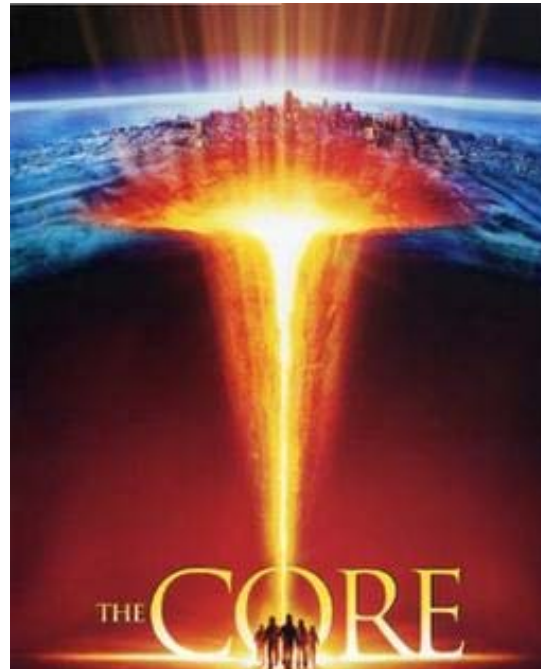




Heidelberg University

# **Mechanically-controlled rock-microstructures: Witnesses of the long-term stress-state in the continental lithosphere**

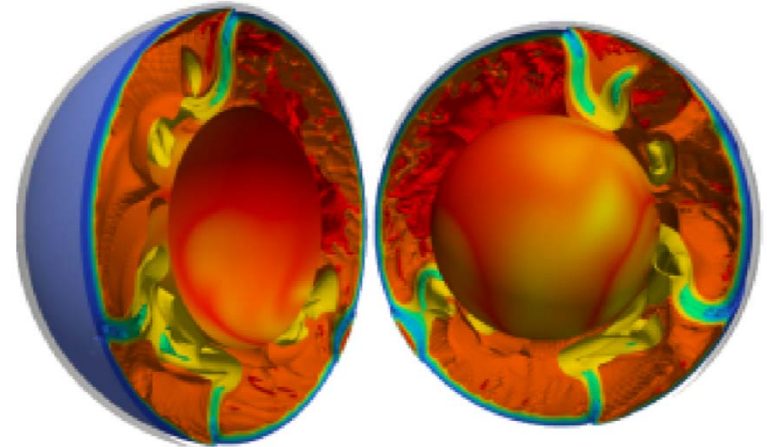
Lucie Tajčmanová, Evangelos Moulas, Yuri Podladchikov



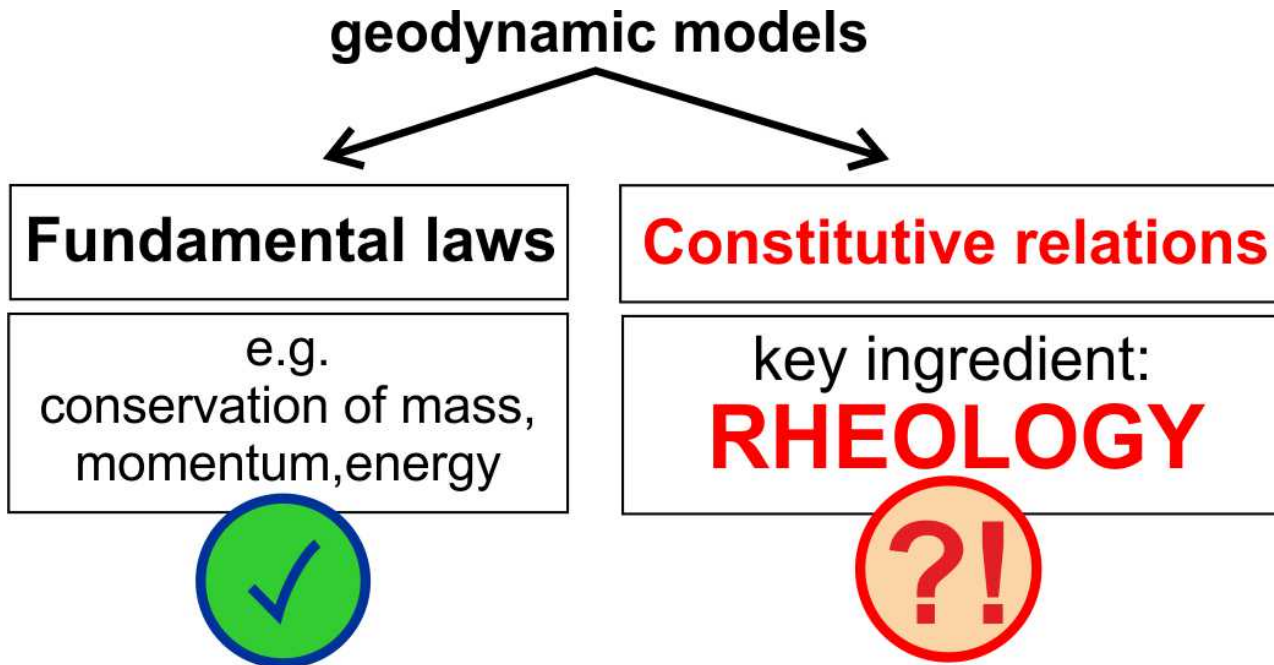
# Motivation: Modelling of geodynamic processes

## CURRENT AIMS:

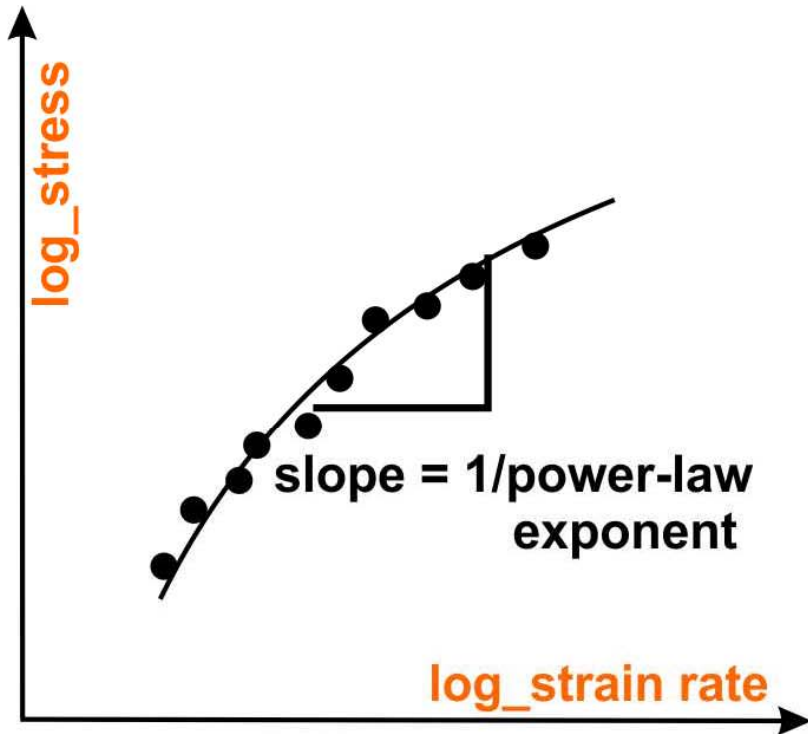
- Understanding processes in the Earth interior
- Mitigation measures, e.g. earthquakes



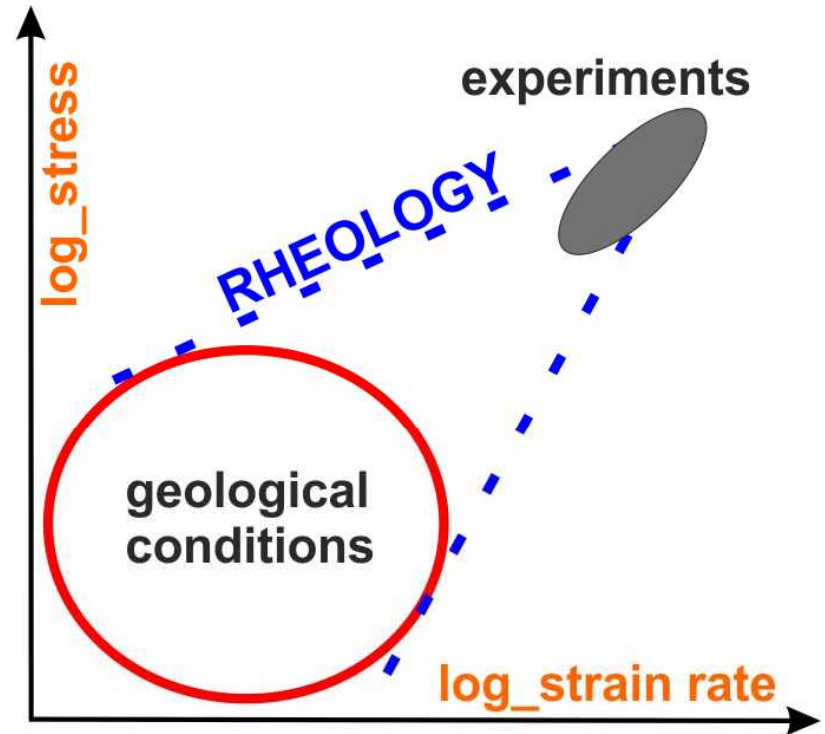
*May et al., 2013 & Cramer et al., 2012*



# Rheology: an extrapolation problem



Log-Log stress-strain rate diagram to determine the stress exponent  $n$



Extrapolation to geological conditions

## CURRENT UNCERTAINTY:

- *Rheology* extrapolated from lab over **10 orders of magnitude**

# Rheology: an extrapolation problem

PLAGIOCLASE – the most abundant mineral in the Earth's crust

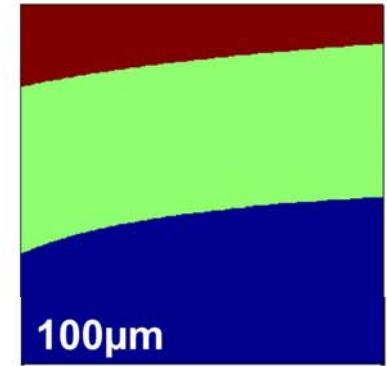
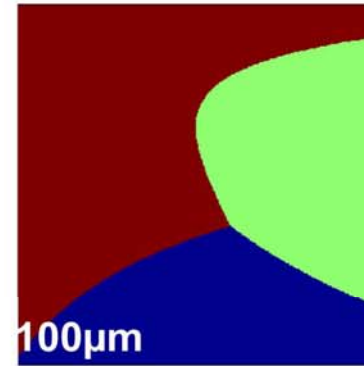
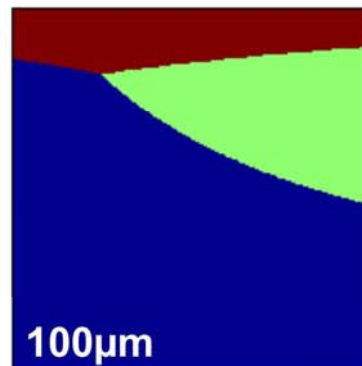
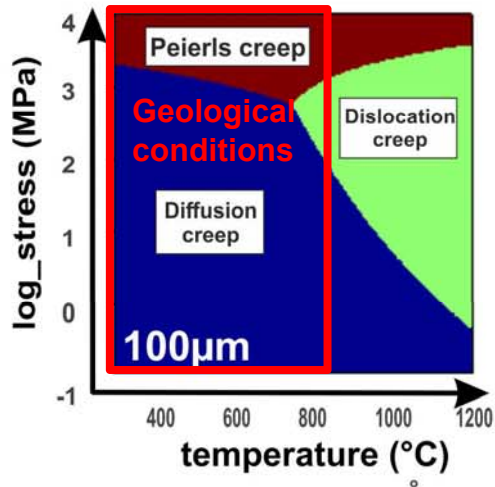
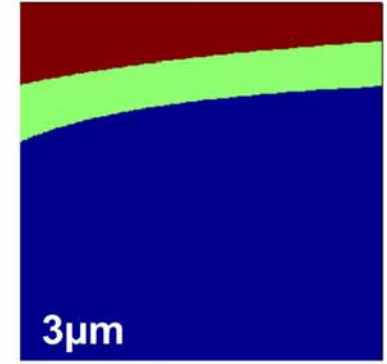
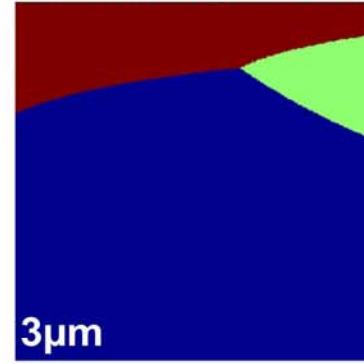
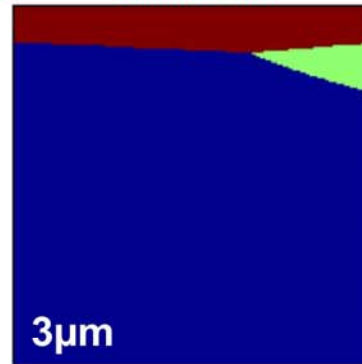
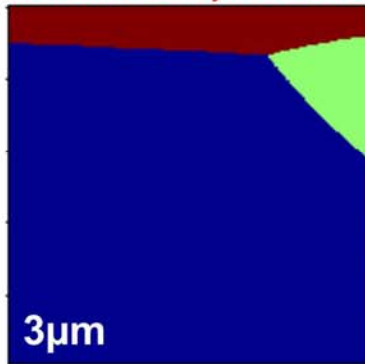
## Differences in grain size, activation energy and dislocation creep data

activation energy for diffusion creep- average wet and dry  
Dislocation c.: Rybacki & Dresen, 2006

Shelton, 1981

activation energy for diffusion creep- only dry  
Dislocation c.: Rybacki & Dresen, 2006

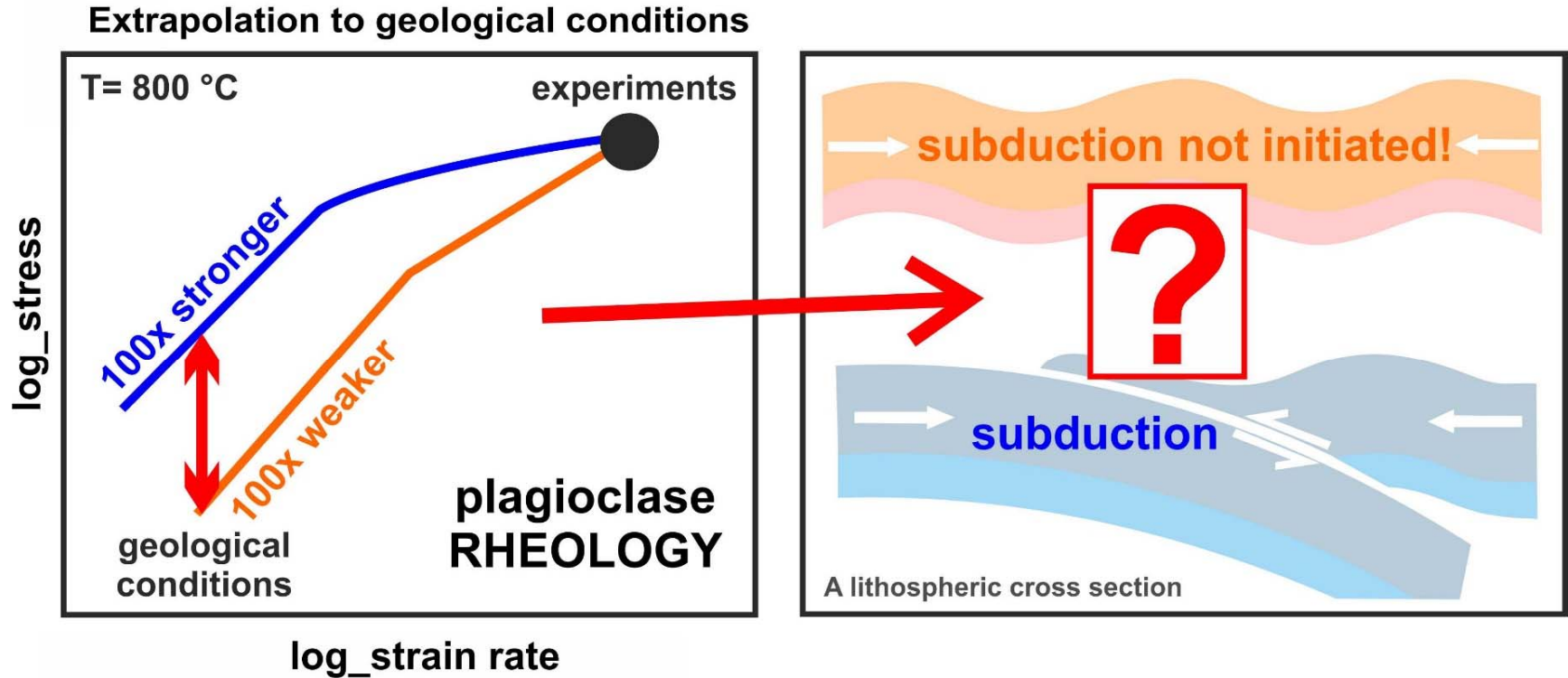
Shelton, 1981



No predictive power for geo-material towards high stresses.  
Constraint under geological conditions is missing.

# Rheology: an extrapolation problem

PLAGIOCLASE – the most abundant mineral in the Earth's crust



***=> This questions our understanding of geodynamic processes!***

# Groningen gas field (Netherlands)

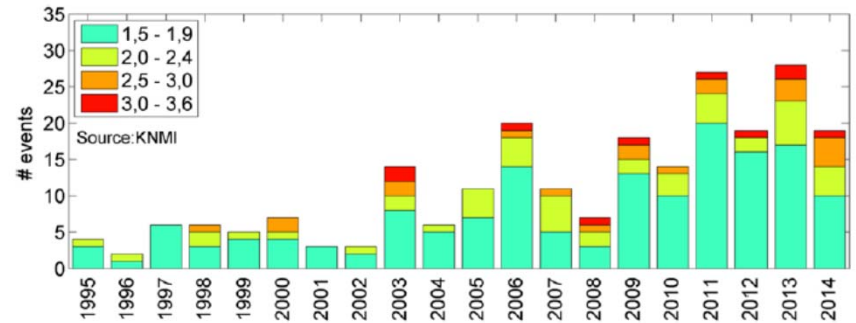
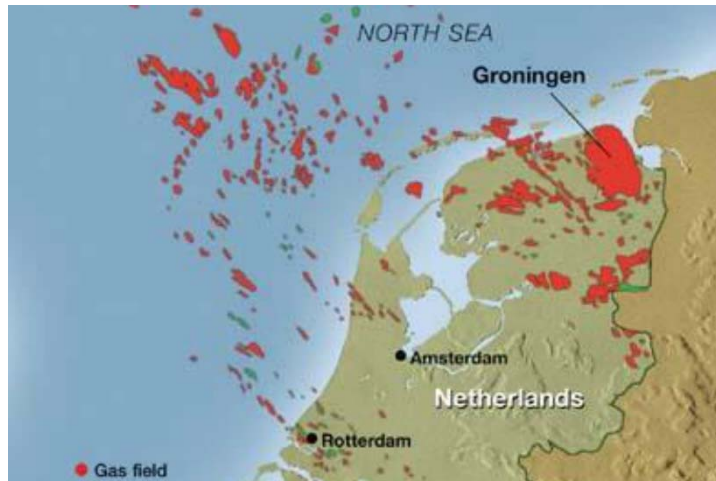
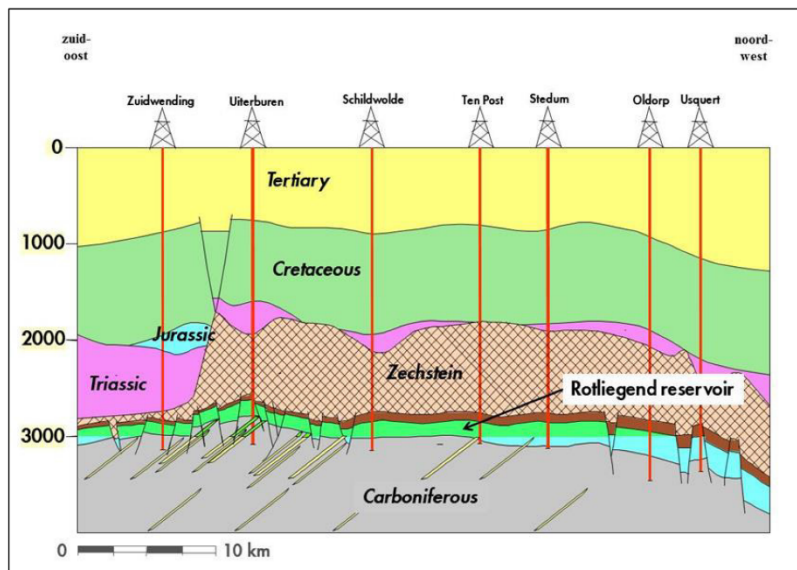


Figure 1. Groningen seismicity ( $M > 1.5$ ) vs. time.



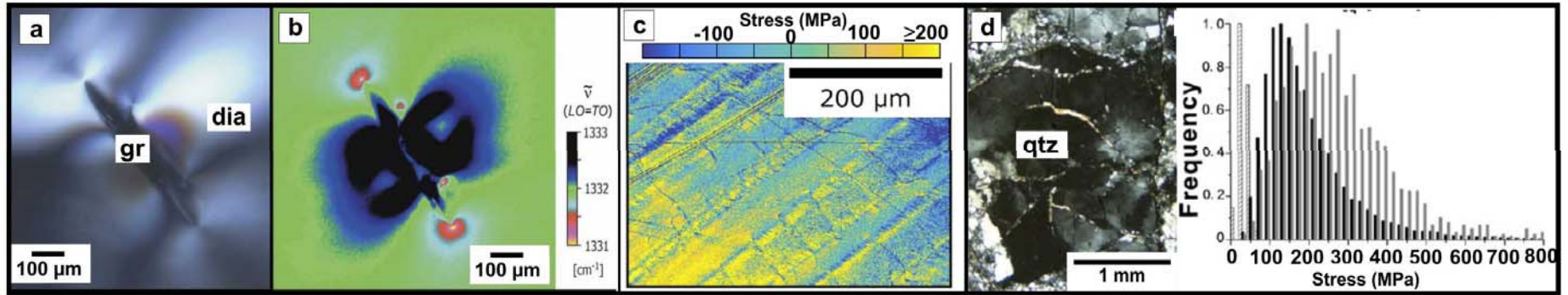
- Only 1-2 orders of magnitude difference in extrapolation, i.e. years vs. months
- Close to the surface
- Elastic vs. **creeping** behavior problem

**=> How confident are we for conditions in the lower crust?**

**Any hope for a constraint at geologically relevant conditions?**

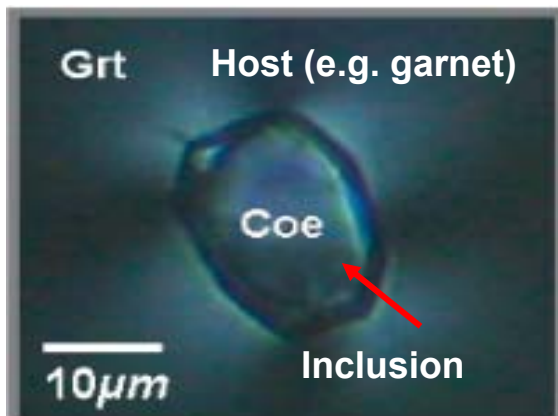


# Mechanically-controlled microstructures

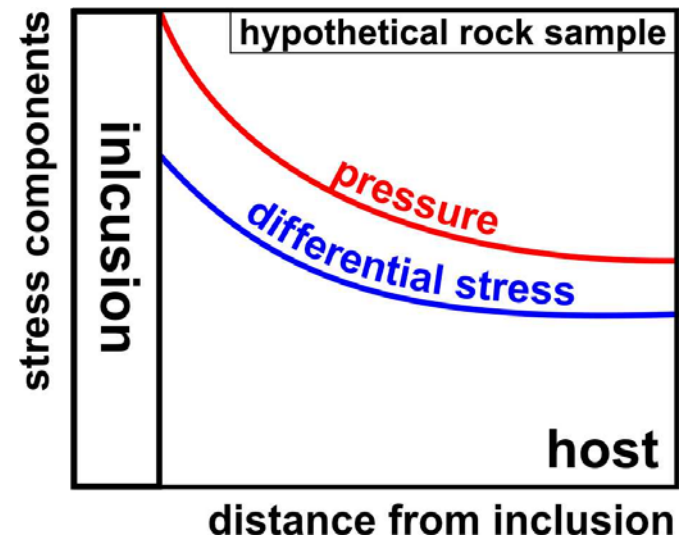


Pressure and stress **directly measured** by state-of-the-art analytical methods.

**Witnesses** of the long-term stress state in the lithosphere!



Inclusion-host environment

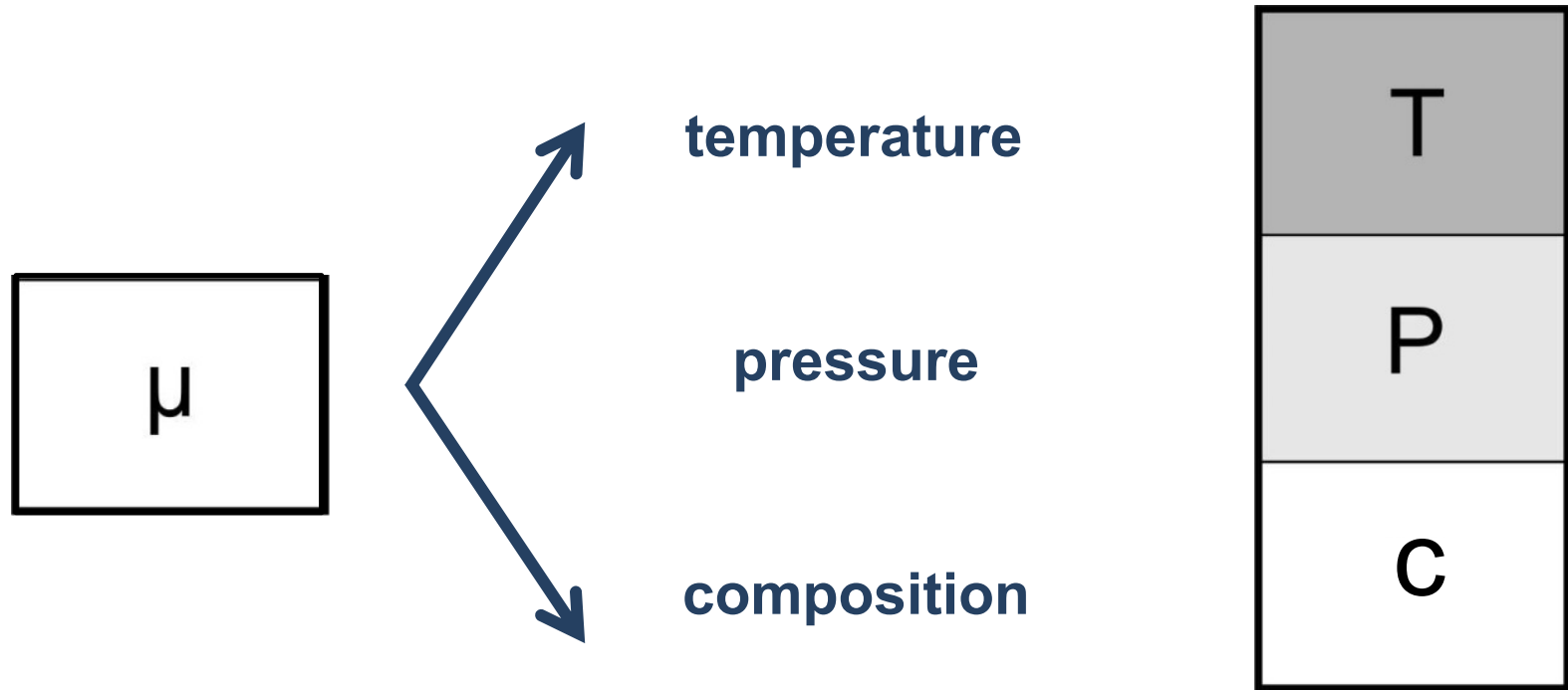




# **Unconventional Rheometry:**

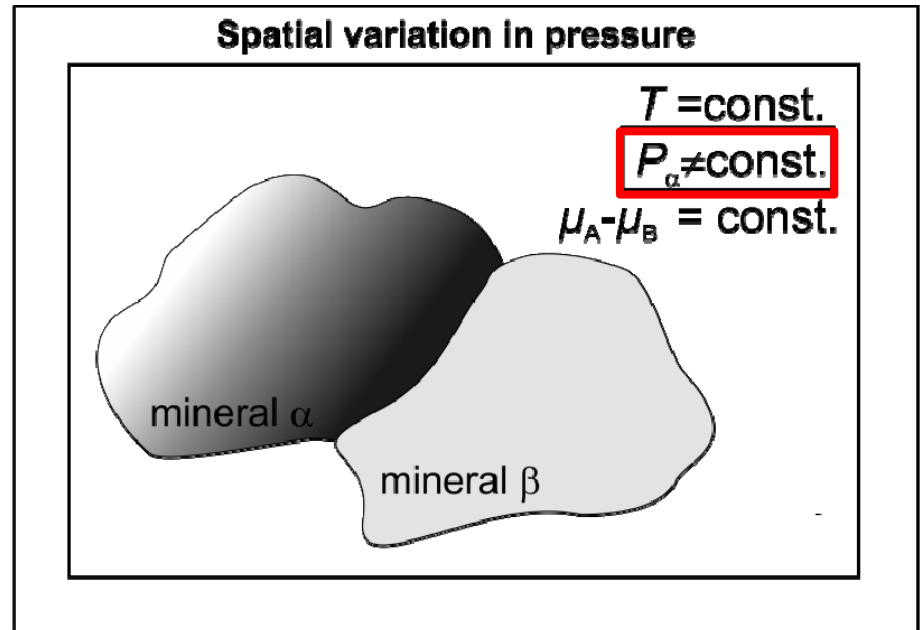
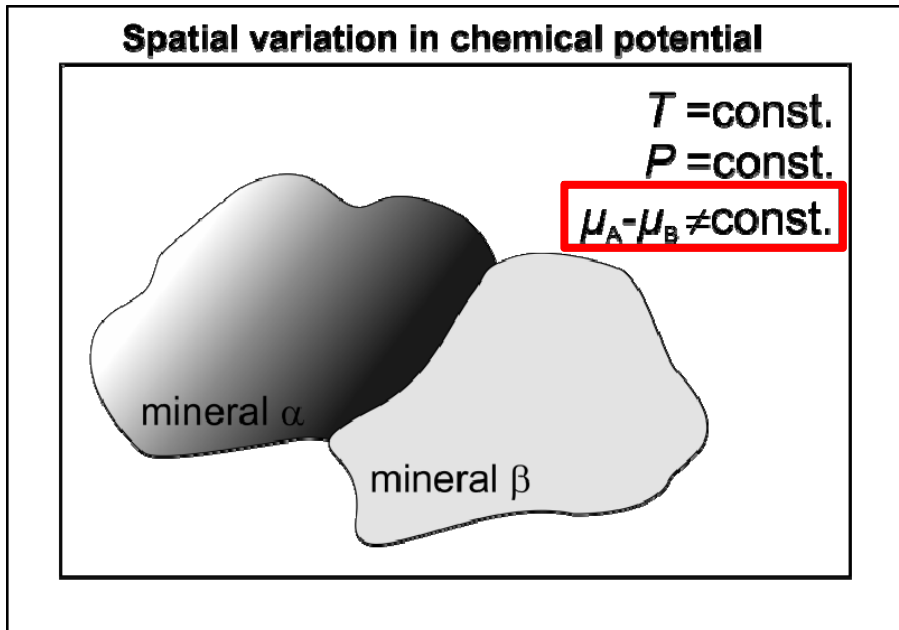
# Little bit of background first

## Chemical potential and equilibrium



$$\mu_i = \mu_0(P, T) + RT \ln a$$

# Compositional vs. pressure variations



## Chemically- controlled

fast viscous relaxation and  
slow chemical diffusion

## Mechanically- controlled

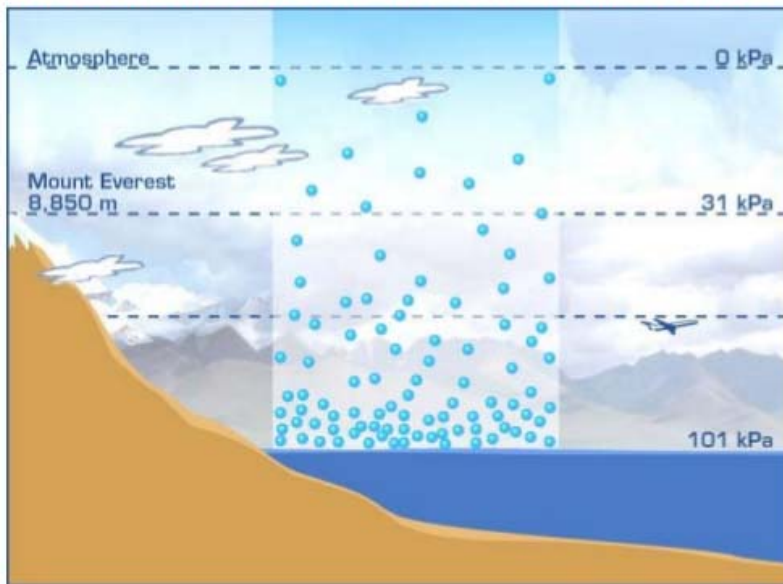
slow viscous relaxation  
and fast chemical diffusion

# The effect of pressure variation on chemical redistribution

Equilibrium under external force: **pressure variations**

Gases

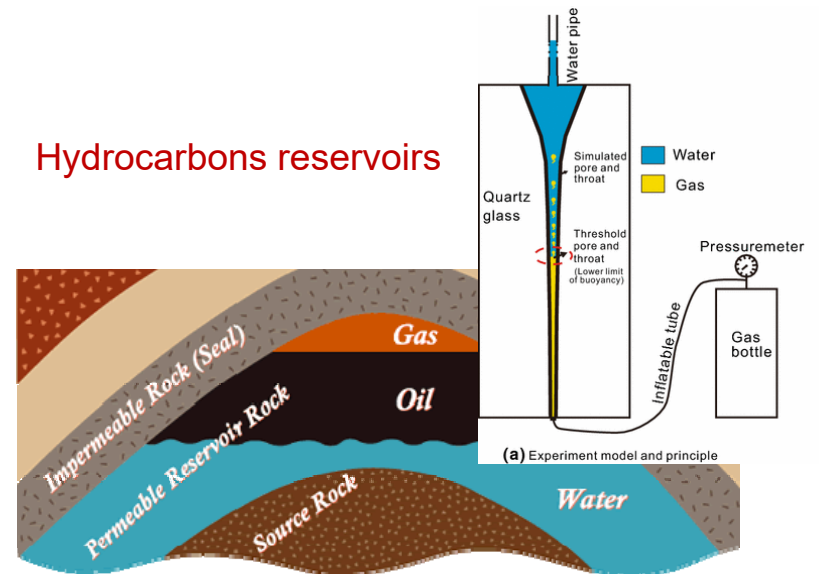
Atmospheric pressure decreases with altitude



PRESSURE

Liquids

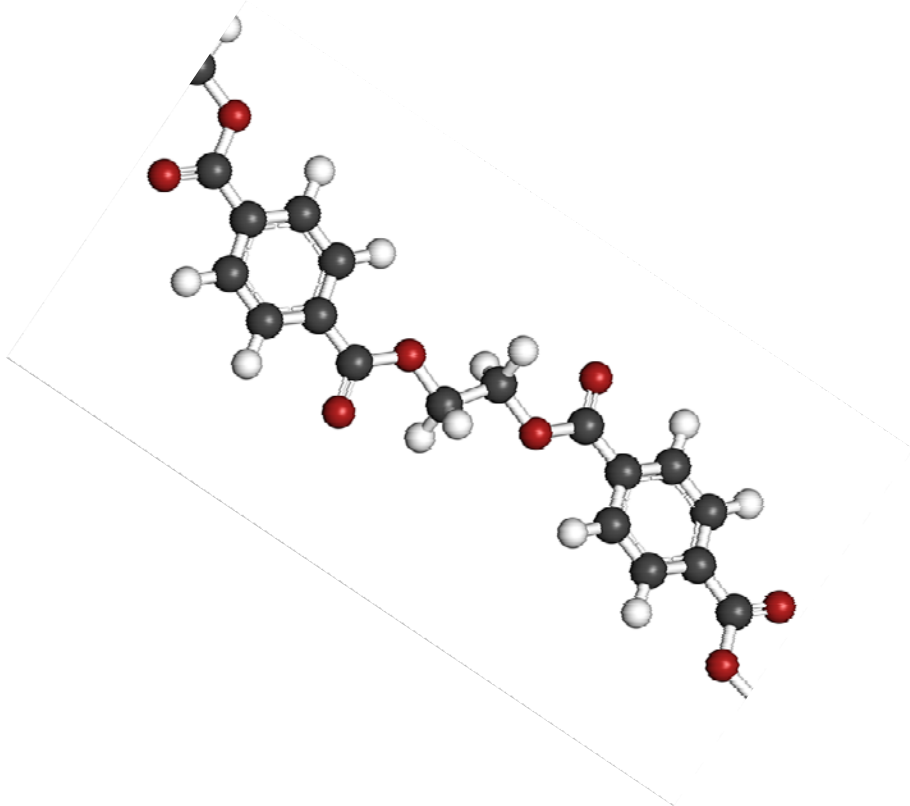
Hydrocarbons reservoirs



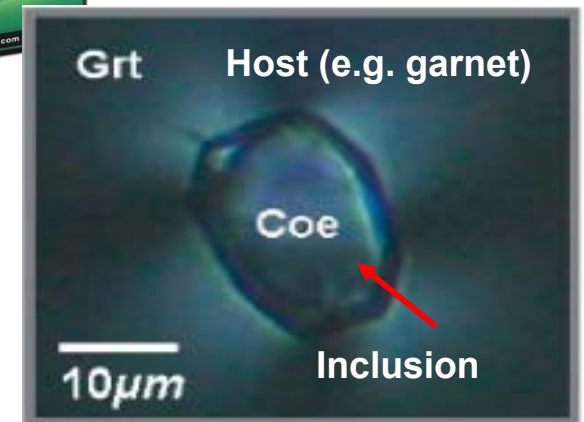
# The effect of pressure variation on chemical redistribution

Equilibrium under external force: **pressure variations**

Polymers

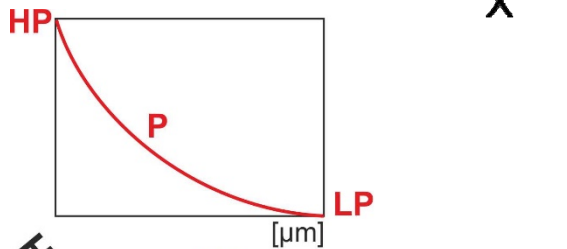
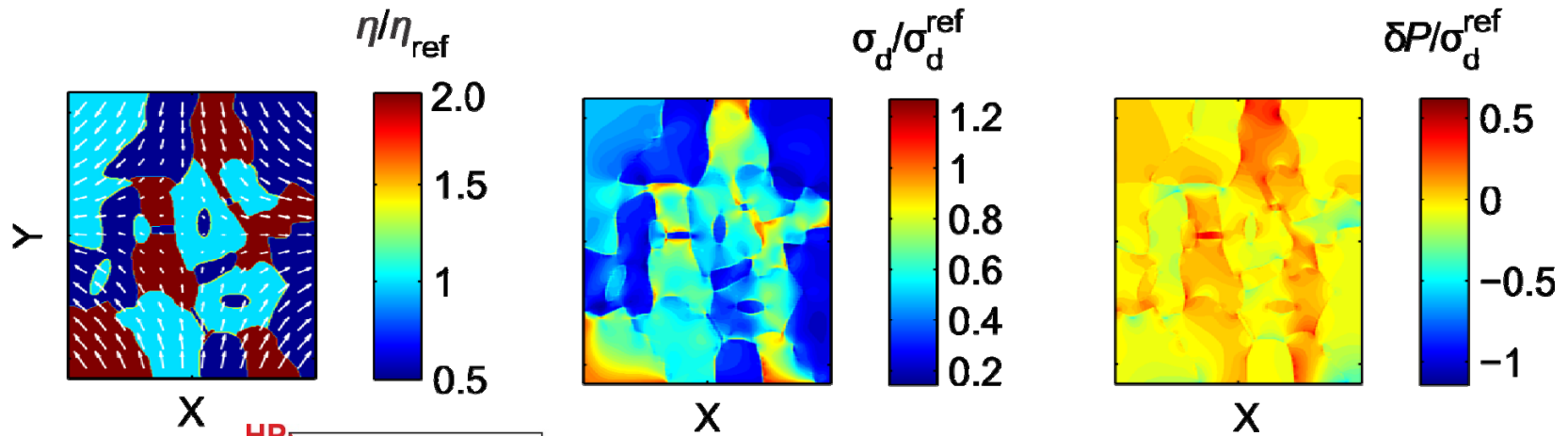


Solids???

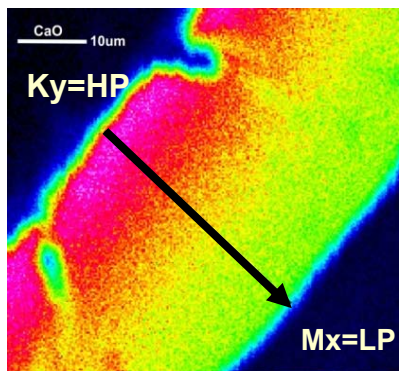
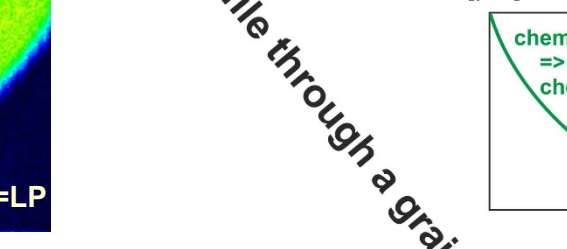
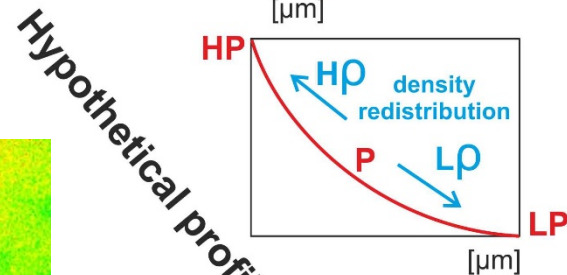


Currently and active research

# Equilibrium under pressure gradients: Unconventional barometry



The effect of pressure variation on the chemical redistribution



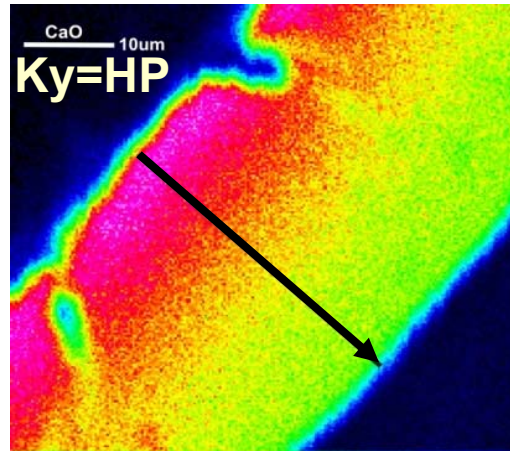
$$\mu_i = \mu_0(P, T) + RT \ln a$$

**Unconventional Rheometry:**  
Key steps of the alternative approach

# Unconventional Rheometry:

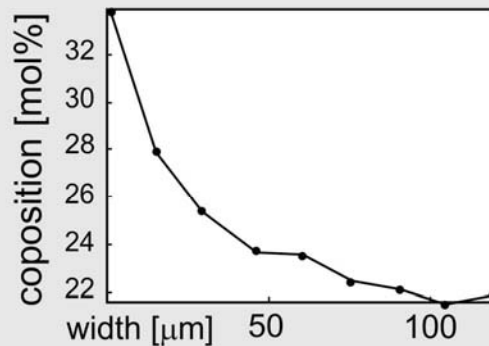
## Key steps of the alternative approach

1/ Chemically zoned mechanically-controlled microstructure, e.g. inclusion-host environment

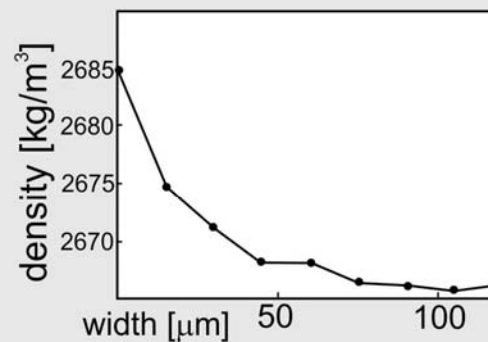


### Composition & Pressure coupling

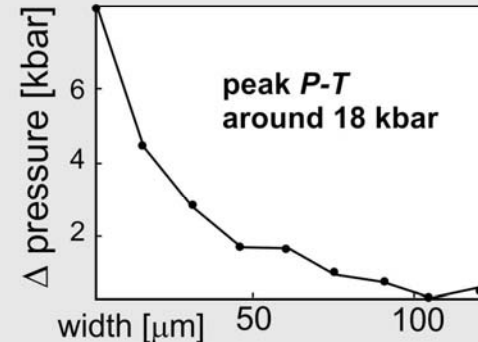
compositional profile



density profile



pressure profile

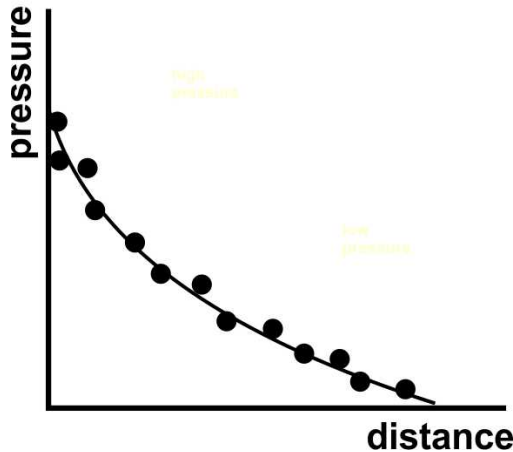




# Unconventional Rheometry:

## 2/ Classical mechanics: From pressure to stress decay

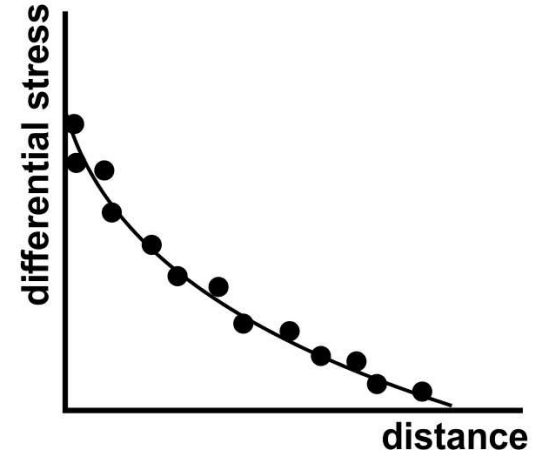
Thermodynamic method to derive pressure from direct observations



force balance →

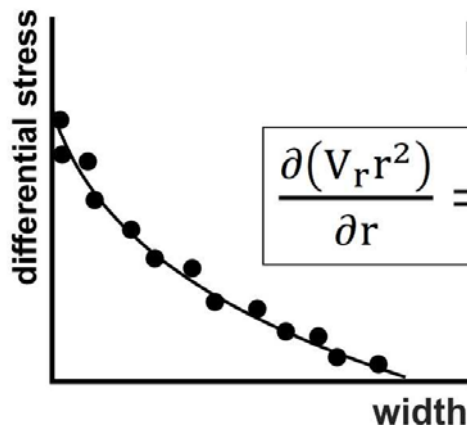
$$-\frac{\partial P}{\partial r} = \frac{2}{3} \frac{\partial \sigma}{\partial r} + 2 \frac{\sigma}{r}$$

Differential stress estimates



## 3/ Classical mechanics: from stress decay to effective stress exponent (n)

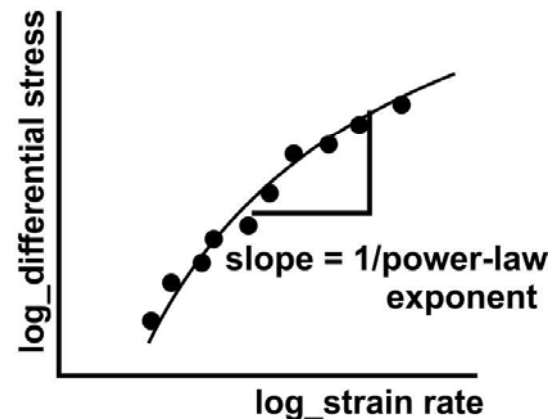
Differential stress estimates



local mass balance →

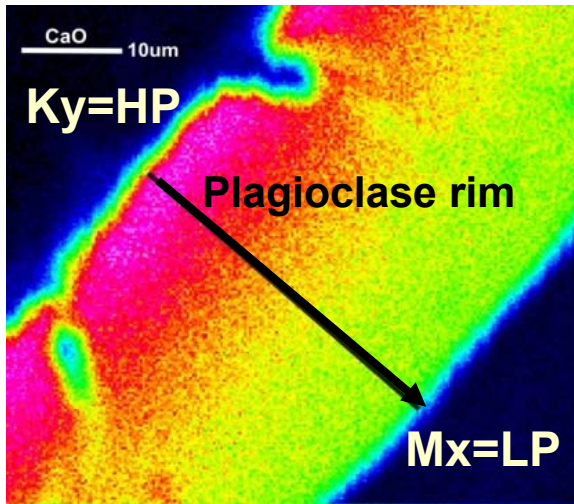
$$\frac{\partial(V_r r^2)}{\partial r} = 0 \quad \& \quad \dot{\epsilon}_{rr} = \frac{\partial V_r}{\partial r}$$

Differential stress and strain rate relation from direct observation



# Unconventional Rheometry:

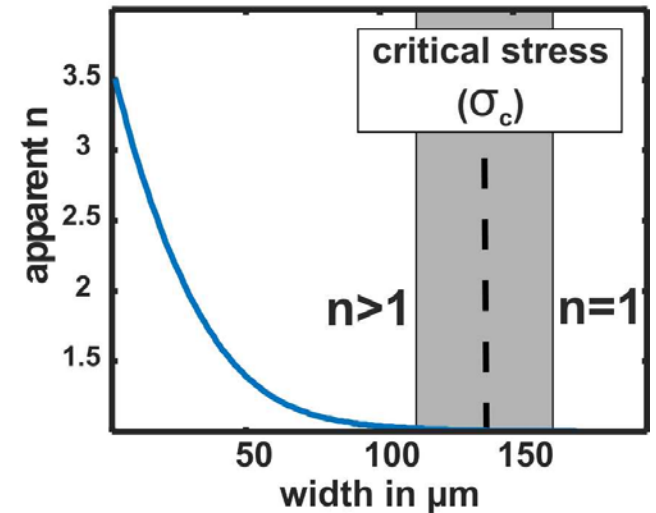
## 4/ Results: From composition to apparent stress exponent and critical stress



$$n_{app} = \frac{\partial(\log \dot{\epsilon}_{rr})}{\partial(\log \sigma)}$$

$$\sigma_c = \frac{\partial(\sigma)}{\partial(\log \dot{\epsilon}_{rr})}$$

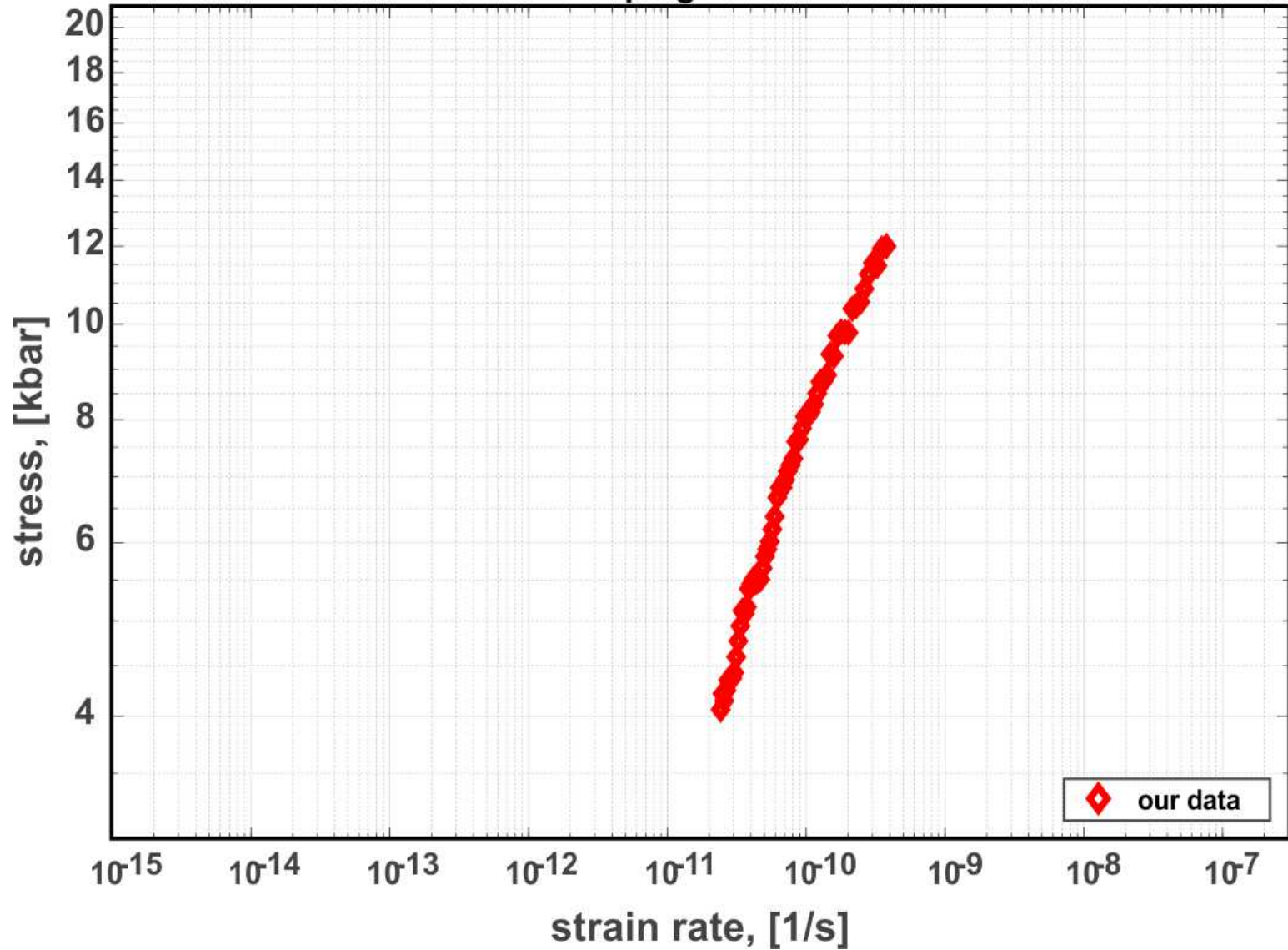
profile of apparent power-law exponent



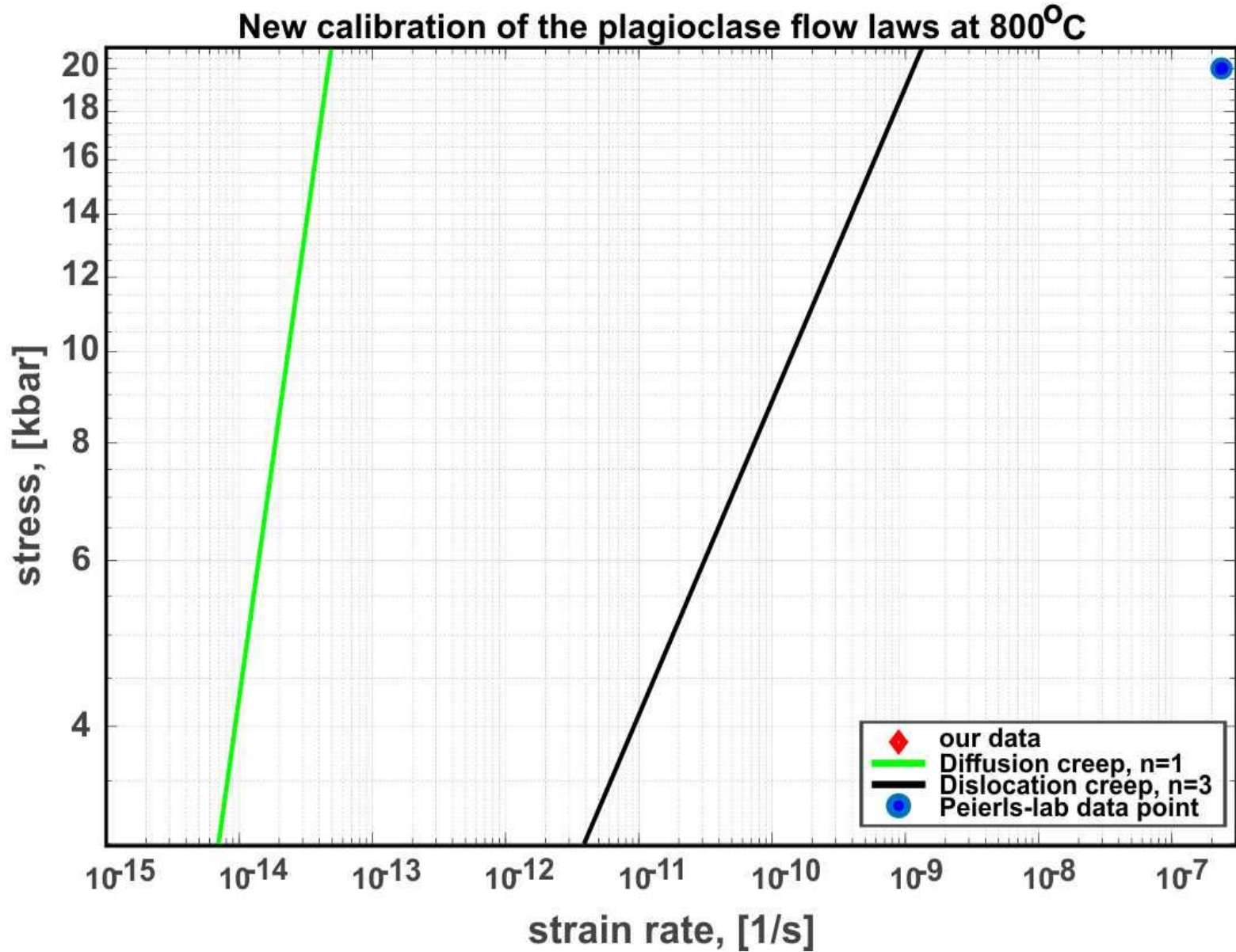
non-constant power-law exponent

# Results & Comparison

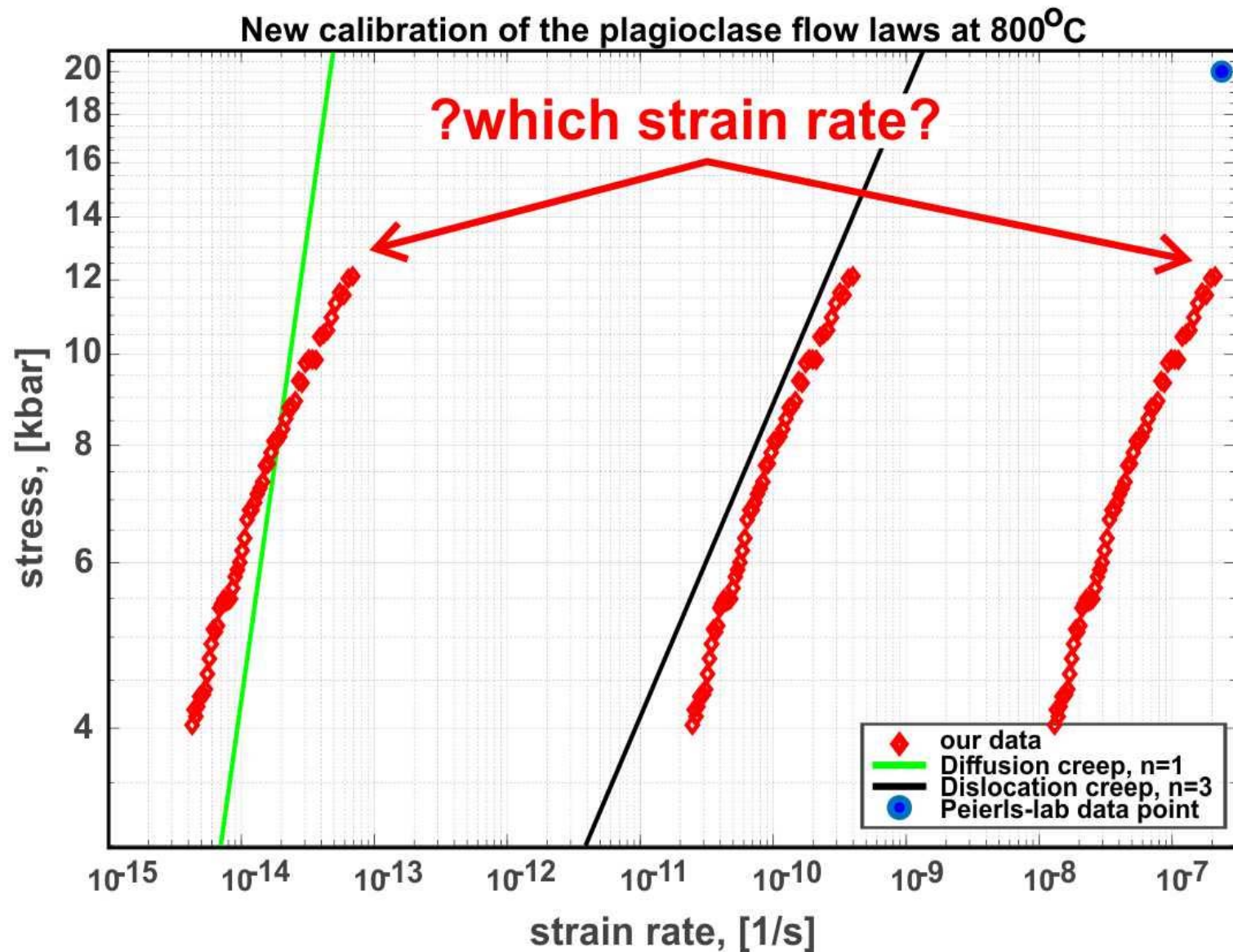
New calibration of the plagioclase flow laws at 800°C



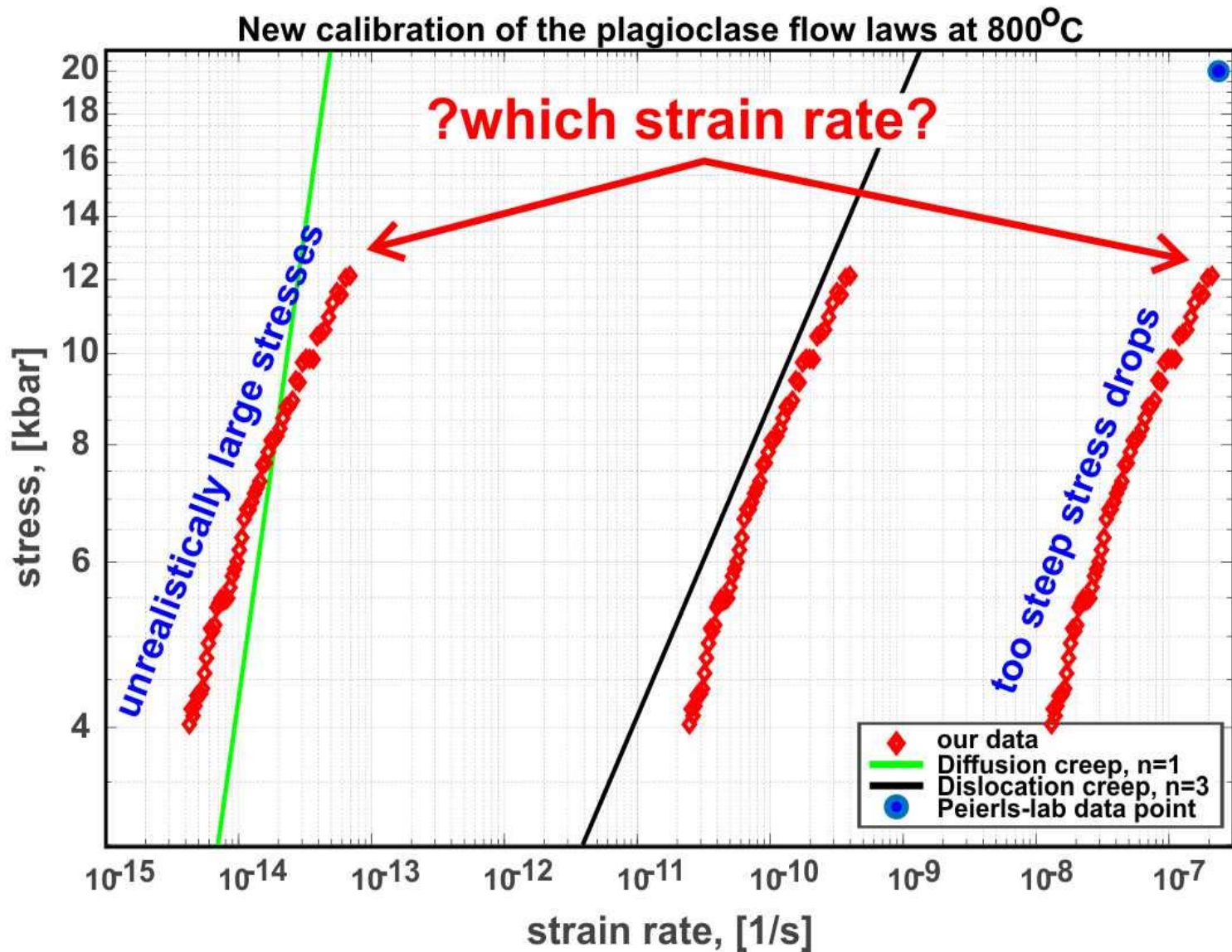
# Results & Comparison



# Results & Comparison

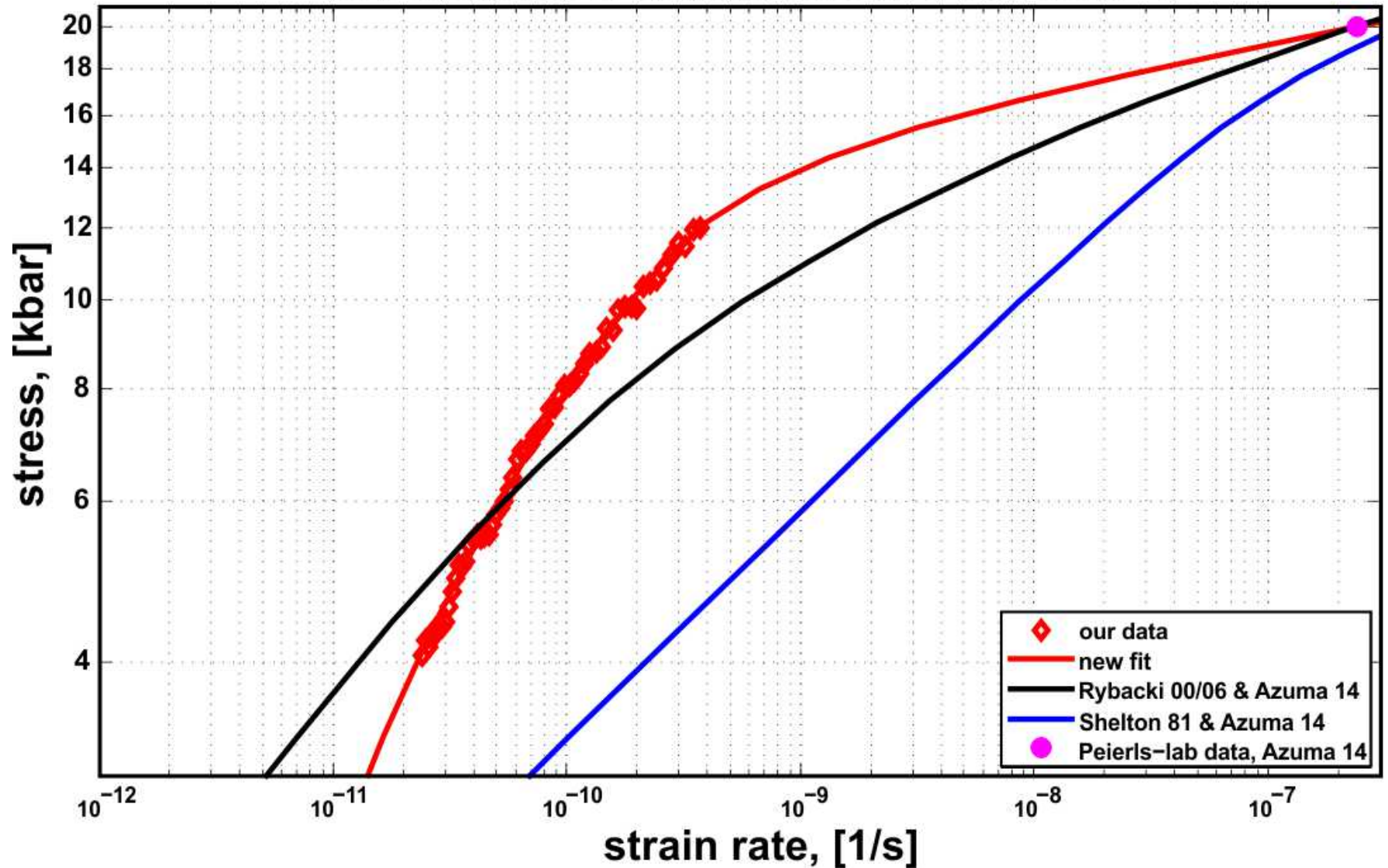


# Results & Comparison



# Results & Comparison

New calibration of the plagioclase flow laws at 800 °C



## Concluding remarks:

- **Mechanically-controlled microstructures provide information on the long-term stress state in the lithosphere**
- **New approach to infer rheology directly from natural samples, for naturally relevant T, grain size and time scale.**
- **Independent of conventional constitutive laws: based on equilibrium thermodynamics and classical mechanics**
- **Constraint for the extrapolation – an inspiration also for olivine and pyroxene**