

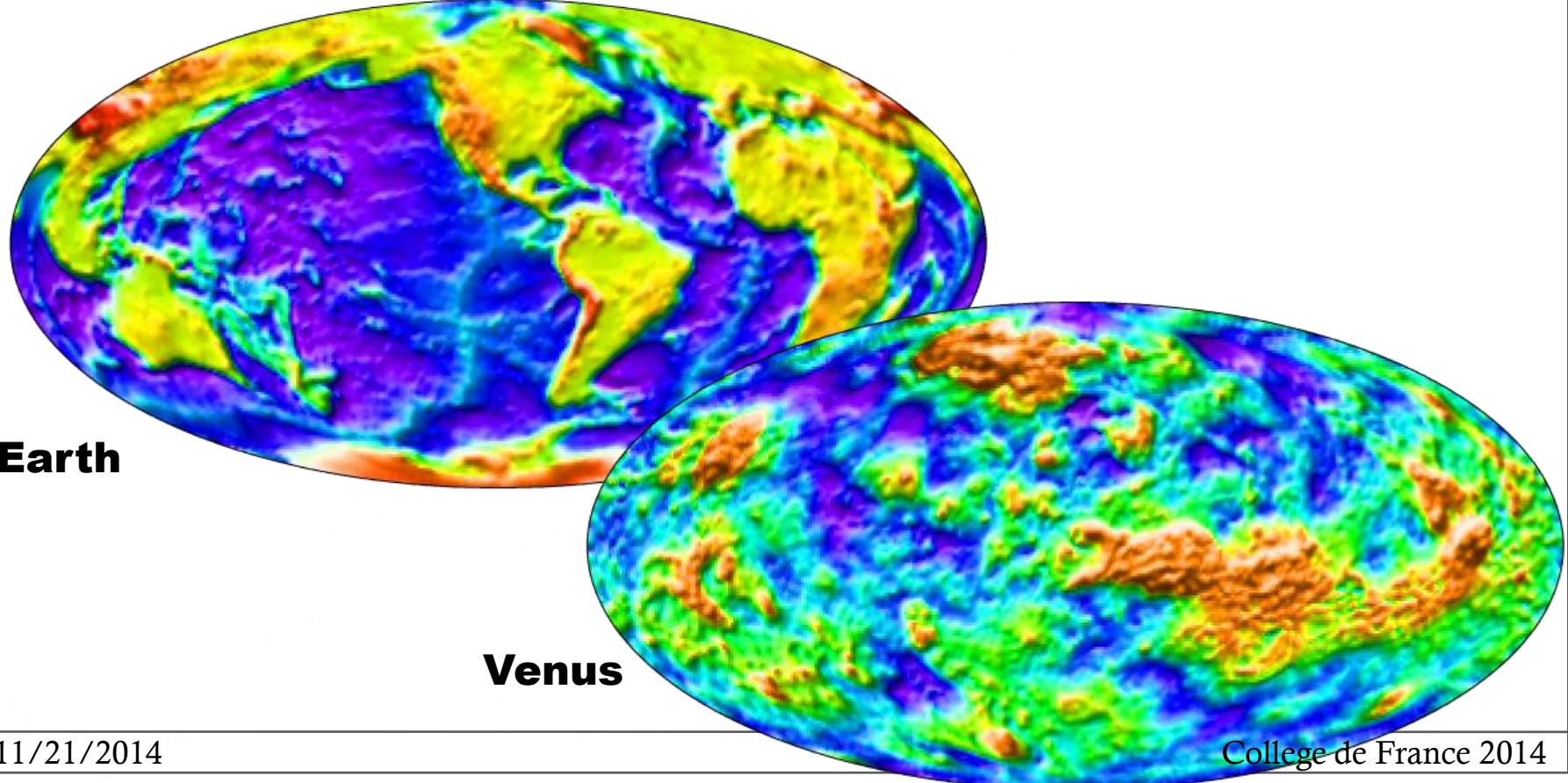


# Strain-dependent strength profiles Implication of planetary tectonics

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# Plate Tectonics: Only on Earth!

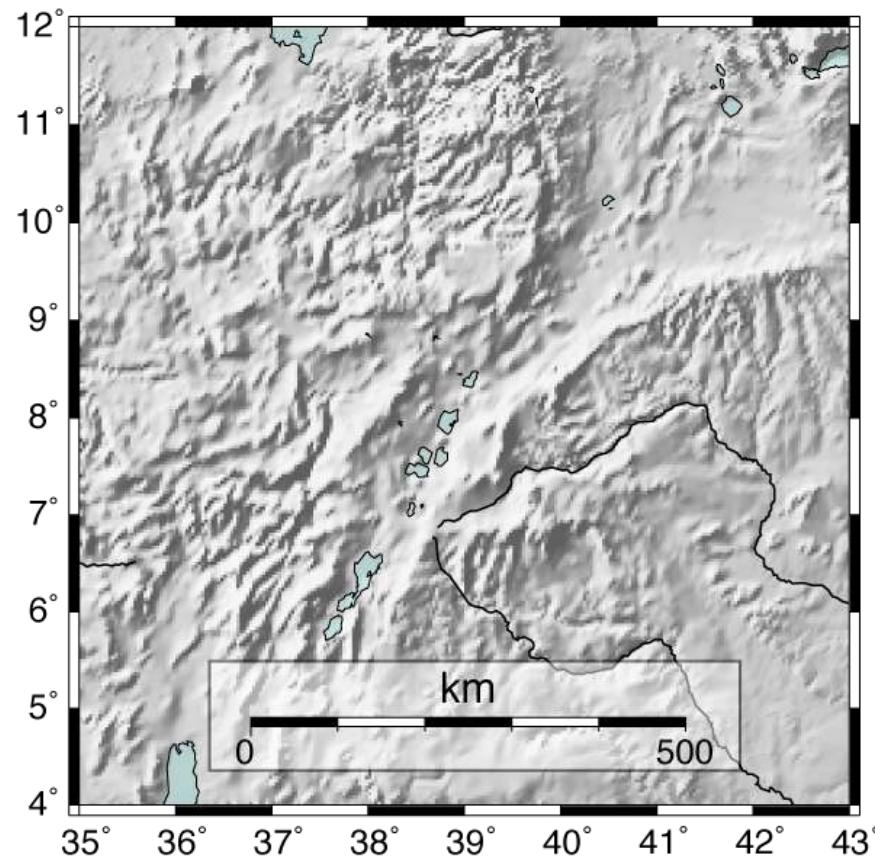
- Grand challenge problem in Solid Earth Sciences (NAS report, 2008)
- Affects planetary evolution, atmosphere, life?



# Extension on Earth and Venus

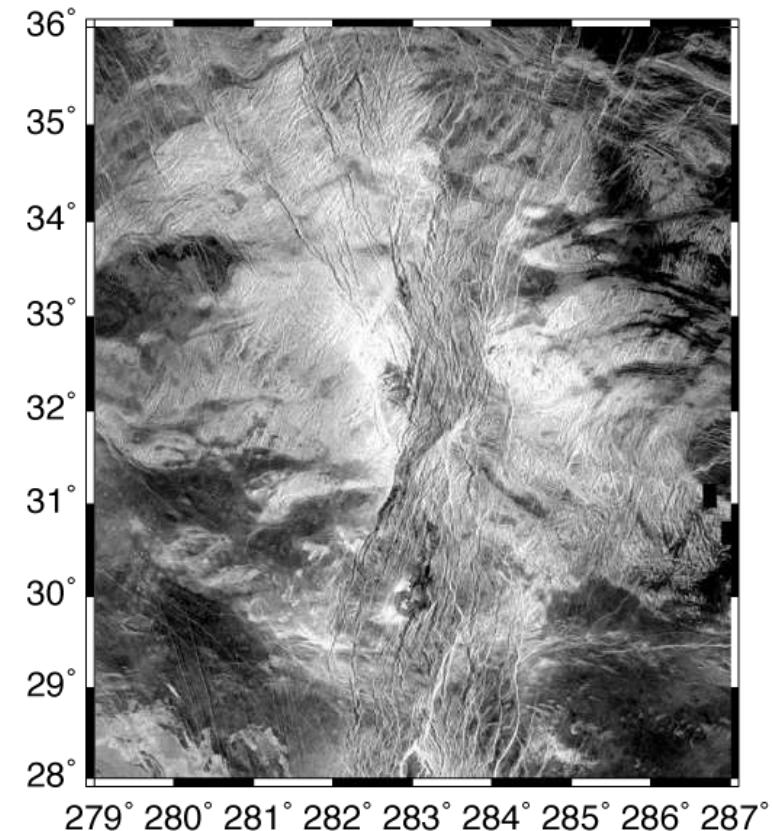
## Extension

### Earth



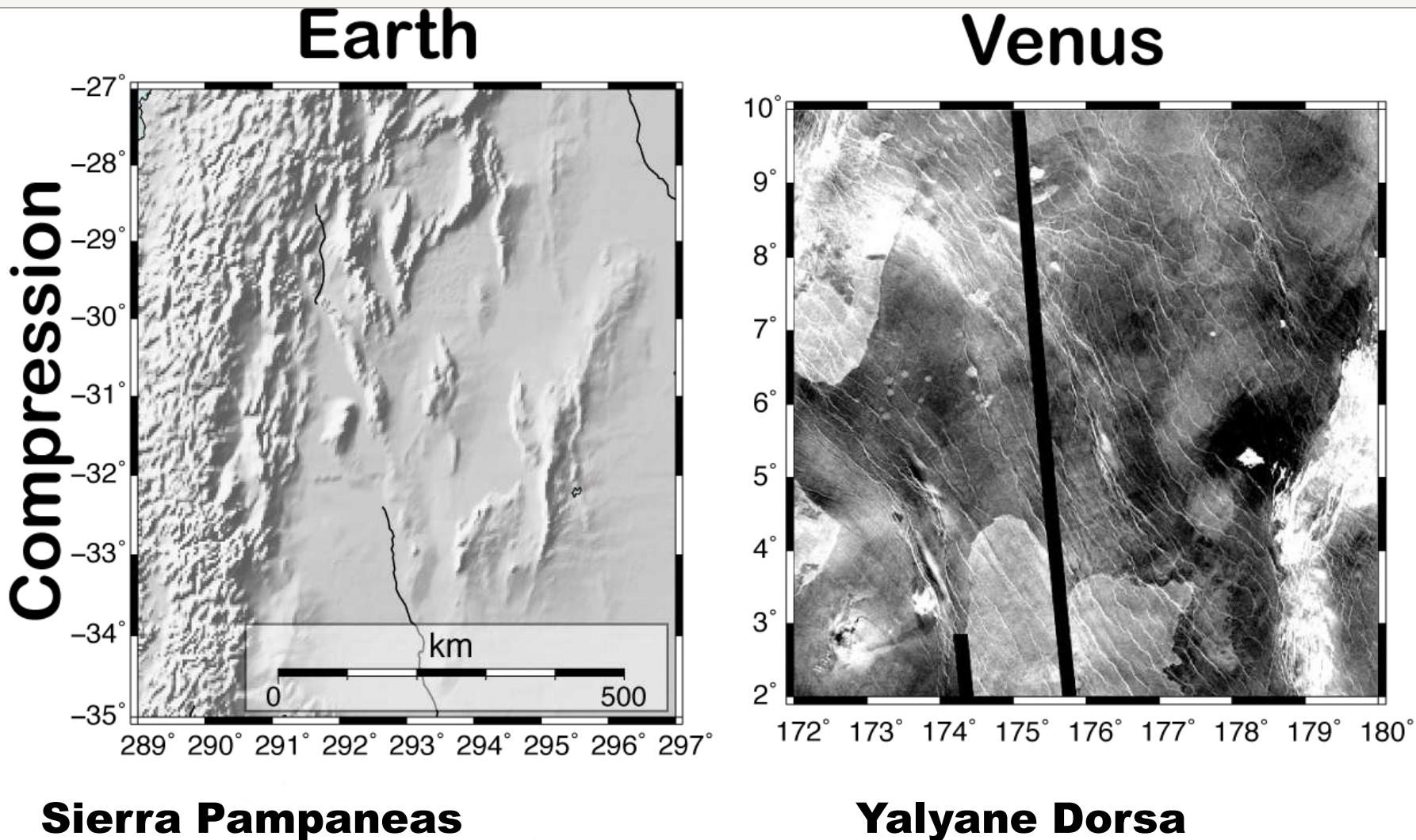
### Main Ethiopian Rift

### Venus

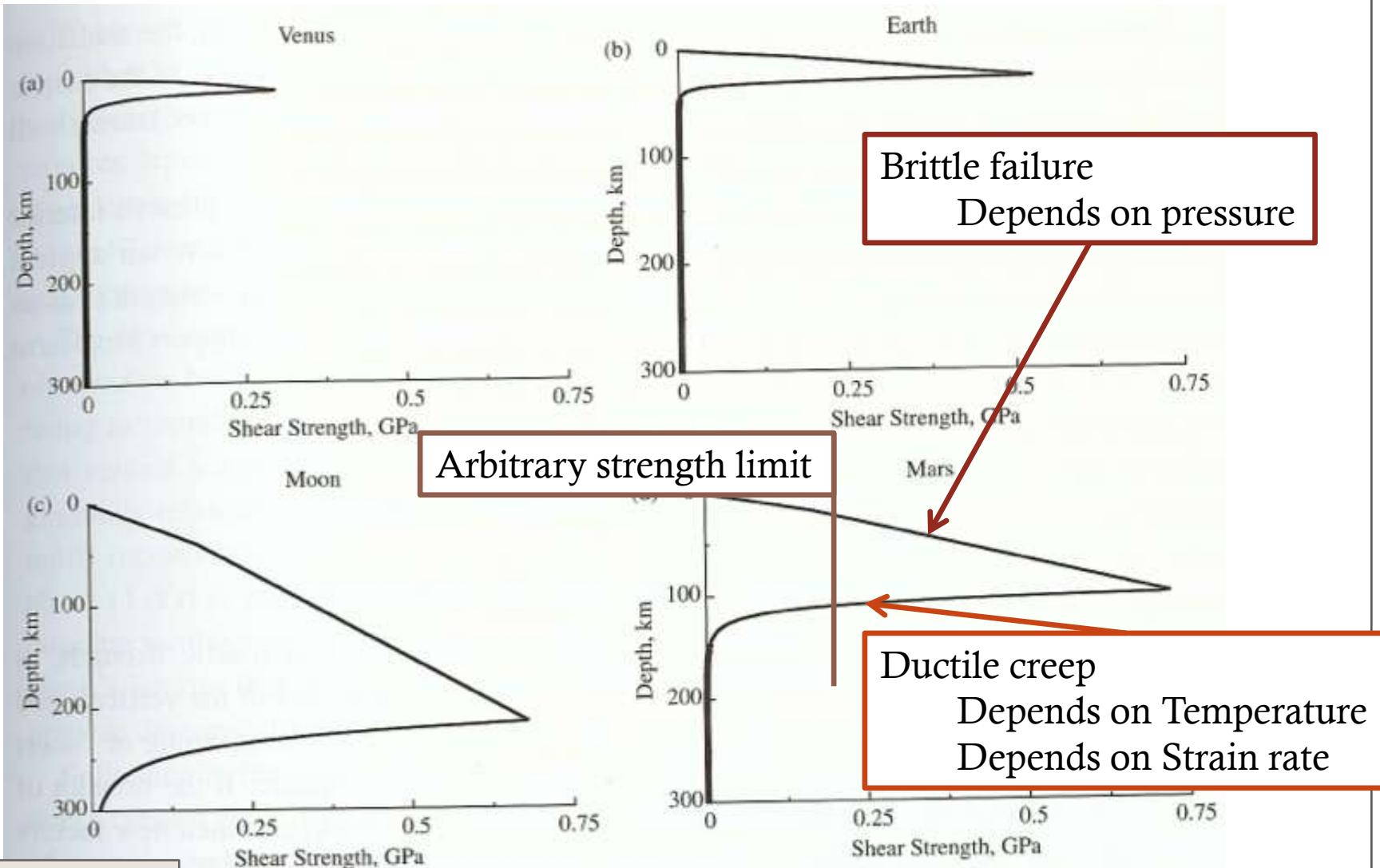


### Devana Chasma

# Shortening on Earth and Venus

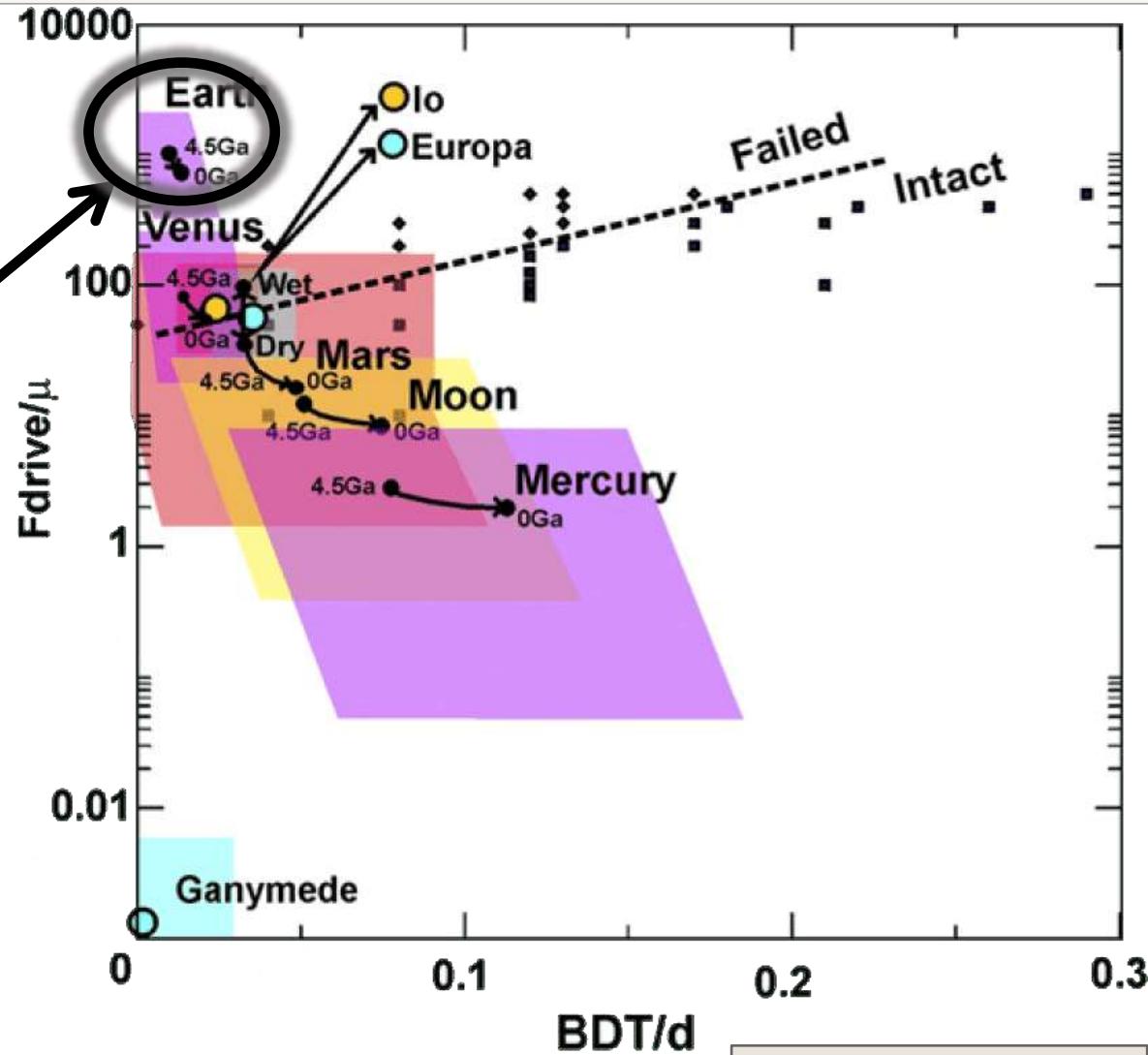


# Strength envelopes (classical)



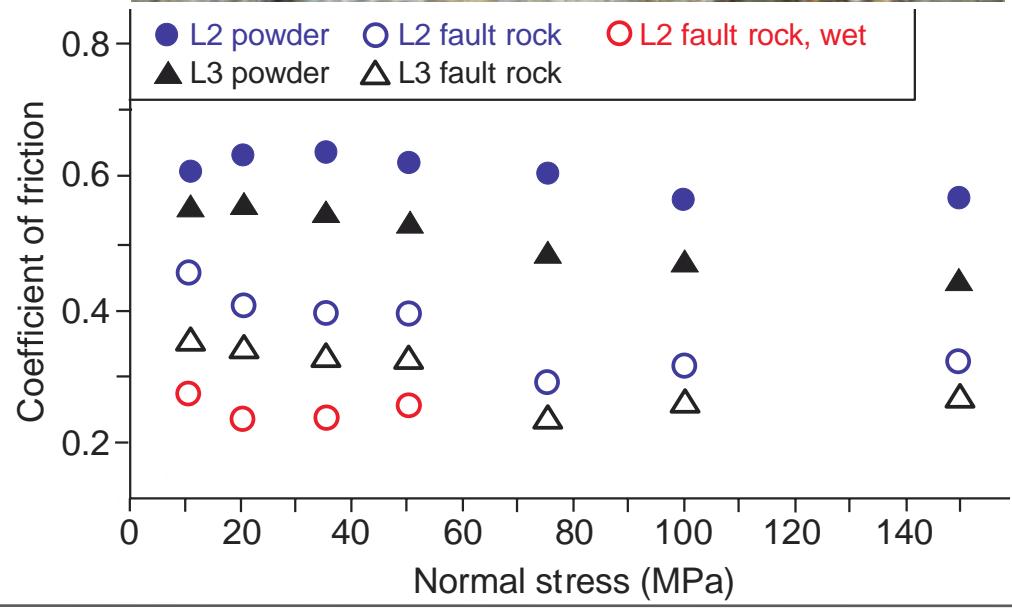
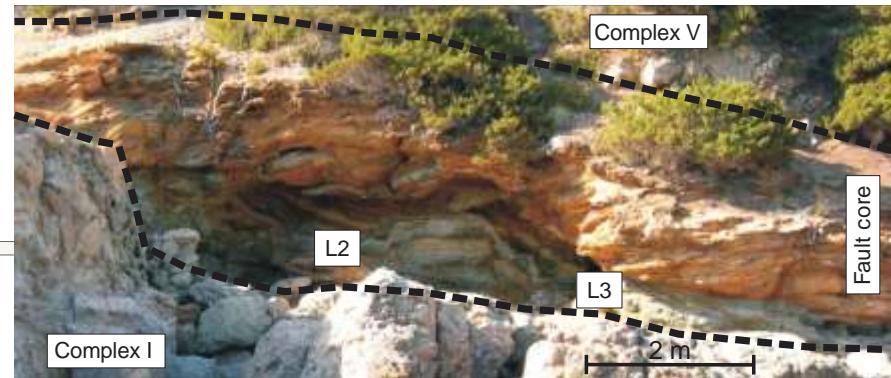
# A plate tectonic recipe

- Vigorous convection
  - Weak deformation zones
- Weakened arbitrarily**
- *What mechanism weakens plate boundaries?*
  - *Which mechanisms are active only on Earth*
  - *Is the “strength” all that is needed?*



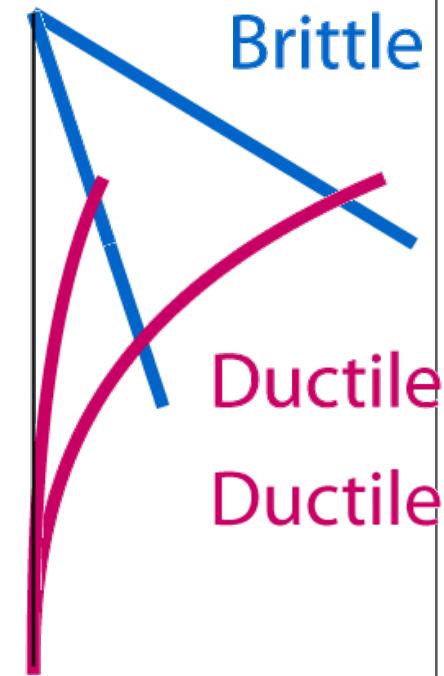
# Weakness of brittle faults

- Fault rocks contain weak minerals (smectite, talc and minor chlorite).
- Low friction results from slip on a network of weak phyllosilicate-rich surfaces that define the rock fabric



# Which level is weak?

- Low yield strength of the lithosphere
  - Solomatov (2004):  $<\sim 3\text{ MPa}$
- Low brittle strength
  - Friction coefficient of 0.15 (O'Neill et al. 2007)
    - Serpentine? (Moore et al., 2007)
    - High pore fluid pressure
- Low ductile strength
  - Necessary to reconcile low coefficient of friction and depth to brittle-ductile transition
  - Ductile shear zones



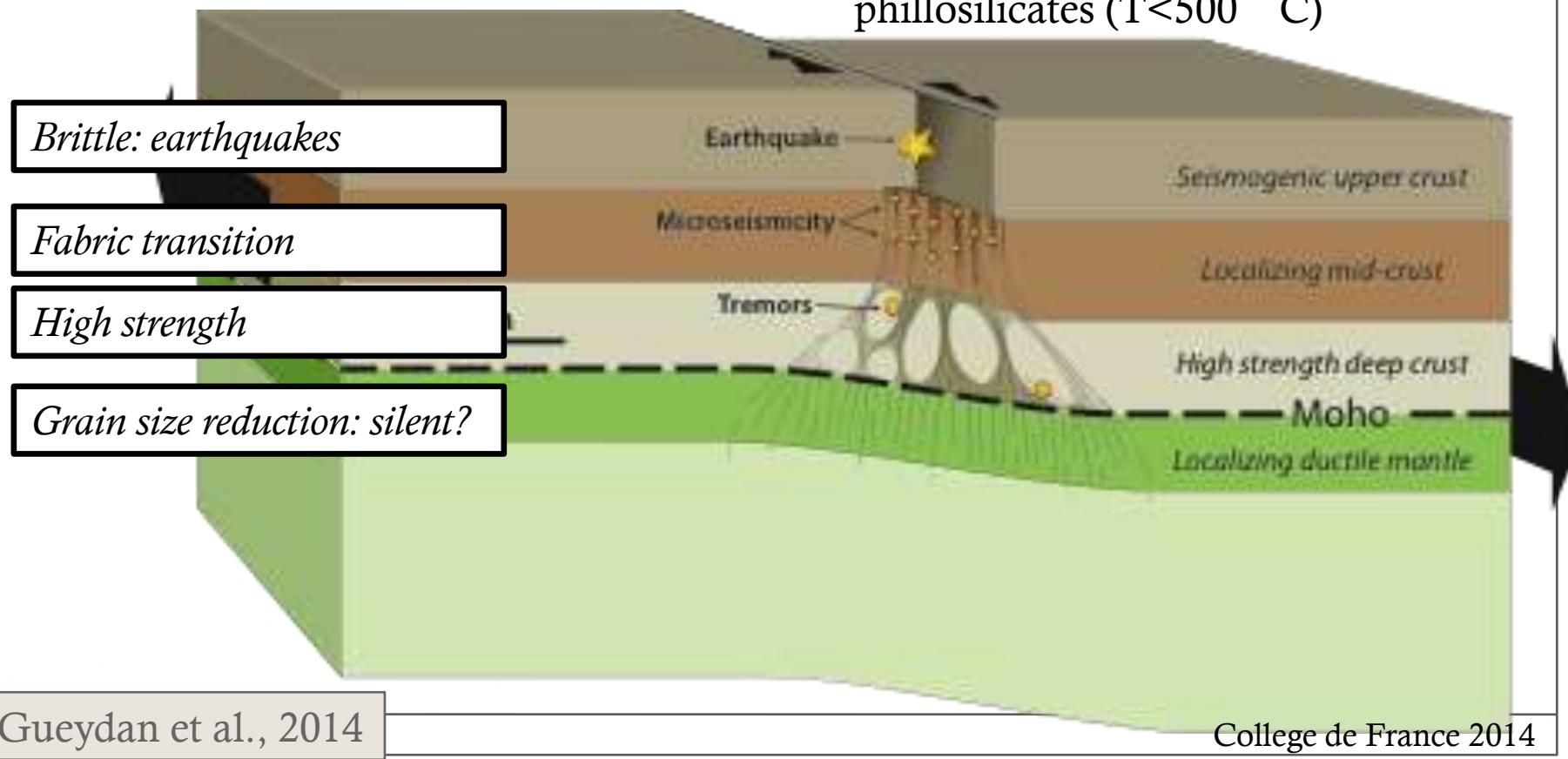
# The San Andreas Fault at depth?

Brittle failure

Need fluids and/or serpentine

Ductile failure

In the mantle when dis-GBS is possible ( $T < 700^\circ \text{ C}$ )  
In the crust in presence of philllosilicates ( $T < 500^\circ \text{ C}$ )



# Ductile shear zone structure

- Requires change in state or environment
  - Temperature
  - Grain size
  - Interconnection of weak phase
  - Abundance of weak phase
  - Composition (metamorphism, melt)

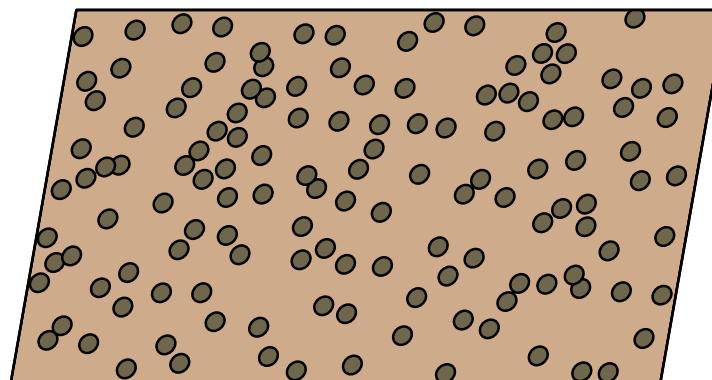


# Fabric and rheology

Protolith (uniform strain)



Shear zone (uniform stress)



Strength controlled by strong phase

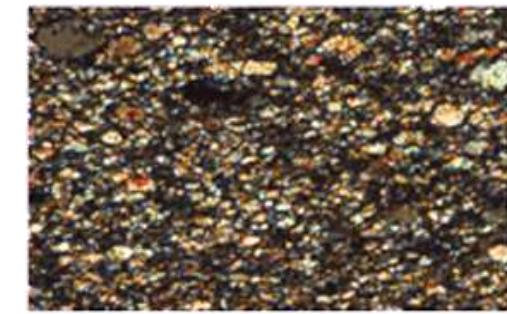
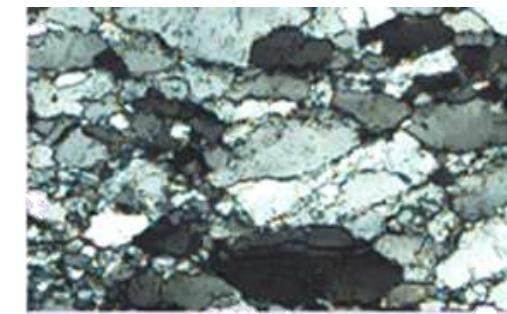
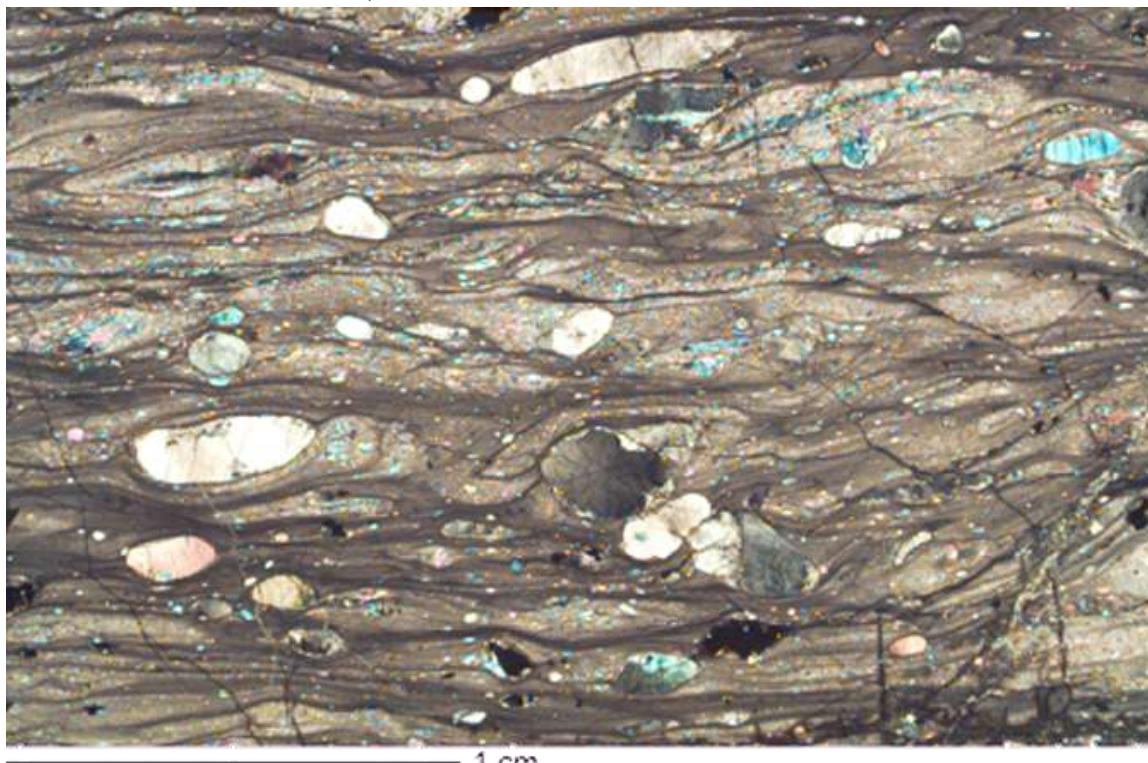


Strength controlled by weak phase

# Grain Size Reduction

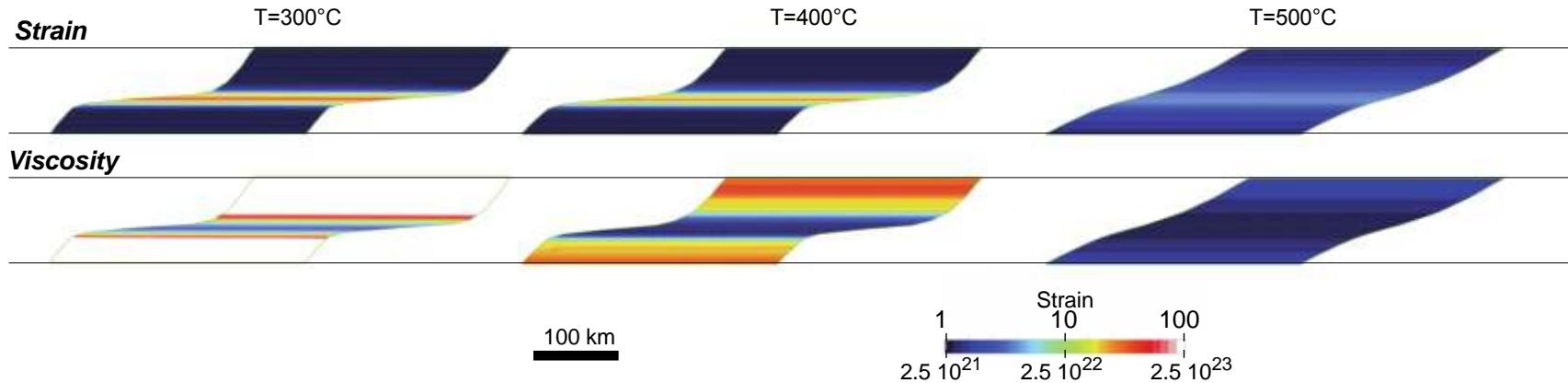
- Mylonite in oceanic peridotite, Shaka Fracture zone

Shear direction  
↔

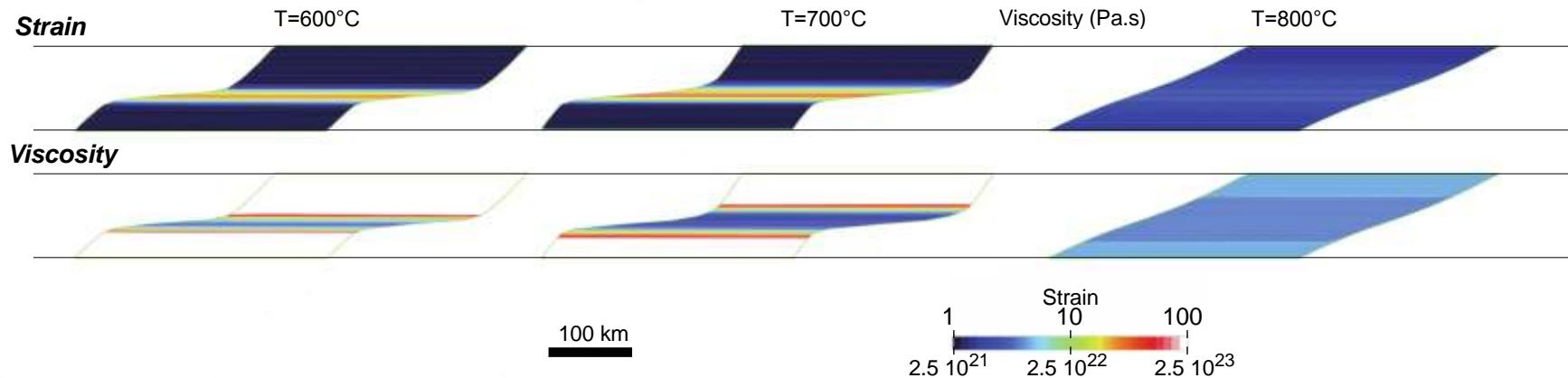


# Shear zone development

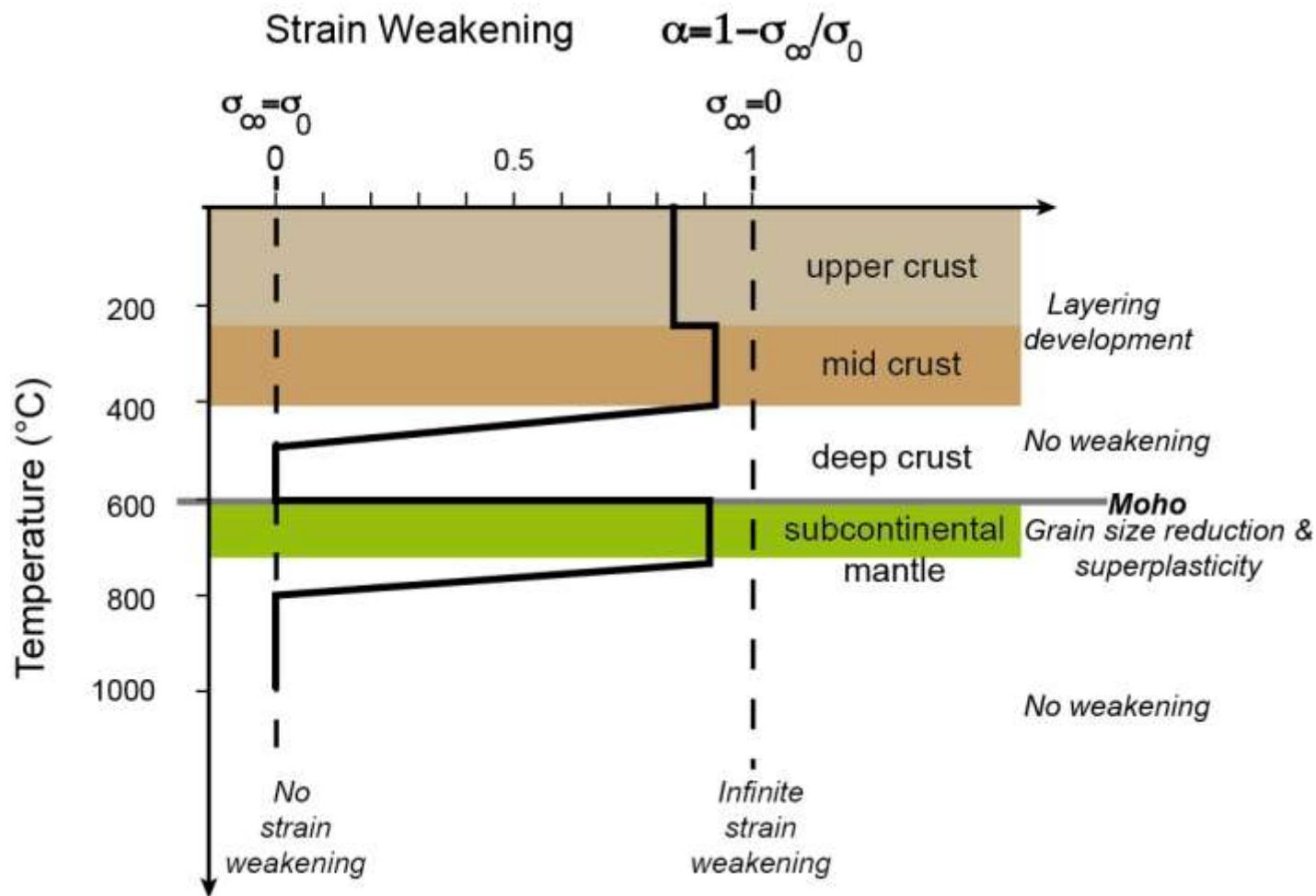
In the crust: Weak phase interconnection if phyllosilicates



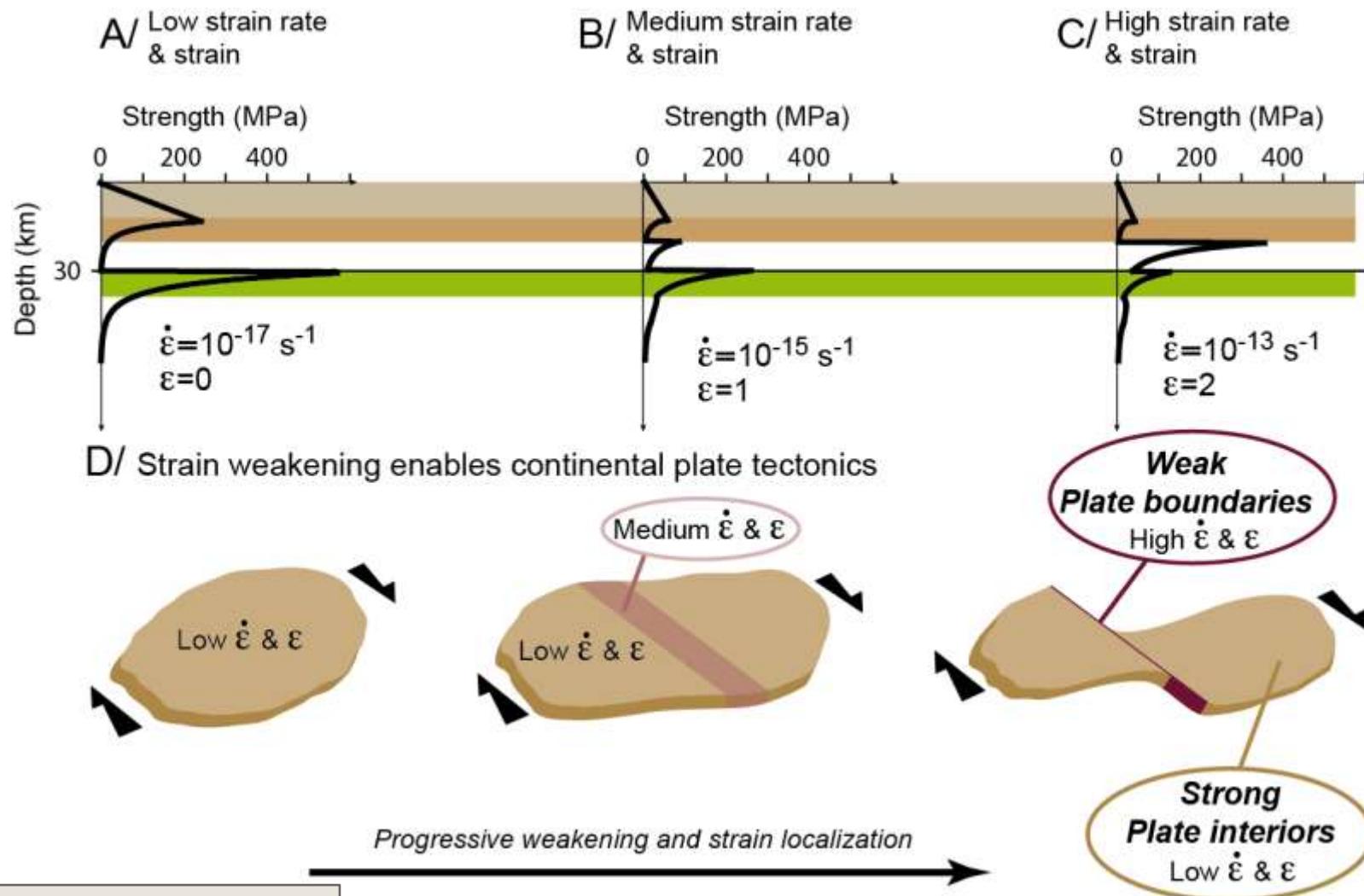
In the mantle: grain size reduction if dis-GBS



# Final weakening

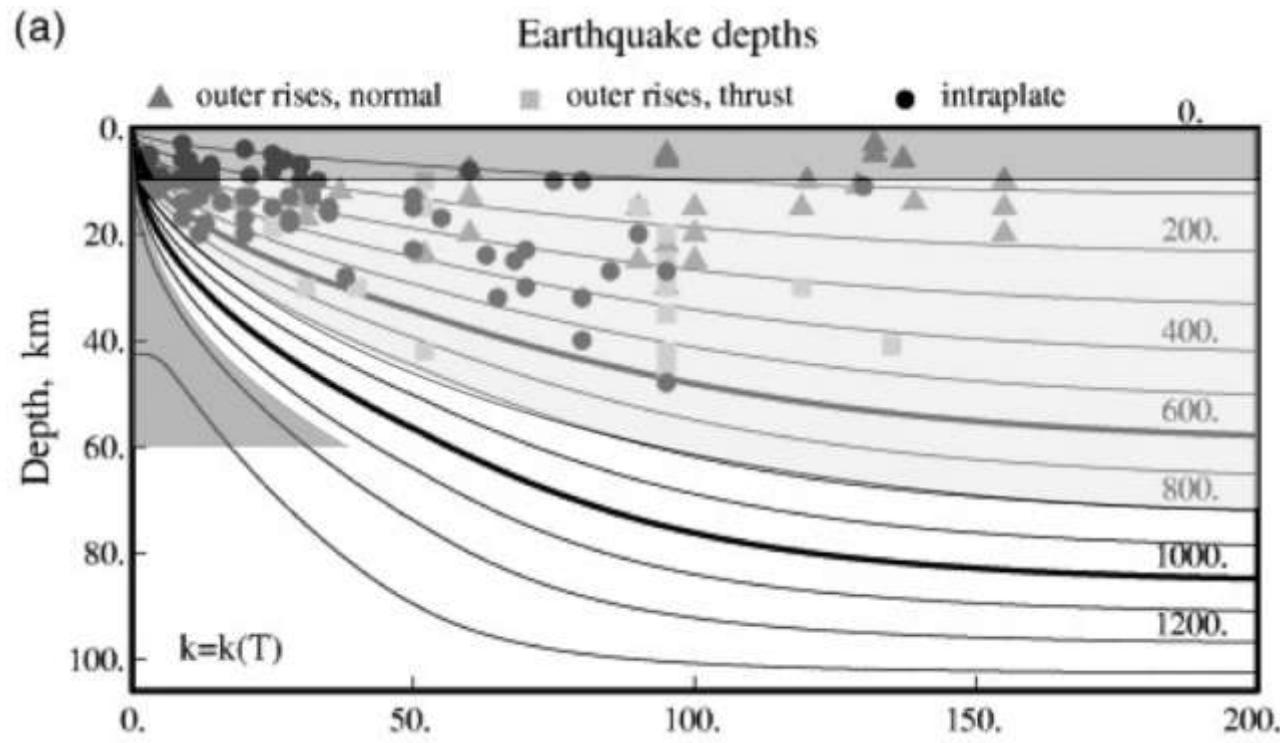
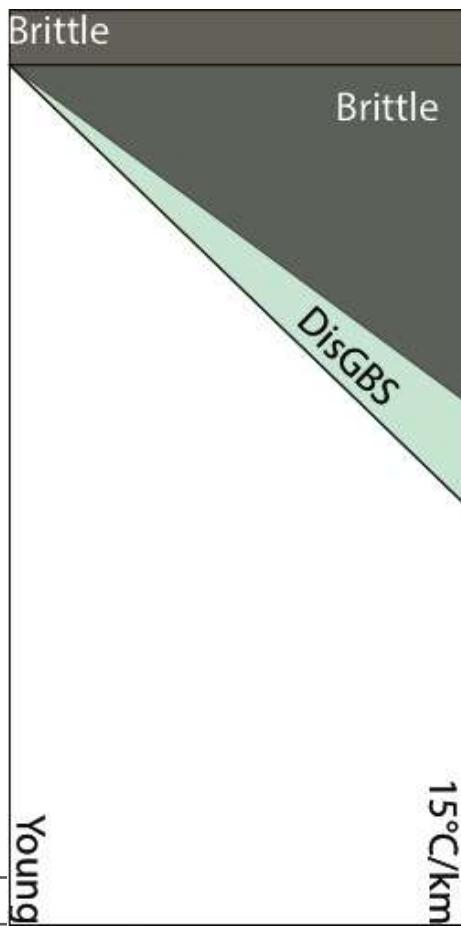


# Strain-dependent strength envelopes



# Oceanic lithosphere

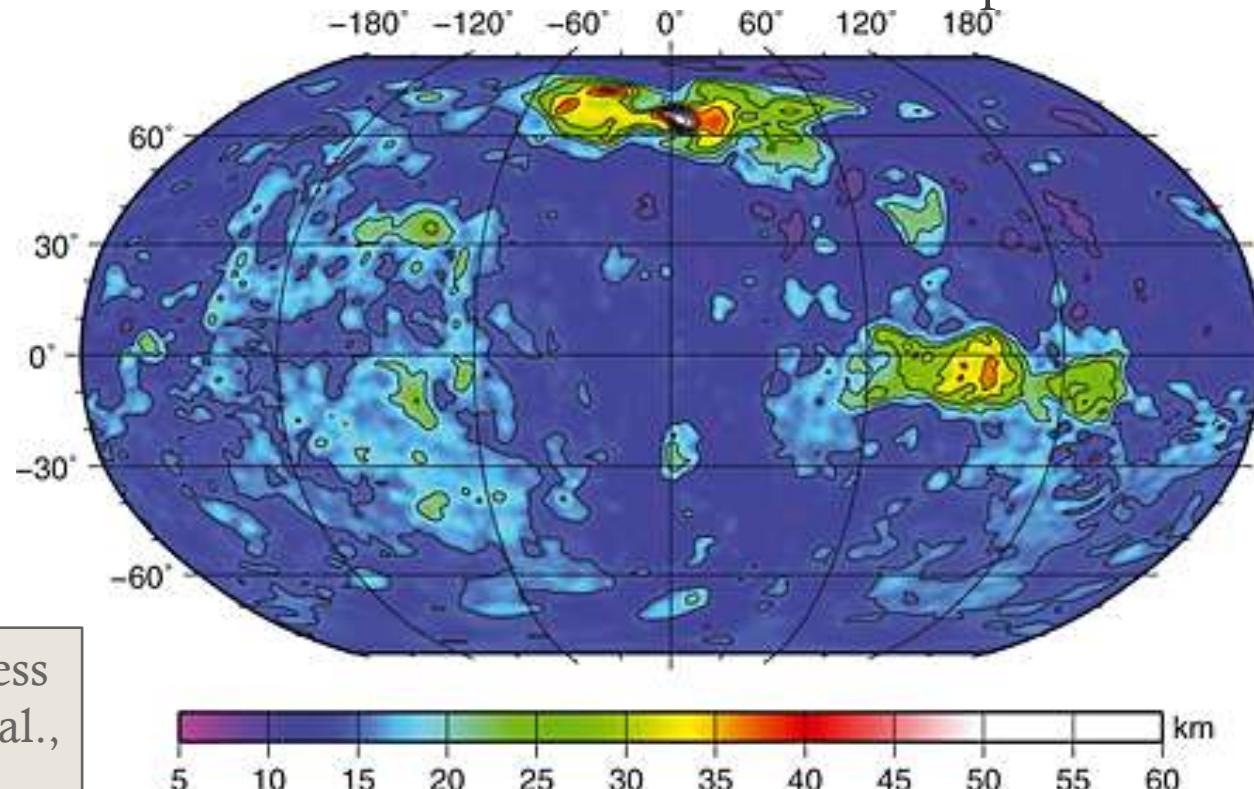
- Crust: mostly brittle
  - Hydrated minerals and high pore fluid pressure
- In the mantle
  - Dis-GBS starts essentially at zero age
  - Older lithosphere has a thicker localizing mantle
- Melt embrittlement?



Thermal structure from McKenzie et al., EPSL 2005

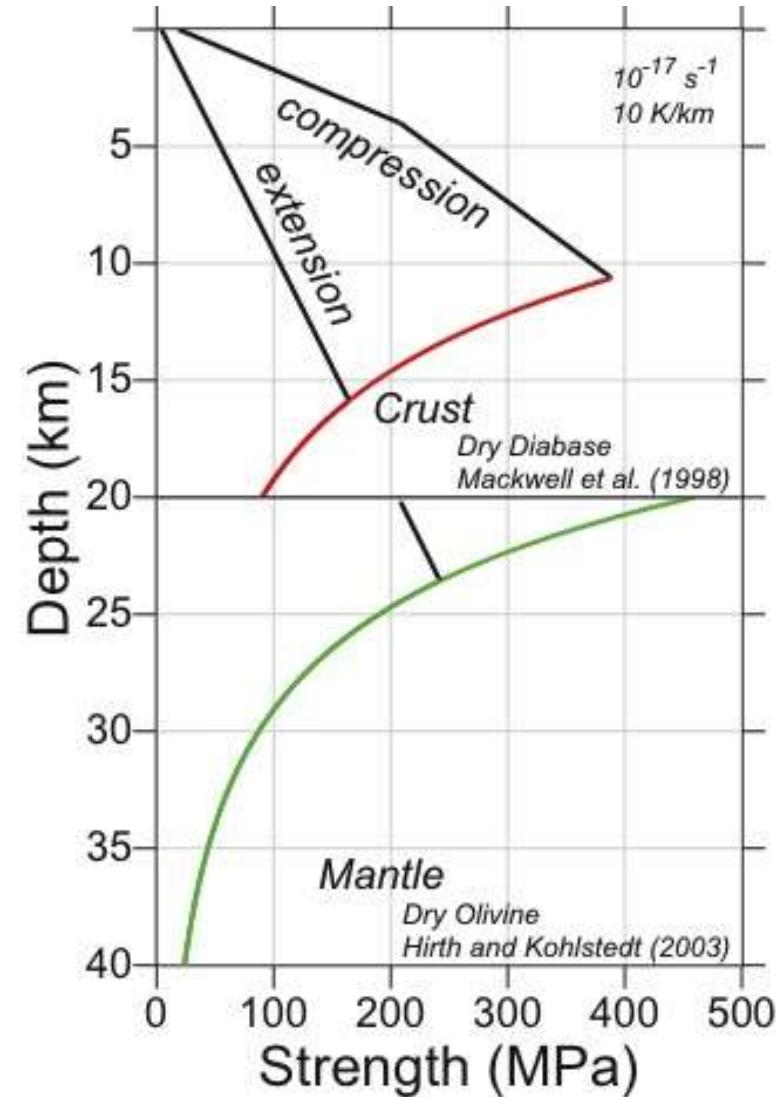
# Venus

- Mantle
  - Dis-GBS possible if low enough geotherm / thin crust
    - 15 K/km: 22km
    - 10 K/km: 34 km thick
- Crust
  - High surface temperature, no water: no phyllosilicates
  - Shallow brittle failure
  - More important in extension



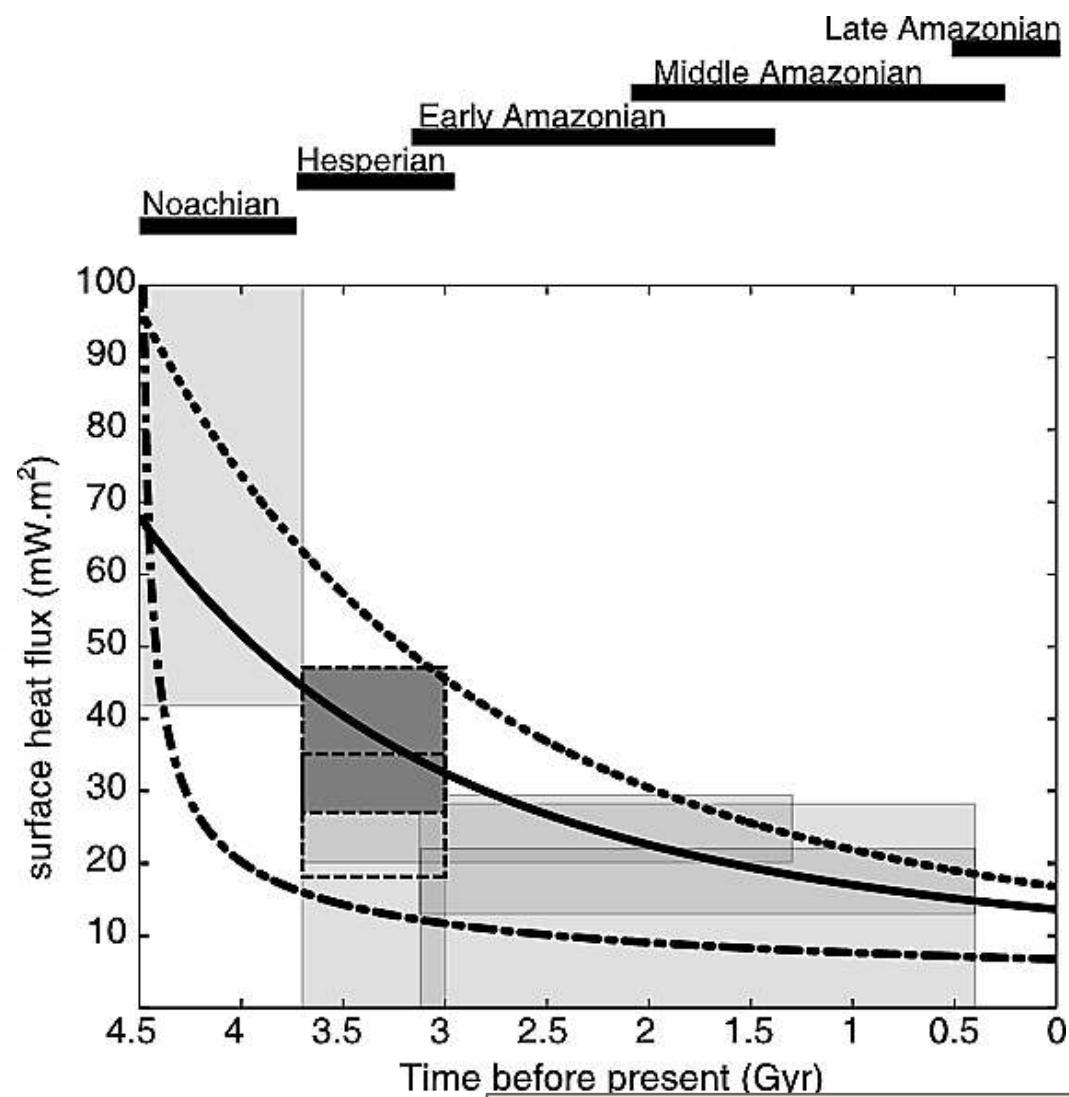
# Localization easier in extension

- Deeper brittle layer => larger fault offsets
- Possibility of melting => melt weakening and melt embrittlement
- Geometrical effects (stress focusing)

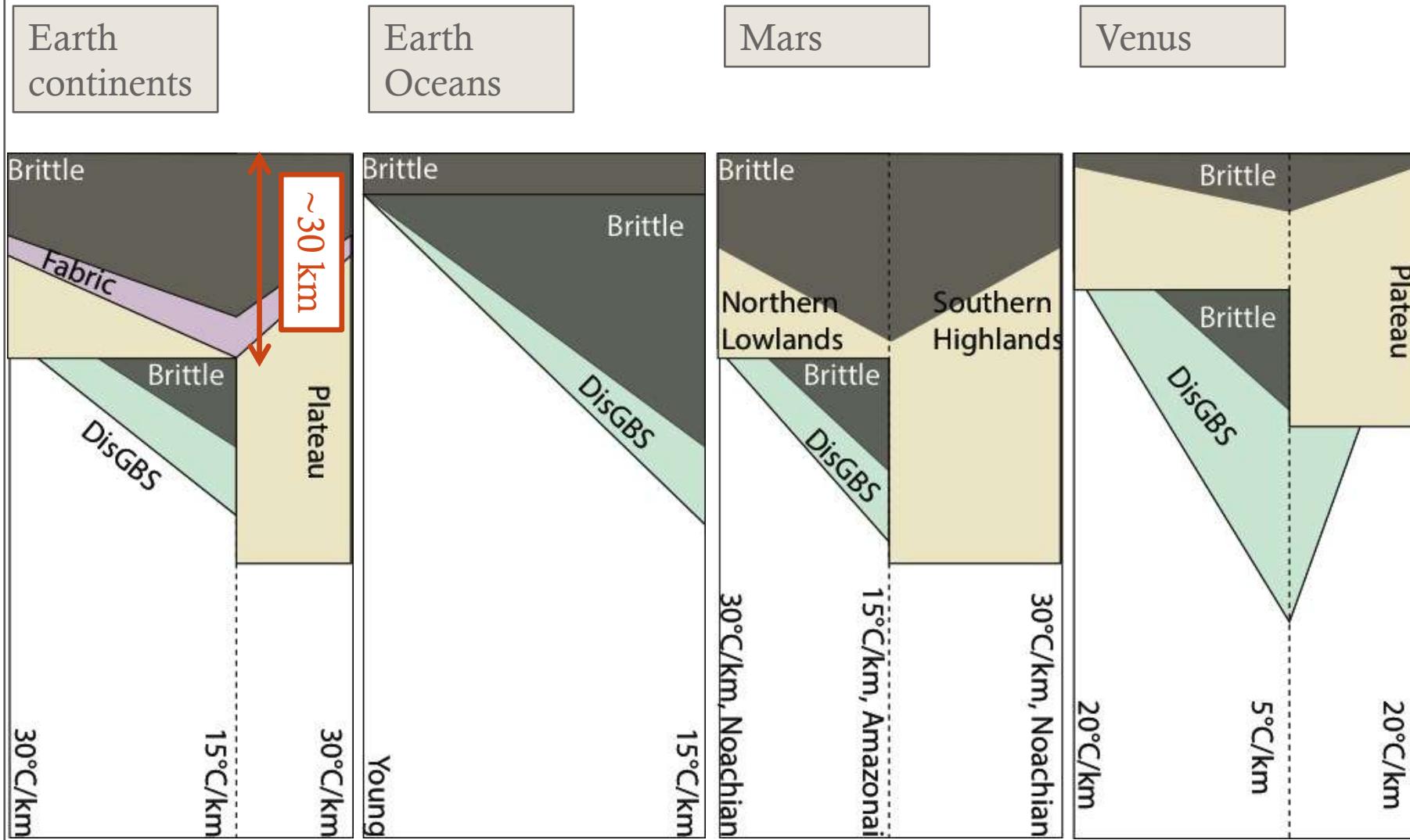


# Mars

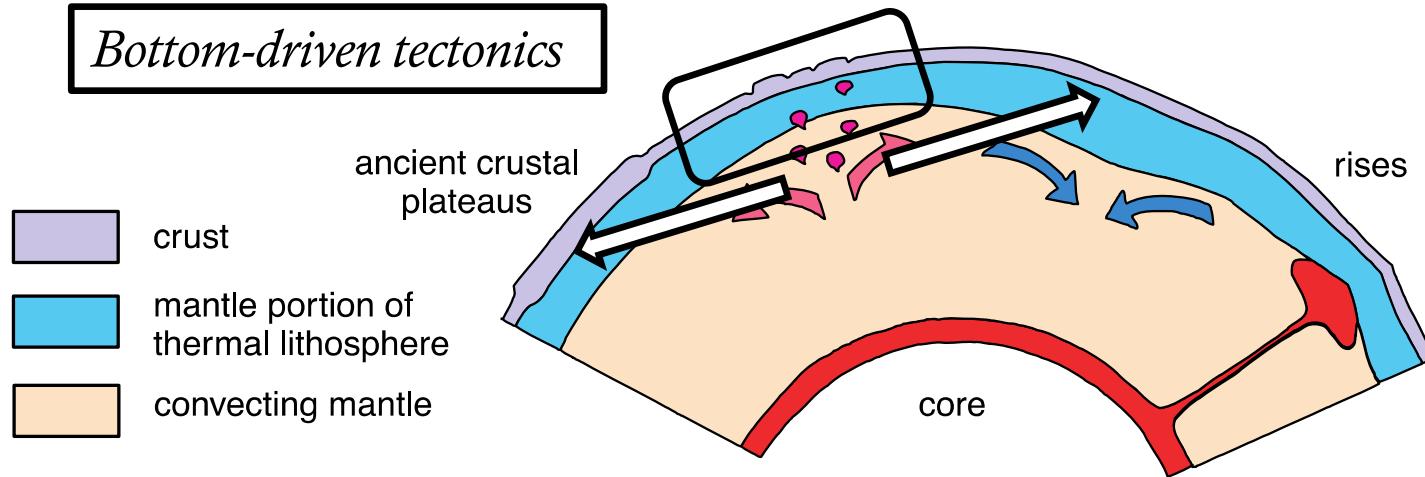
- Mantle:
  - Dis-GBS if low enough heat flux/ thin crust
    - 30km crust (lowlands):  $100 \text{ mW/m}^2$
    - 60km crust (highlands):  $50 \text{ mW/m}^2$
- Crust:
  - Water likely: brittle weakening
  - Mafic composition does not favor phyllosilicates at depth: no layering effect



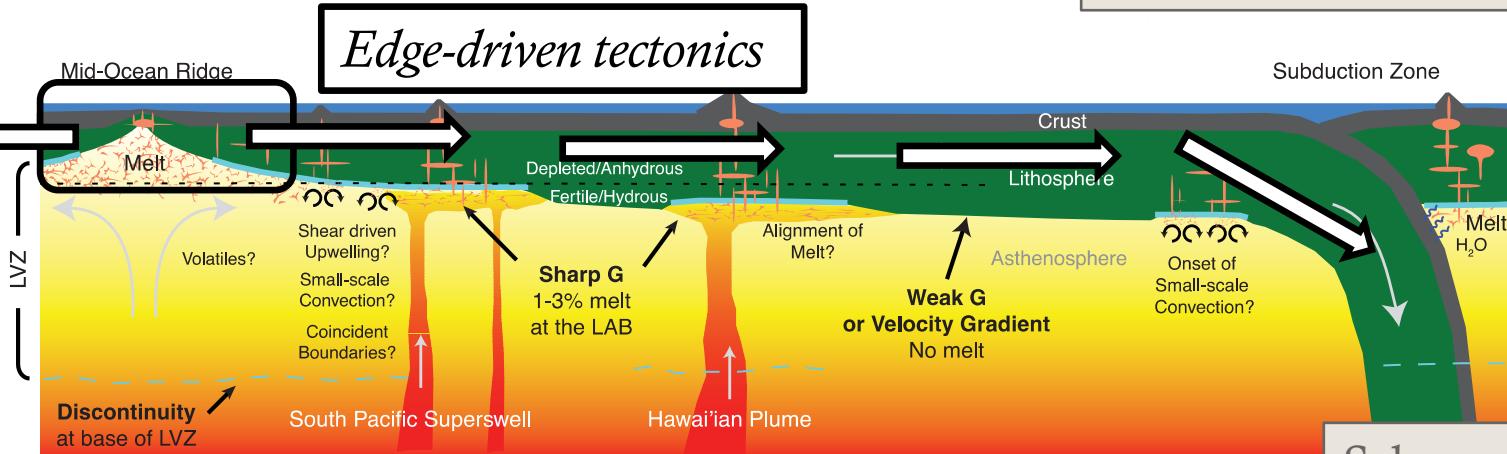
# Localization in planetary lithospheres



# Convection and tectonics

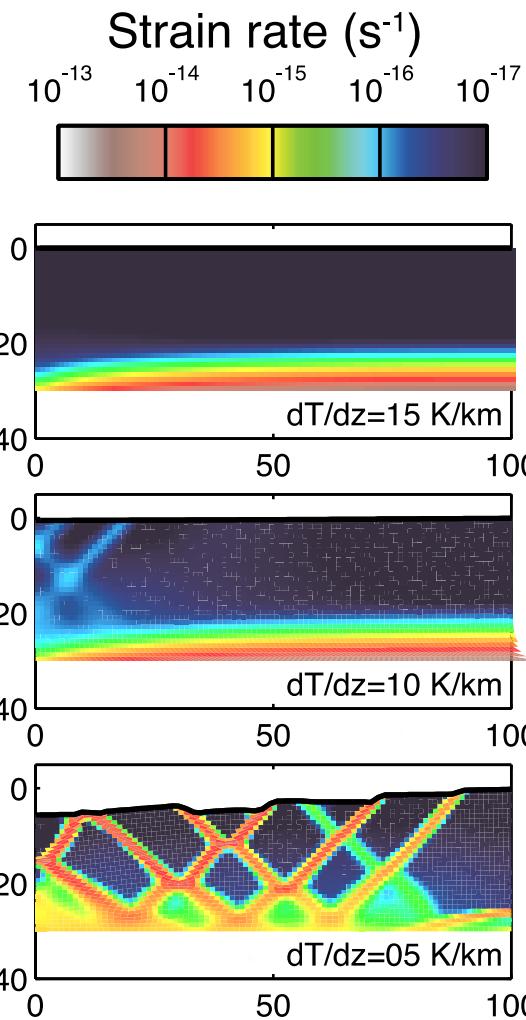


Phillips and Hansen, 1998

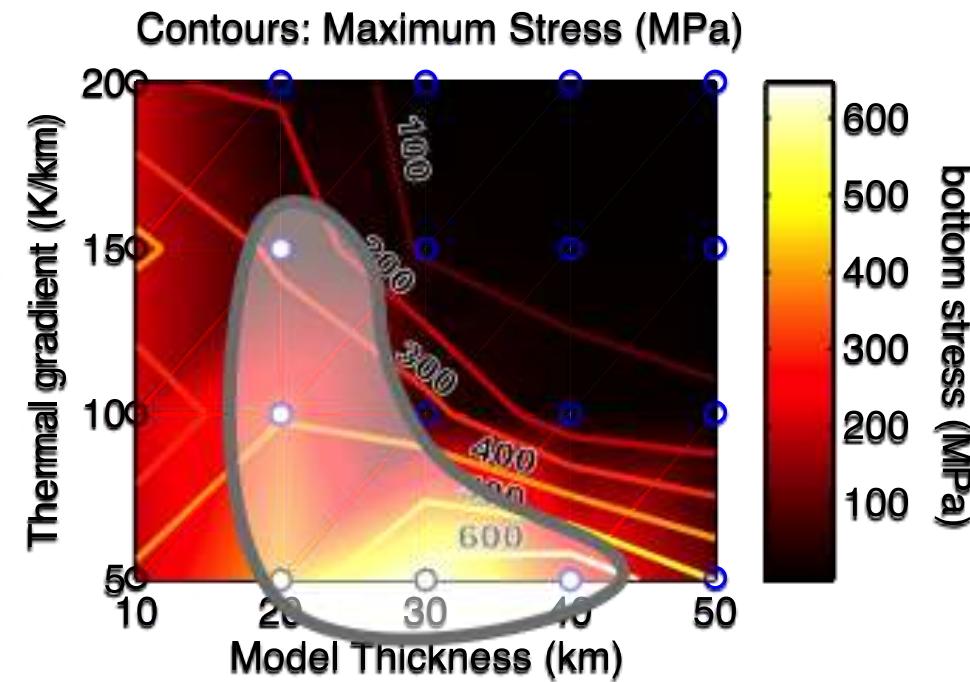


Schmerr, 2012

# Bottom-driven tectonics

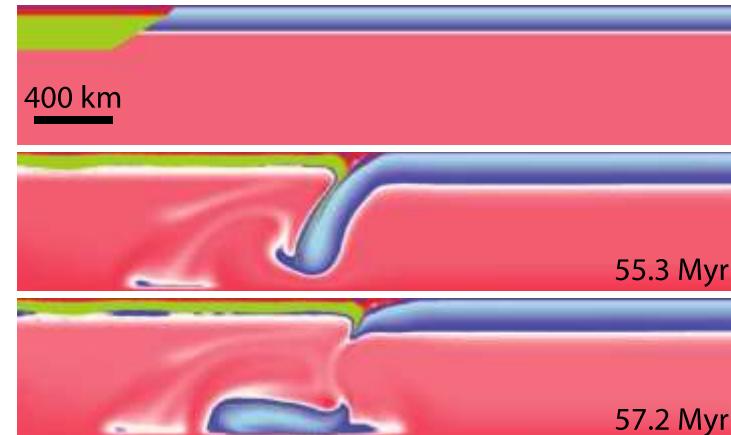


- Need high basal stress: convection reaches the brittle-ductile transition
- Need ductile layer for accommodation of shear zones



# Buoyancy-driven tectonics

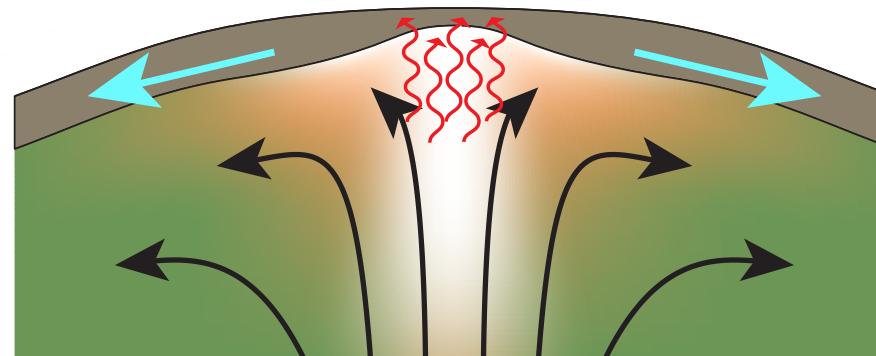
- Continent spreading (Earth's Archean)



Rey et al., 2014

Less efficient on Mars

- Upwellings (Beta Regio, East African Rift?)



# Summary

- The strength of the lithosphere changes with strain
  - Layering in brittle regime or in presence of micas
  - Grain size reduction in presence of dis-GBS creep
- compared to the Earth
  - Mars is similar to Earth's continents
  - Venus is dominated by brittle localization unless low heat flux
  - Issue of driving forces
- inefficient to drive tectonics
  - Continents dominated by buoyancy-driven flows?
  - Buoyancy less efficient on Mars

