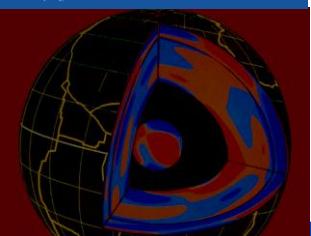




ISTerre

Institut des Sciences de la Terre

Seismic anisotropy of the inner core deduced from geodynamical and mineralogical models



Ph. Cardin, ISTerre, Grenoble, France

A. Lincot (ISTerre), S. Merkel (UMET-Lille)

R. Deguen (JHU, US), R. Lebenhson (LANL, US)

OSUG

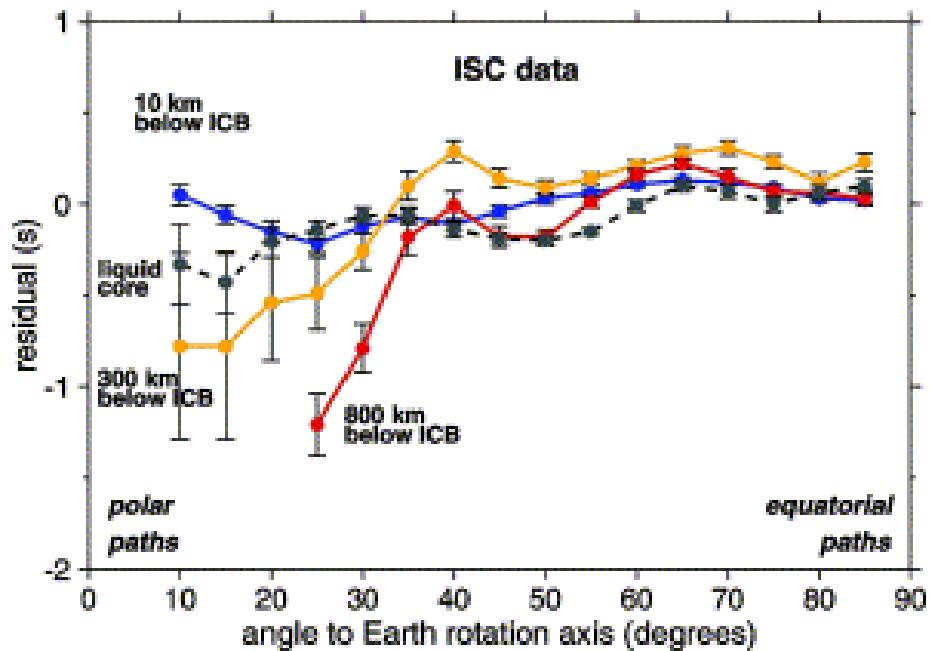


Observatoire des
Sciences de l'Univers
de Grenoble

Outline

1. Seismological observations, anisotropy
2. Review of proposed mechanism
3. Dynamical state of the inner core
4. Thermal convection in the inner core
5. Stratified inner core and equatorial growth
6. Perspectives

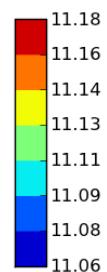
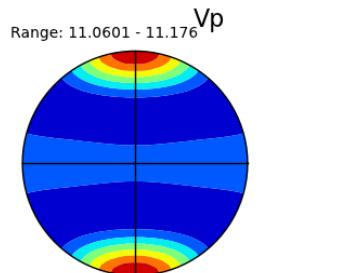
Seismological evidences?



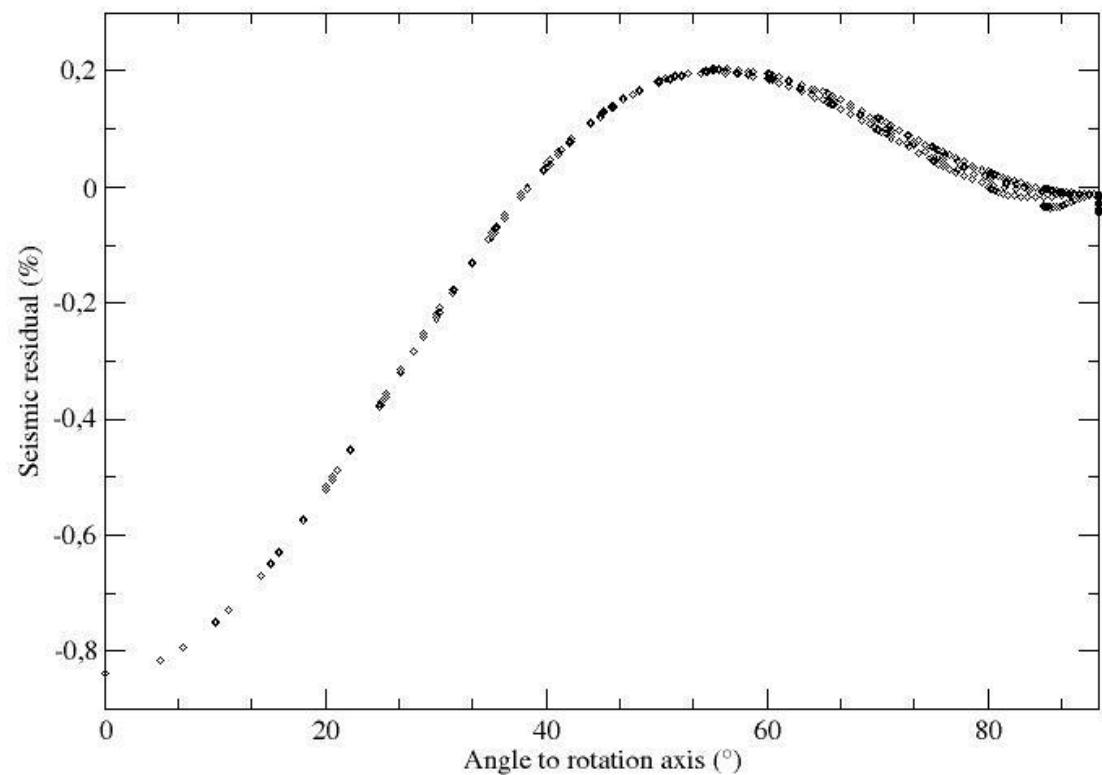
- 3% central anomaly
- Fast velocity almost NS ($\sim 10^\circ$)
- Isotropic superficial layer (100 to 400 km)
- With a hemispheric variation
- And perhaps an innermost innercore

Souriau, Garcia et Poupinet 2003

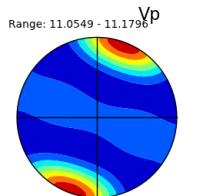
Direct model of anisotropy



Axisymmetrical inner core

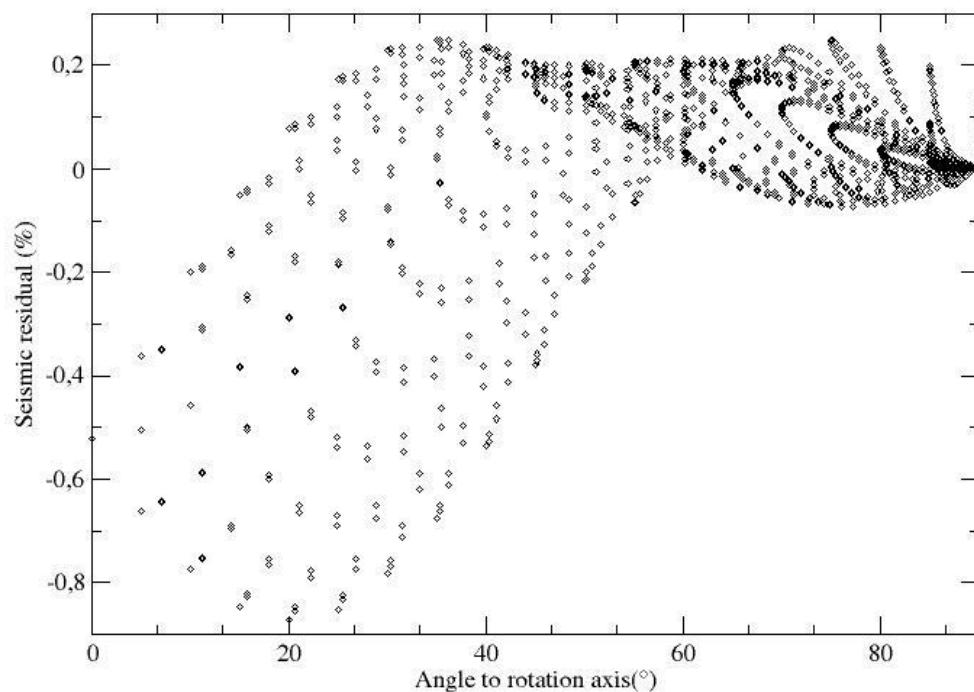


Model of anisotropy



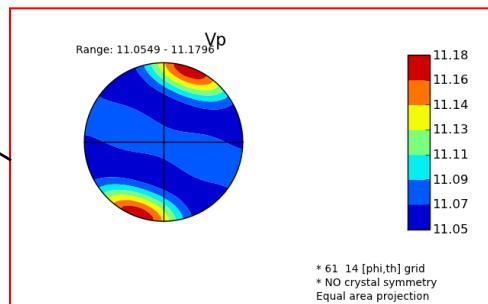
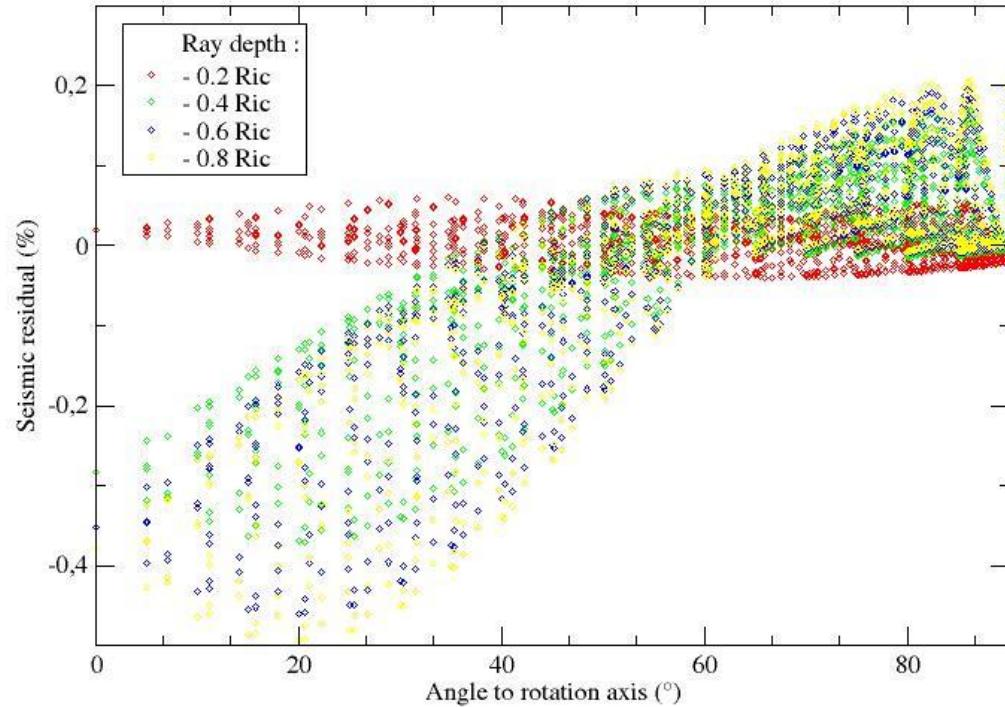
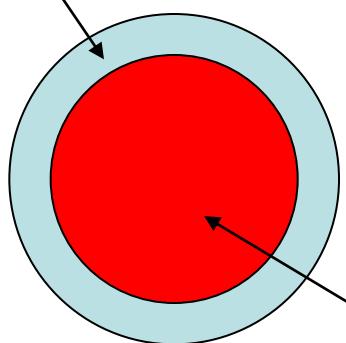
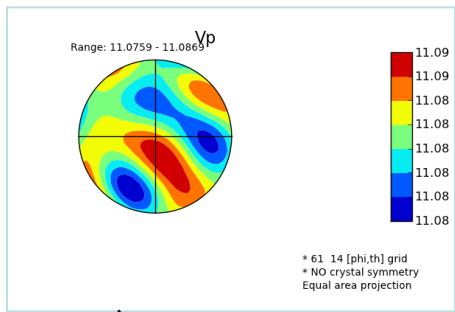
* 61 14 [phi,th]
* NO crystal symmetry
Equal area projection

Non-axisymmetrical and monocrystalline inner core



Anisotropy for a bilayered inner core

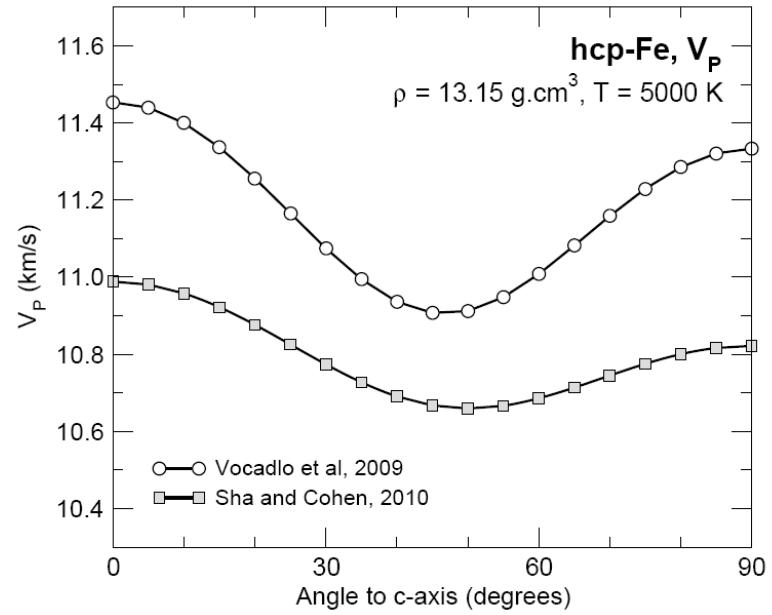
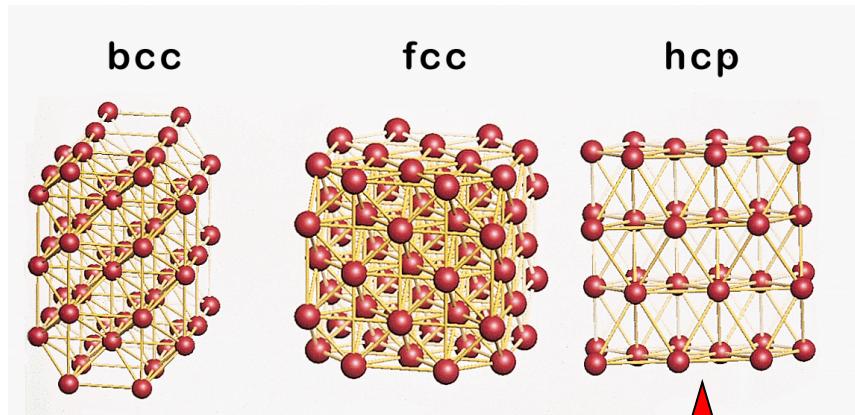
Bilayered and monocrystalline inner core



Outline

1. Seismological observations, anisotropy
2. Review of proposed mechanism
3. Dynamical state of the inner core
4. Thermal convection in the inner core
5. Stratified inner core and equatorial growth
6. Perspectives

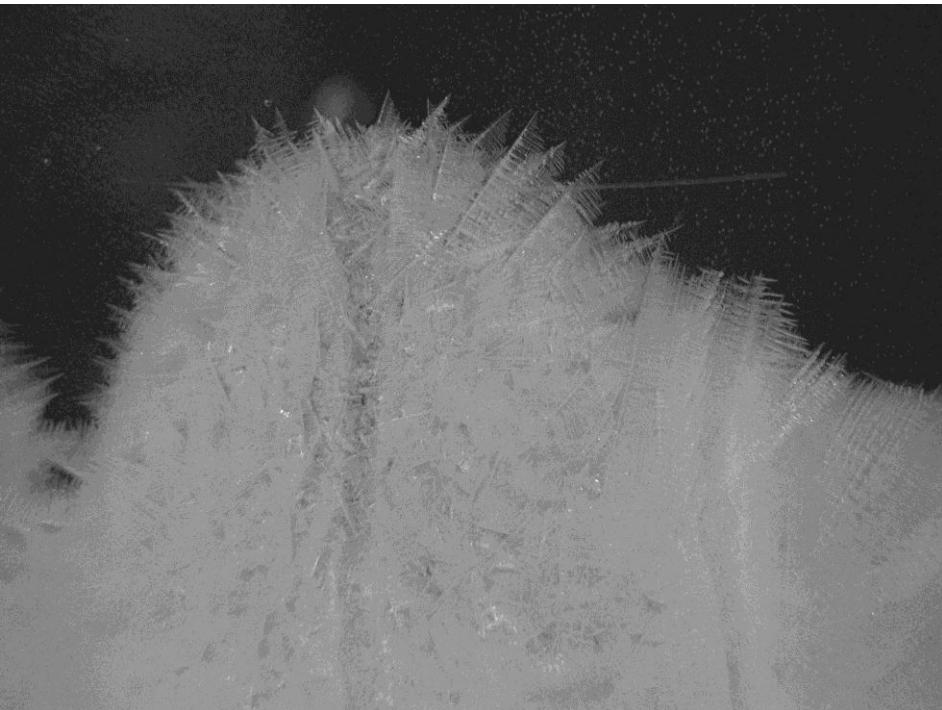
Anisotropic crystals of iron



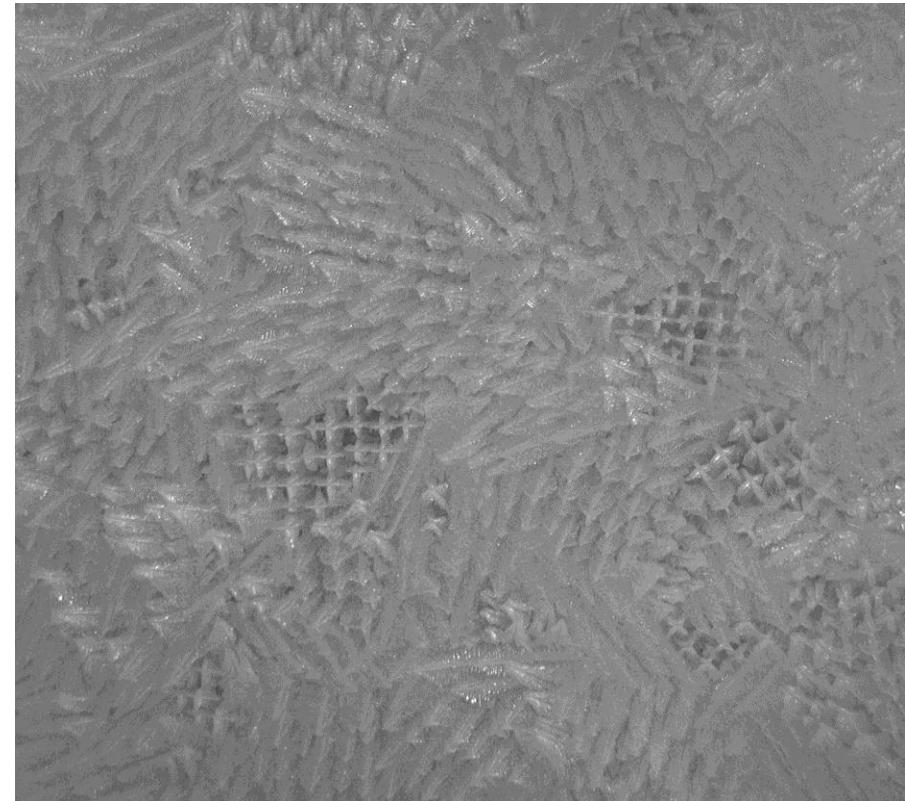
Experimental versus ab initio?

1-Texture by crystallisation

Dendrites

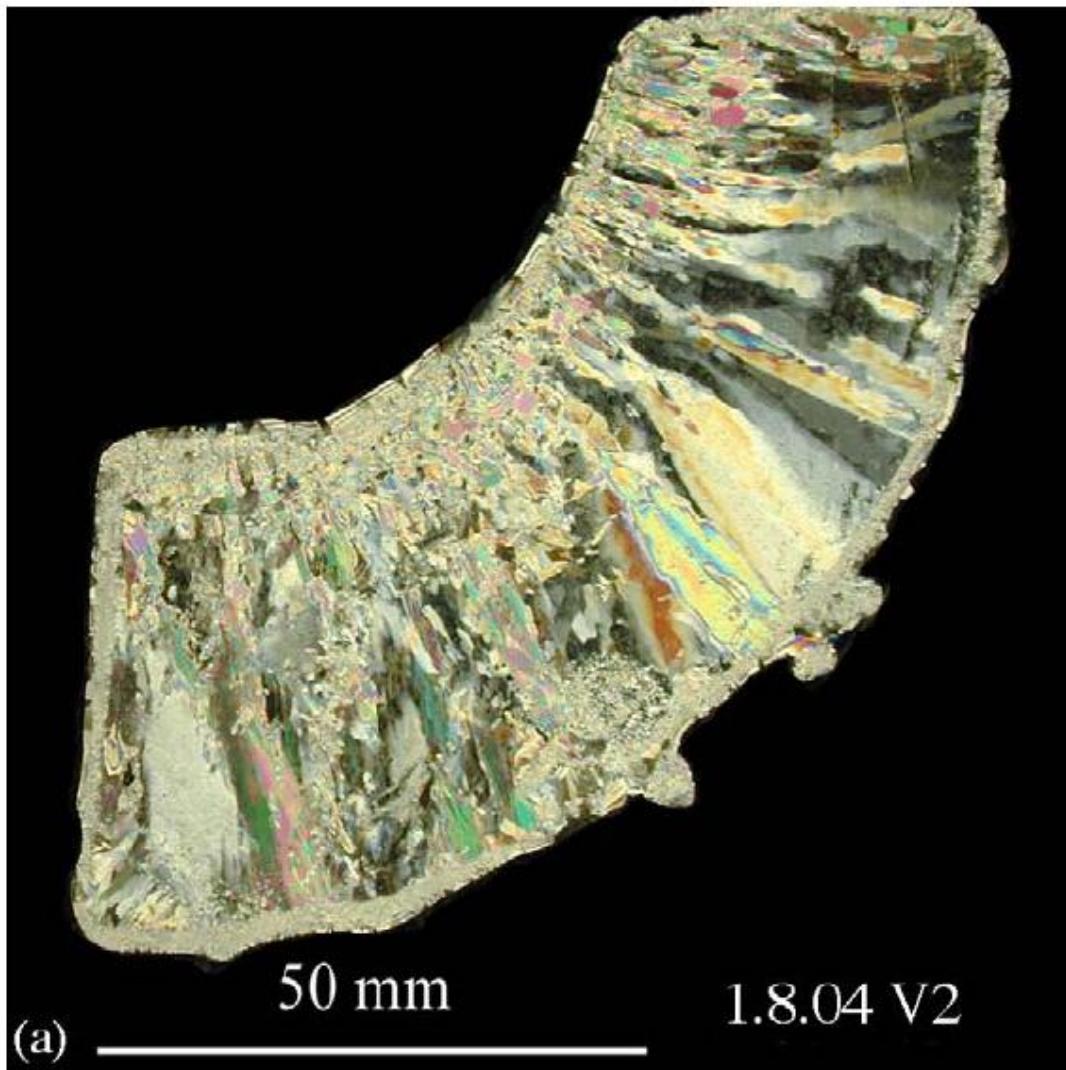


NH4Cl experiments



Deguen, 09

Solidification texturing



Salt water experiments

Radial symmetry
of the cristals

Effect of rotation,
magnetic fields,
heat flux variations

Bergman et al, 05

2- Texture by deformation

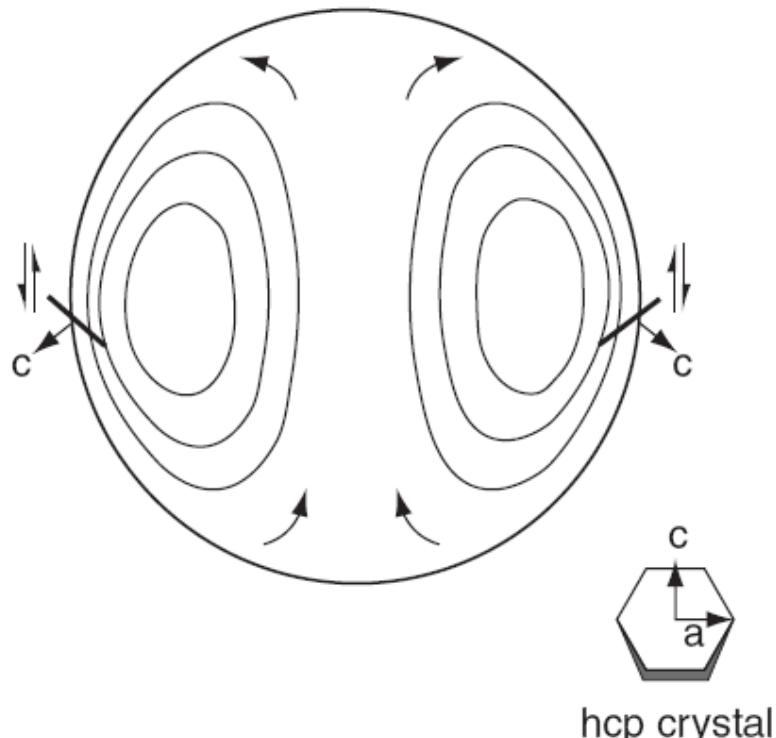
Plastic deformation

Or recrystallisation



Deformation mechanism (1)

- Thermal convection in the inner core



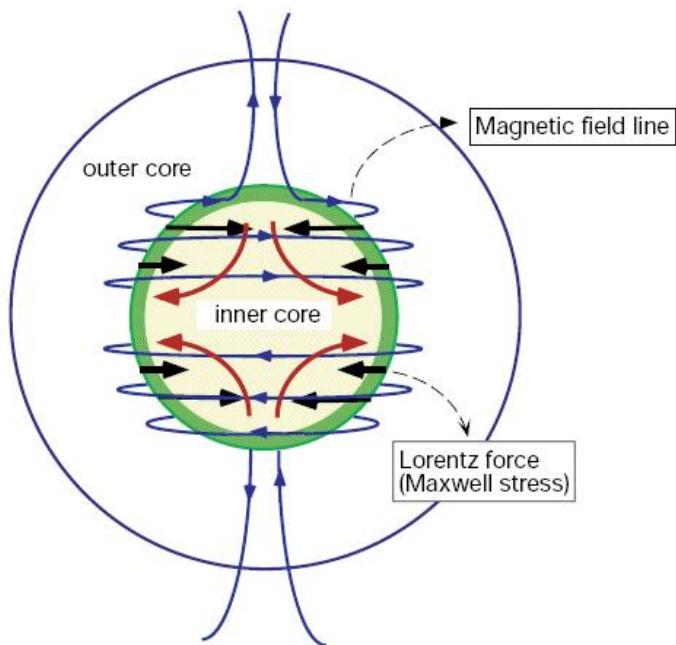
Weber et Machetel, 92

Wenk et al, 2000

Deformation mechanism (2)

- Maxwell stress

Poloidal motions



Karato, 1999

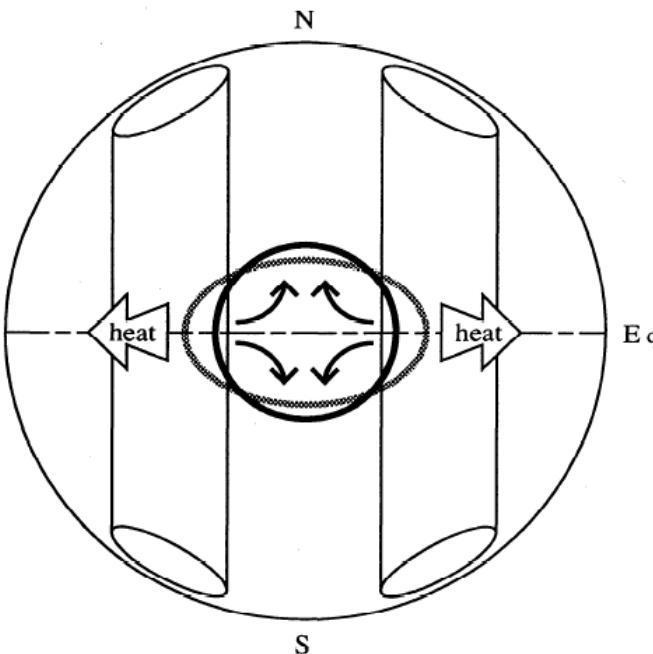
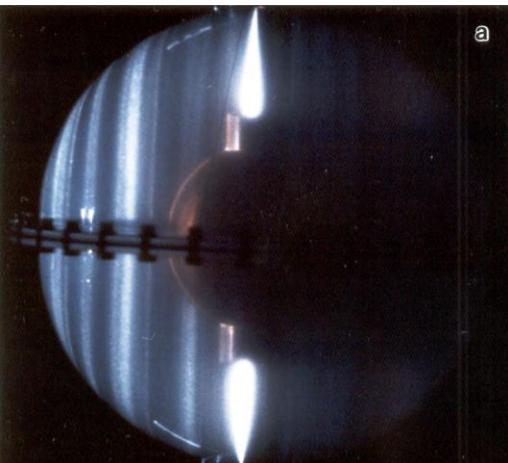
Toroidal motions

Differential motions
between the gravitational
torque at the ICB and the
Lorentz forces in the bulk

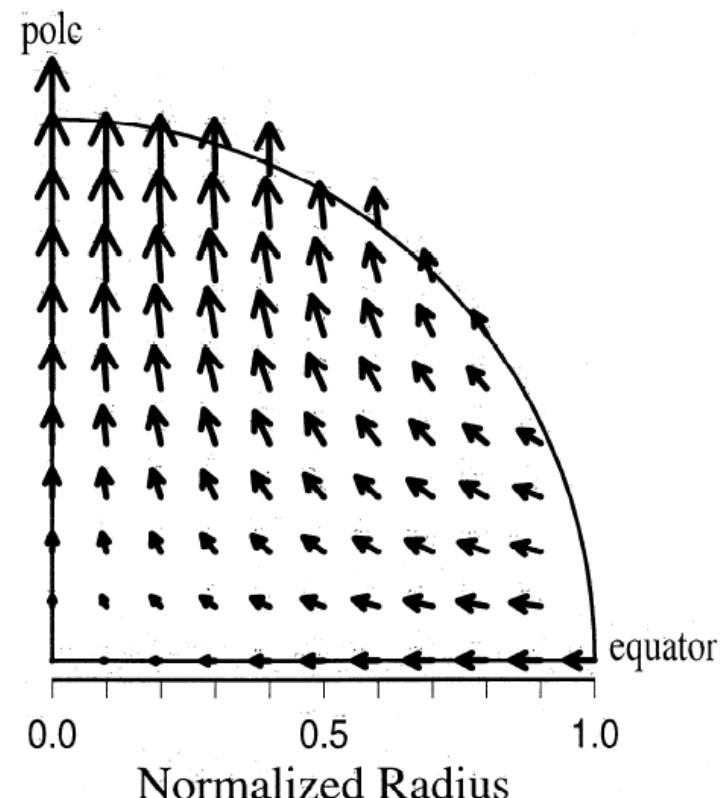
Buffet and Wenk, 2001

Deformation mechanism (3)

- Differential growth of the inner core



Yoshida et al, 1996



Review of proposed mechanism

summary

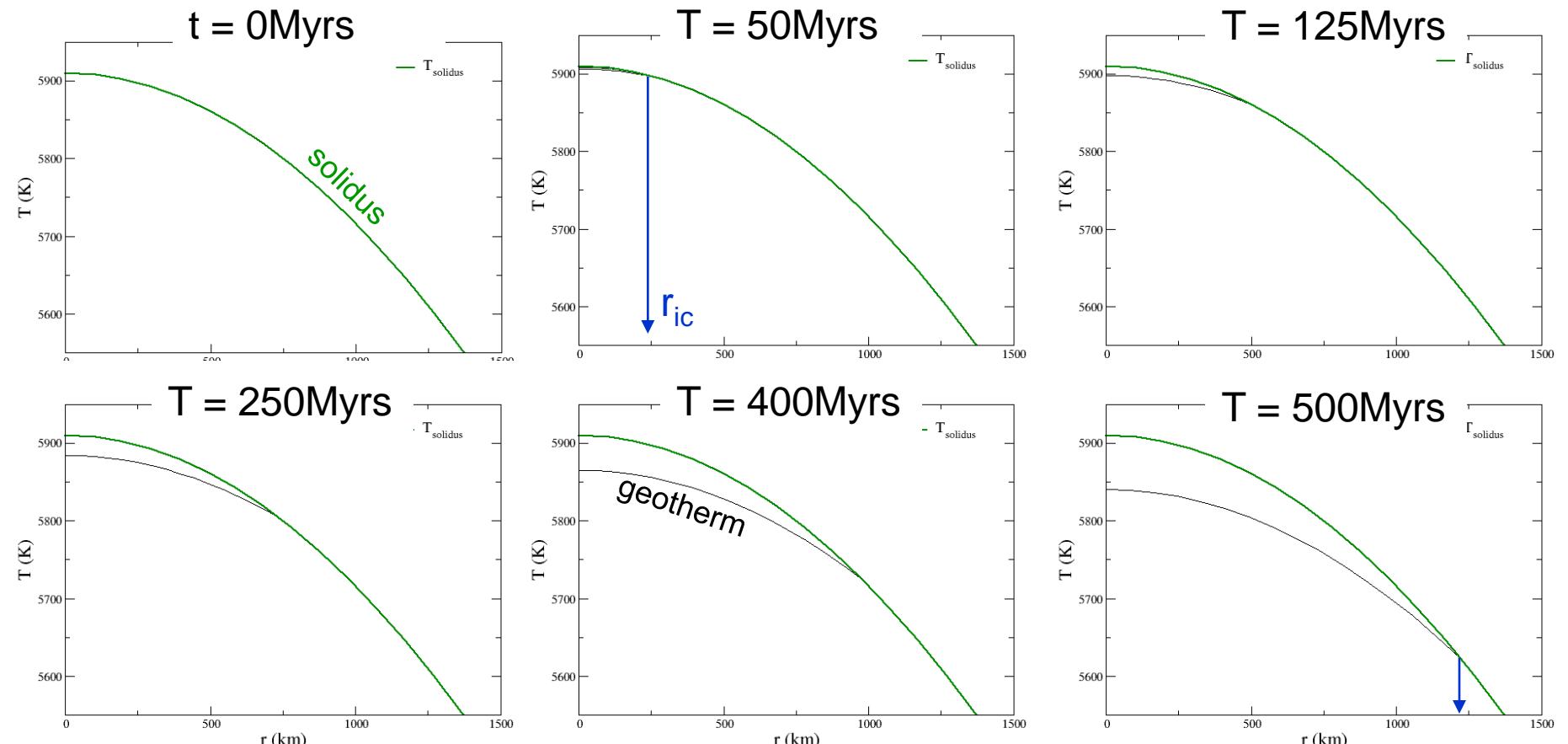
- HCP Fe crystal is anisotropic
- Texture by solidification (spherical)
- Texture by plastic deformation
 - thermal convection
 - Maxwell stress
 - differential growth

Outline

1. Seismological observations, anisotropy
2. Review of proposed mechanism
3. **Dynamical state of the inner core**
4. Thermal convection in the inner core
5. Stratified inner core and equatorial growth
6. Perspectives

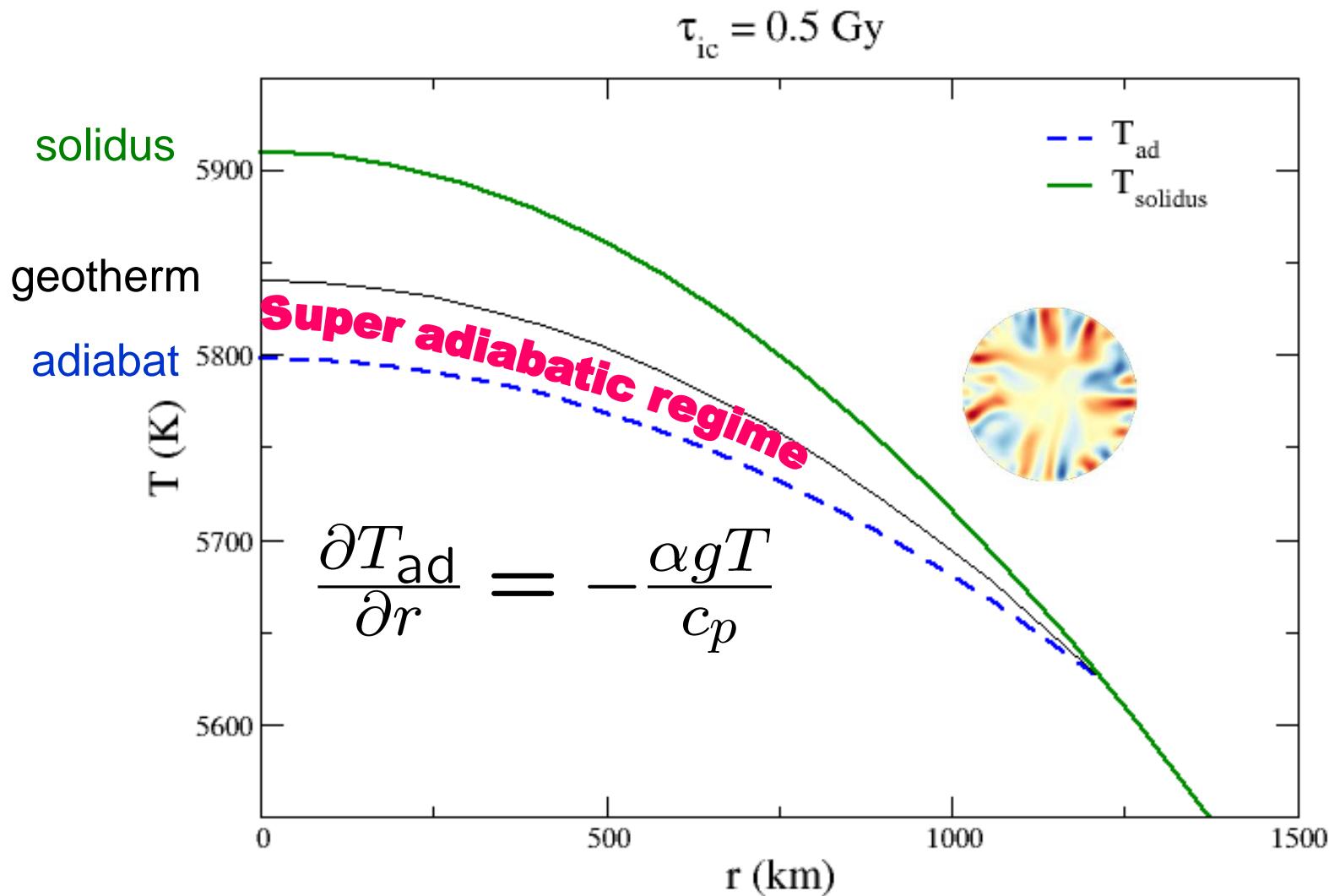
Time evolution of thermal state of the inner core

- Take an inner core of 500Myrs

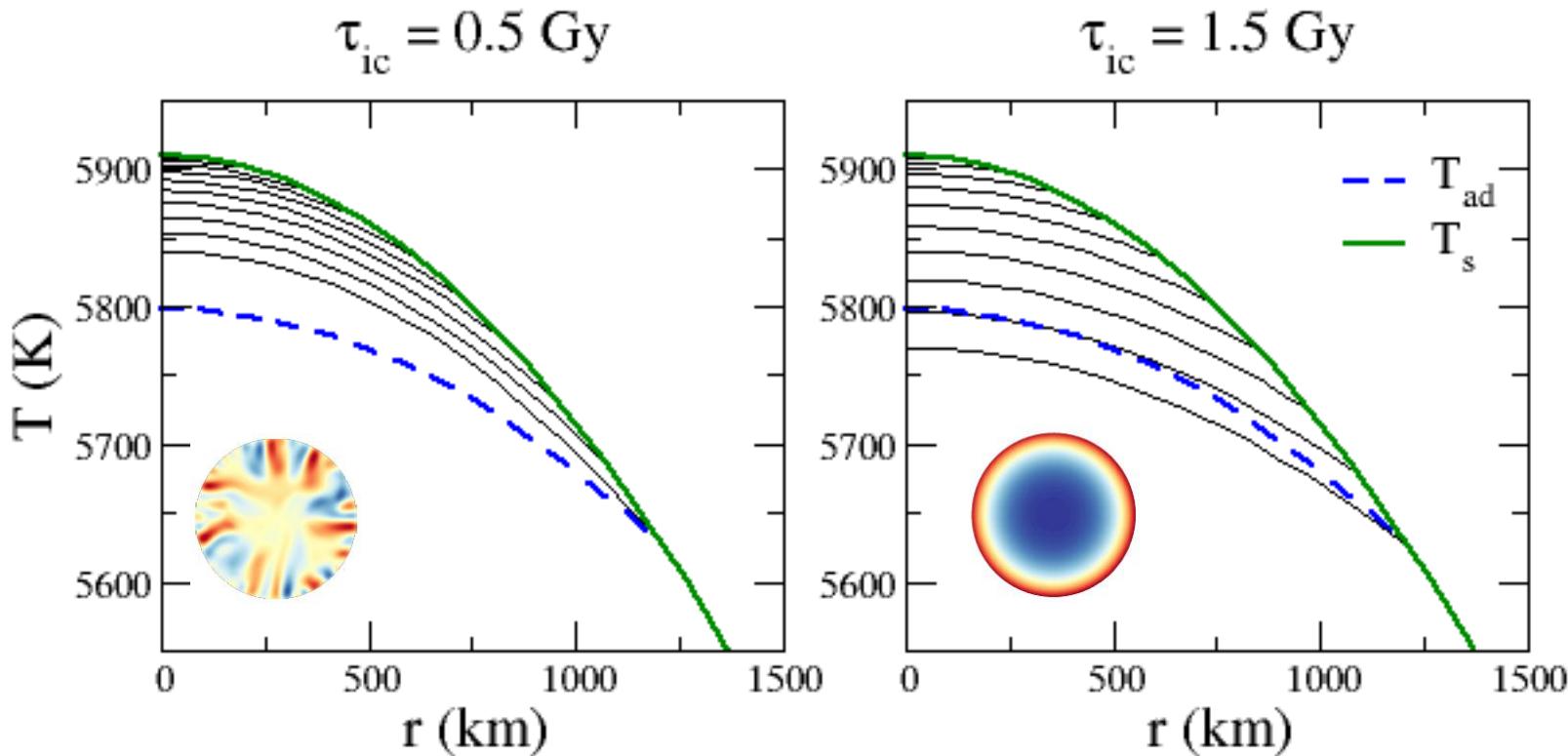


$$r_{\text{ic}} = 1220 \text{ km}$$

Dynamical thermal state ?



Inner core age



Inner core cooling (age) is the key parameter to predict the dynamical state.

CMB heat flux

Thermal conductivity

Dynamical state of the inner core

- Age of IC is the key parameter:

$$\tau_{ic} < \tau_\kappa \left(\frac{dT_S}{dT_{ad}} - 1 \right) = \frac{r_{ic}^2}{3\kappa} \left(1 - \frac{1}{3\gamma} \right) \left(\frac{1}{\gamma} + \alpha T \right)$$

$$\boxed{\tau_{ic}^{crit} = 0.9 \pm 0.6 Gy}$$

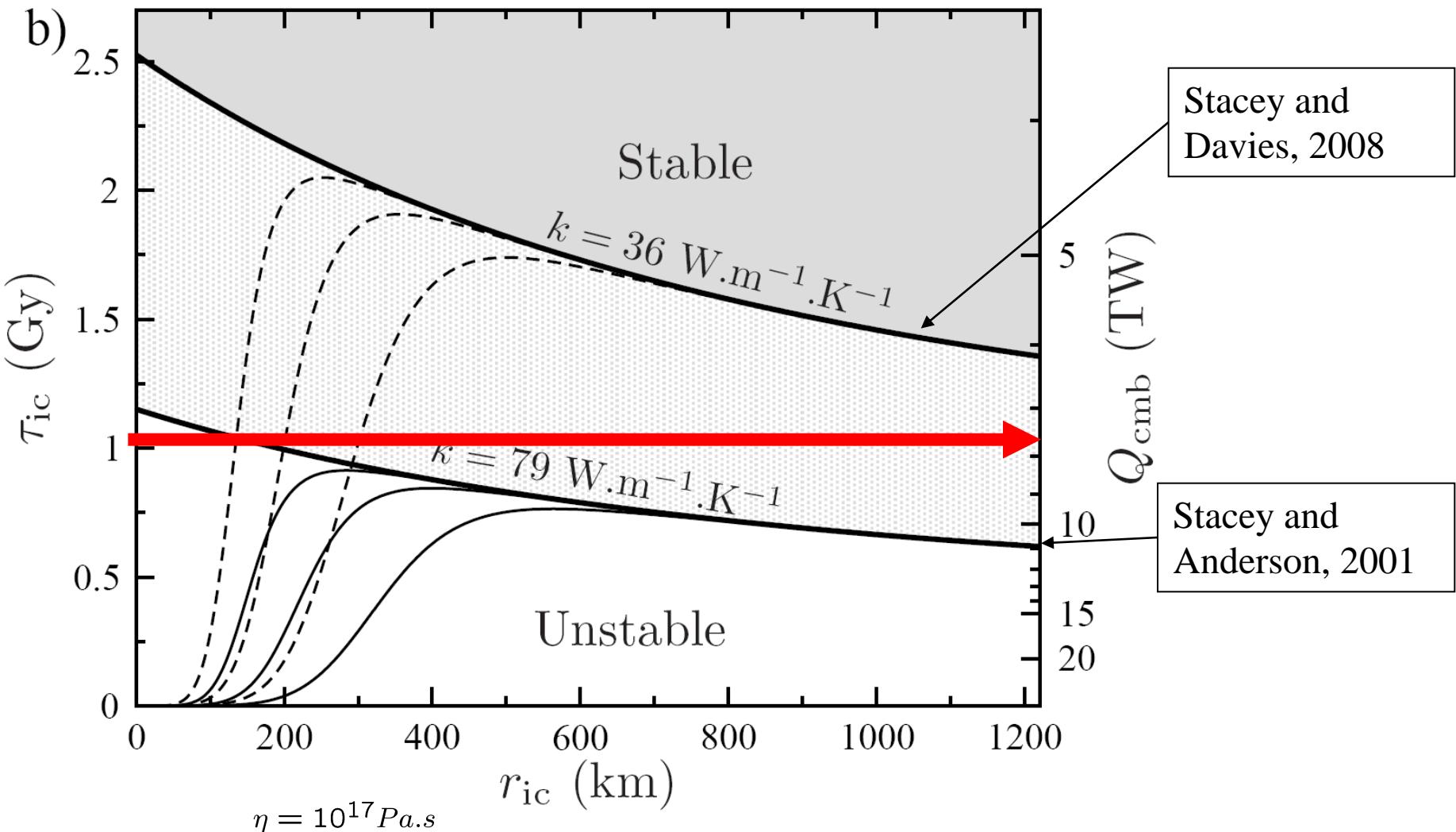
- Controled mainly by the CMB heat flux

$$\tau_{ic} = \frac{E_{tot}}{Q_{cmb}} = \frac{\int_0^{r_{ic}} P_c + P_g + P_L dr}{Q_{cmb}} = \frac{29 \pm 18 \ 10^{28} \ J}{Q_{cmb}}$$

$$\boxed{\tau_{ic} = 0.6Gy(12TW) - 2.3Gy(4TW)}$$

Labrosse et al, 2003

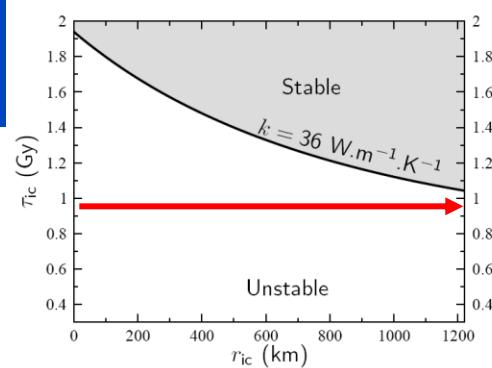
Adiabatic state of the inner core



Outline

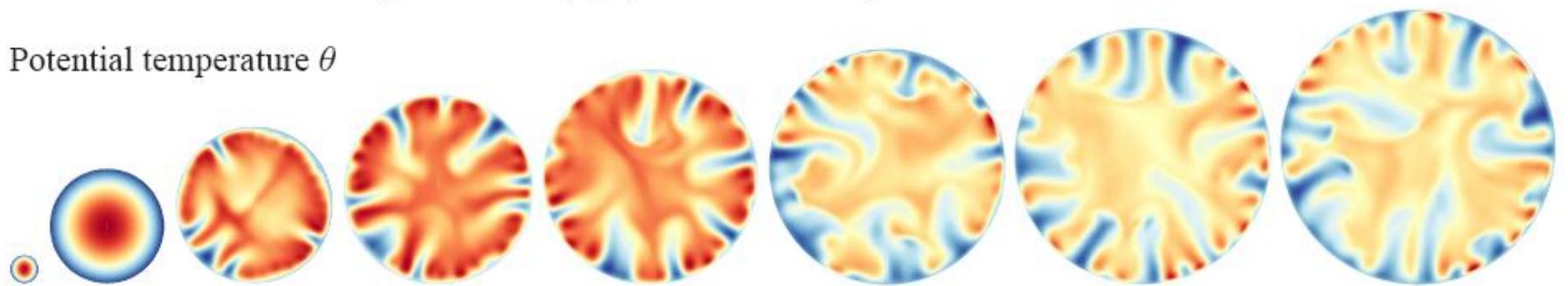
1. Seismological observations, anisotropy
2. Review of proposed mechanism
3. Dynamical state of the inner core
4. Thermal convection in the inner core
5. Stratified inner core and equatorial growth
6. Perspectives

Thermal convection

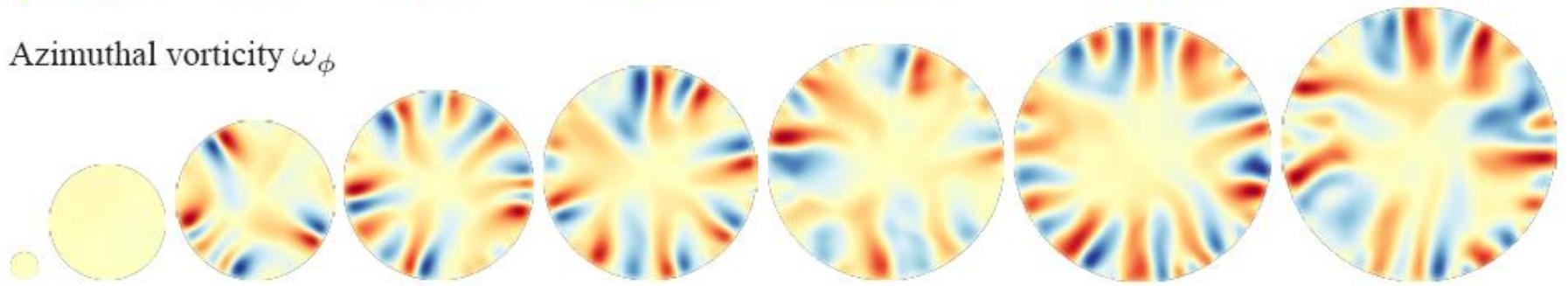


a. $\tau_{ic} = 0.9$ Gy , $\eta = 10^{18}$ Pa.s, $k = 36$ W.m $^{-1}$.K $^{-1}$.

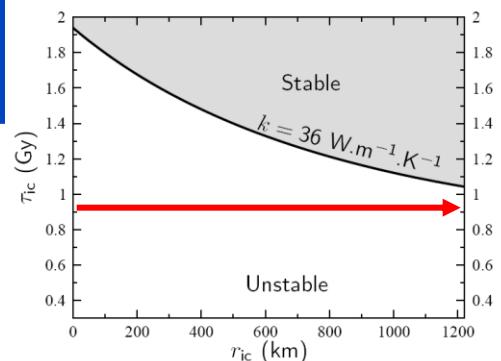
Potential temperature θ



Azimuthal vorticity ω_ϕ

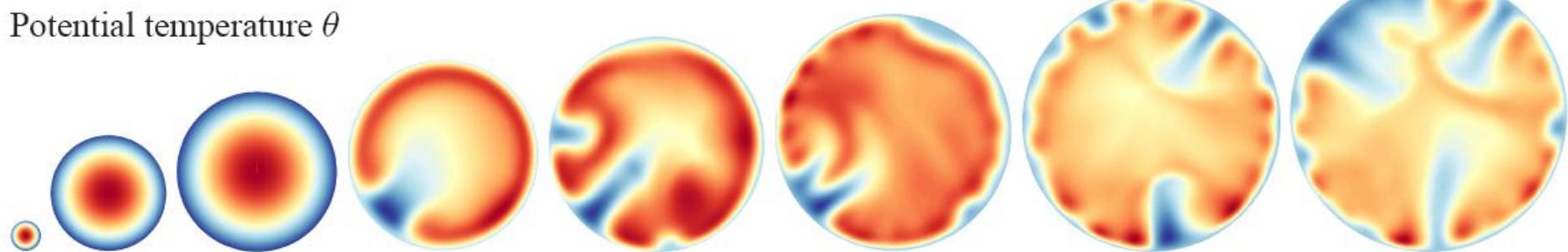


Thermal convection

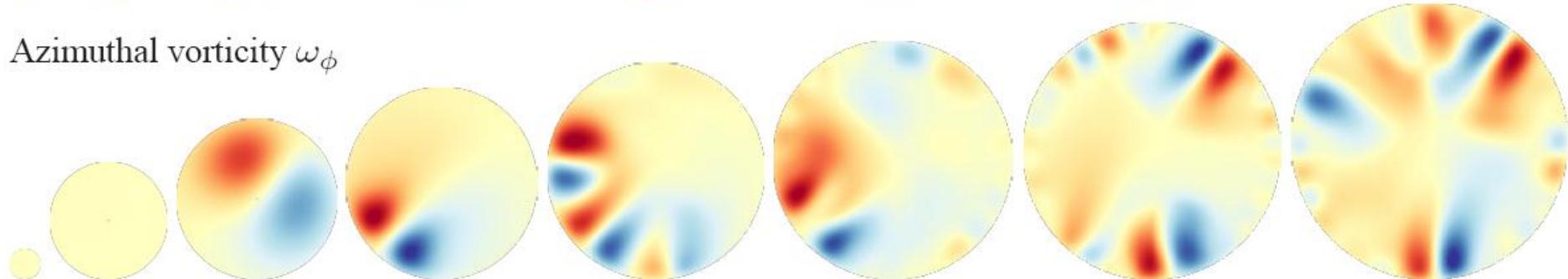


b. $\tau_{ic} = 0.9$ Gy , $\eta = 10^{19}$ Pa.s, $k = 36$ W.m $^{-1}$.K $^{-1}$.

Potential temperature θ



Azimuthal vorticity ω_ϕ



Scaling laws

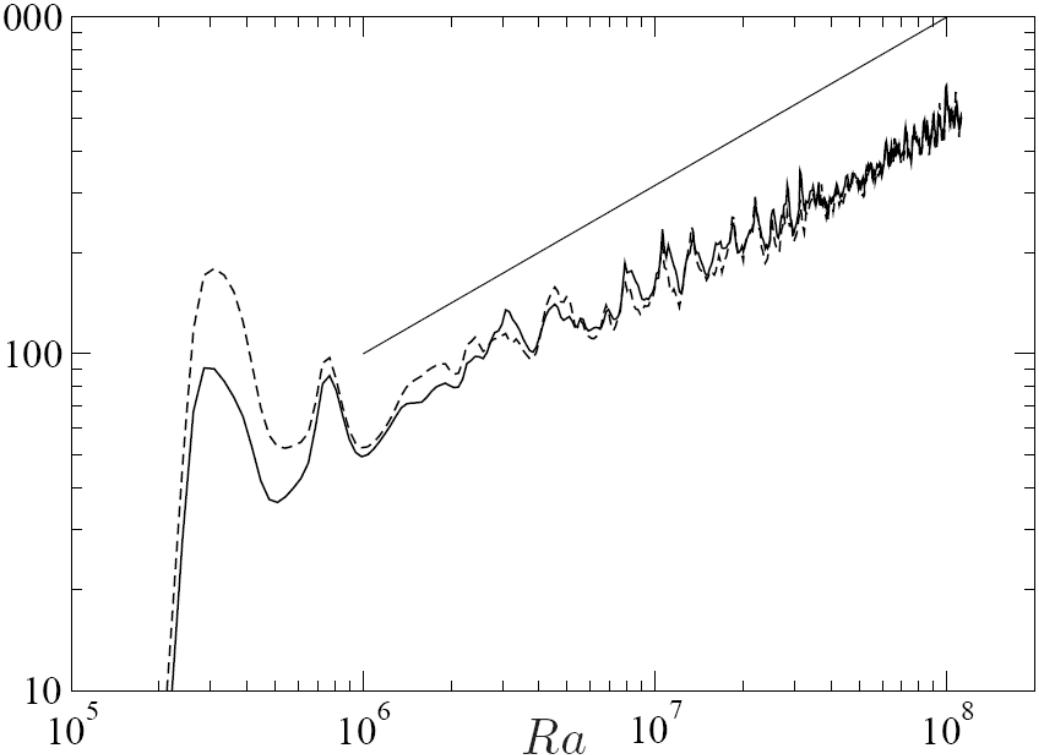
$$u_{rms} \approx 0.04 \frac{\kappa}{r_{ic}} Ra^{0.51}$$

$$u_{rms} \propto Ra^{1/2}$$

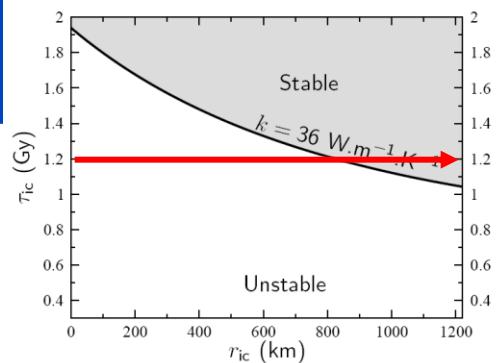
Parmentier and Sotin, 2000

$$\delta\theta \approx 1.4 \frac{Sr_{ic}}{6\kappa} Ra^{0.27}$$

$$\delta\theta \approx 1K$$

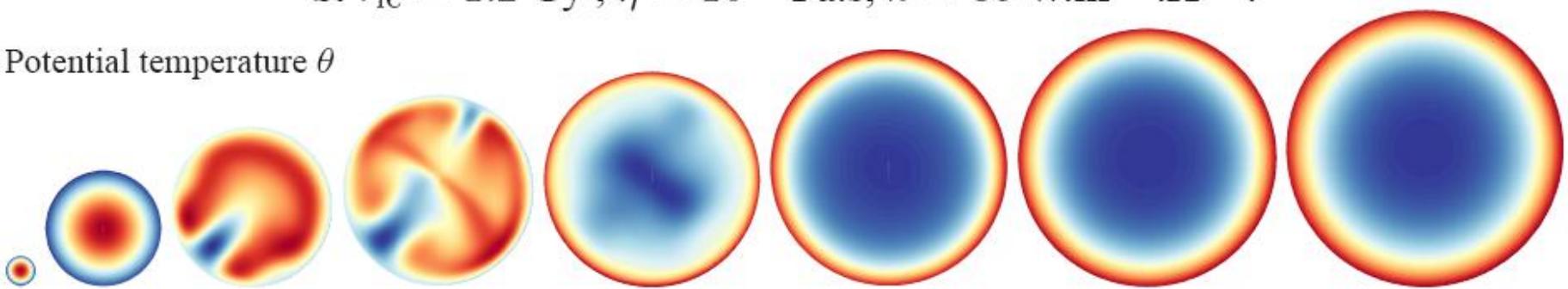


A convective window

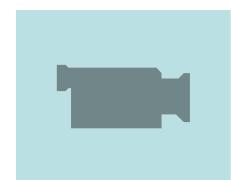
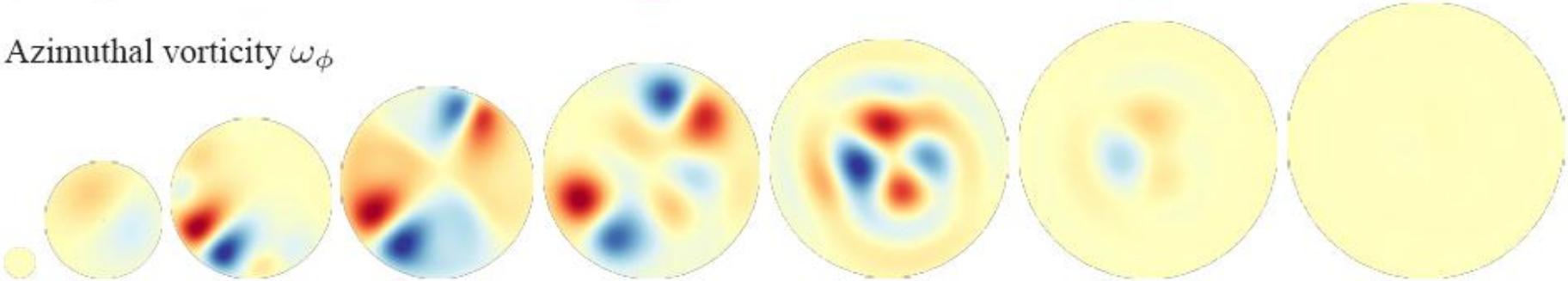


b. $\tau_{ic} = 1.2$ Gy , $\eta = 10^{18}$ Pa.s, $k = 36$ W.m⁻¹.K⁻¹.

Potential temperature θ

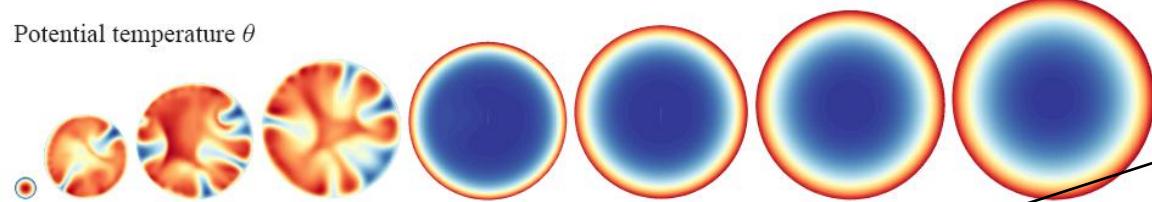


Azimuthal vorticity ω_ϕ

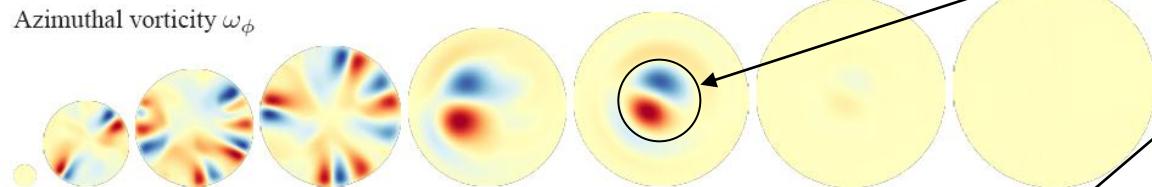


a. $\tau_{ic} = 1.2 \text{ Gy}$, $\eta = 10^{17} \text{ Pa.s}$, $k = 36 \text{ W.m}^{-1}.\text{K}^{-1}$.

Potential temperature θ



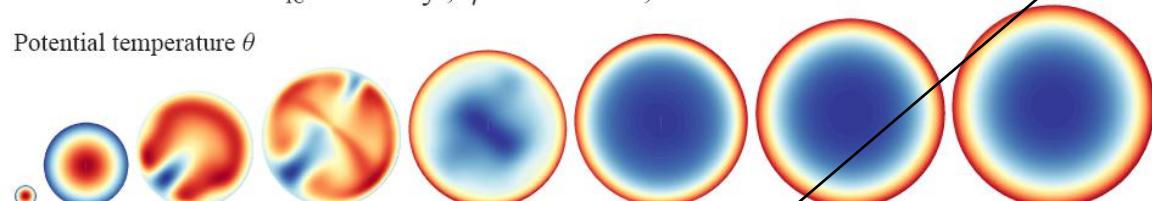
Azimuthal vorticity ω_ϕ



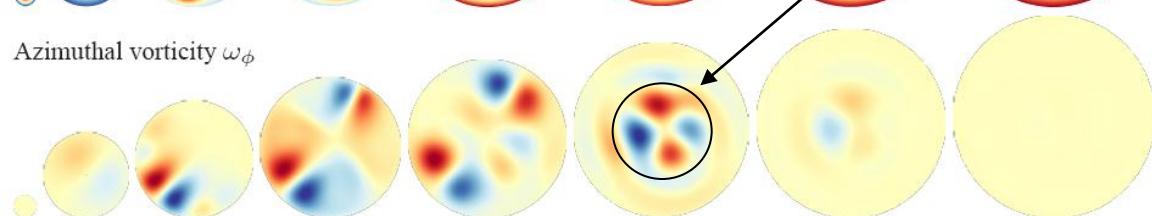
Innermost
Inner Core ?

b. $\tau_{ic} = 1.2 \text{ Gy}$, $\eta = 10^{18} \text{ Pa.s}$, $k = 36 \text{ W.m}^{-1}.\text{K}^{-1}$.

Potential temperature θ



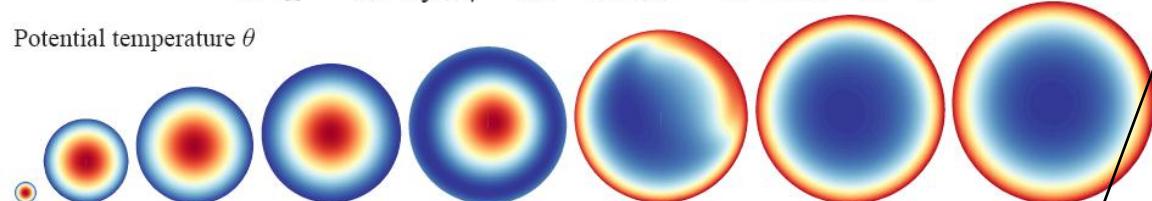
Azimuthal vorticity ω_ϕ



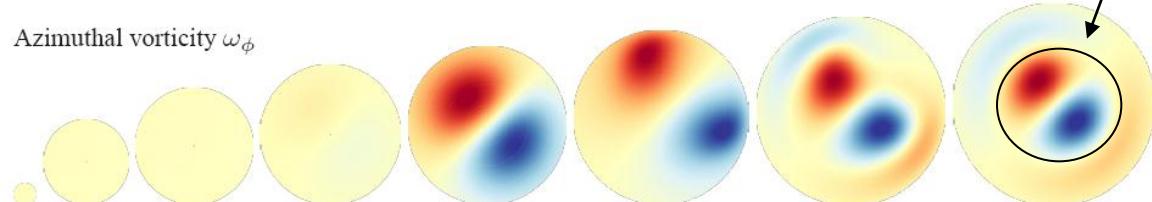
Higher viscosity

c. $\tau_{ic} = 1.2 \text{ Gy}$, $\eta = 10^{19} \text{ Pa.s}$, $k = 36 \text{ W.m}^{-1}.\text{K}^{-1}$.

Potential temperature θ

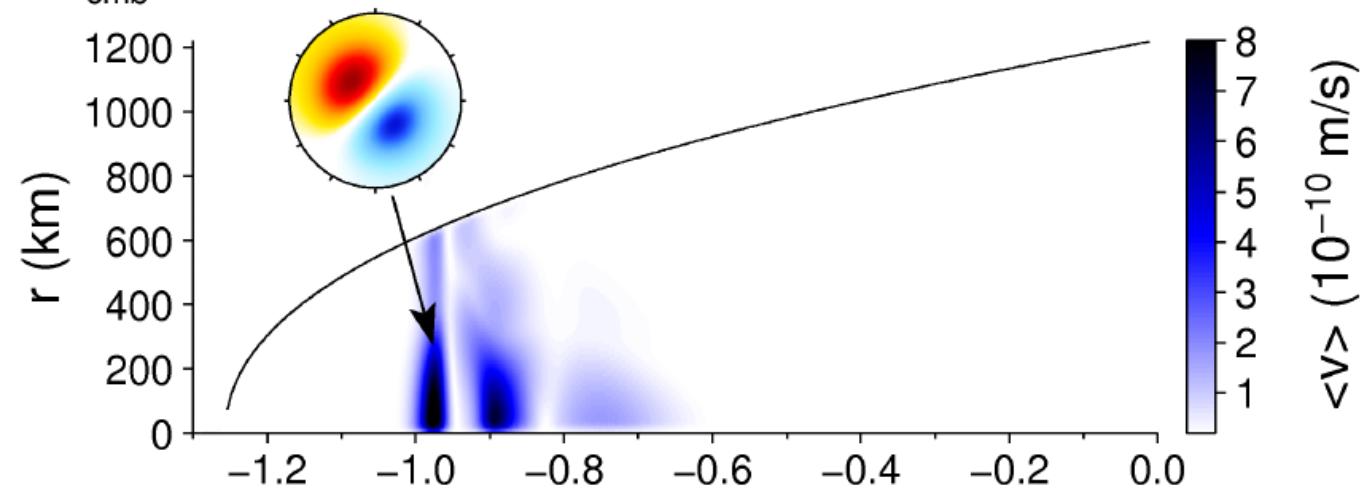


Azimuthal vorticity ω_ϕ

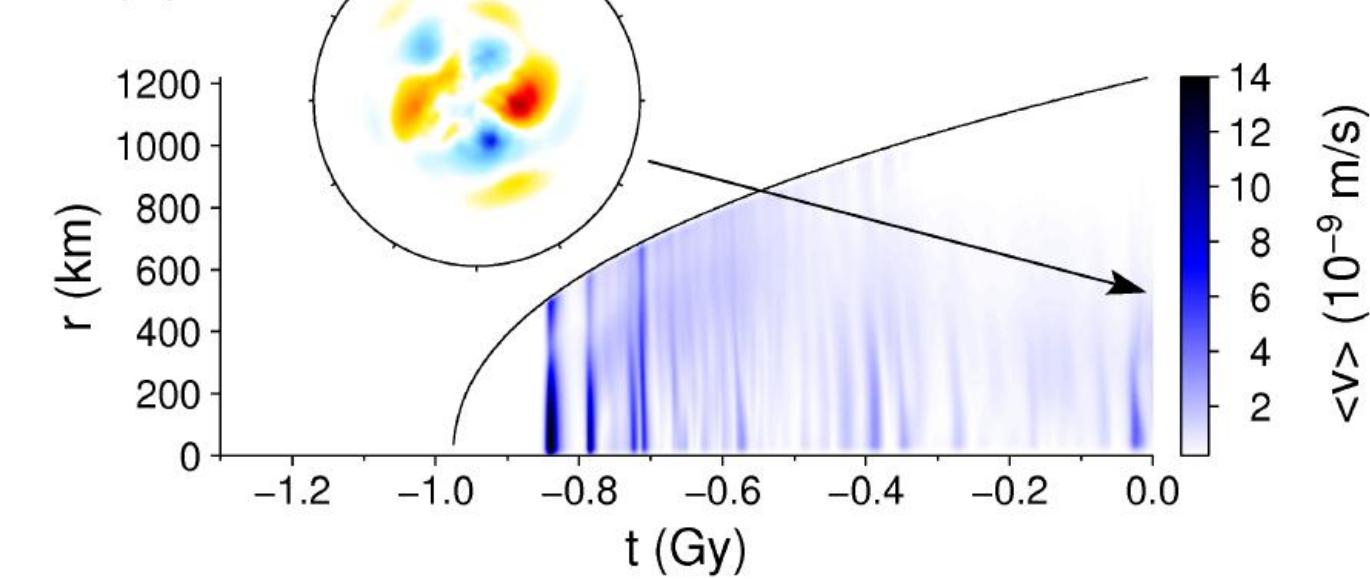


IIC: a thermal signature?

A $Q_{\text{cmb}}=7 \text{ TW}$

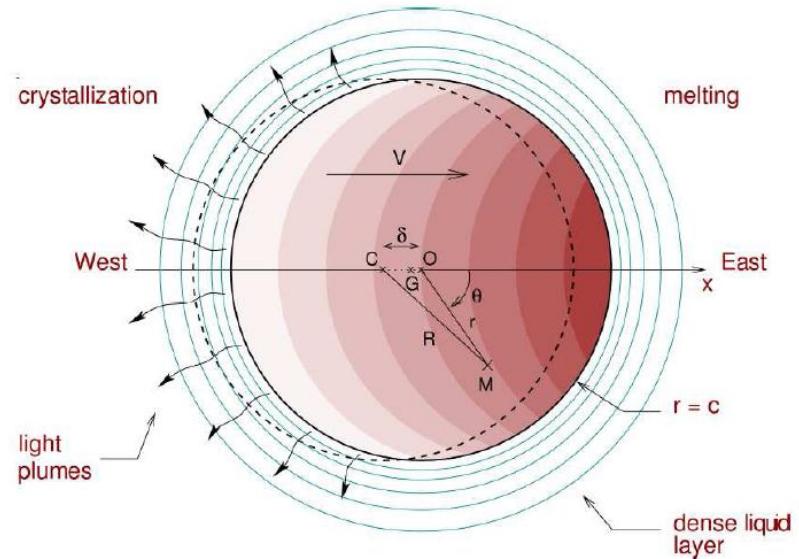


B $Q_{\text{cmb}}=9 \text{ TW}$



Translation of the inner core

- A rigid inner core
- EW perturbation of T, change the density field and the center of mass.
- Allow quick fusion and solidification.
- Misfit between center of mass and “center” of T.
- Archimedean force translate the inner core.



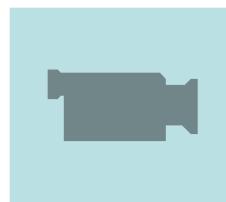
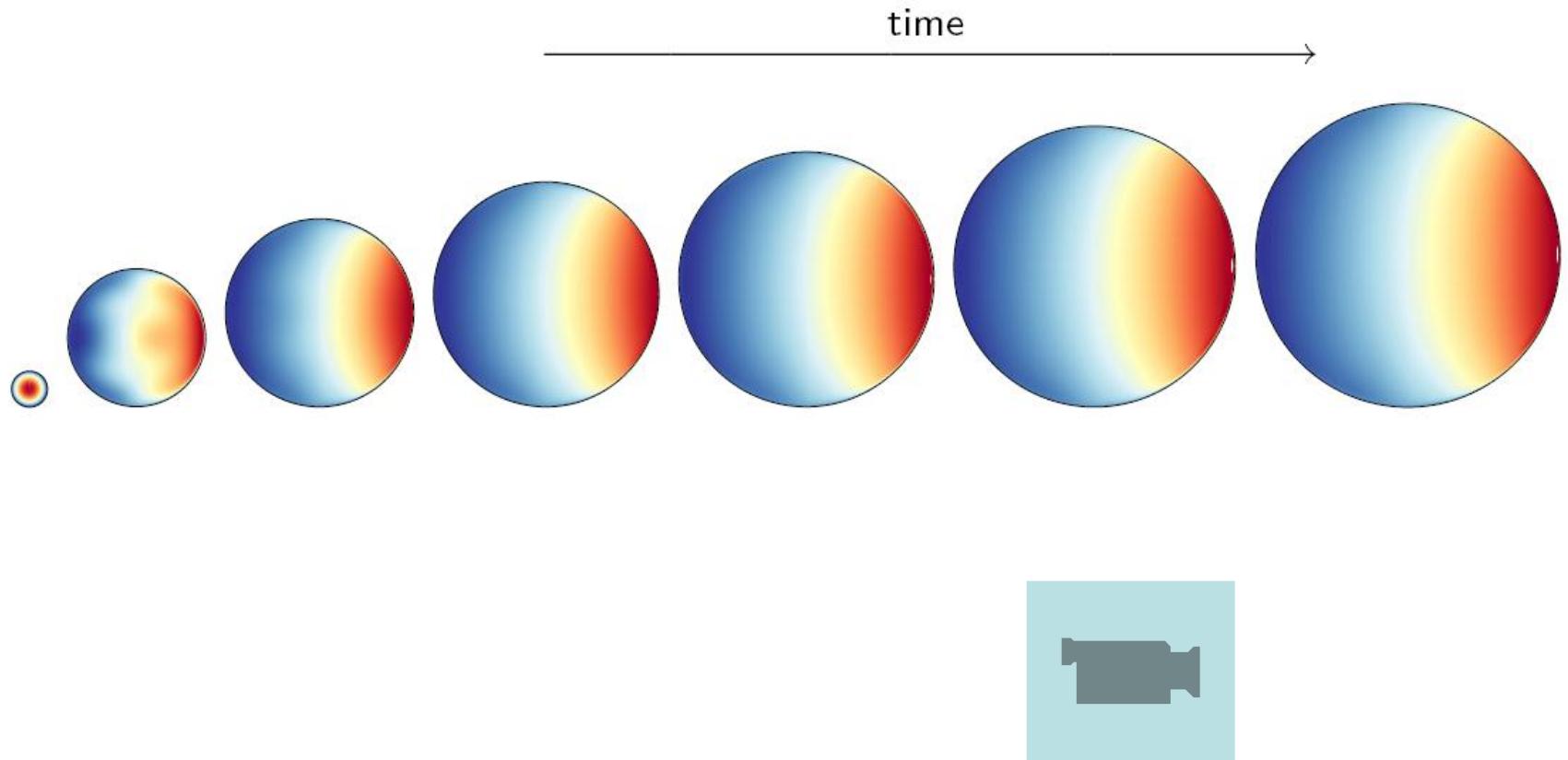
Alboussière et al, 2010

Monnereau et al, 2010

translation

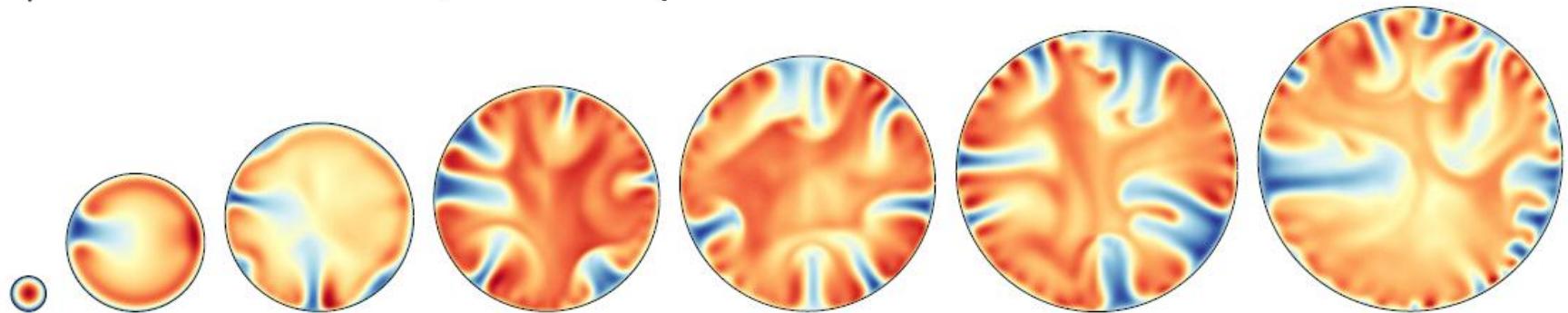
$\tau_{ic} = 0.9 \text{ Gy}$, $k = 36 \text{ W.m}^{-1}.K^{-1}$, potential temperature :

- ▶ $\eta = 10^{19} \text{ Pa.s}$, translation regime :

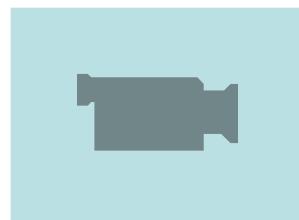


Convection and translation

- ▶ $\eta = 3 \times 10^{17}$ Pa.s, chaotic plume convection :



For moderate viscosity, thermal convection and translation together.



Thermal convection in the inner core

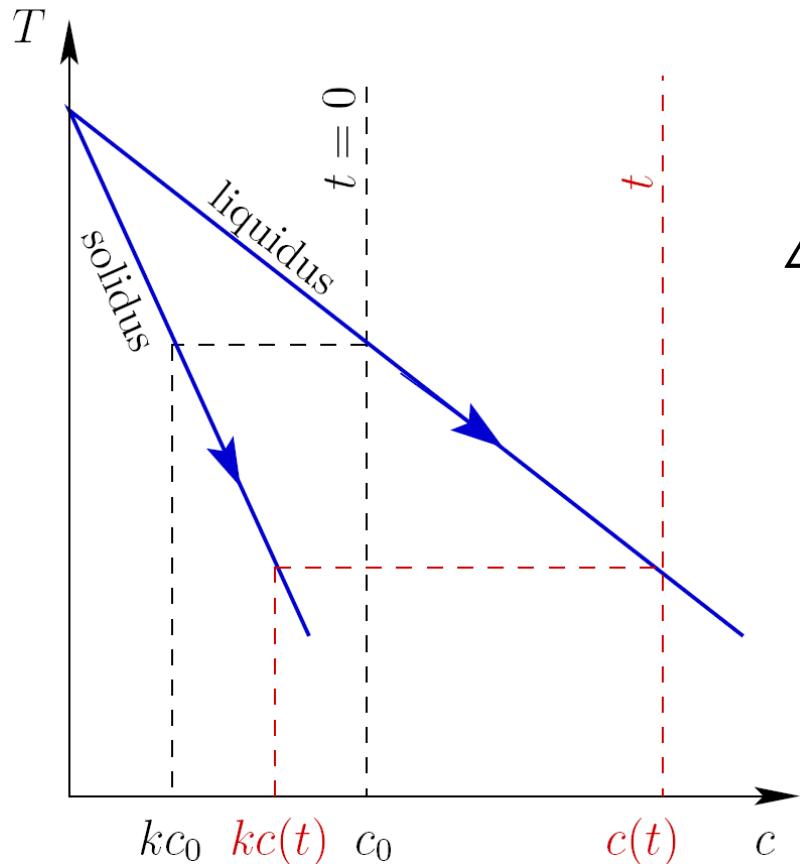
- Whole sphere convection with active cold plumes
- Convective time window: signature of the innermost inner core?
- Fusion and solidification may generate a translation regime for a very viscous inner core.
- Convection and translation may coexist

Outline

1. Seismological observations, anisotropy
2. Review of proposed mechanism
3. Dynamical state of the inner core
4. Thermal convection in the inner core
5. Stratified inner core and equatorial growth
6. Perspectives

Ligth element in the core

$$c(t) = c_0 \left[\frac{r_c^3 - r_{ic}^3}{r_c^3} \right]^{k-1}$$



k , partition coefficient

$$\begin{aligned}\Delta\rho(t) &= \beta\rho k c_0 \left[\left(1 - \left(\frac{r_{ic}}{r_c} \right)^3 \right)^{k-1} - 1 \right] \\ &\sim \beta\rho c_0 k (1 - k) \left(\frac{r_{ic}}{r_c} \right)^3\end{aligned}$$

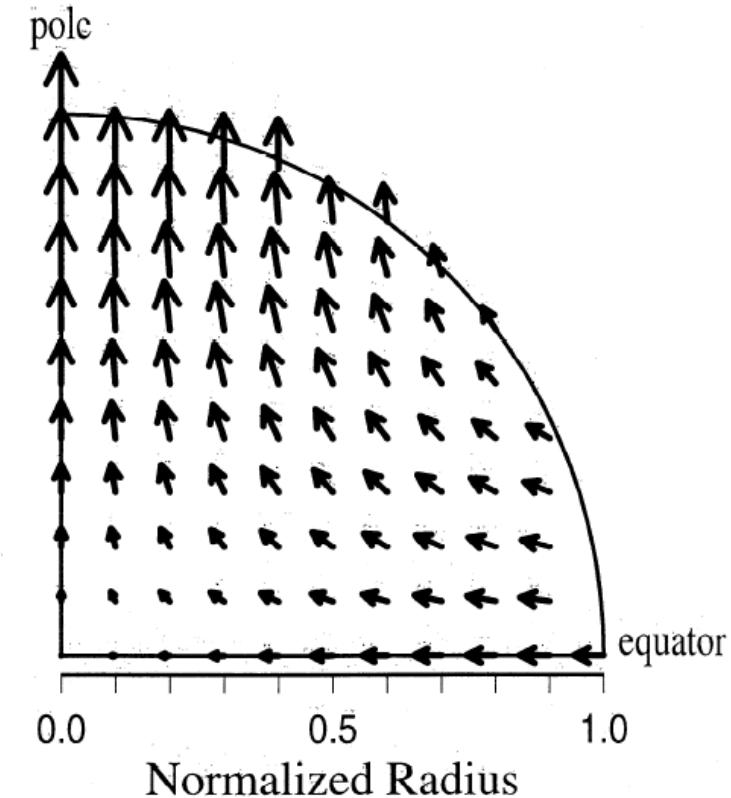
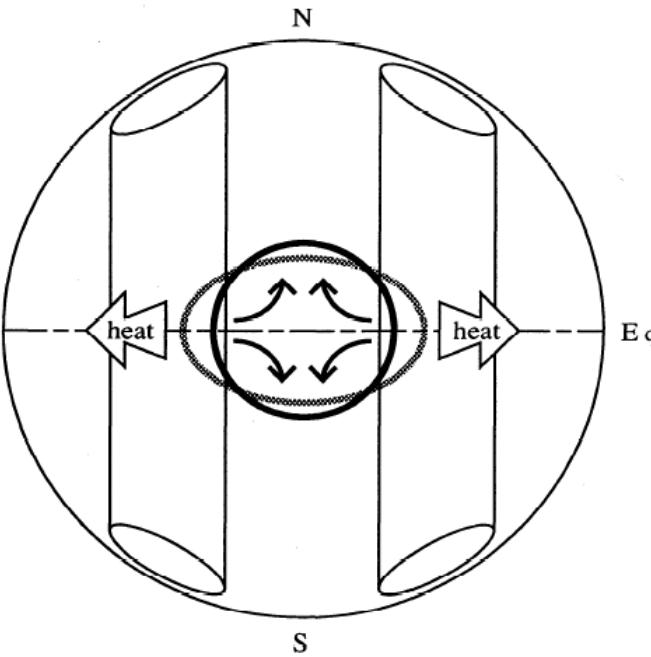
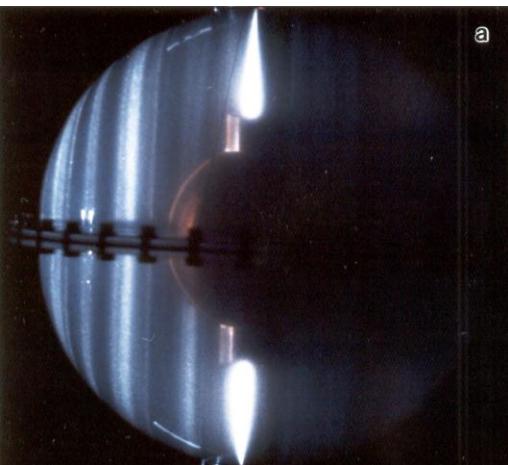
$$\beta = \frac{1}{\rho} \frac{\partial \rho}{\partial c}$$

$$k_{Si} \sim k_S \sim 0.8, k_O \sim 2.5 \cdot 10^{-3}$$

(Alfè et al, 2002)

$$\boxed{\Delta\rho \sim -5 kg/m^3}$$

Equatorial growth



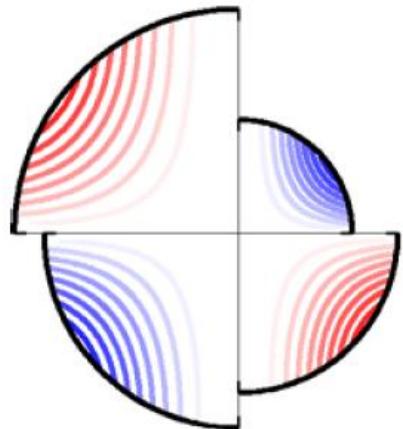
Yoshida et al, 1996

$$B = \frac{g(t) \beta \rho \Delta c(t) r_{ic}(t)^2}{\eta u_{ic}(t)}$$

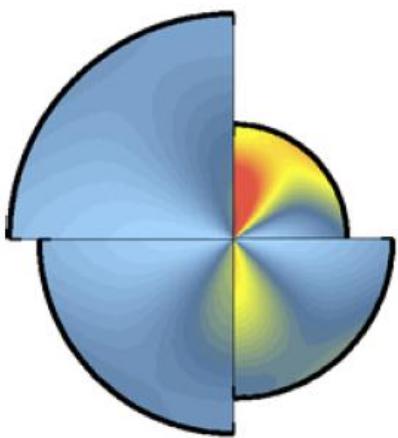
$$S_2 = 2/5$$

$$u_r(r_{ic}, \theta) = -S_2 u_{ic} P_2(\cos \theta)$$

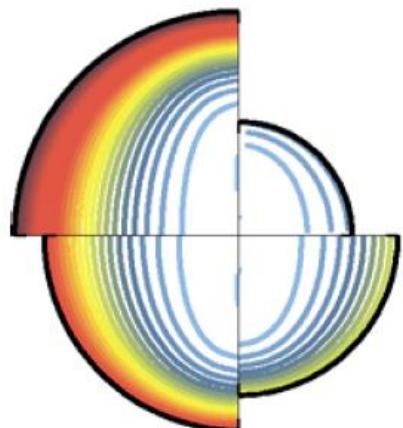
The rate of cristalisation in the equator is twice the polar one



Ψ

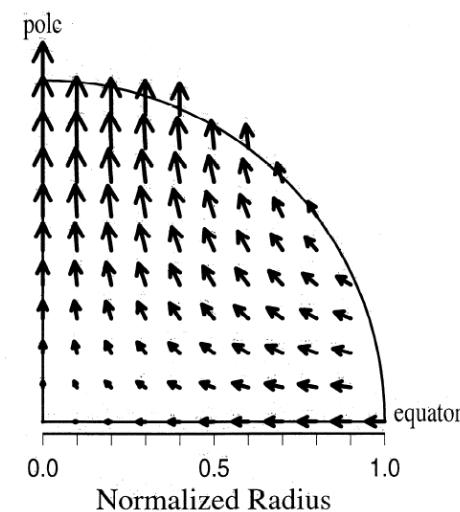


$\dot{\varepsilon}$ ($M_{\odot} \text{Myr}^{-1}$)

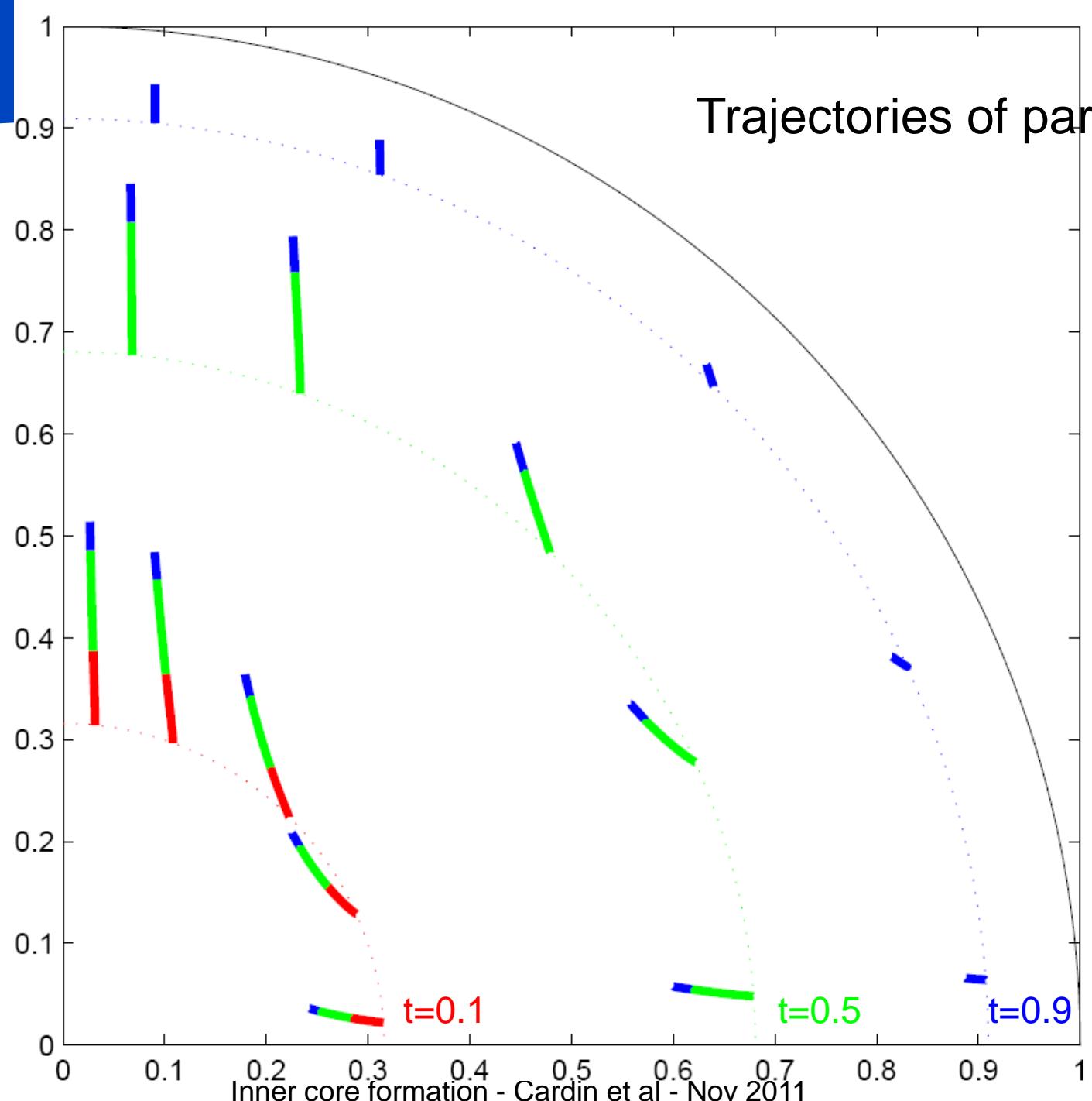


χ

No stratification



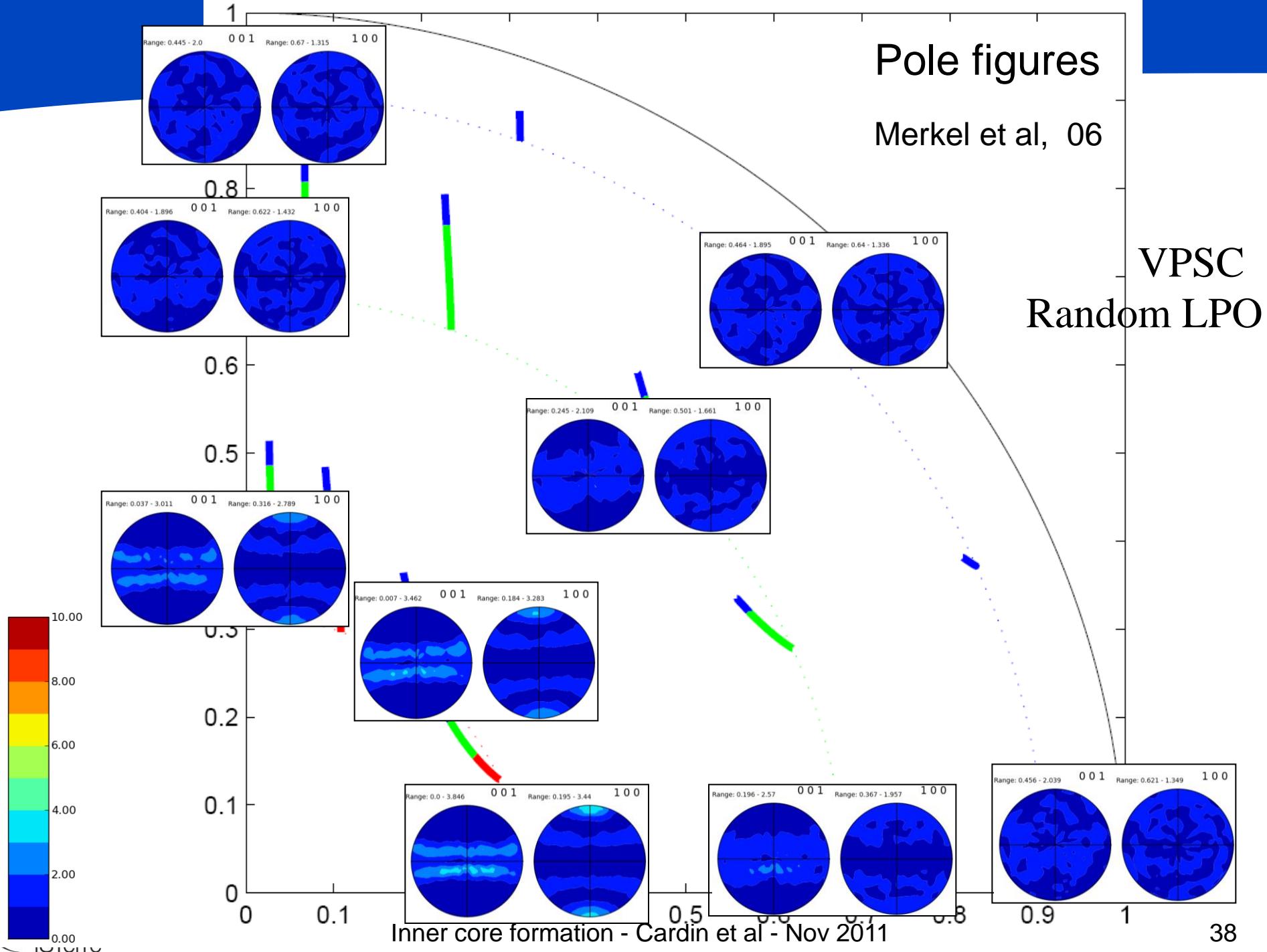
Trajectories of particles



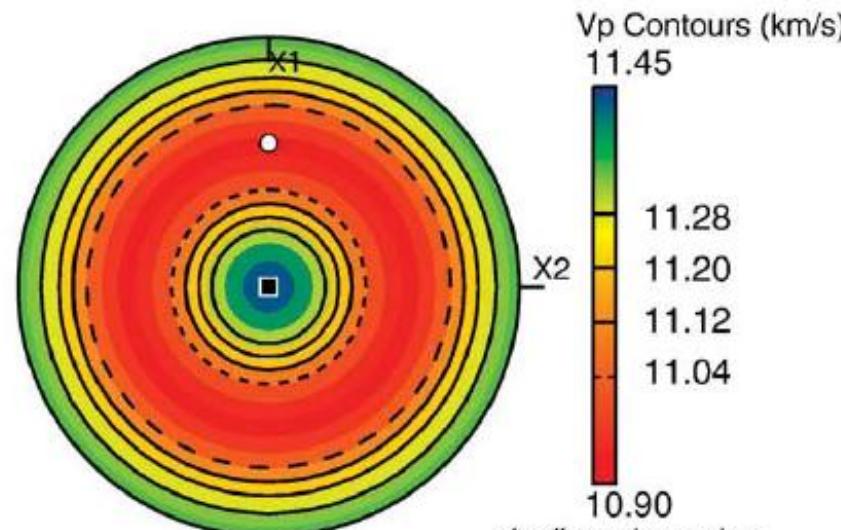
Pole figures

Merkel et al, 06

VPSC
Random LPO



5000K



Vočadlo et al., 09

L. Vočadlo et al. / Earth and Planetary Science Letters 288 (2009) 534–538

535

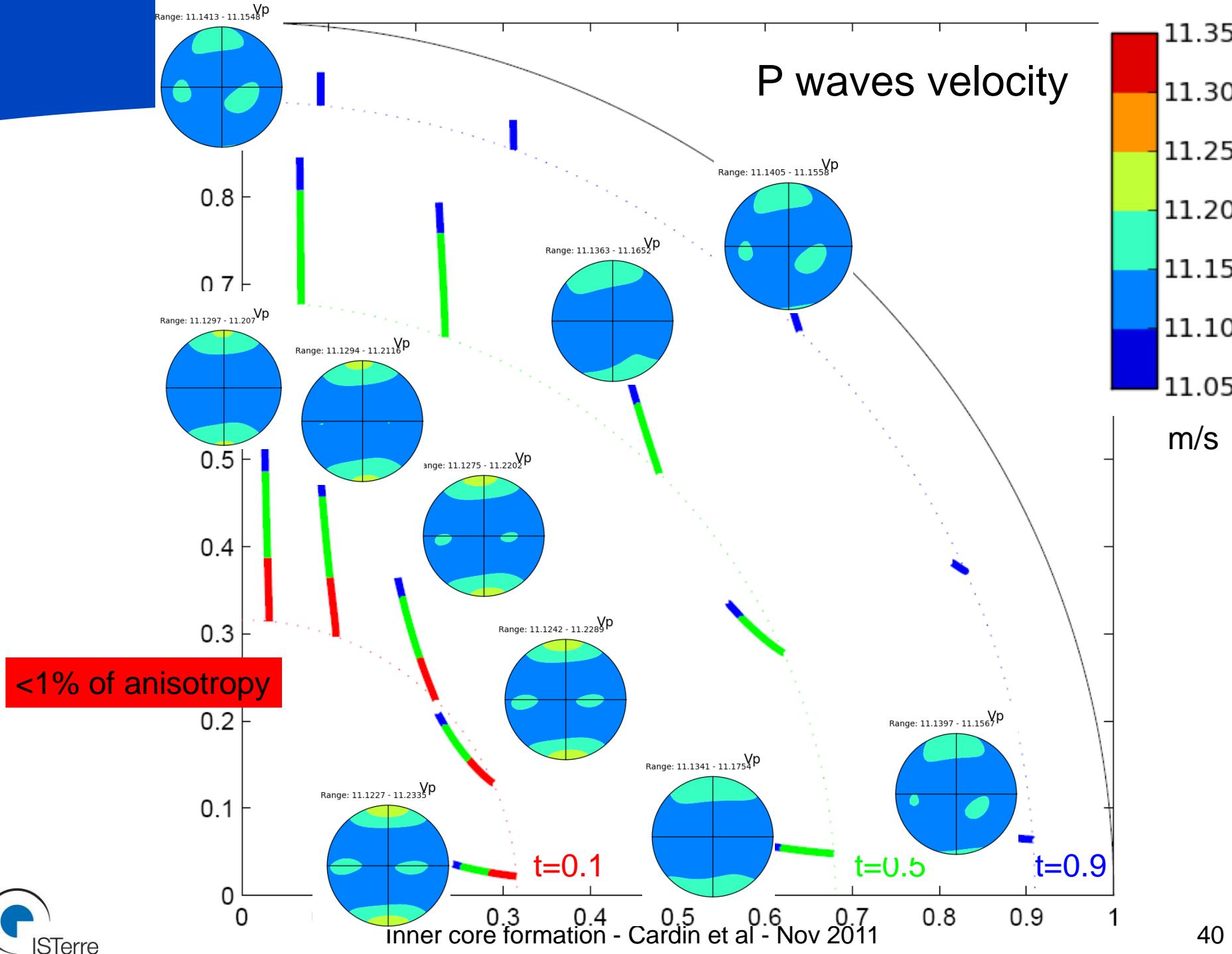
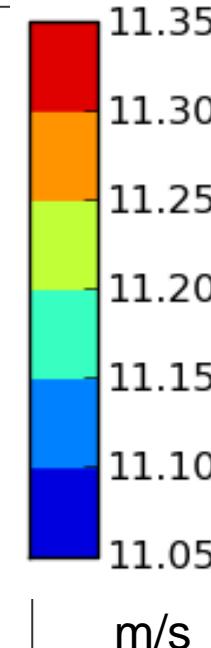
Table 1

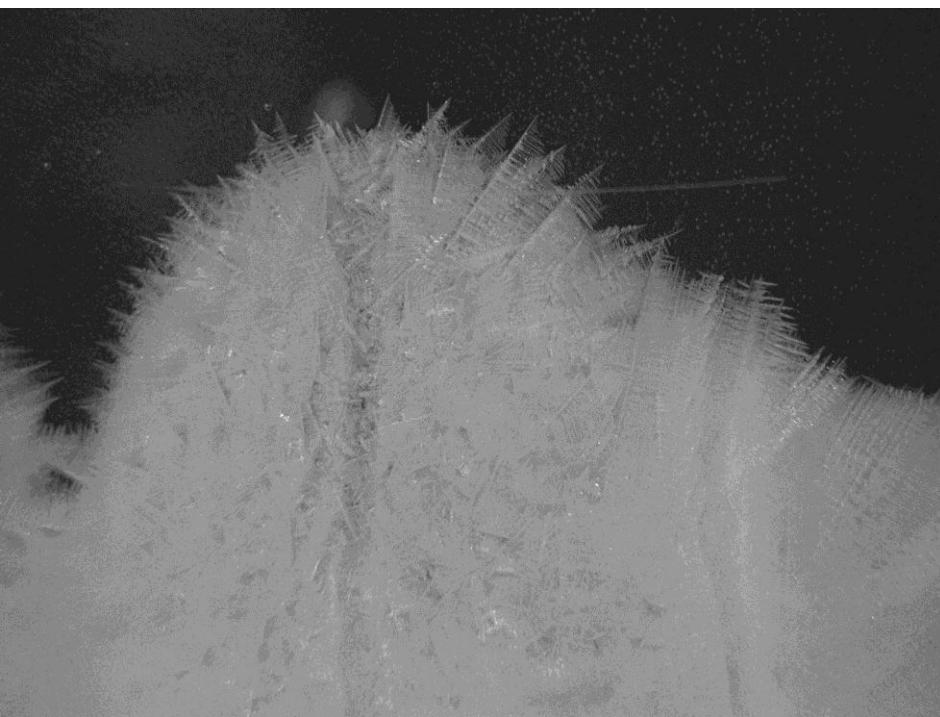
Isothermal elastic constants (in GPa), adiabatic incompressibility, shear modulus and sound velocities for hcp-Fe as a function of temperature.

T (K)	c/a	ρ (kgm ⁻³)	P (GPa)	c ₁₁	c ₃₃	c ₁₂	c ₂₃	c ₄₄	c ₆₆	K _S (GPa)	G (GPa)	V _P (km s ⁻¹)	V _S (km s ⁻¹)
0	1.585	13698	295	2205	2418	1053	950	479	576	1419	564	12.59	6.42
2000	1.59	13401	290	1998	2183	1111	986	375	444	1412	444	12.23	5.75
3000	1.60	13324	293	1915	2109	1121	986	313	397	1407	394	12.05	5.44
4000	1.61	13236	298	1830	1905	1150	1004	248	340	1399	328	11.78	4.98
5000	1.62	13154	308	1689	1725	1186	990	216	252	1365	266	11.44	4.50
5500	1.62	13155	316	1646	1559	1253	995	153	197	1363	208	11.17	3.97

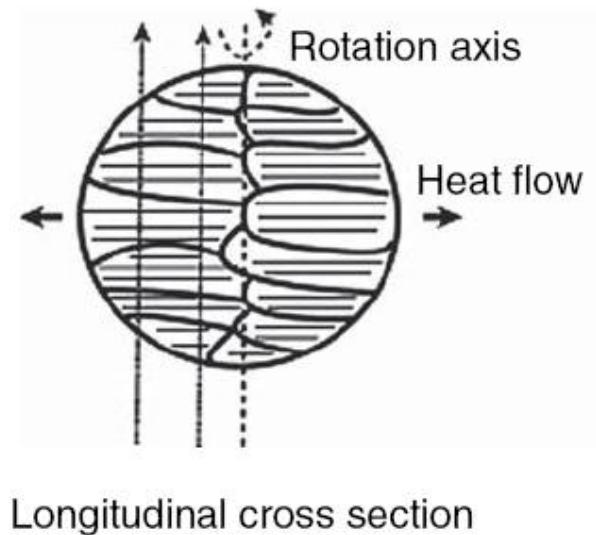
c₆₆ is not independent and is equal to 1/2 (c₁₁ − c₁₂). Uncertainties in c_{ii} < 2.5%; uncertainties in V_P and V_S are ~1%. The values for c/a were taken from Gannarelli et al. (2005).

P waves velocity





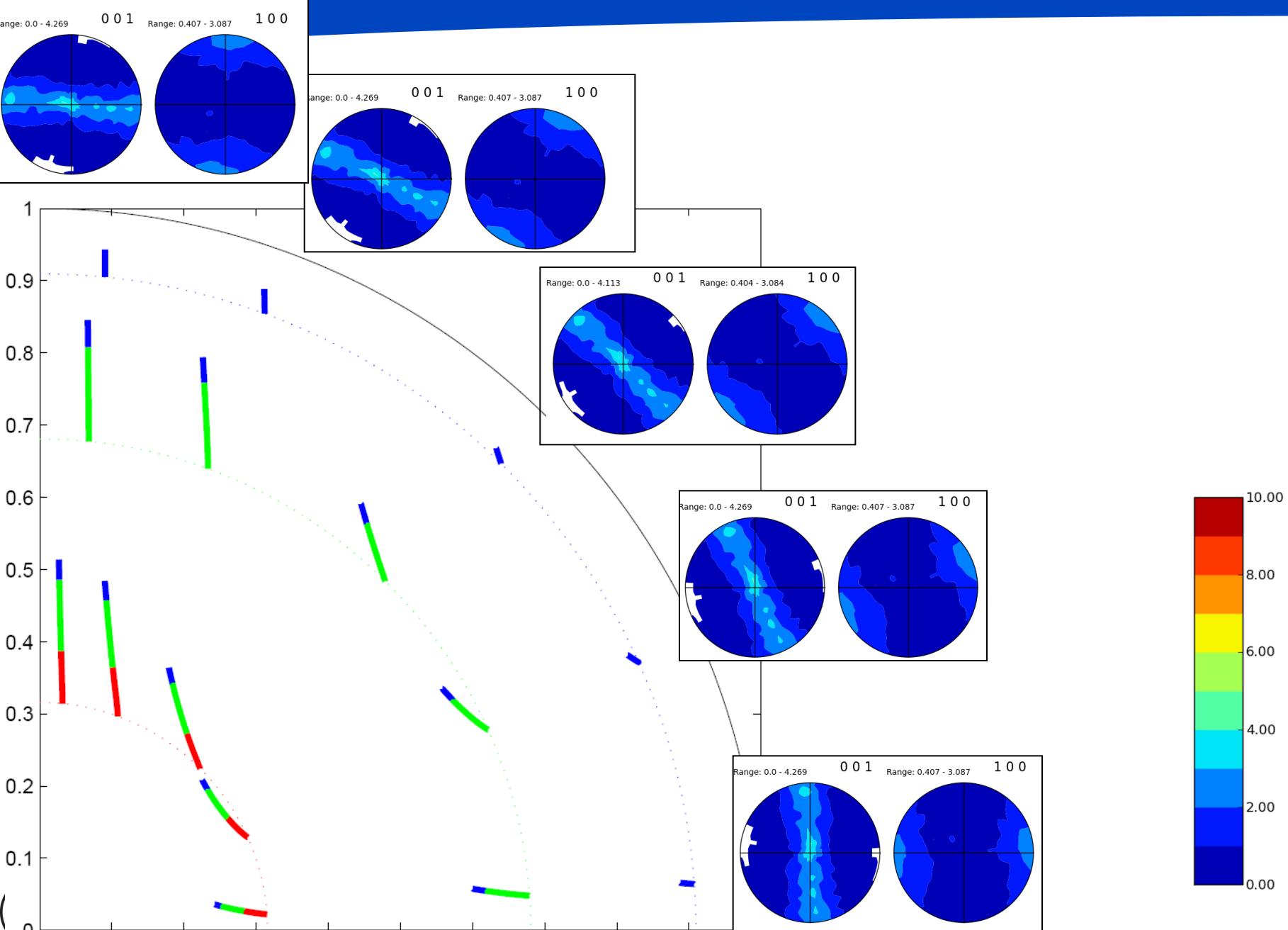
Deguen, 09



Bergman, 97

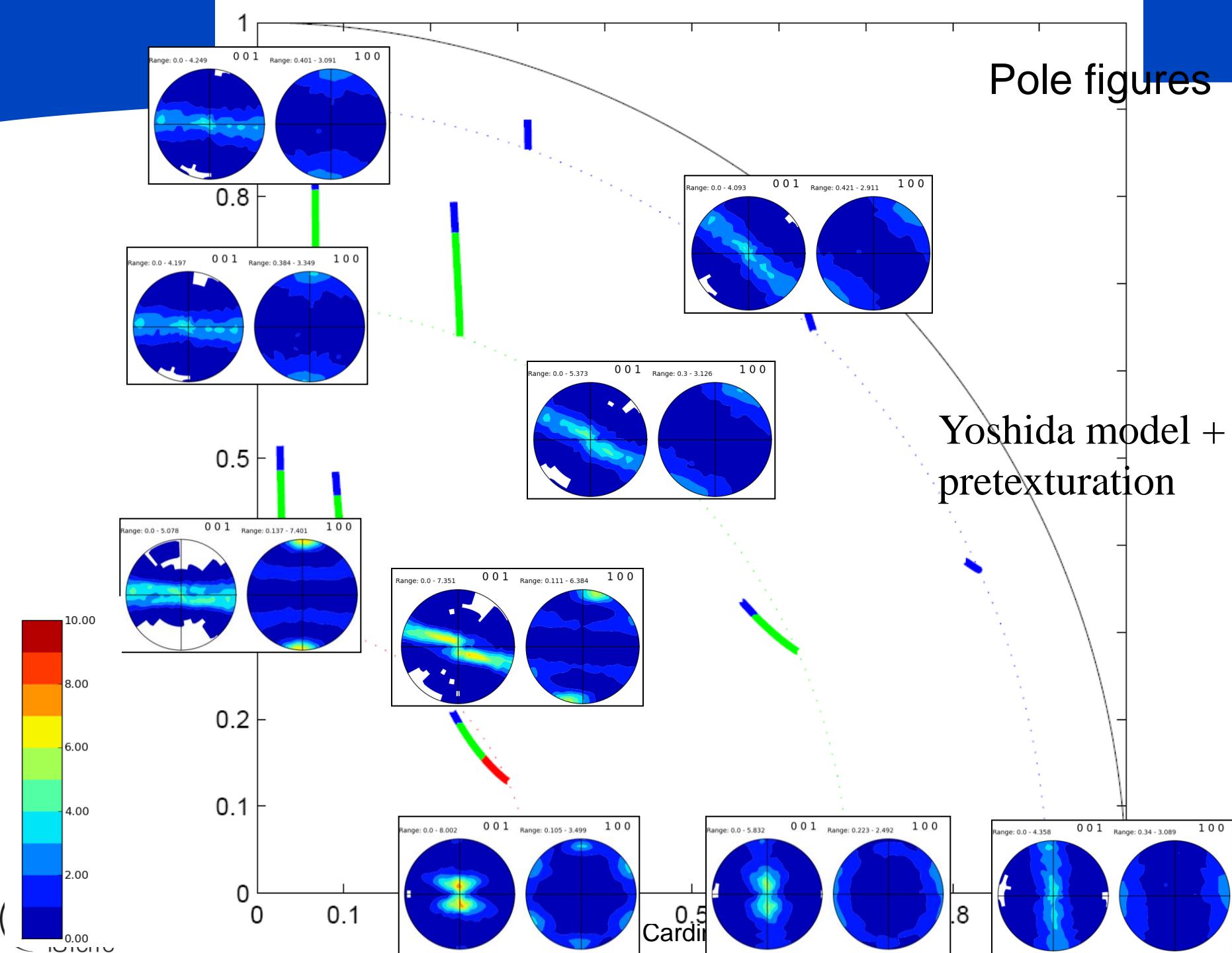
C axes parallel to crystallisation front

Surface pretexturbation

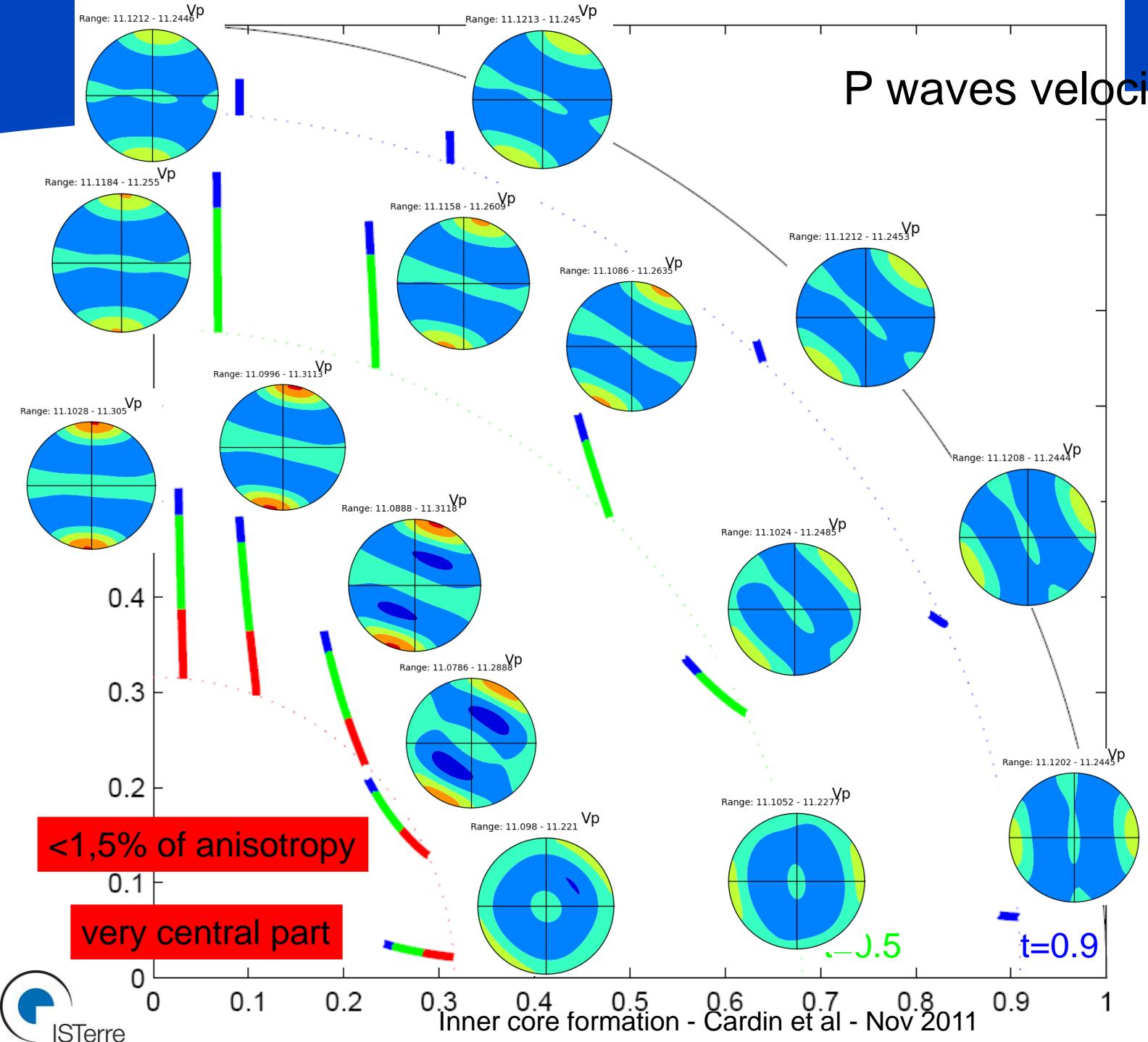
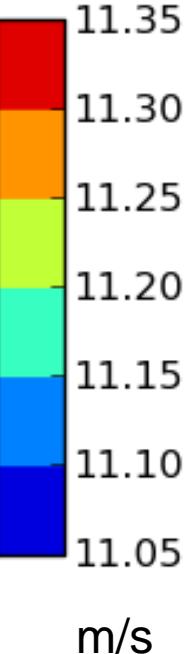


Pole figures

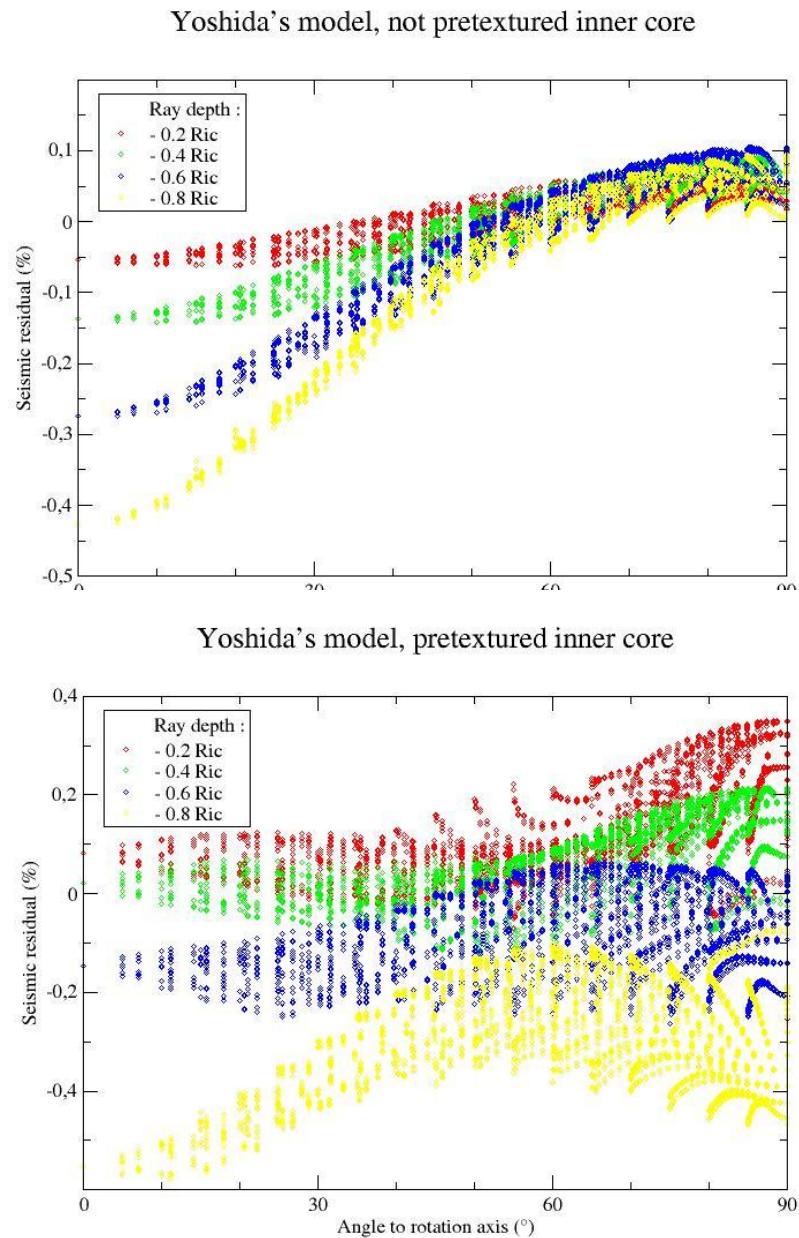
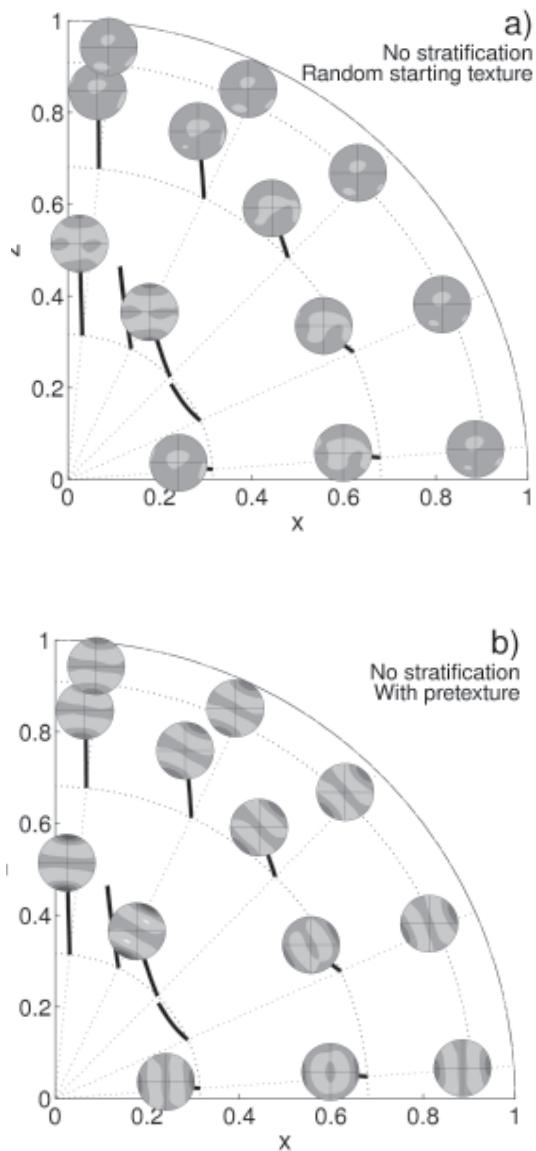
Yoshida model +
pretexturation



P waves velocity

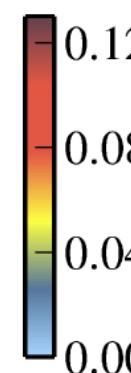
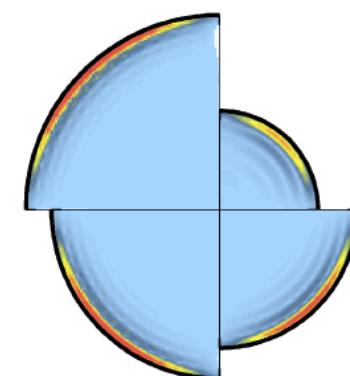
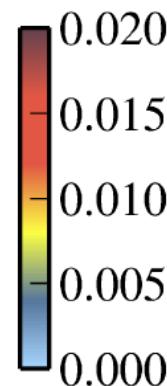
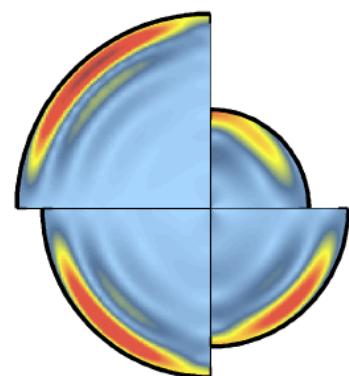
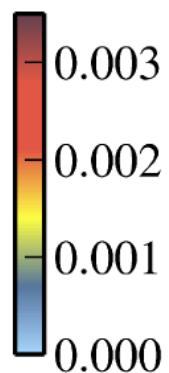
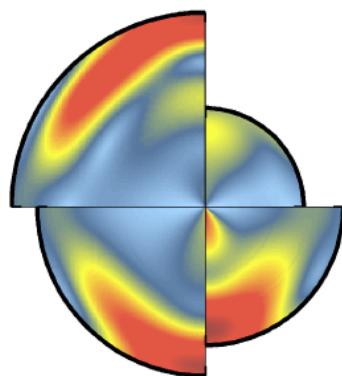
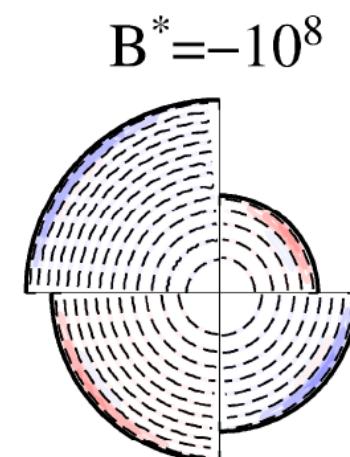
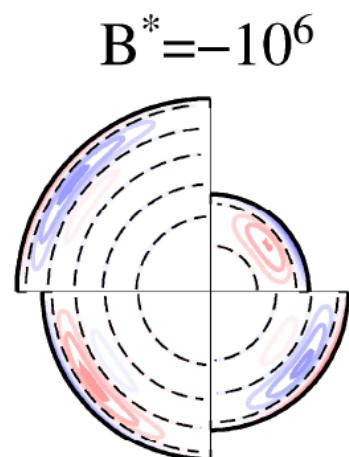
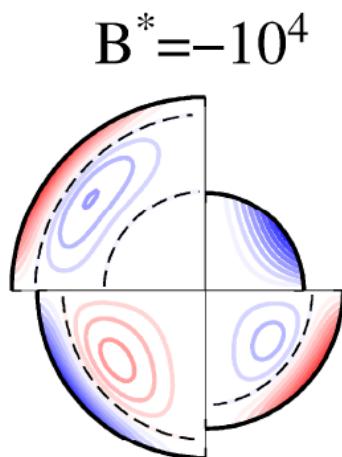


Residuals for Yoshida model



Stratified models

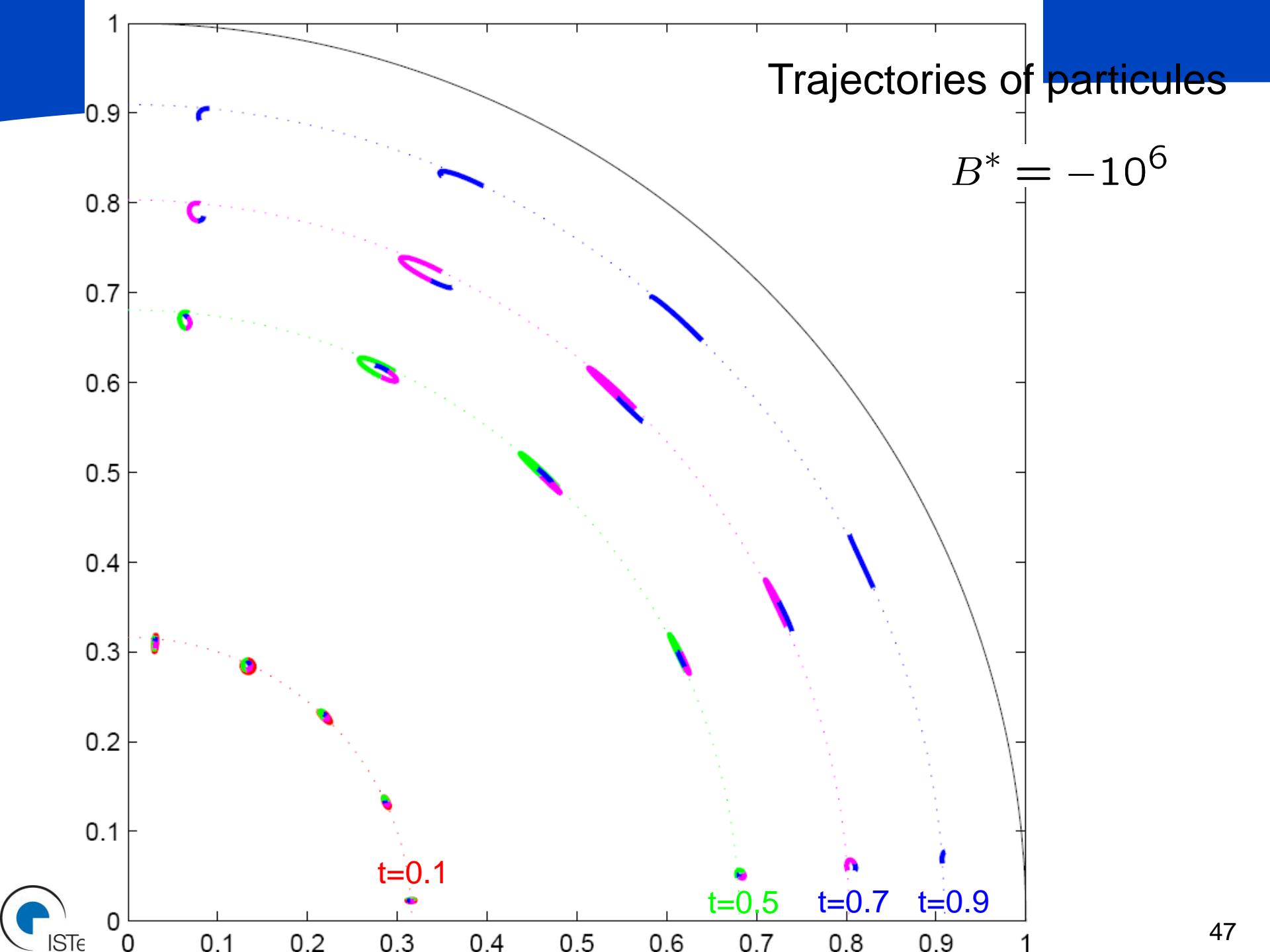
More and more stratification



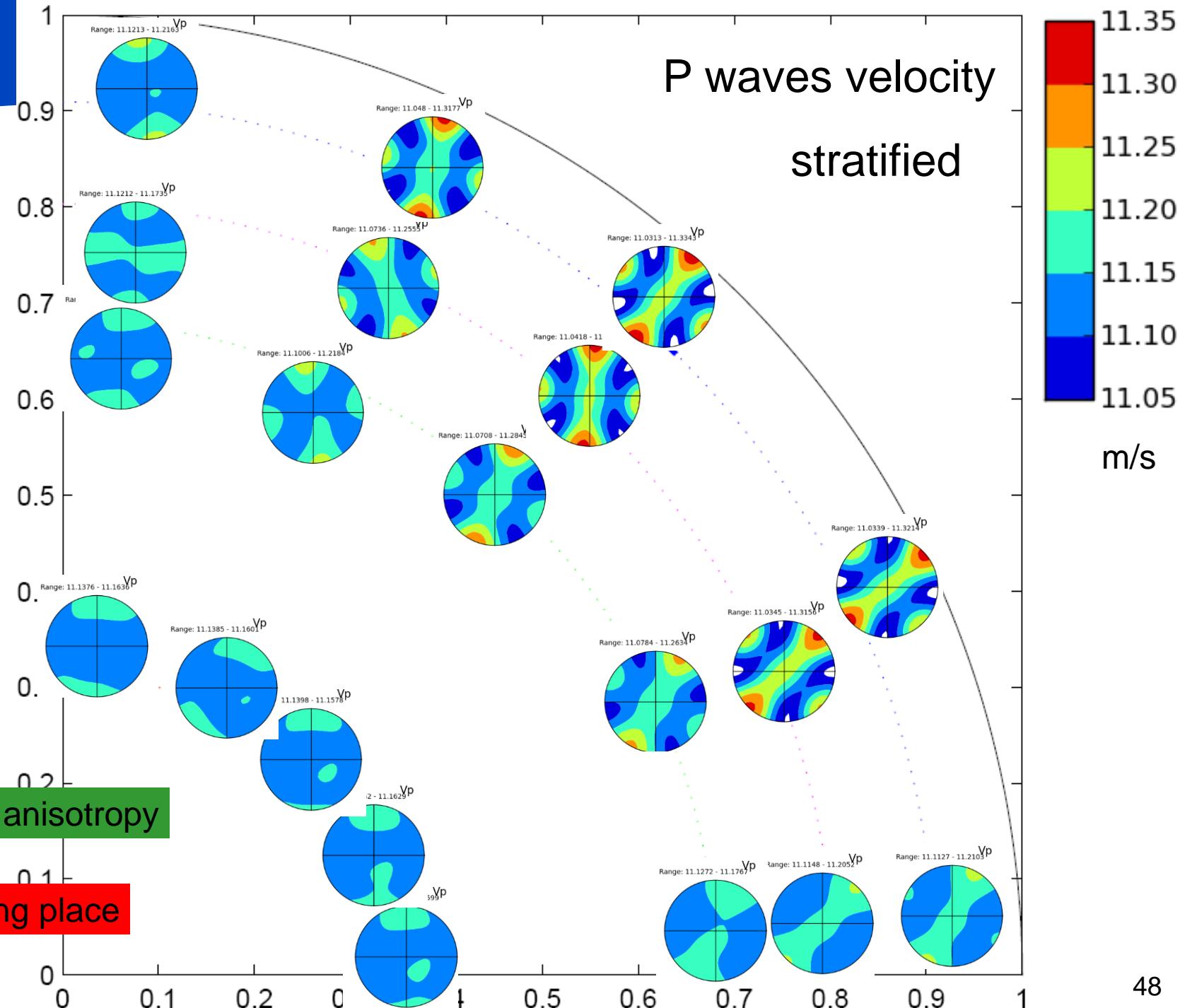
$\dot{\varepsilon} (\text{My}^{-1})$

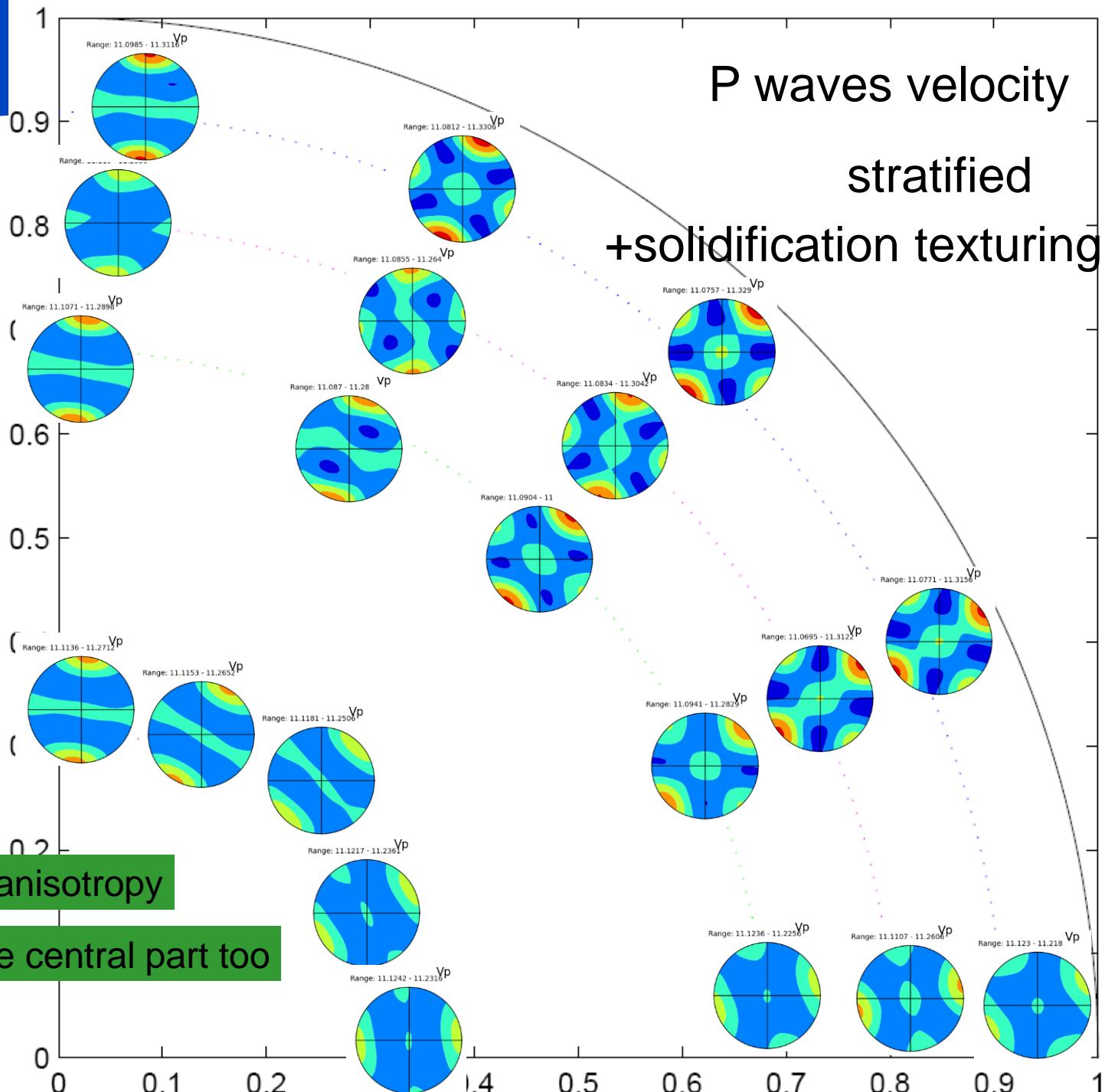
$$B^* = -10^6$$

Trajectories of particles



P waves velocity stratified





Stratified inner core and equatorial growth

- Stratification inhibits radial motion
- Superficial layer with largest stresses
- With no stratification, anisotropy is too small
- Stratification and solidification texture produce V_p maps compatible to seismic observation.

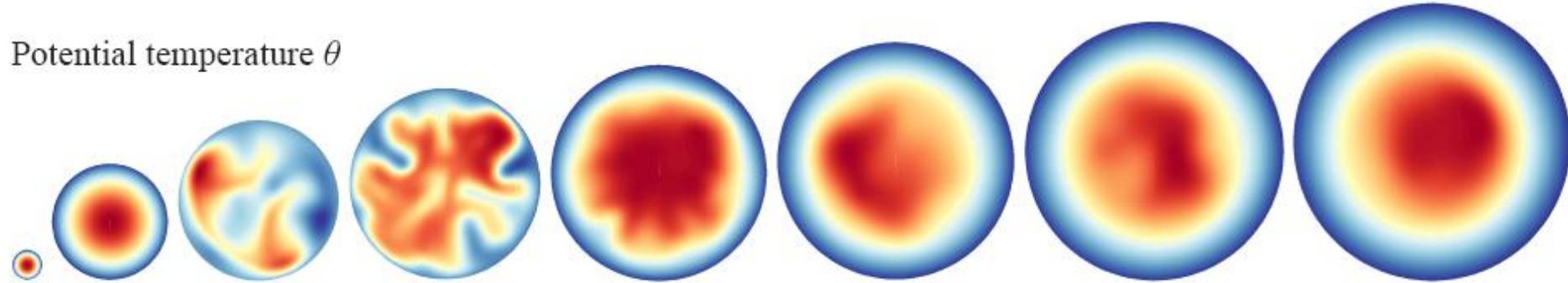
Outline

1. Seismological observations, anisotropy
2. Review of proposed mechanism
3. Dynamical state of the inner core
4. Thermal convection in the inner core
5. Stratified inner core and equatorial growth
6. Perspectives

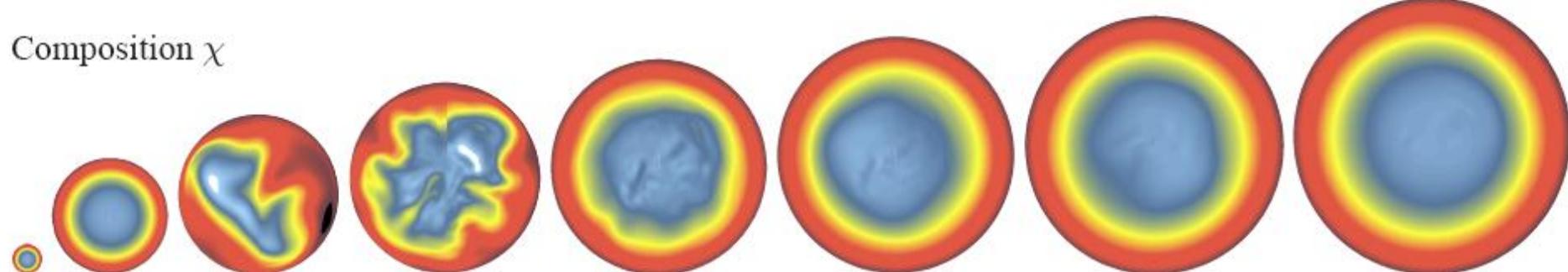
Double diffusive convection

a. $\tau_{\text{ic}} = 1 \text{ Gy}$, $\eta = 10^{18} \text{ Pa.s}$, $k = 36 \text{ W.m}^{-1}.\text{K}^{-1}$, $c_0 = 5.6 \text{ wt.\%}$.

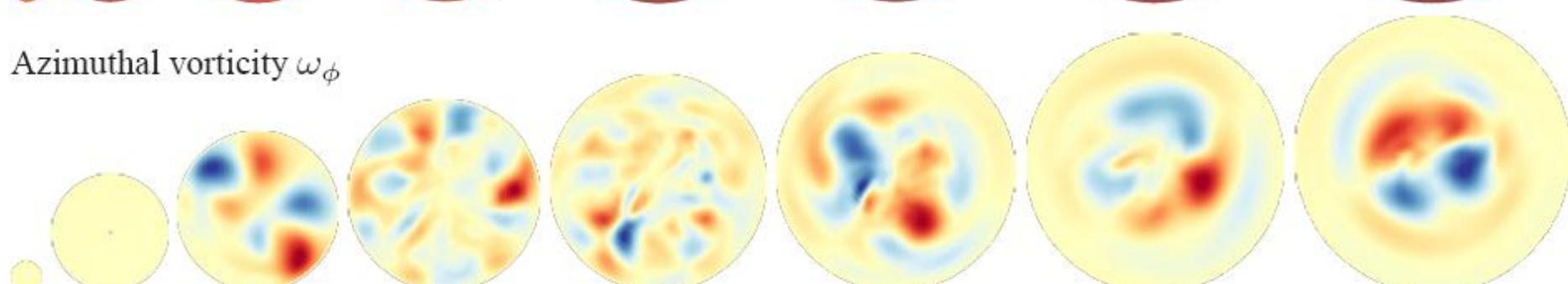
Potential temperature θ



Composition χ



Azimuthal vorticity ω_ϕ



References

- Tectonic history of the Earth's inner core preserved in its seismic structure. Deguen, R., Cardin, P. *Nature Geoscience* 2, 419–422, 2009.
- Texturing in the Earth's inner core due to preferential growth in its equatorial belt, R. Deguen, Ph. Cardin, S. Merkel, R. Lebensohn, *PEPI*, 2011
- Thermo-chemical convection in the Earth's inner core, R. Deguen and P. Cardin, *Geophysical Journal International*, 2011
- Thèse en cours Ainhoa Lincot