

# Chasing the lithosphere-asthenosphere transition in dynamical models of the Earth's mantle

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Workshop “**Global Scale Seismic Imaging and Dynamics of the Earth's Mantle**”

Collège de France

7-8 Oct 2021



COLLÈGE  
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— 1530 —





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DE FRANCE  
1530

## CHAIRE DE PHYSIQUE DE L'INTÉRIEUR DE LA TERRE

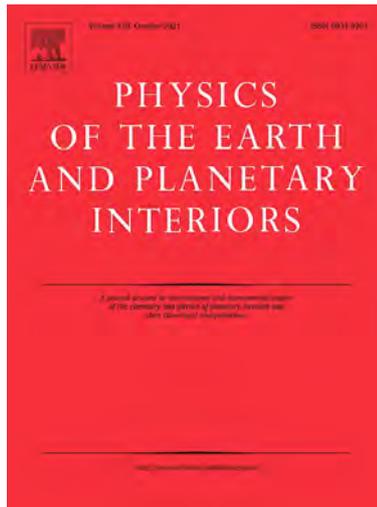
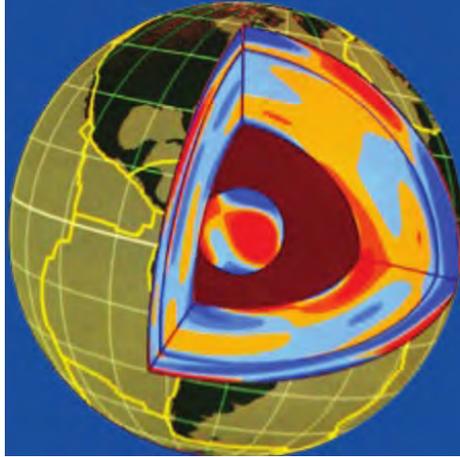
Année académique 2013-2014

Pr Barbara ROMANOWICZ

# Structure and Dynamics of the Lithosphere/Asthenosphere System

**Colloque en anglais - Workshop in English**

Mardi 19 et mercredi 20 novembre 2013.  
Amphithéâtre Maurice Halbwachs.



PEPI special issue 20-21

## **Physical properties and observations of the lithosphere-asthenosphere system**

Editors : Rick Aster, Saskia Goes, Derek Schutt

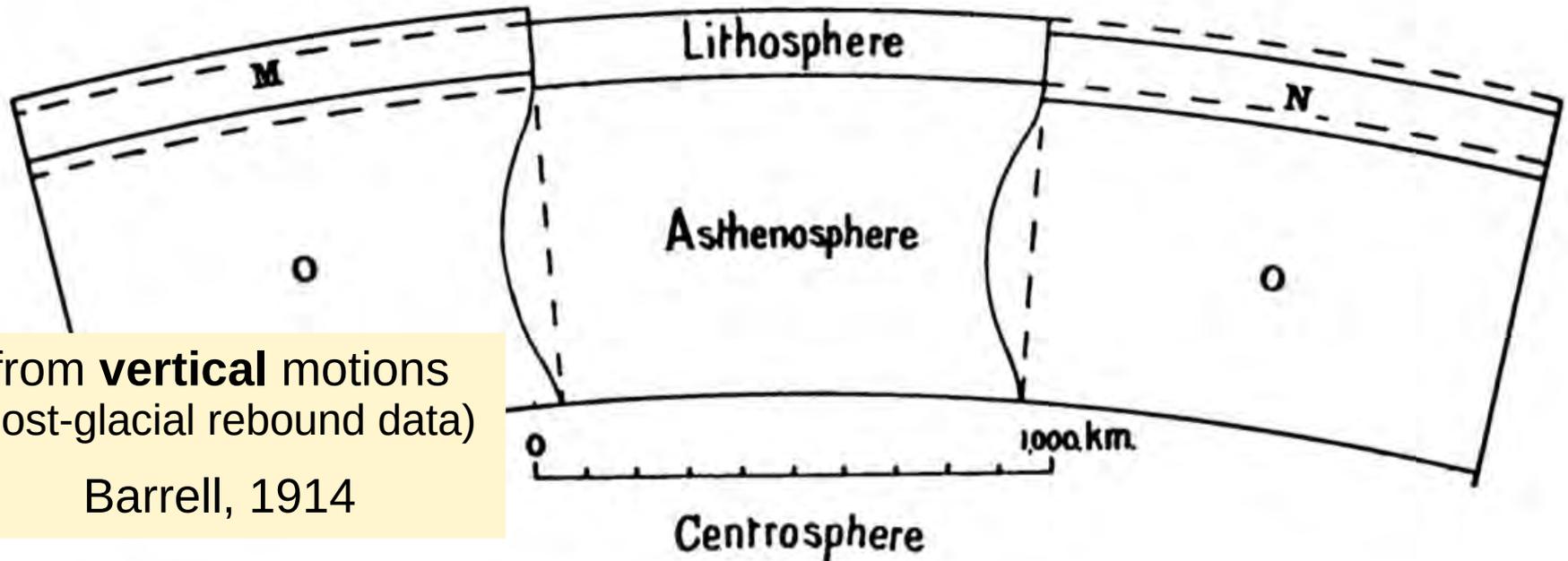
# Strong “lithosphere” vs. soft “asthenosphere”

## LITHOSPHERE

**rigid layer**  
(crust + uppermost mantle)

## ASTHENOSPHERE

relatively “softer”  
mantle layer

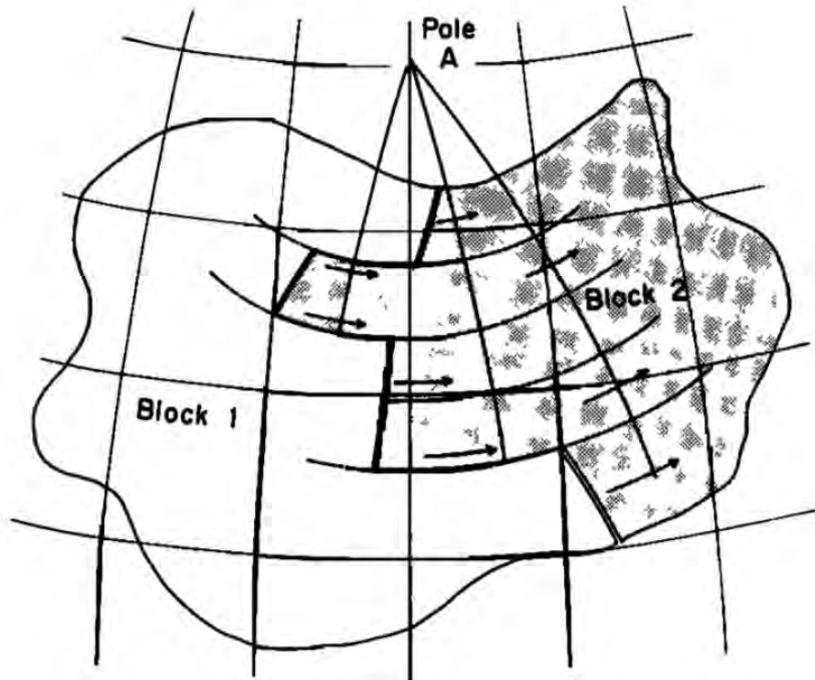


from **vertical** motions  
(post-glacial rebound data)

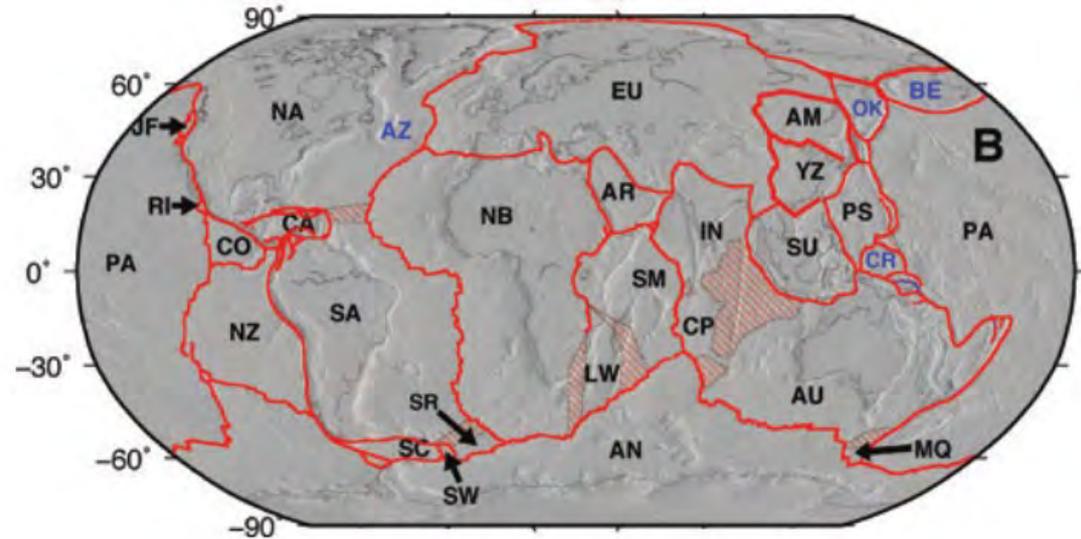
Barrell, 1914

# Plate tectonics' theory : lithosphere $\equiv$ dynamic plate

rigid blocks moving horizontally (2-D) at the surface of the Earth's sphere  
(no reference to asthenosphere)

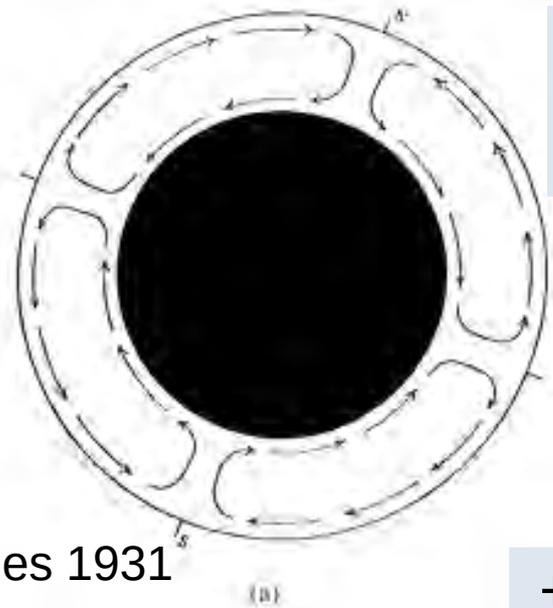


Morgan, 1968



25 plates' model MORVEL (DeMets et al., 2010)

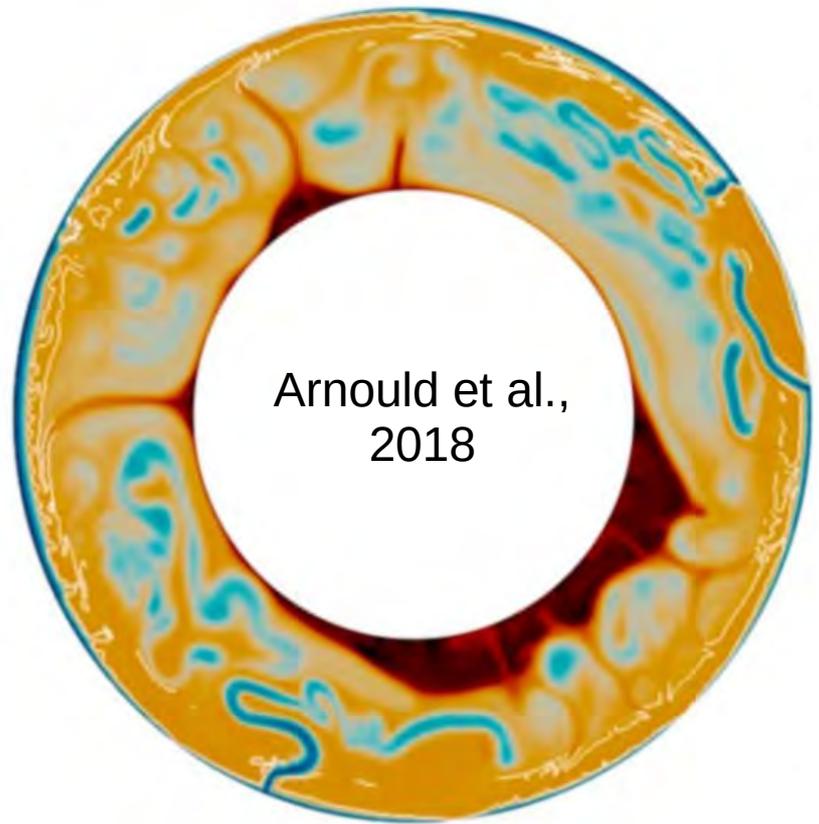
# Cold lithosphere vs. hot asthenosphere



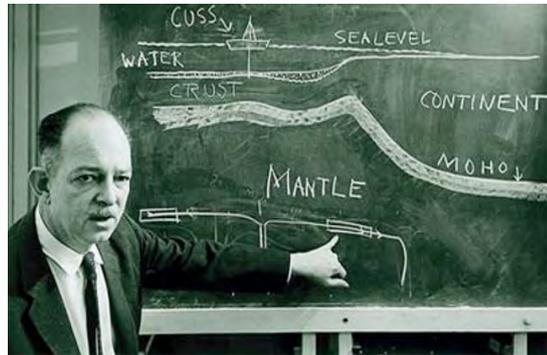
Lithosphere as the conducting and moving cold boundary layer of **convective mantle**

Holmes 1931

+ lithospheric plates sinking in the mantle

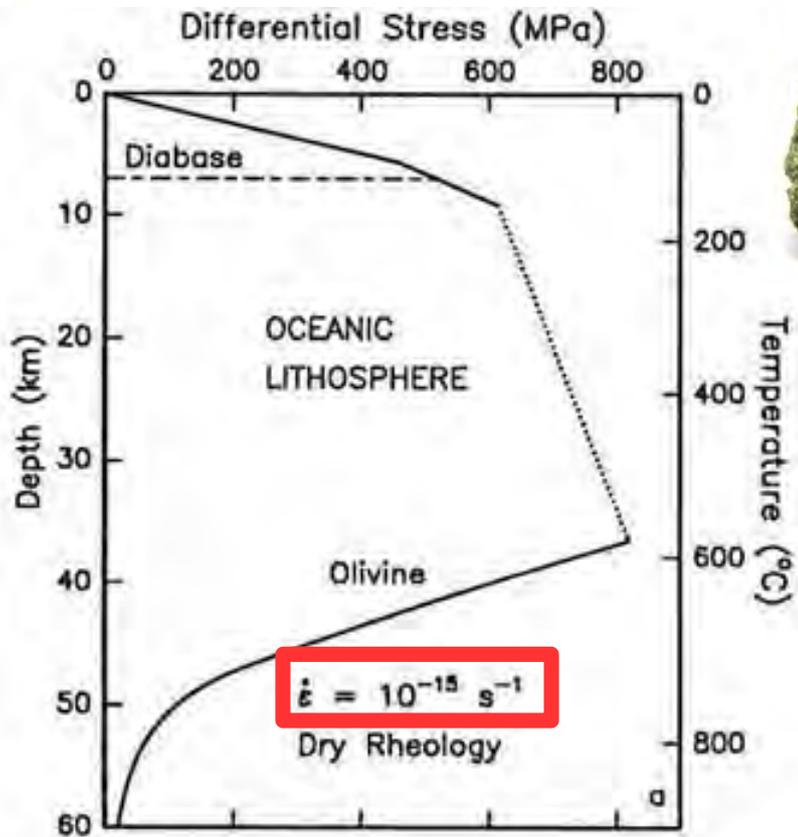


Arnould et al.,  
2018

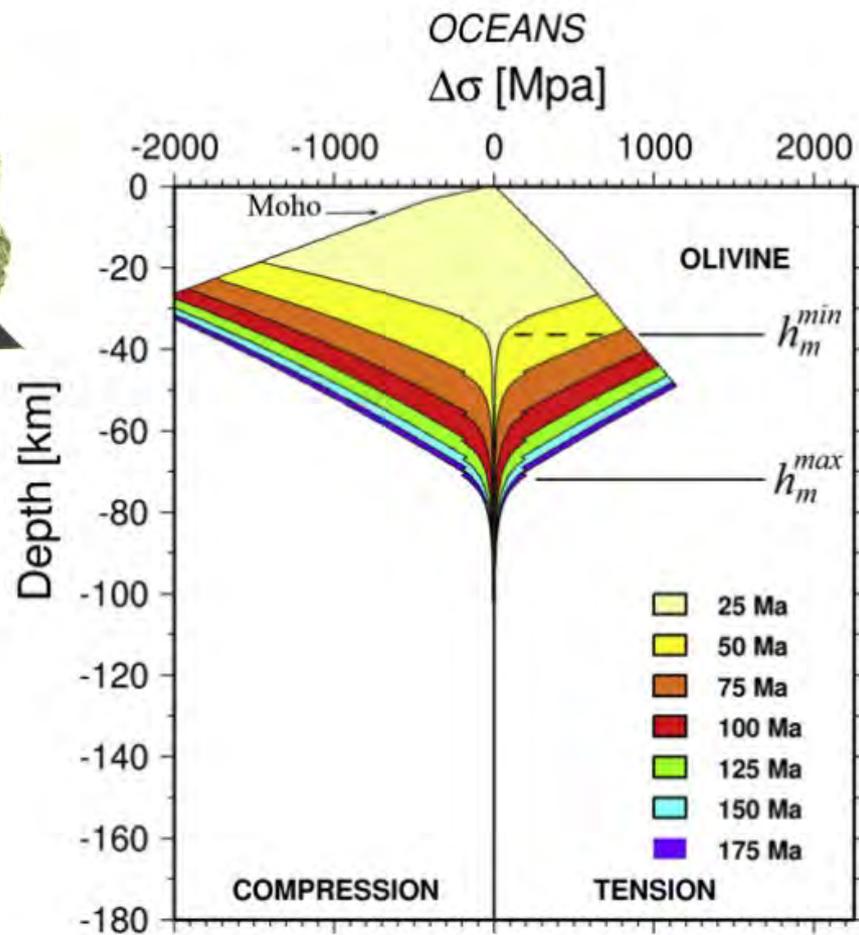


Hess, 1962

# Temperature-dependence of strength and viscosity : same peridotite either cold/strong or hot/weak ?

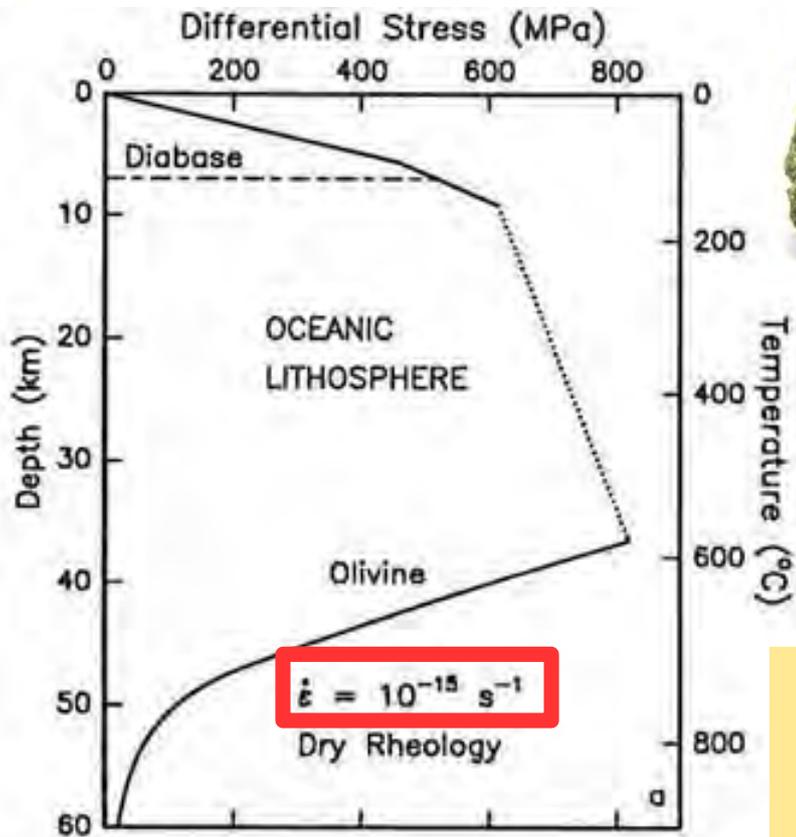


Kohlstedt et al., 1995



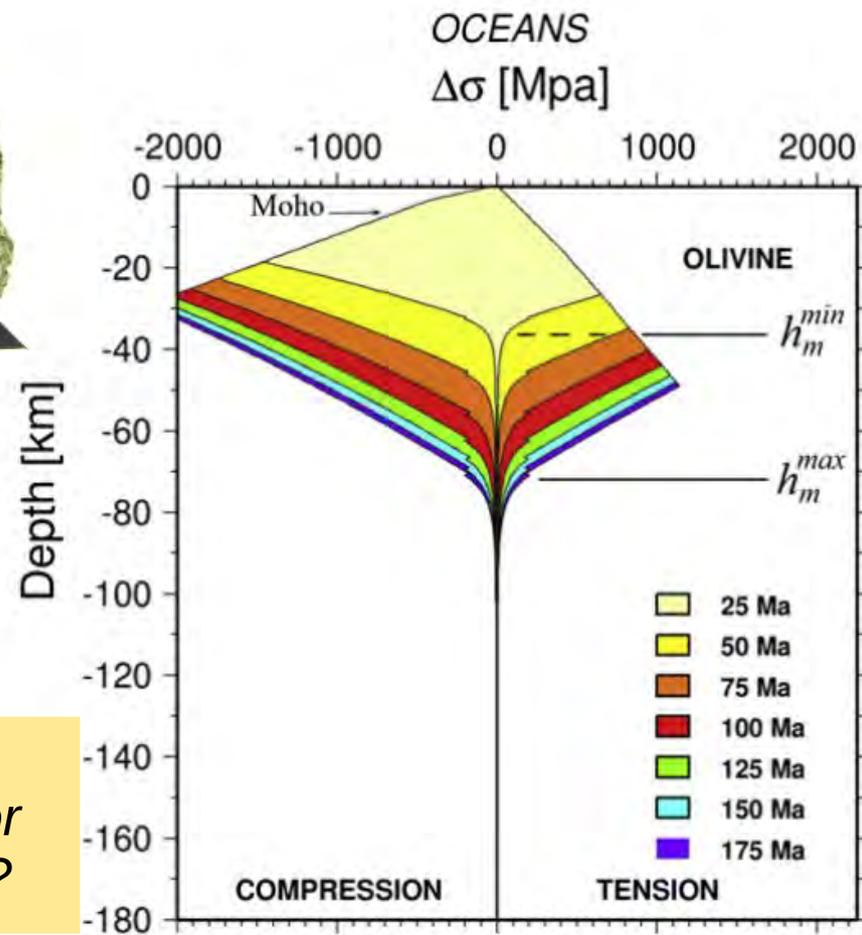
Burov, 2011

# Temperature-dependence of strength and viscosity : same peridotite either cold/strong or hot/weak ?



Kohlstedt et al., 1995

*lith-asth  
transition or  
boundary ?*



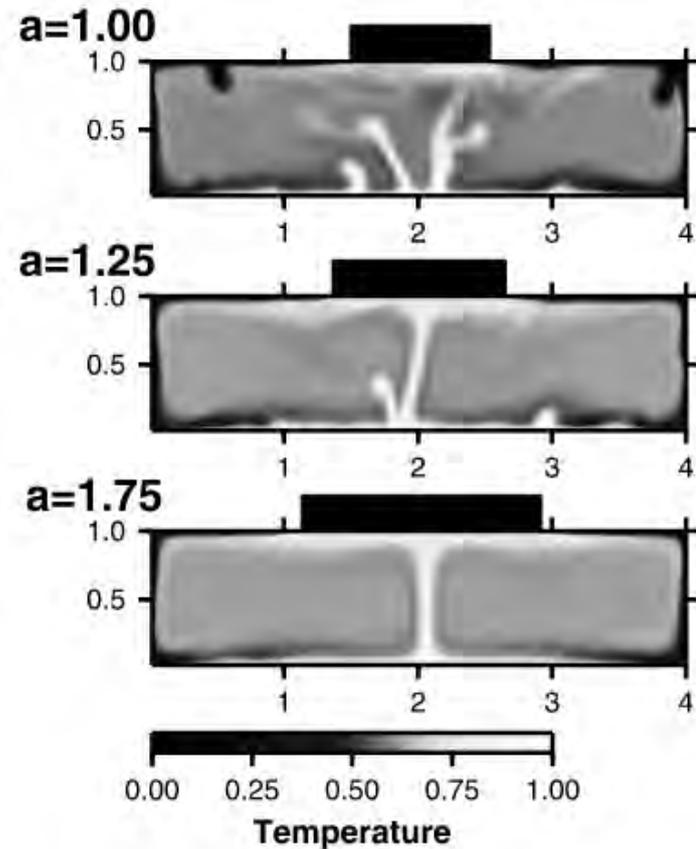
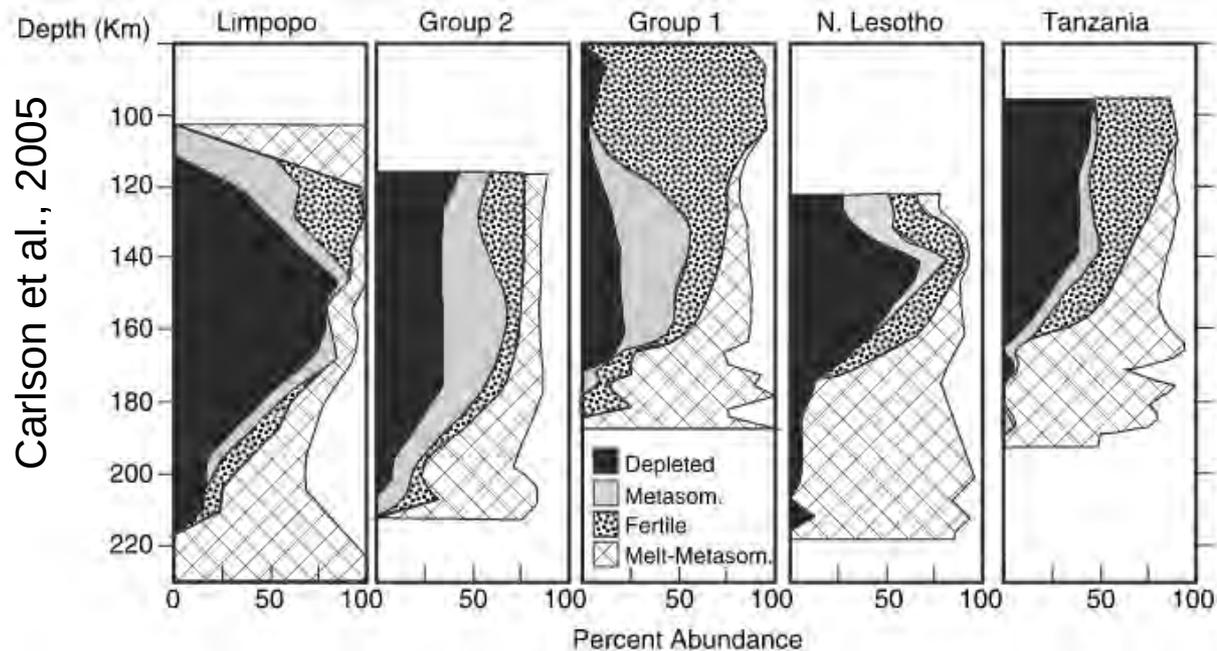
Burov, 2011

# Lith.-asth. compositional boundary below continents ?

Continental lithosphere melt-depleted / dehydrated ?

→ effect on viscosity and conductivity ?

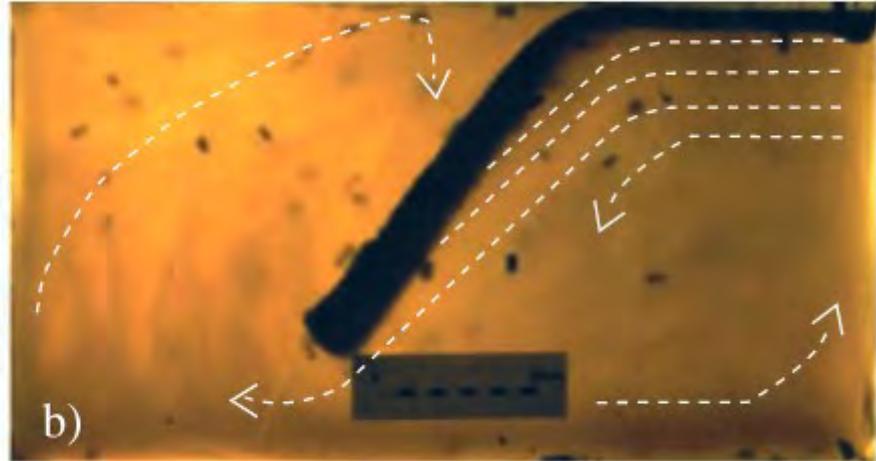
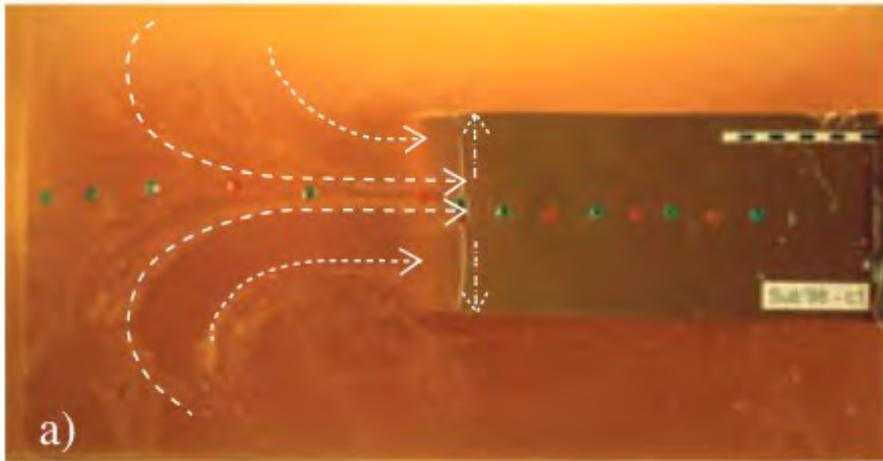
(*implications for mantle flow and heat transfer*)



Mantle types in the African lithosphere interpreted from xenoliths

Grigné et al.,  
2007

# Dynamic models with a lithosphere-asthenosphere boundary



Funiciello et al., 2003

## Analogue experiments

- lithosphere  $\equiv$  silicone putty
- asthenosphere = glucose syrup

> prescribed viscosity contrast ( $\geq 1000$ )

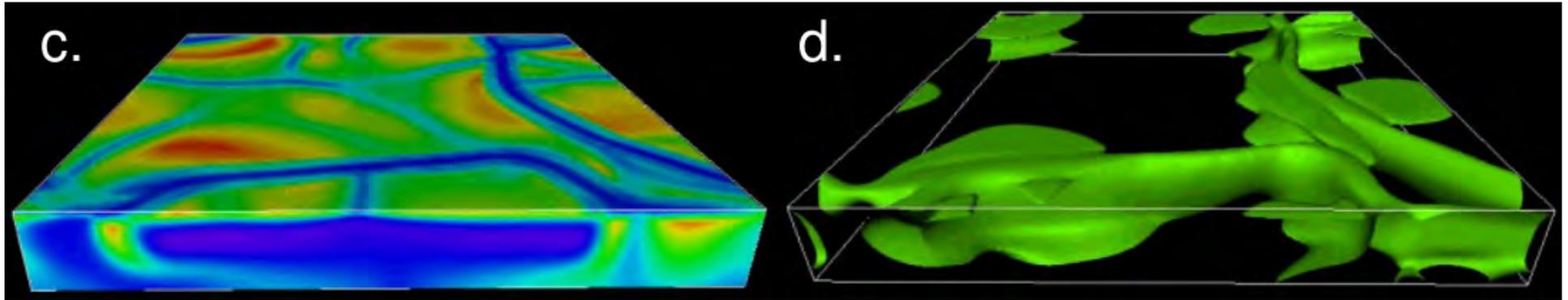
## “Compositional” numerical models



Capitanio et al., 2010

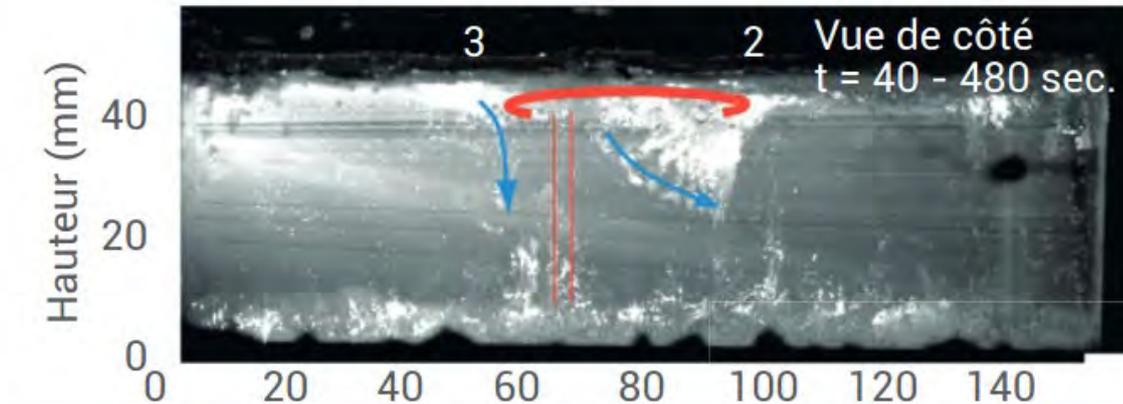
# Dynamic models with a lithosphere-asthenosphere transition

Tackley, G3, 2000



“thermo-mechanical” numerical models

> temperature-dependent viscosity



**Analogue experiments**

*Ludox* (silica particle suspension)

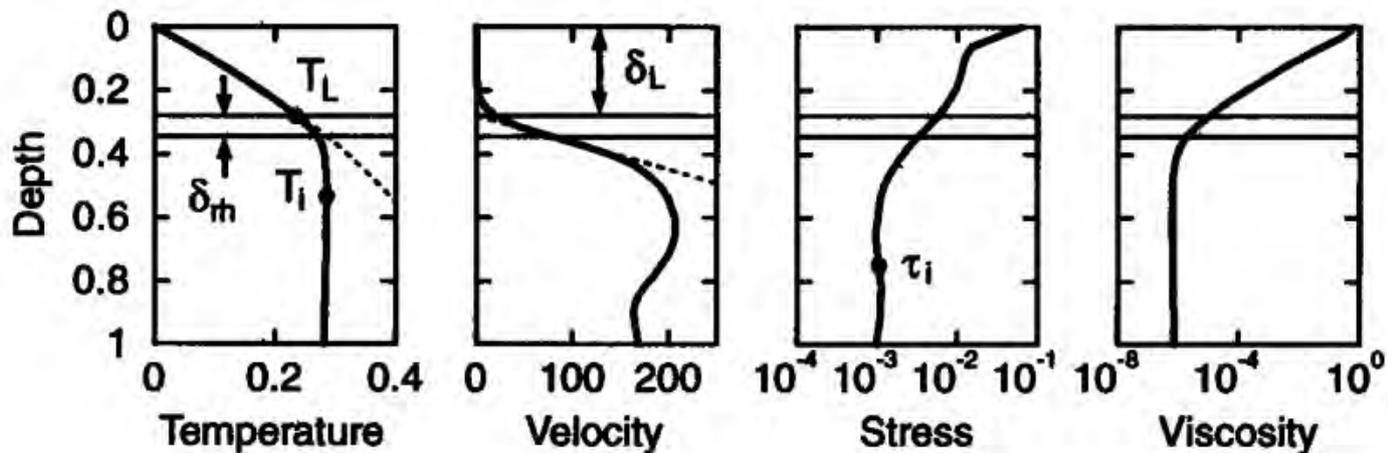
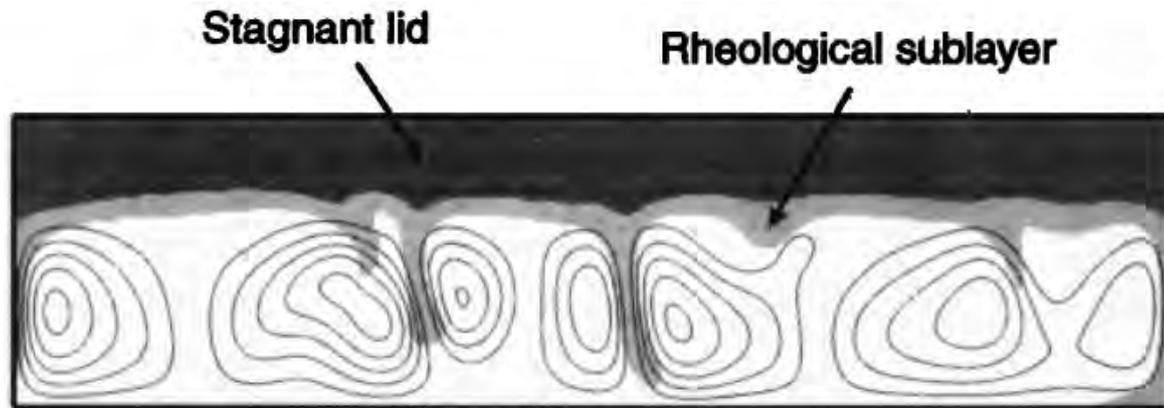
> rheology =  $f(\text{water content})$

- dry “lithosphere”
- wet “asthenosphere”

Davaille et al., 2020

# Thermal, mechanical, viscosity lith-asth transition... but what about the velocity transition ?

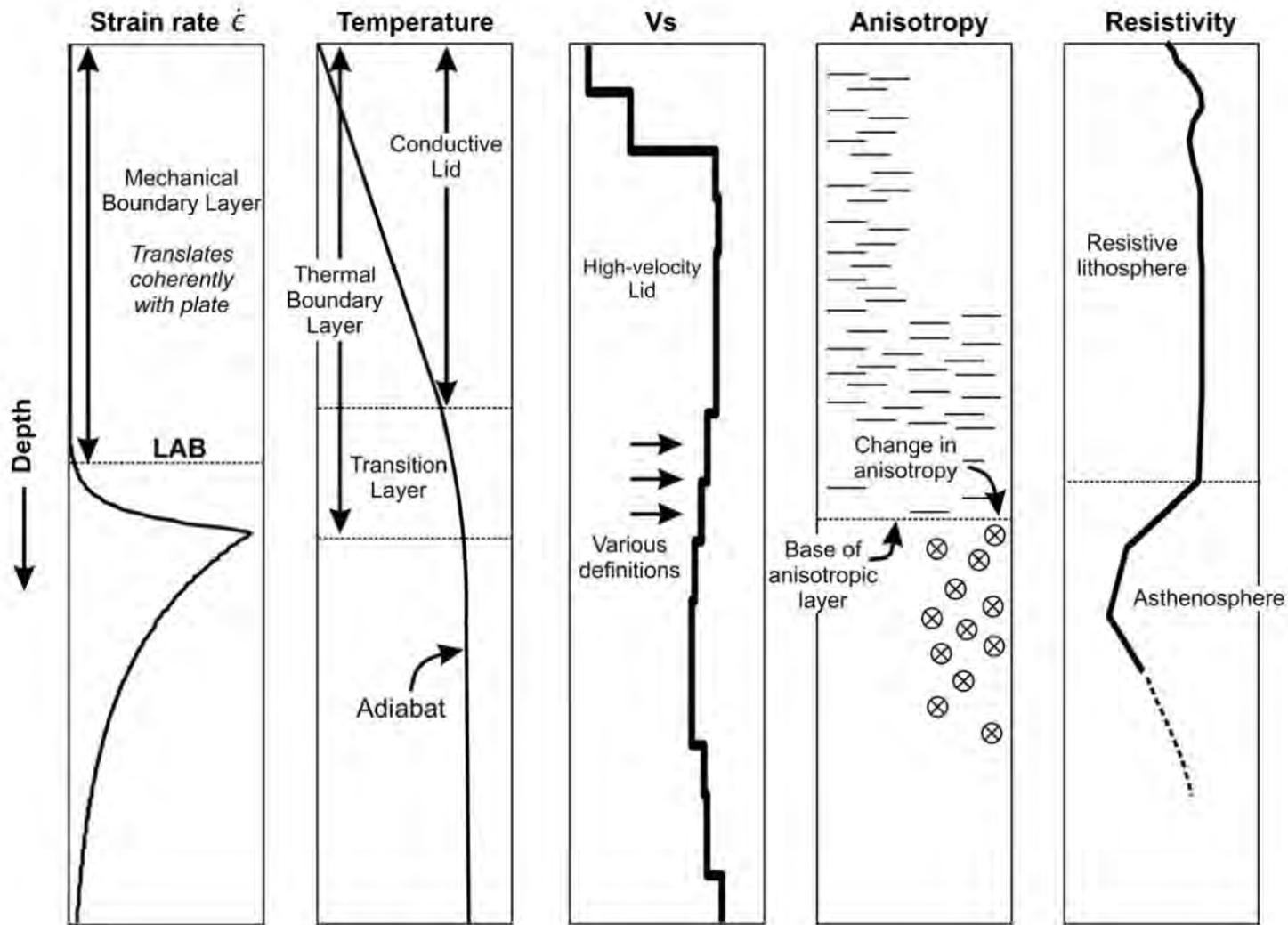
Lithosphere  
superimposed to  
moving tectonic plates  
→ **how does velocity  
vary vertically ?**



*Stagnant-lid  
simulation*

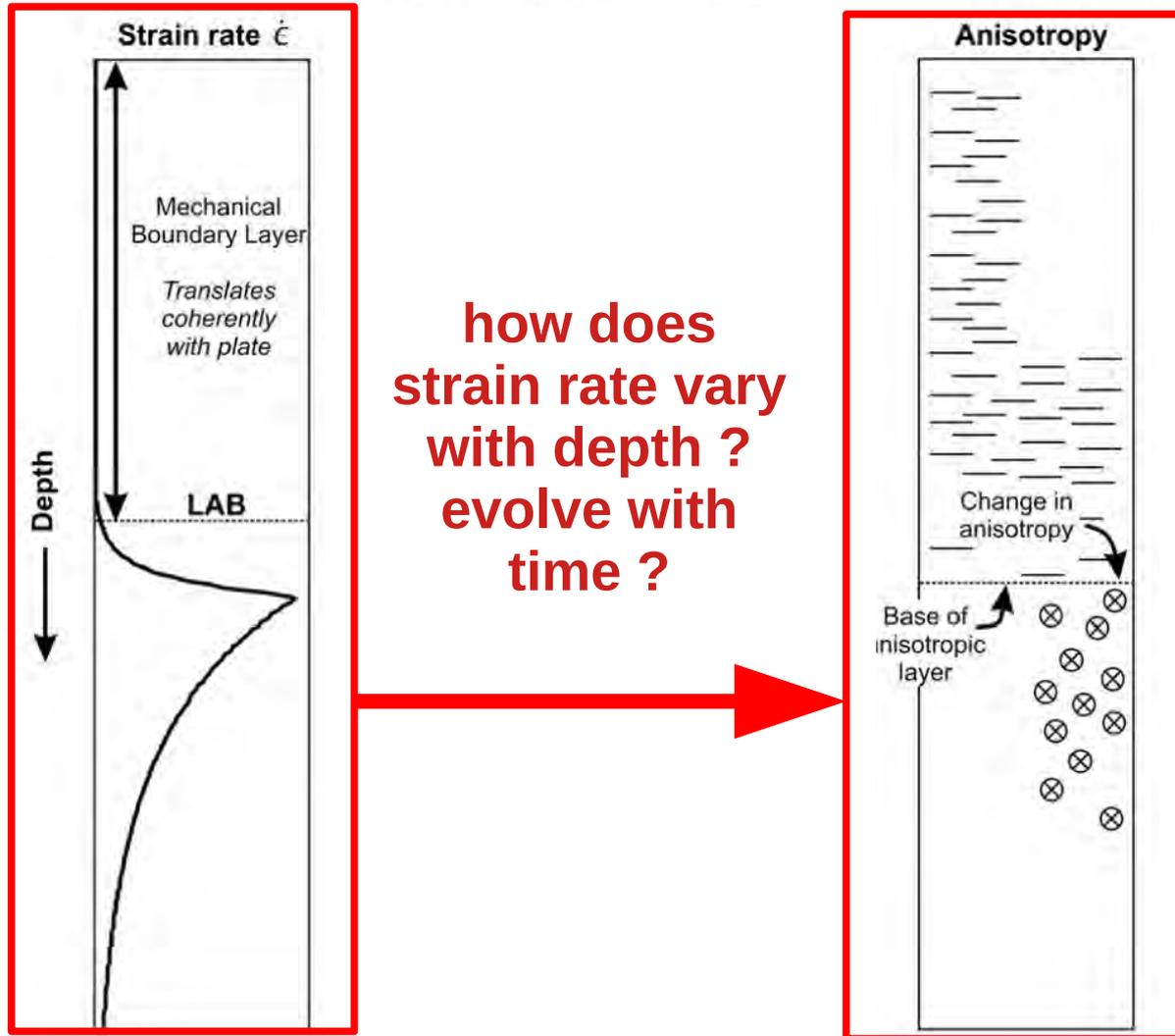
Solomatov & Moresi, 2000

# Multiple definitions and proxies of the LAB



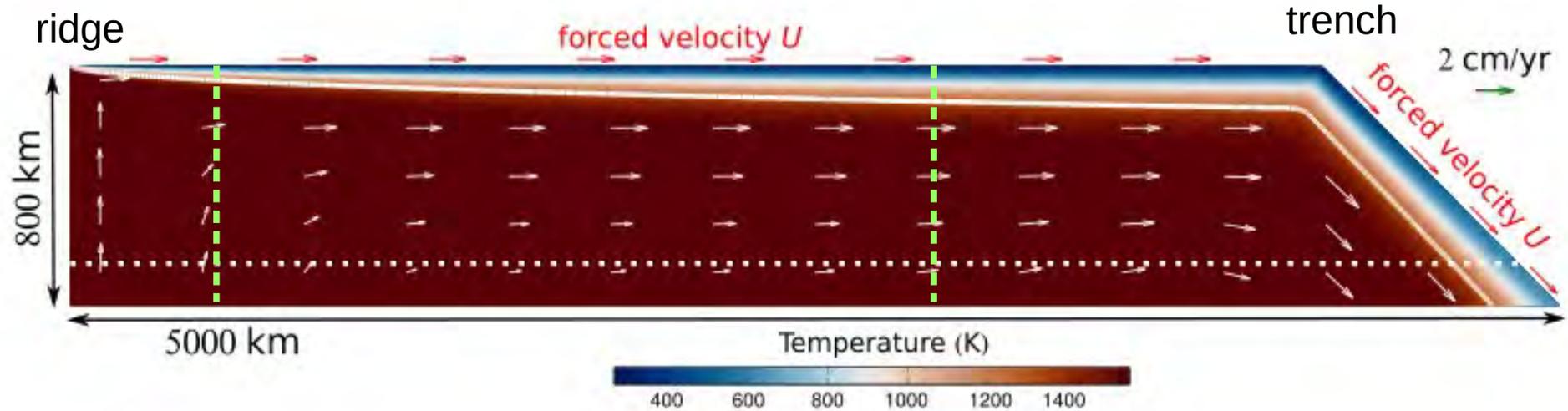
Eaton et al.,  
2009

# Seismic anisotropy as a proxy for the lith-asth transition ?



*Eaton et al.,  
2009*

# Asthenospheric flow driven by surface plate



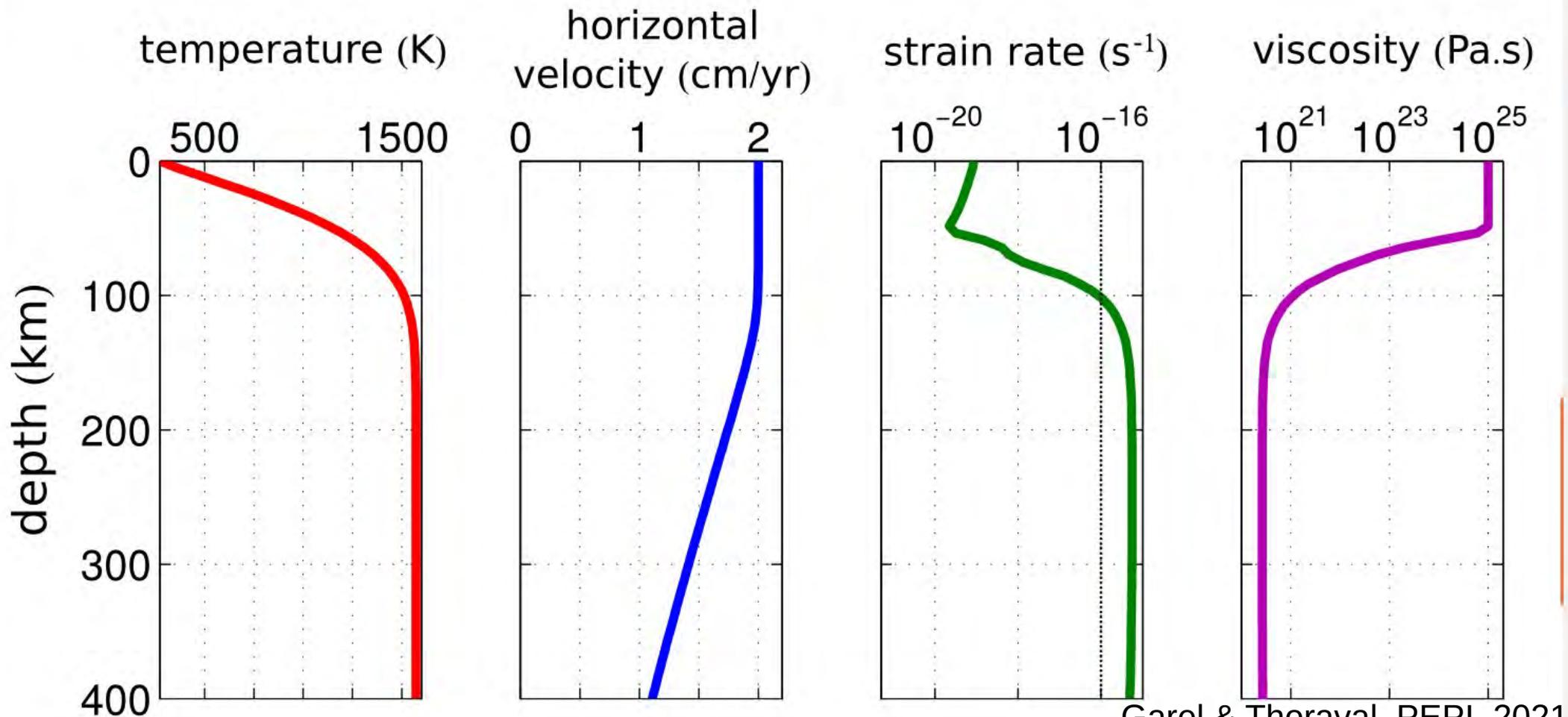
**Forced velocity  $U$  : 1-2-5-10-20 cm/yr**

**Vertical profiles at different ages**

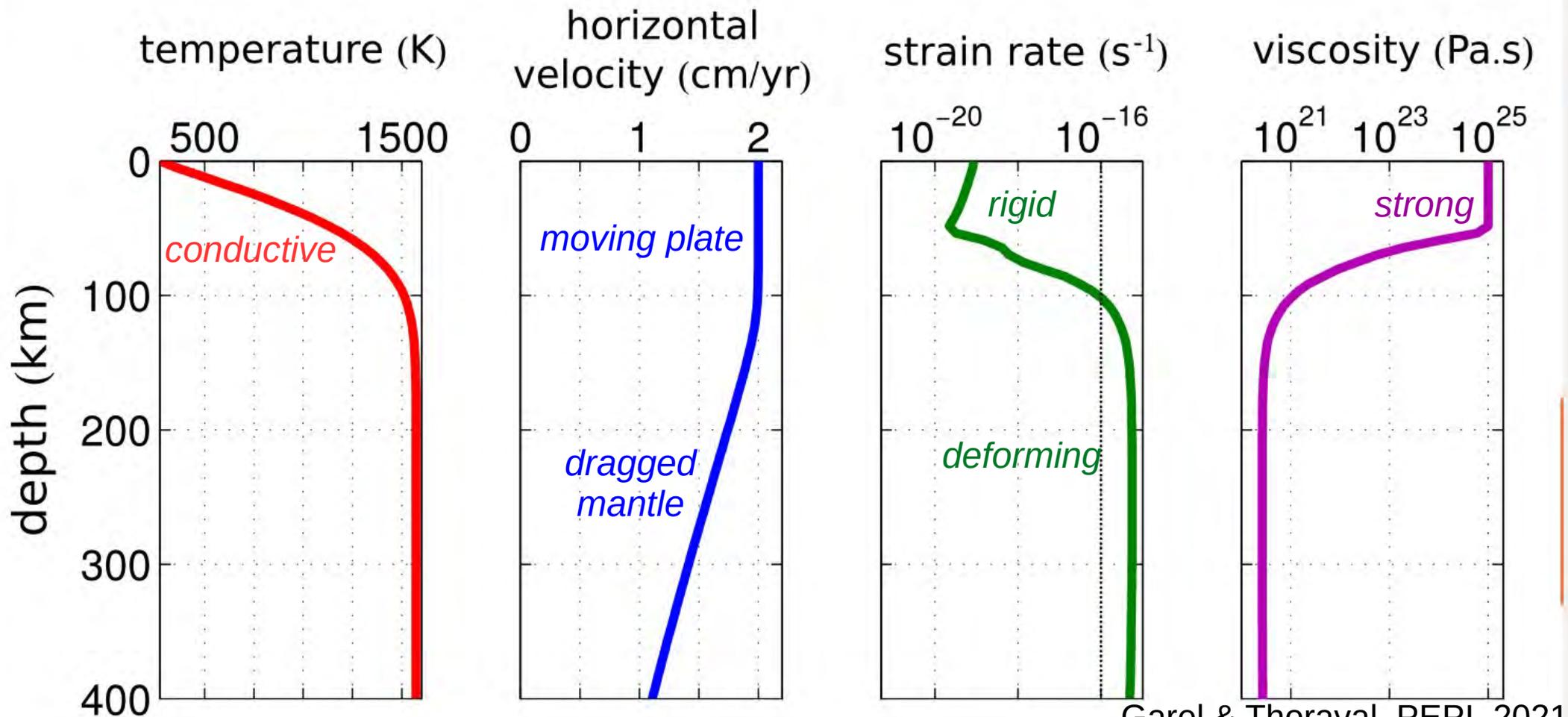
- **one single material for both lithosphere and asthenosphere**  
(no pre-imposed discontinuity)

- steady-state

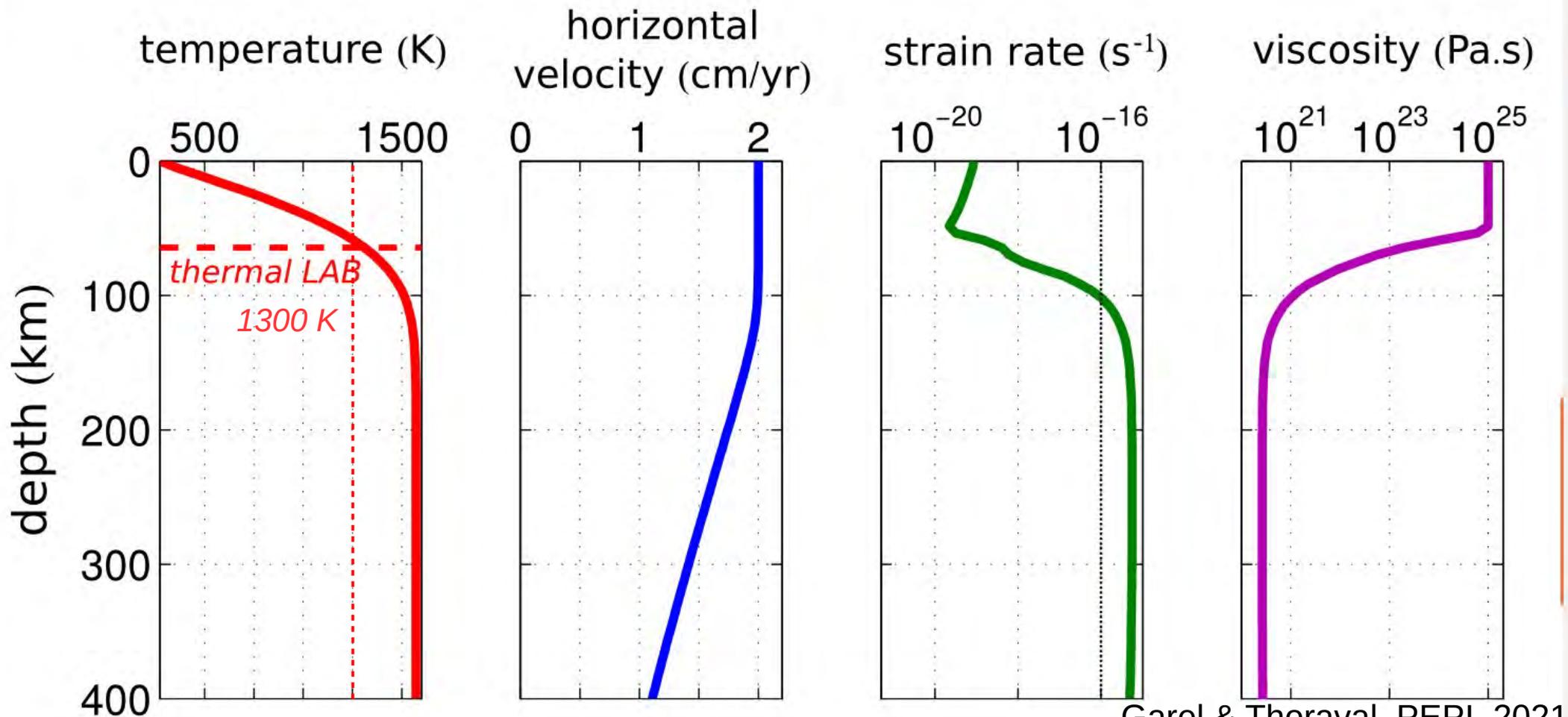
# Vertical profiles below a moving plate (2 cm/yr)



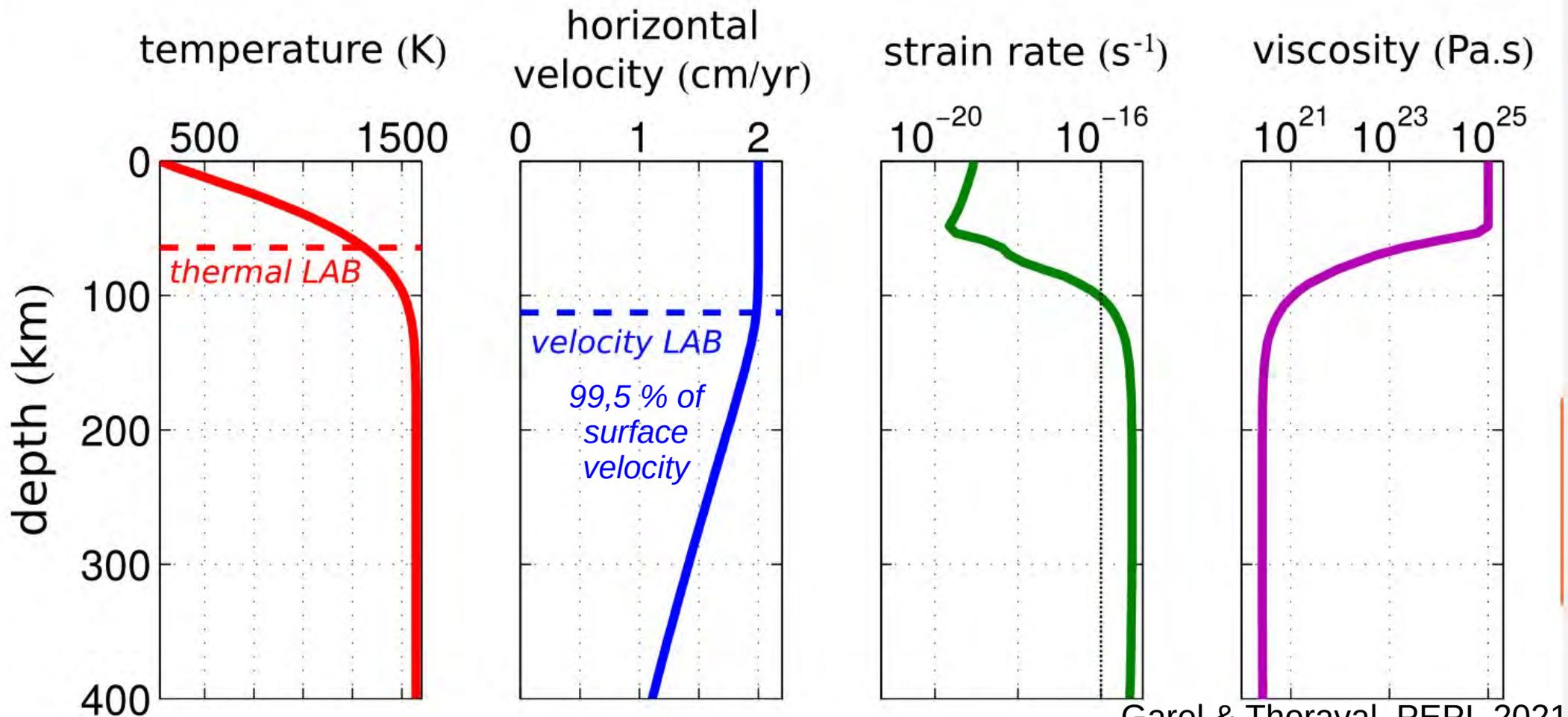
# Vertical profiles below a moving plate (2 cm/yr)



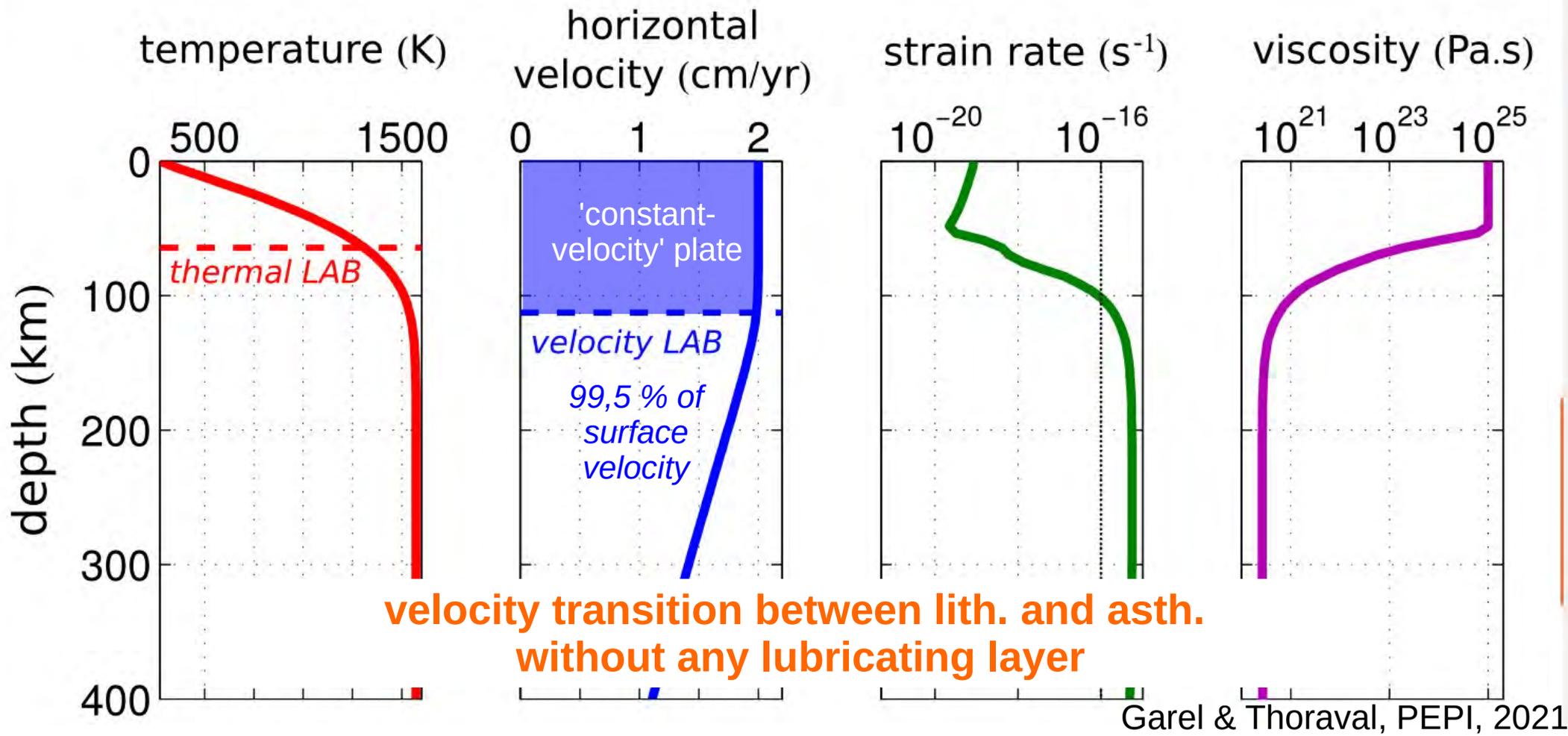
# Vertical profiles below a moving plate (2 cm/yr)



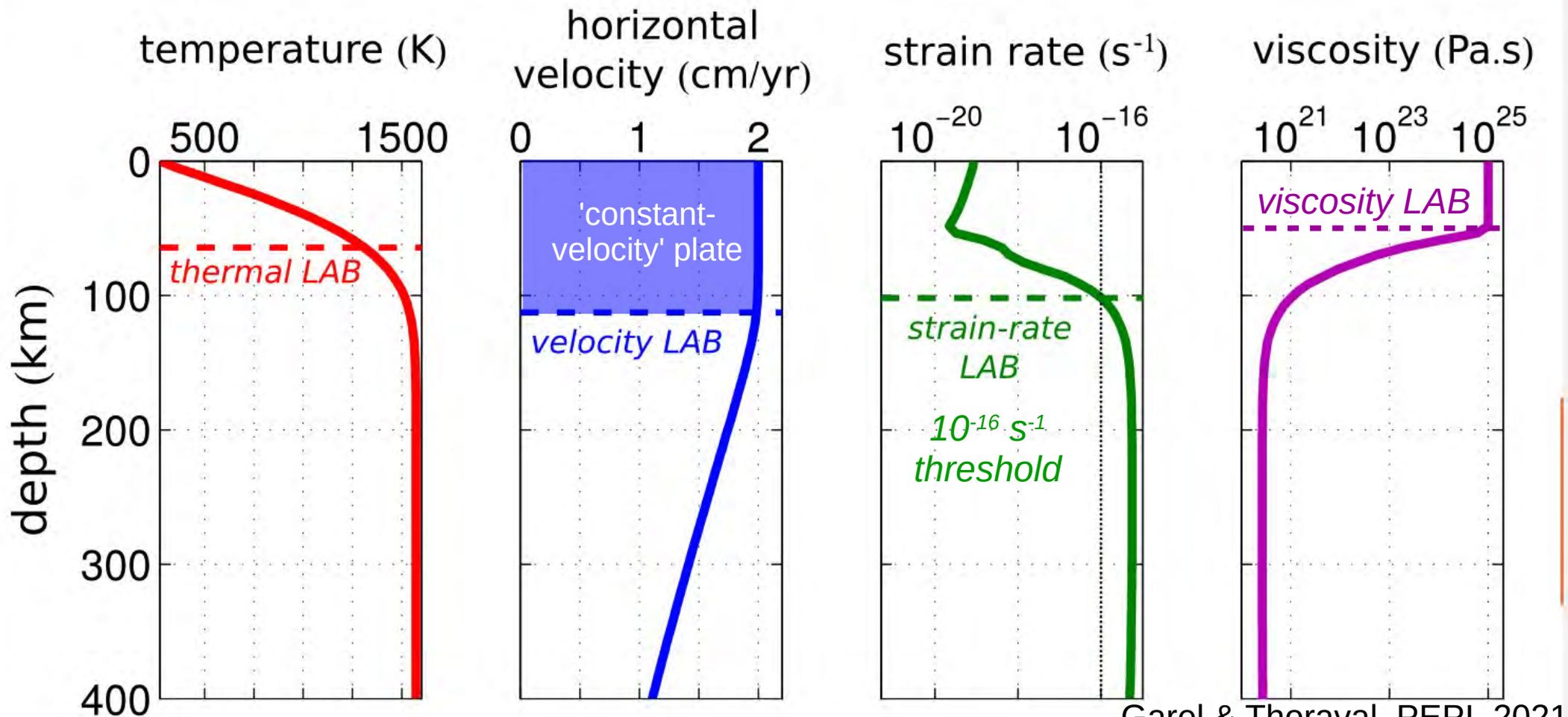
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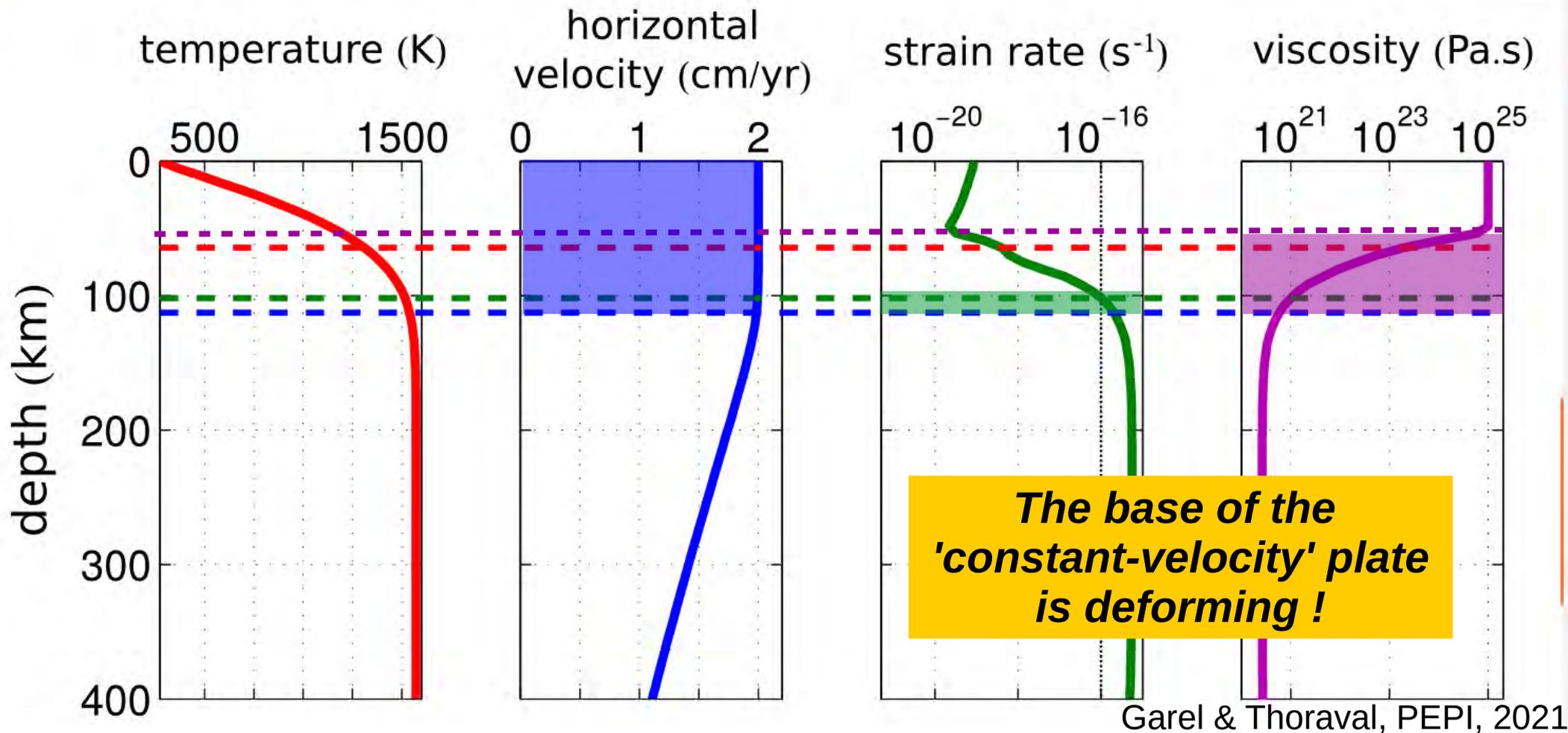
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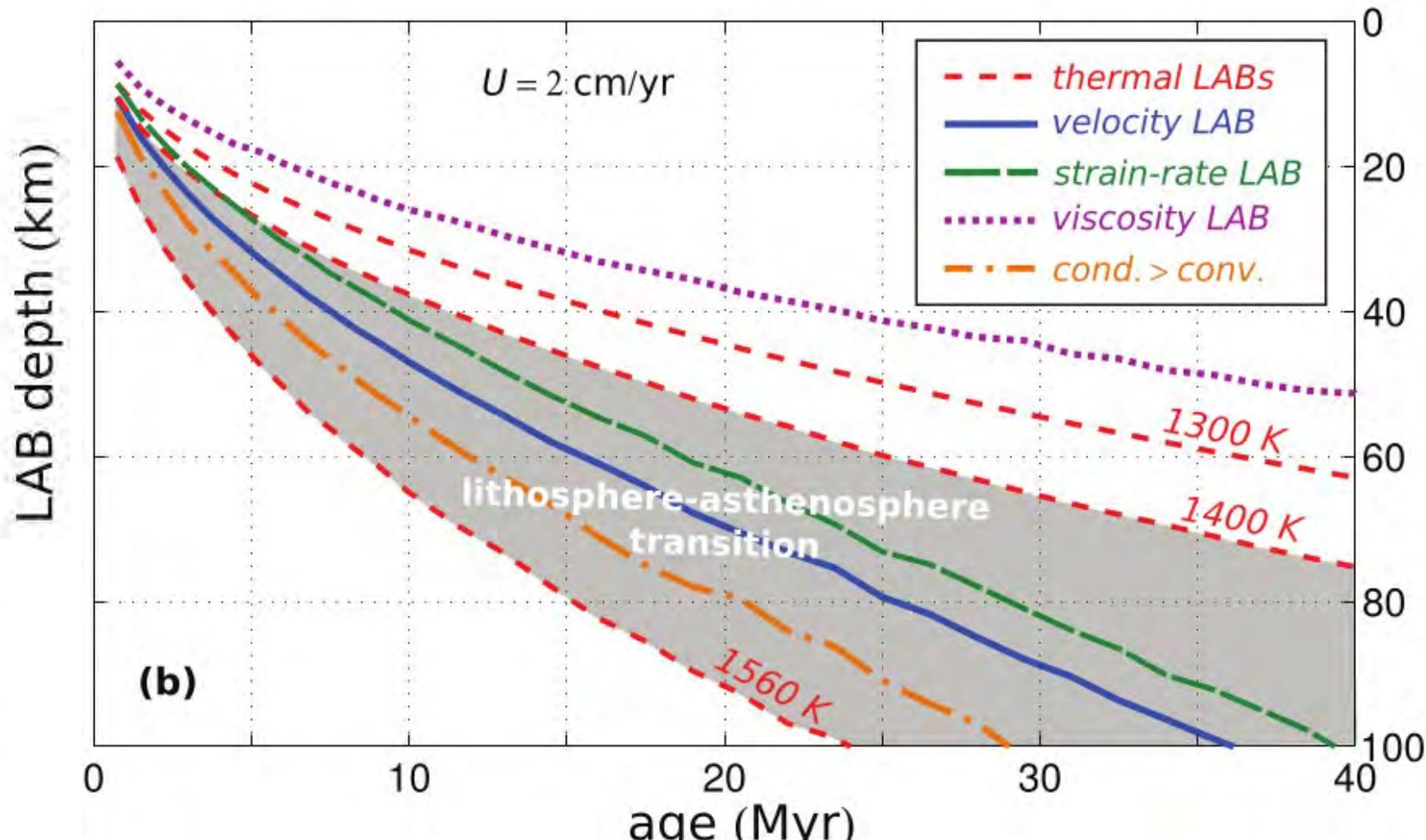
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# Vertical profiles below a moving plate (2 cm/yr)

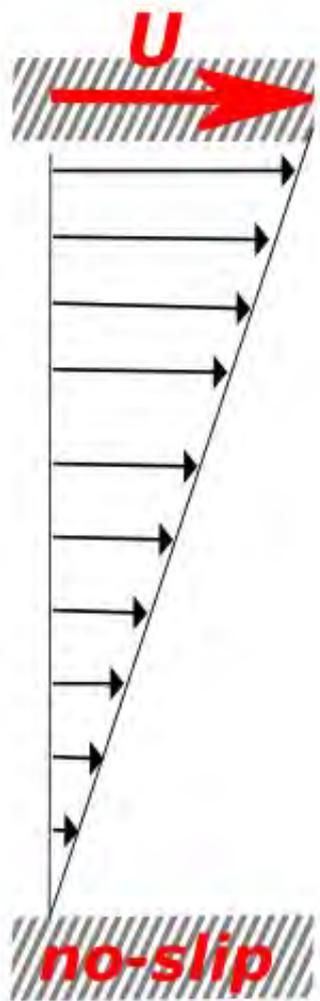


# Progressive transition from lithosphere to asthenosphere for a homogeneous material with temperature-dependent viscosity

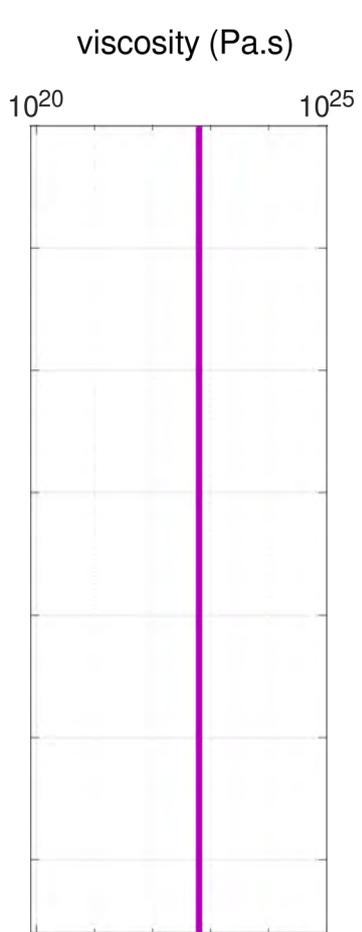
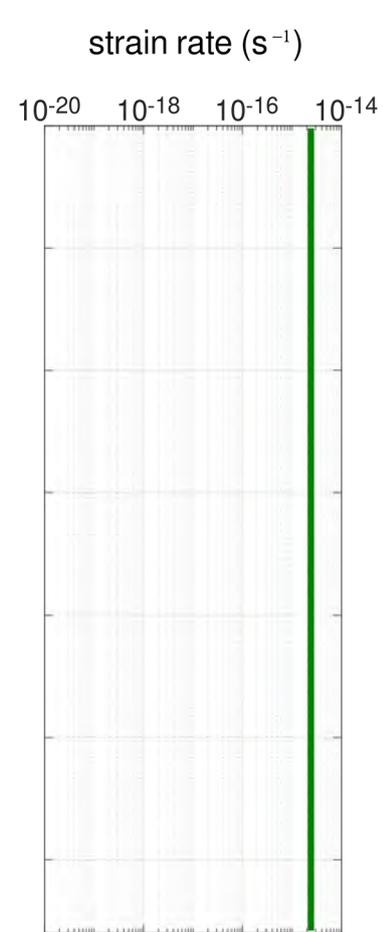
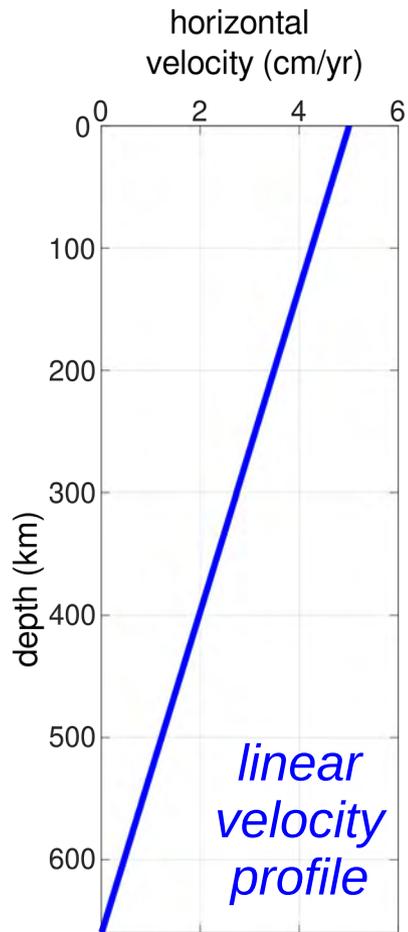


Garel & Thoraval, PEPI, 2021

# Explaining the 'constant-velocity' plate using a Couette flow in $n$ layers with variable viscosity



constant  
viscosity



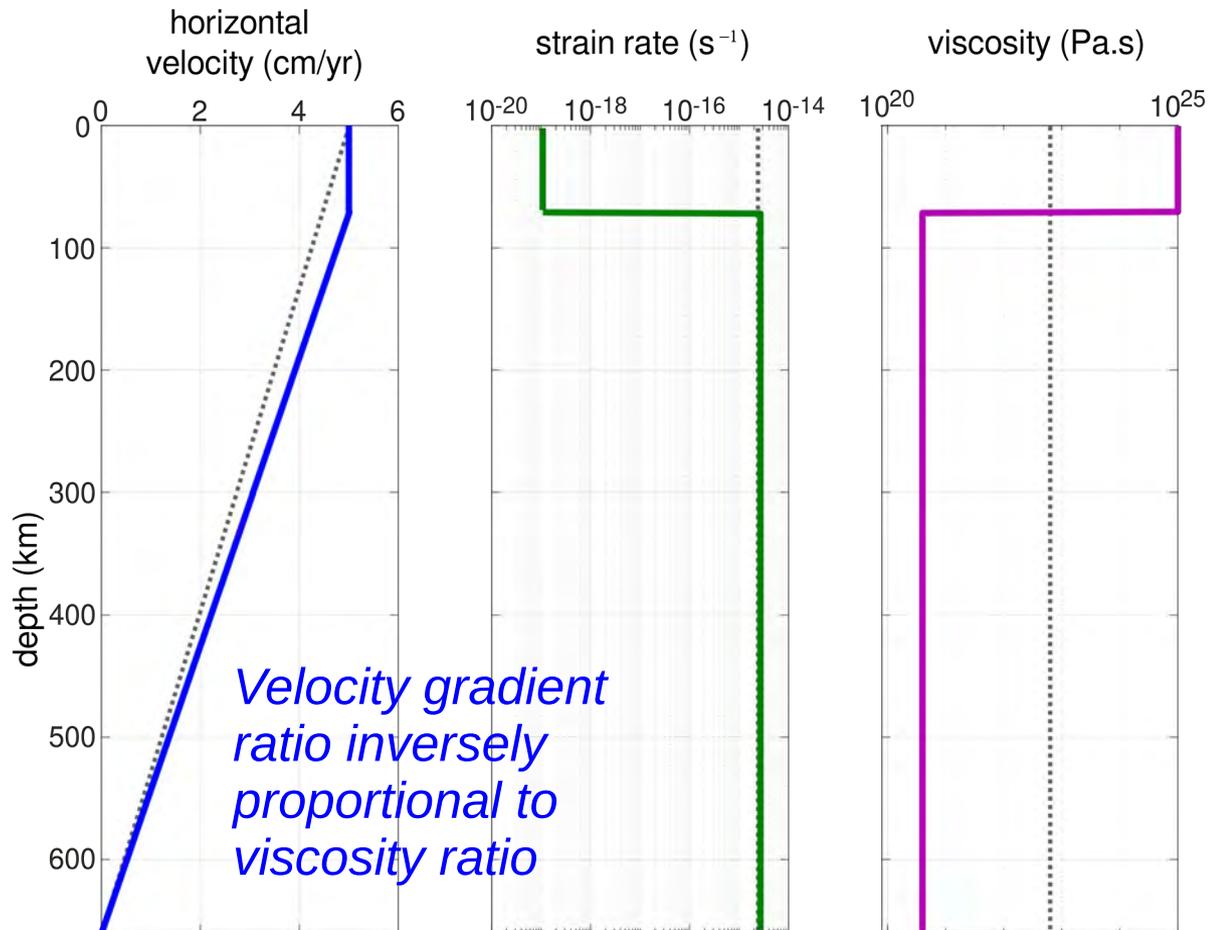
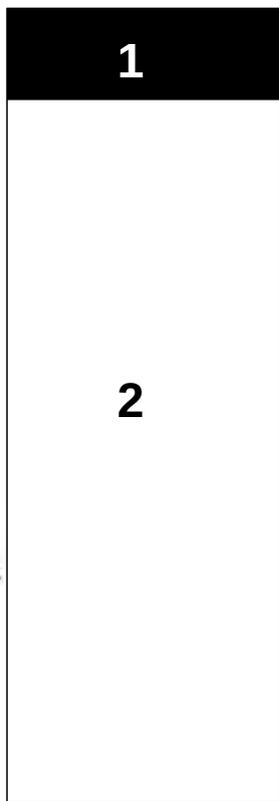
# Explaining the 'constant-velocity' plate using a Couette flow in $n$ layers with variable viscosity

2 layers with different viscosities

- continuity of viscous shear stress at the interface :

$$\mu_1 \frac{\partial U_1}{\partial z} \Big|_{z=h} = \mu_2 \frac{\partial U_2}{\partial z} \Big|_{z=h}$$

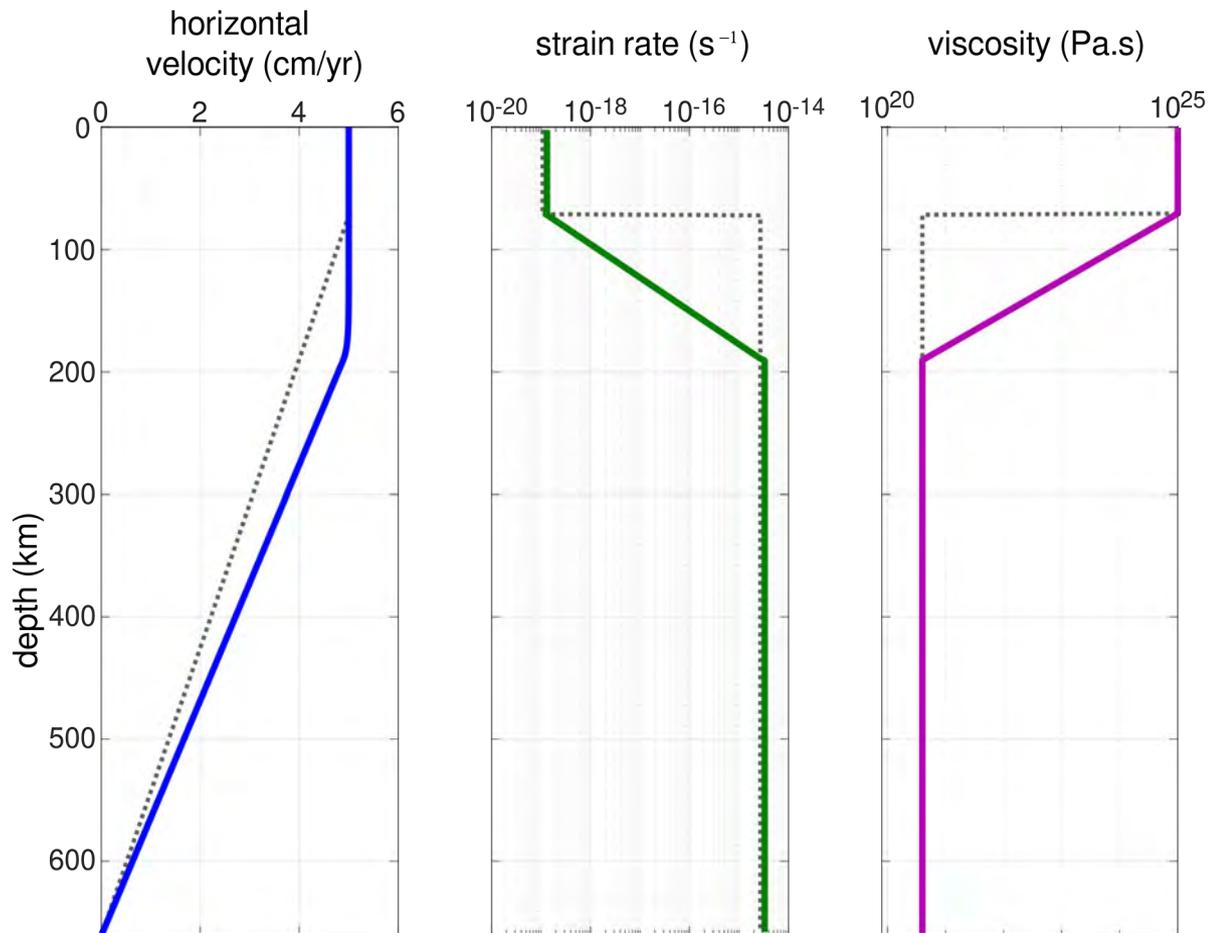
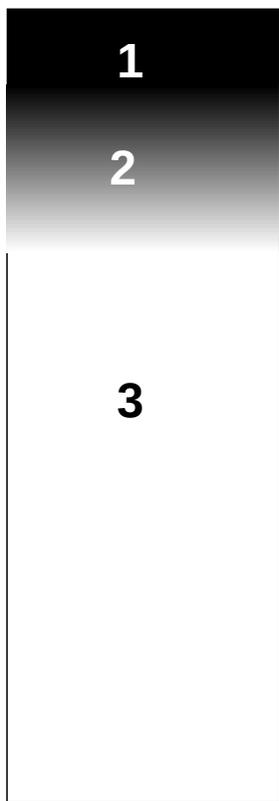
- BCs at surface and bottom



# Explaining the 'constant-velocity' plate using a Couette flow in $n$ layers with variable viscosity

Smooth viscosity-log transition between rigid surface and weak bottom layer

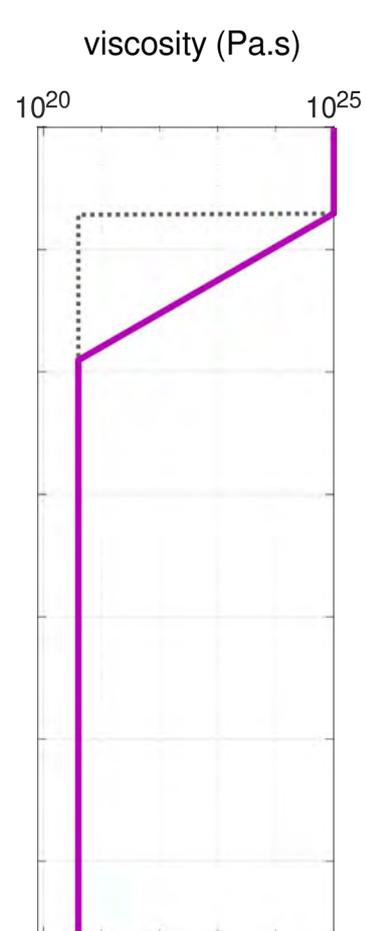
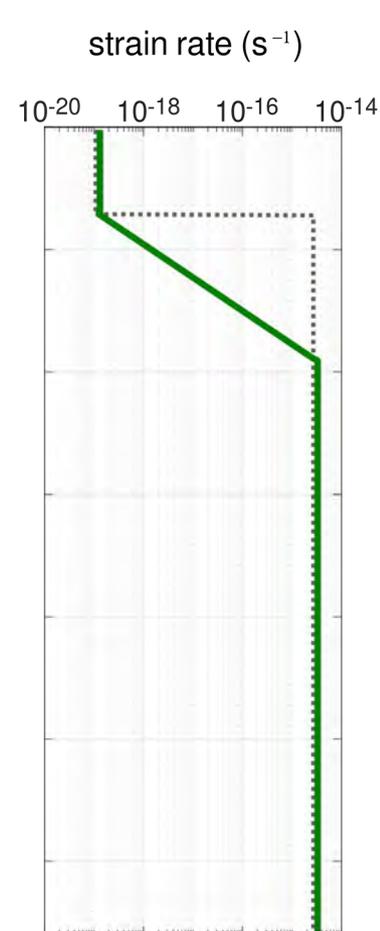
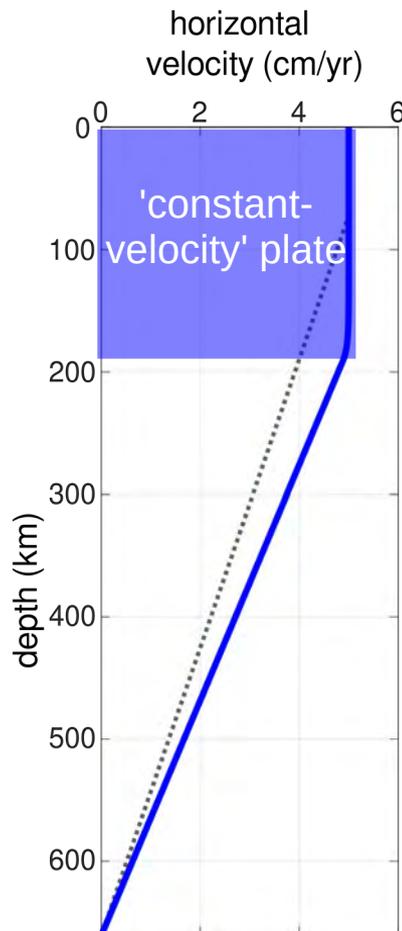
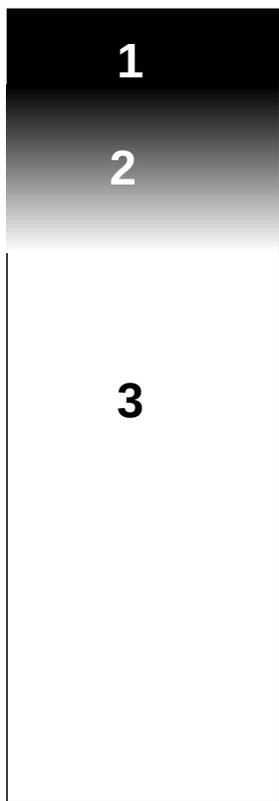
- continuity of viscous shear stress at layer interfaces
- BCs at surface and bottom



# Explaining the 'constant-velocity' plate using a Couette flow in $n$ layers with variable viscosity

Smooth viscosity-log transition between rigid surface and weak bottom layer

- continuity of viscous shear stress at layer interfaces
- BCs at surface and bottom

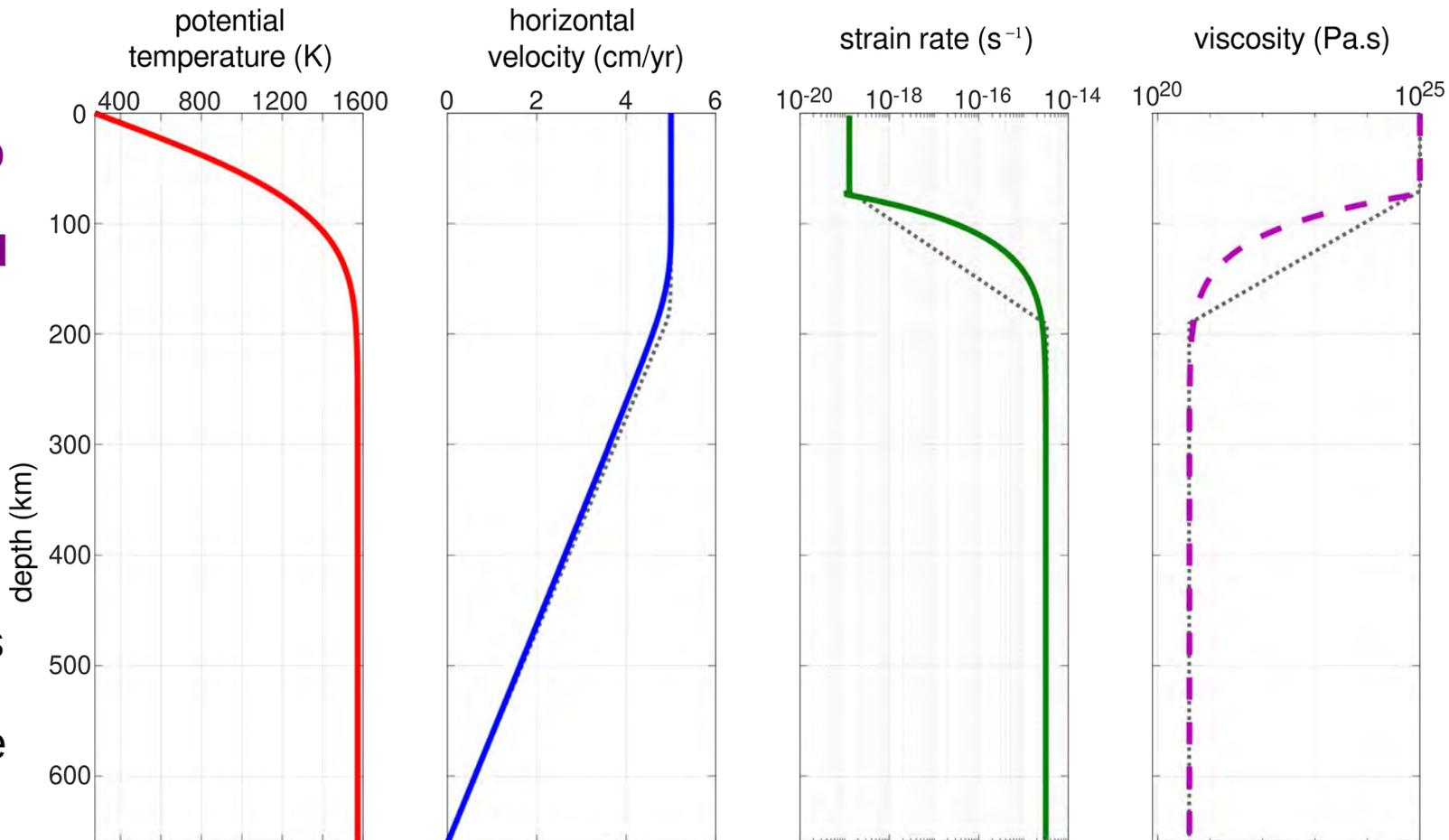


# Explaining the 'constant-velocity' plate using a Couette flow in $n$ layers with variable viscosity

Diffusion creep  
viscosity  
for a 80 Myr-old  
plate

- continuity of  
shear stress  
at layer interfaces

- BCs at surface  
and bottom

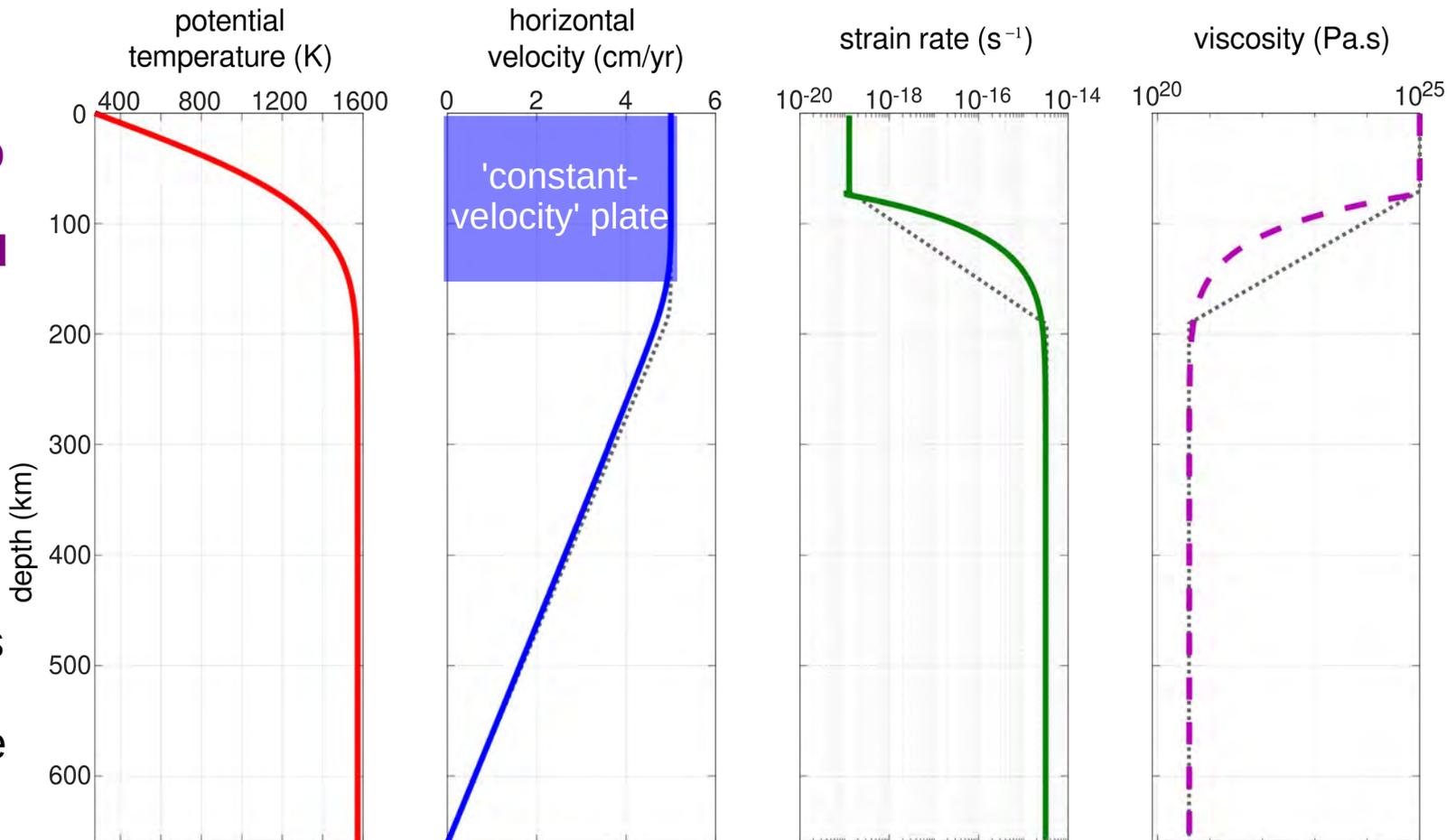


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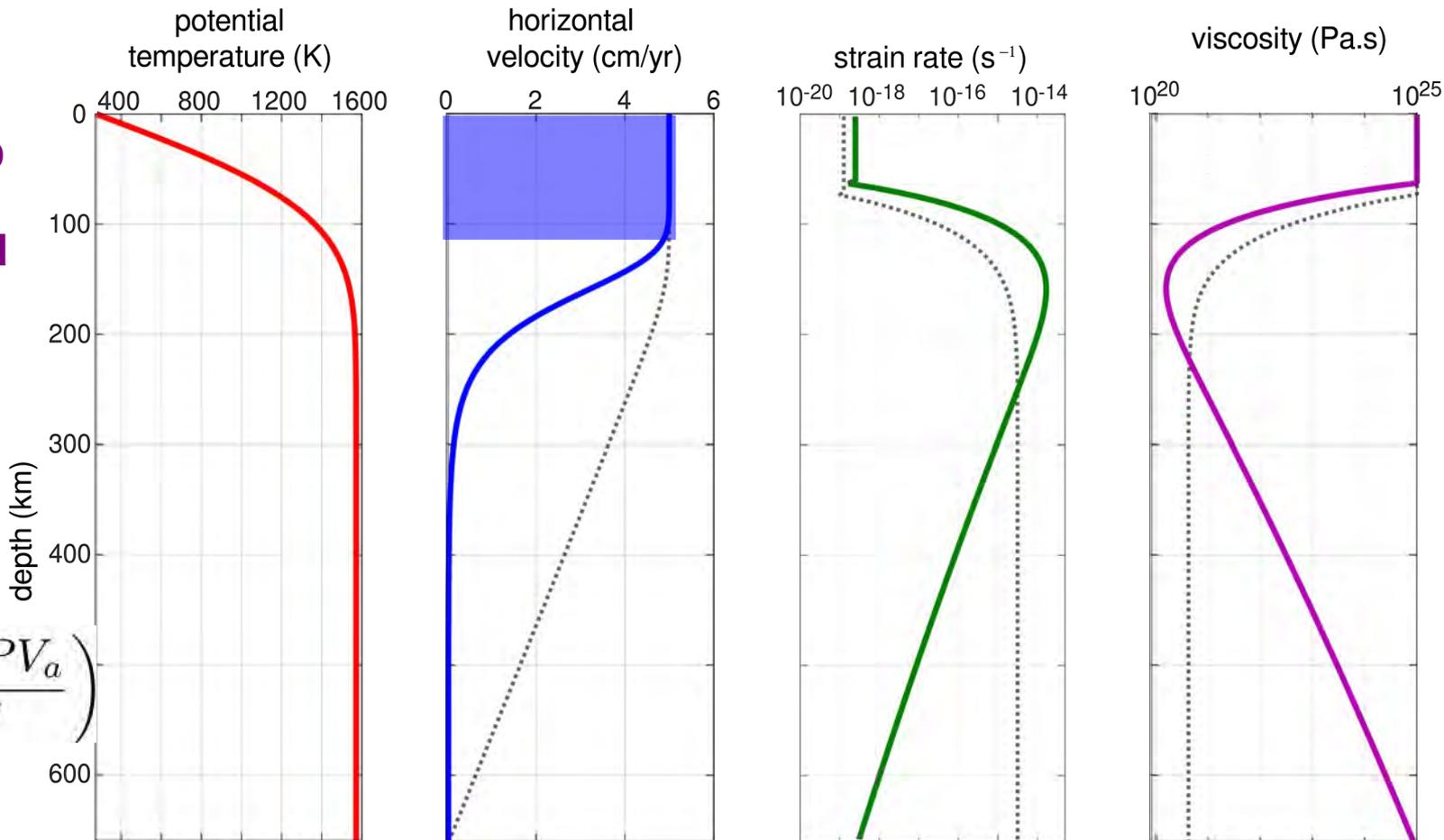


# Explaining the 'constant-velocity' plate using a Couette flow in $n$ layers with variable viscosity

Diffusion creep  
viscosity  
for a 80 Myr-old  
plate

larger  
activation  
volume

$$\mu \propto \exp\left(\frac{E_a + PV_a}{RT}\right)$$

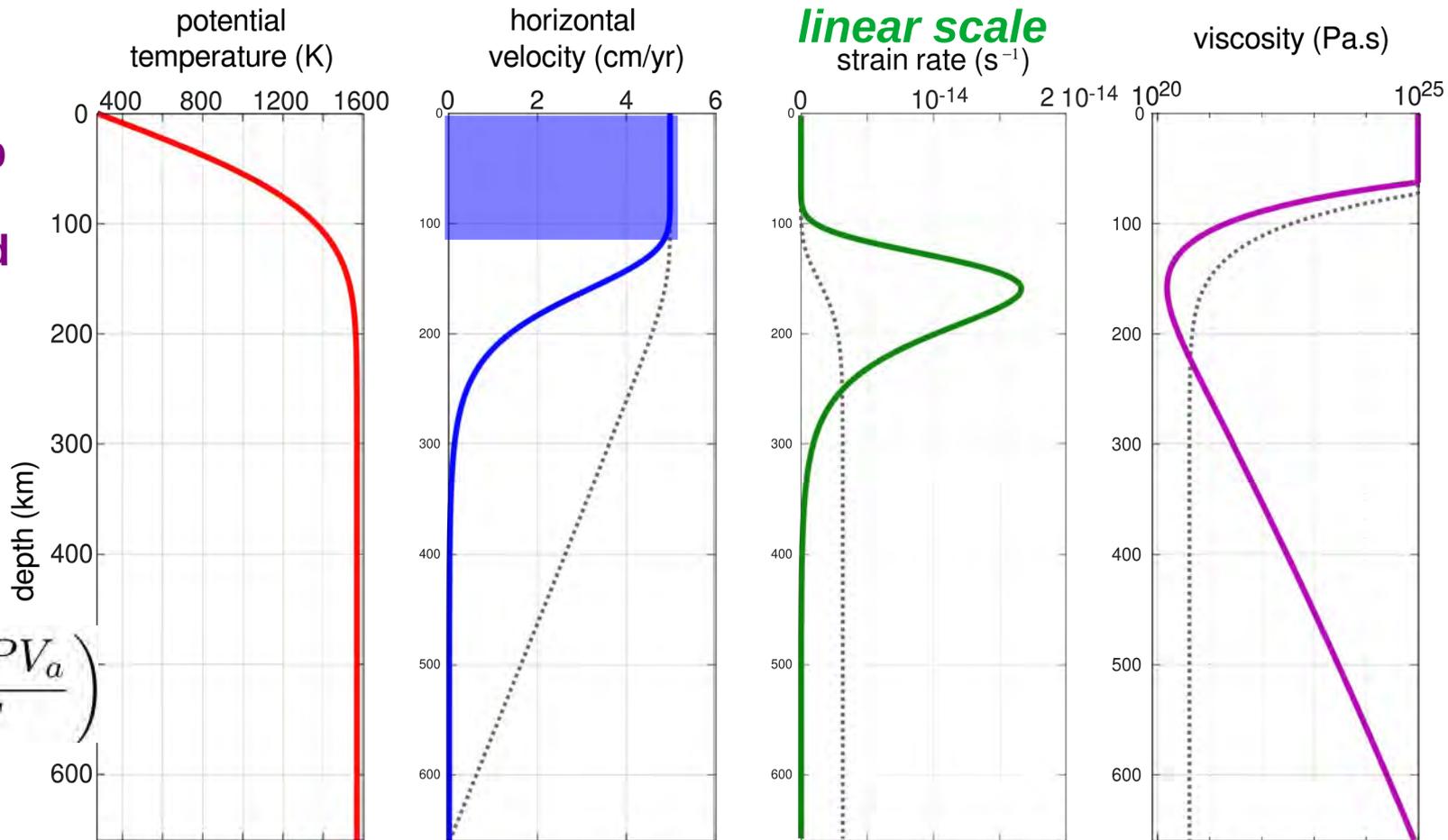


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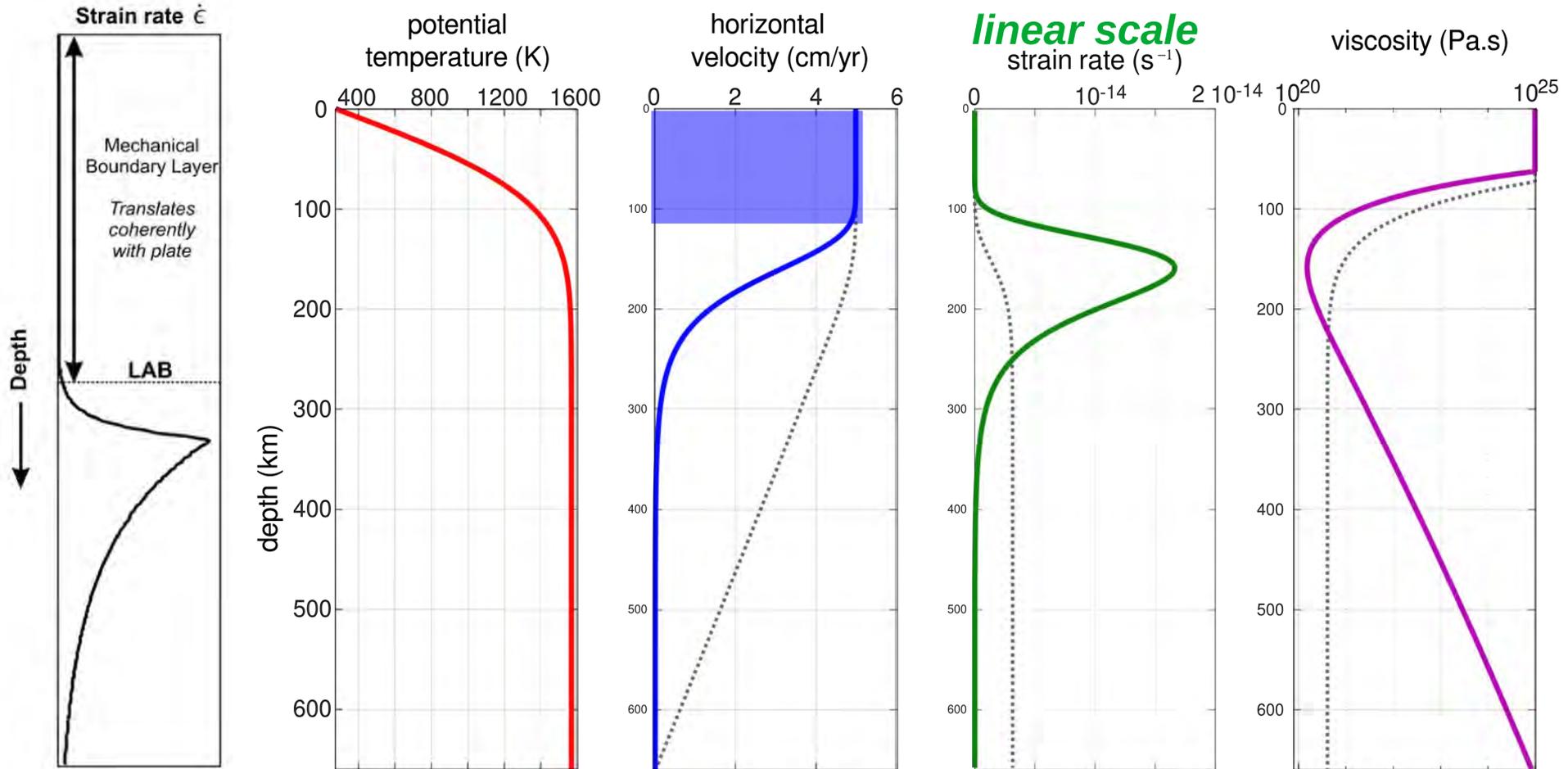
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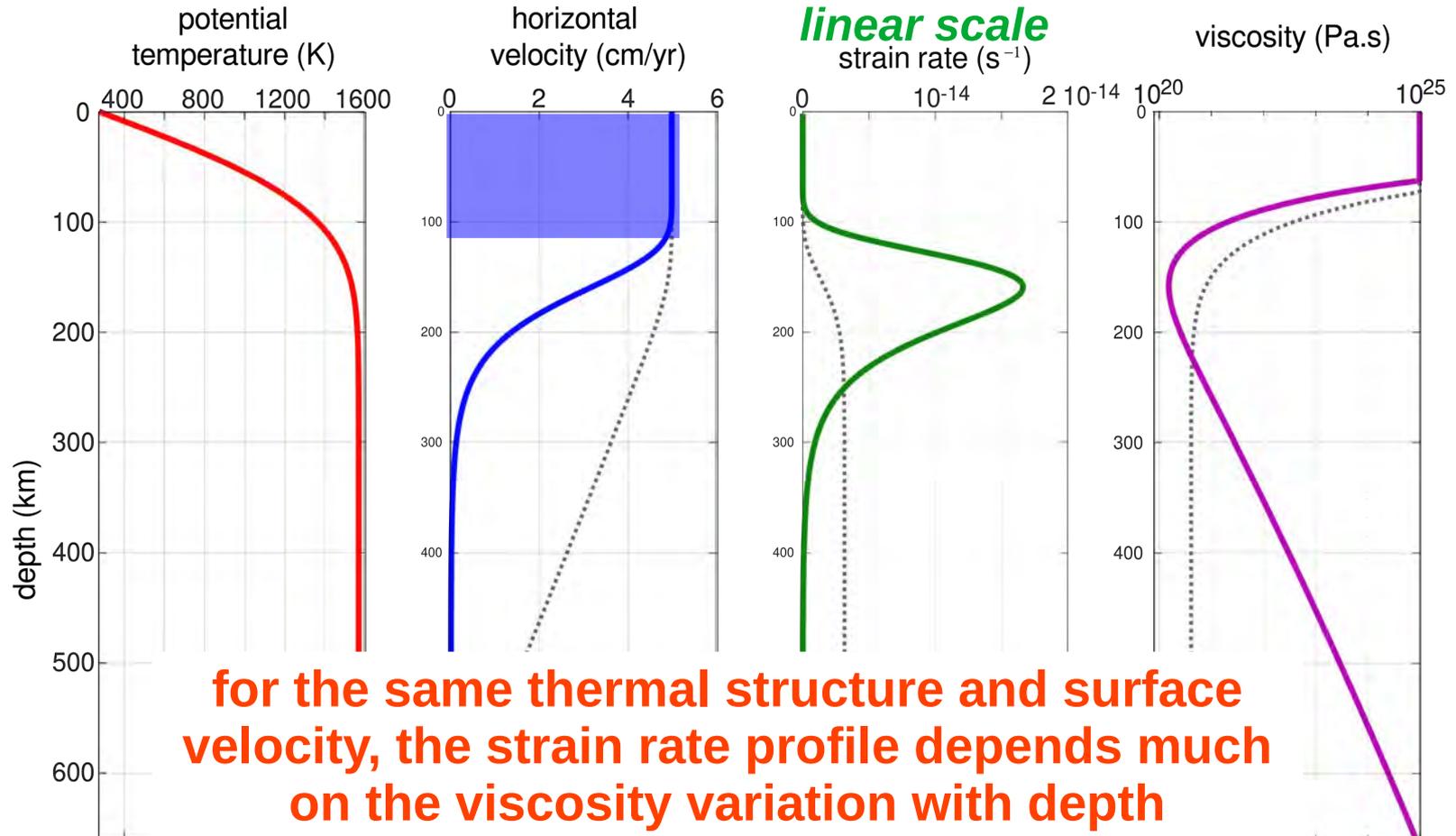
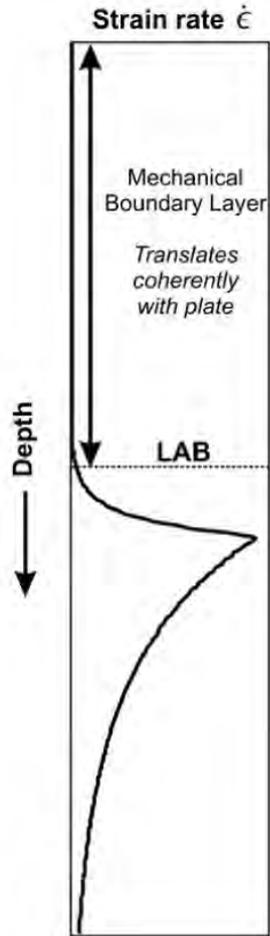
# Explaining the 'constant-velocity' plate using a simple Couette flow in $n$ layers

Eaton et al., 2009



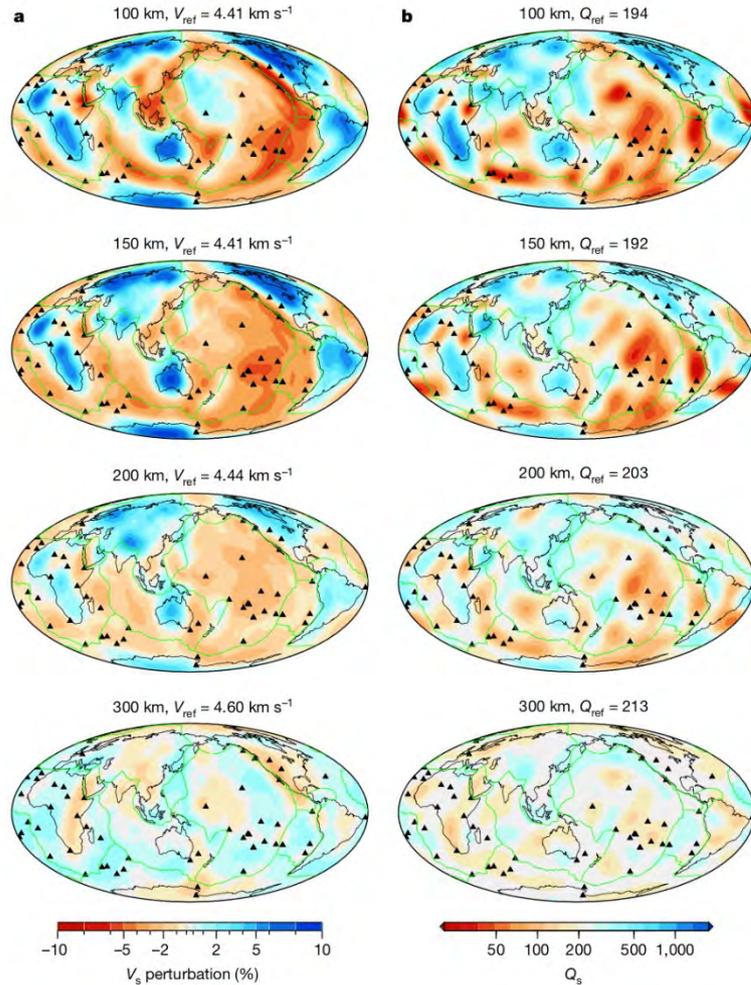
# Explaining the 'constant-velocity' plate using a simple Couette flow in $n$ layers

Eaton et al., 2009

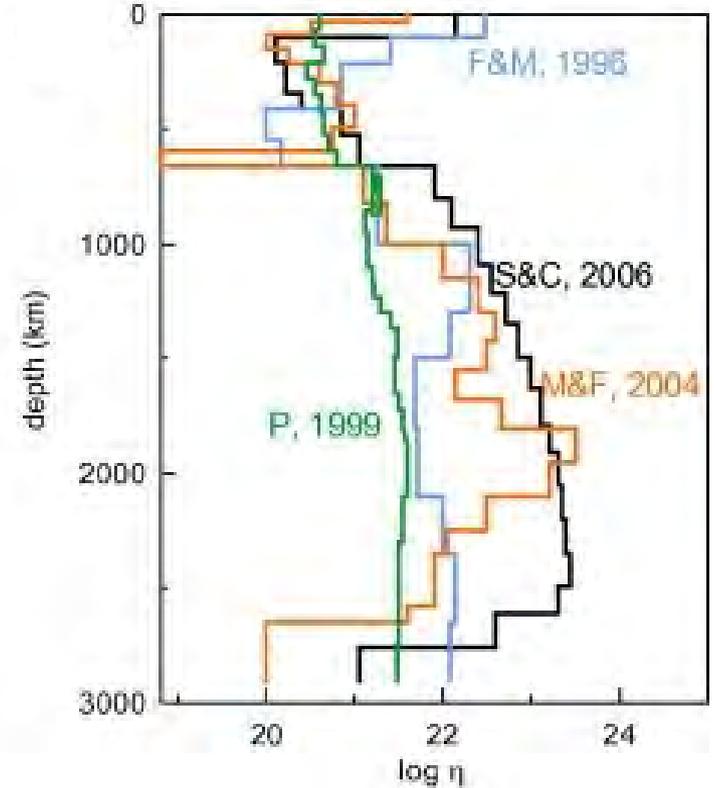


# A low-viscosity layer below the plates ?

## Shear velocity, attenuation



## joint inversion of geoid and postglacial rebound data

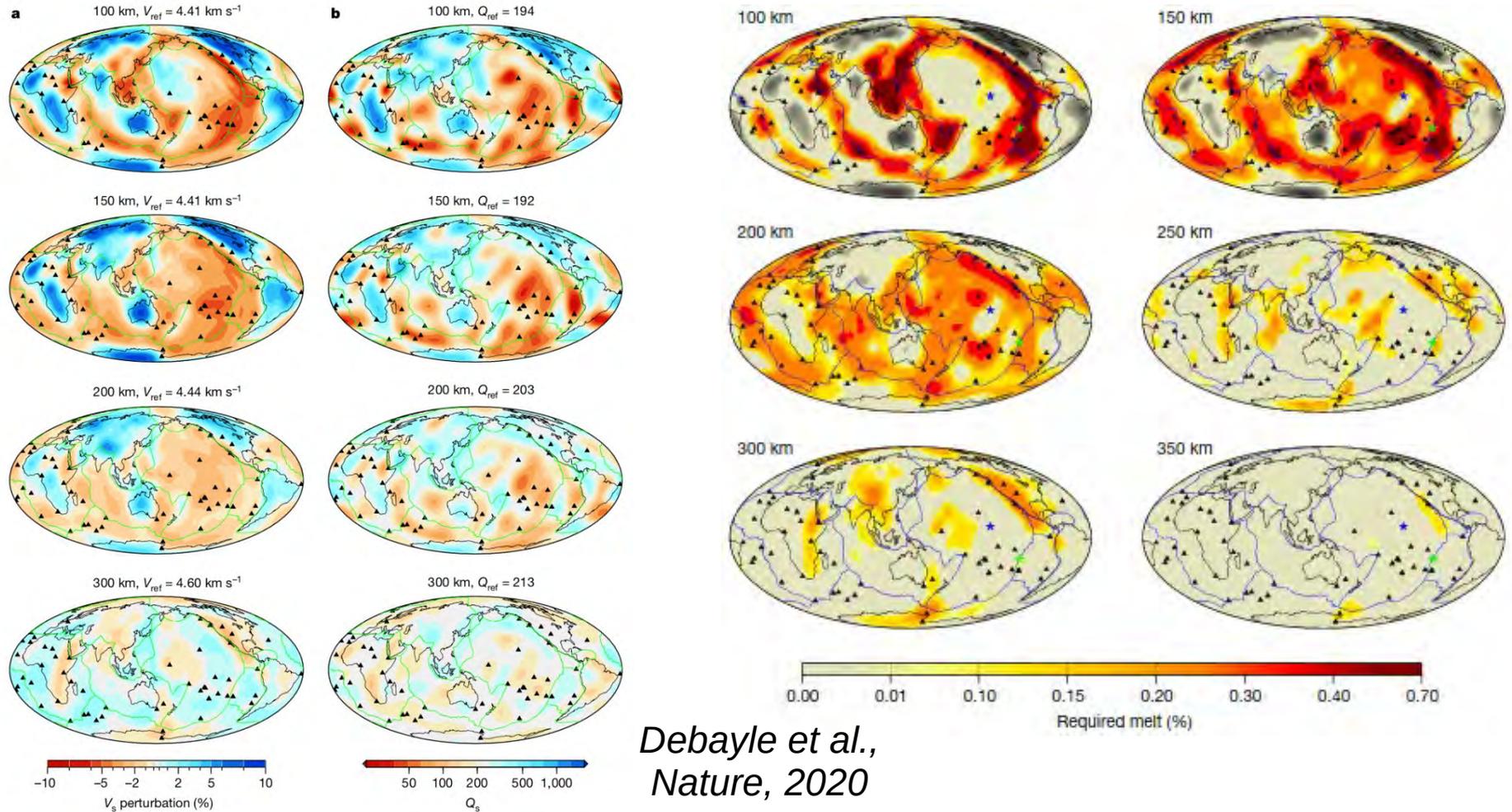


Debaille et al.,  
Nature, 2020

Compilation in Cizkova  
et al., PEPI, 2012

# A low-viscosity layer below the plates ?

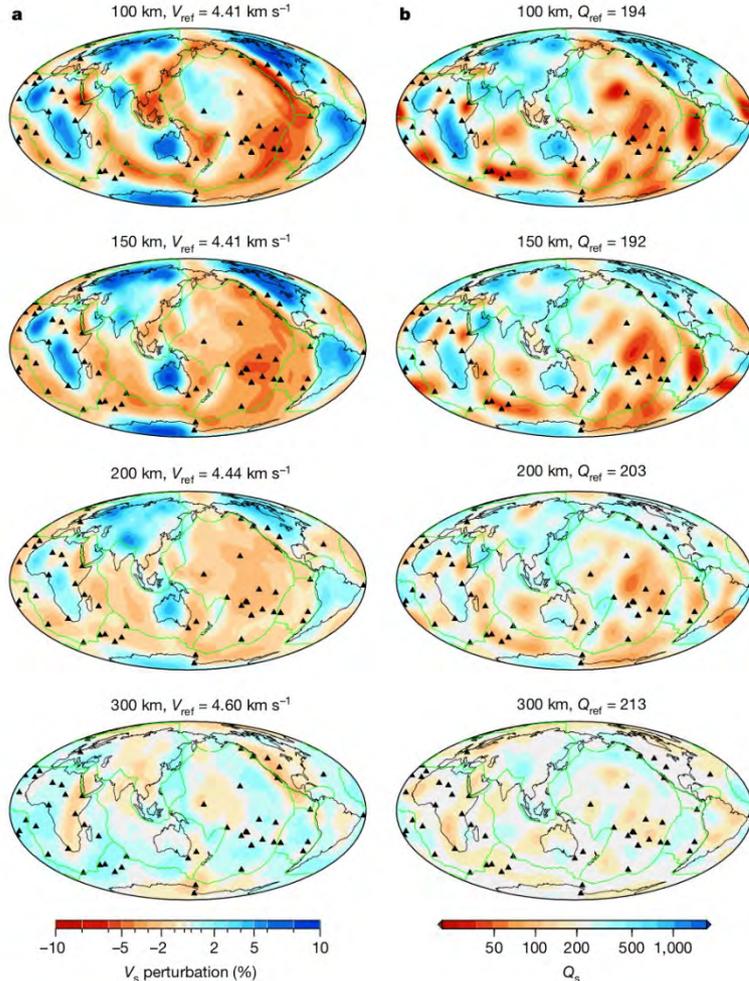
Shear velocity, attenuation -----▶ partial melt ?



Debayle et al.,  
Nature, 2020

# A low-viscosity layer below the plates ?

Shear velocity, attenuation -----► partial melt ?



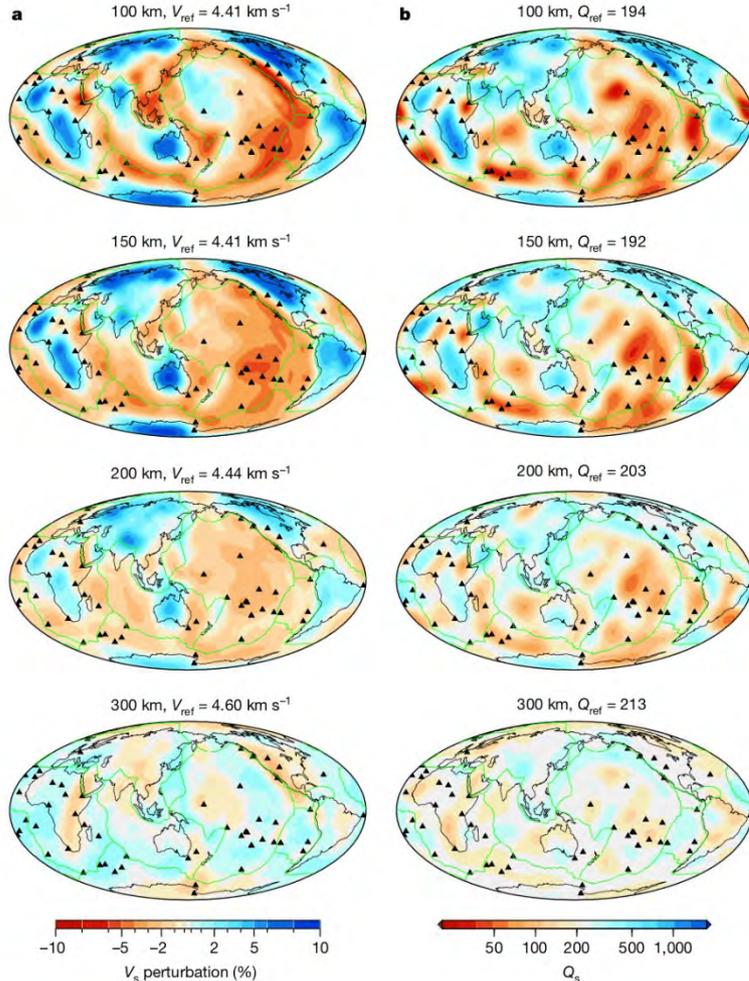
-----► grain size variations ?

creep laws → low-viscosity layer possible with certain rheological parameters

*Debayle et al.,  
Nature, 2020*

# A low-viscosity layer below the plates ?

Shear velocity, attenuation -----► partial melt ?



-----► grain size variations ?

creep laws → low-viscosity layer possible with certain rheological parameters

**how weak?**

$10^{19} \text{ Pa.s ?}$

**how much weaker?**

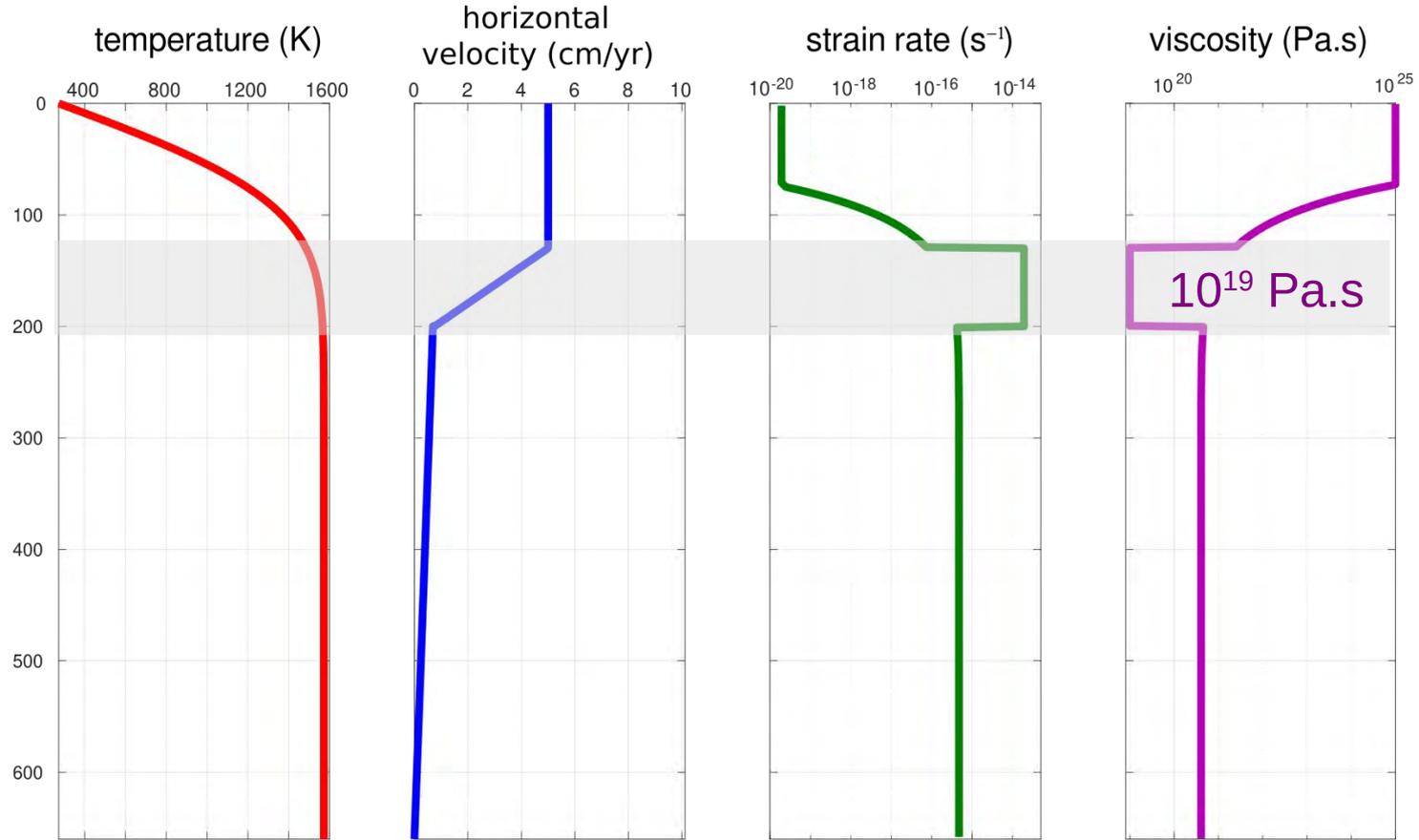
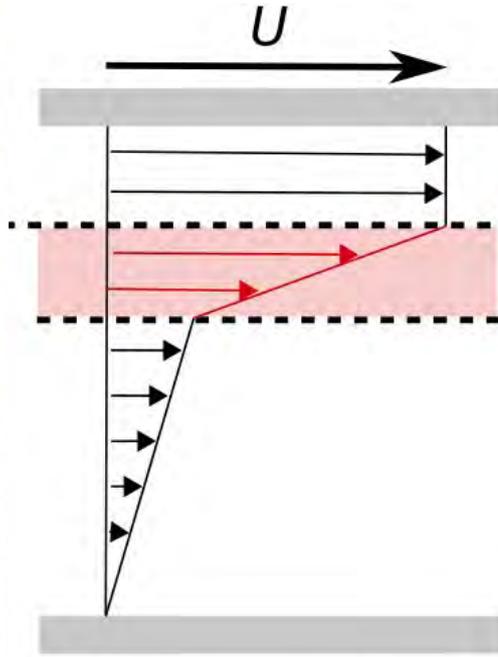
10-100 x ?

**how thick?**

< 250 km depth ?

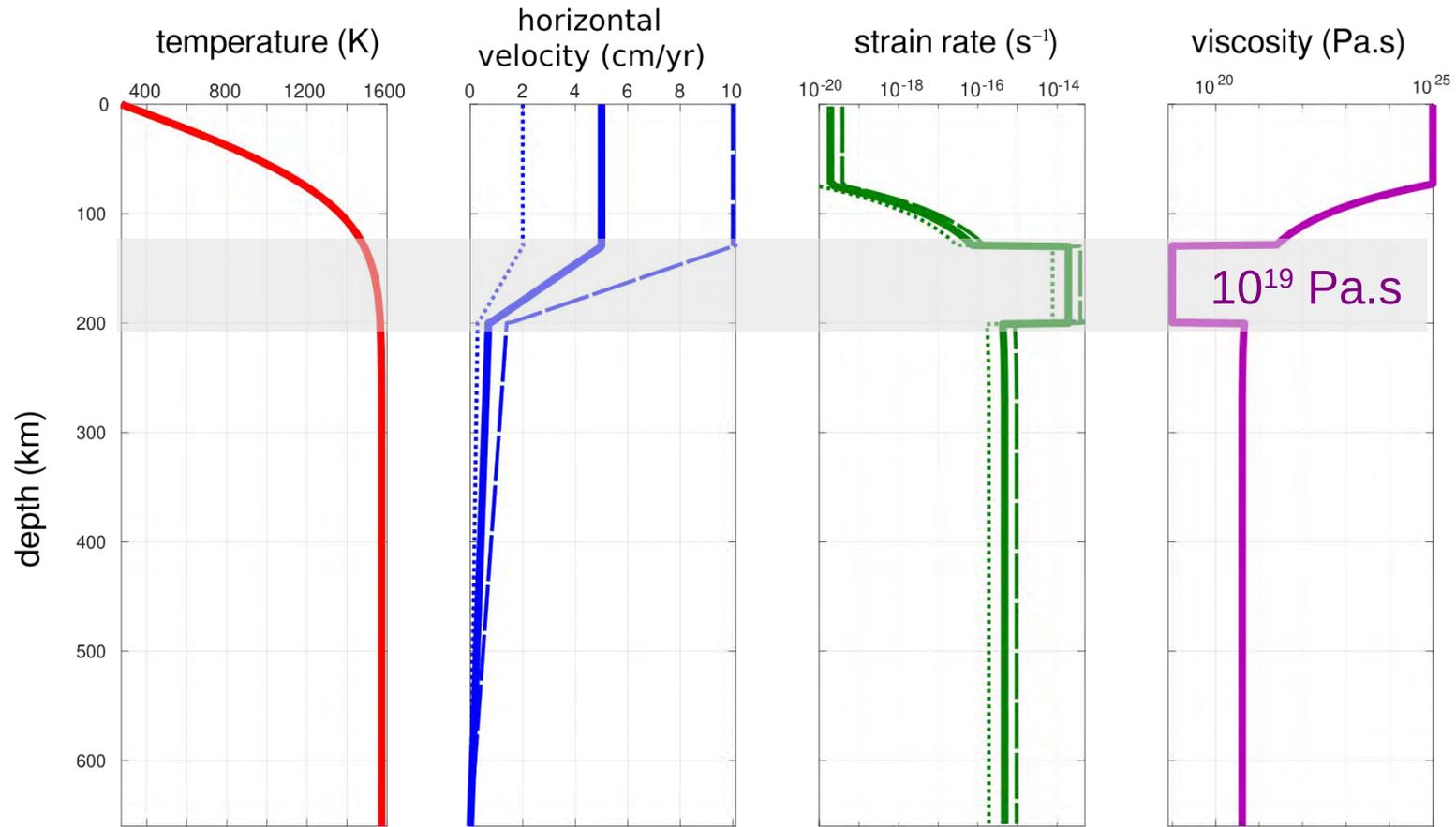
Debayle et al.,  
Nature, 2020

# A decoupling low-viscosity layer ?



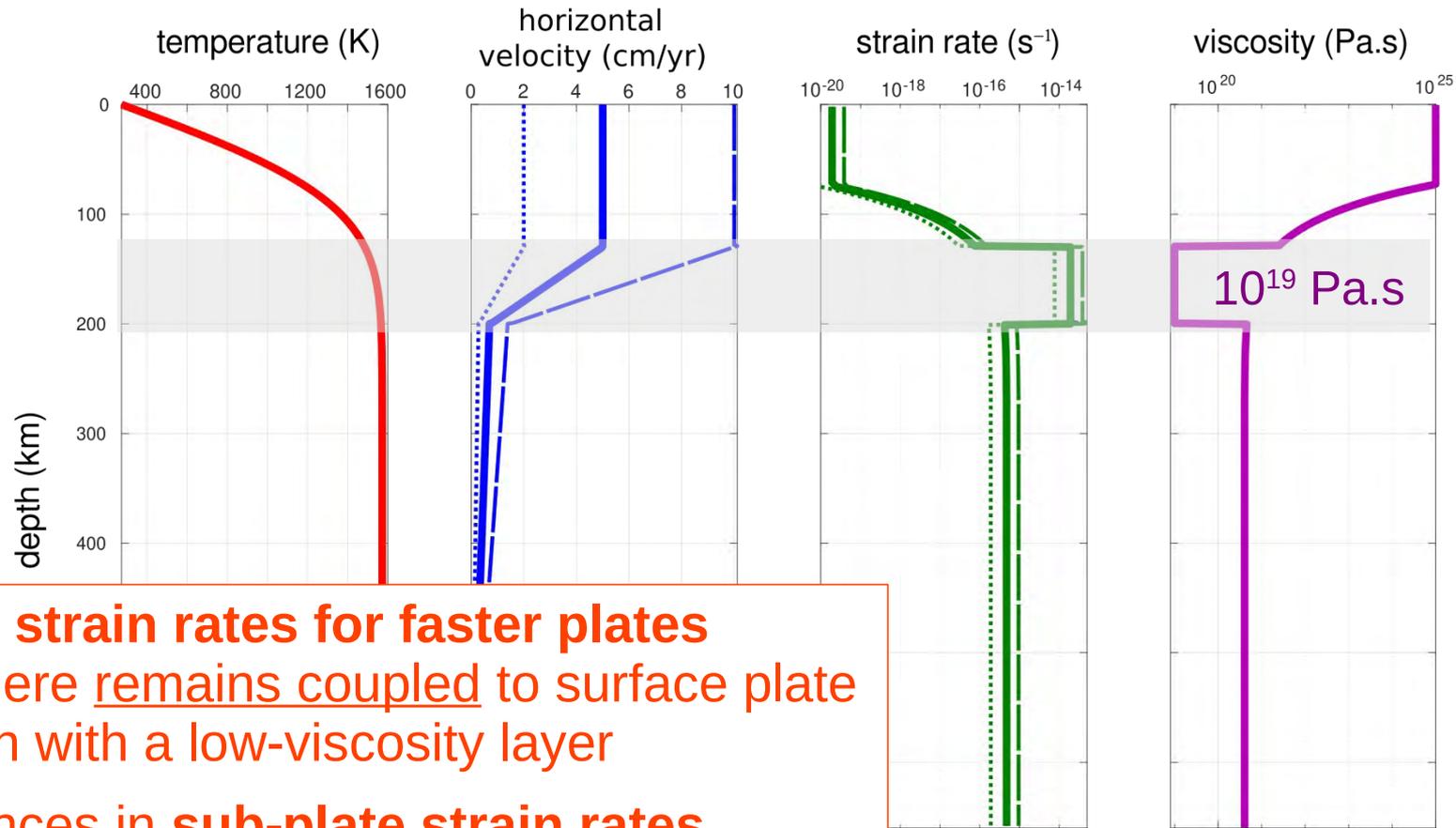
# A decoupling low-viscosity layer ?

Various  
surface plate  
velocities :  
2, 5, 10 cm/yr



# A decoupling low-viscosity layer ?

Various surface plate velocities : 2, 5, 10 cm/yr



- **higher strain rates for faster plates**  
→ asthenosphere remains coupled to surface plate even with a low-viscosity layer
- differences in sub-plate strain rates proportional to surface velocity ratio

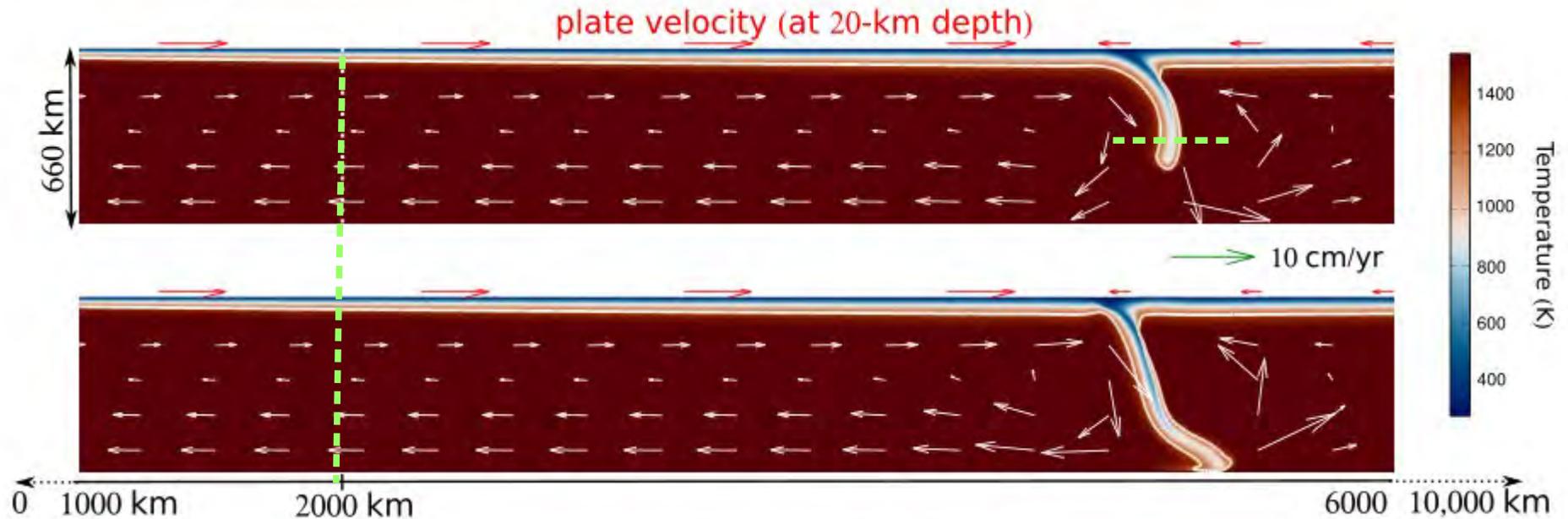
# Estimating the time required to develop crystal-preferred orientation under a constant velocity field

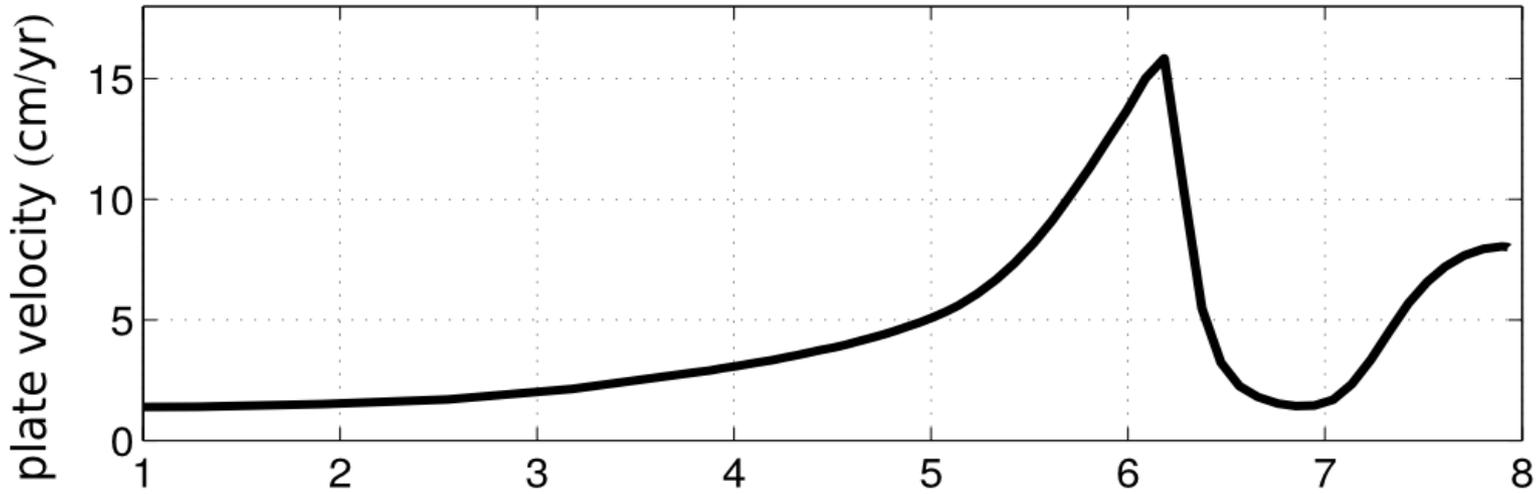
- hyp.: cumulative strain  $\varepsilon$  of 100 % needed to produce CPO retrievable from seismic anisotropy
- time = strain / strain rate =  $\varepsilon / \dot{\varepsilon}$

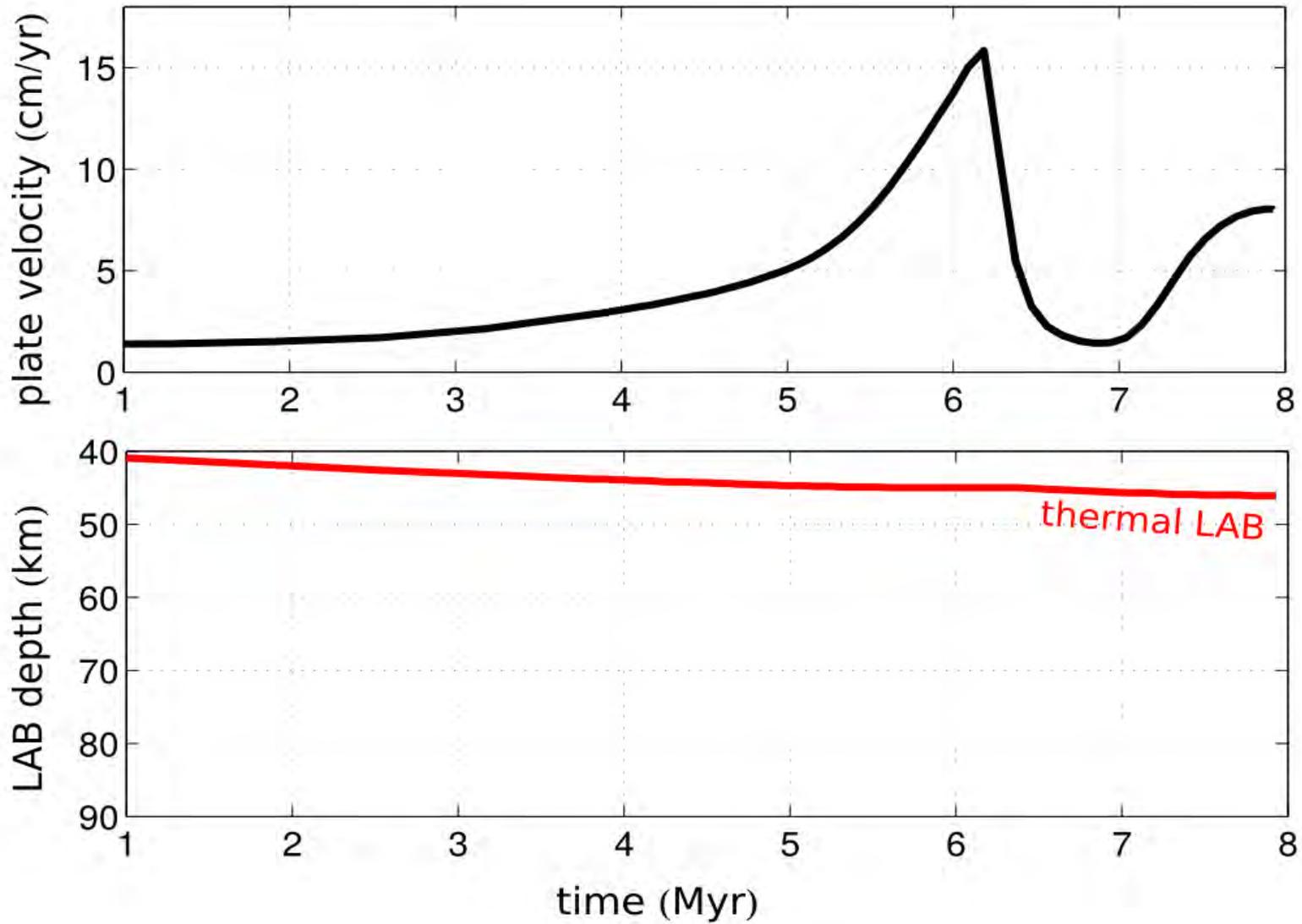
$10^{-14} \text{ s}^{-1}$	$\leftrightarrow$	3 Myr
$10^{-15} \text{ s}^{-1}$	$\leftrightarrow$	30 Myr
$10^{-16} \text{ s}^{-1}$	$\leftrightarrow$	300 Myr

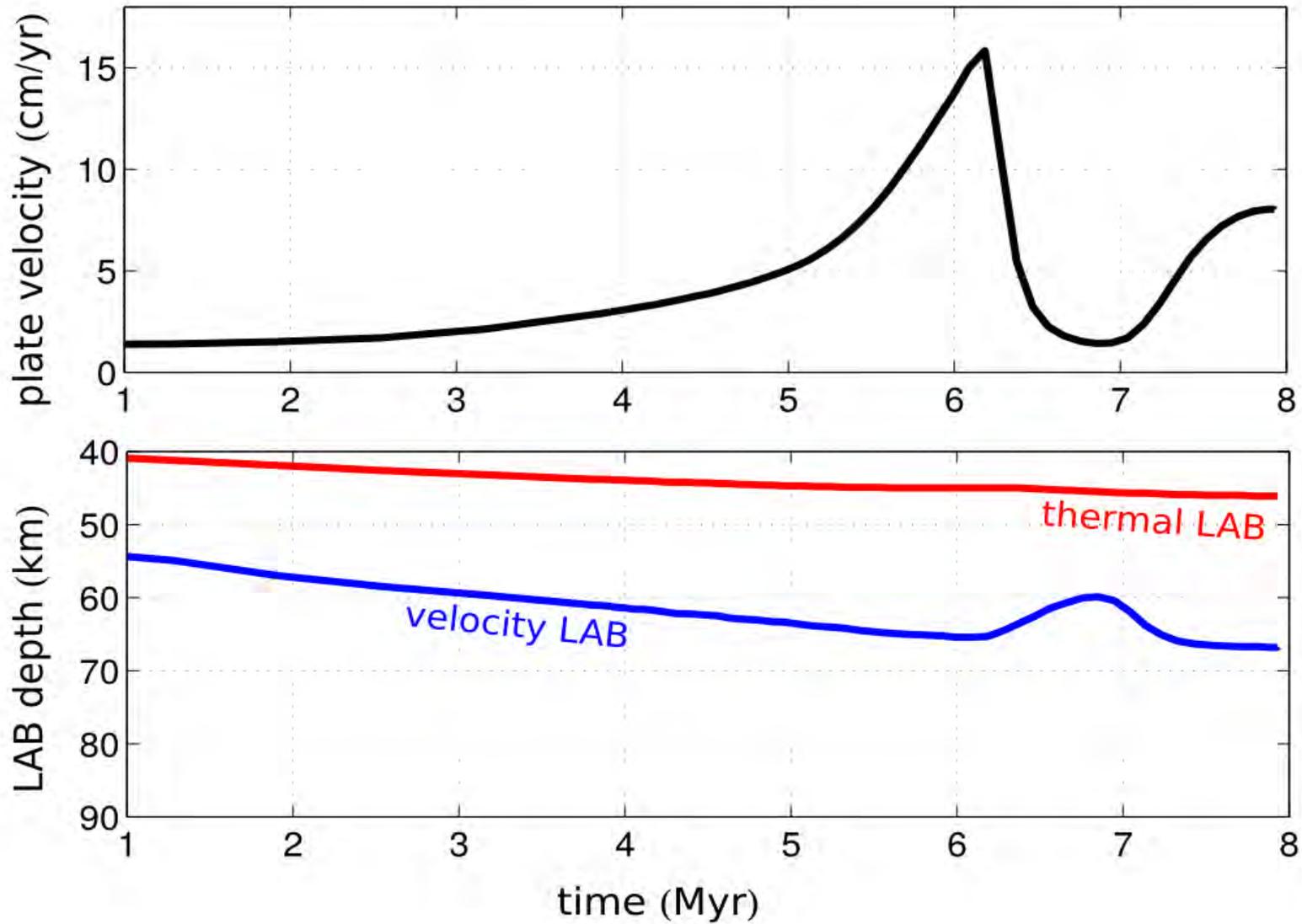
*But plate velocity magnitude and direction are not fixed...*

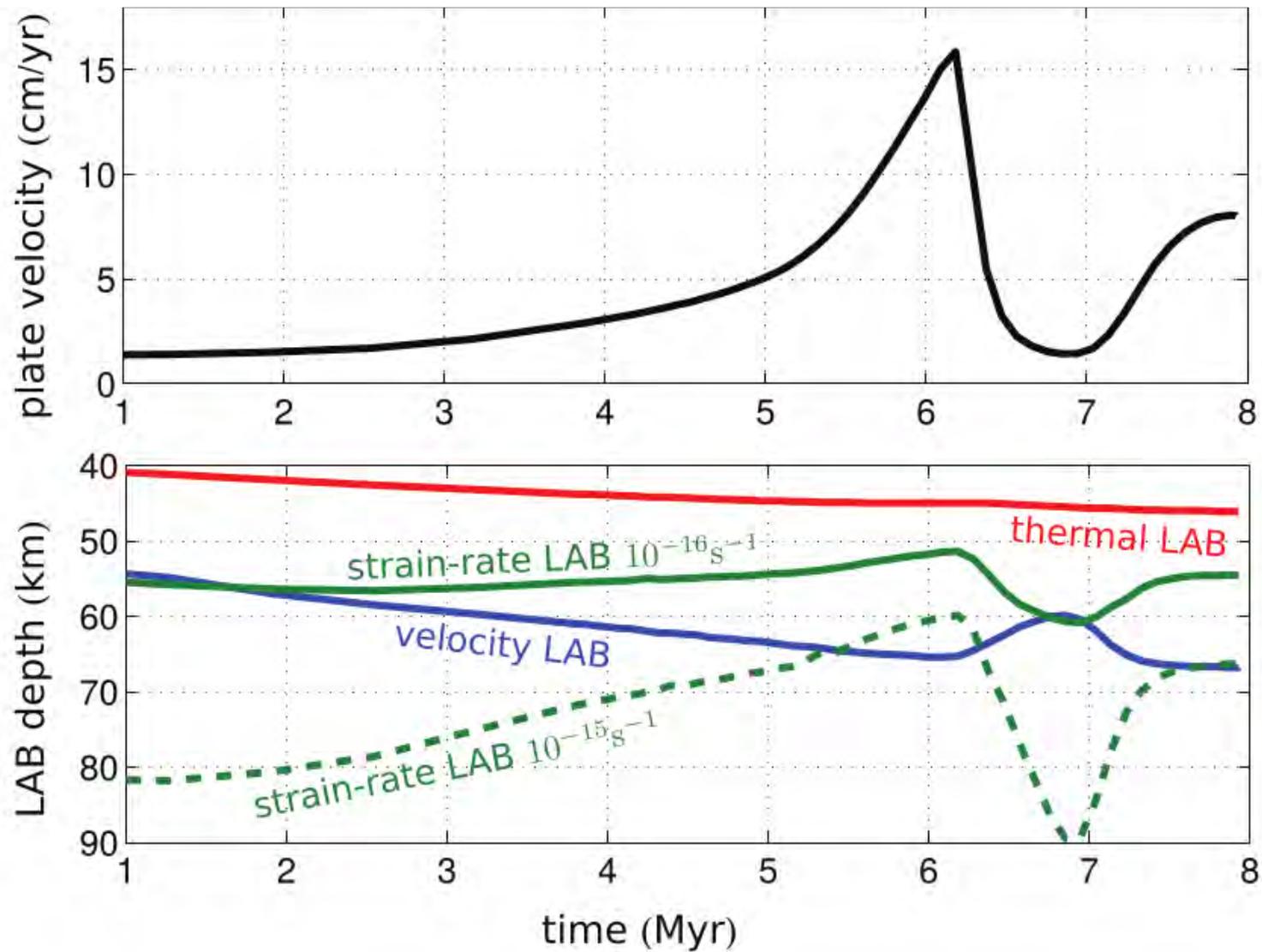
# Transient asth. flow driven by subduction





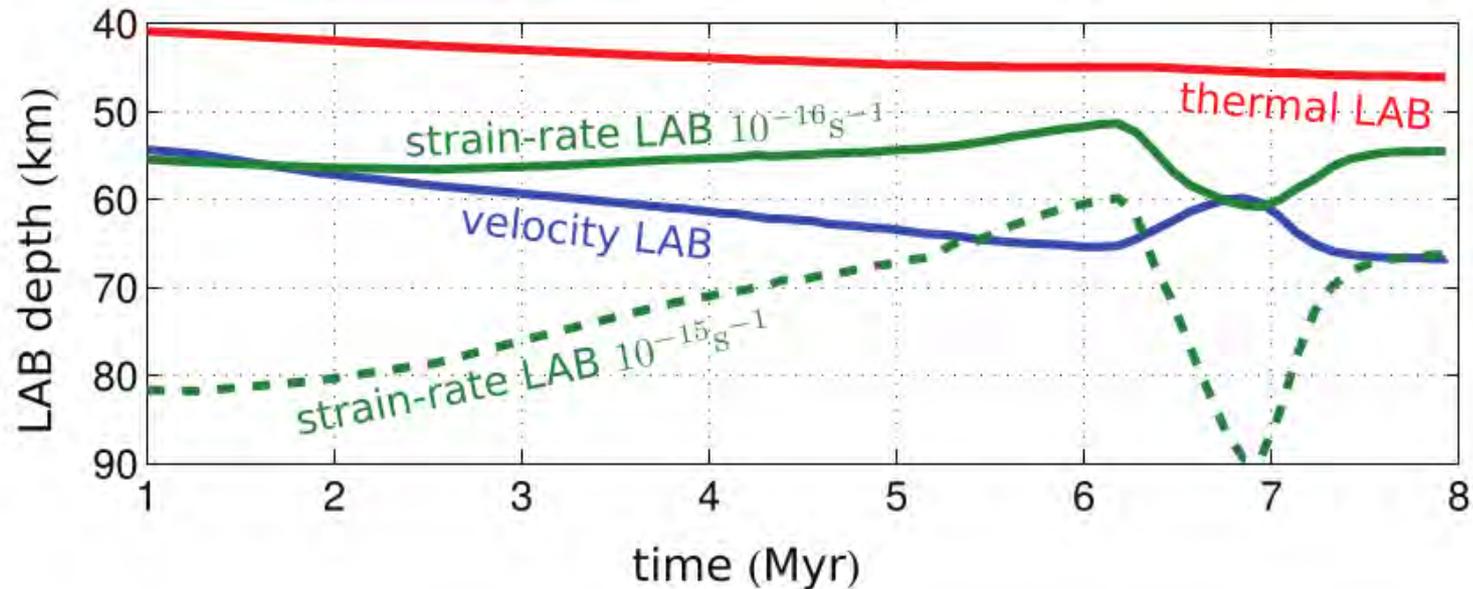
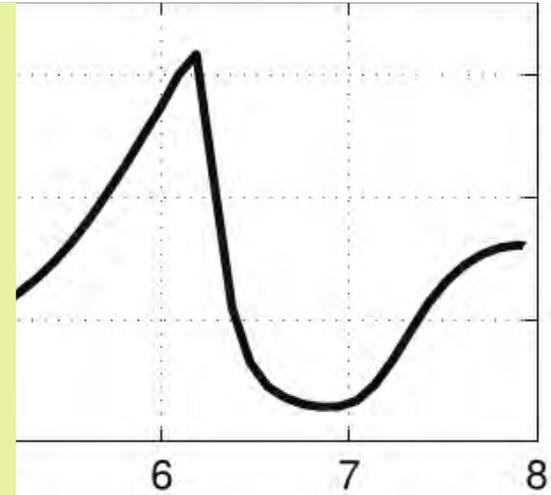






*lower strain rates reached at depth when plate decelerates (and vice-versa)*

→ if a plate slows down, past CPO will persist longer for lower strain rate (slower plate)

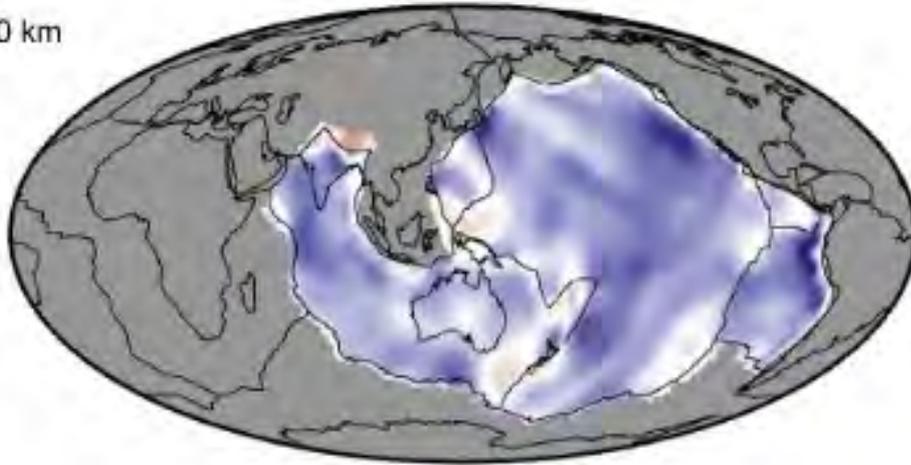


# Agreement between fast direction of $S_v$ waves and present-day absolute plate motion (from NUVEL-1A)

blue = parallelism

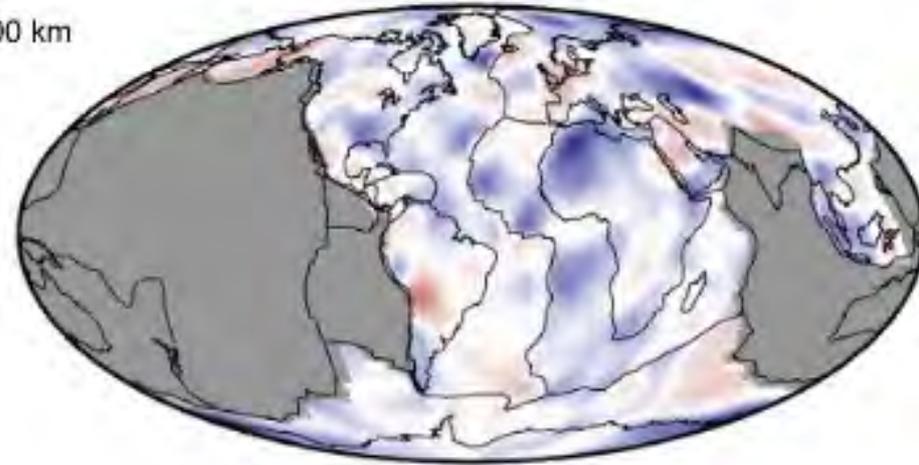
red = orthogonality

200 km



Fast-moving plates

200 km



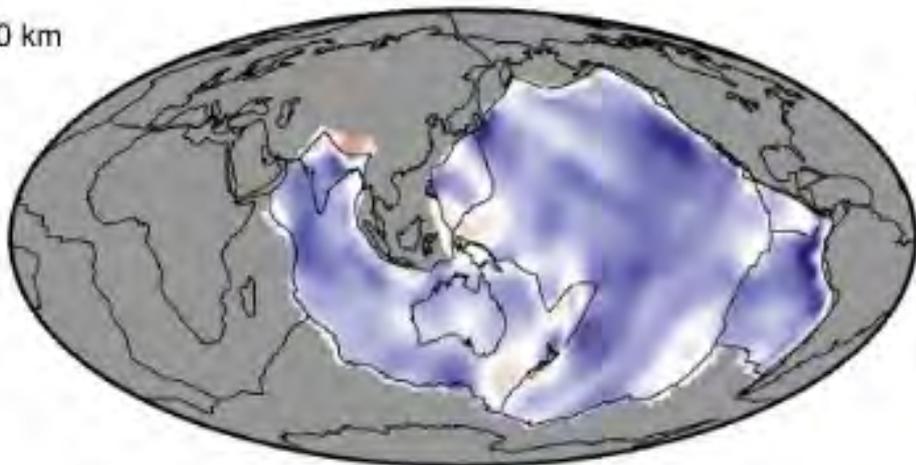
Slow-moving plates

# Agreement between fast direction of $S_v$ waves and present-day absolute plate motion (from NUVEL-1A)

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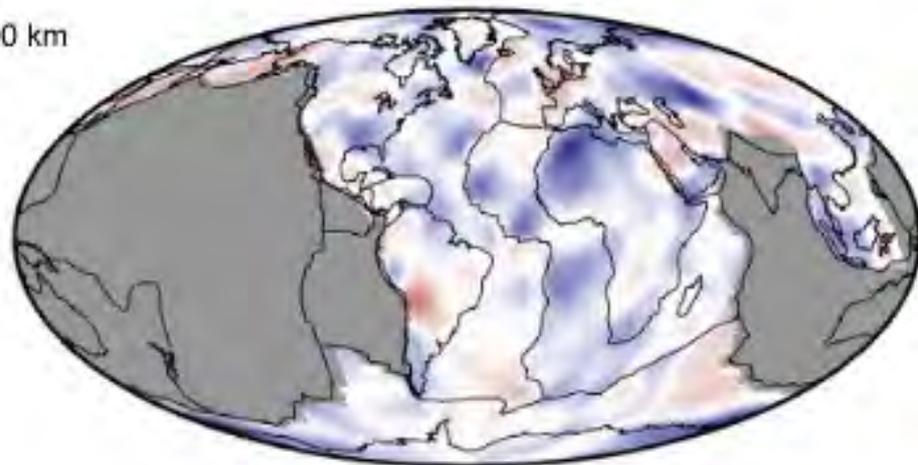
red = orthogonality

200 km



Fast-moving plates

200 km



Slow-moving plates

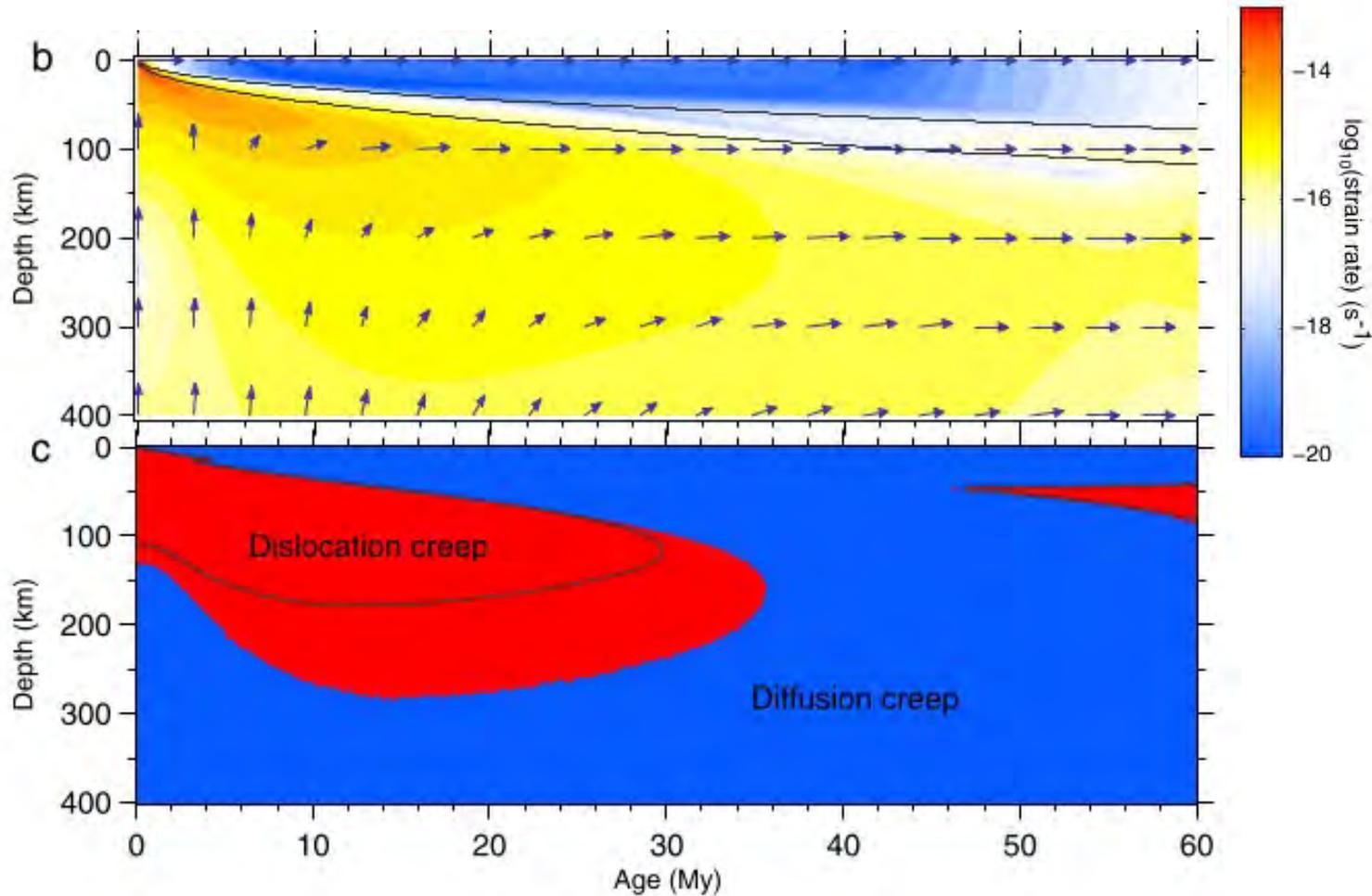
*insufficient strain rates in the asthenosphere to generate CPO aligned with plate motion in less than 30 Myr*

Debaille & Ricard, 2013

# First conclusions

- From thermo-mechanical models :
  - a dynamical transition from a **constant-velocity plate** to the underlying mantle appears self-consistently due to viscosity decrease with depth
    - the constant-velocity plate is **not fully rigid** and deforms at its base
    - the constant-velocity plate is **transient** and adjusts to flow field evolution
- Even with a low-viscosity layer, asthenosphere flow and strain rates depend on surface plate velocities
- **Below slow plates anisotropy fast axis direction may align with past mantle flow** because of the long time-scales required to develop mantle CPO under small strain rates

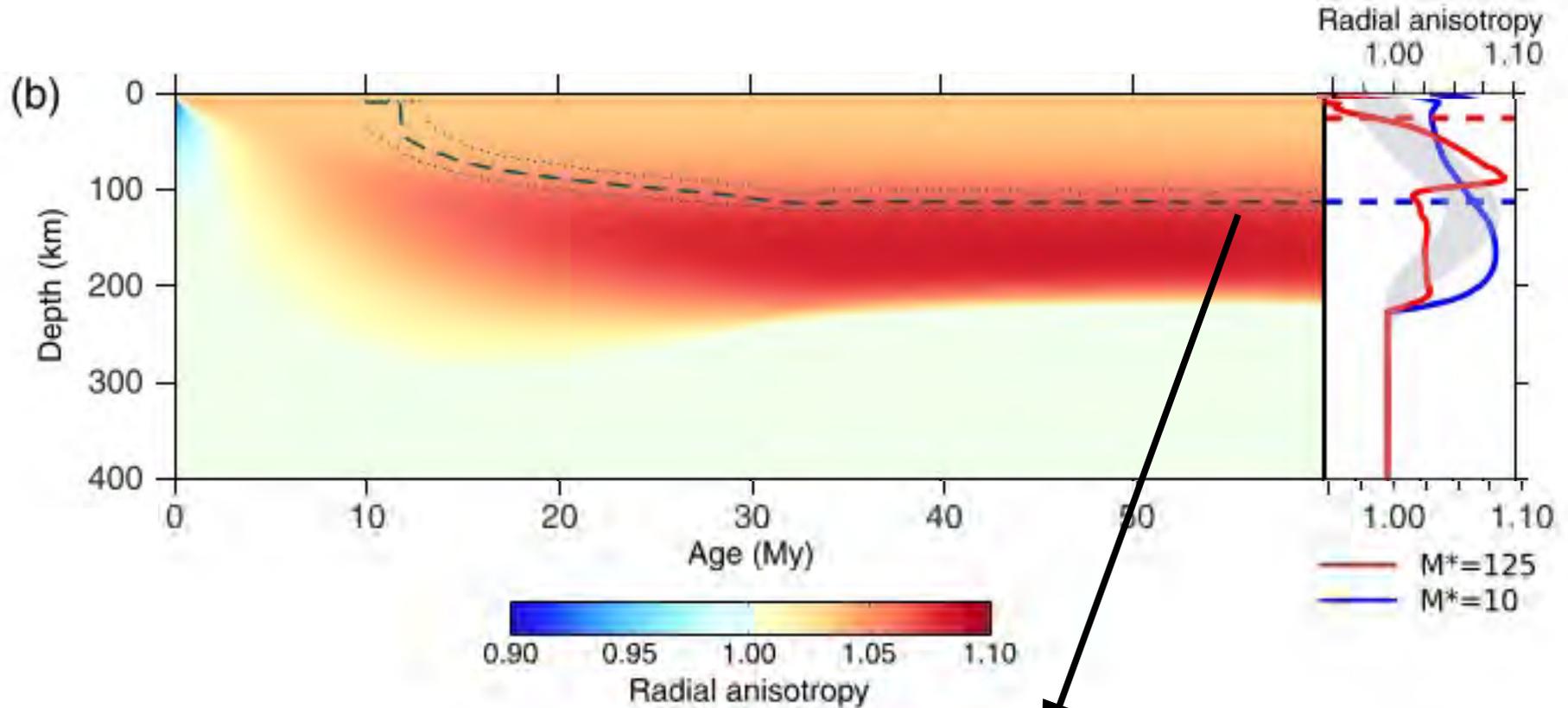
# Complexifying models towards the real Earth : dislocation vs. diffusion creep regimes



Steady-state  
flow model  
below a plate  
+ composite  
rheology  
(diff + disl creep)  
+ D-Rex calculation  
of anisotropy

Hedjazian et al.,  
EPSL 2017

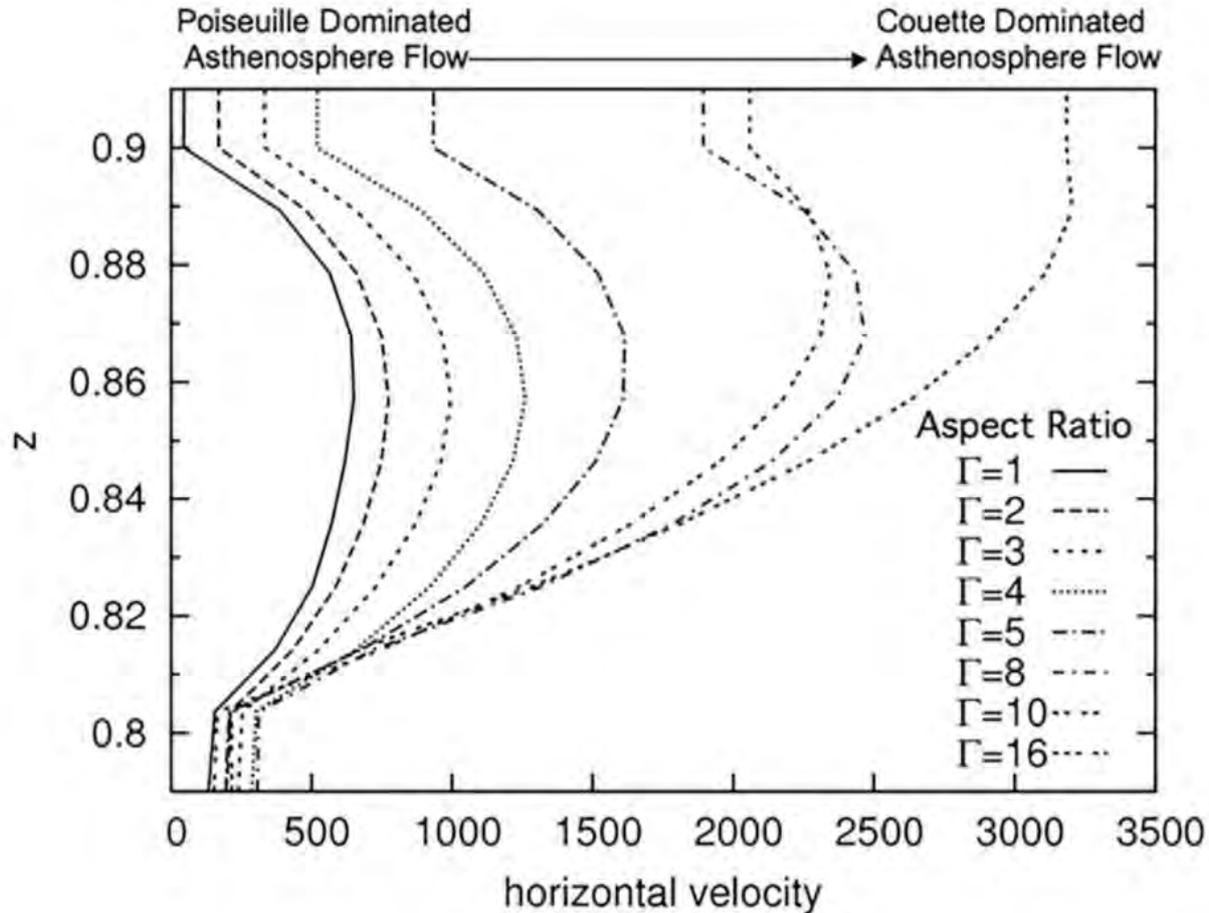
# Complexifying models towards the real Earth : dislocation vs. diffusion creep regimes



**age-independent radial seismic anisotropy !**

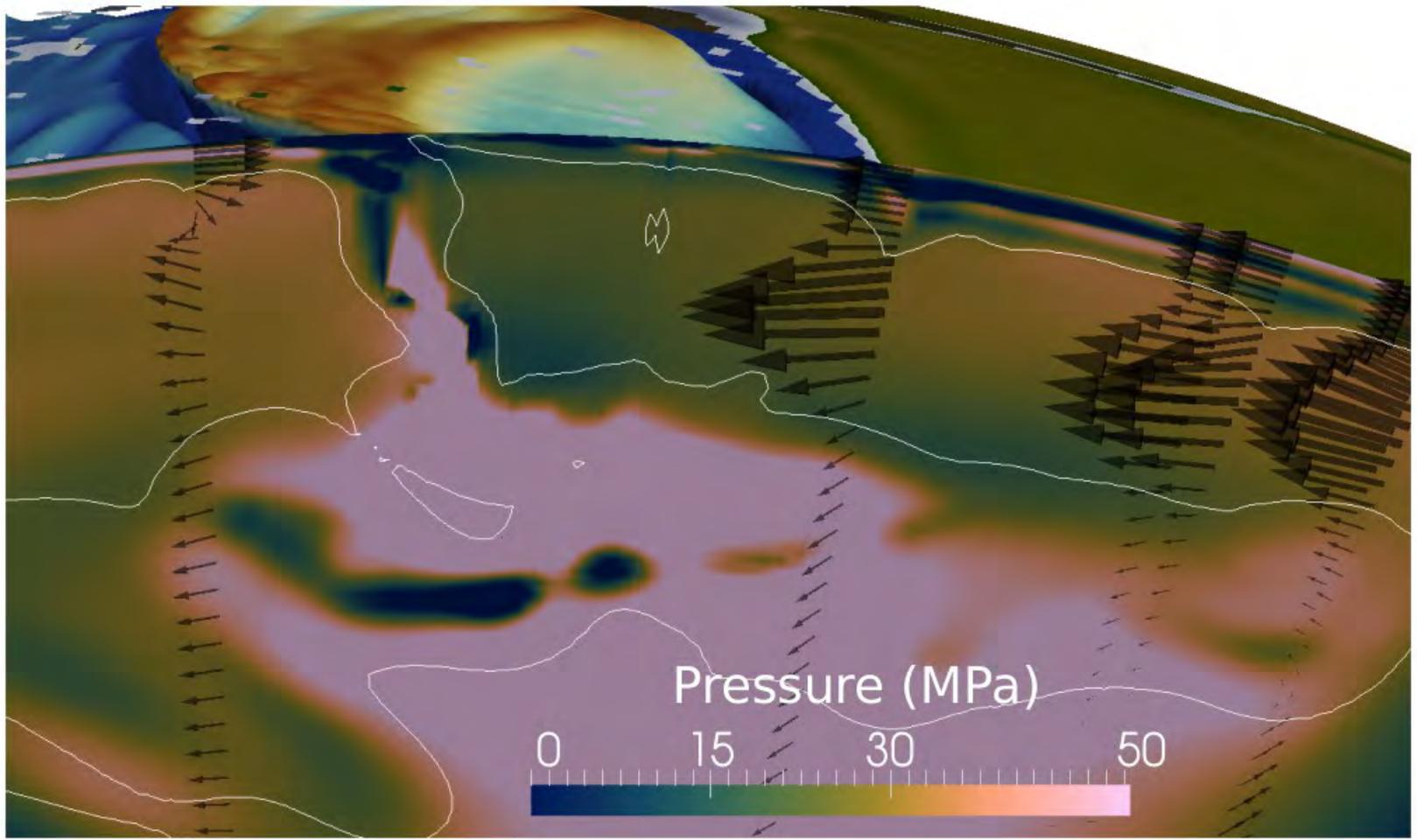
Hedjazian et al.,  
EPSL 2017

# Complexifying models towards the real Earth : Couette vs. Poiseuille flows



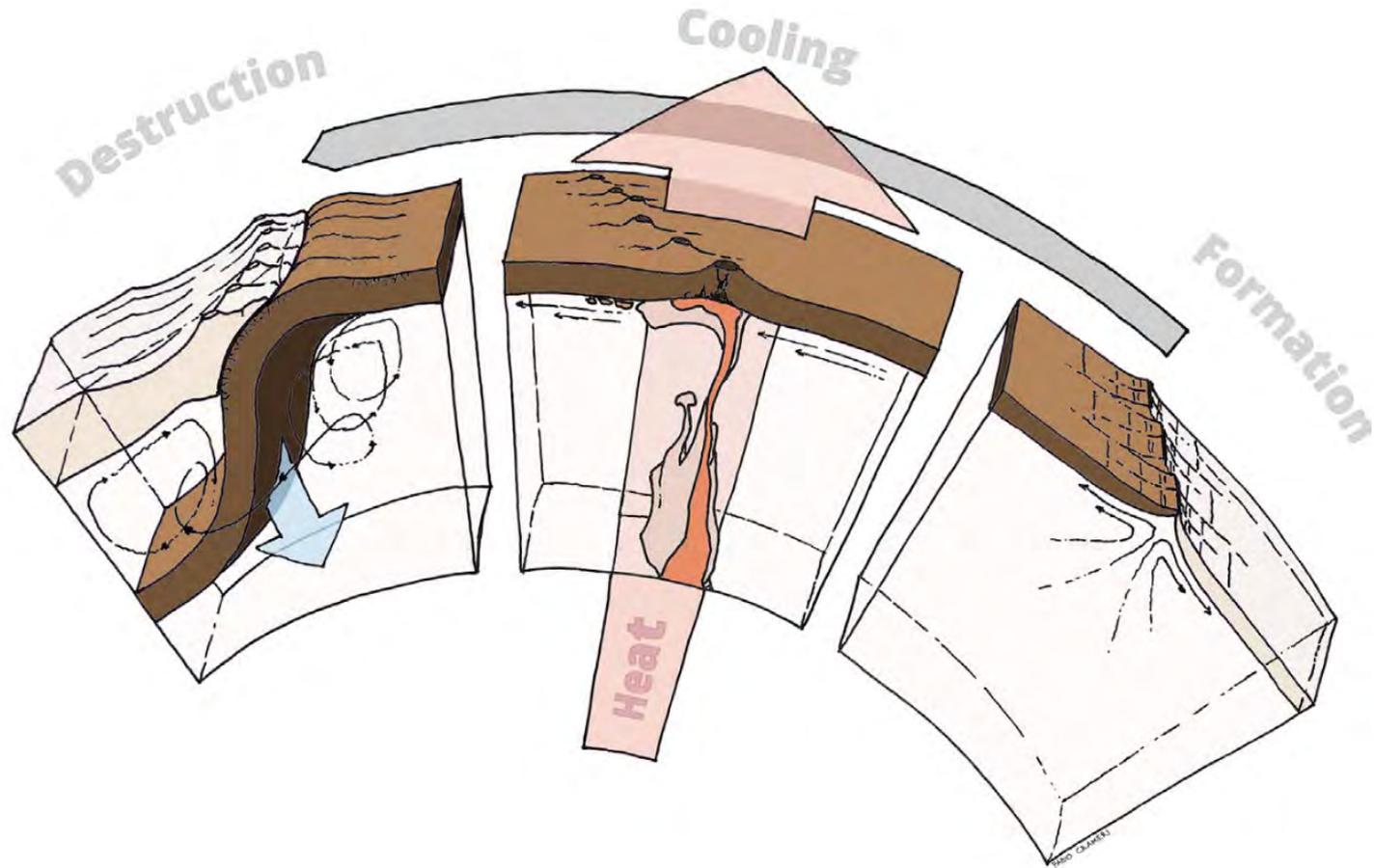
**constant-velocity plate  
also observed  
for 'active'  
asthenosphere flow**

Hoink and Lenardic, 2010  
Richards and Lenardic, 2018



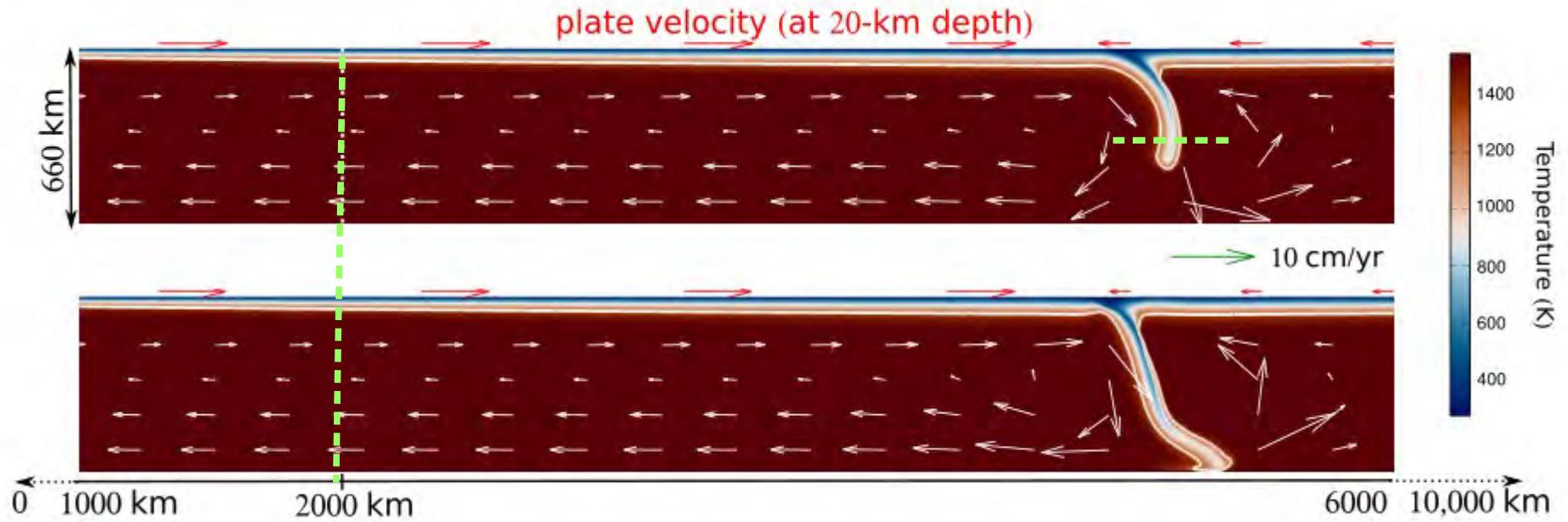
*Coltice et al., Science Adv., 2019*

# Complexifying models towards the real Earth : LAB in sinking slabs



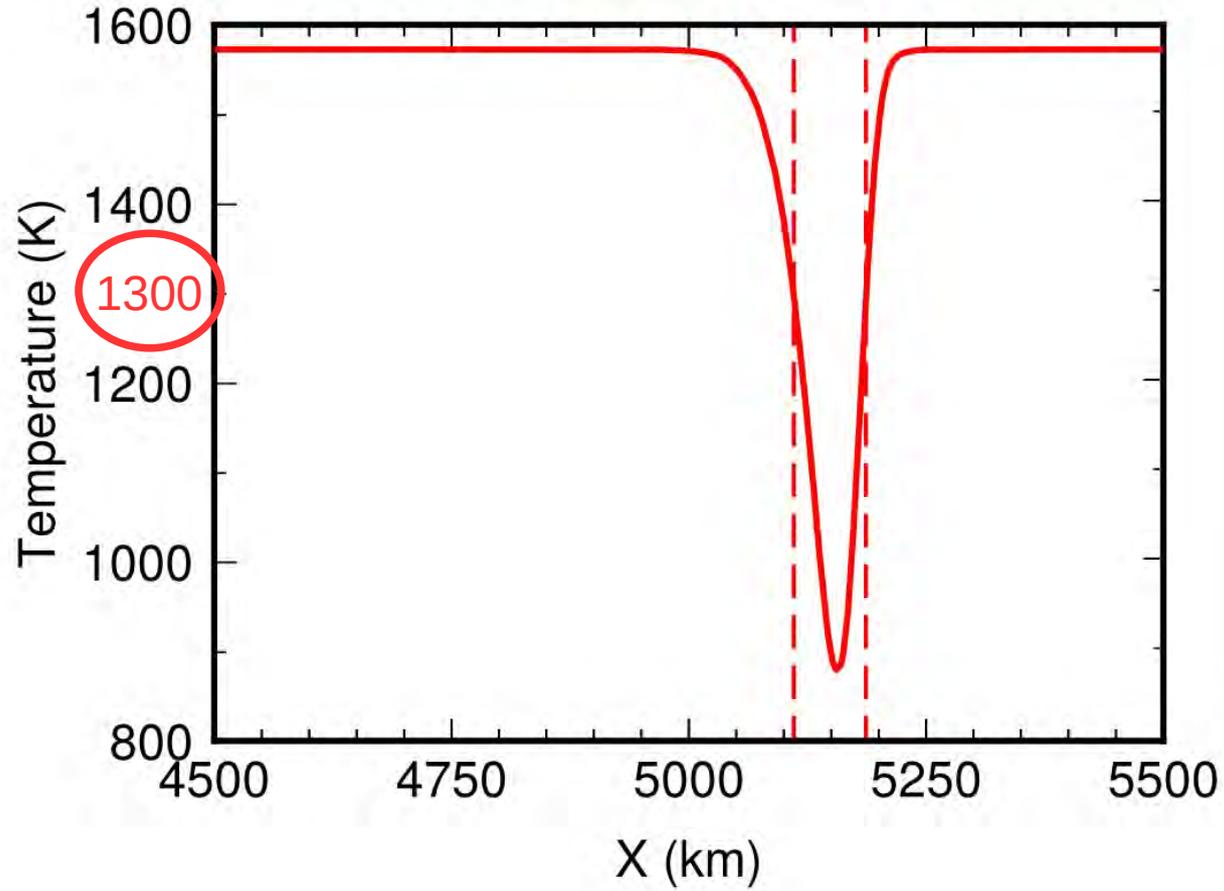
*Crameri et al., Tectonophysics, 2019*

# Transient asth. flow driven by subduction

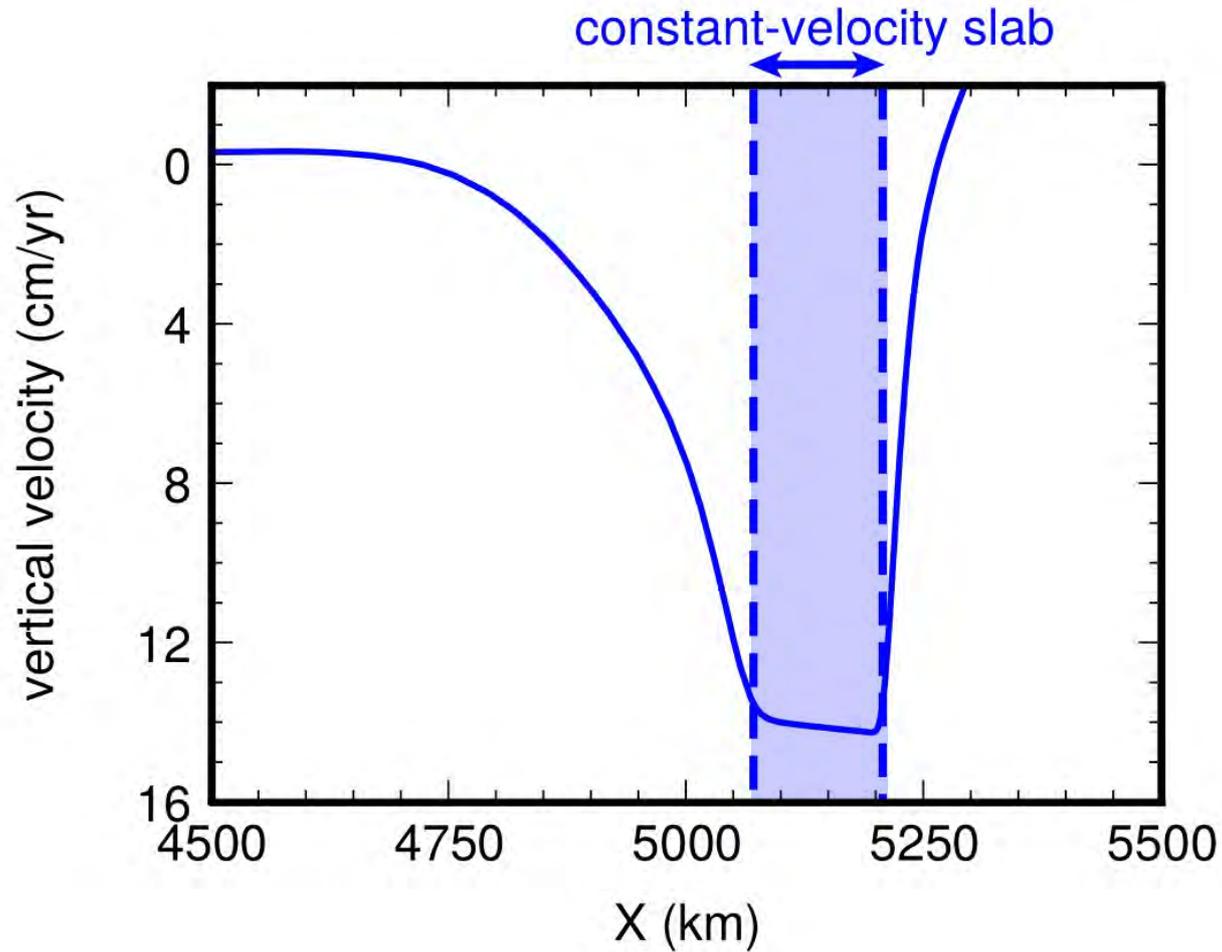


Profile across the fast-sinking  
plate in the upper mantle

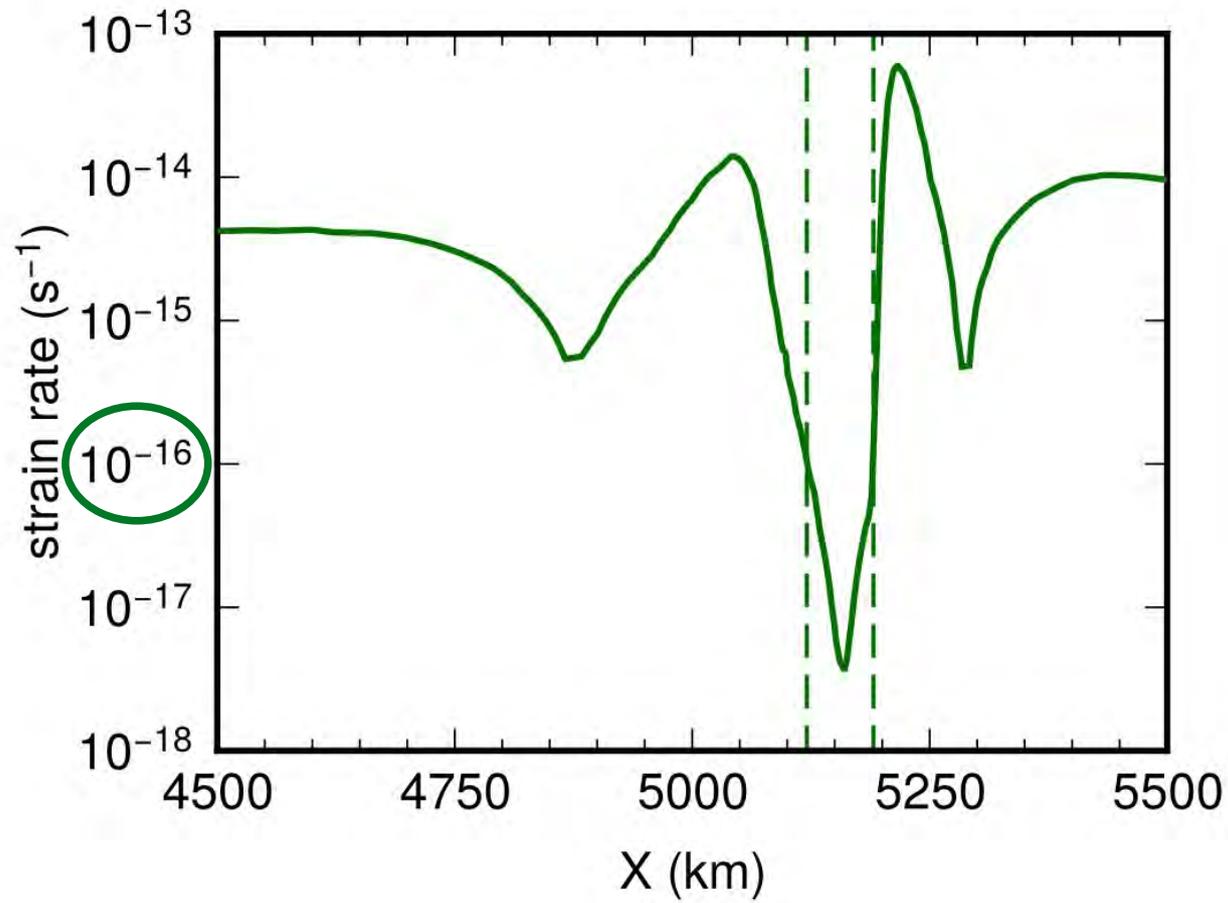
# Transient asth. flow driven by subduction



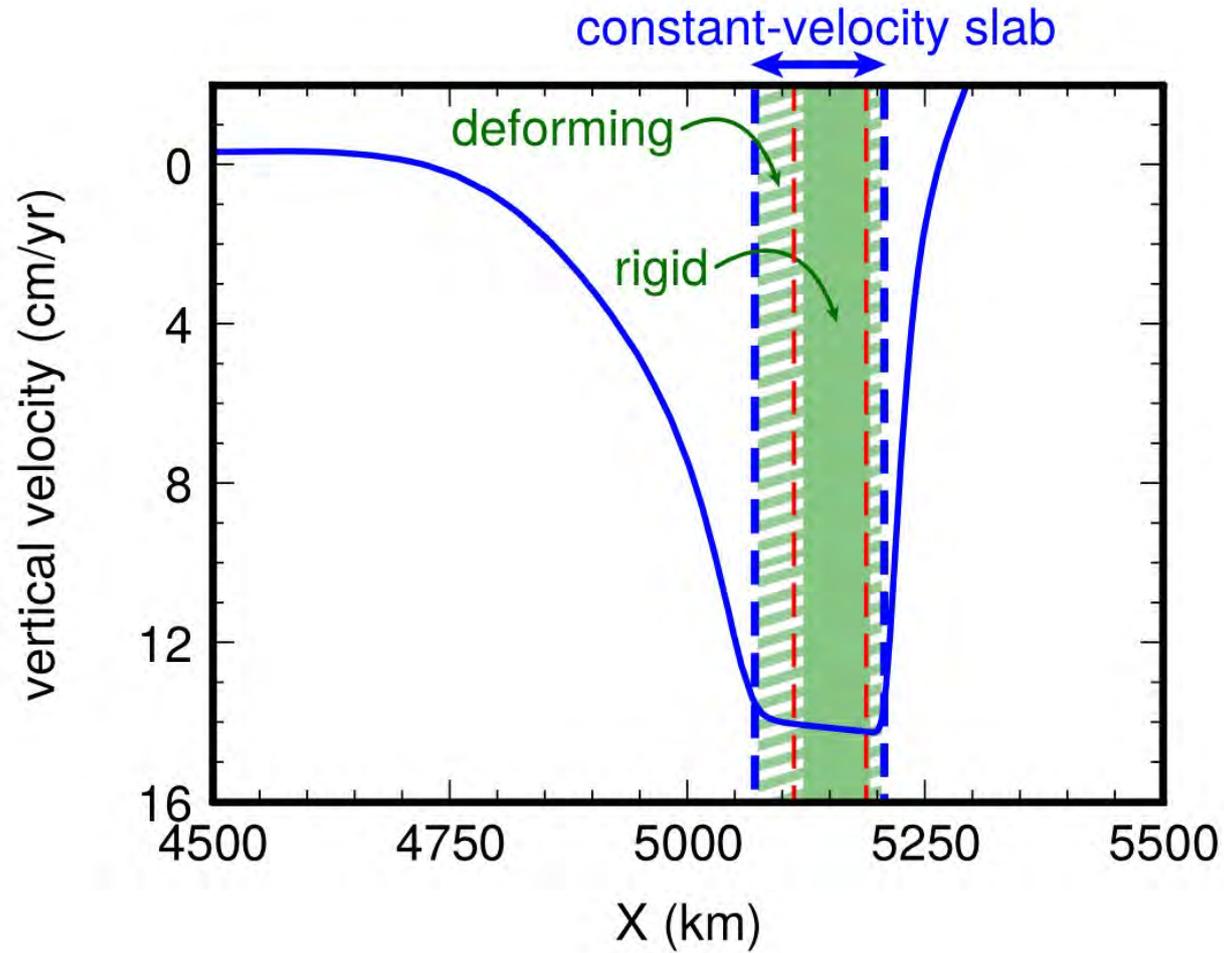
# Transient asth. flow driven by subduction



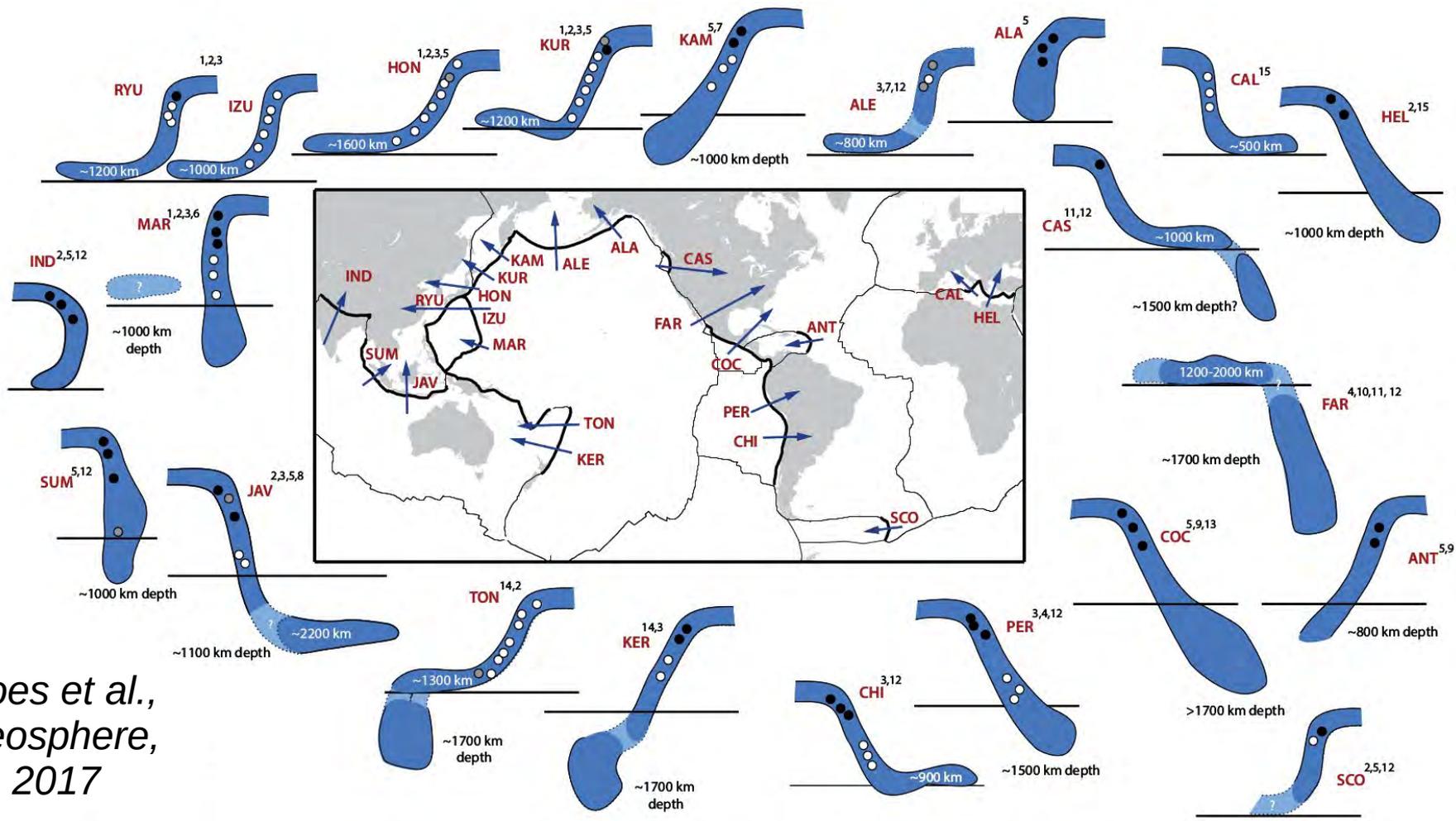
# Transient asth. flow driven by subduction



# Transient asth. flow driven by subduction



# Mass transfers between upper and lower mantle through asthenosphere dragged by cold slabs ?



Goes et al.,  
Geosphere,  
2017