

# Structural seismology constraints for mantle dynamics

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Rob Porritt, Lapo Boschi, Ludwig Auer

*Global Scale Seismic Imaging and Dynamics of the Earth's Mantle*

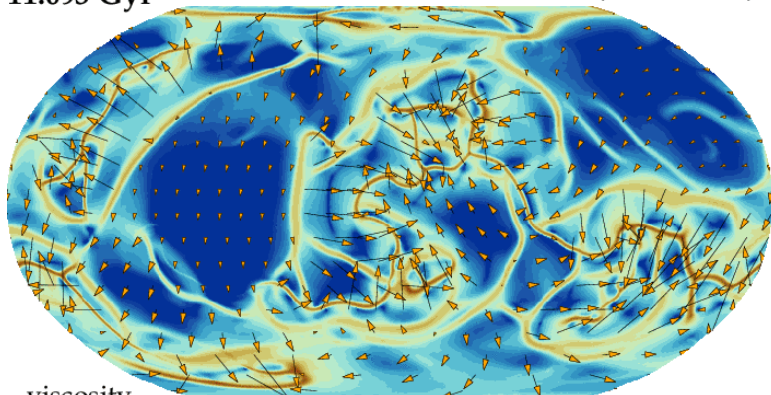
College de France

October 7, 2021

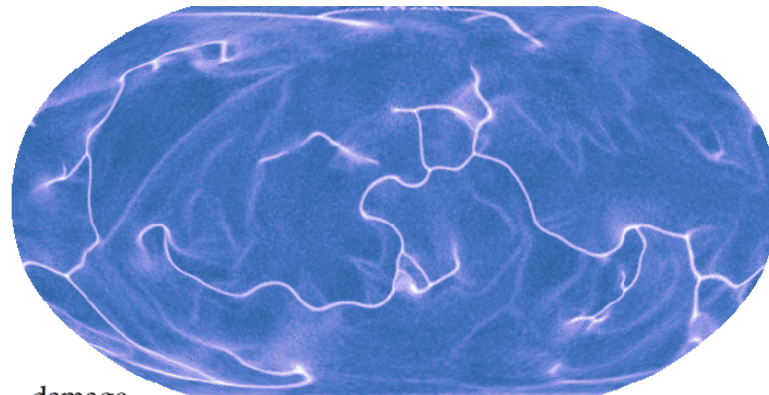
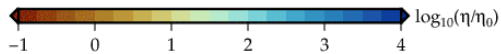
# Role of rheology and memory for plate tectonics

11.093 Gyr

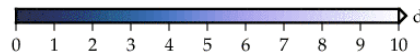
Rate8P1e6\_ddy\_151e7\_10\_244e9\_et\_129xyz\_6



viscosity

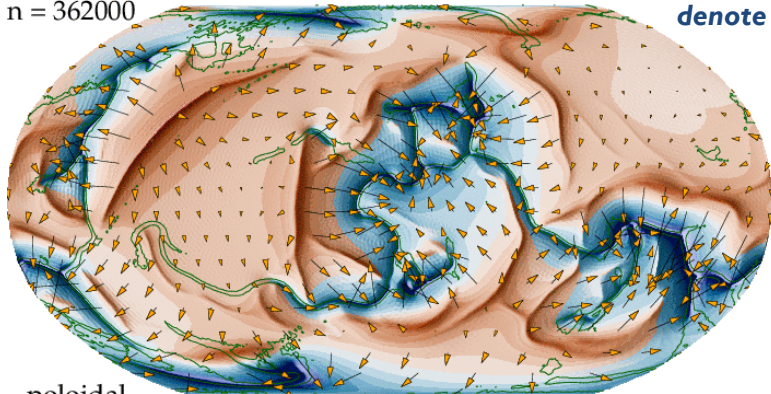


damage

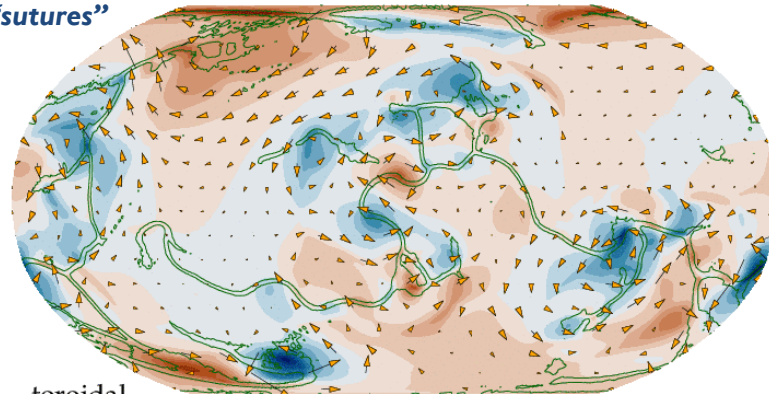
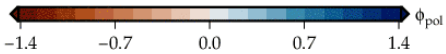


green contours  
denote "sutures"

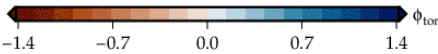
$n = 362000$



poloidal



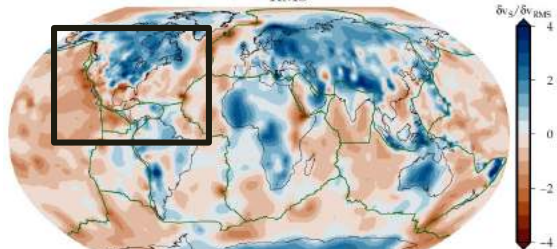
toroidal



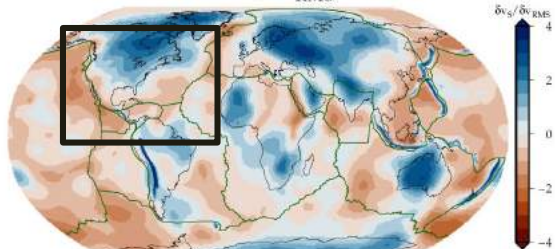
Global mantle convection computation  
using modified *CitcomS* (cf. Zhong et al.,  
2000), convective vigor  $\sim 0.1$  Earth

# Shear wave mantle tomography

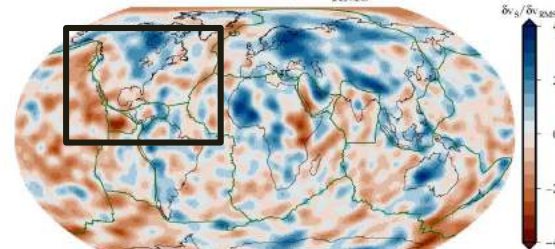
RSAVANI @ 250 km,  $\delta v_{\text{RMS}} = 1.20\%$



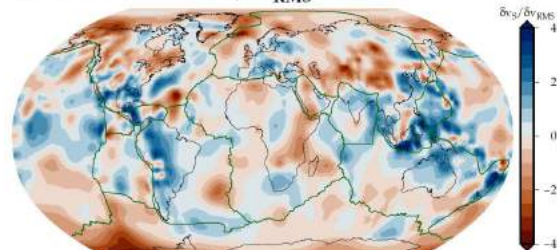
TX2019SLAB @ 250 km,  $\delta v_{\text{RMS}} = 0.80\%$



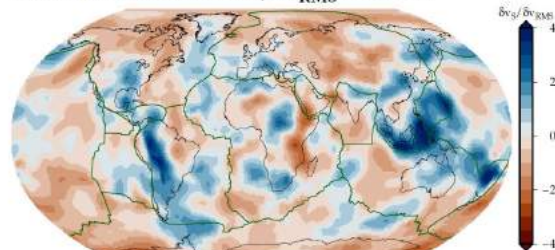
SEMUCB-WM1 @ 250 km,  $\delta v_{\text{RMS}} = 1.36\%$



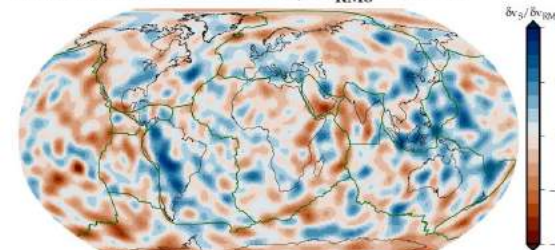
RSAVANI @ 750 km,  $\delta v_{\text{RMS}} = 0.57\%$



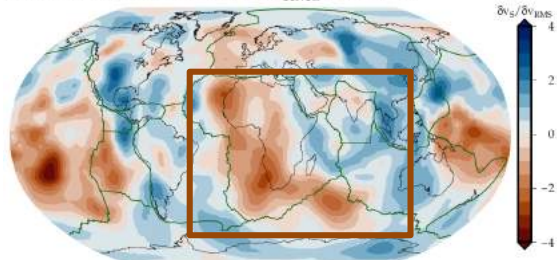
TX2019SLAB @ 750 km,  $\delta v_{\text{RMS}} = 0.51\%$



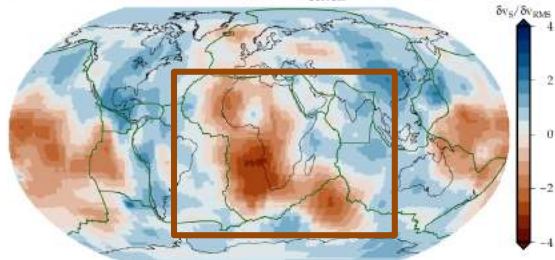
SEMUCB-WM1 @ 750 km,  $\delta v_{\text{RMS}} = 0.61\%$



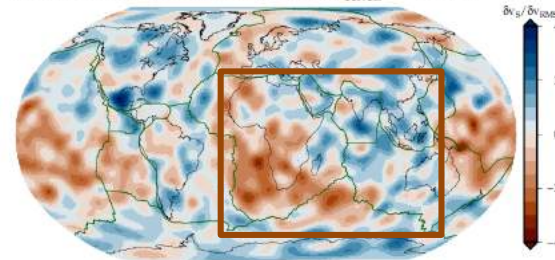
RSAVANI @ 2850 km,  $\delta v_{\text{RMS}} = 0.67\%$



TX2019SLAB @ 2850 km,  $\delta v_{\text{RMS}} = 1.08\%$



SEMUCB-WM1 @ 2850 km,  $\delta v_{\text{RMS}} = 1.16\%$



Porritt et al. (2021)

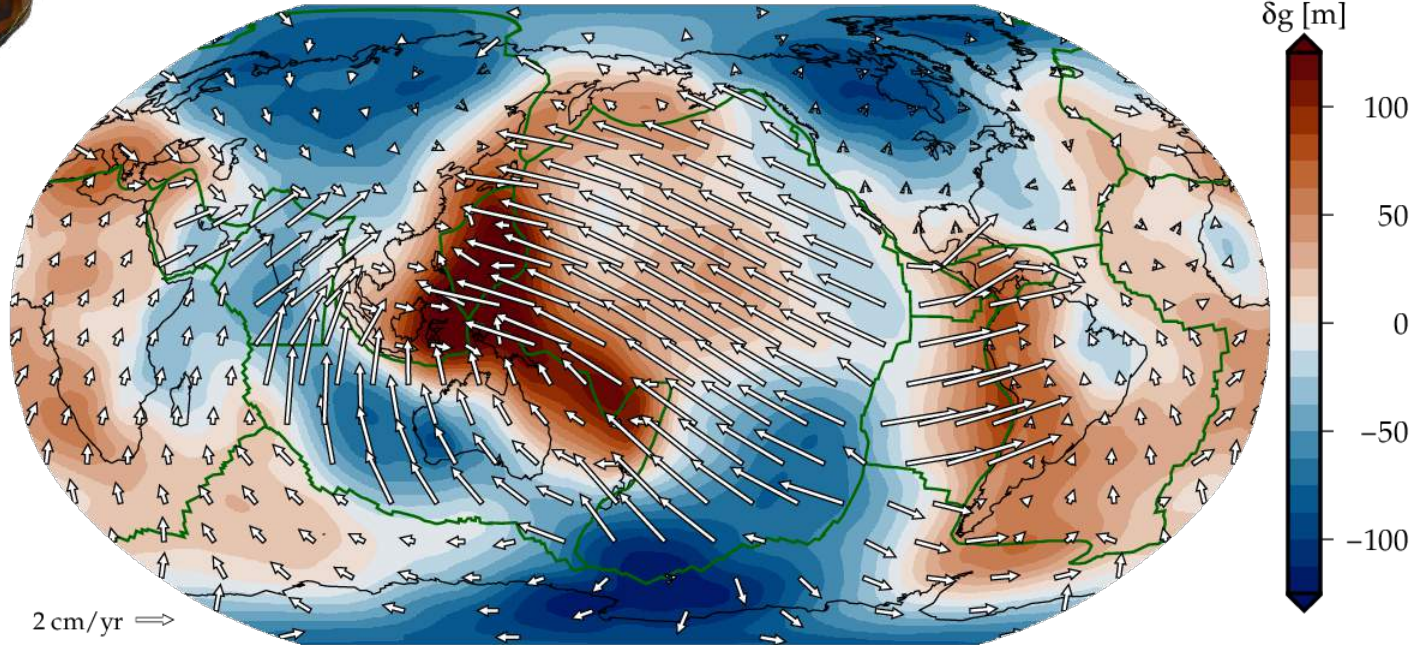
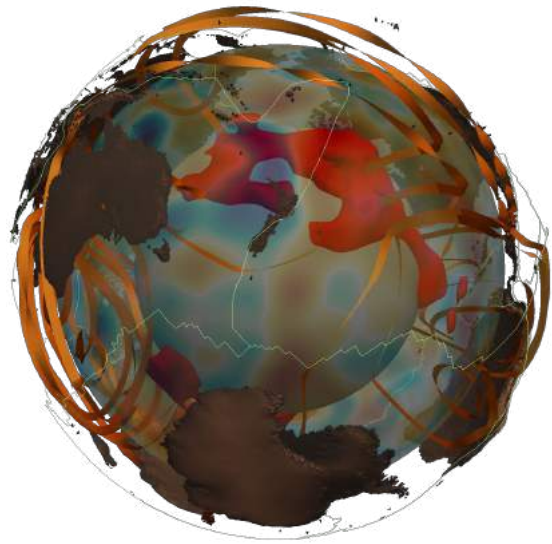
Lu et al. (2019)

French et al. (2015)

# Global mantle circulation models based on seismic tomography

$$\text{velocities} \propto \frac{\text{density anomaly}}{\text{viscosity}}$$

$$\text{geoid and topography} \propto \text{density anomaly}$$



Correlation with  
plate velocities: 0.89  
geoid: 0.82

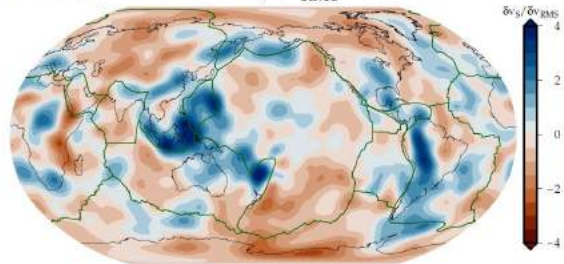


Dynamically consistent  
computation with prescribed  
weak zones and lateral  
viscosity variations based on TX2019

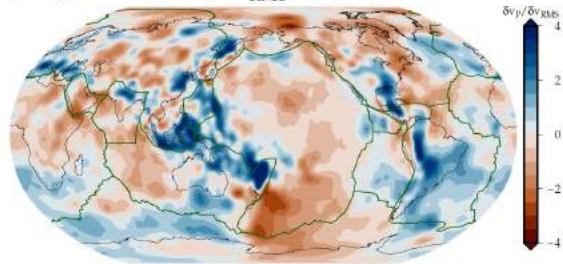
2 cm/yr →

# S vs P tomography: thermo-chemical effects

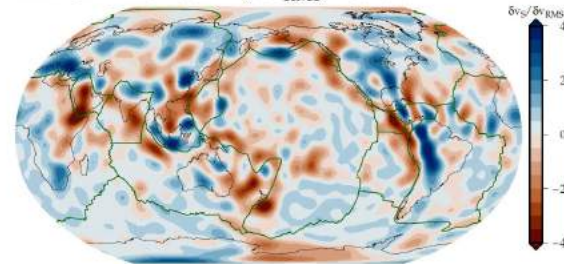
TX2019SLAB @ 750 km,  $\delta v_{RMS} = 0.50\%$



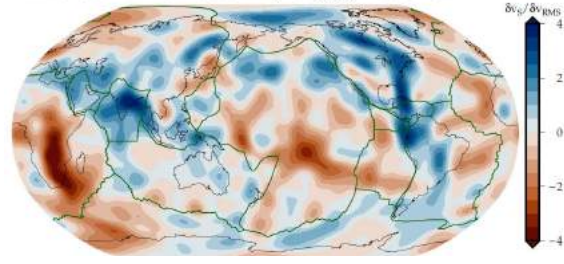
GAP\_P4 @ 750 km,  $\delta v_{RMS} = 0.30\%$



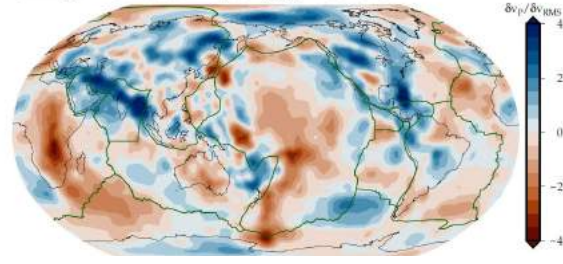
DETOX-P2 @ 750 km,  $\delta v_{RMS} = 0.19\%$



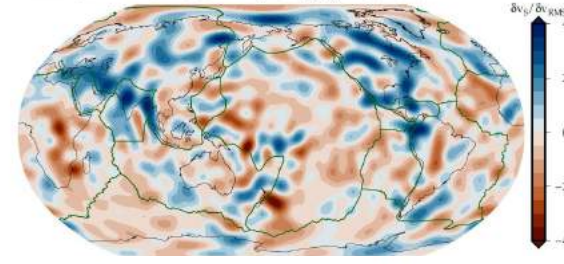
TX2019SLAB @ 1500 km,  $\delta v_{RMS} = 0.35\%$



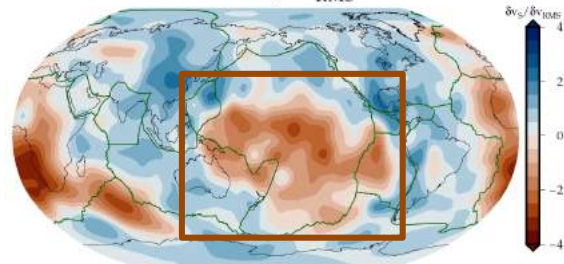
GAP\_P4 @ 1500 km,  $\delta v_{RMS} = 0.18\%$



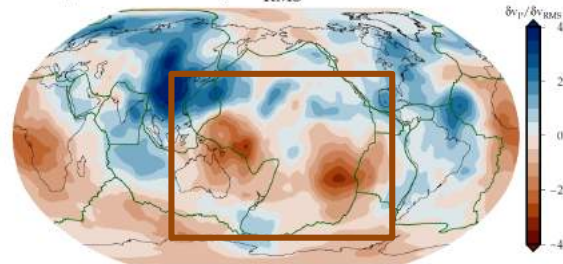
DETOX-P2 @ 1500 km,  $\delta v_{RMS} = 0.18\%$



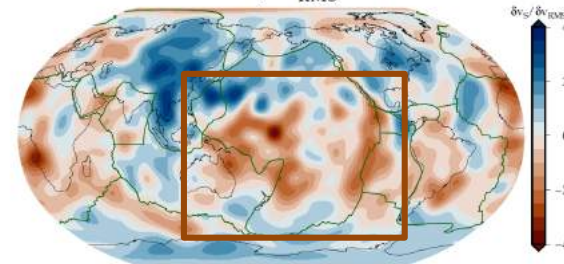
TX2019SLAB @ 2850 km,  $\delta v_{RMS} = 0.98\%$



GAP\_P4 @ 2850 km,  $\delta v_{RMS} = 0.33\%$



DETOX-P2 @ 2850 km,  $\delta v_{RMS} = 0.28\%$

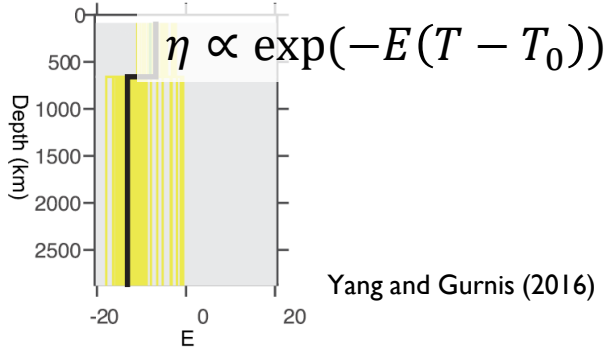


Lu et al. (2019)

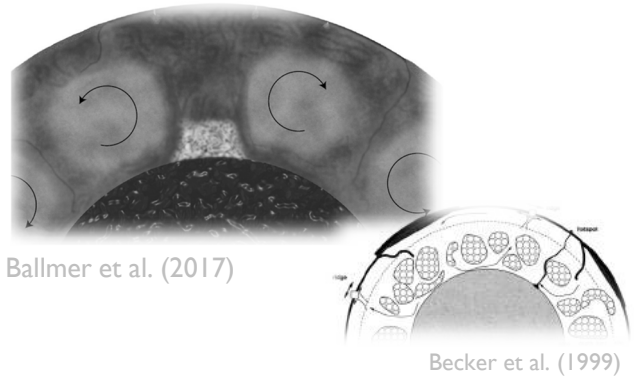
Obayashi et al. (2013)

Hosseini et al. (2020)

# Next level challenges: Inferring mantle rheology



Yang and Gurnis (2016)

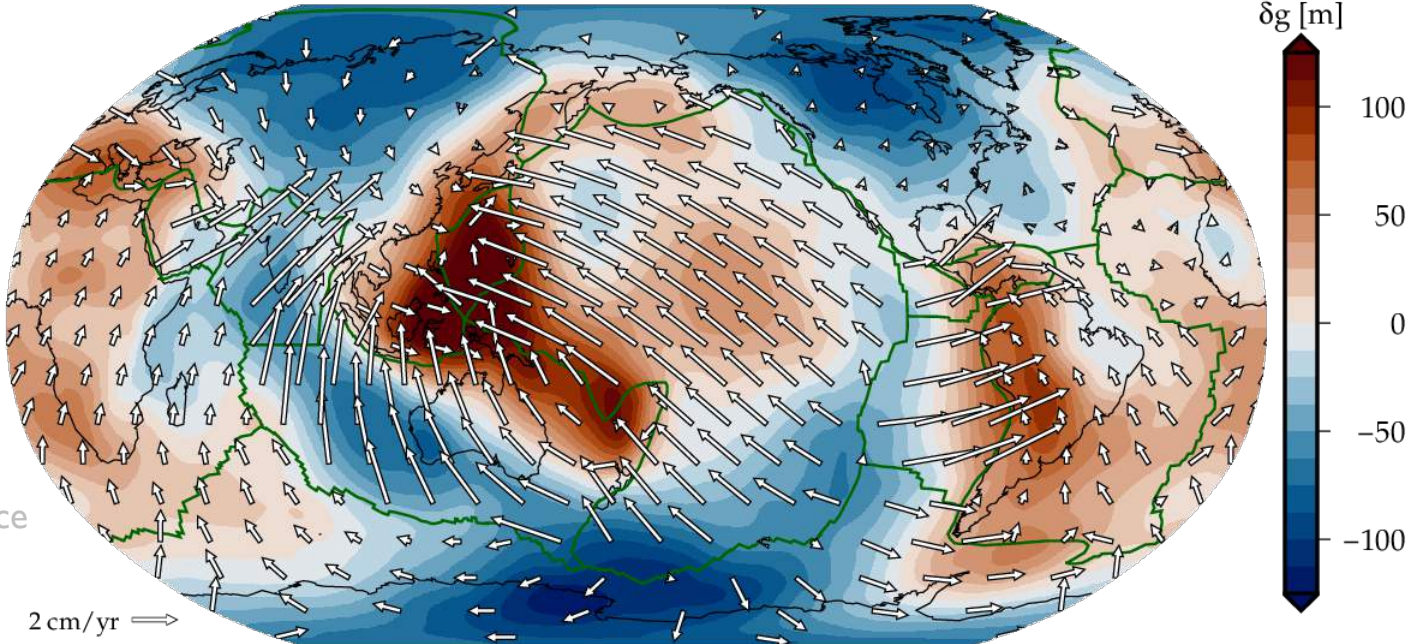


Ballmer et al. (2017)

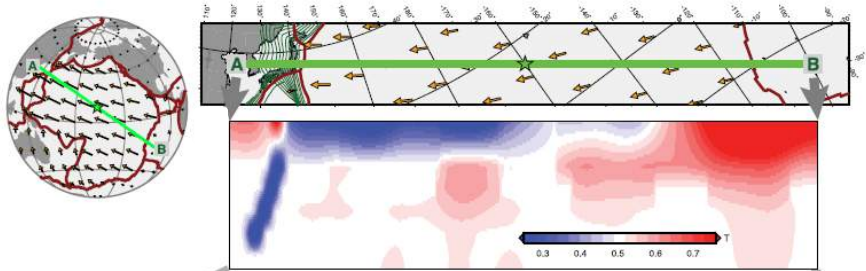
Becker et al. (1999)

Correlation with  
plate velocities: 0.84 (0.89)  
geoid: 0.84 (0.82)

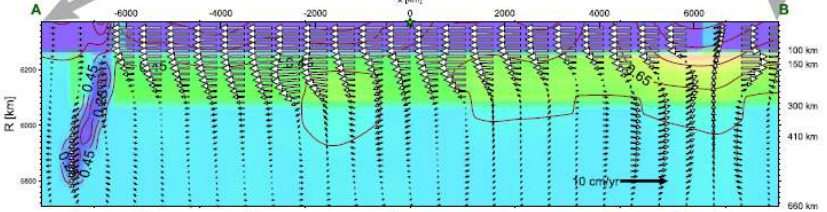
$E = -7$   
Inverse temperature dependence  
interpreting slow regions  
as BEAMS



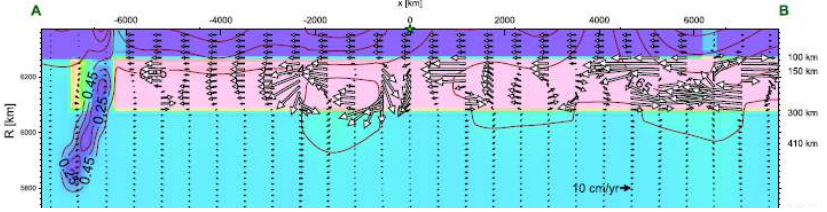
# How weak is the asthenosphere?



background asthenosphere, tomography and slabs



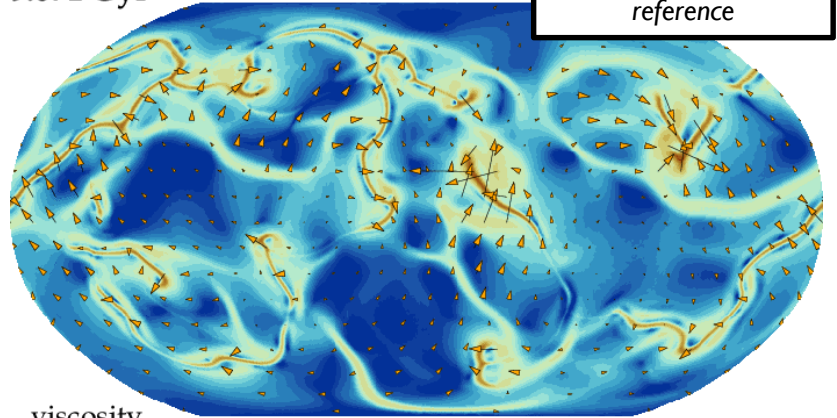
weak asthenosphere



Becker (2017)

9.071 Gyr

reference



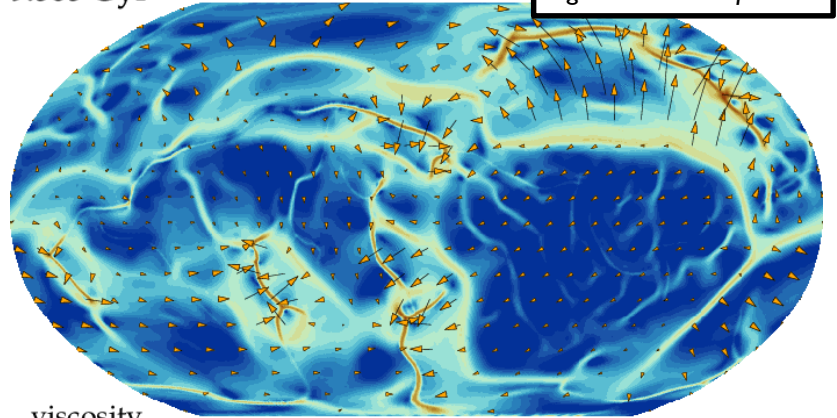
viscosity

$n = 286000$

$\log_{10}(\eta/\eta_0)$

9.585 Gyr

global asthenosphere

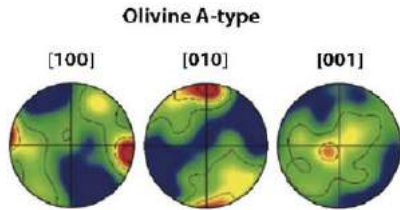
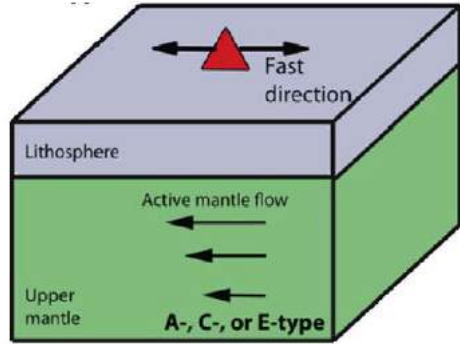


viscosity

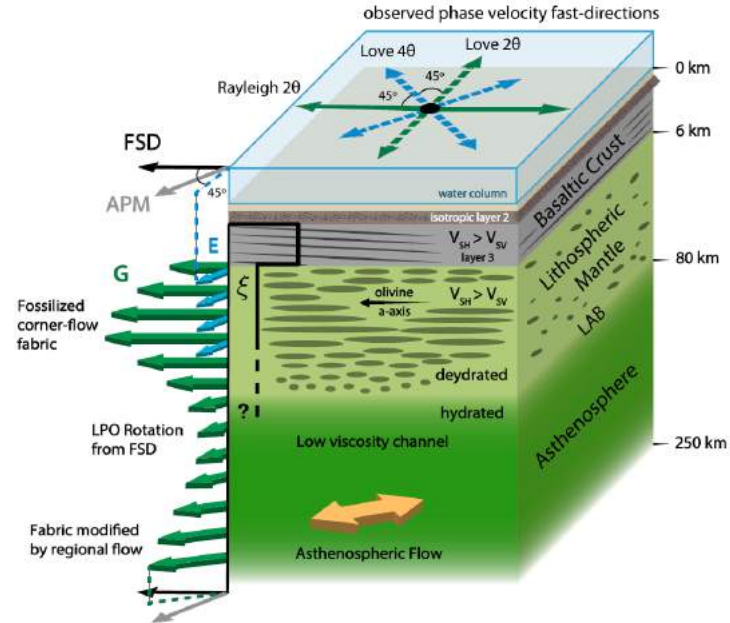
$n = 330000$

$\log_{10}(\eta/\eta_0)$

# Upper mantle seismic anisotropy as a constraint for rheology



Long and Becker (2010)

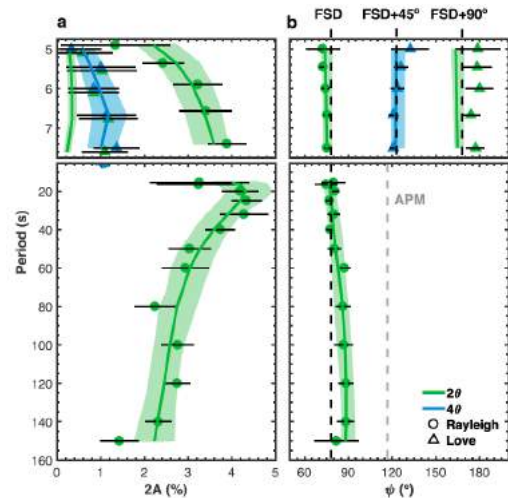
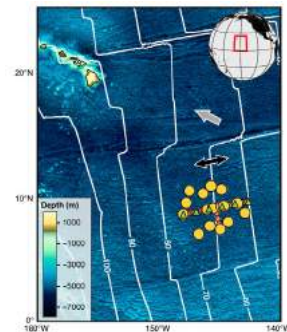
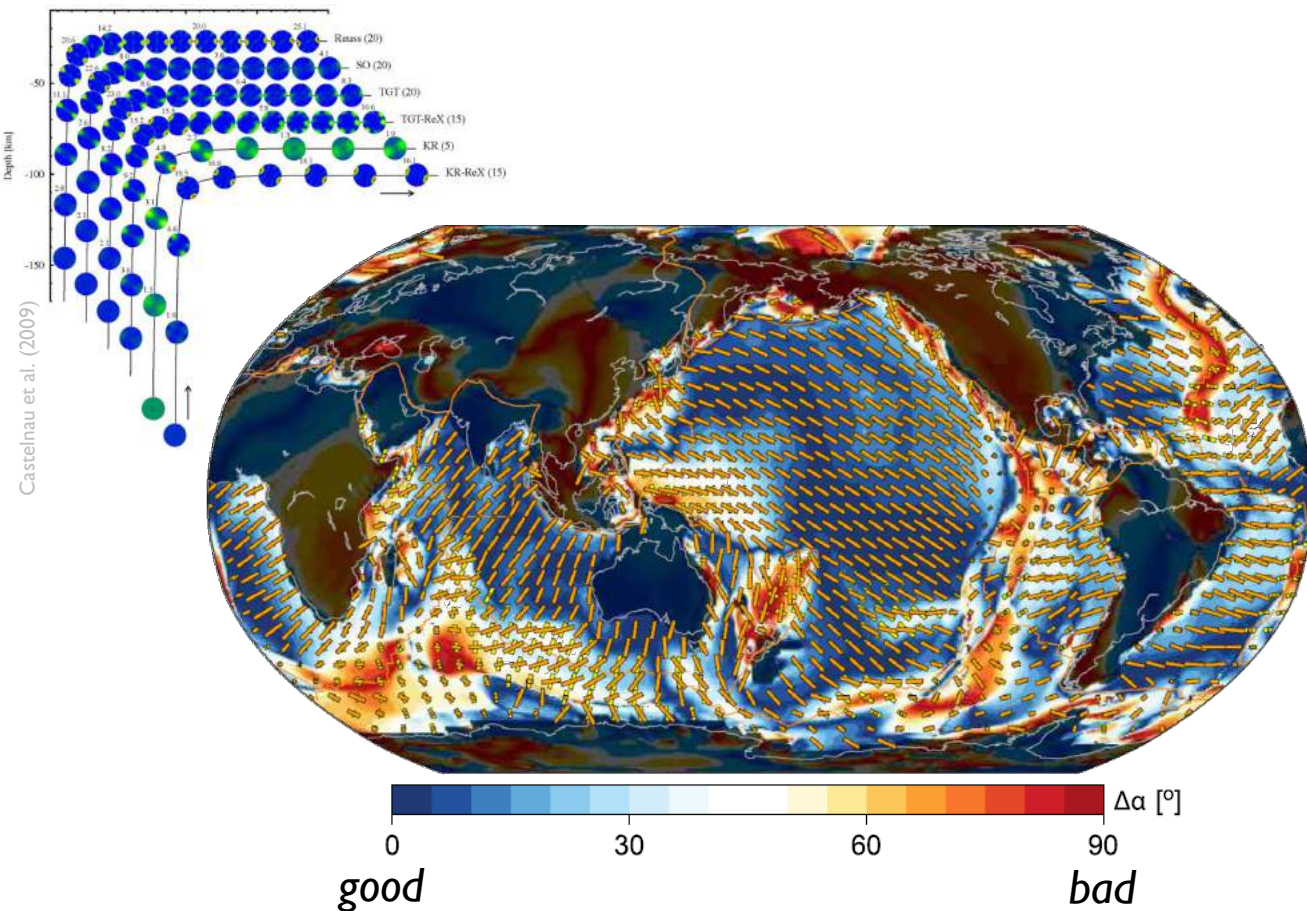


Russell et al. (2019)

- crystallographic preferred orientation (CPO) of olivine sensitive to time-integral of mantle circulation



# Azimuthal anisotropy in oceanic basins

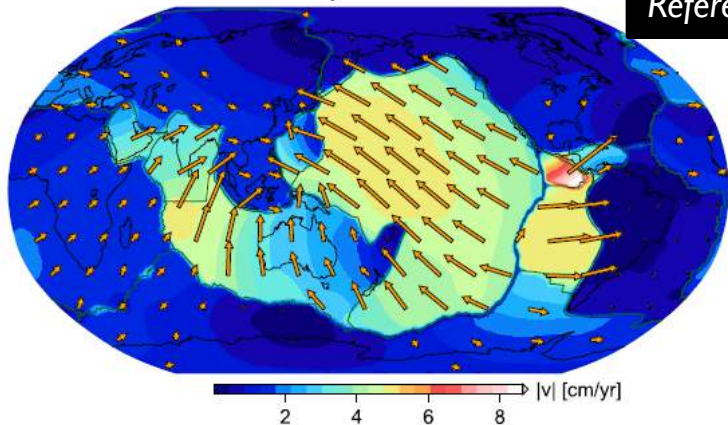


## Plate velocity predictions

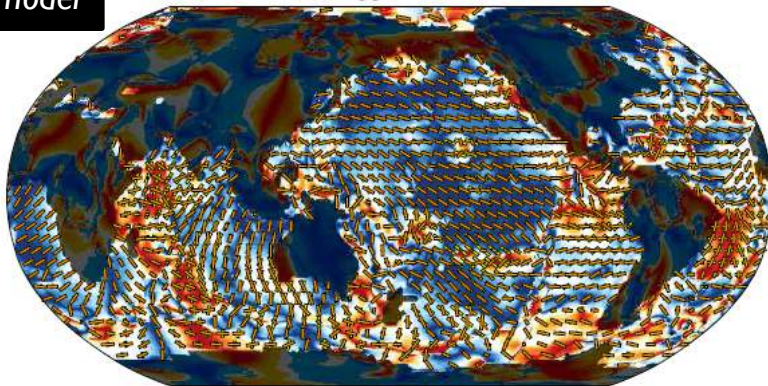
## Azimuthal anisotropy match

$r_v = 0.916$

Reference model

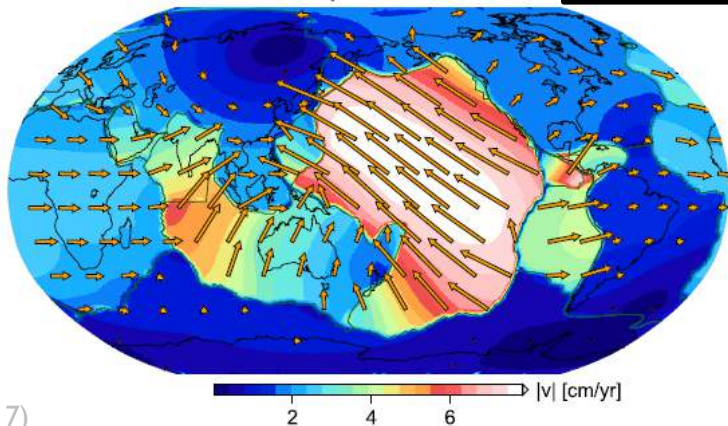


$\langle \Delta\alpha \rangle_{oc} = 32.9^\circ$

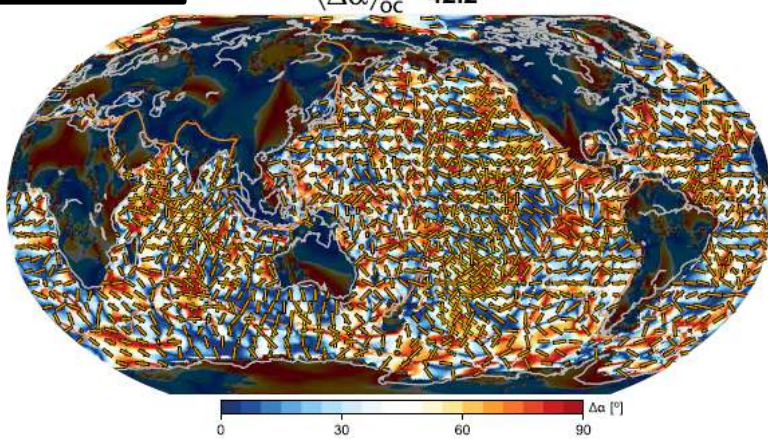


$r_v = 0.910$

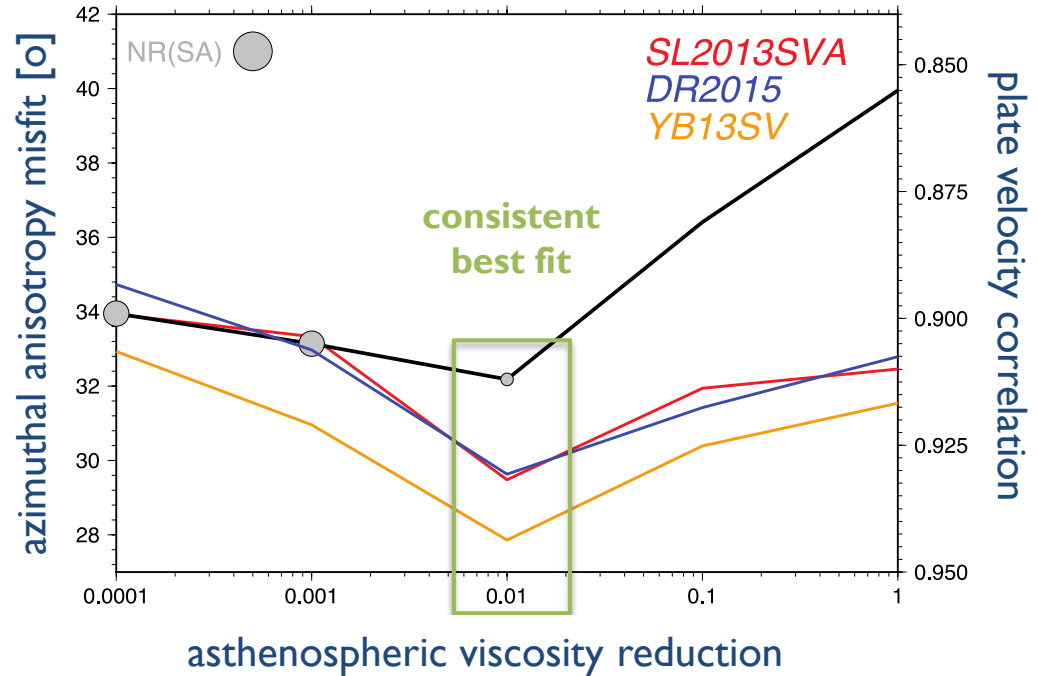
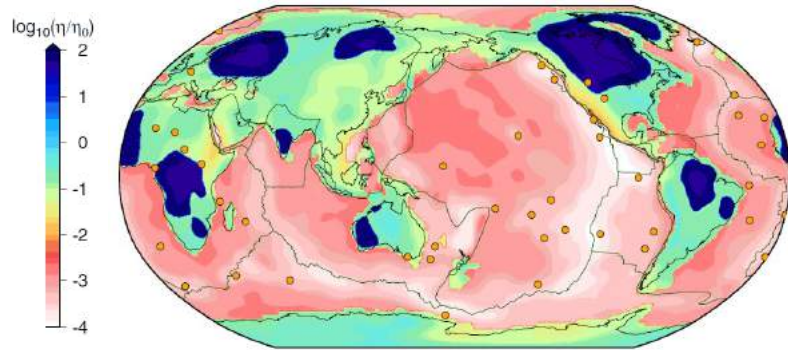
Super weak asthenosphere



$\langle \Delta\alpha \rangle_{oc} = 42.2^\circ$

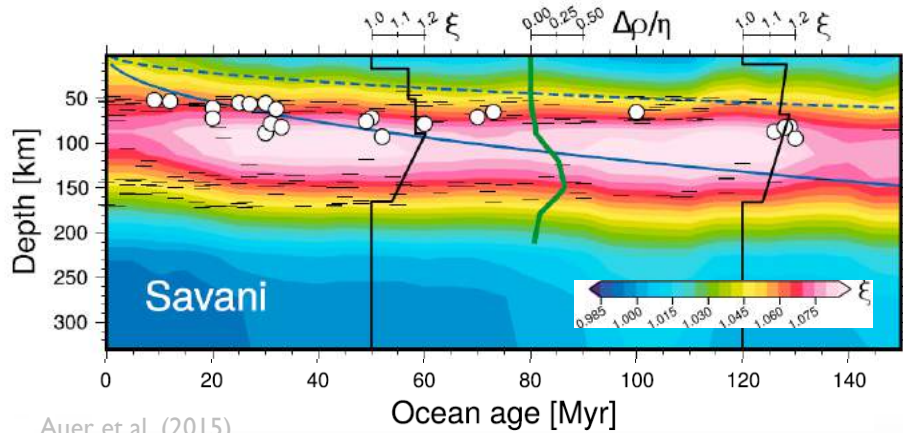


# Predicting plate velocities and azimuthal anisotropy

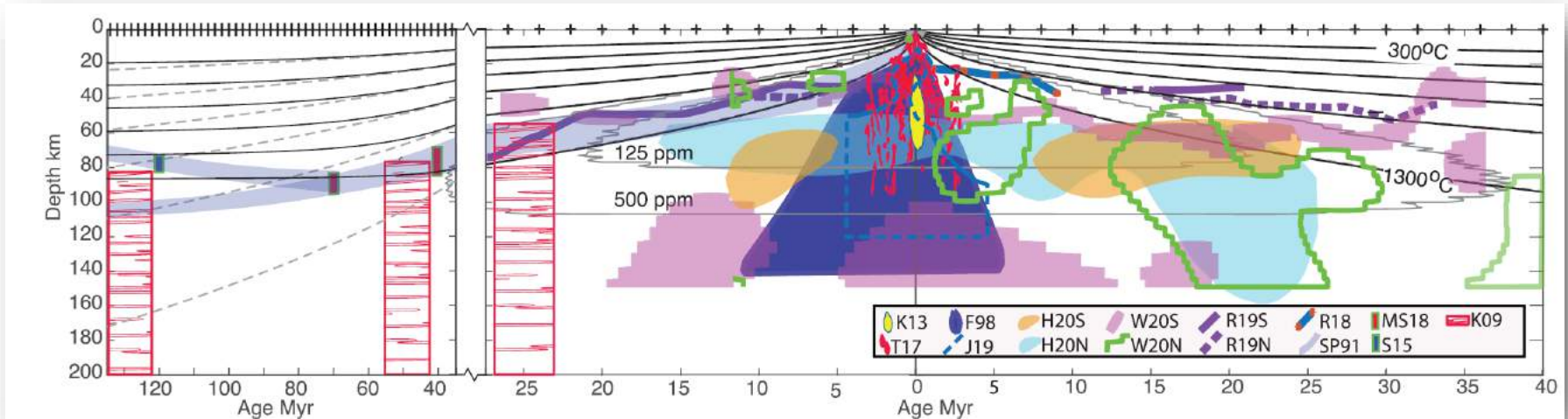


- sweet spot for oceanic asthenosphere viscosity reduction
- trade-off between thickness and strength of asthenosphere
- if melt affects rheology, should be local/not connected, else major disruption of anisotropy

# Yet, melt may be required to explain seismic features?



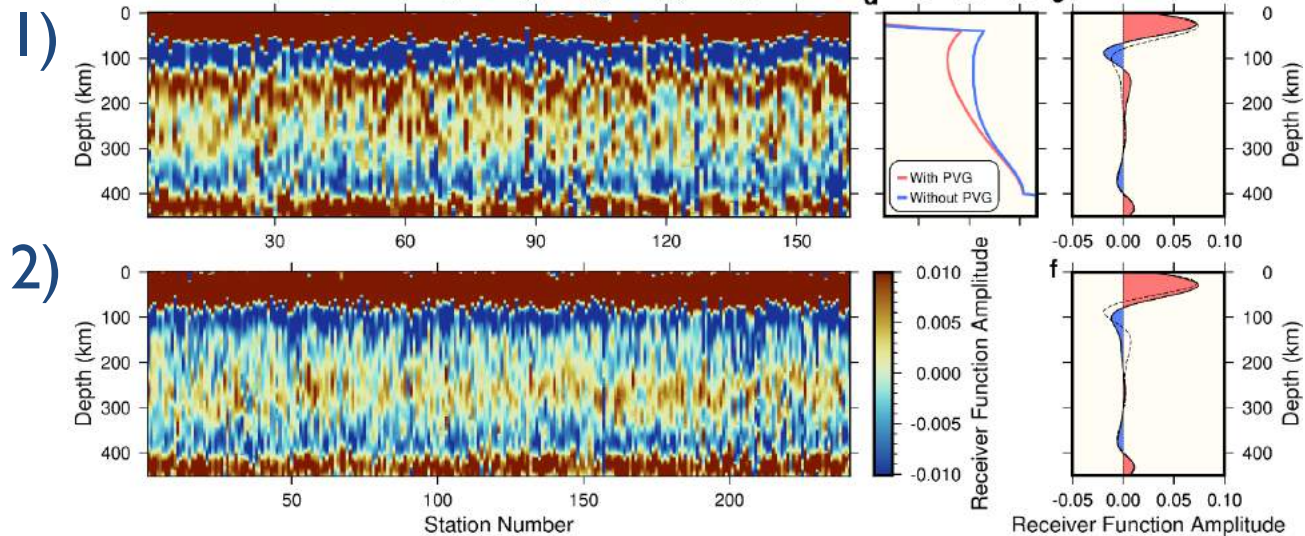
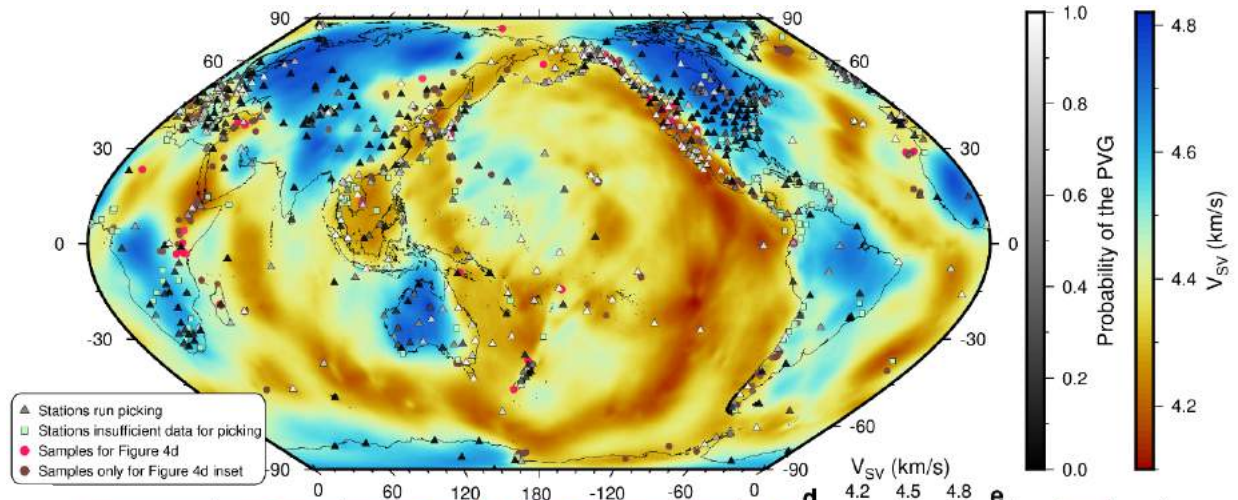
Auer et al. (2015)



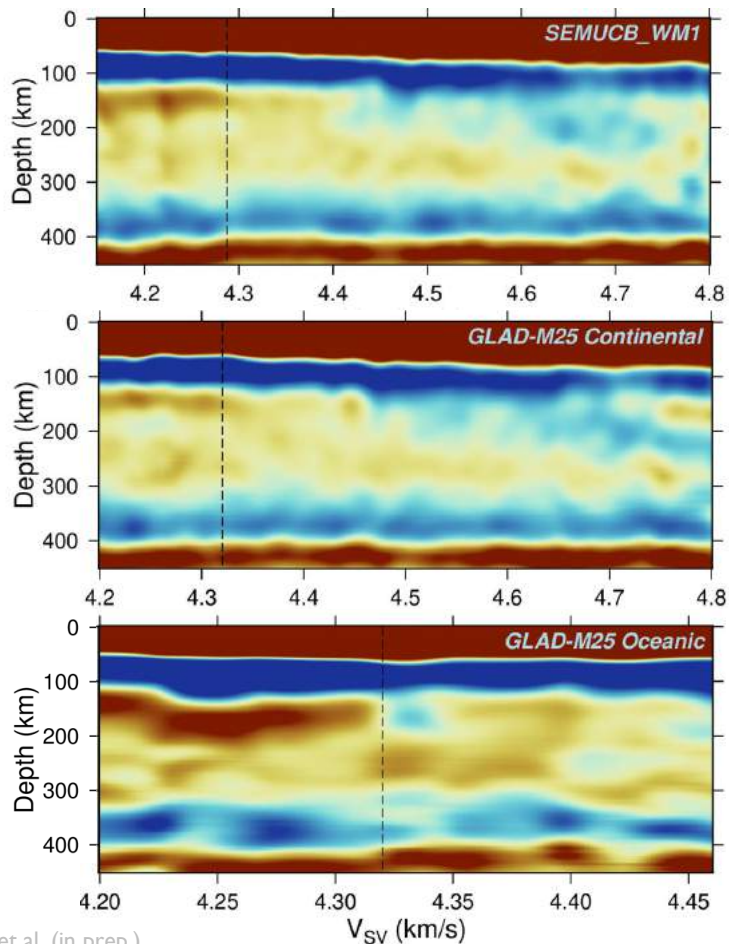
Rychert et al. (2020)

# Role of melt

Two classes of velocity profiles from receiver function analysis

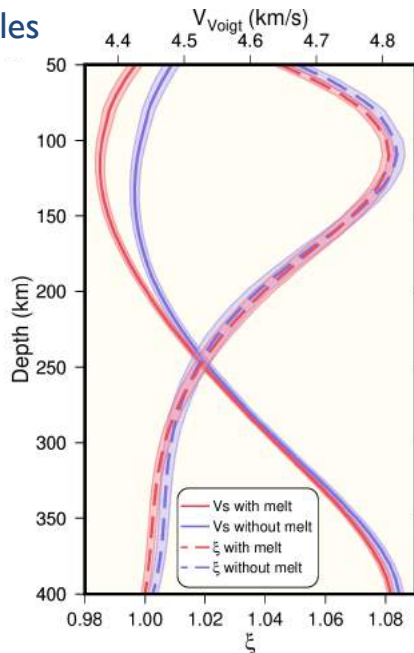


receiver functions sorted by  $v_S$

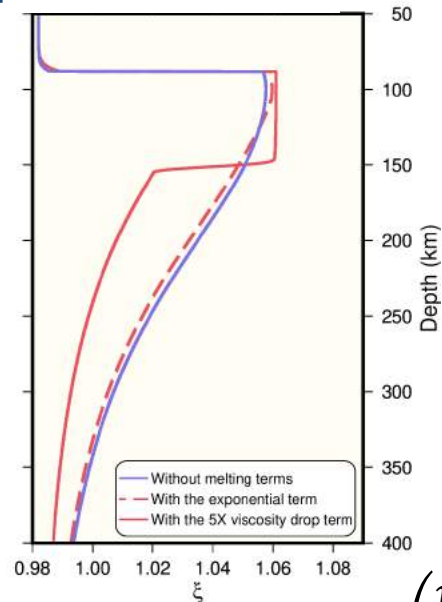


# Partial melting associated with relative hot regions

Observed profiles

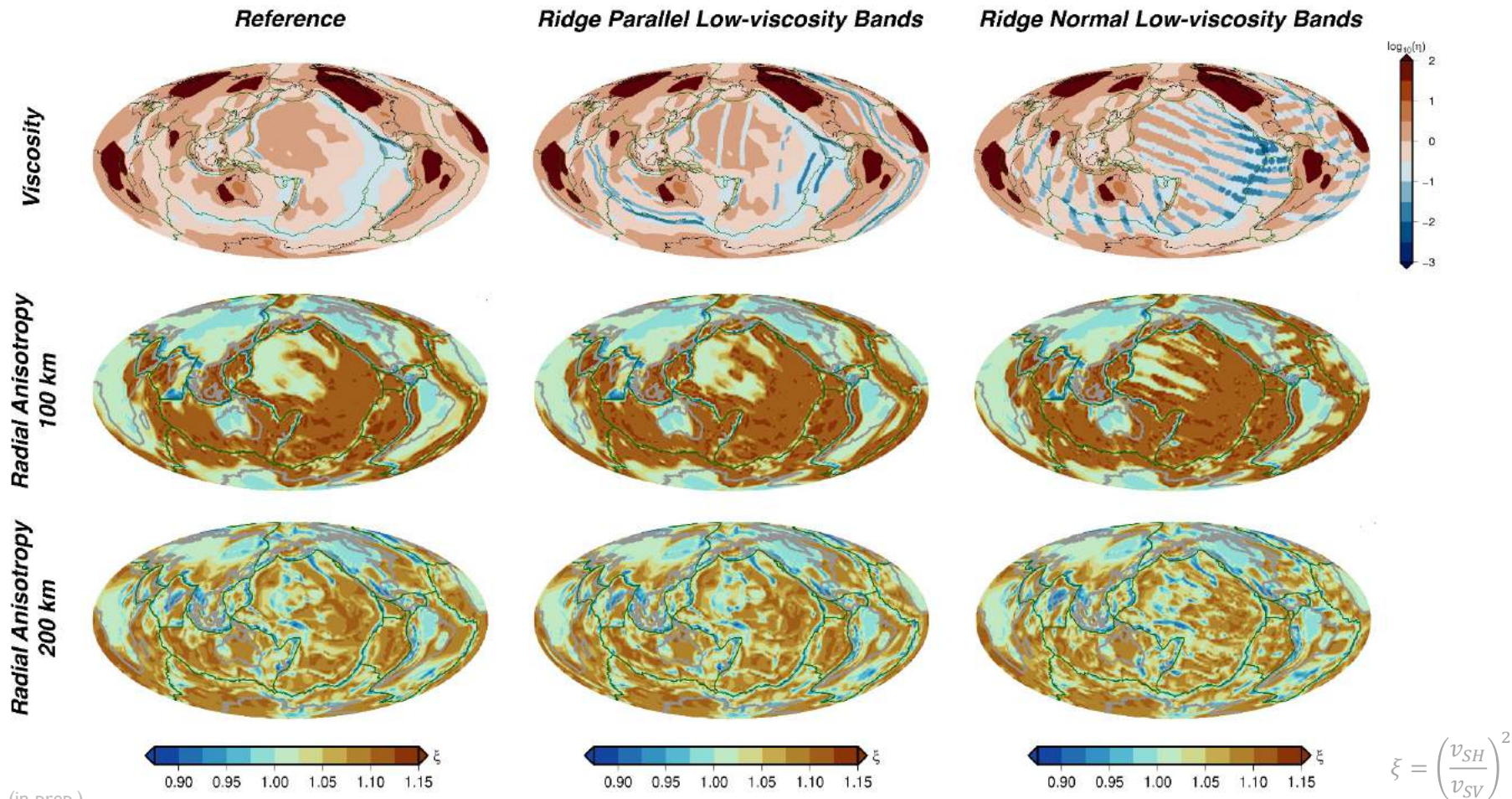


I-D radial anisotropy prediction



$$\xi = \left( \frac{v_{SH}}{v_{SV}} \right)^2$$

# Asthenospheric low viscosity channels and radial anisotropy



# Upper mantle radial anisotropy

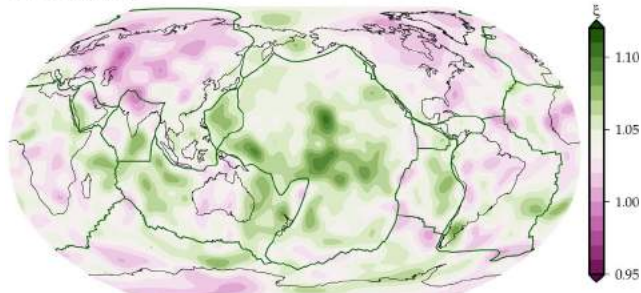
For isotropic shear wave speeds:

$$\langle r_8 \rangle \geq 0.7$$

$$\xi = \left( \frac{v_{SH}}{v_{SV}} \right)^2$$

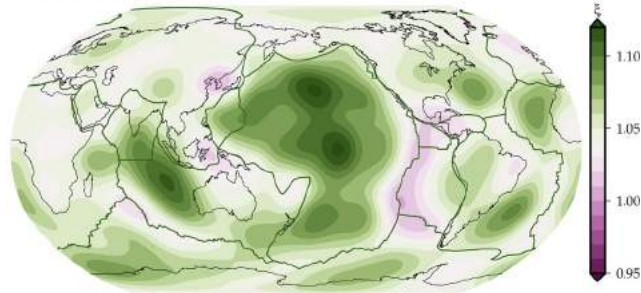
150 km depth

RSAVANI



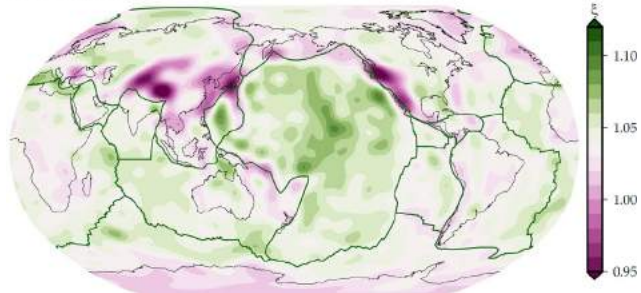
Porritt et al. (2021)

SEMUCB-WM1



French et al. (2015)

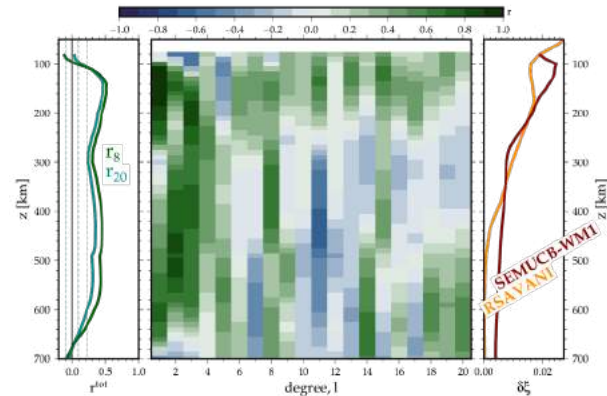
SGLOBE-RANI



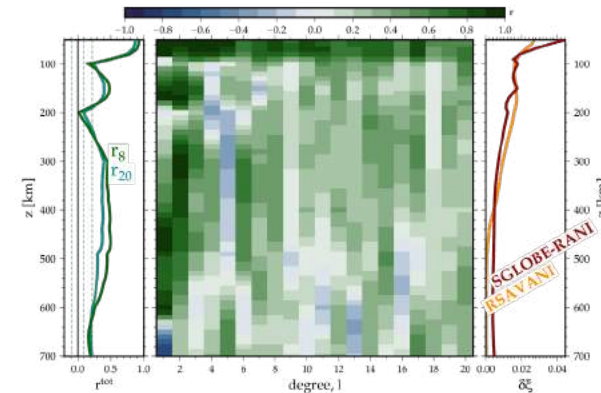
Chang et al. (2015)

## radial anisotropy correlation

RSAVANI vs. SEMUCB-WM1,  $\langle r_{20} \rangle = 0.28$ ,  $\langle r_8 \rangle = 0.34$

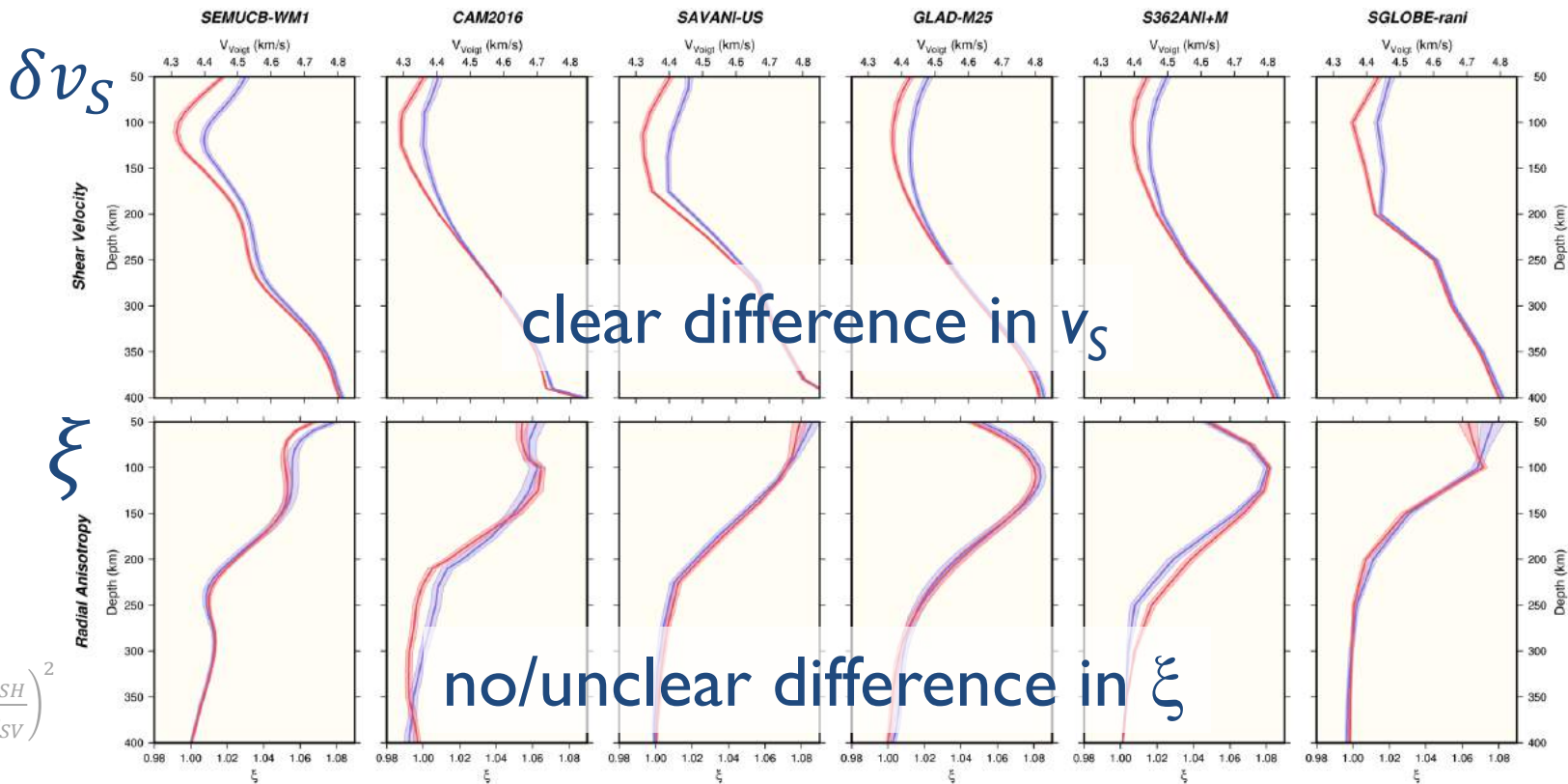


RSAVANI vs. SGLOBE-RANI,  $\langle r_{20} \rangle = 0.34$ ,  $\langle r_8 \rangle = 0.39$





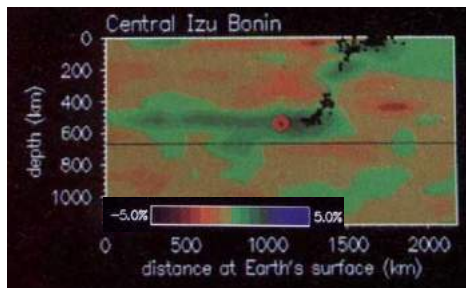
# Asthenospheric melt != low viscosity



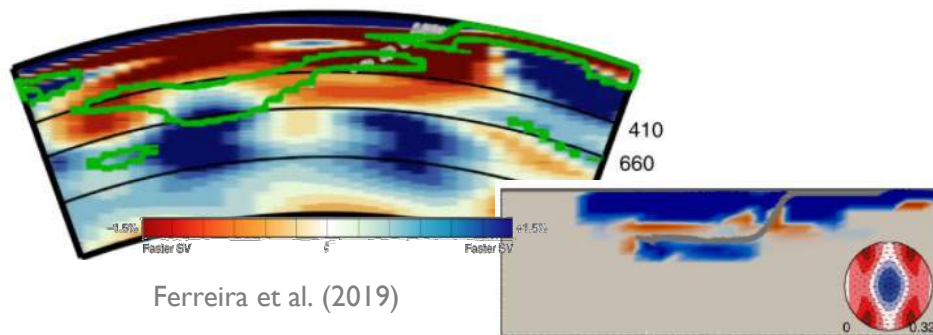
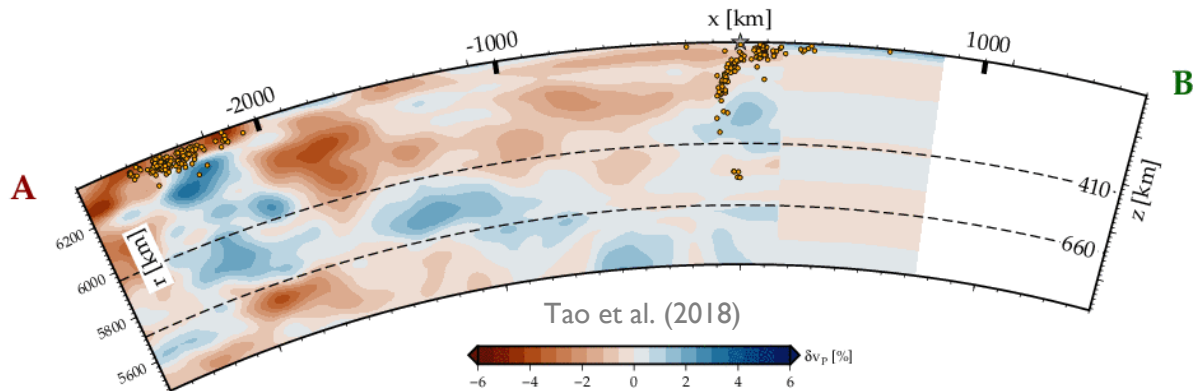
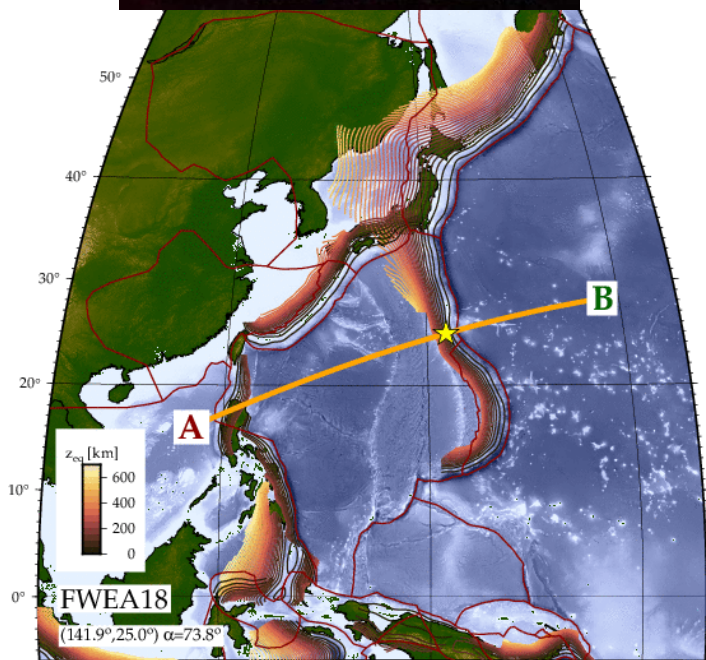
red = RF profile with melt phase

blue = RF profile without melt phase

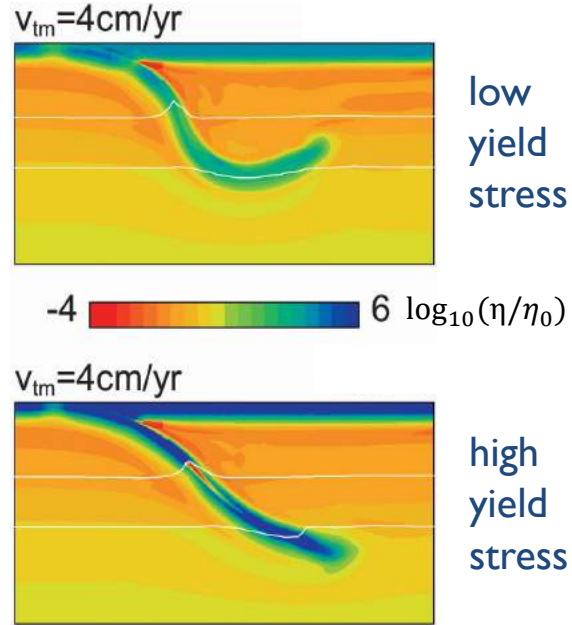
Hilst et al. (1991)



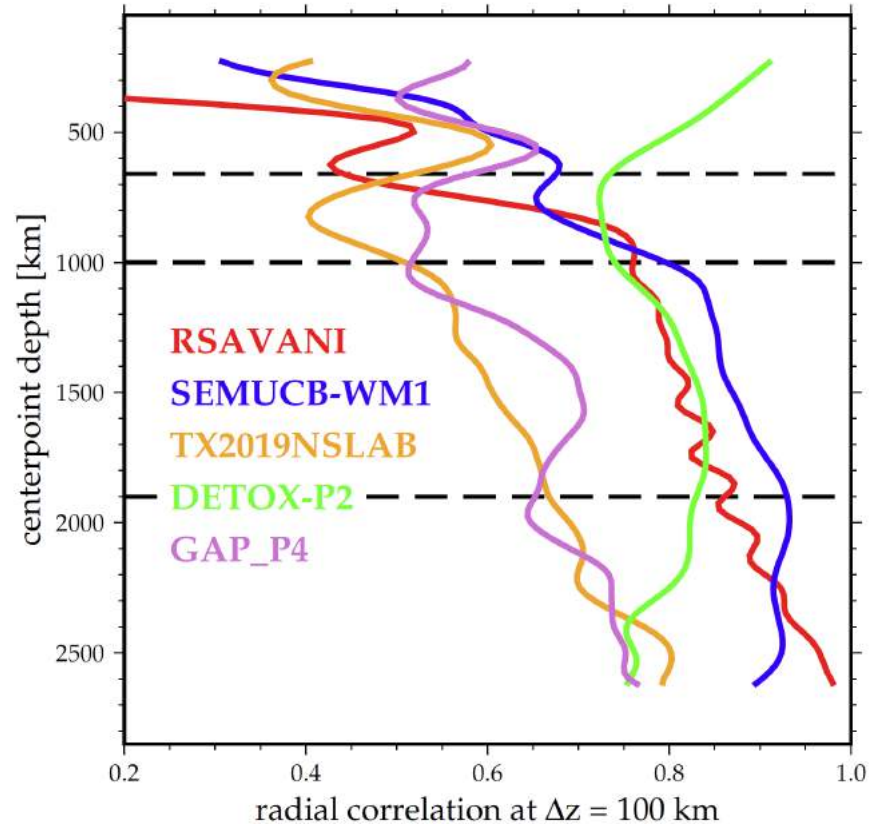
# Subduction engine



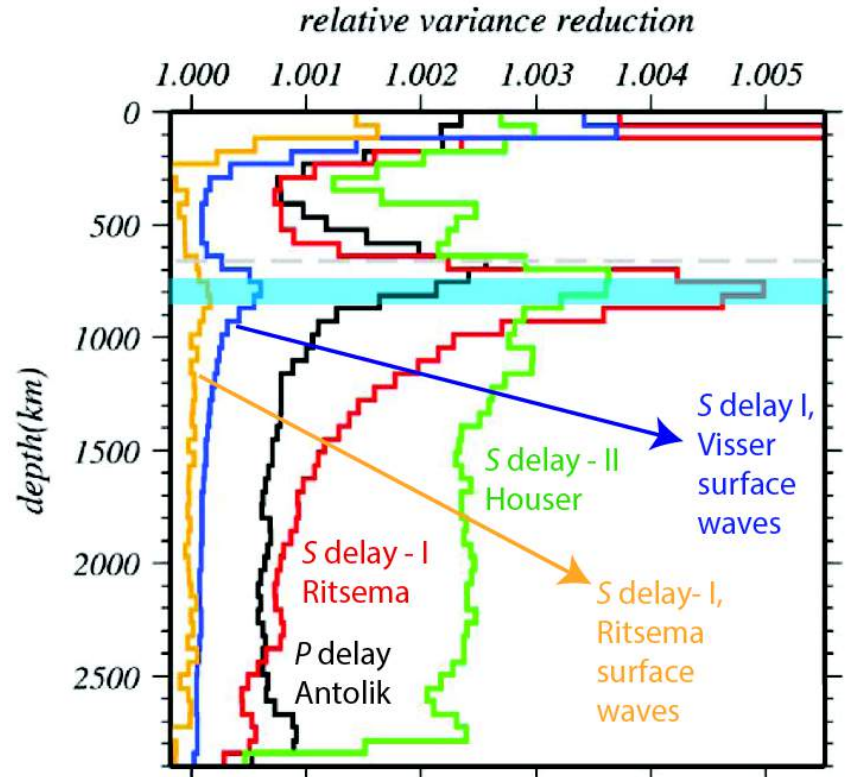
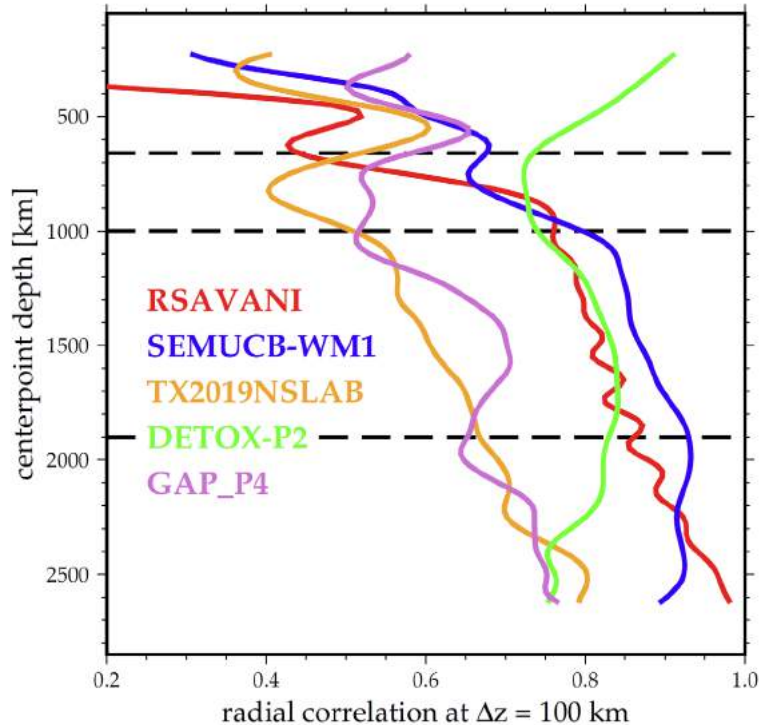
# Role of slab rheology for transition zone dynamics



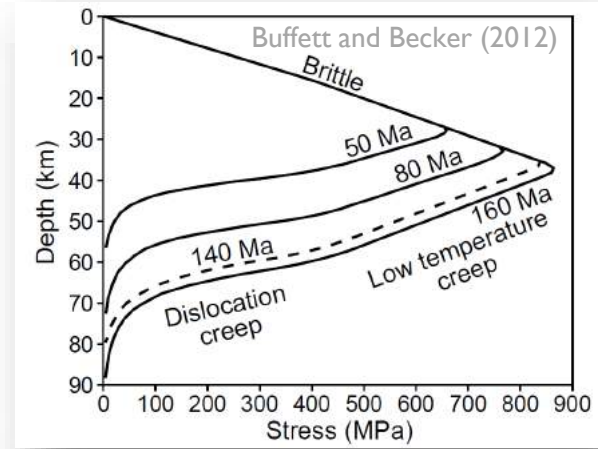
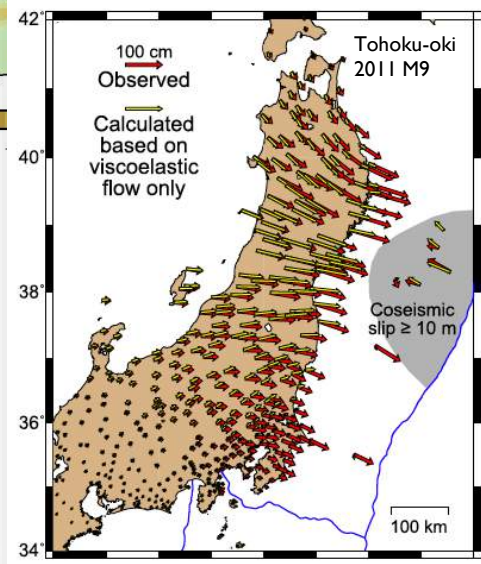
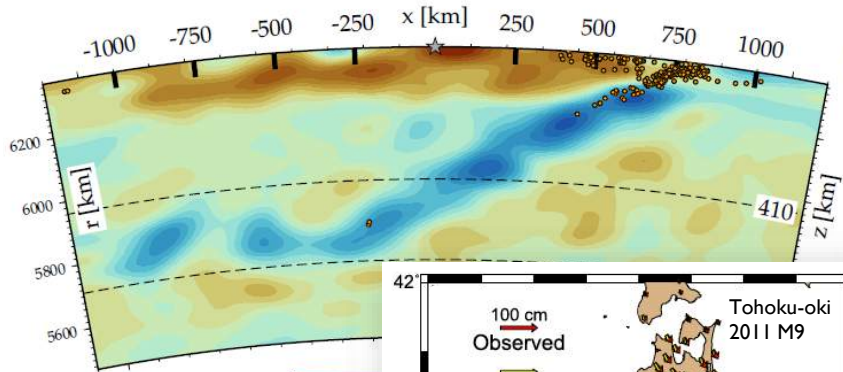
Čížková et al. (2002)



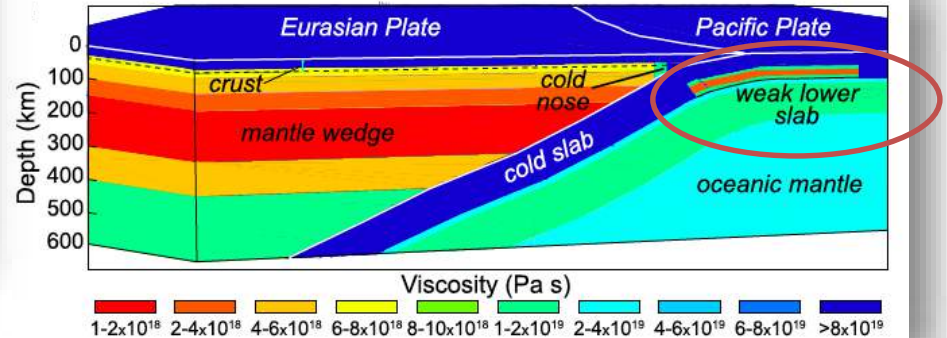
# Seismic data prefers decorrelation at ~800 km



# How are slabs weakened?

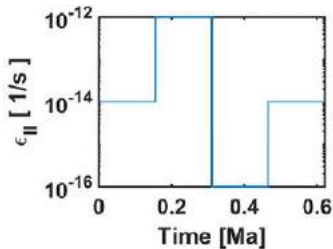


Freed et al. (2017)

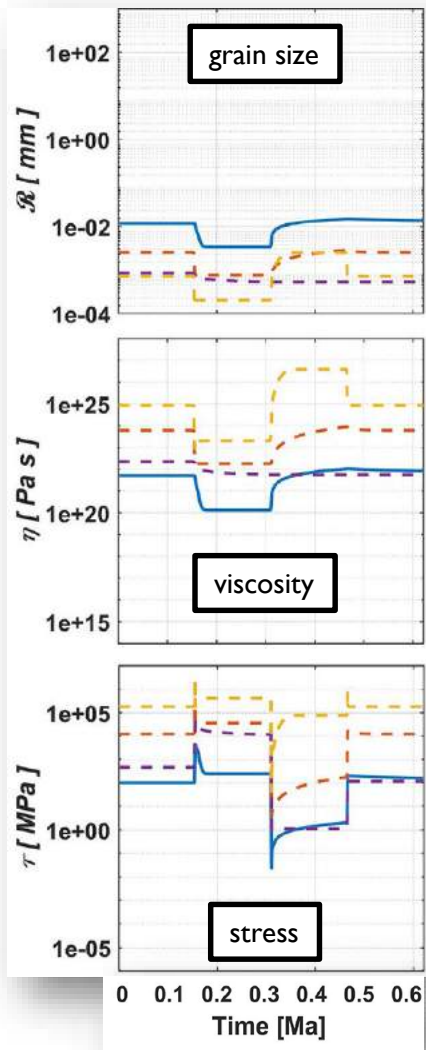


➤ Short timescale constraints on viscosity

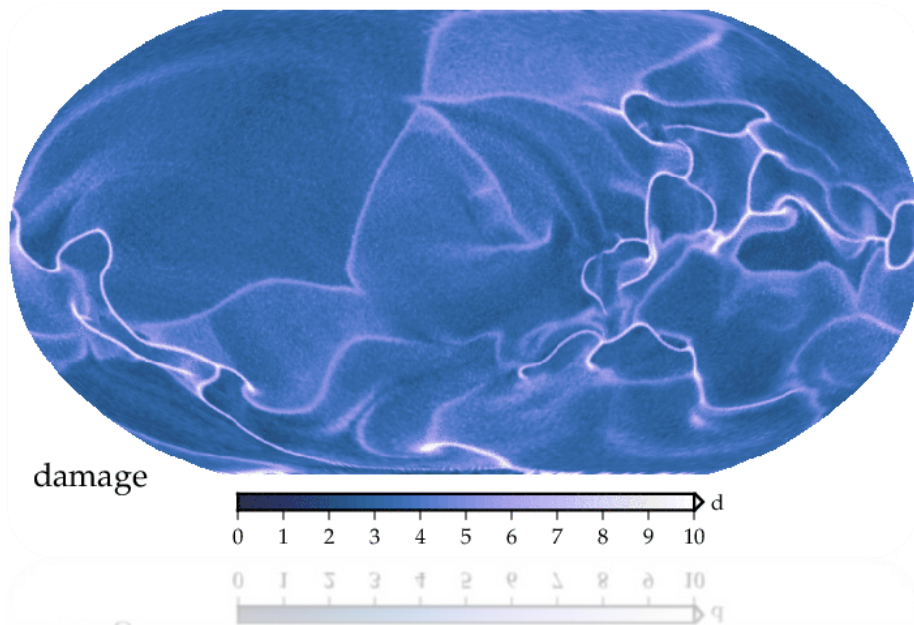
Slide-hold-slide test  
for different grain  
size evolution laws



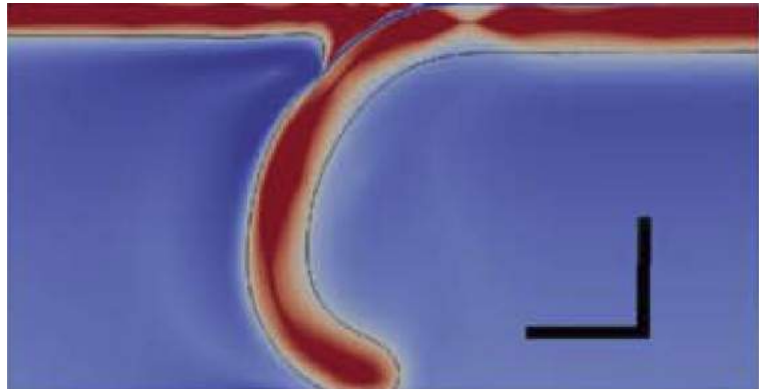
- Br99
- Be09
- Ro11
- Dan17



# Grain size evolution and ductile damage

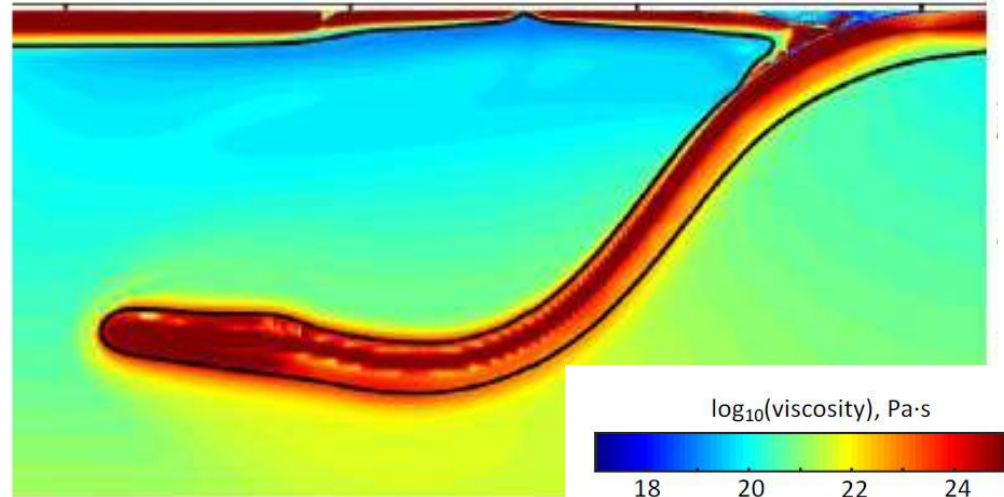


# Free subduction models with brittle and ductile damage



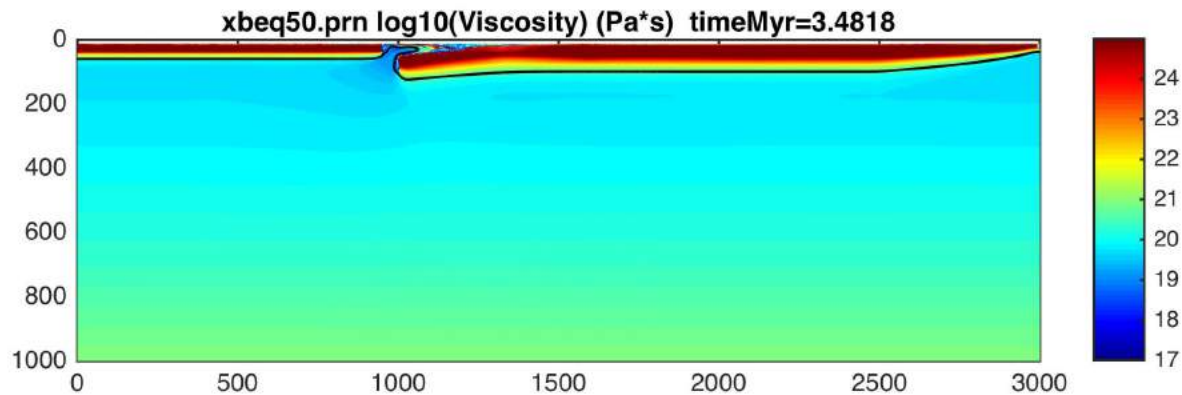
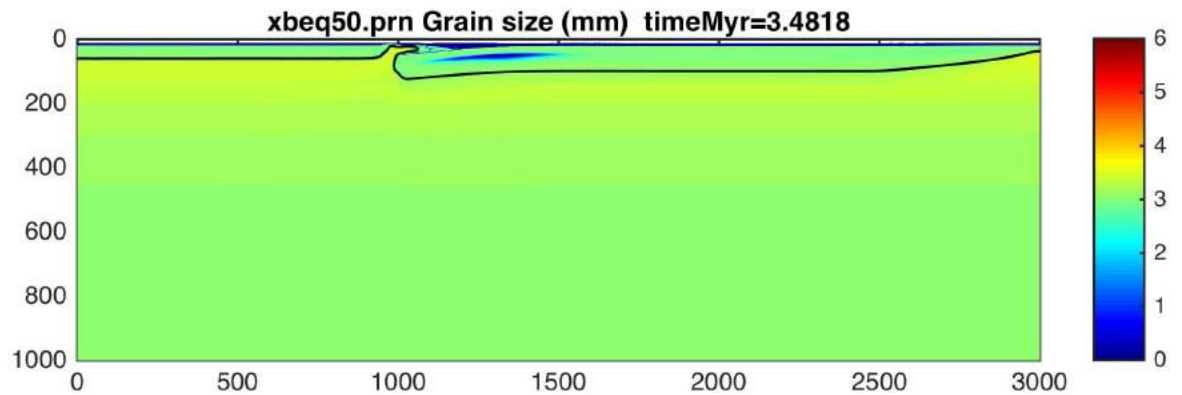
Garel et al. (2014)

Viscosity (Pa.s)

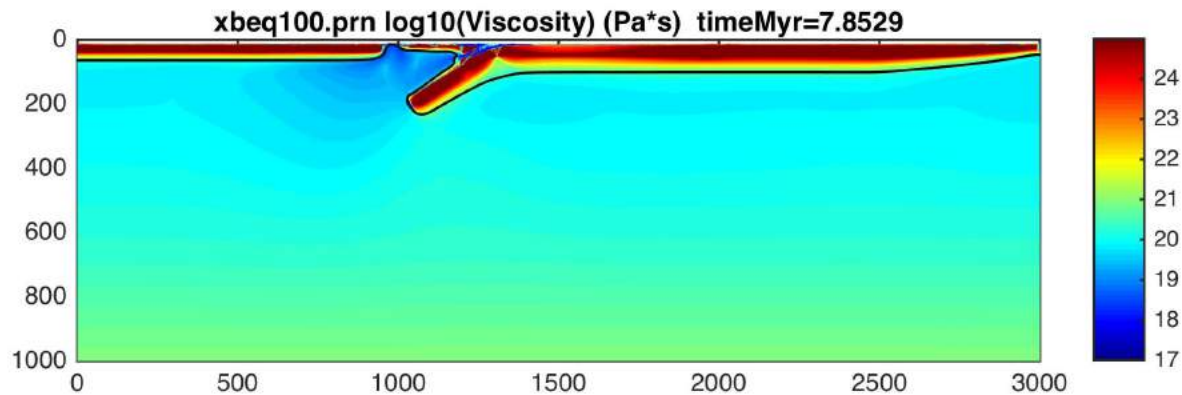
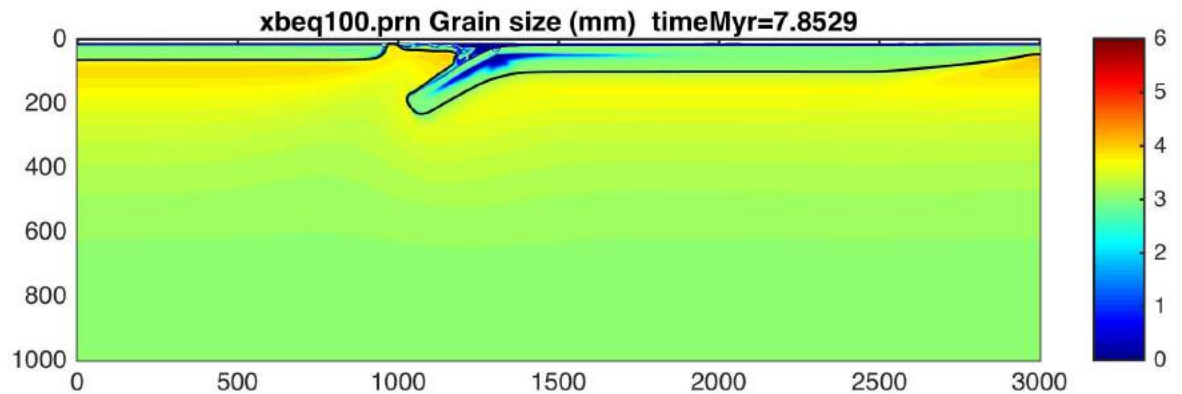


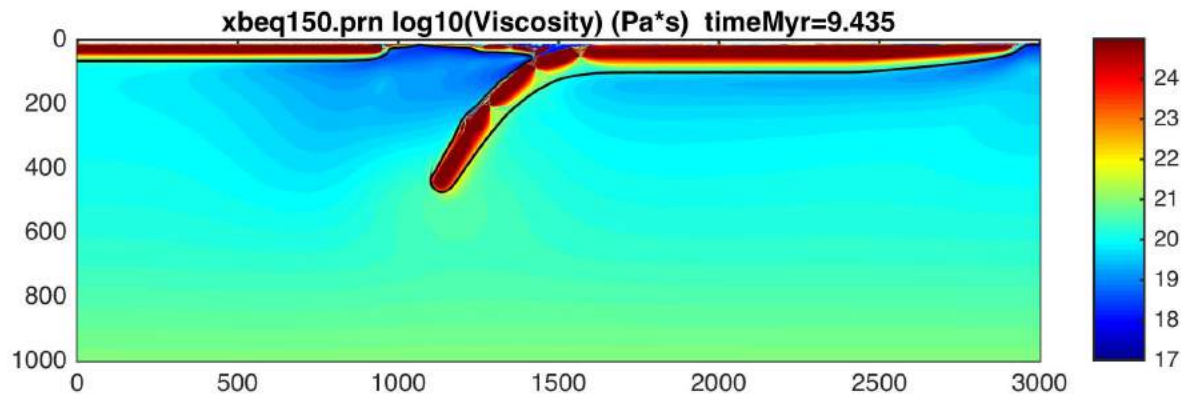
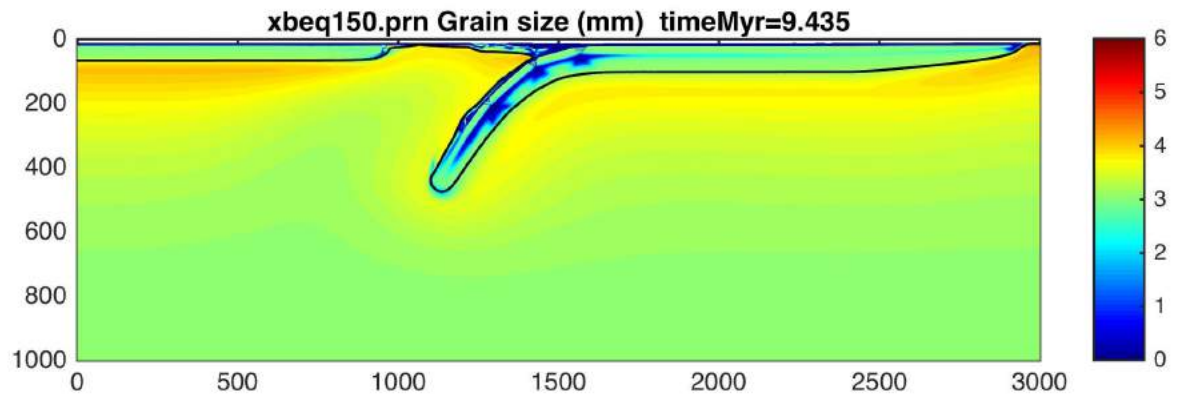
- Brittle domain damage by fault weakening
- Ductile domain grain-size evolution following Rozel, Bercovici et al. including Zener pinning

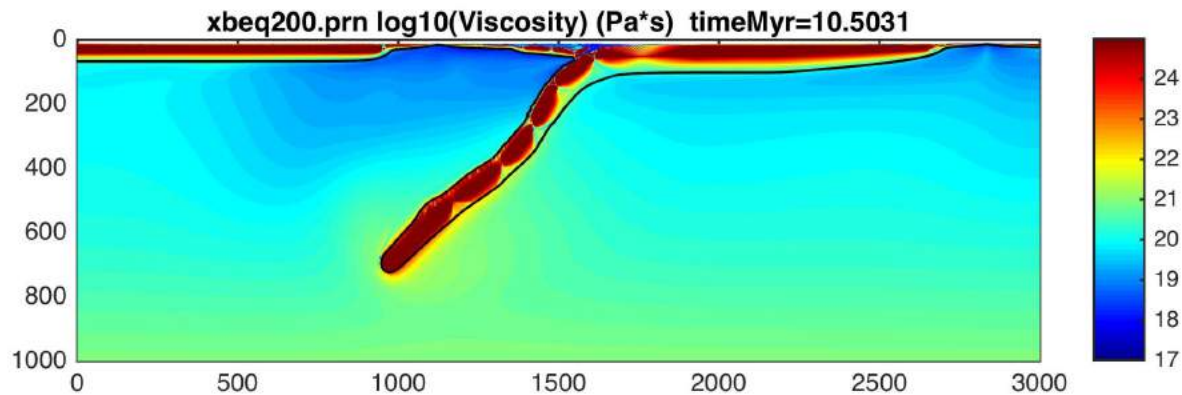
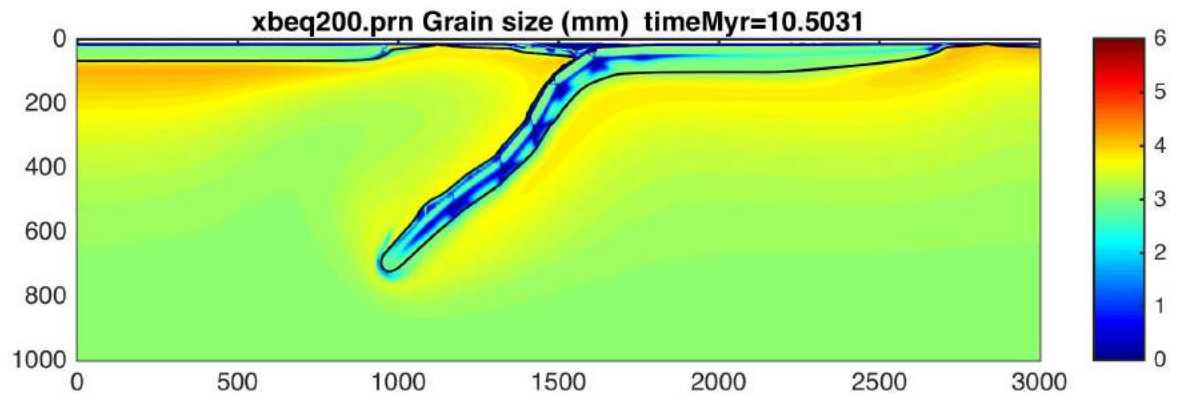
**plate age = 40 Myr, initial grainsize = 3 mm**



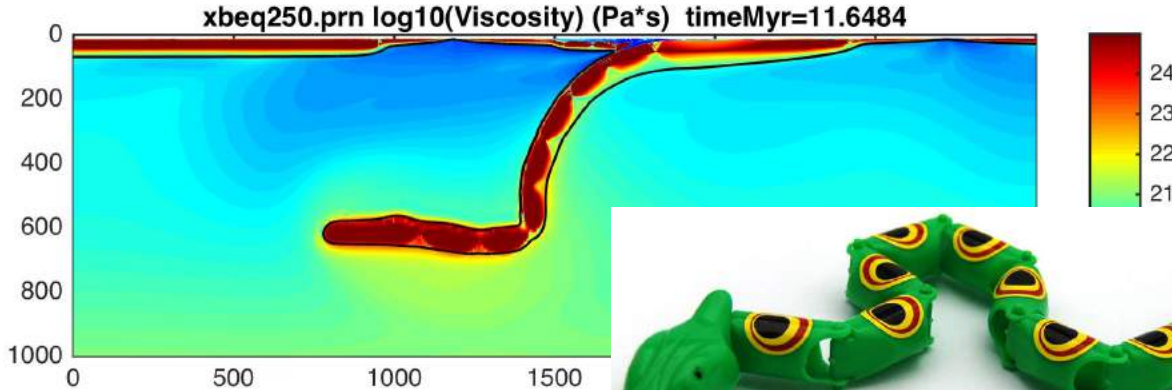
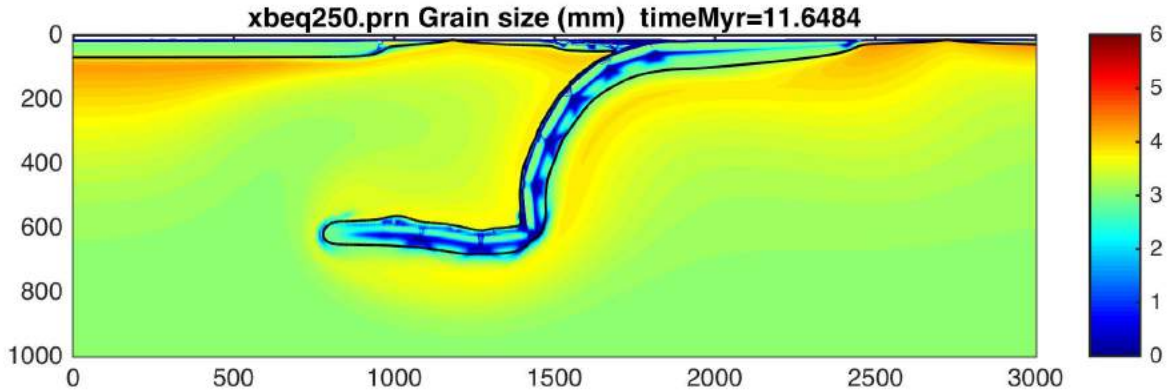








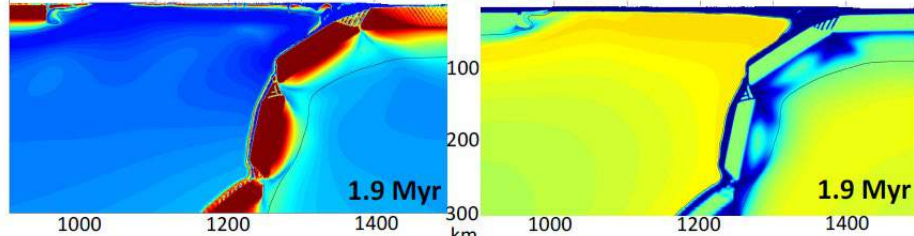
➤ Segmentation of slab facilitates bending and slab contortion, but allows for efficient force transmission



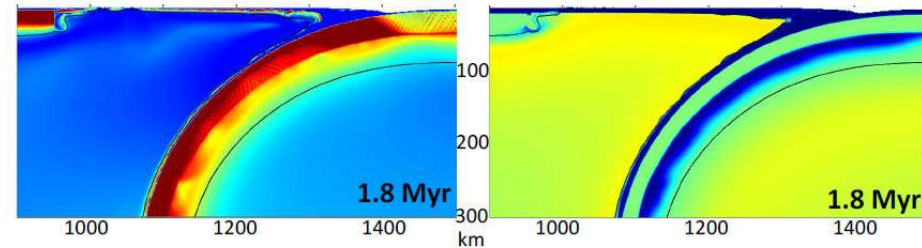
# Interaction between brittle and ductile damage required for segmentation

**Reference model:**

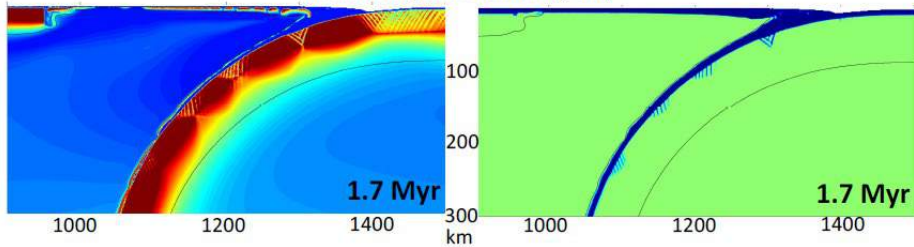
*both faults weakening and grain size evolution*



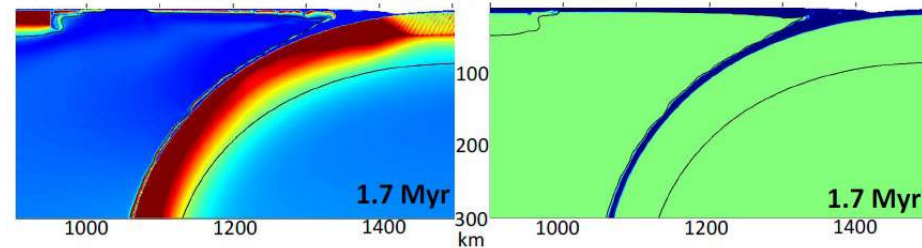
*model with grain size evolution only*



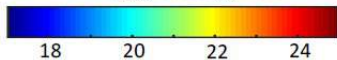
*model with faults weakening only*



*model with neither faults weakening nor grain size evolution*



$\log_{10}(\text{viscosity}), \text{Pa}\cdot\text{s}$



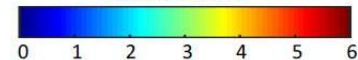
grain size, mm



$\log_{10}(\text{viscosity}), \text{Pa}\cdot\text{s}$

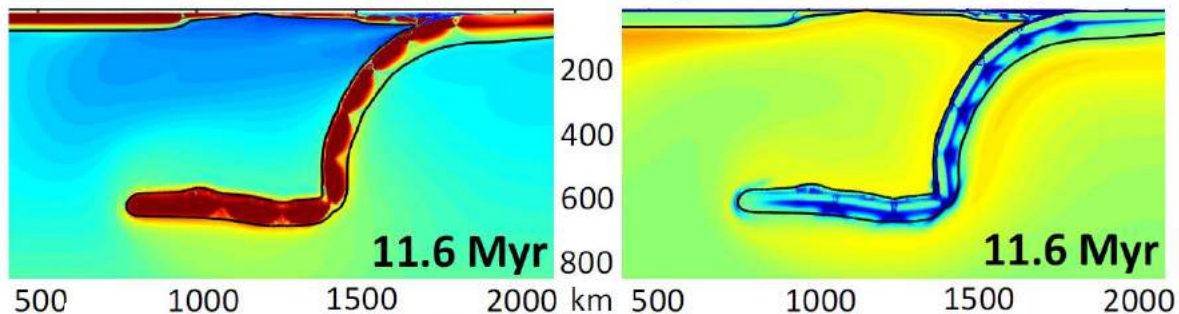


grain size, mm

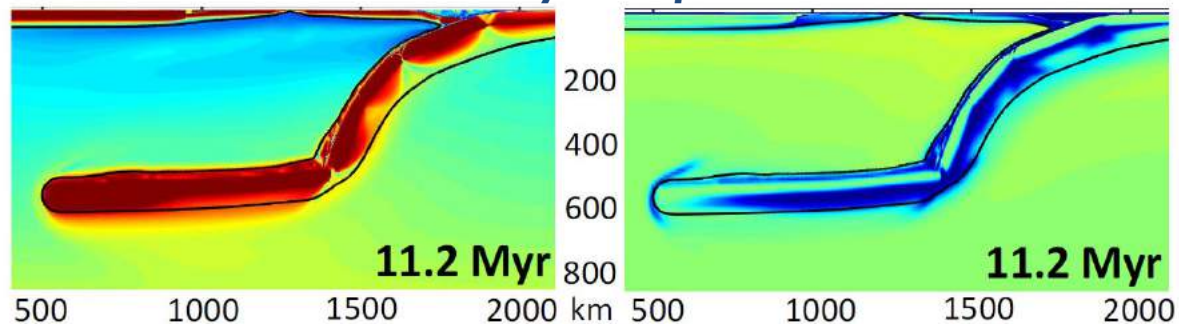


# Plate age controls spacing of segments

*40 Myr old plate*



*100 Myr old plate*



$\log_{10}(\text{viscosity}), \text{Pa}\cdot\text{s}$



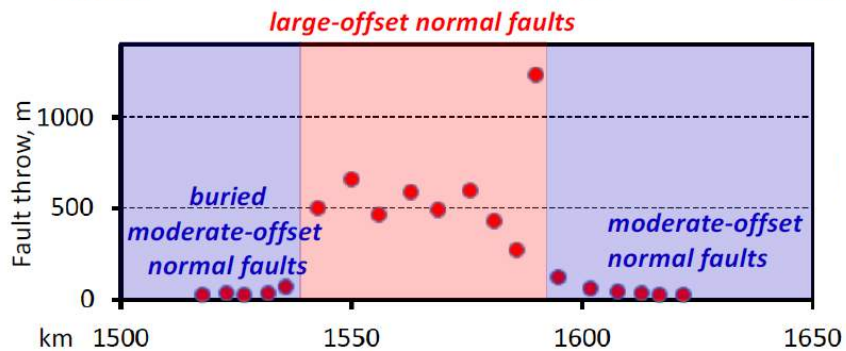
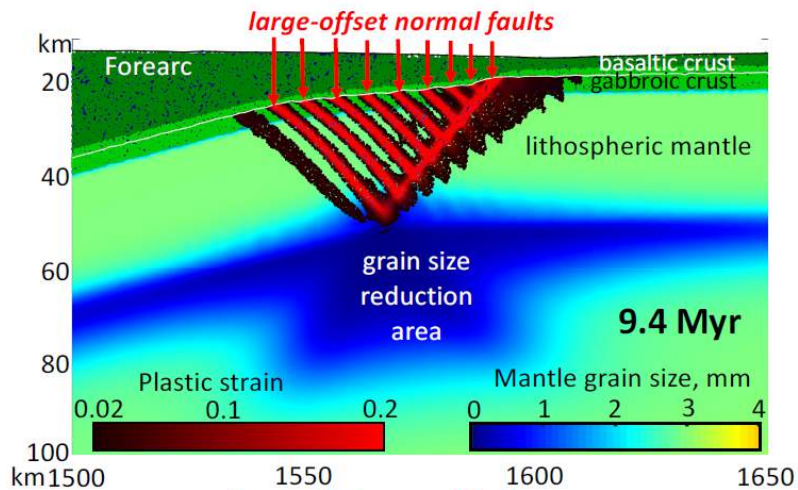
18 20 22 24

grain size

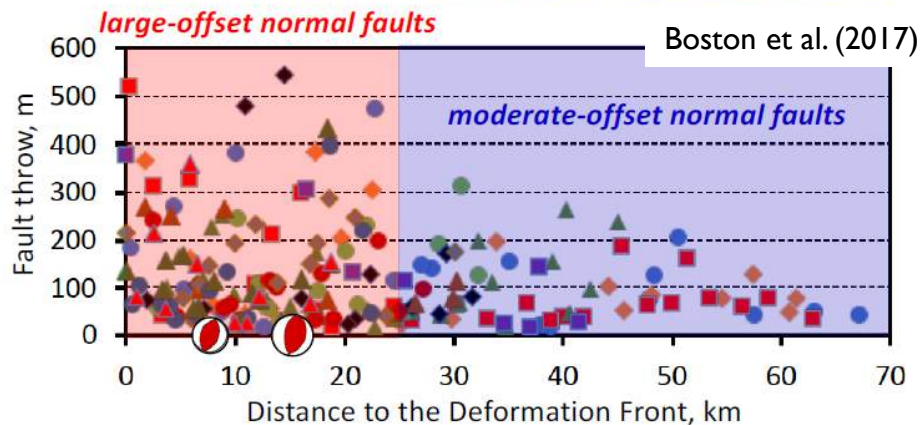
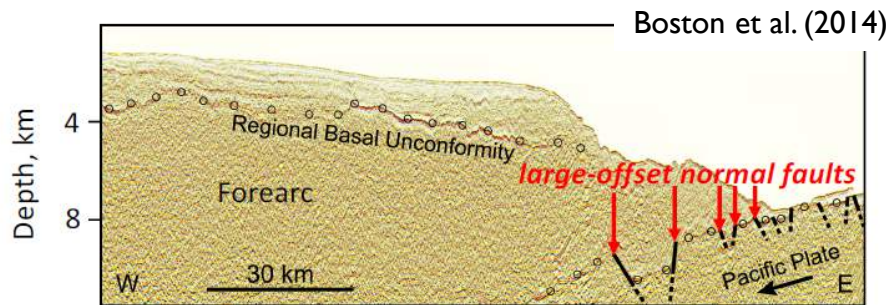


0  $h_{\max}$

# Large offset normal faults in bending region



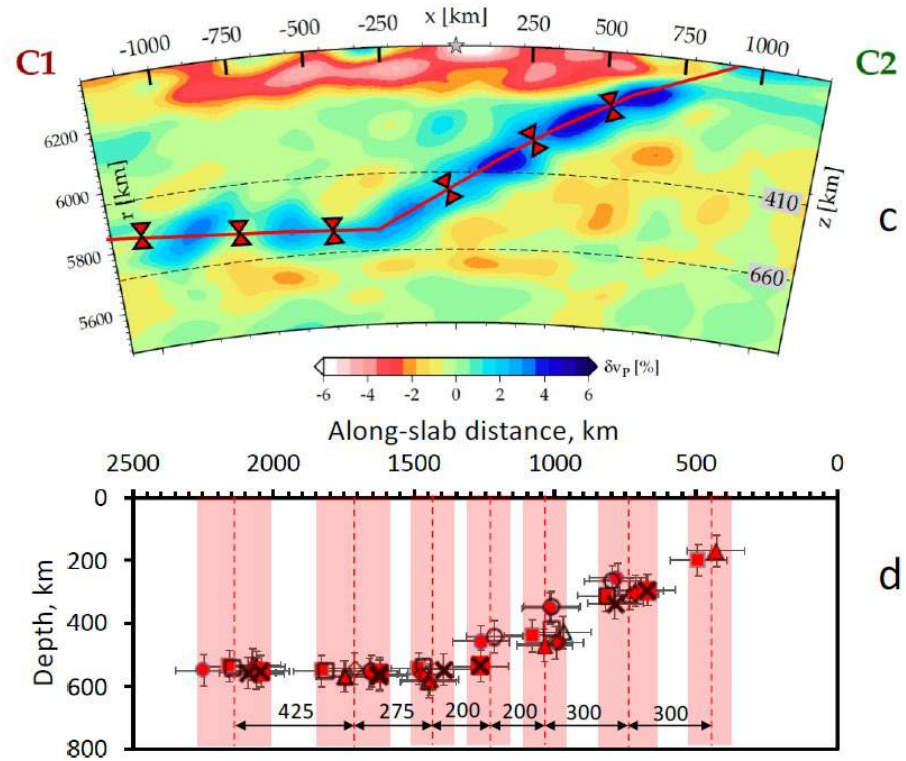
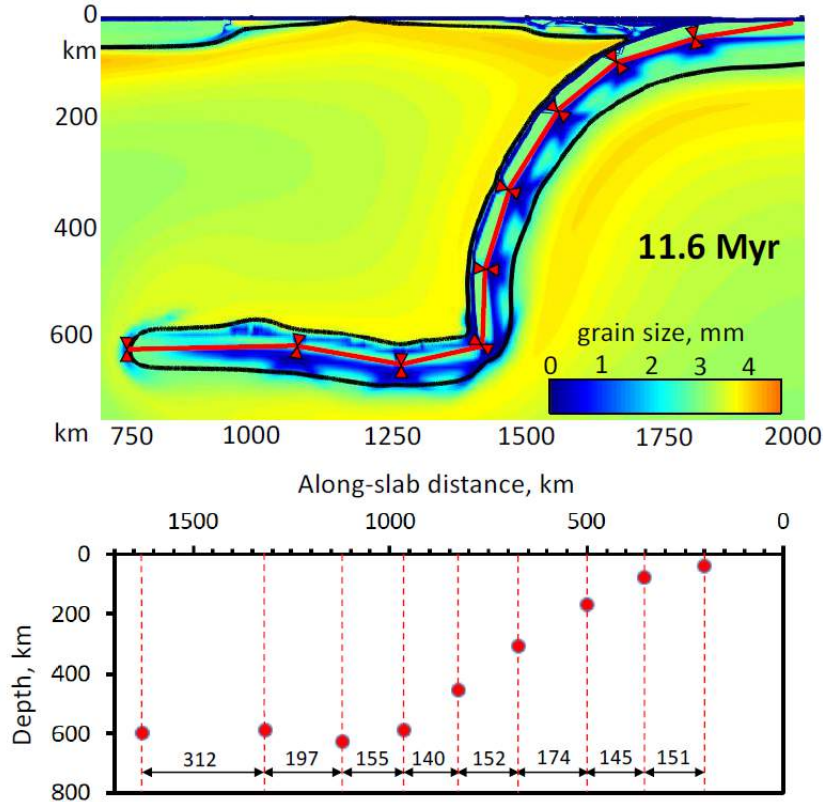
## Japan trench





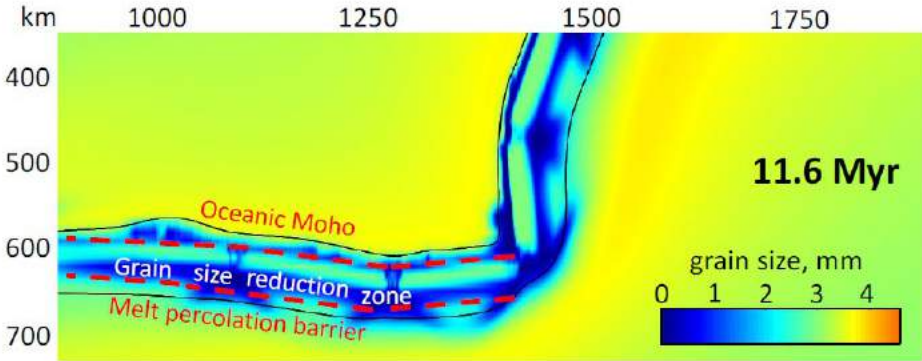
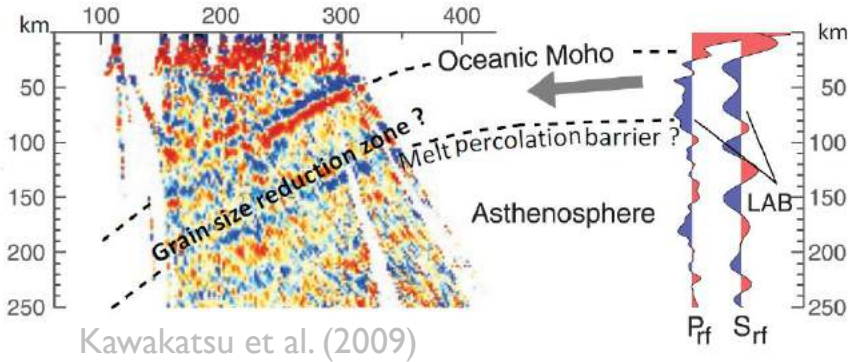
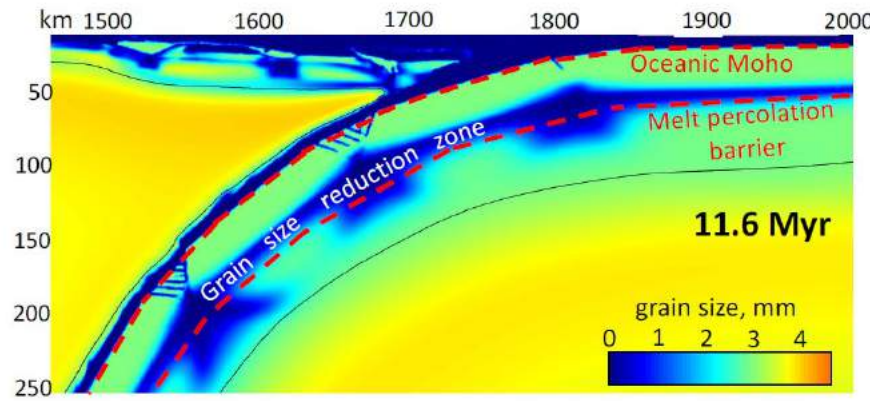


# Comparable segmentation signatures?

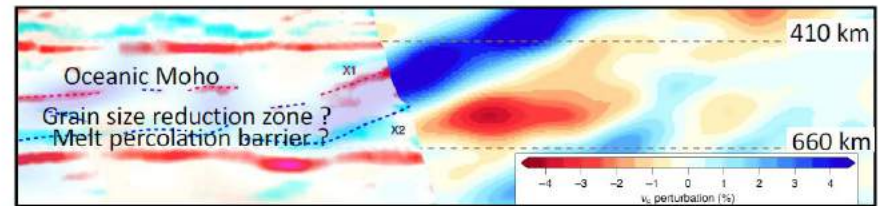


# Seismic interface signature

- Melt as marker: control on seismic wave speed but not viscosity



Wang et al. (2020)  
Tao et al. (2018)



# Slabs are segmented due to coupled brittle-ductile damage



Slab segmentation explains:

1. dichotomy of strong plates and weak slabs
2. large-offset normal faults near trenches
3. segmented slab seismic velocity anomaly
4. low-viscosity region below the outer rise

# Conclusions

- Global geodynamics/seismology models allow hypothesis testing
- Regional observations holds important clues for refinement
- Melt effects secondary for viscosity
- Full waveform hypothesis testing promising