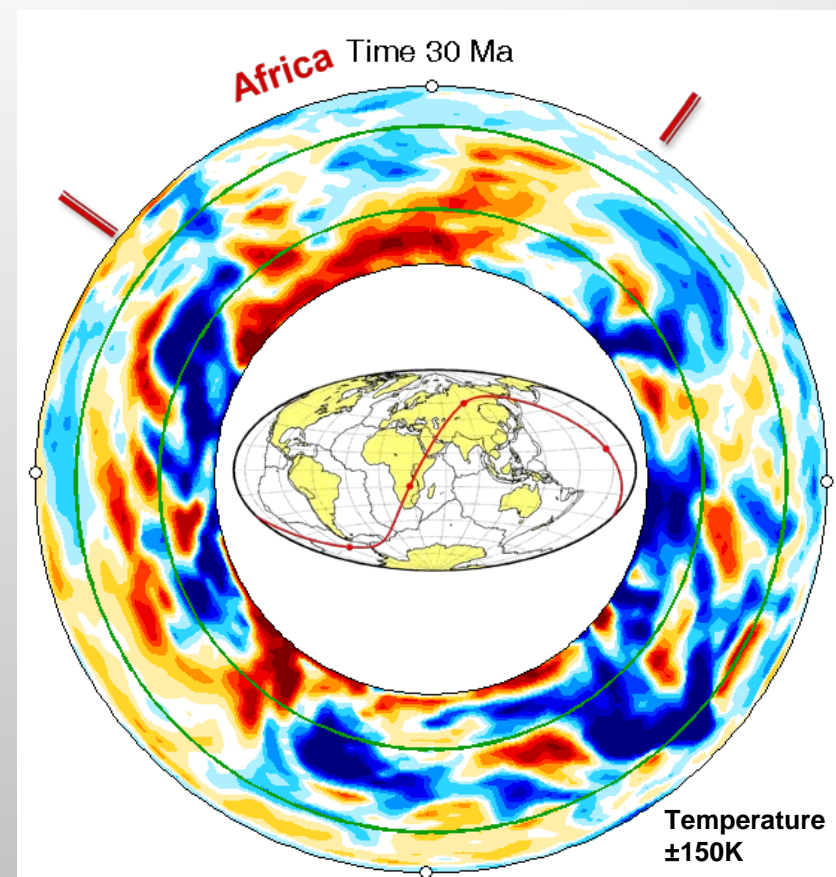
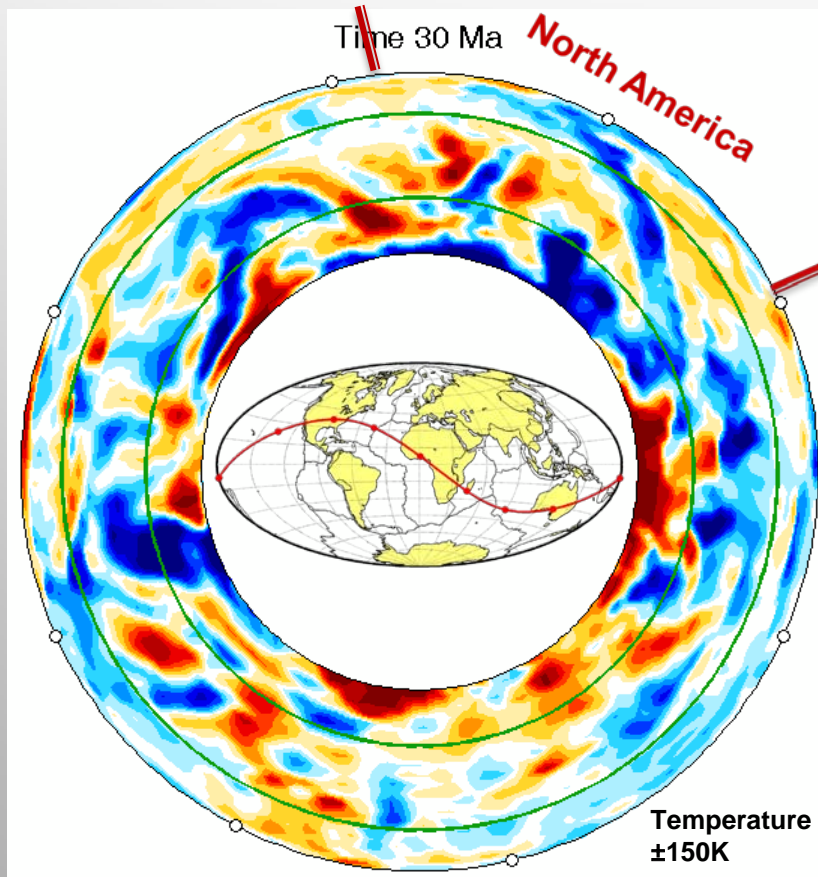
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Joint models are used for flow simulations



Geodynamics/Tectonophysics Collaborators:

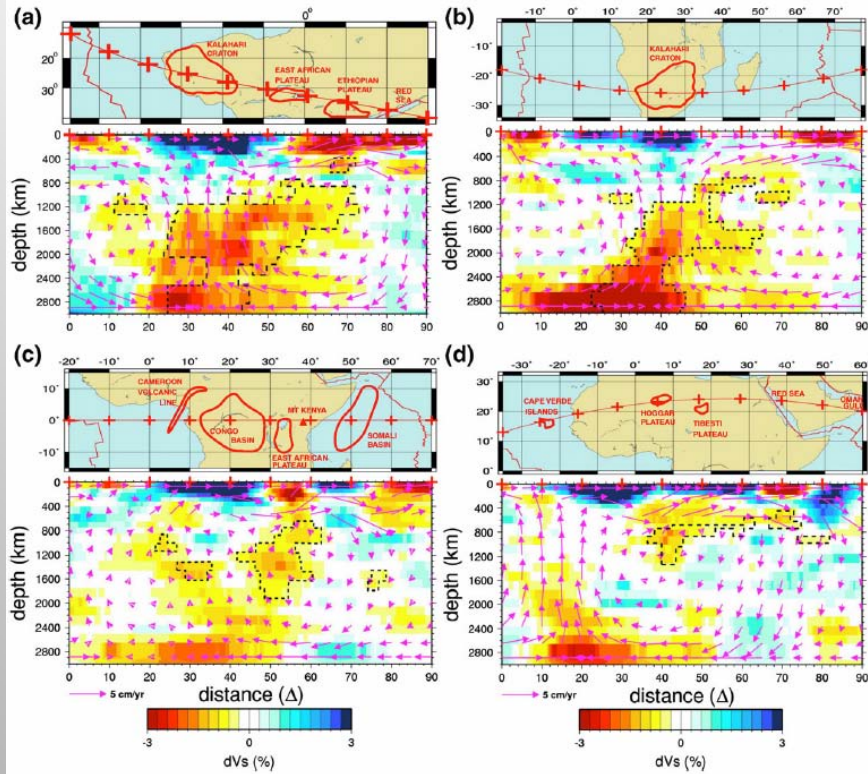
Alessandro Forte *Université du Québec*, **Robert Moucha** *Syracuse University*, **Jerry Mitrovica** *Harvard*,
David Rowley *University of Chicago*, **Sandrine Quéré** *Utrecht University*

Linking the deep mantle to the surface



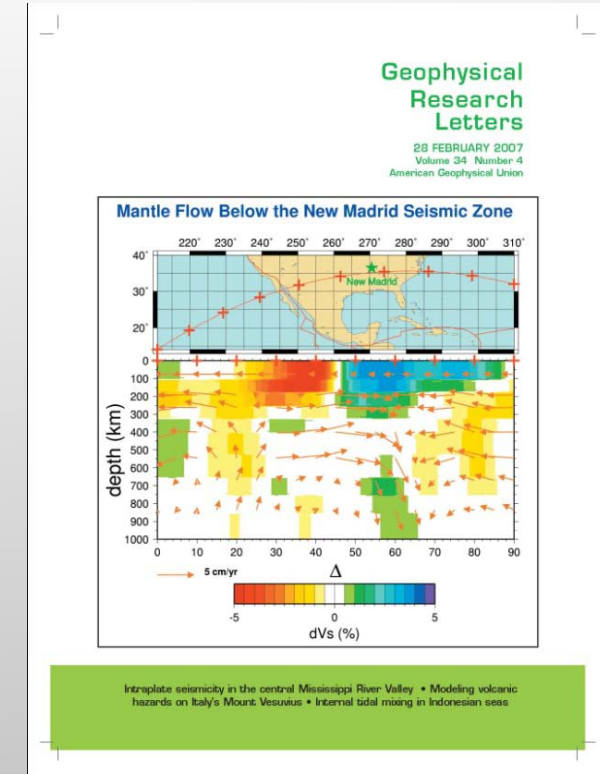
Mantle flow beneath Africa and surface manifestations

Forte et al. [2010b] in EPSL



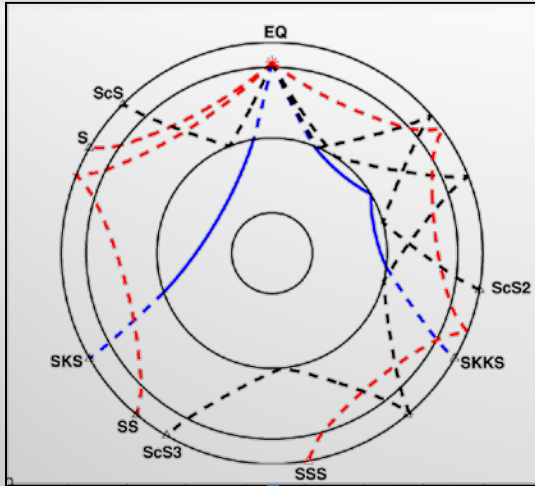
Possible contribution to New Madrid EQs

Forte et al. [2007] in GRL



- ❑ Uplift of the Colorado Plateau *Moucha et al. [2008a] in Geology, Moucha et al. [2009] in GRL*
- ❑ Instability of the “stable” Eastern US *Moucha et al. [2008b] in EPSL, Rowley et al. [2012, submitted]*
- ❑ Global plate decelerations *Forte et al. [2009] in GRL*
- ❑ Deep-mantle contributions to North American surface dynamics *Forte et al. [2010a] in Tectonophysics*
- ❑ African topography driven by mantle convection *Moucha and Forte [2012] in Nature Geoscience*

GyPSuM: What is it?



Geodynamic Observations

Global free-air gravity (EGM96)
Tectonic plate motions (NUVEL-1)
Dynamic Topography (*Forte & Perry 2000*)
CMB ellipticity (*Mathews et al. 2002*)

G = Geodynamic
y
P = P waves
S = S waves
u
M = Mineral physics

Density

S-wave Data

S, ScS, SKS, SKKS,
sS, sScS, and
multiples
*Grand (2002);
Simmons et al. (2006)*

S-wave
Velocity

Mineral Physics Constraints

$$R_{\rho/S} = \frac{d \ln \rho}{d \ln V_S} \quad R_{P/S} = \frac{d \ln V_P}{d \ln V_S}$$

Karato & Karki (2001)
Cammarano et al. (2003)

P-wave
Velocity

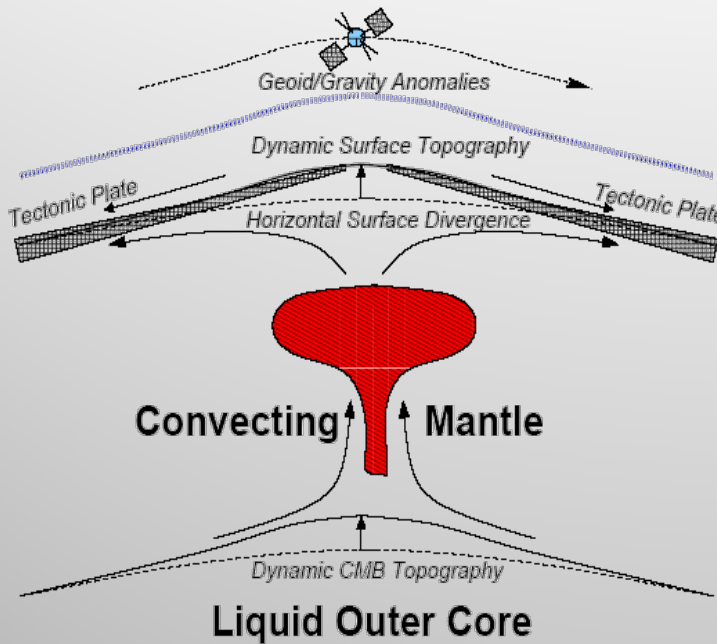
P-wave Data

~626,000 teleseismic
summary rays
Antolik et al. (2003)

3-D Global Model

S-wave velocity
P-wave velocity
Density

$$\begin{bmatrix} \mathbf{L}_S \\ \mathbf{L}_P(R_{P/S}) \\ \mathbf{G}(R_{\rho/S}) \end{bmatrix} \Delta \mathbf{m}_S = \begin{bmatrix} \mathbf{r}_S \\ \mathbf{r}_P \\ \mathbf{g} \end{bmatrix}$$



Geodynamic Observations

- Density anomalies drive flow
- Fields are dynamically coupled
- Numerical description in:
 - Richards & Hager [1984]*
 - Ricard et al. [1984]*
 - Forte & Peltier [1987]*

Observation

Spherical harmonic component

Up to degree = 16

Density-velocity conversion

$$R_{\rho/S} = \frac{d \ln \rho}{d \ln V_s}$$

Based on mineral physics

$$s_l^m = \int_b^a K_l(\eta(r)/\eta_0; r) R_{\rho/S} \left(\frac{\delta V_s}{V_s} \right)_l^m d r$$

Viscous Flow Response

Response of internal density loads on the observation

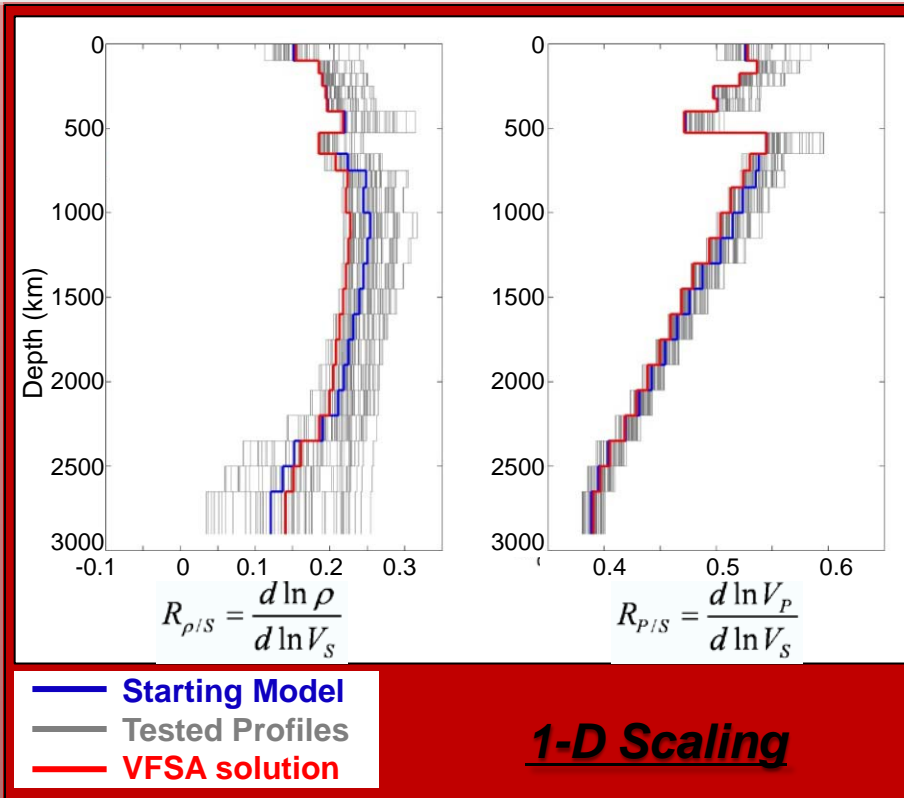
Seismic velocity

S-wave speed in this case

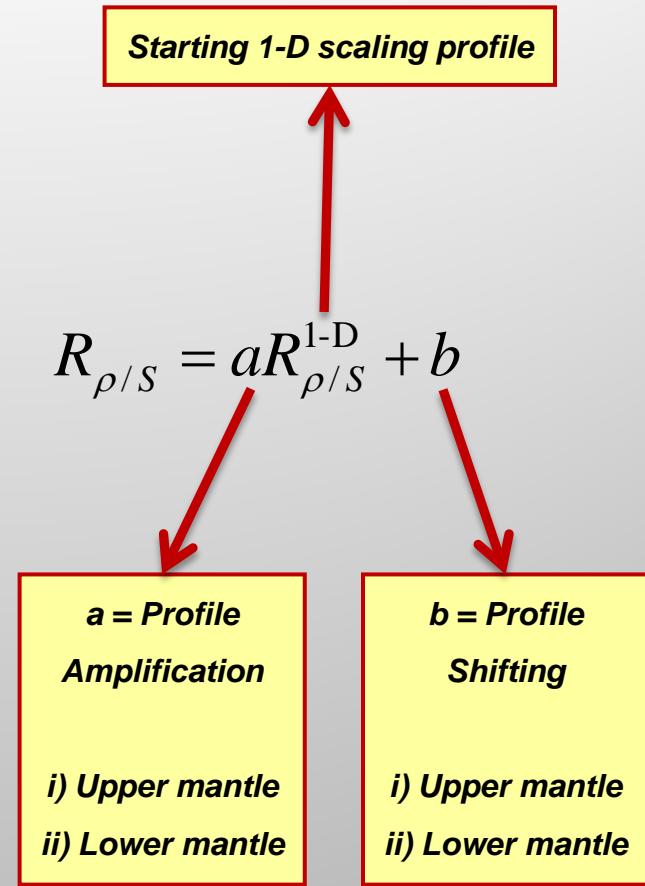
$$\Rightarrow R_{P/S} = \frac{(\delta_s - 1)}{2} R_{\rho/S} (1 - \gamma) + \gamma$$

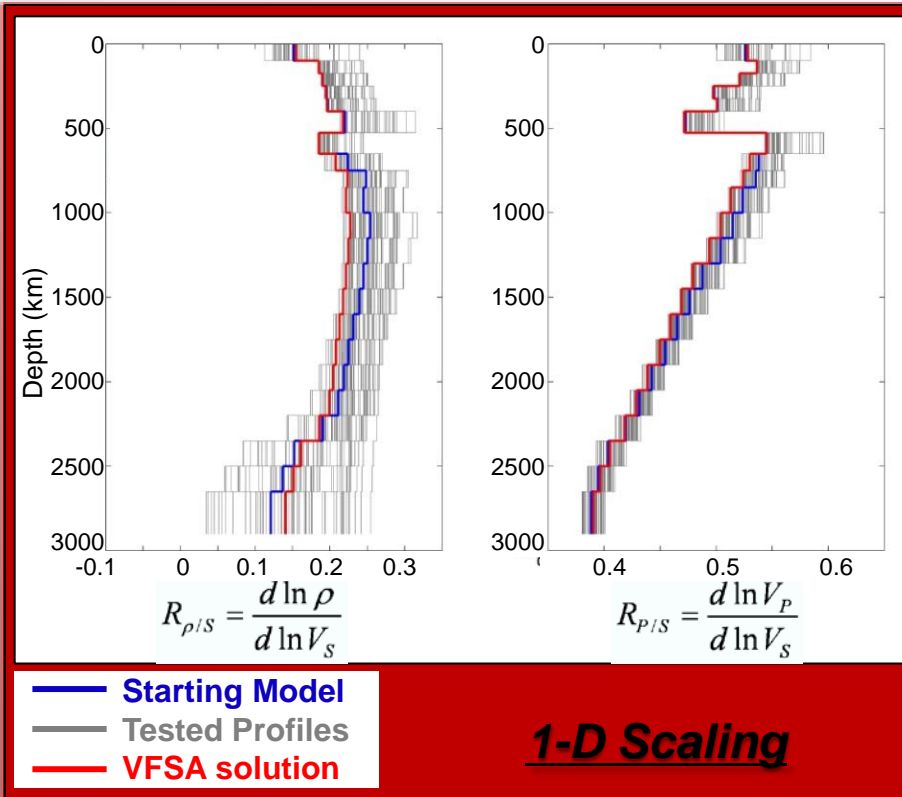
$$\gamma \equiv \frac{4 V_s^2}{3 V_P^2}$$

$$\delta_s \equiv \frac{\delta \log K}{\delta \log \rho}$$

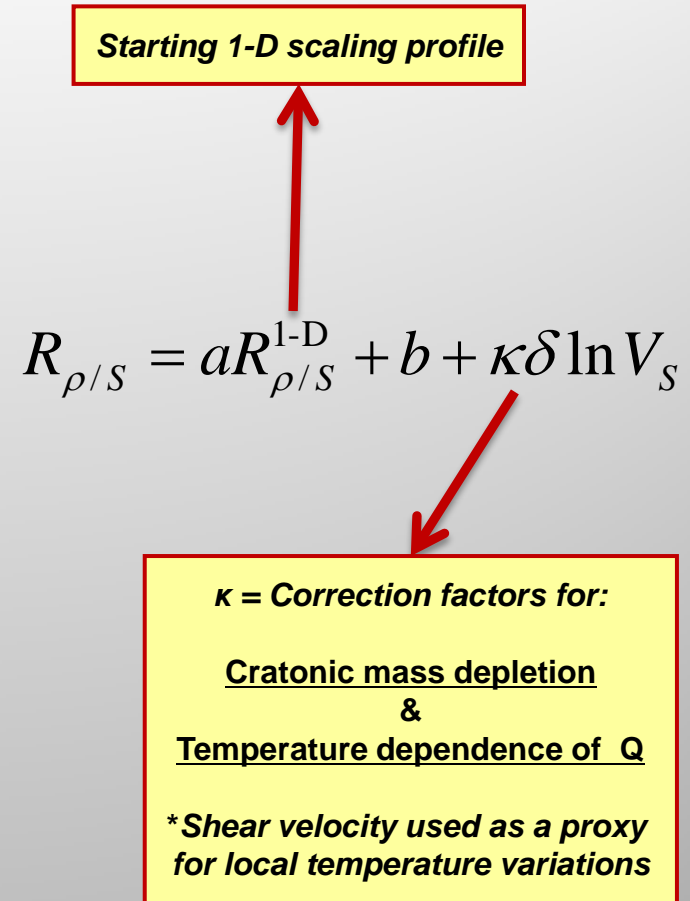


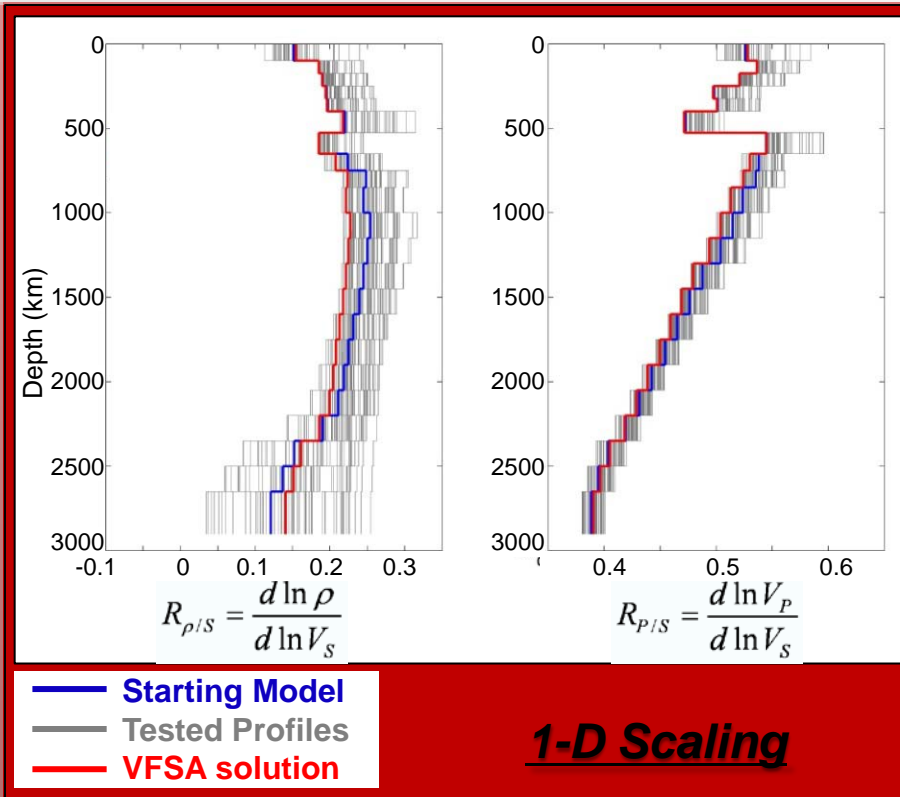
- ❑ Range defined by mineral physics
 - *Expected thermal values*
- ❑ Simulated annealing (VFSA)
- ❑ Full joint inversion performed with each update



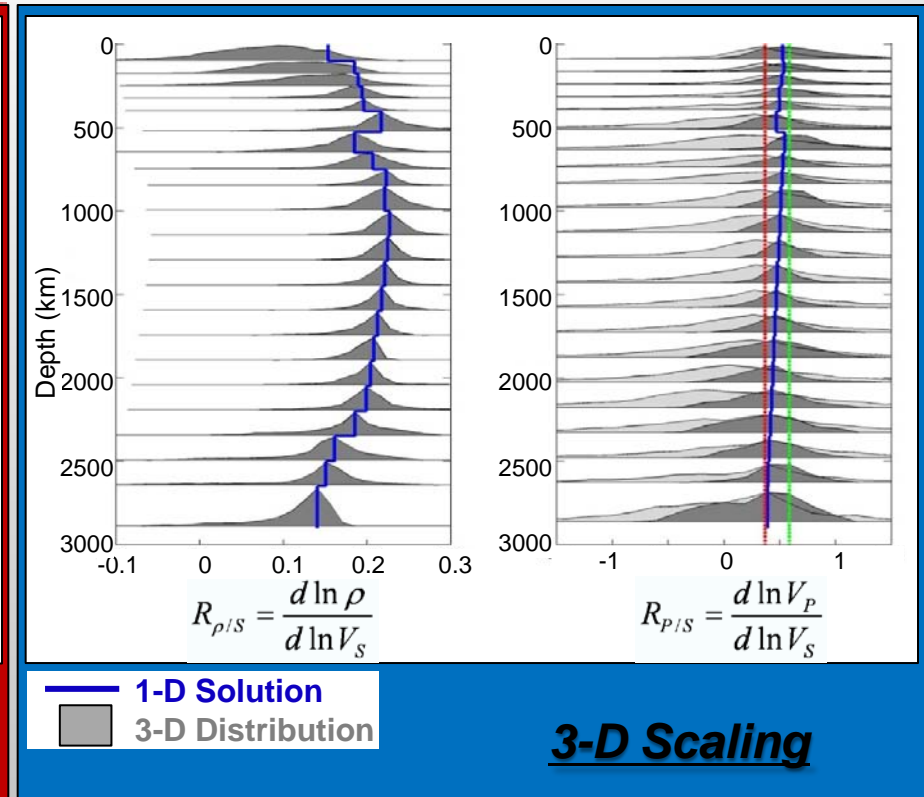


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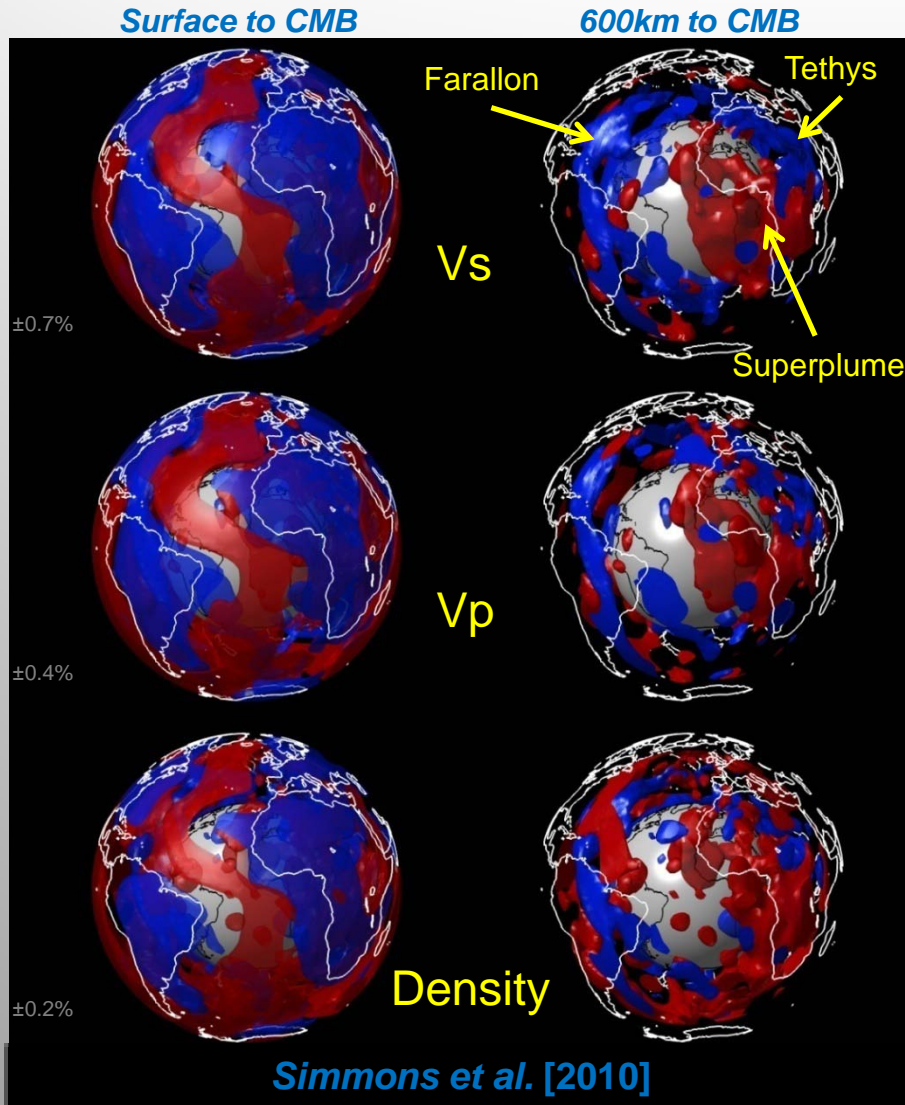




- Range defined by mineral physics
 - Expected thermal values
- Simulated annealing (VFSA)
- Full joint inversion performed with each update



- Scaling factors allowed to diverge from thermal
- Non-linear inversion process
- Allows reconciliation of all seismic/geodynamic data
- Produces model most closely resembling thermal

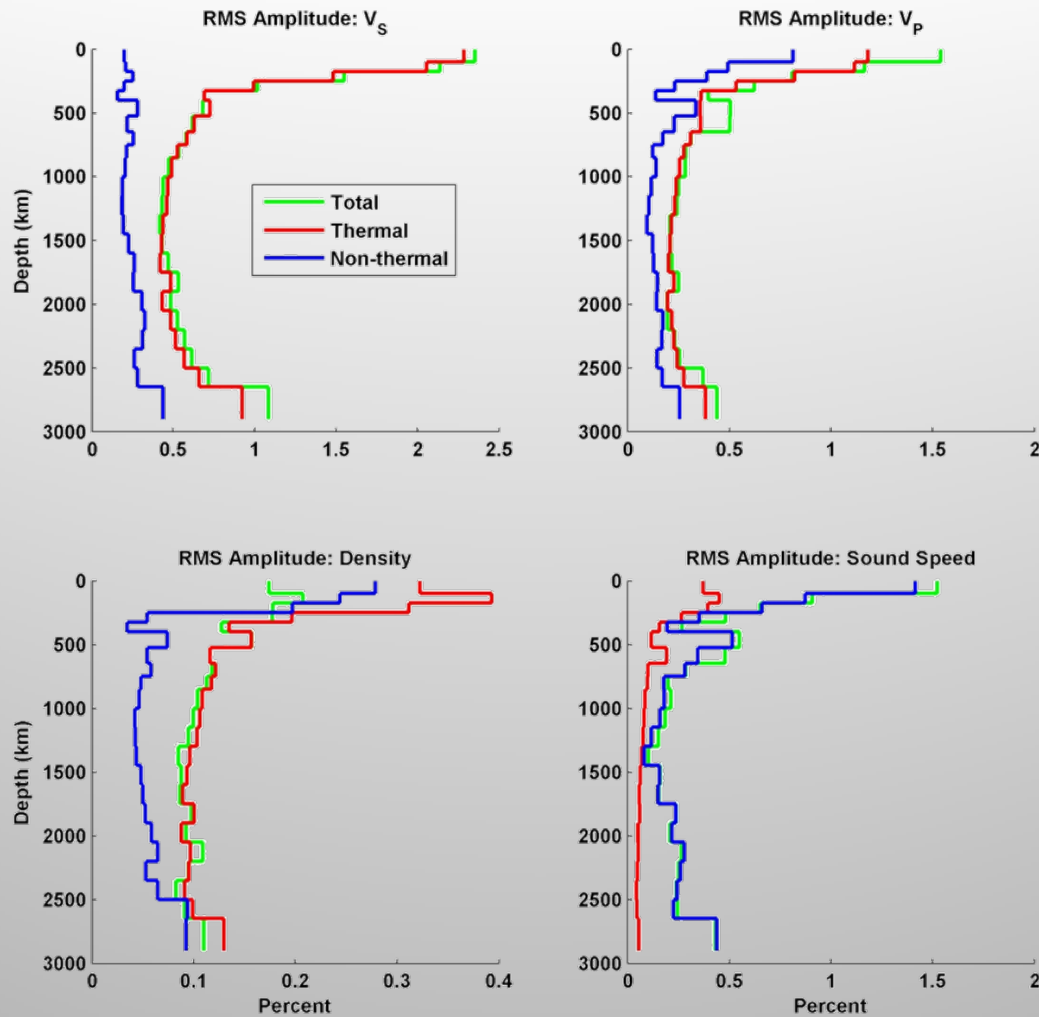


Joint model of mantle S-wave velocity, P-wave velocity and density

- ❑ Built with the hypothesis that temperature variations dominate
- ❑ The role of composition is minimized
- ❑ Detailed, (more) dynamically consistent density model of the mantle

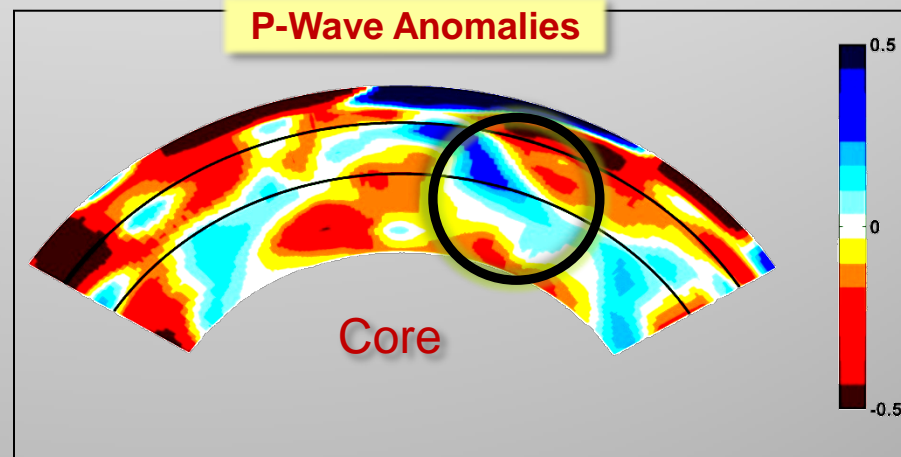
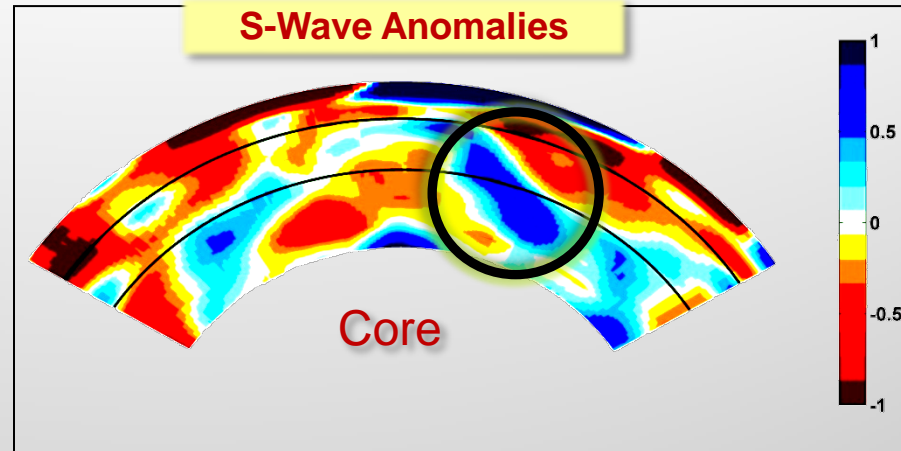
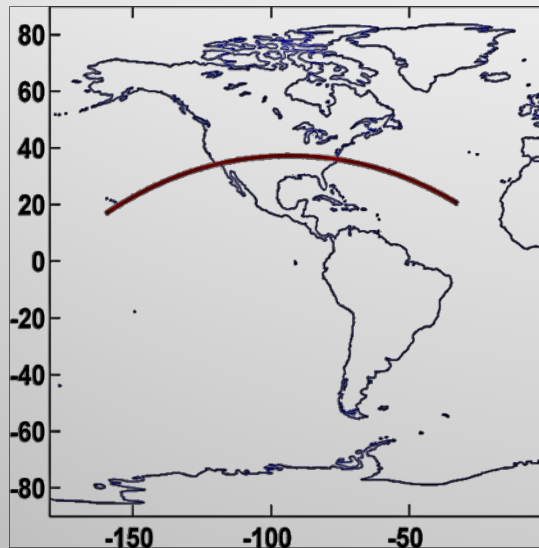
S-wave Data	P-wave Data	Free-air Gravity	Plate Div.	Dynamic Topo.	CMB Excess Ellip.
93%	31%	88%	99%	72%	100% (0.4 km)
	Variance: 2.6s → 1.8s				

GyPSUM: Temperature is the primary contributor to P, S, and Density



**Thermal and Non-thermal components may be constructive or destructive.*

GyPSUM: Systematic de-correlation of S- and P-velocity within ancient subducted slabs ... compositional variations?



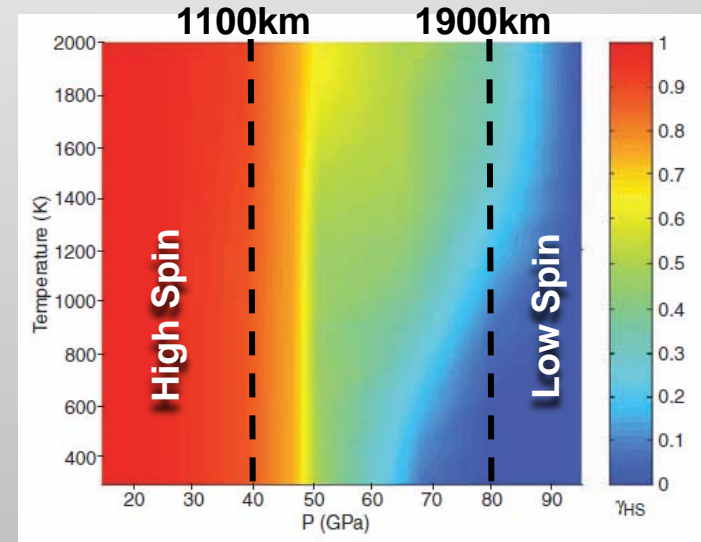
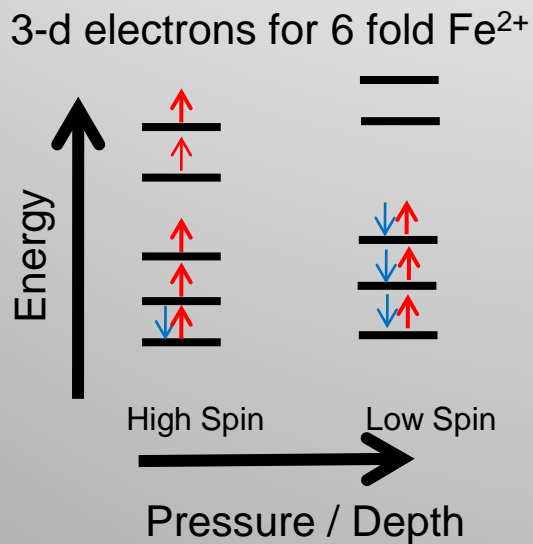


Electronic Spin Transitions:

➤ Fe^{2+} undergoes a pressure induced transition from a high spin state to a low spin state... *under mid-mantle P/T conditions*

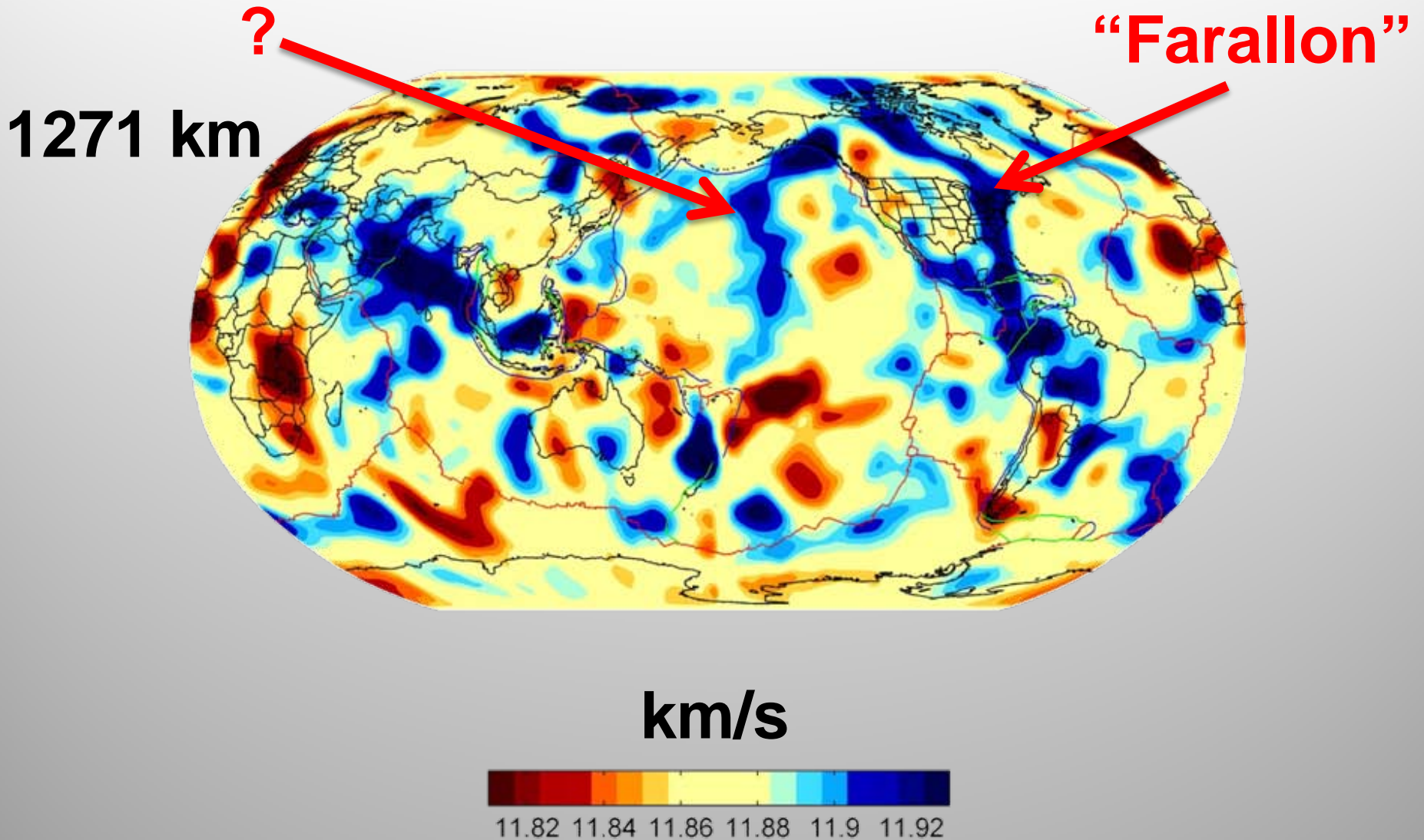
Affects major mantle minerals $(\text{Mg,Fe})\text{SiO}_3$ and $(\text{Mg,Fe})\text{O}$:

- Elastic properties
- Density
- Iron partitioning
- Thermal conductivity



From *Lin et al. (2007)*

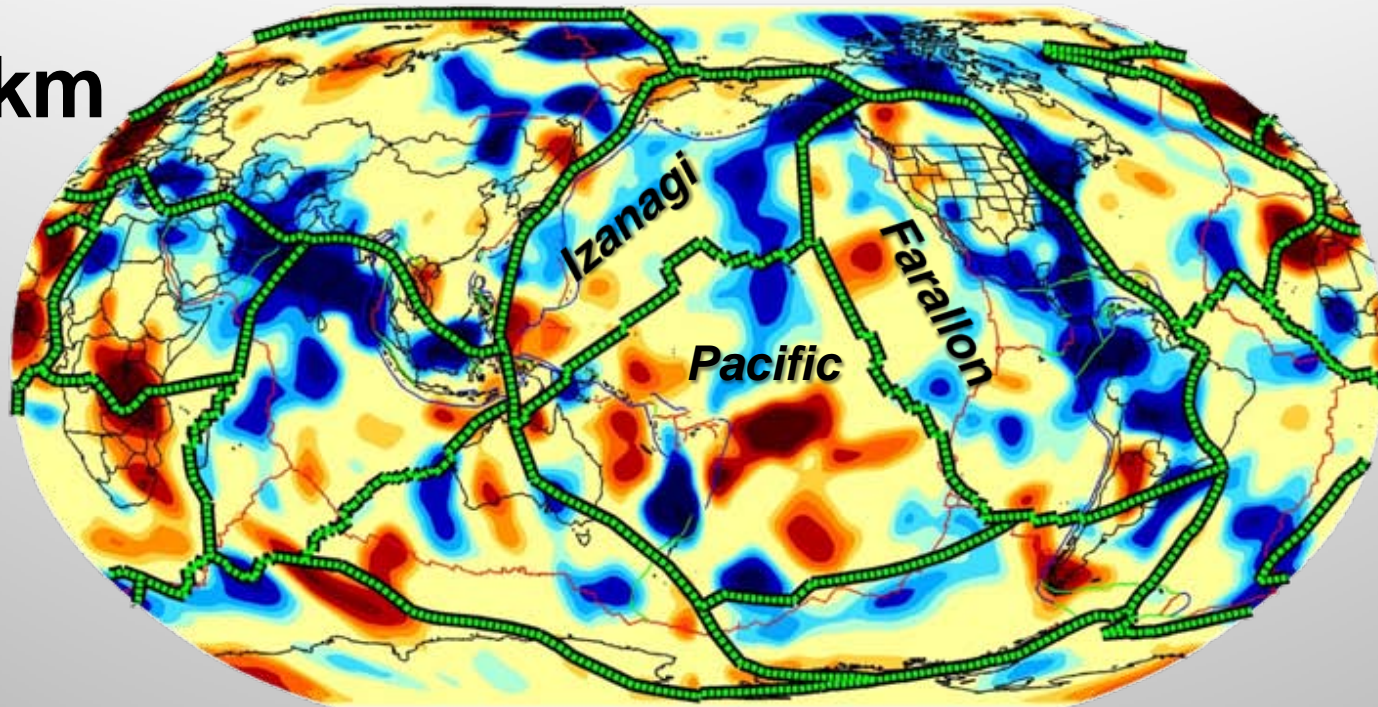
GyPSuM: An ancient slab remnant beneath the Pacific Ocean?





Paleo-plate boundaries @ 100 Ma: *Torsvik, Steinberger, Gurnis & Gaina [2010]*

1271 km



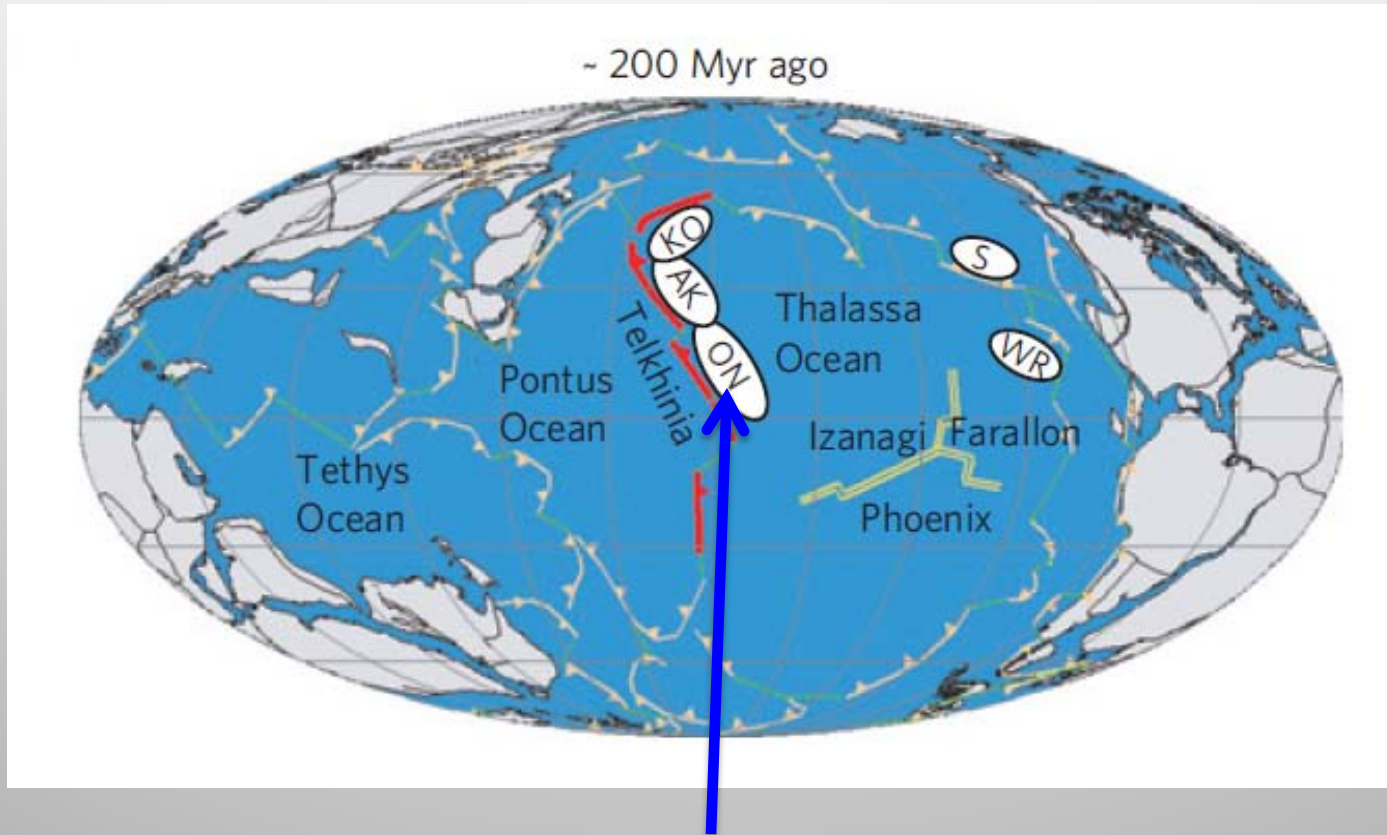
km/s



11.82 11.84 11.86 11.88 11.9 11.92

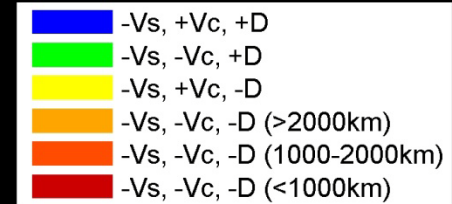
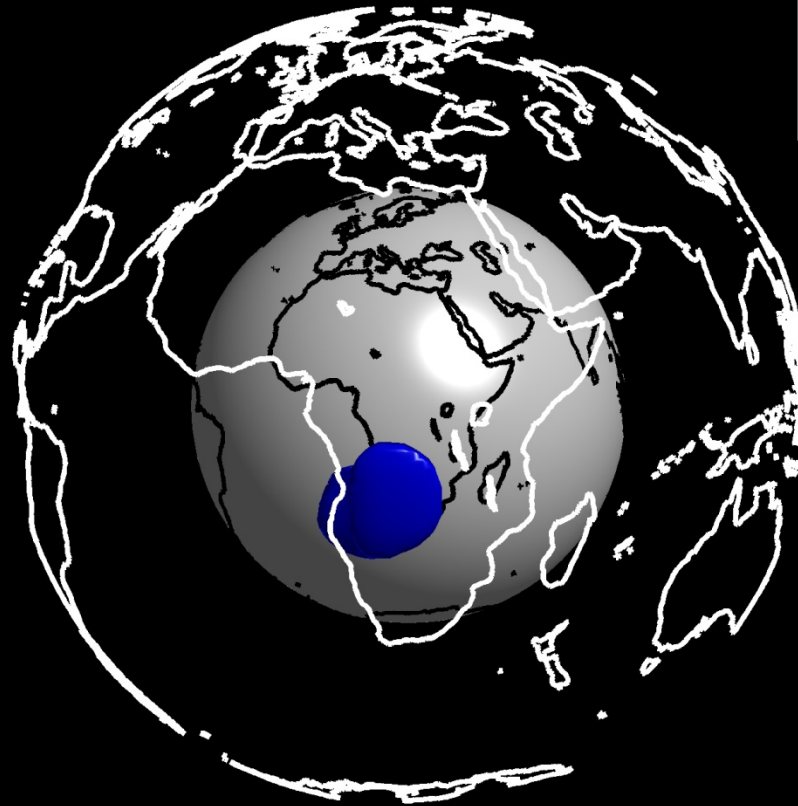


Paleo-plate boundaries @ 200 Ma: *van der Meer et al. [2012]*



Extinct intra-oceanic volcanic arcs formed above ancient subduction zones

***Locations consistent with Paleomagnetism and Biostratigraphy**

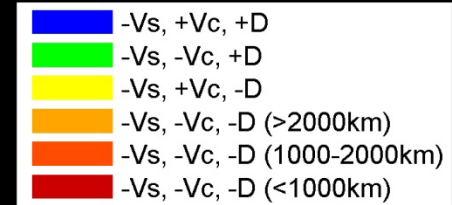
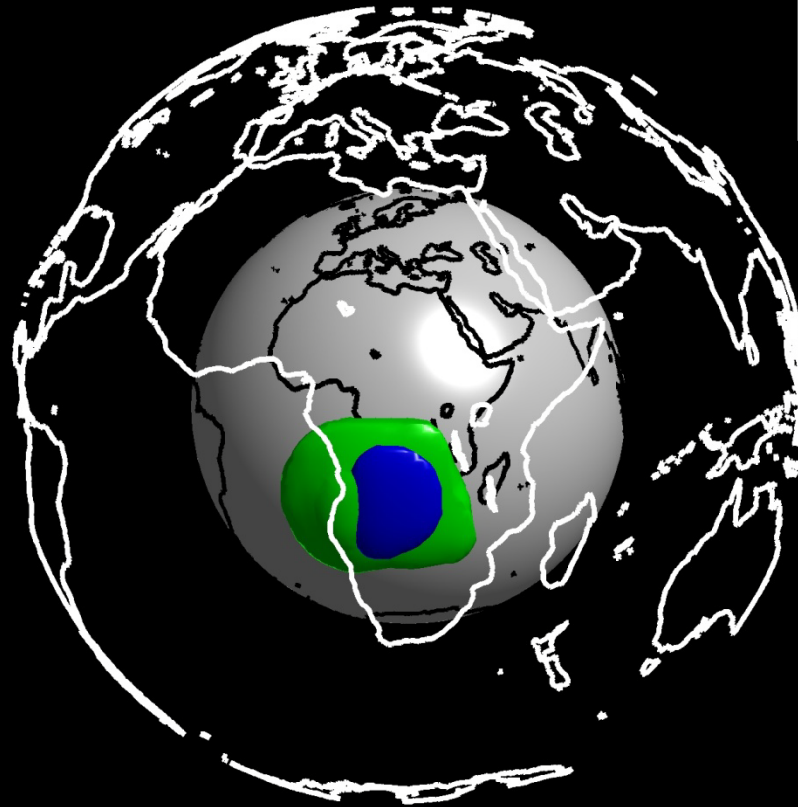


Dense Pile Part I

Properties:

- Low Shear Speed
- High Sound Speed
- High Density

Intrinsically dense, hot material.

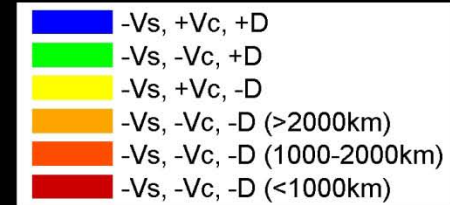


Dense Pile Part II

Properties:

- Low Shear Speed
- **Low** Sound Speed
- High Density

Hotter outer shell. Hot enough to reduce sound speed...dense, partial melt?

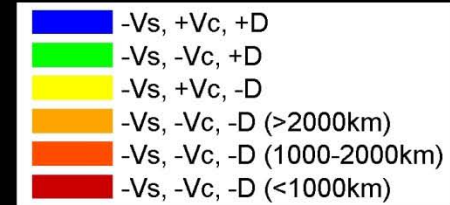


High Entrainment Zone

Properties:

- Low Shear Speed
- **High** Sound Speed
- **Low** Density

Upwelling partly composed of intrinsically dense material entrained from the dense pile. Seen in the SASP, but not the WASP.

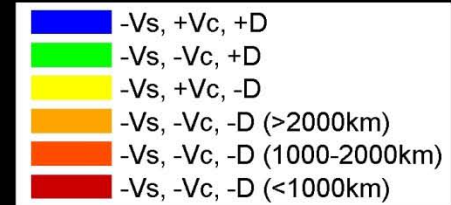
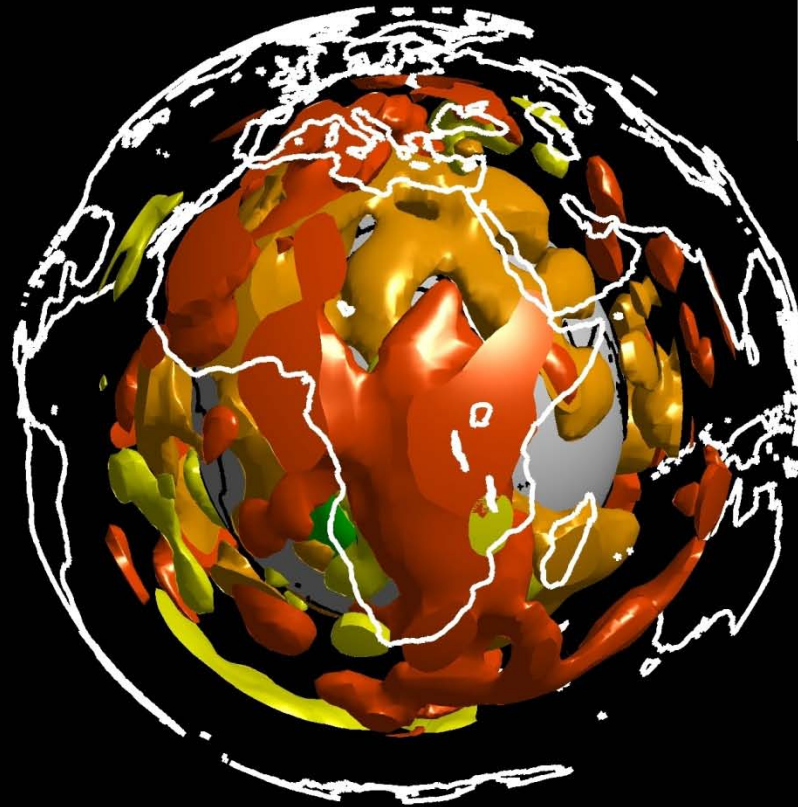


Deep Negative Zones

Properties:

- Low Shear Speed
- **Low** Sound Speed
- Low Density

Buoyant material without significant chemical signatures. Comprises the remaining low shear zones deep beneath Africa.

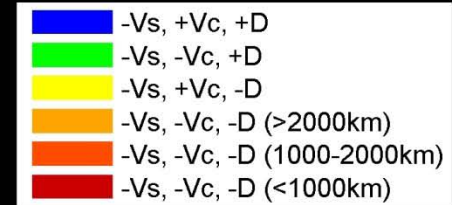
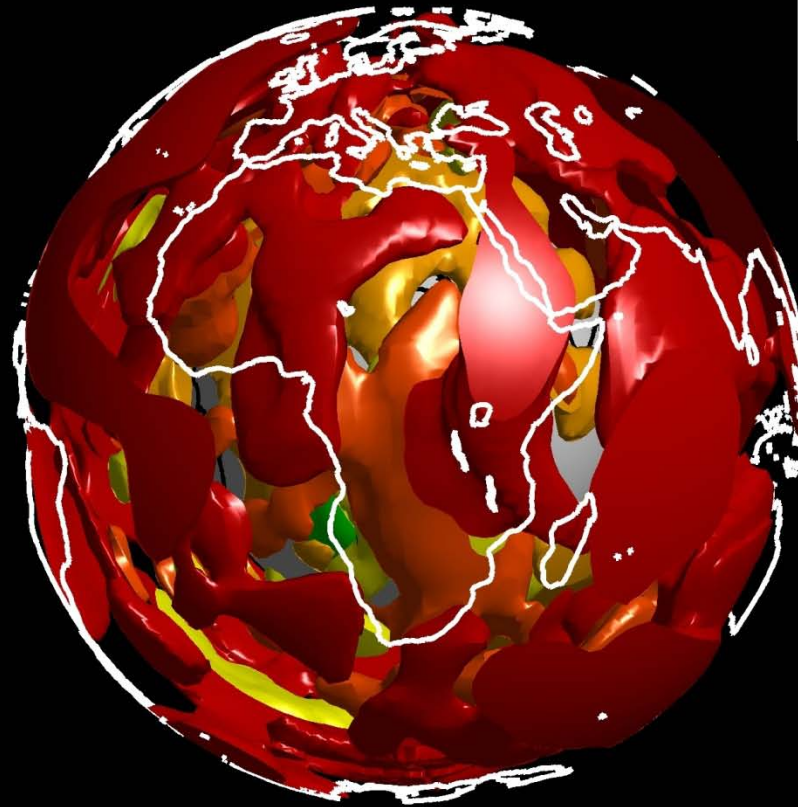


Mid-mantle Negative Zone

Properties:

- Low Shear Speed
- Low Sound Speed
- Low Density

Buoyant material rising towards the EARZ, Cameroon, and Cape Verde. High-density chemical signatures seen in extensions from the SASP.



Shallow Negative Zone

Properties:

- Low Shear Speed
- Low Sound Speed
- Low Density

Buoyant material rising towards the EARZ, Cameroon, and Cape Verde. Extension toward Hoggar with compositional signature. Possible SASP fingerprint.

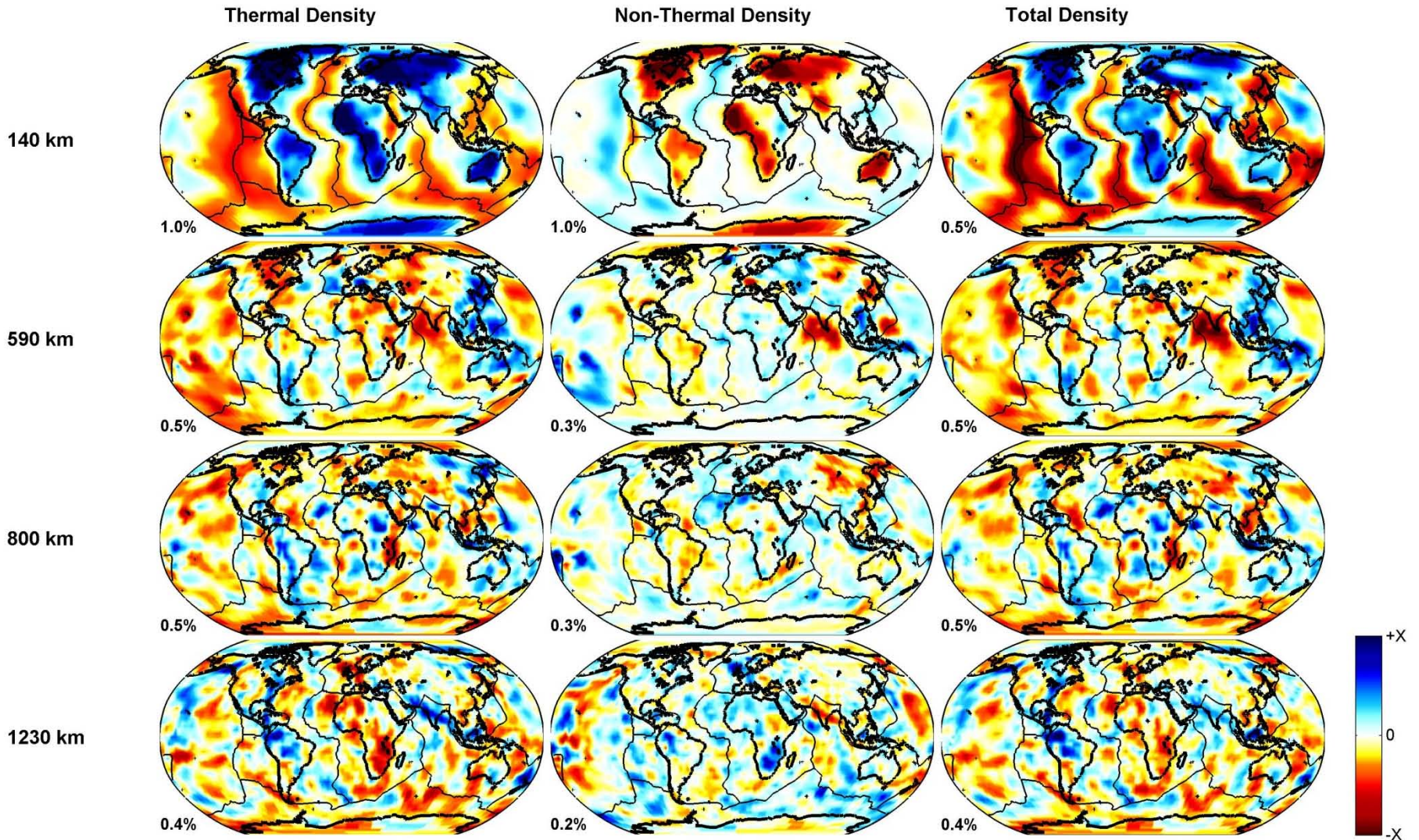
Summary

- We have constructed a global-scale joint seismic-geodynamic-mineral-physics model (GyPSuM)
 - With a “minimum composition” approach
 - Except for cratonic roots and parts of D”, temperature seems to dominate

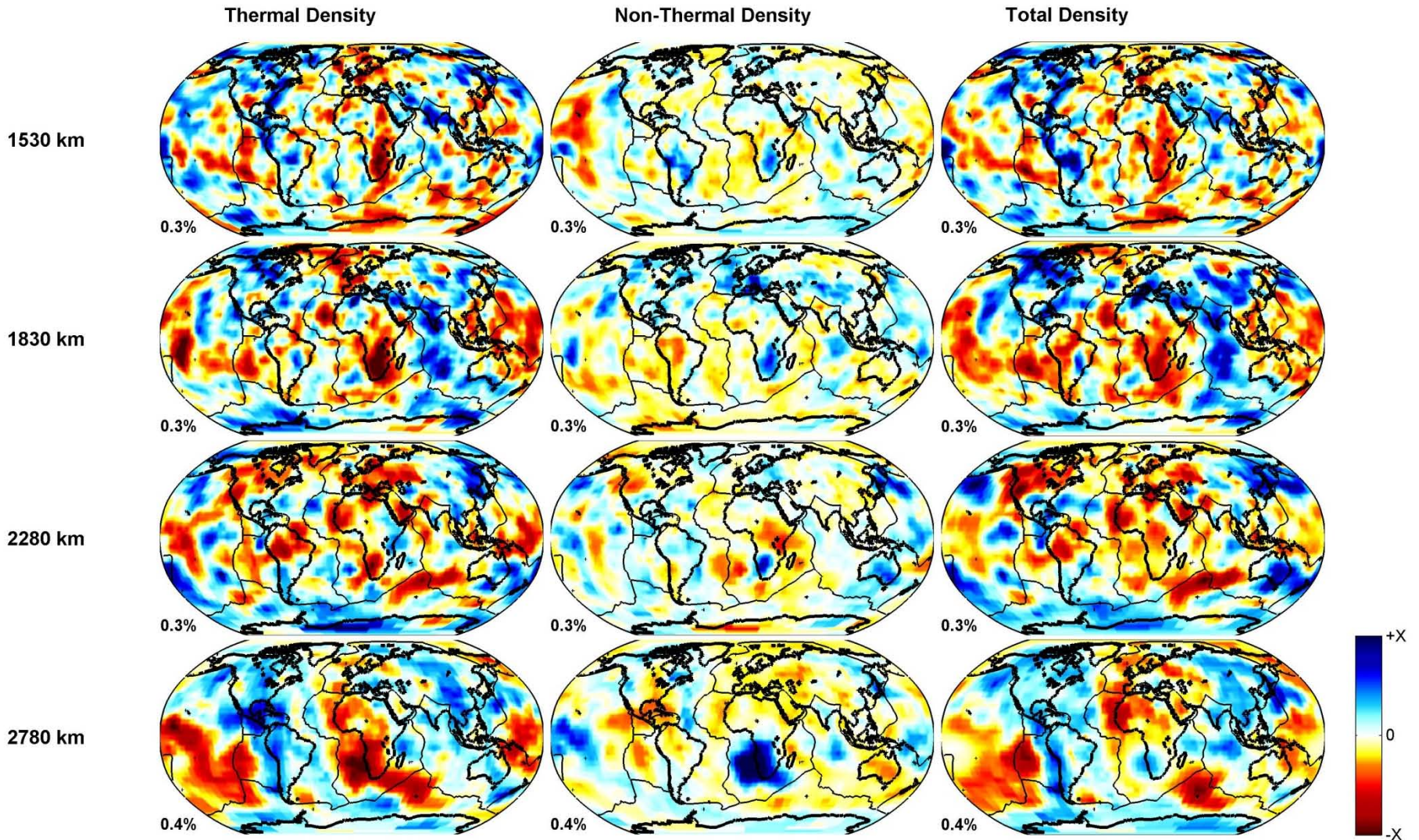
- We have started with broadly-defined mineral physics constraints
 - Future models should incorporate more recent and complete mineral physics relationships
 - Trade-offs are problematic

- Model available for download on the IRIS website:
 - <http://www.iris.edu/dms/products/emc/>

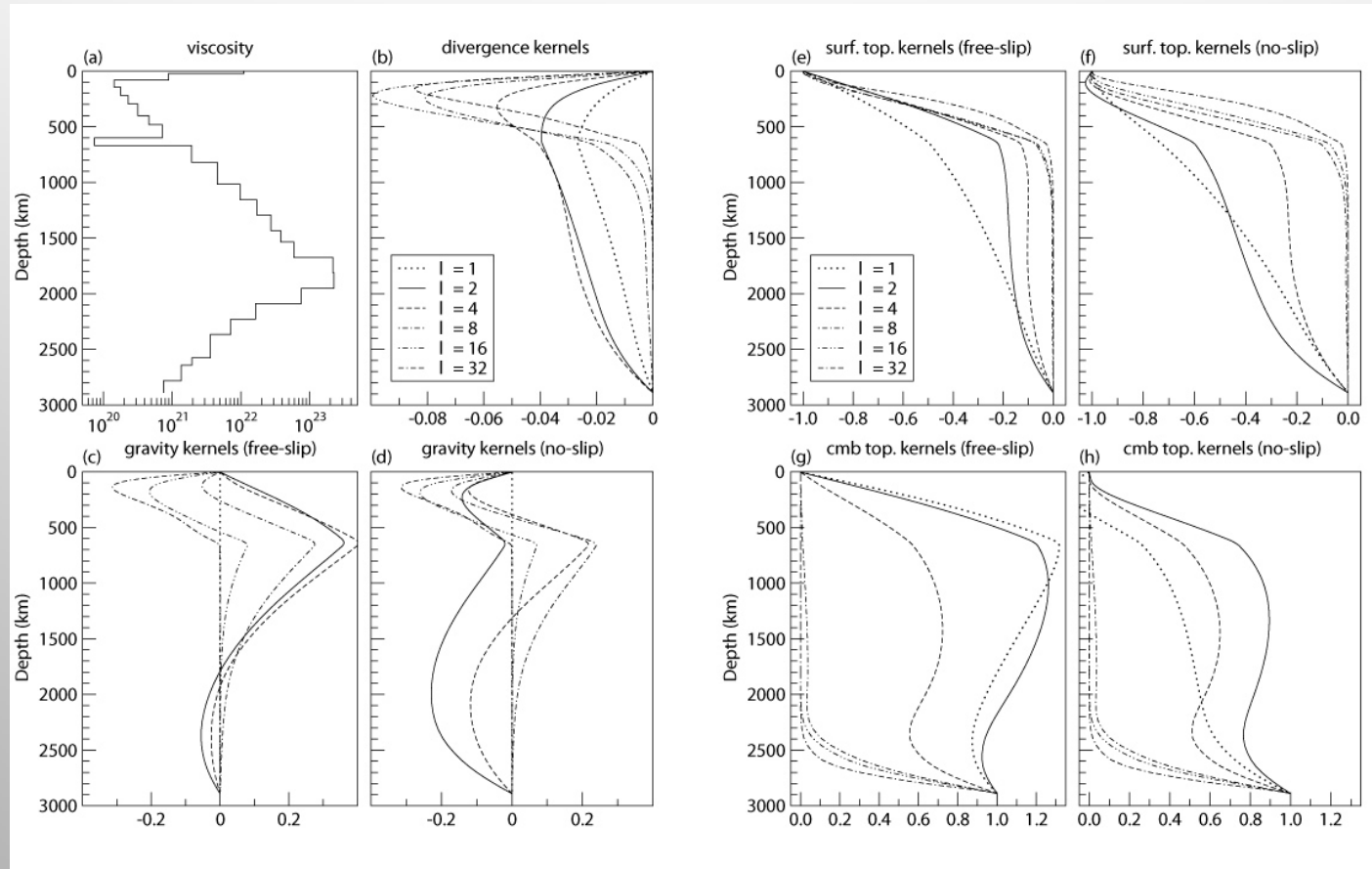
GyPSuM: Thermal VS non-thermal heterogeneity



GyPSuM: Thermal VS non-thermal heterogeneity



GyPSuM: Viscous Flow Response Kernels



GyPSuM: Density-P-S Coupling

$$R_{\rho/S} = \frac{2}{(\Gamma-1) + 2Q_S^{-1}X(\xi)C}$$

anharmonic anelastic

$$\Gamma \equiv \frac{\delta \log \mu}{\delta \log \rho}$$

Anderson-Grüneisen
parameter

$$R_{\phi/S} = \frac{(\delta_s - 1)}{(\Gamma-1) + 2Q_S^{-1}X(\xi)C} = \frac{(\delta_s - 1)}{2} R_{\rho/S}$$

anharmonic anelastic

$$\delta_s \equiv \frac{\delta \log K}{\delta \log \rho}$$

Anderson-Grüneisen
parameter

$$\frac{\delta V_P}{V_P} = \gamma \frac{\delta V_S}{V_S} + (1-\gamma) \frac{\delta V_\phi}{V_\phi}$$

$$\gamma \equiv \frac{4 V_S^2}{3 V_P^2}$$

