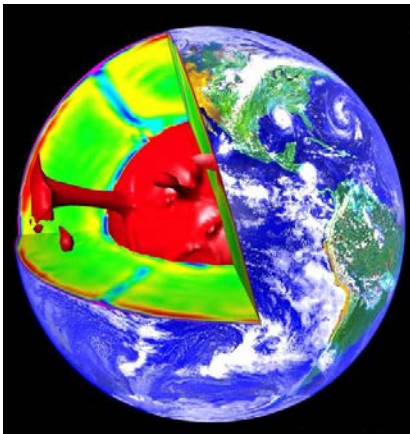


Dynamics and Structure of the Mantle

Paul J. Tackley
(ETH Zürich)





Mantelkonvektion mit Plattentektonik und Kontinentaldrift auf der Erde
Mantle Convection with Plate Tectonics and Continental Drift on Earth

Movie by:

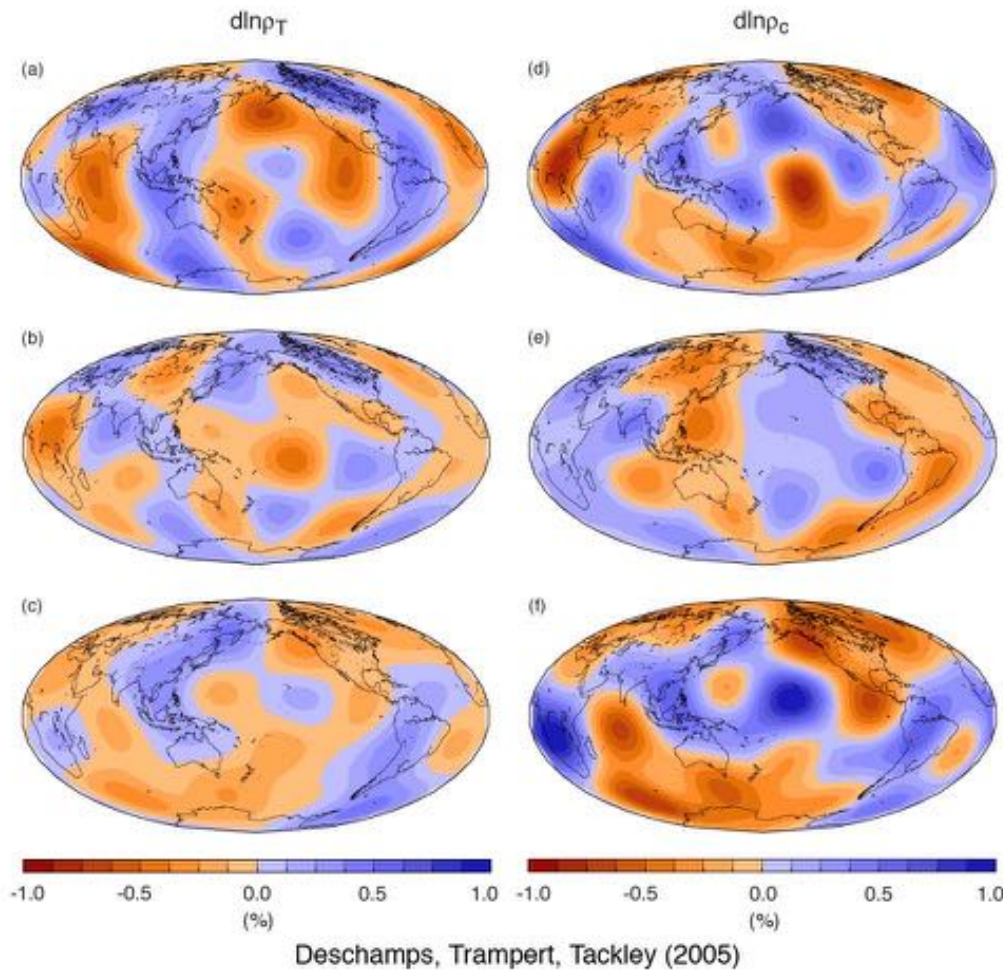
Tobias Rolf, Antoine Rozel, Paul Tackley

<https://gfd.ethz.ch>

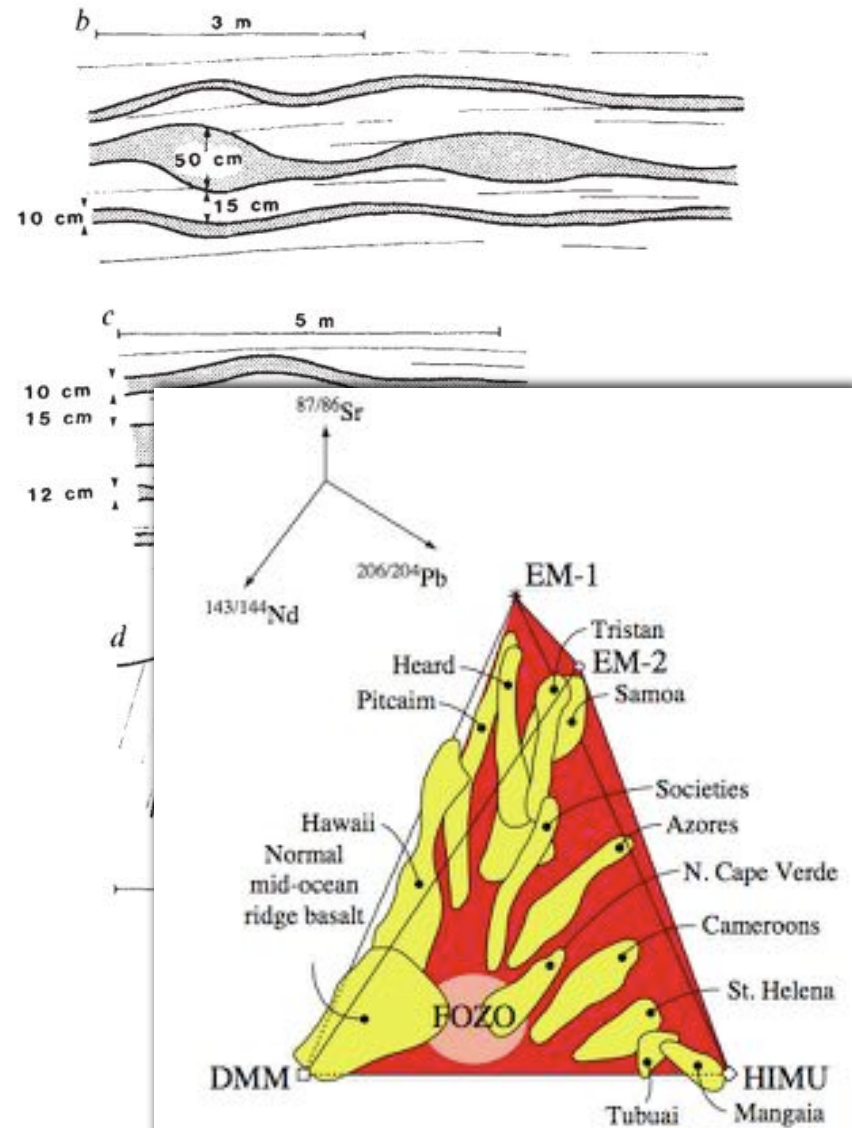
Overview

- Concepts & cartoon models
- Tectonics: recent and early Earth
- Recycled crust & Earth evolution
- Importance of compositional viscosity contrast on deep mantle structure & dynamics

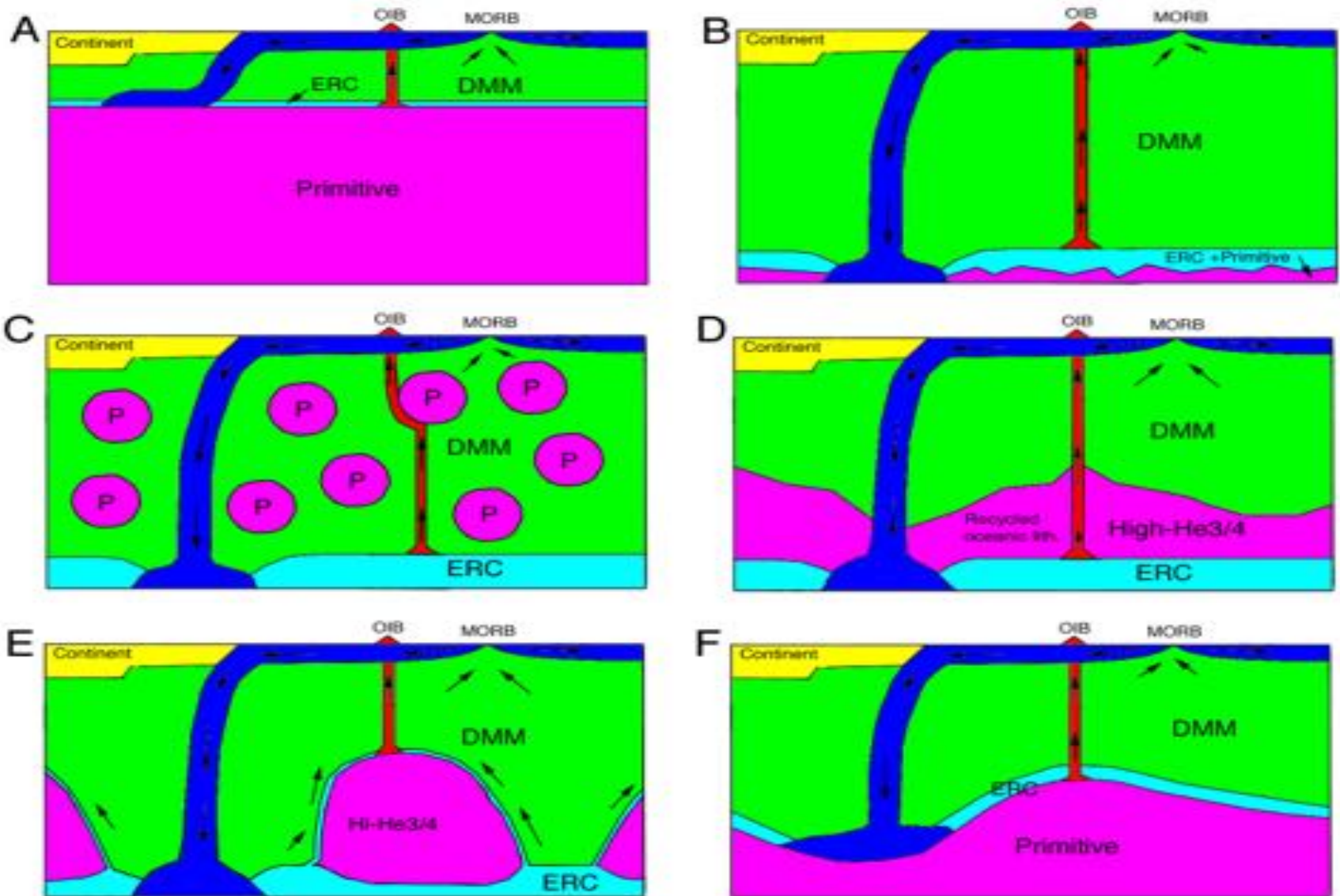
Compositional variations exist at all scales! ...and are a result of (partial) melting



Deschamps, Trampert, Tackley (2005)



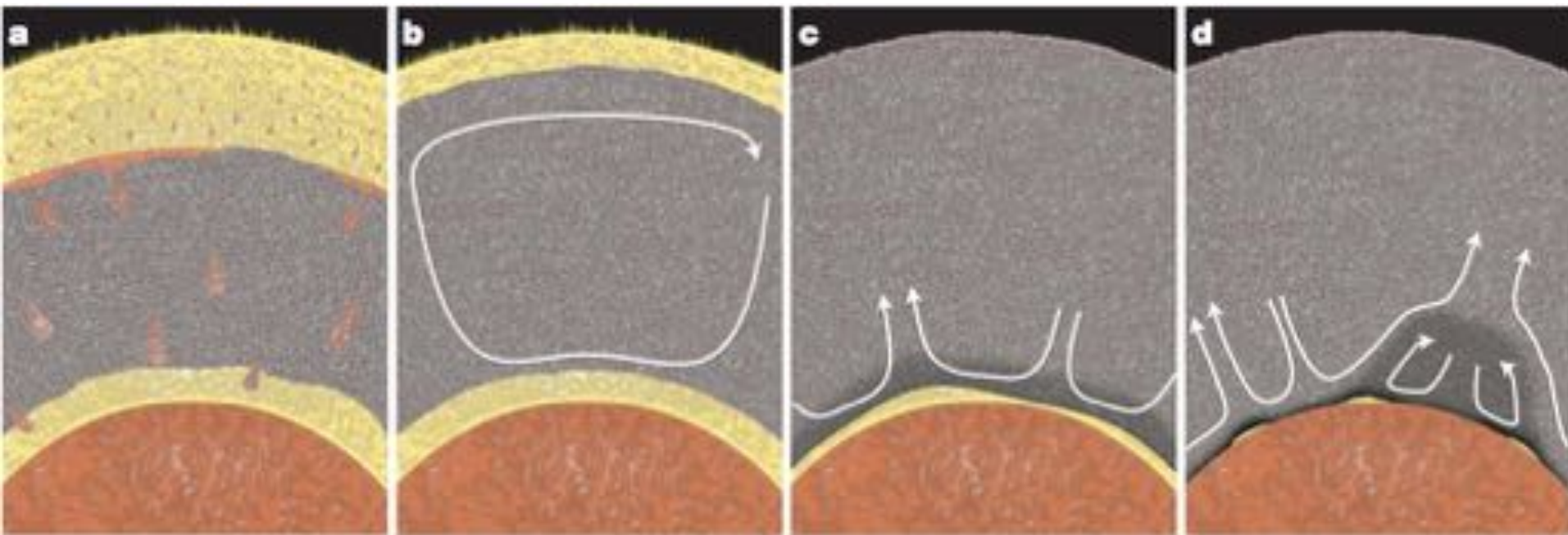
Geochemical mantle: Old cartoons (2000)



from Tackley, Science, 2000: Figure 2

Long-term persistence of melting: Basal Magma Ocean

Labrosse et al., 2007

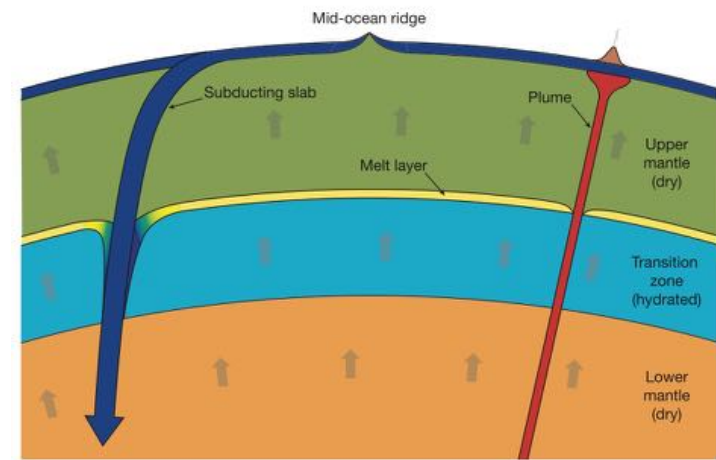


Early Earth

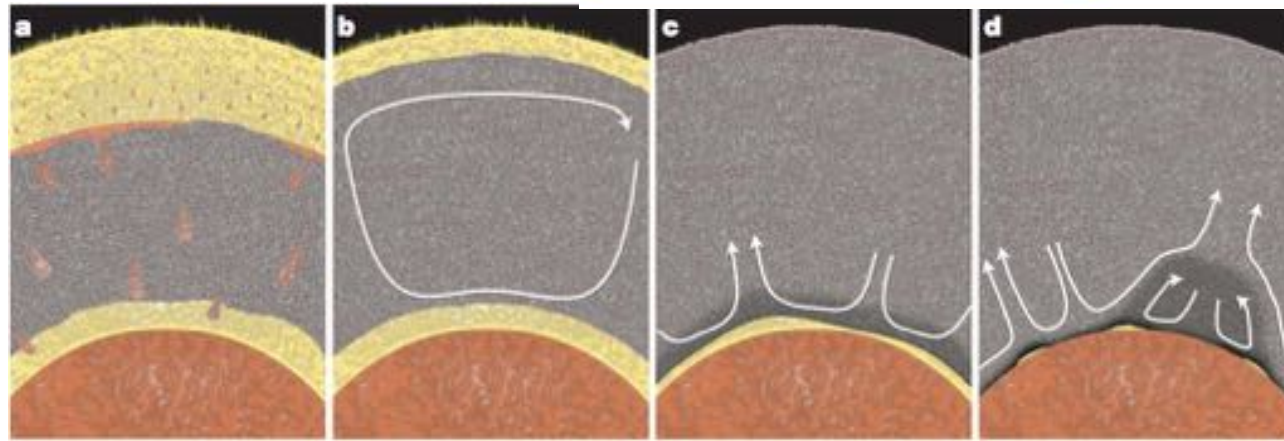
Present day

Deep melting: cartoon models

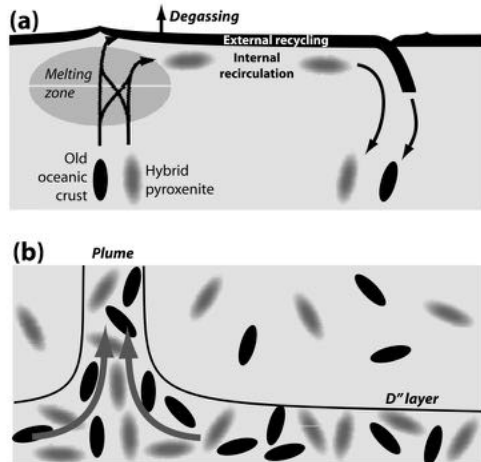
Transition Zone Water Filter
Bercovici & Karato 2003



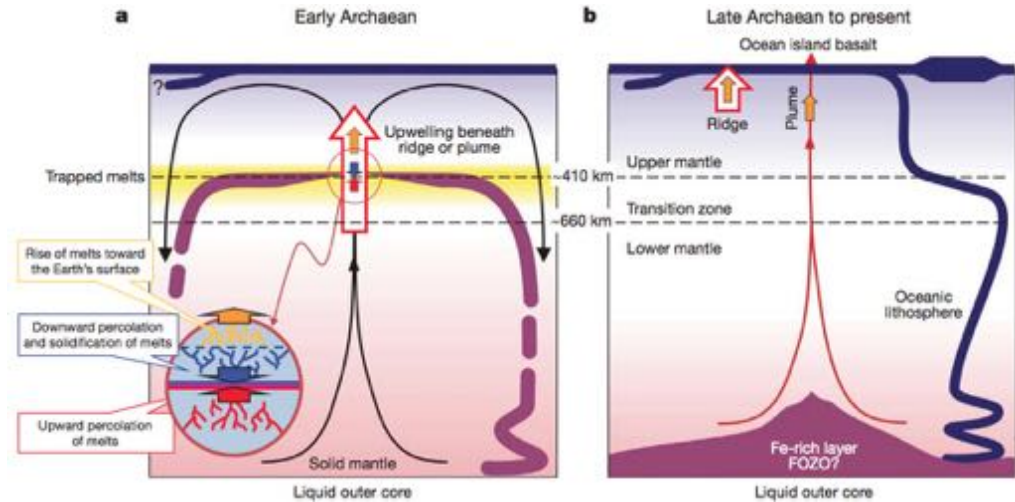
Basal Magma Ocean
Labrosse et al., 2007



Davies 2009

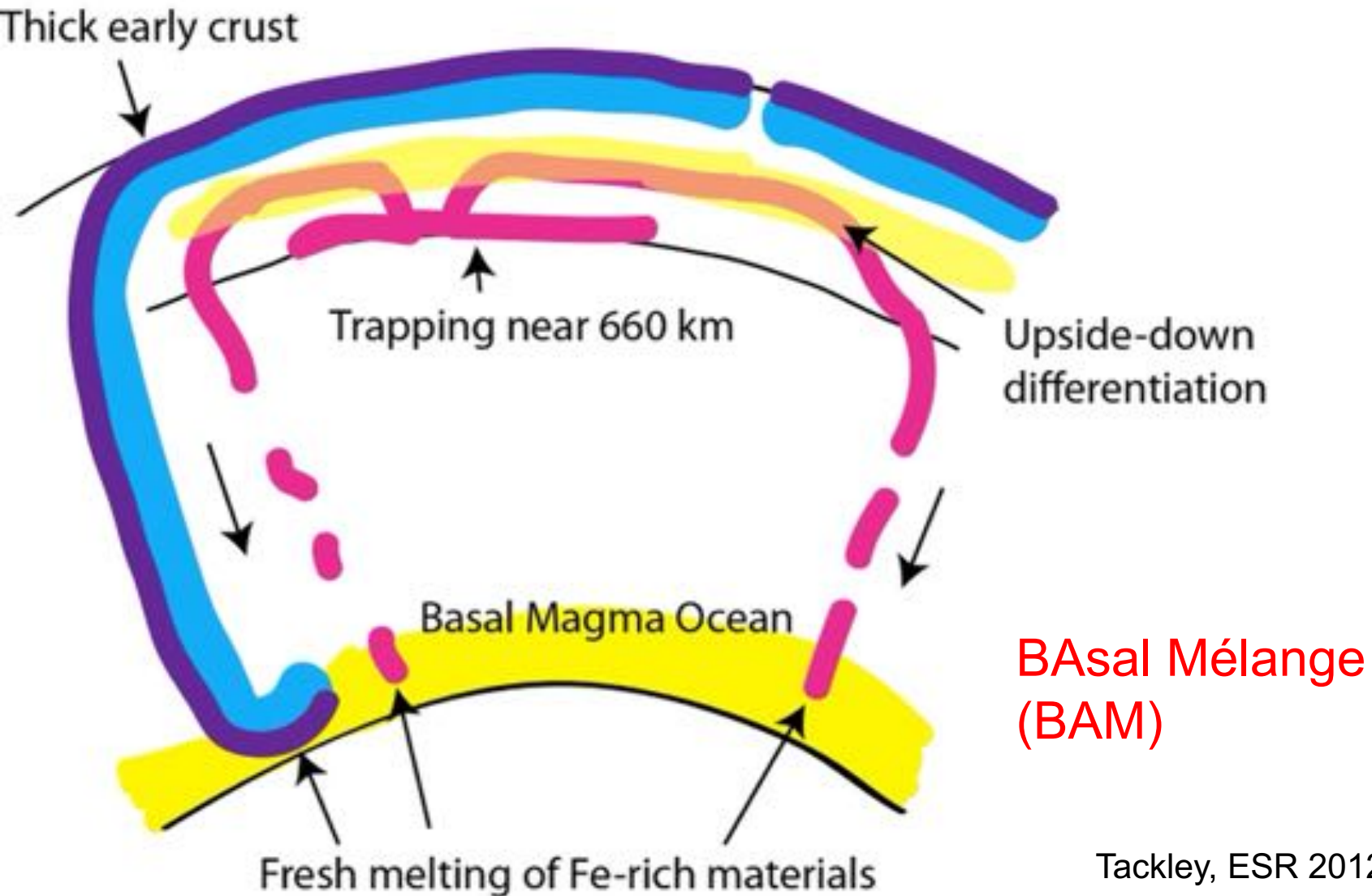


Upside-down differentiation
Lee et al 2010



More than one process operating!

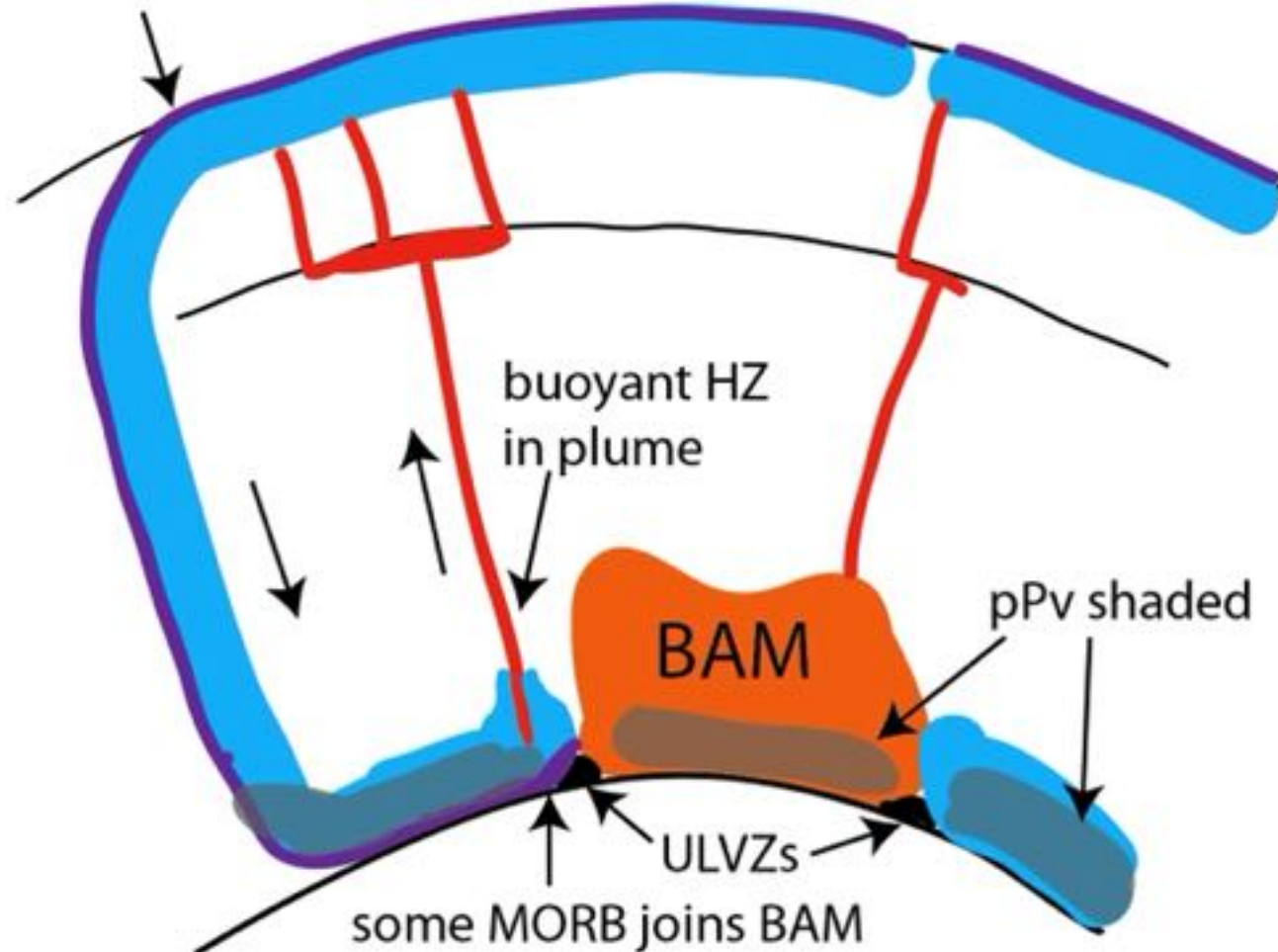
a. Early Earth



More than one process operating!

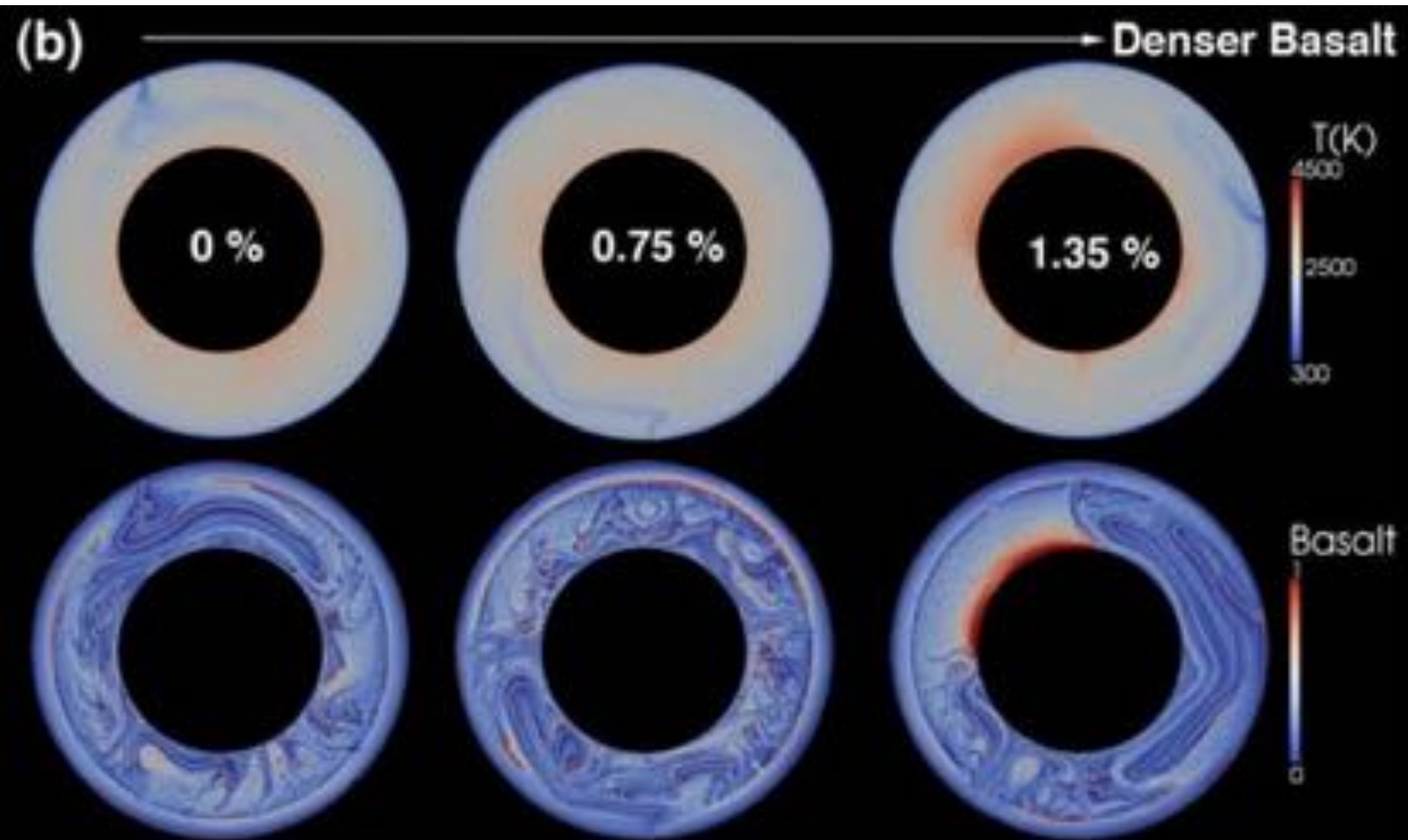
b. Present day

Thin oceanic crust



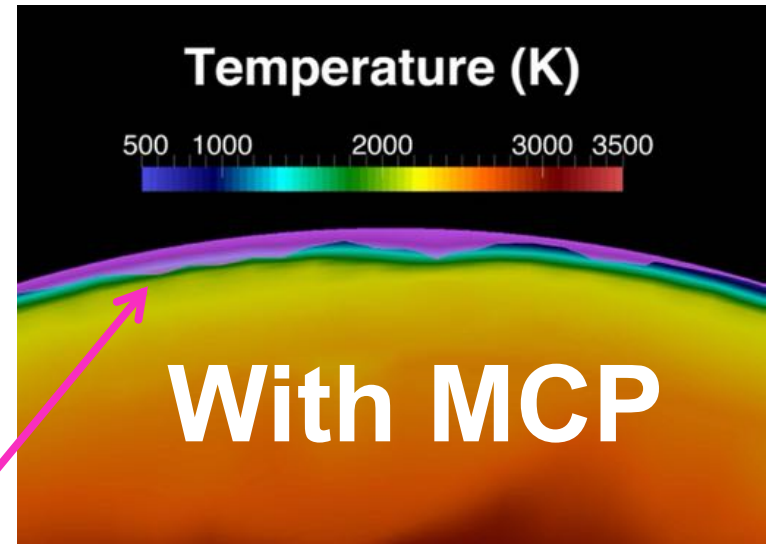
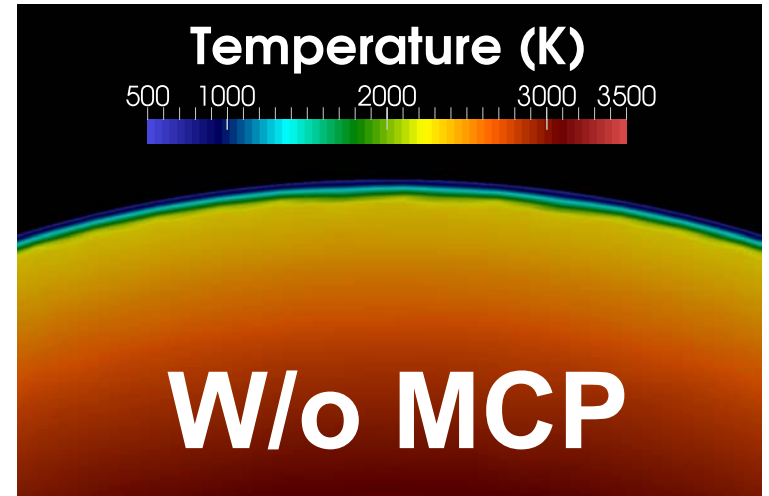
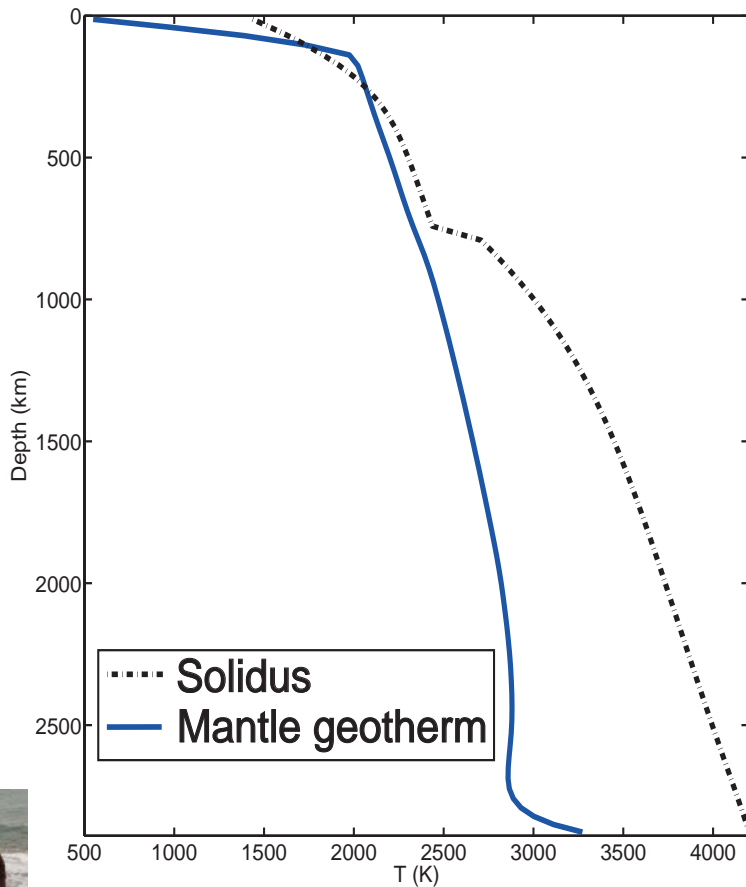
Basal Mélange
(BAM) mix:
BMO remnants
UM differentiated products
Recycled crust
...

Production of oceanic crust by partial melting



Numerical and physical model

Melting-induced crustal production (MCP)



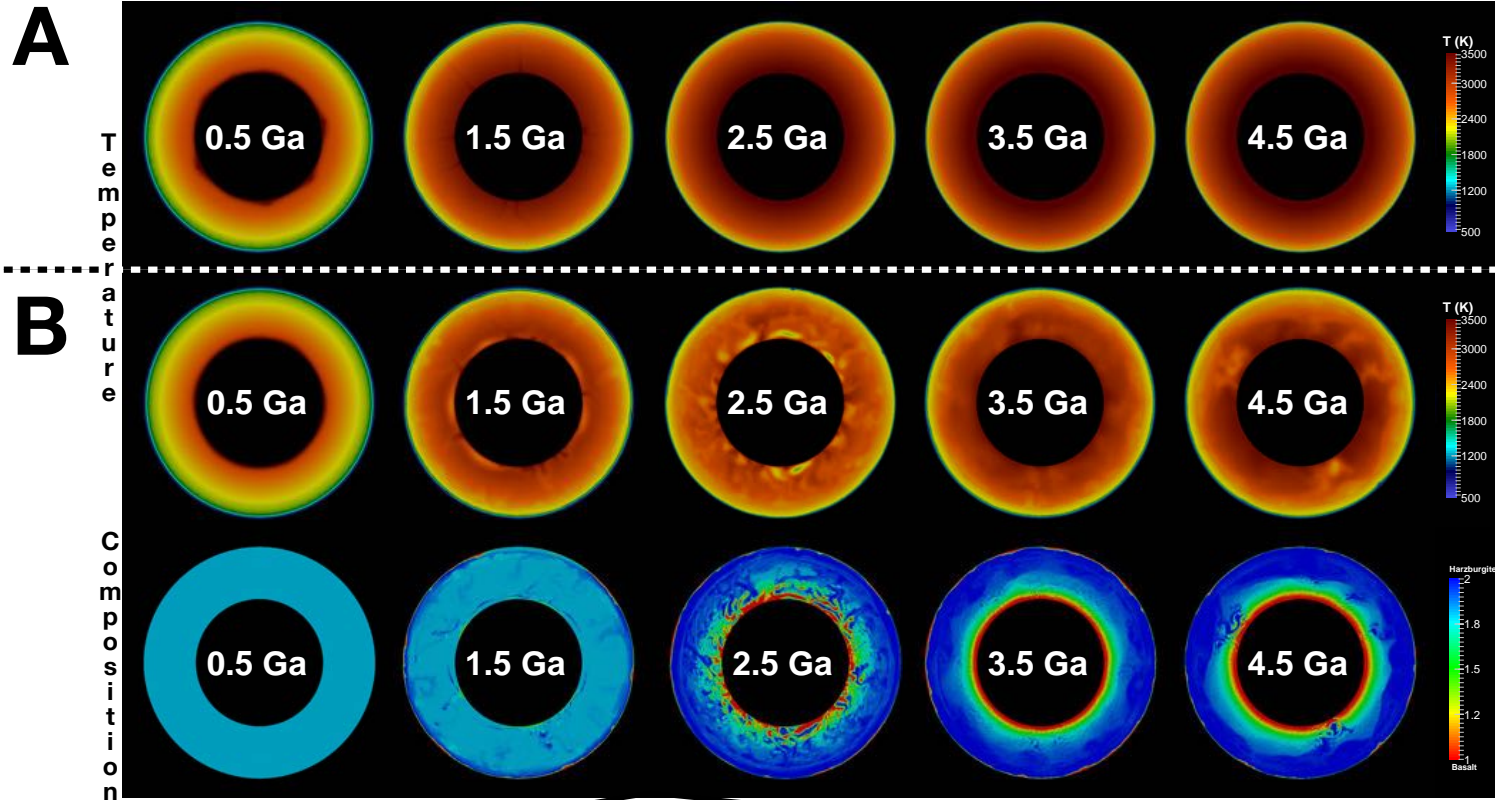
Diogo Lourenco

Basaltic crust

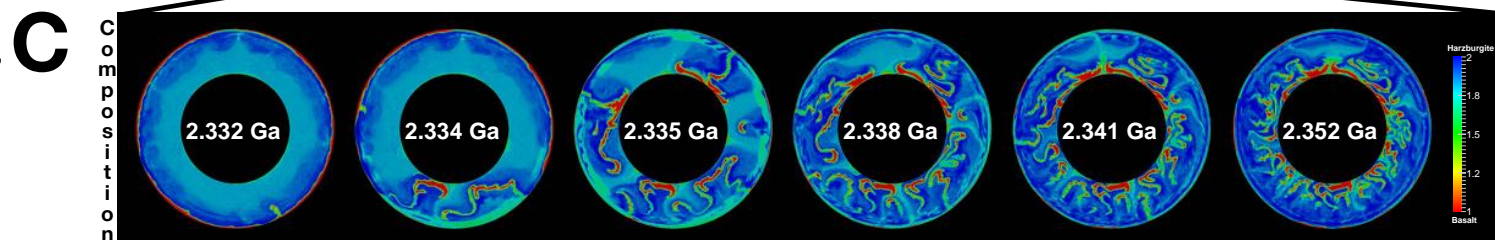
Magmatism->crust helps plate tectonics

Purely thermal
-> Stagnant

With magma &
crust
Episodic plate
tectonics



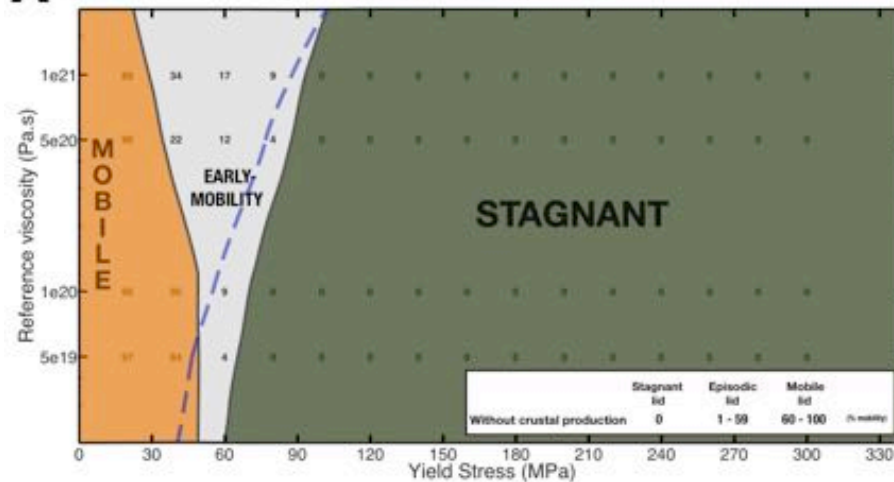
Diogo Lourenco A. Rozel & Tackley,
EPSL 2016



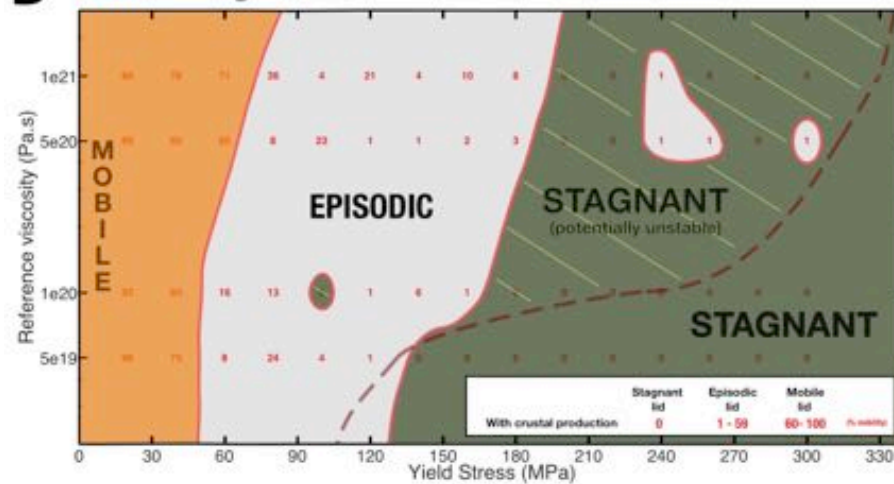
Melting +
 crustal
 production
 makes
 stagnant lid
 less likely

Lourenco et al.,
 EPSL 2016

A - Without Melting and Crustal Production

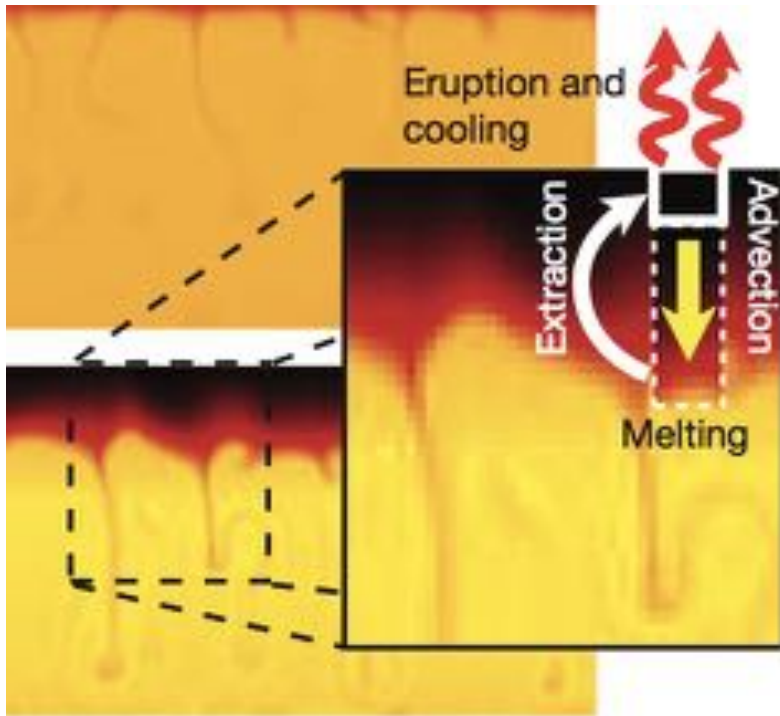


B - With Melting and Crustal Production



08/04/2014

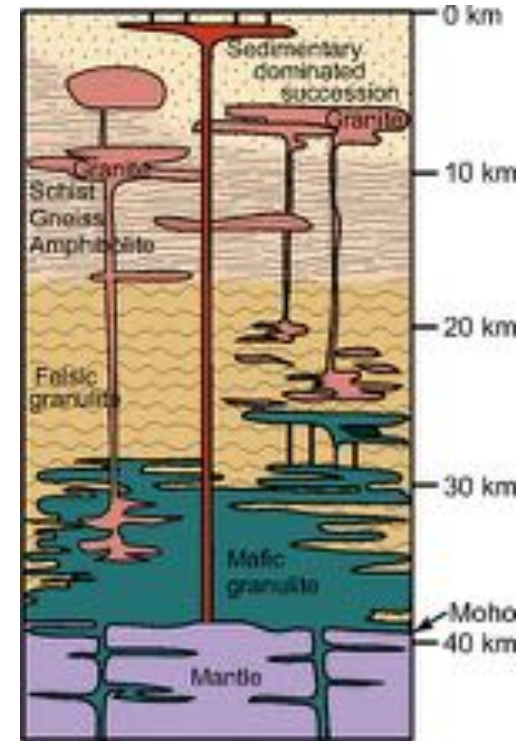
Extrusive heat pipe magmatism



(picture from Moore&Webb 2013)

-> COLD, STRONG crust/lithosphere

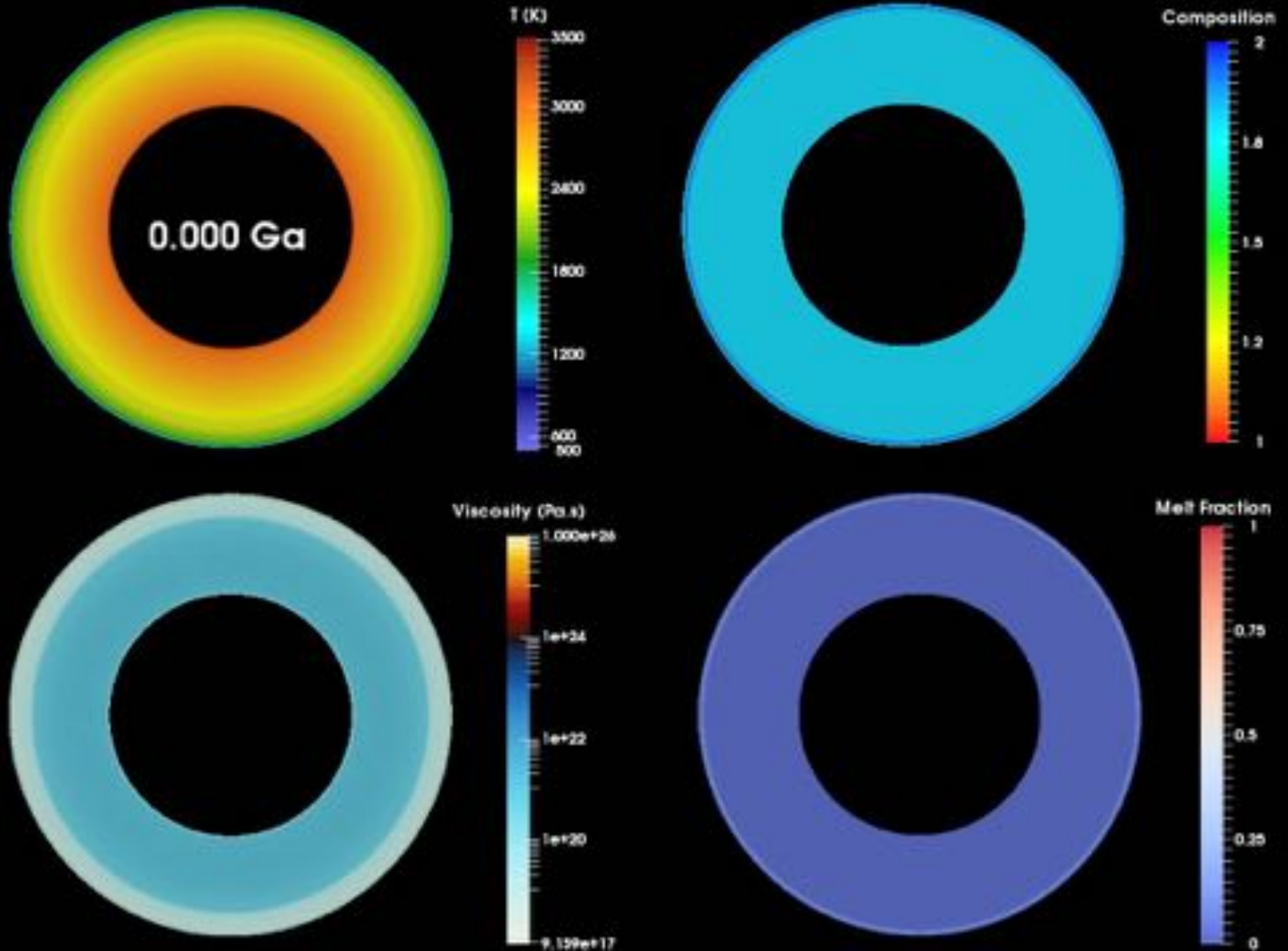
But probably **most** magmatism is intrusive



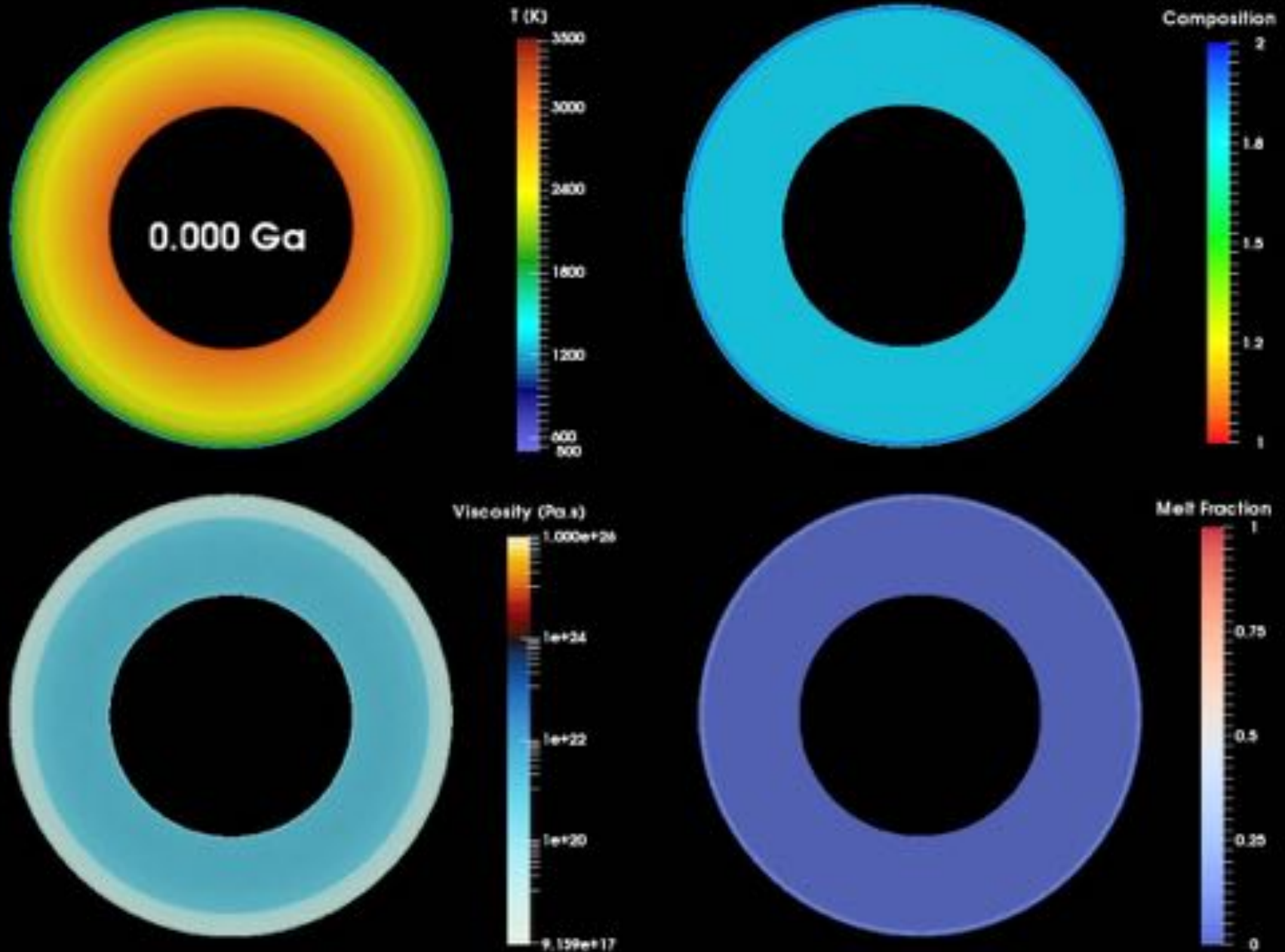
(picture from Cawood et al 2013)

-> WARM, WEAK crust/lithosphere

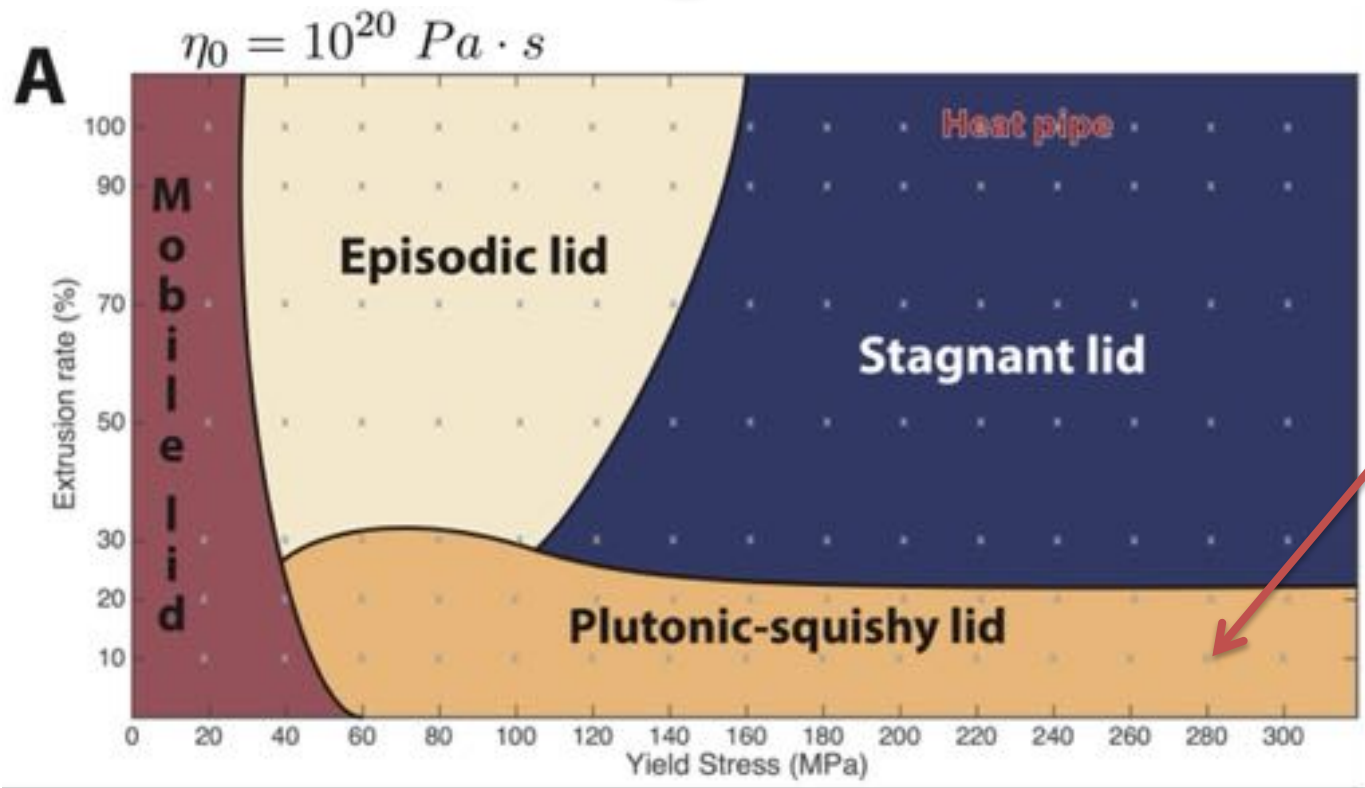
Typical episodic evolution - **extrusive**



In comparison – 90% **intrusive**



New regime found!

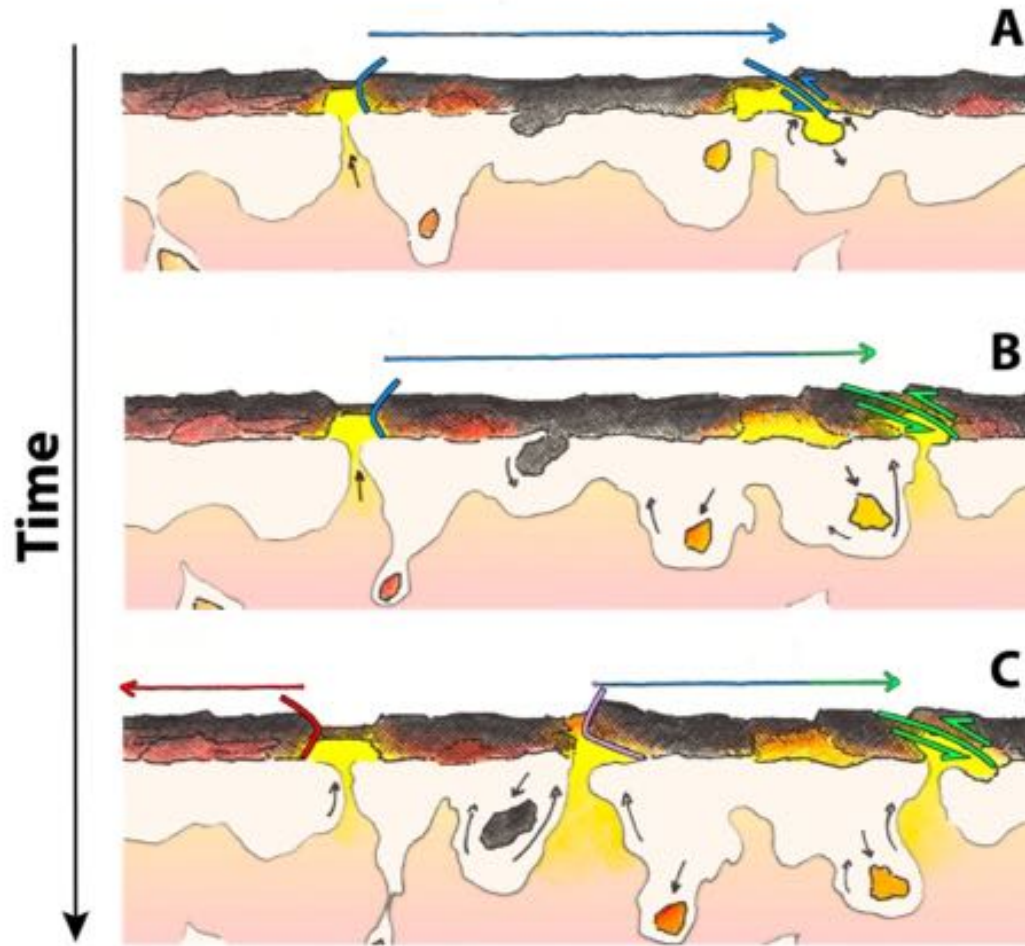


New mode

Diogo Lourenco et al., 2020 G-Cubed

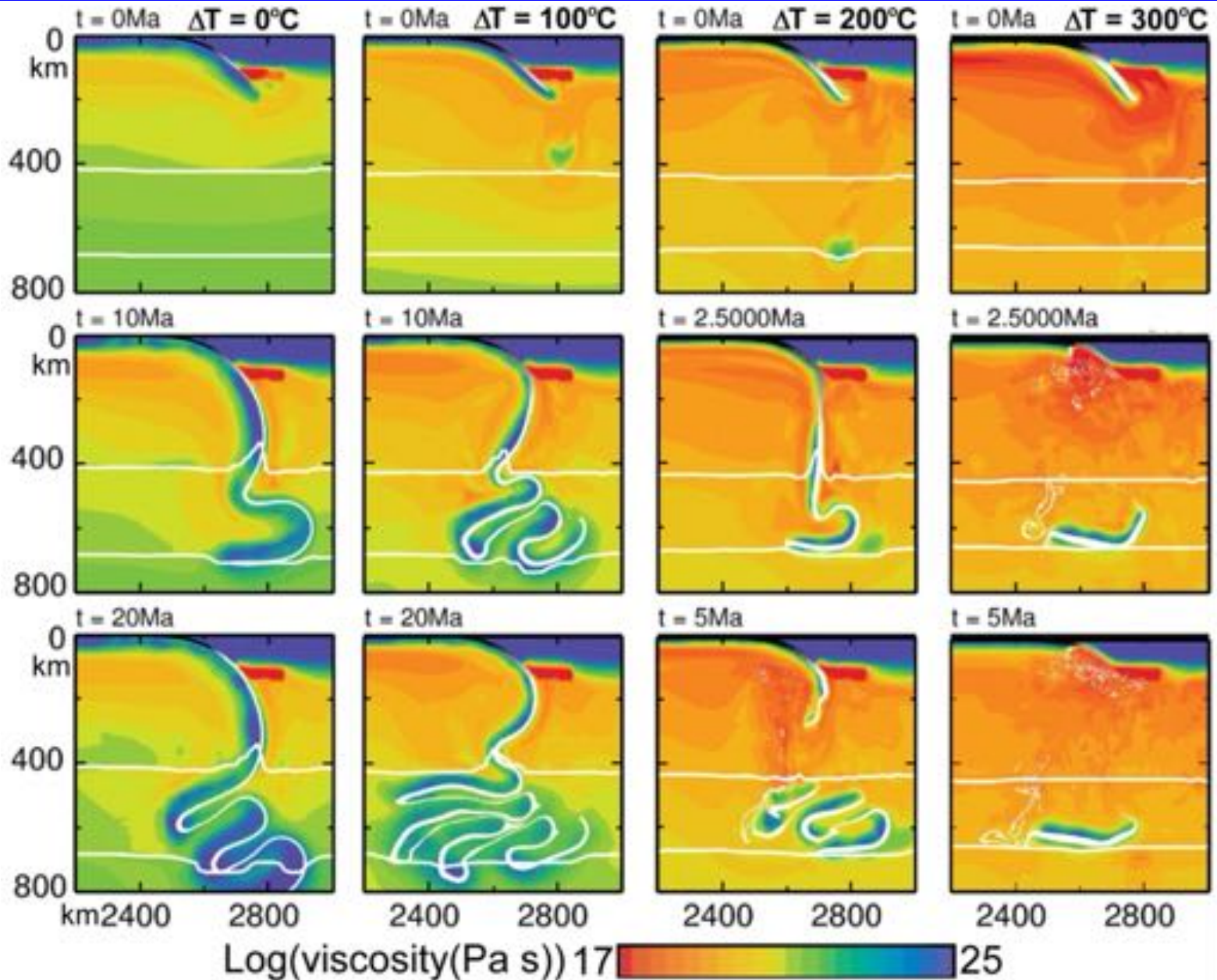


“Plutonic Squishy Lid” mode



Lourenco et al., 2020

Subduction doesn't work on a hotter Earth

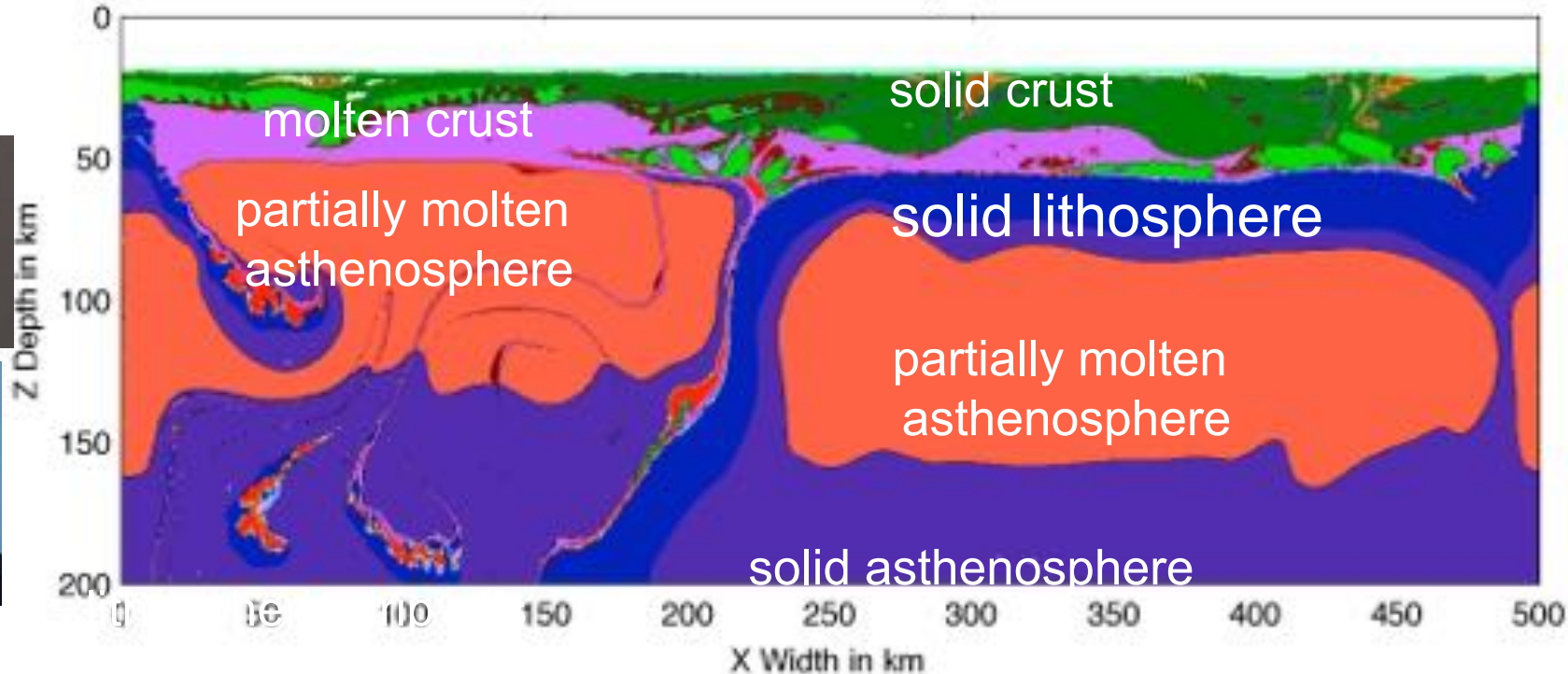


(van Hunen & van den Berg, 2008)

PSL in early Earth

- weak **deformable** plates with low topography
- mantle-flows-driven orogeny (Sizova et al., in progress)
- magma-assisted **crustal convection**

Elena
Sizova

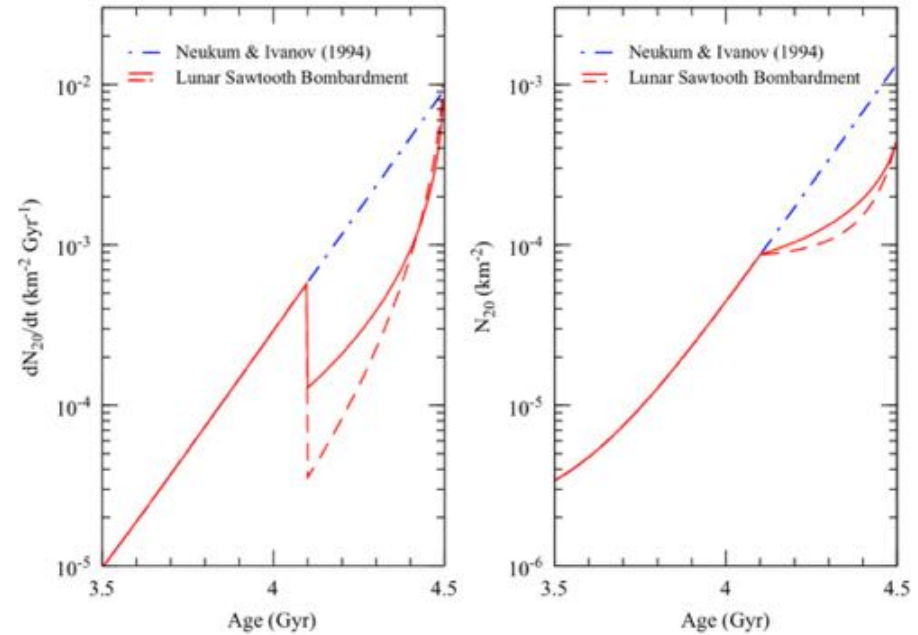


No plate tectonics but not a rigid lid either!

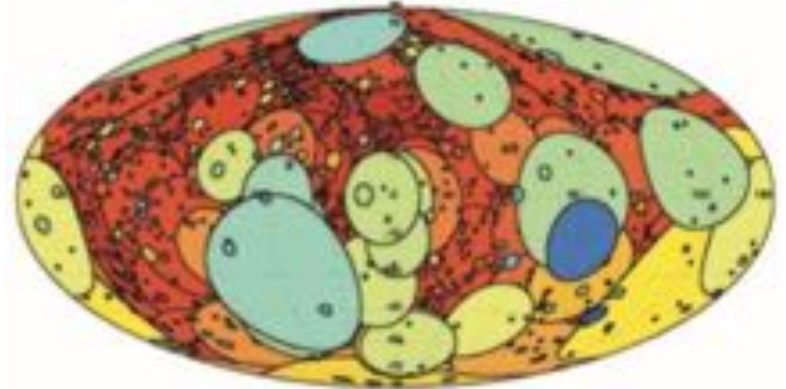
-> **Plutonic Squishy-Lid tectonics**

Impacts *(late heavy bombardment, late veneer)*

A. Morbidelli et al. / *Earth and Planetary Science Letters* 355–356 (2012) 144–151



>3.5 Gyr ago



- Sawtooth bombardment (Morbidelli et al. 2012)
- Supplies late veneer as in Marchi et al. (2014 Nature)
- O'Neill et al. (2017 NGeo) presented 2D models, here we explore 3D models
- BSc project of Xavier Borgeat

Greatly influence tectonics & crust!

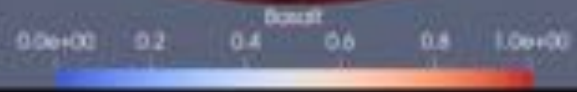
Without Impacts



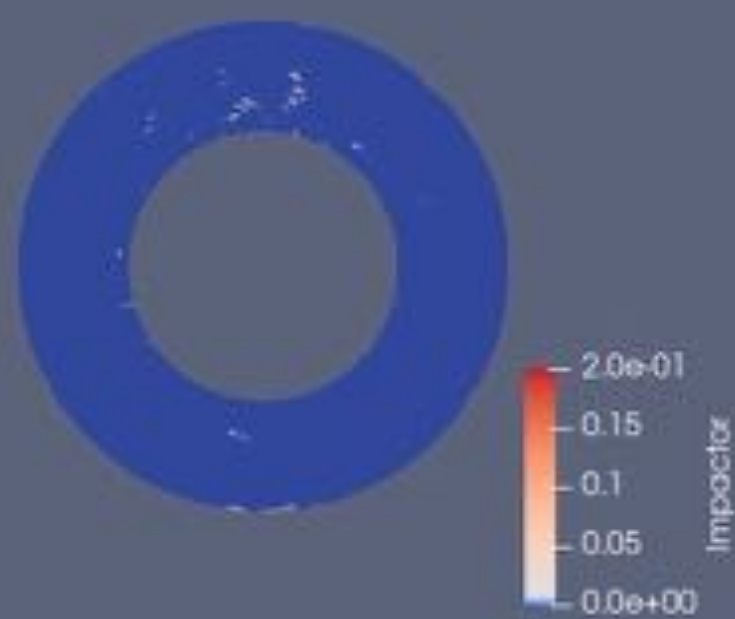
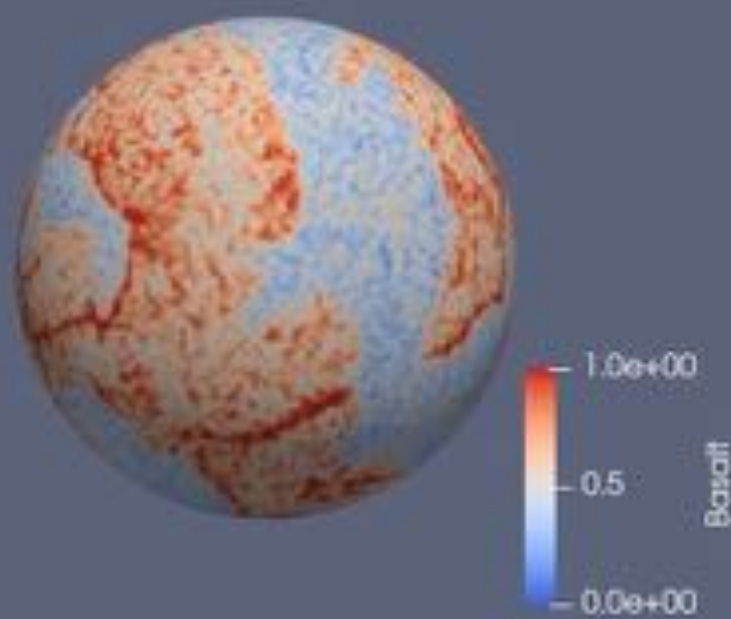
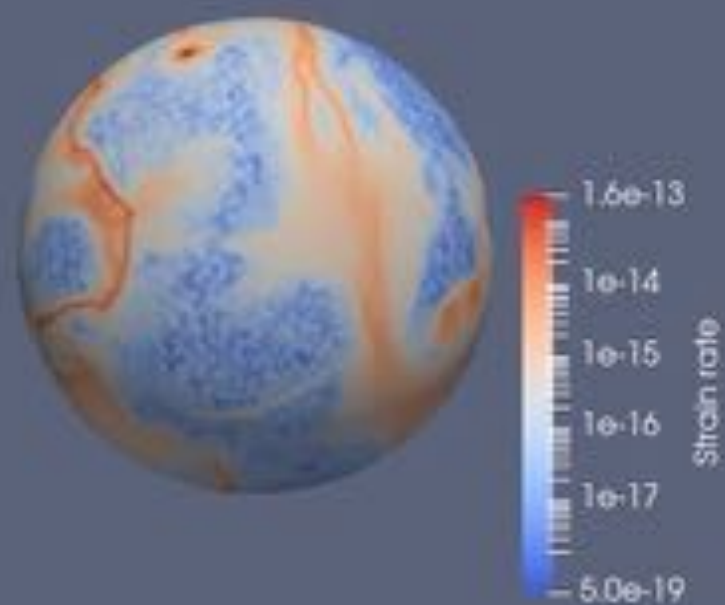
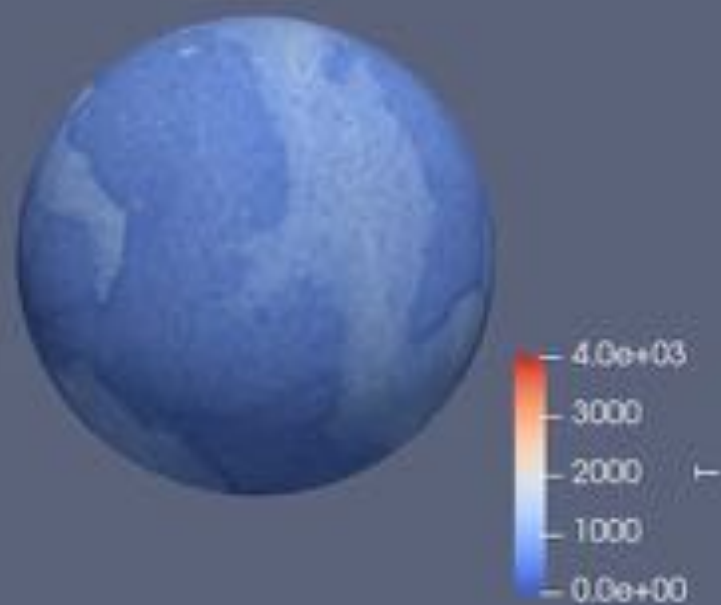
With Impacts



Time: 0.000000



Time: 100.418190



Surface mobility

Impacts can break a stagnant lid, giving temporary mobility, BUT when the impacts stop, stagnant lid returns.

If anyway a mobile-lid case, does not influence average mobility.

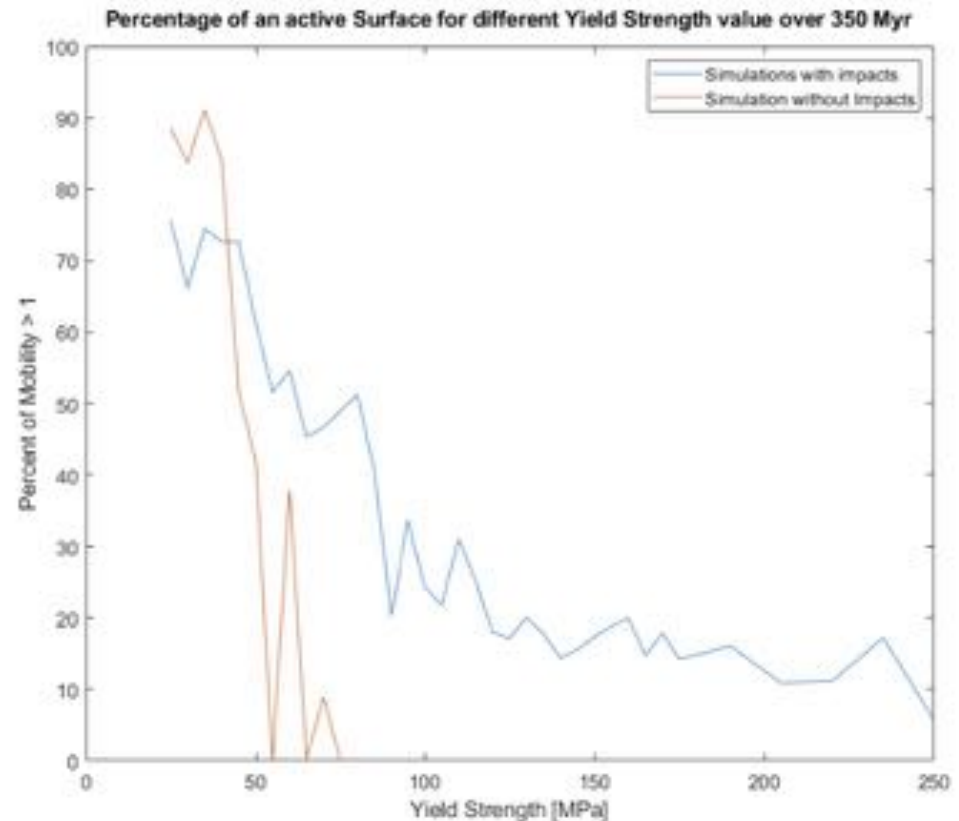


Figure 4.6: Comparison of the surface mobility for the cases with and without impacts (blue cases 3-38, orange 39-49)

Early Earth: Summary

- Archean tectonics likely characterized by hot, weak, deformable lithosphere undergoing delamination and horizontal motion.
- Intrusive magmatism dominant during Archean (as opposed to “heat pipe” extrusive magmatism)
- Subduction does not appear to be necessary for production of early TTG crust
- Impacts can play a major role in promoting mobility and melting in first ~600-700 Myrs

Coupled mantle-core evolution

- The **mantle** controls the heat flow from the core
- Run mantle convection simulations for 4.5 Gyr of Earth history, coupling CMB heat flux to core evolution
- Which mantle evolution scenarios give a reasonable core evolution?
 - Geodynamo for at least 3.5 Gyr
 - Correct final inner core size / T_{cmb}
- Constrains mantle evolution & indicates what is possible

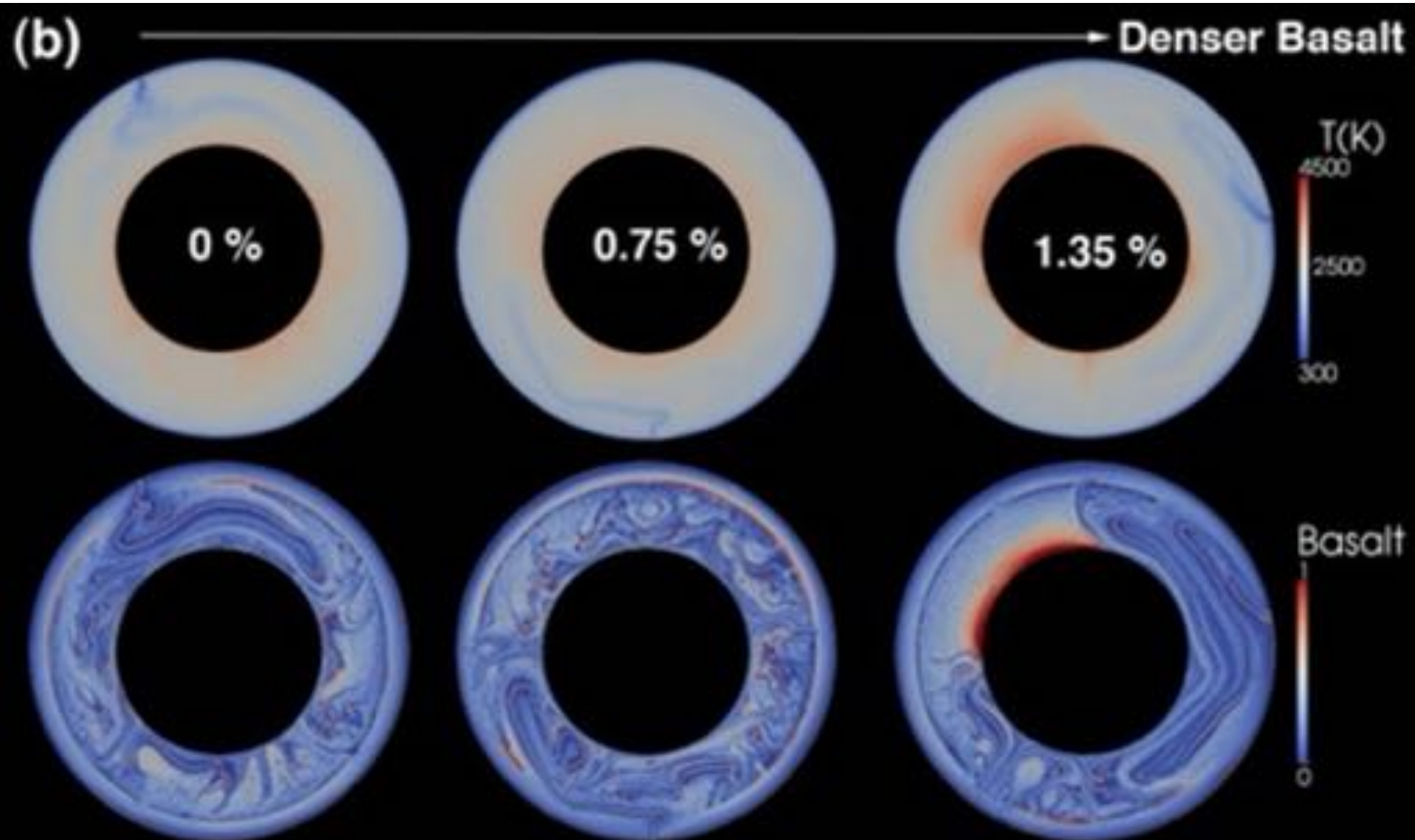
Calculations of mantle thermo-chemical evolution over 4.5 Gyr

- Include melting->crustal production,
 - viscosity dependent on T, d, and stress,
 - self-consistent plate tectonics,
 - decaying radiogenic elements and cooling core,
 - compressible anelastic approximation
- Many papers by Takashi Nakagawa & me

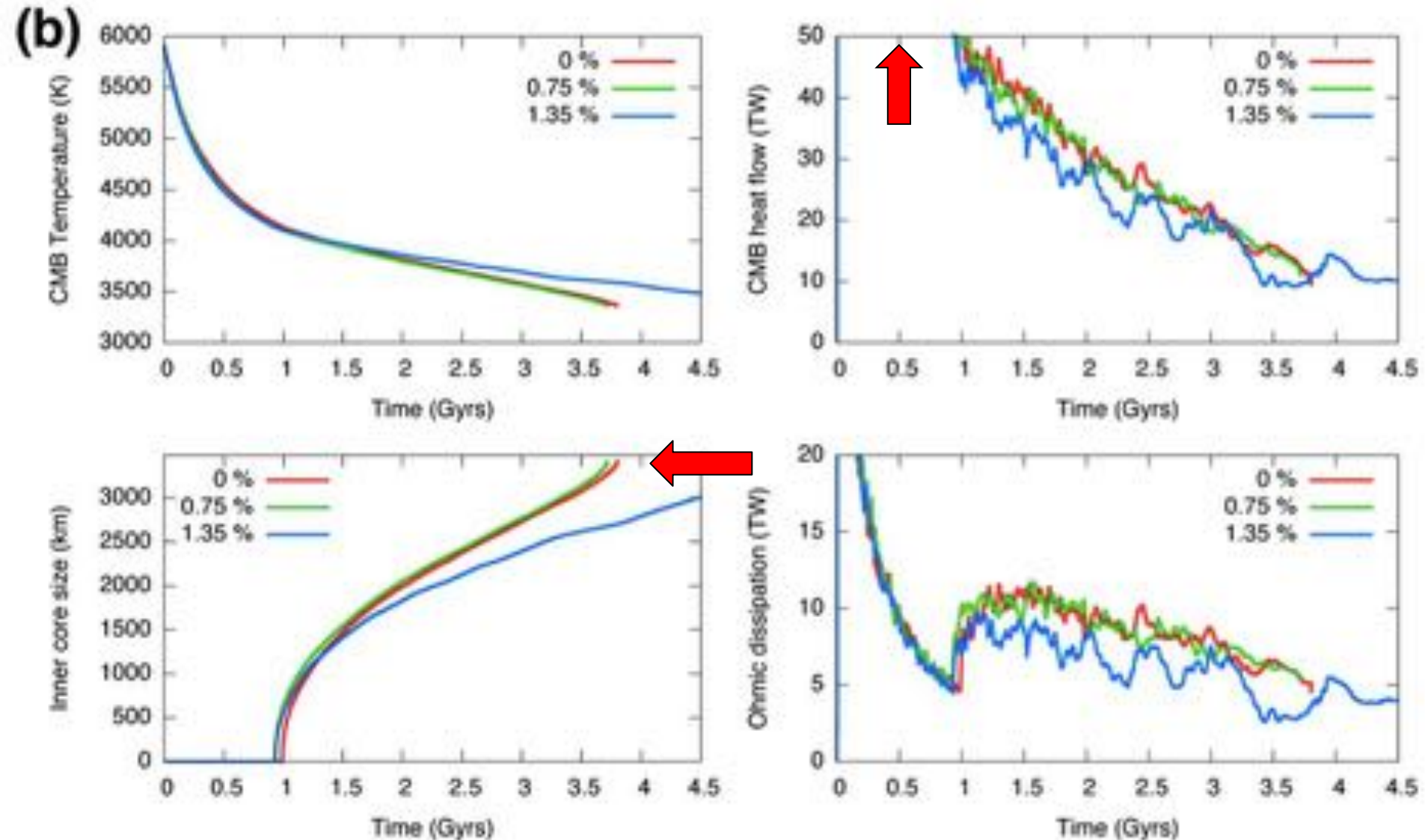


Nakagawa & Tackley 2014 G3

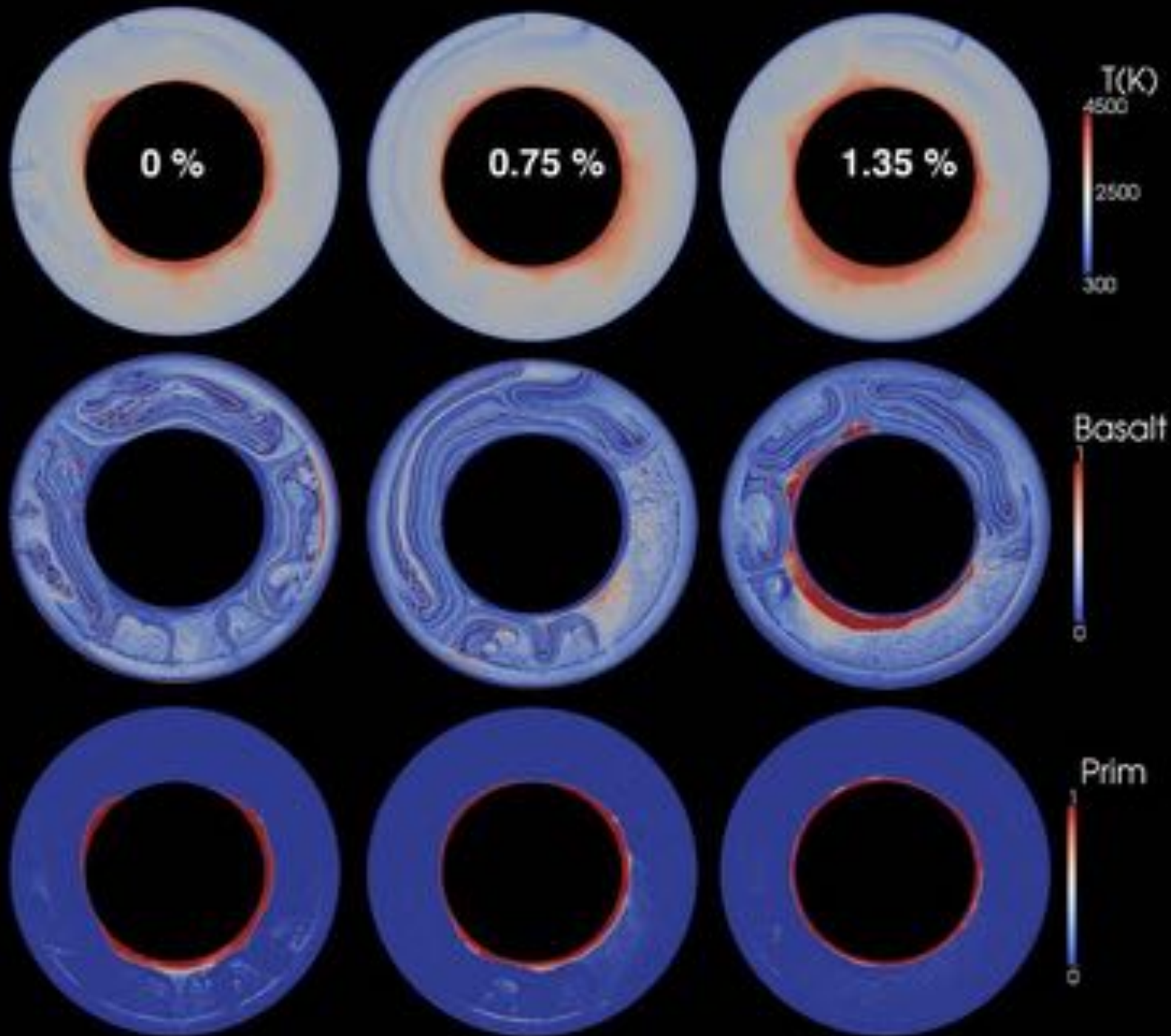
Only segregating MORB



Too-large inner core! (very high early CMB heat flow)

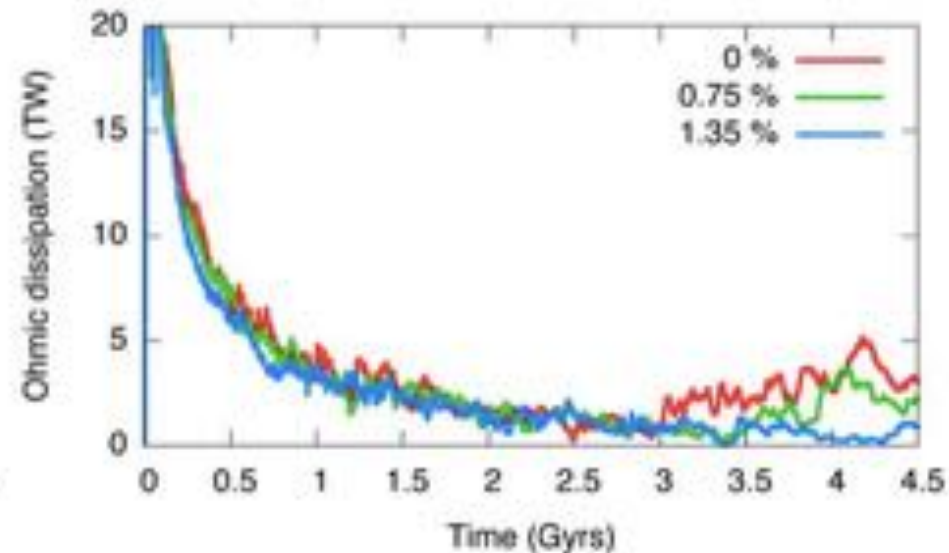
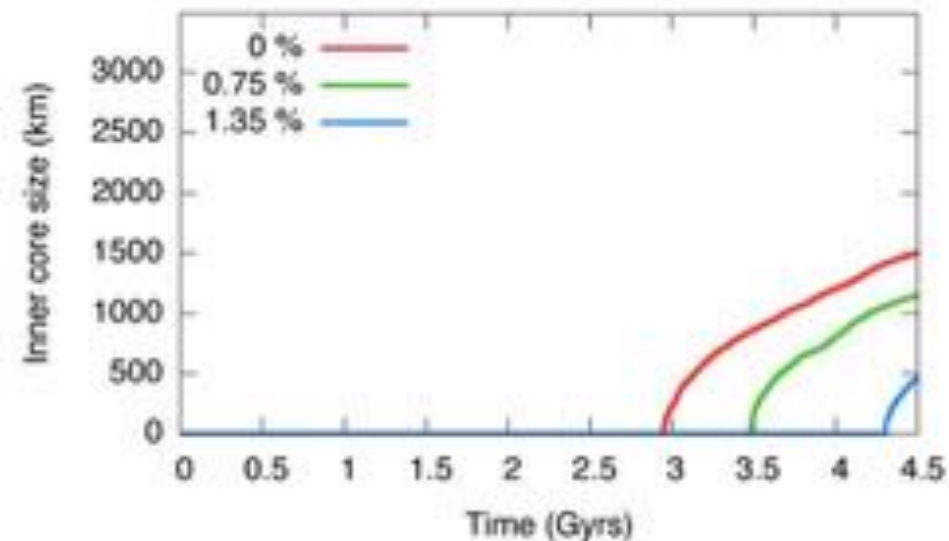
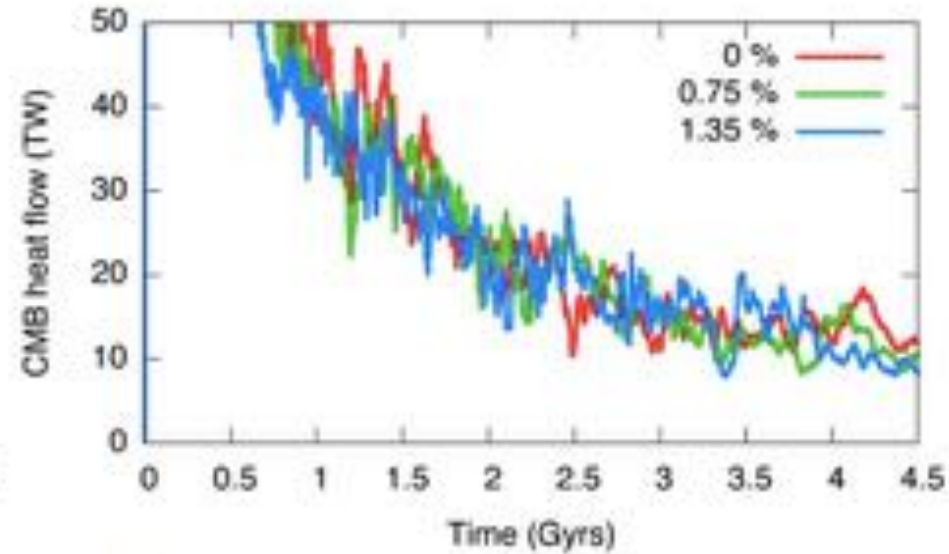
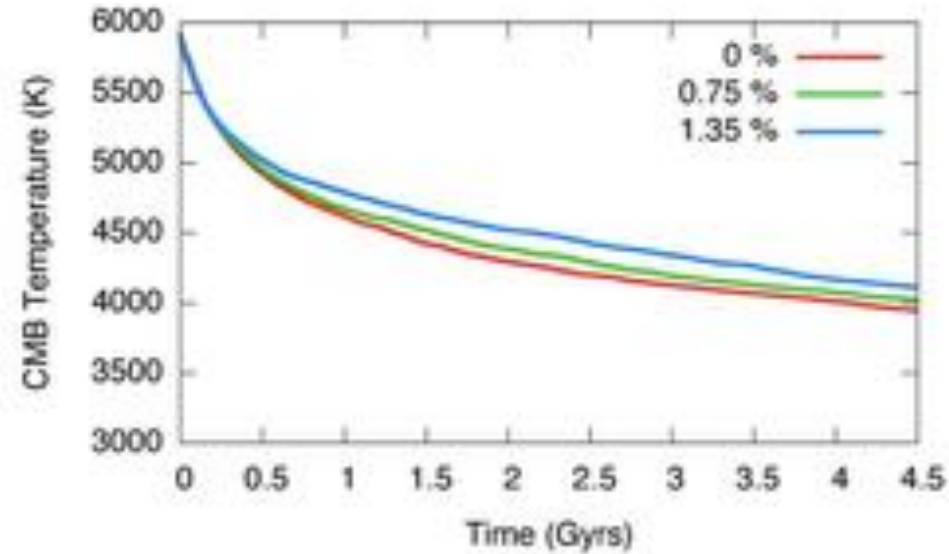


With
prim.
Layer
+
MORB



Successful core evolution

Deep dense layer reduces core cooling

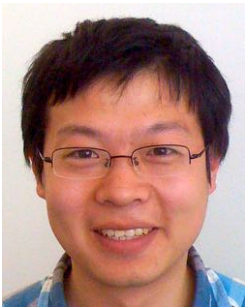


Primordial dense material: Effect of viscosity contrast and plate tectonics

- Langemeyer, Lowman & Tackley (2020 GJI)



- Li, Deschamps, Yang, Chen, Zhao & Tackley (2019 GRL)



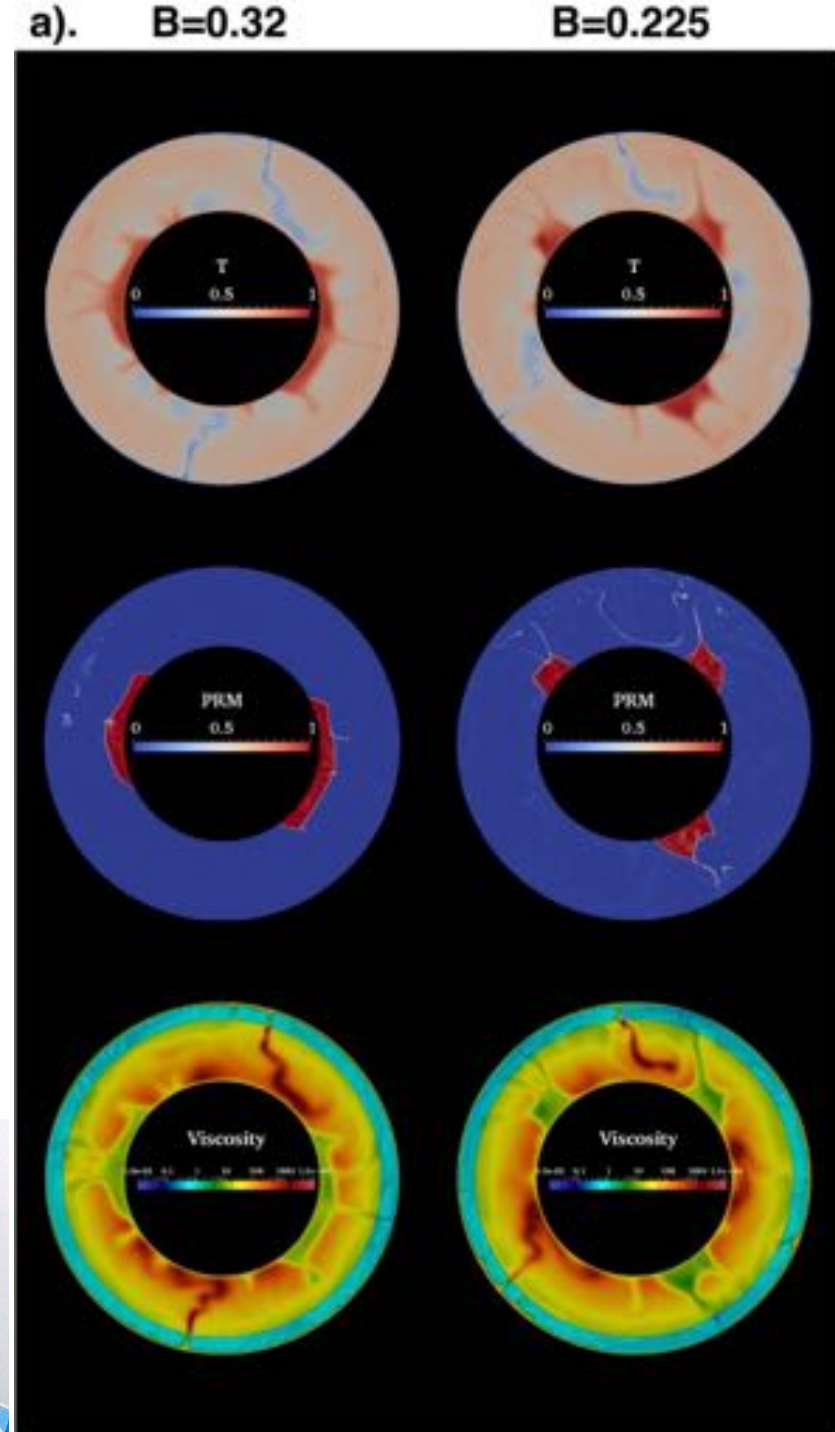
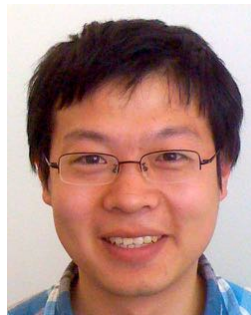
Why should there be an intrinsic (chemical) viscosity contrast?

- Different composition -> different mineralogy (brigmanite vs. magnesiowüstite)
- Different water content
- Different iron content
- Different grain size (grains grow with time, recrystallise in phase transitions)

Buoyancy ratio

(chemical:thermal) has a first-order influence on pile topography

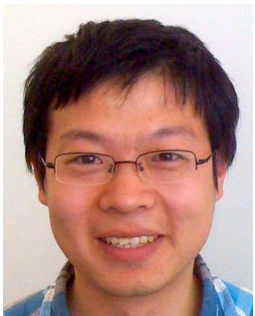
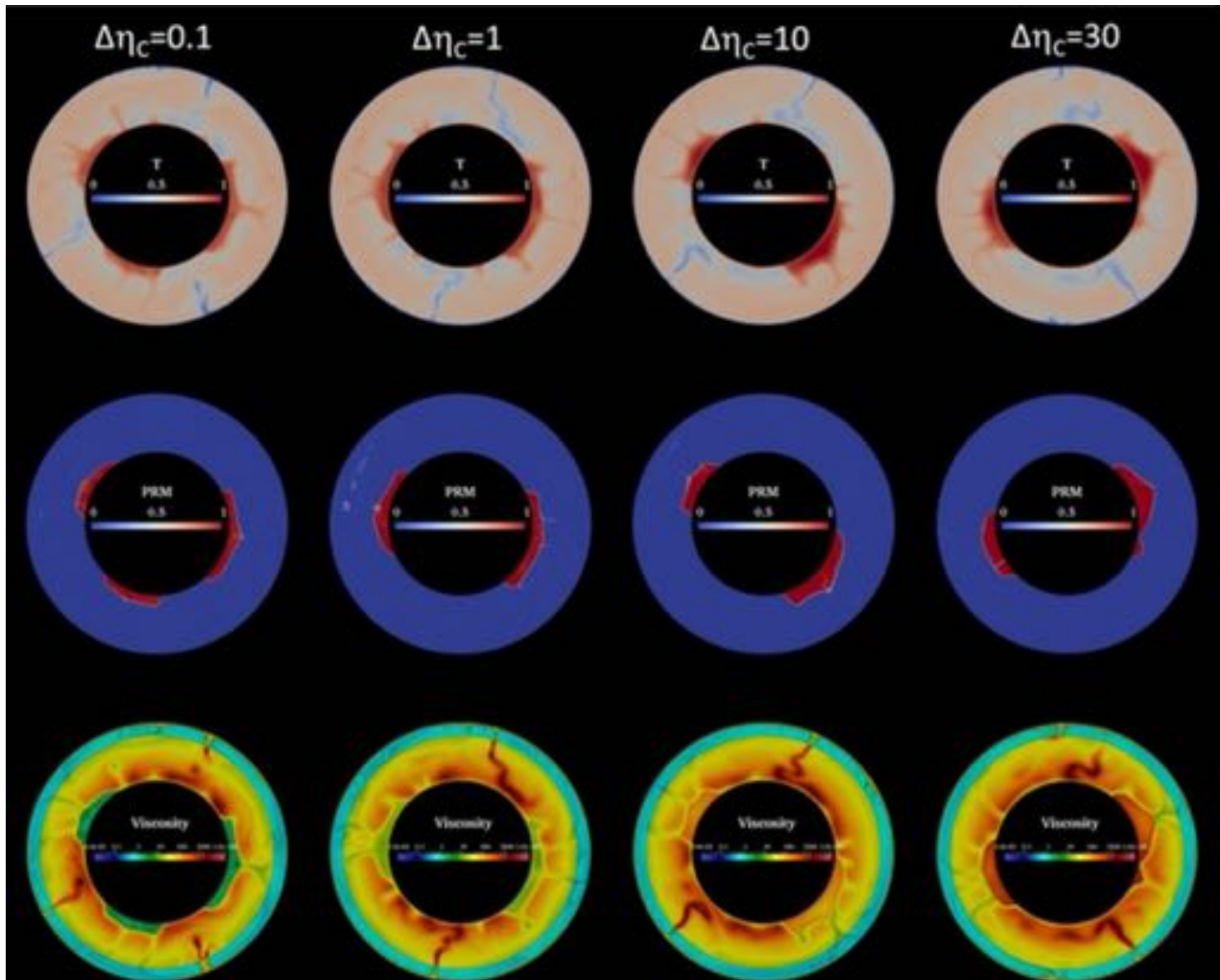
- Well-known from previous studies; this is just a reminder



Viscosity contrast doesn't hugely affect the dynamics

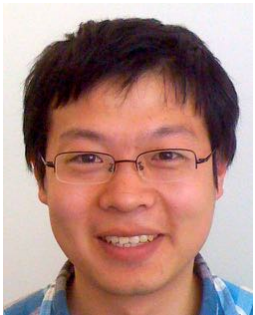
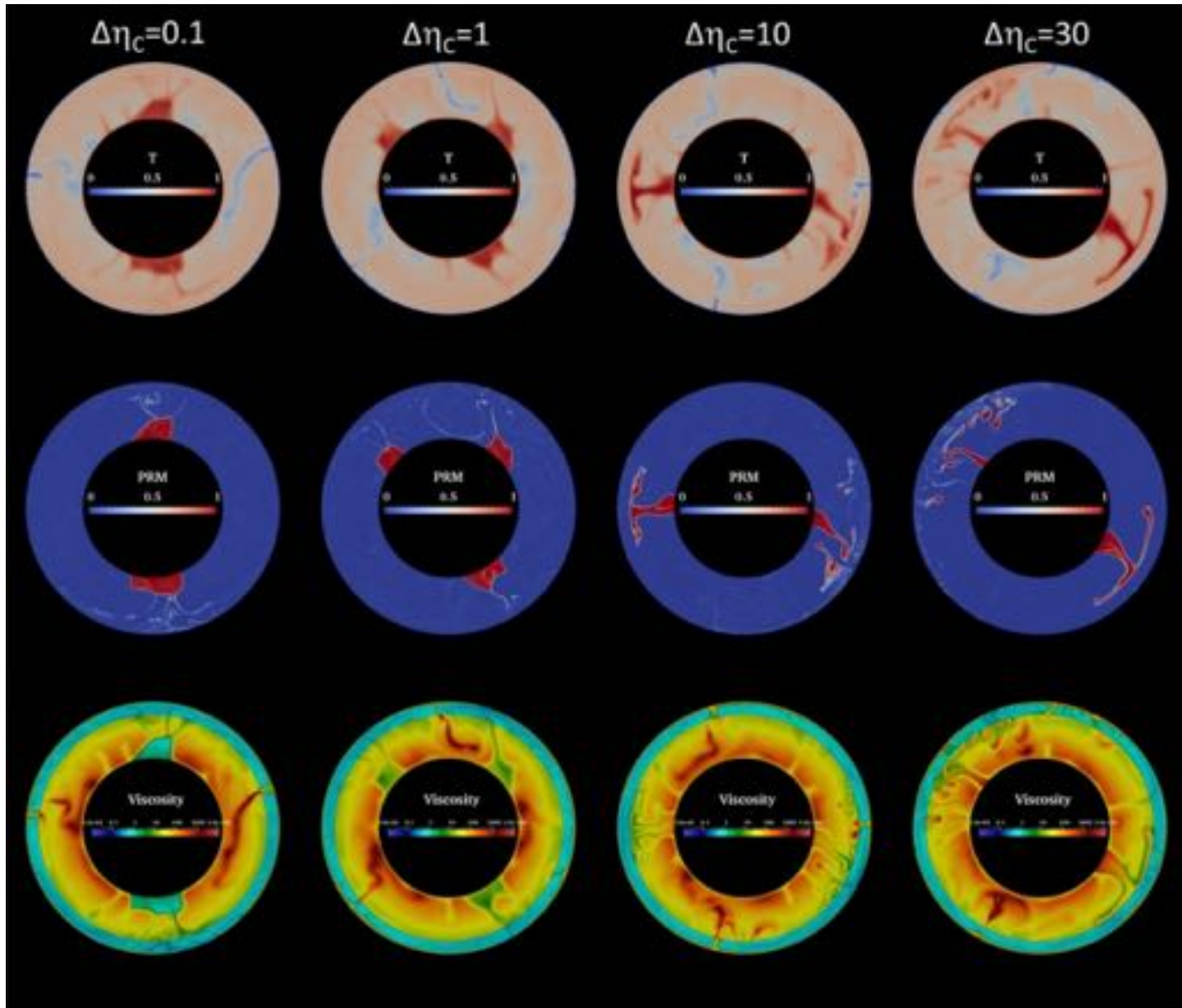
Temperature contrast increases, so does pile topography

$B=0.32$



But near the threshold, it makes a key difference
High-viscosity piles are more unstable because they become hotter

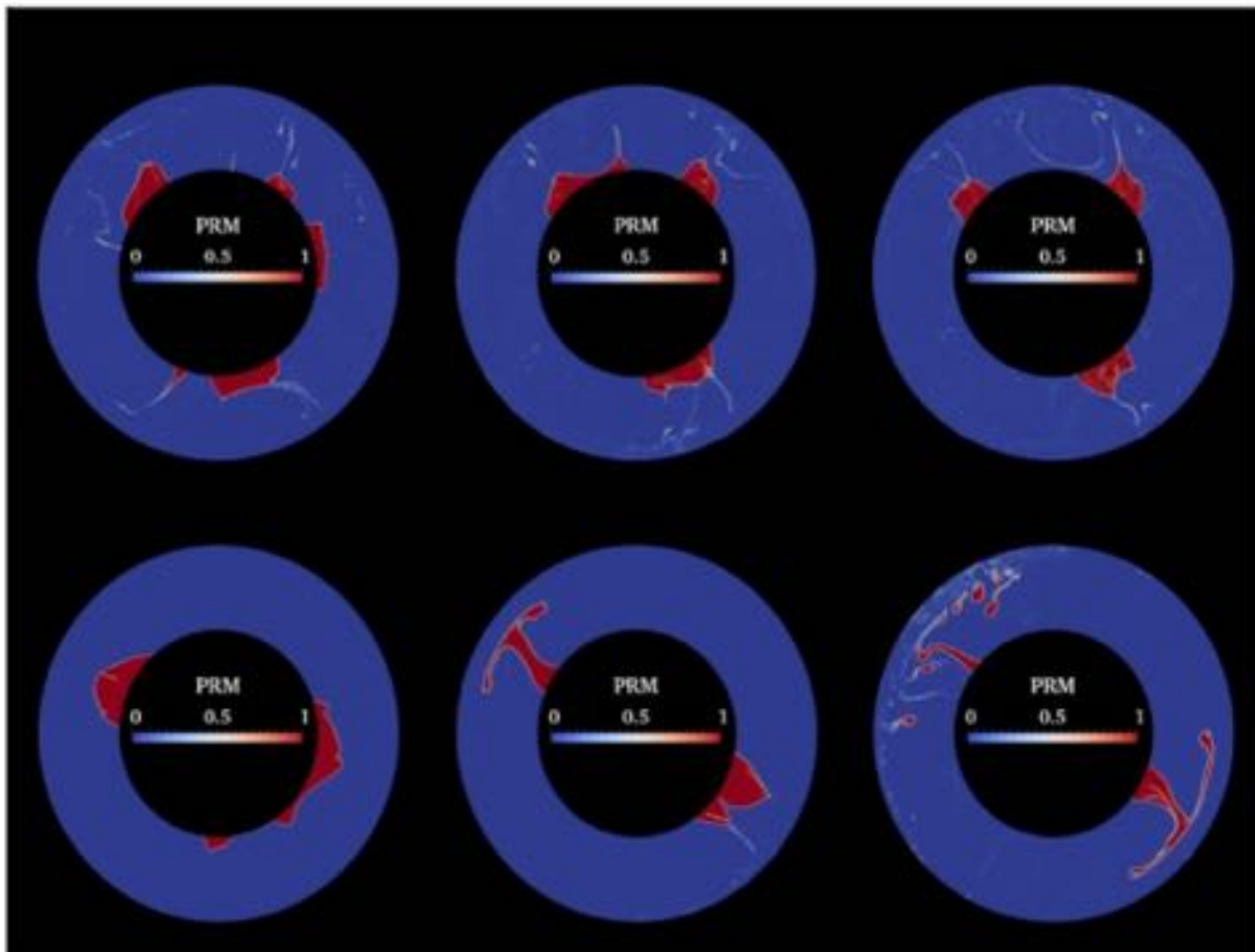
$B=0.225$



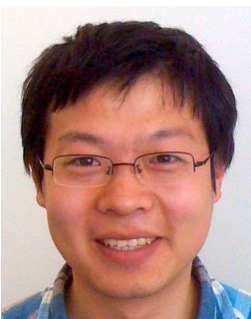
It takes a while for the layer to become unstable

a). $t=2.3\text{Gyrs}$ $t=3.5\text{Gyrs}$ $t=4.5\text{Gyrs}$

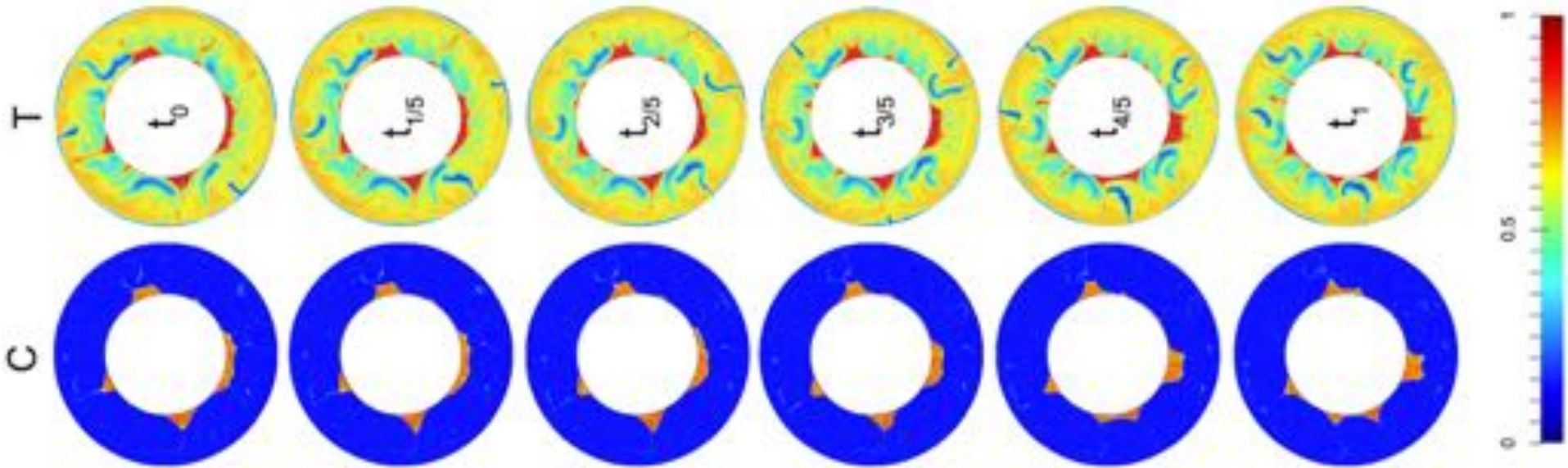
$B=0.32$



$B=0.225$



Piles don't stay fixed over billions of years

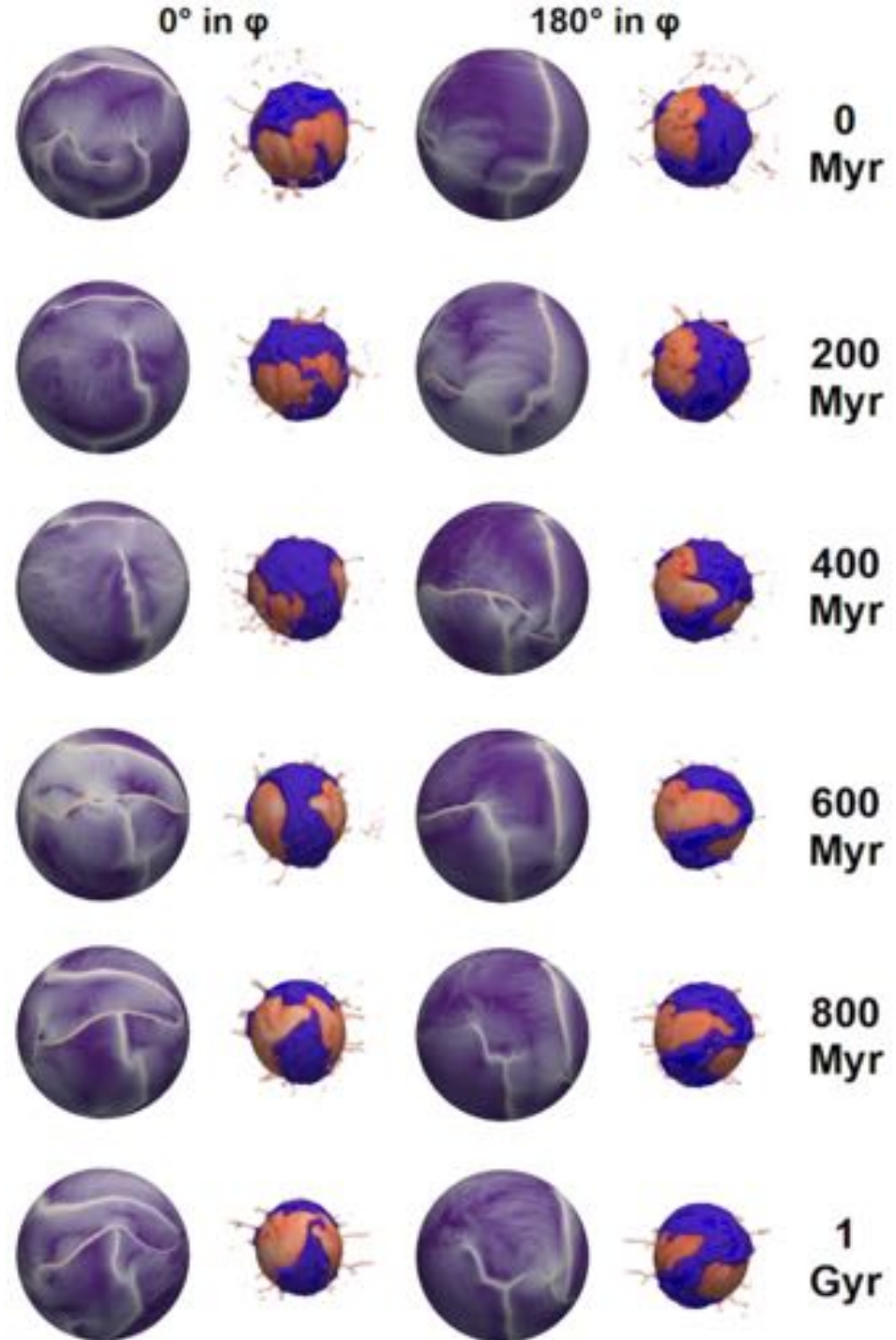


- Downwelling slabs move them around, split them, merge them, change their topography

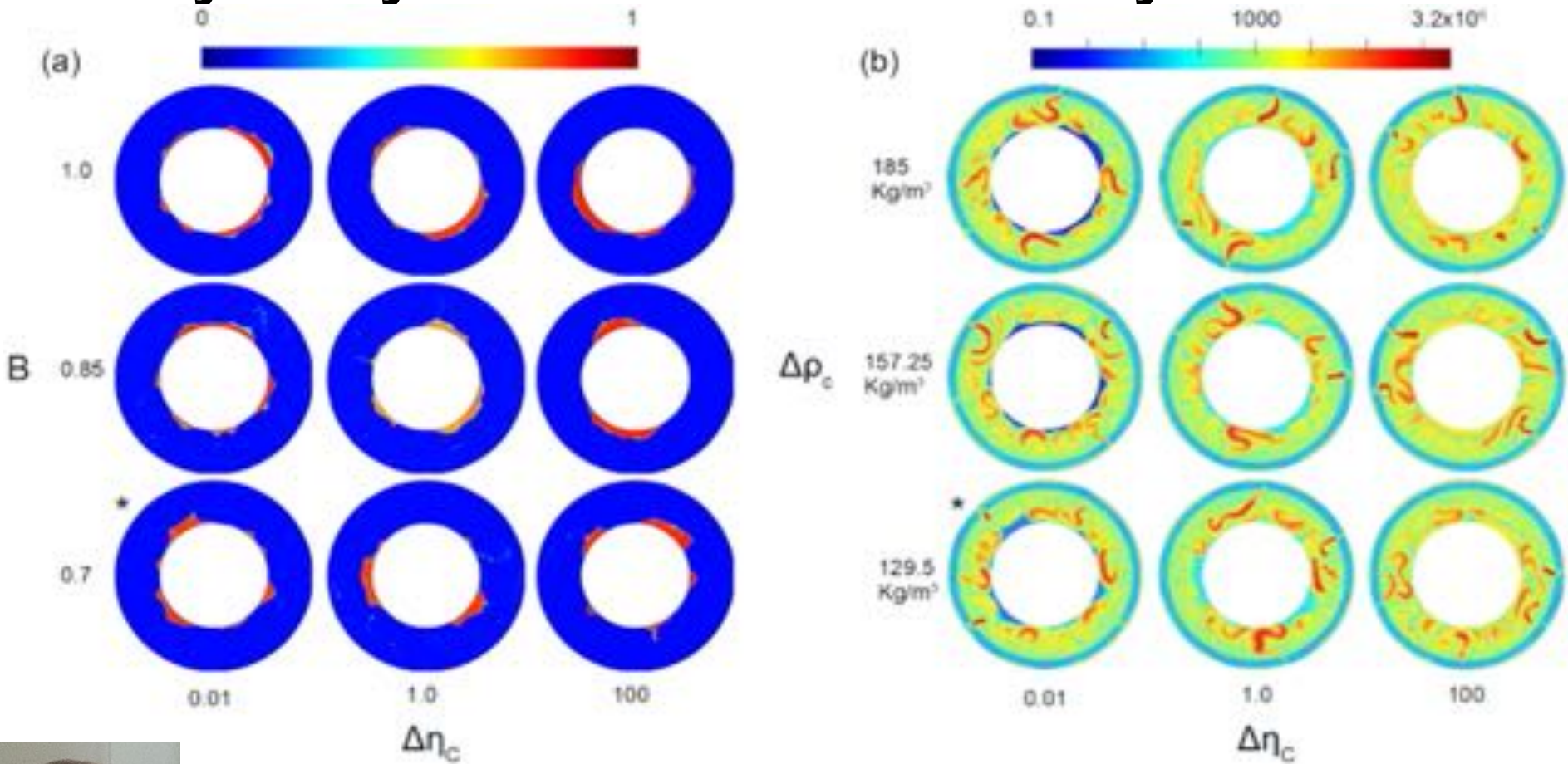
3D piles are also time-dependent



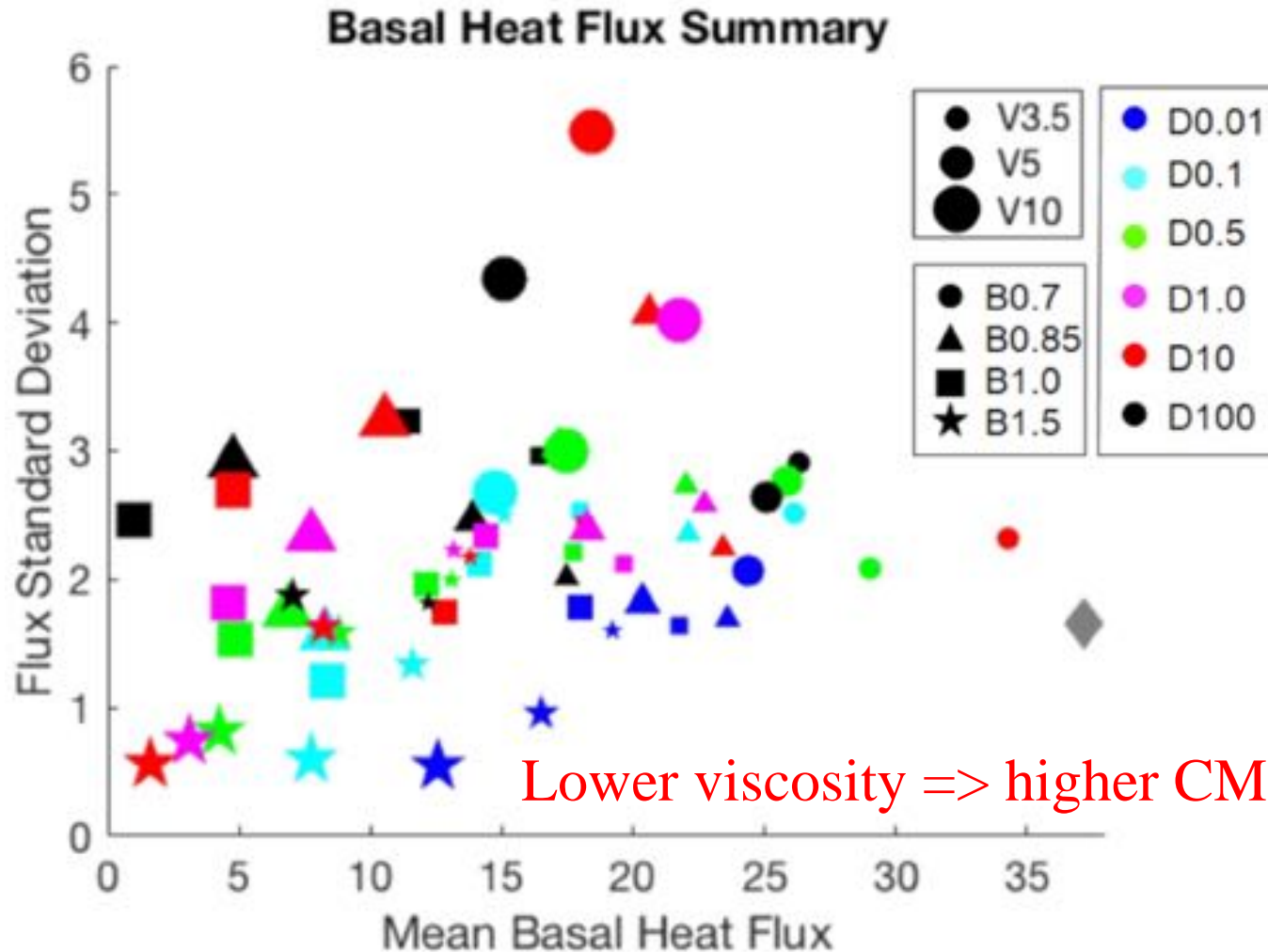
Langemeyer, Lowman &
Tackley, GJI submitted



Pile morphology depends more on buoyancy ratio than viscosity contrast



Viscosity contrast greatly affects CMB heat flux



Lower viscosity => higher CMB heat flux



Summary

- Basal Melange (BAM): any “piles” are likely a mixture of materials, much of which subducted
- Early Earth tectonics: Plutonic Squishy-Lid, also impacted by impacts
- CMB piles have a strong influence on CMB heat flux; may even be needed for a successful geodynamo evolution
- Intrinsic viscosity contrast of piles influences heat flux and stability

THE END

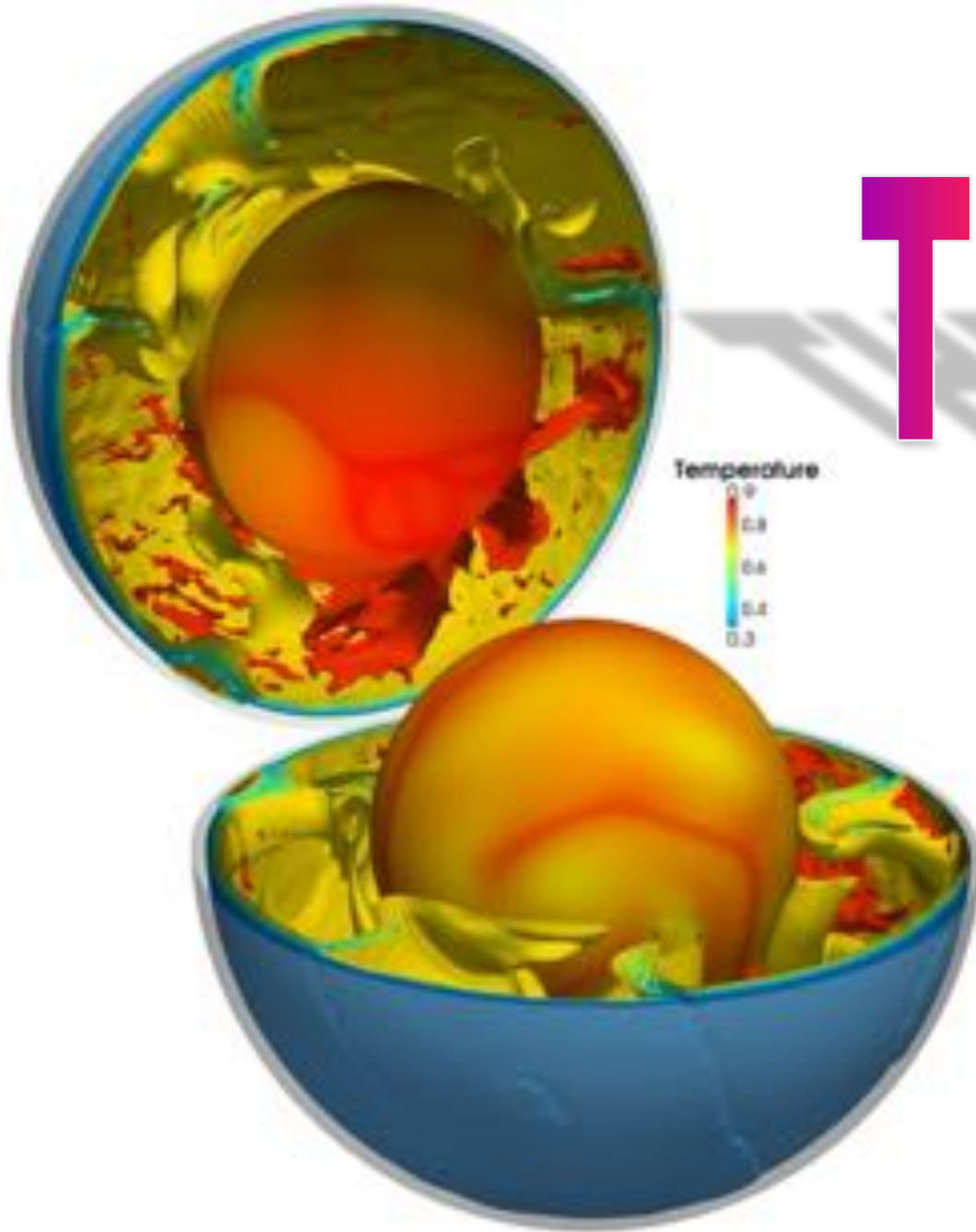


Image by Fabio Crameri