

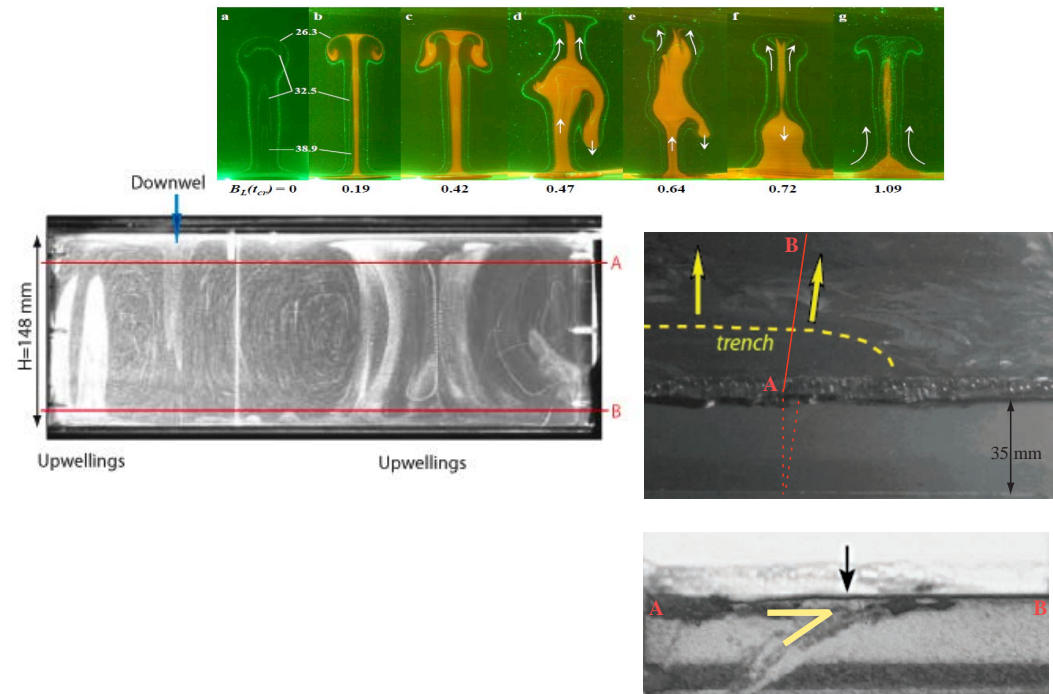
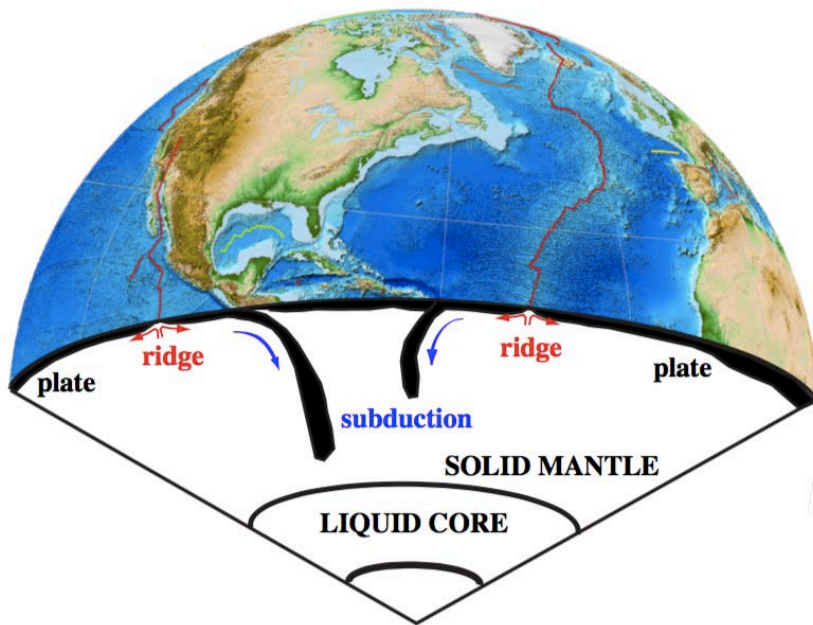
# Structure of Mantle Convection: Plumes and Plates

**Anne Davaille**, (FAST, CNRS / Univ. Paris-Sud, Orsay, France)

Michael Le Bars, Sophie Androvandi, Cecilia Cadio, Judith Vatteville, Valérie Vidal, Eléonore Stutzmann, Jean Besse, Isabelle Panet, Vincent Courtillot (IPGP)

Erika Di Giuseppe, Eric Mittelstaedt, Aurore Sibrant (FAST)

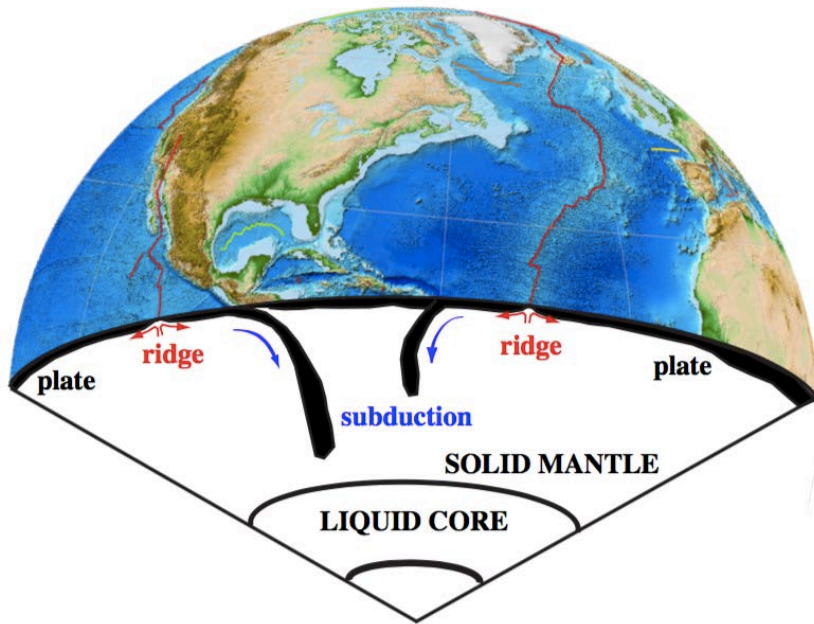
, Ichiro Kumagai, Kei Kurita (ERI, Tokyo), Sue Smrekar (JPL, USA), Nick Arndt (ISTerre, Grenoble)



**The view of a physicist ...**

**= provide a framework to predict and to interpret the observations**

# Evolution of a planet = cooling ⇔ Thermal convection

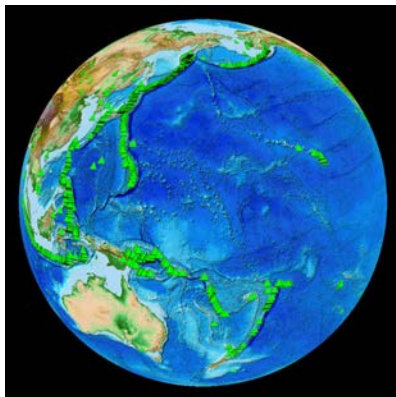


1-What are the conditions for thermal convection to produce plate tectonics ?

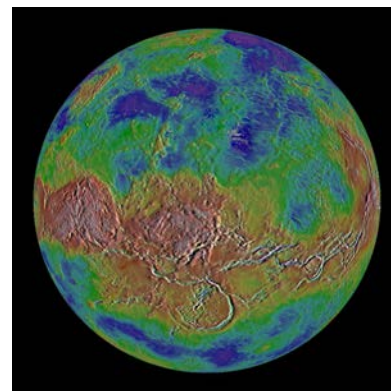
2- Structure of convection

3-Evolution of a planet?

4- What type of planet ? Exoplanet ?



Continuous Resurfacing



Episodic Resurfacing



No Resurfacing

## 2- Observations :

### A- Seismic Tomography:

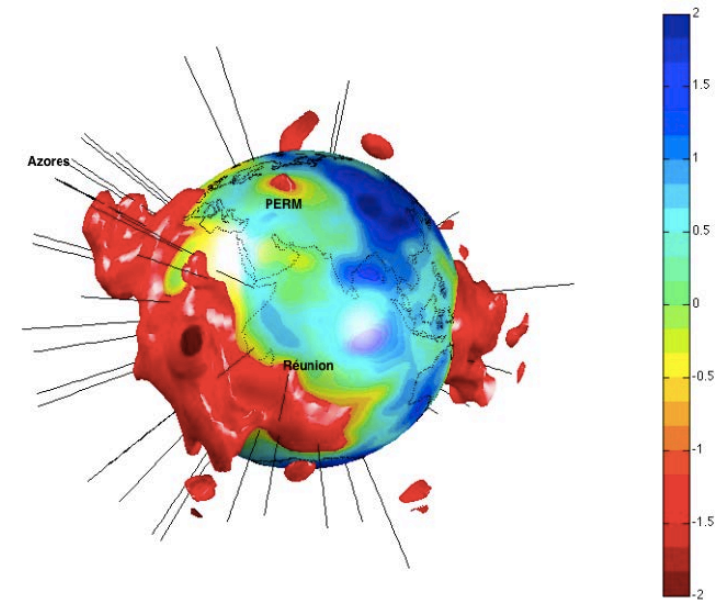
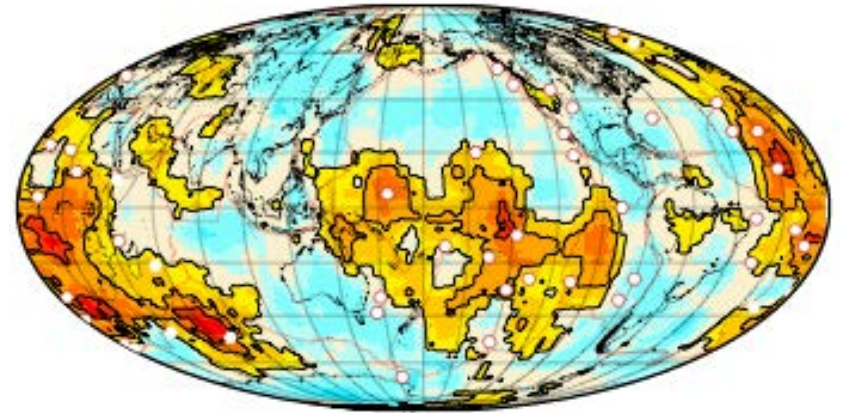
- Ridges are shallow features
- Blue (COLD) tracable at all depth

✂ Delimit 2 main « boxes »  
and 2 LLSVPs  
(Large Low S-Velocity Provinces)

### B- Geochemistry:

- MORB  $\neq$  OIB
  - Existence of long-lived reservoirs
    - primitive material
    - recycled crust
- => density heterogeneities

TXBW (Grand, 1997)

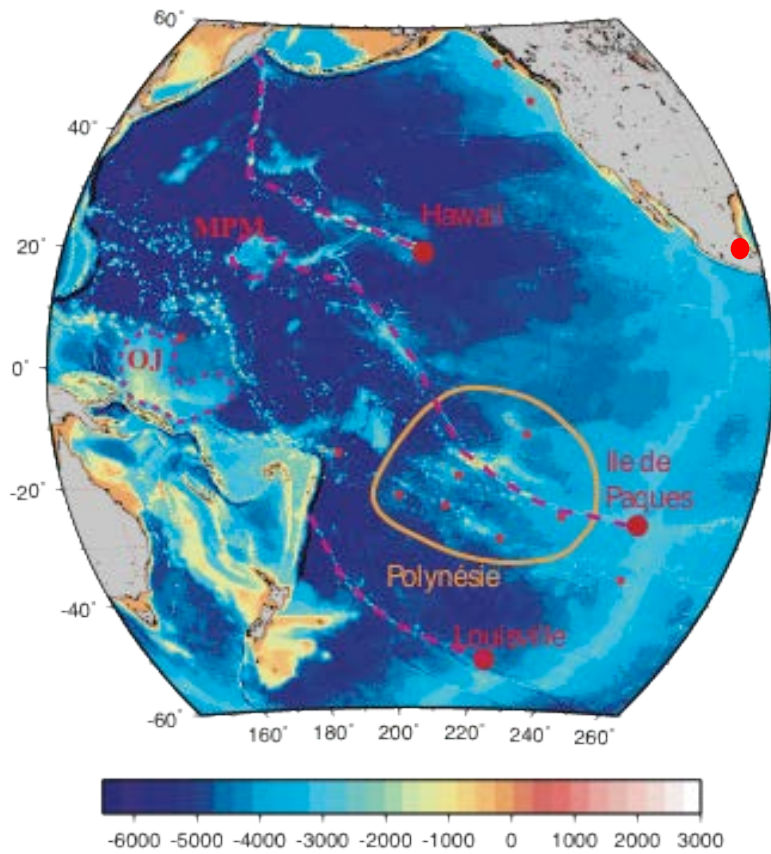


SMEAN (Boschi & Becker, 2002)

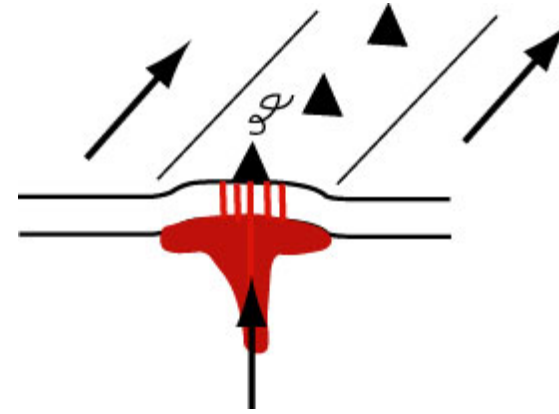
## 2- Observations :

### C- the « Hot spots » Zoo:

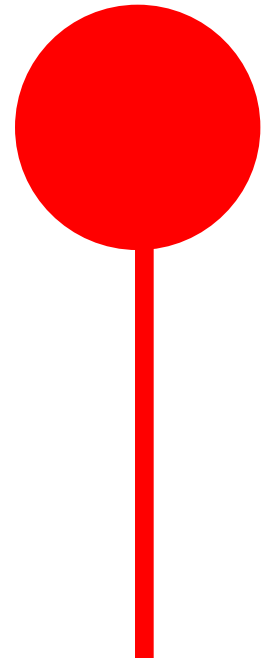
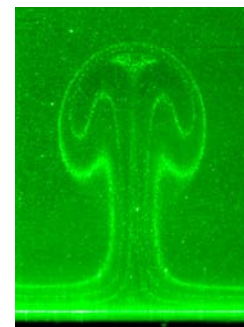
(Wilson 1963; Morgan 1971, 1972)



(Clouard & Bonneville 2000)



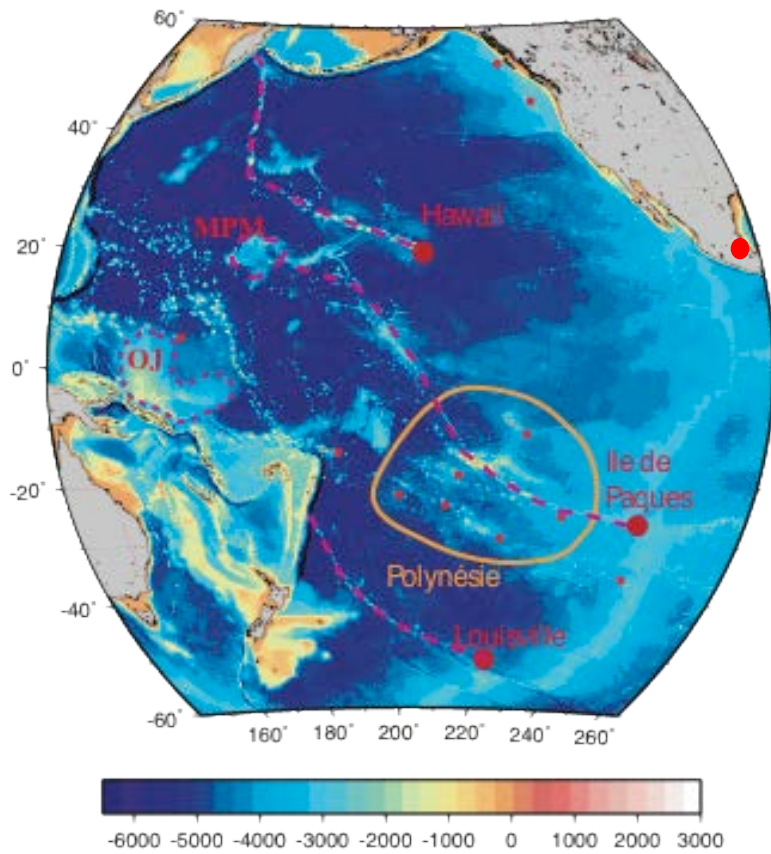
- Volcanism fixed / moving plate
- can start on by an oceanic plateau



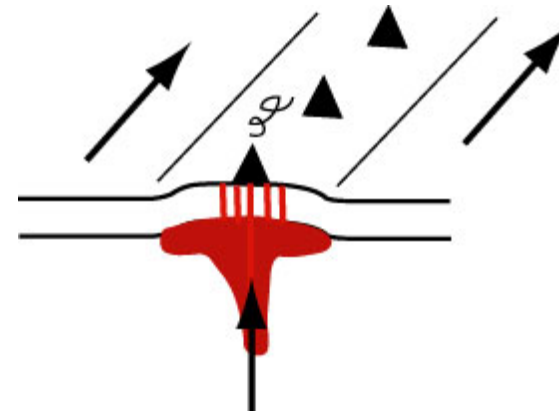
## 2- Observations :

### C- the « Hot spots » zoo:

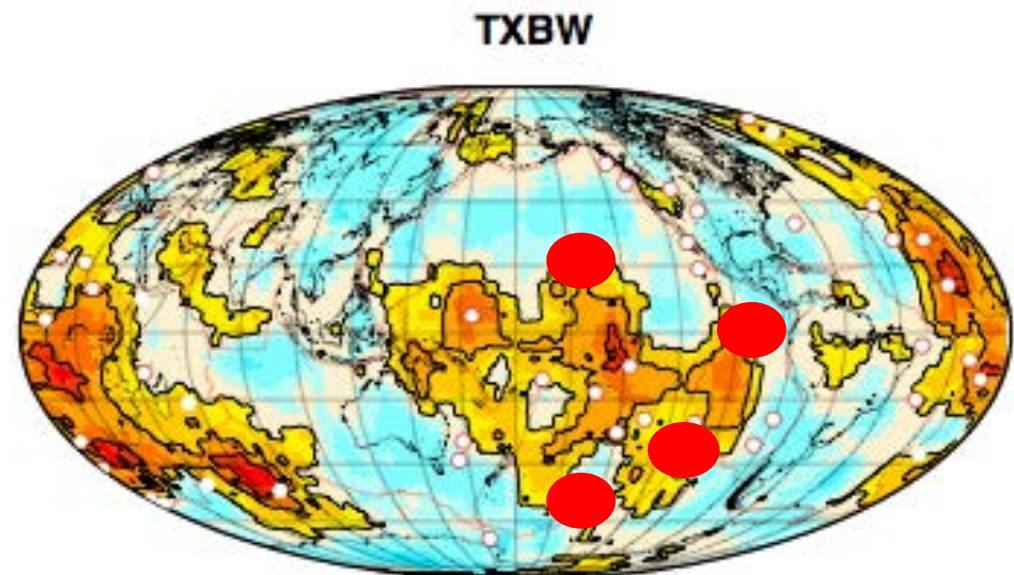
(Wilson 1963; Morgan 1971, 1972)



(Clouard & Bonneville 2000)



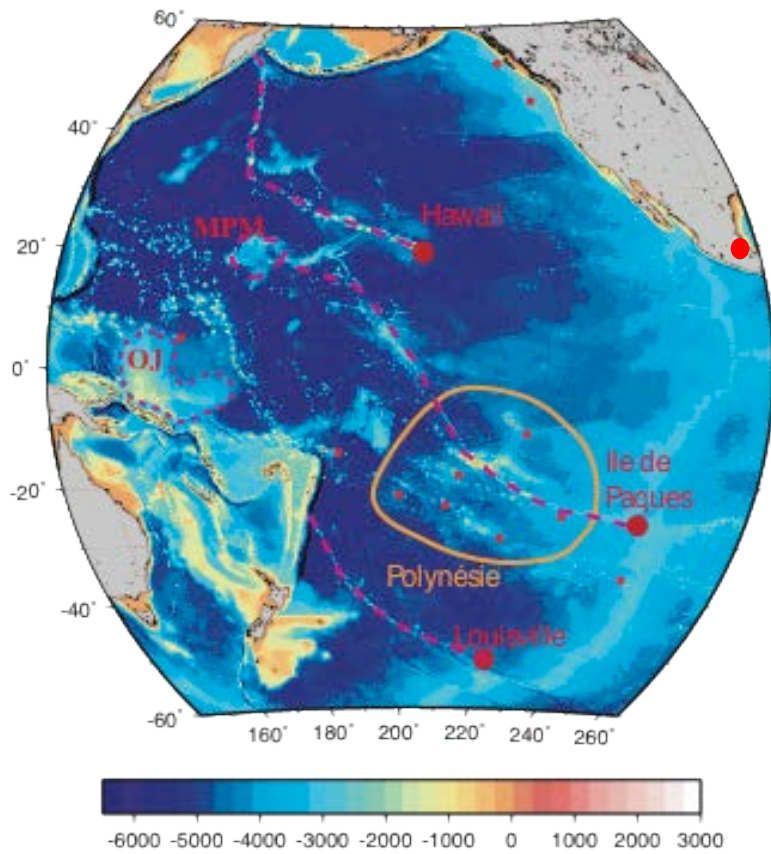
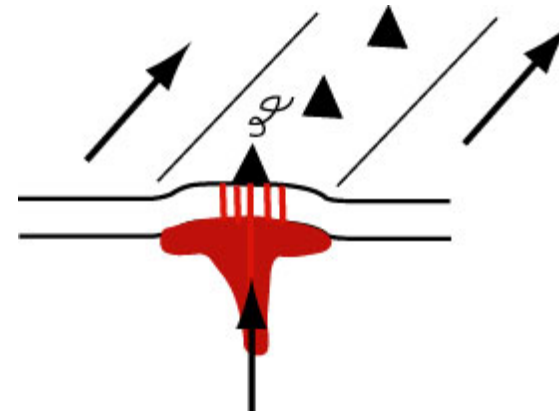
- Volcanism fixed / moving plate
- can start by an oceanic plateau
- Long tracks on the edges of slow anomaly



## 2- Observations :

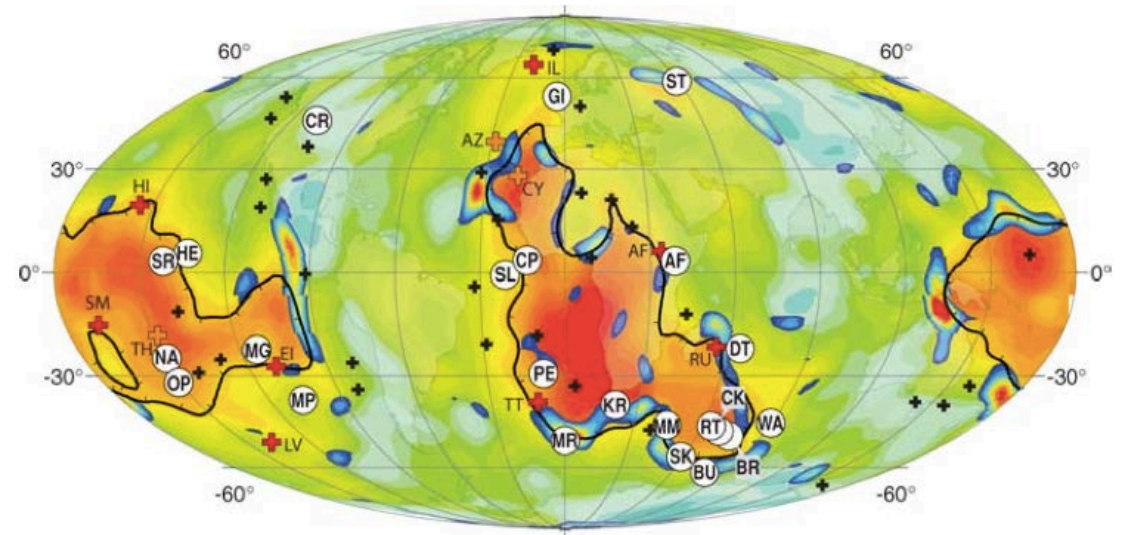
### C- the « Hot spots » zoo:

(Wilson 1963; Morgan 1971, 1972)



(Clouard & Bonneville 2000)

- Volcanism fixed / moving plate
- can start by an oceanic plateau
- Long tracks/LIPs on the edges of slow anomaly



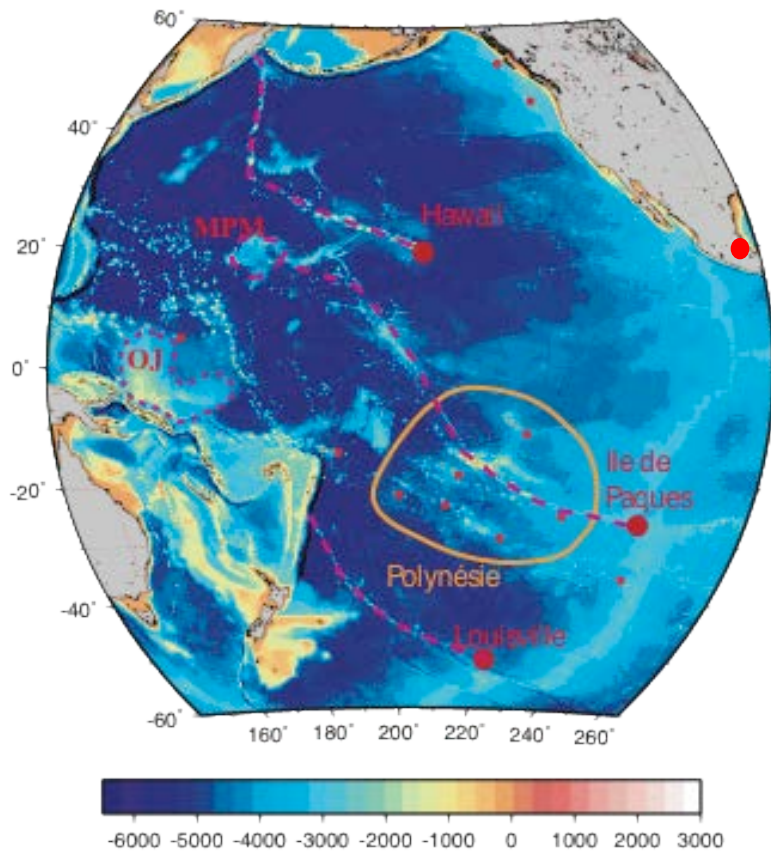
(Torsvik et al, 2004)

## 2- Observations :

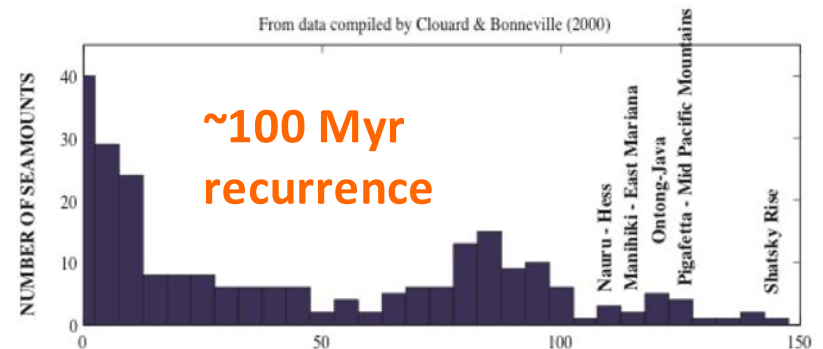
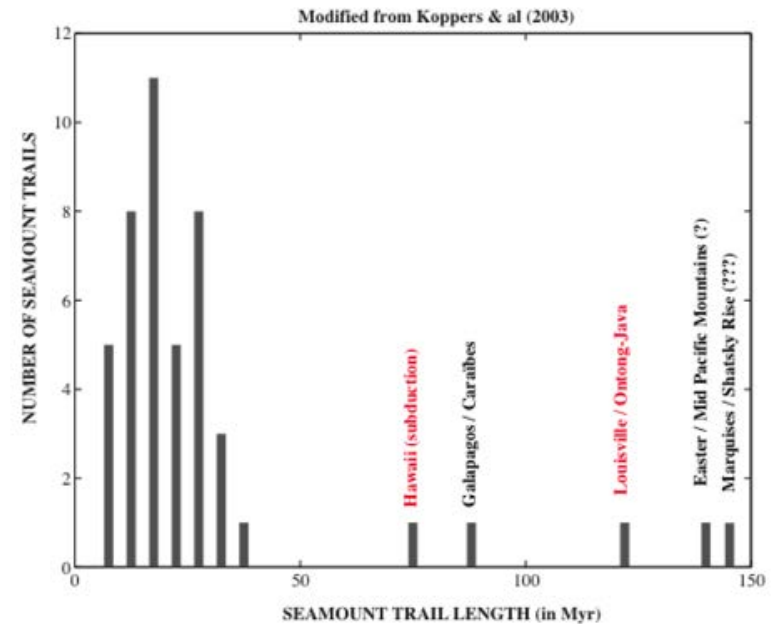
### C- the « Hot spots » zoo:

(Wilson 1963; Morgan 1971, 1972)

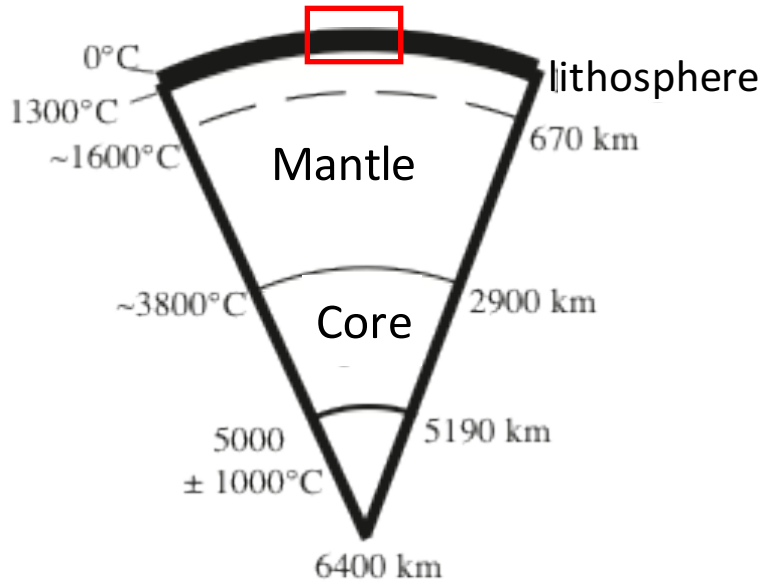
+ another type  
= shorter tracks in « cluster »



(Clouard & Bonneville 2000)



### 3-Thermal convection:



+ Boundary conditions= fixed temperatures

$T \sim 0^\circ\text{C}$  at surface

$T \sim 3500^\circ\text{C}$  at CMB

+ Free surfaces

+ **Rheology** : in cold Thermal Boundary Layer depends on  $T_p$ ,  $P$ , stress, composition

**Evolution of a planet = cooling  $\Leftrightarrow$  Thermal convection**

Continuity:  $\vec{\nabla} \cdot \vec{u} = 0$

Momentum conservation:

$$\frac{1}{\text{Pr}} \left( \frac{d\vec{u}}{dt} + (\vec{u} \cdot \vec{\nabla}) \vec{u} \right) = -\vec{\nabla} P + \vec{\nabla} \cdot \left( \eta (\vec{\nabla} \vec{u} + \vec{\nabla} \vec{u}^t) \right) - Ra T \vec{n}$$

Heat conservation:

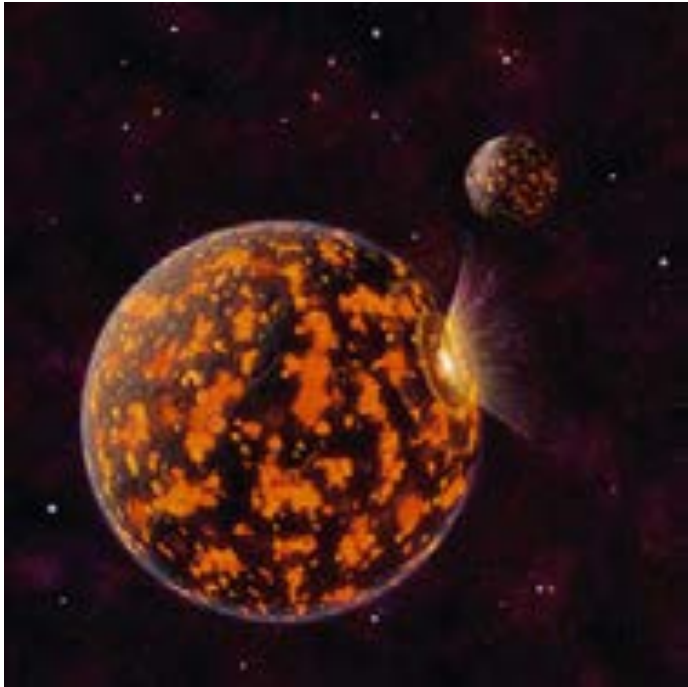
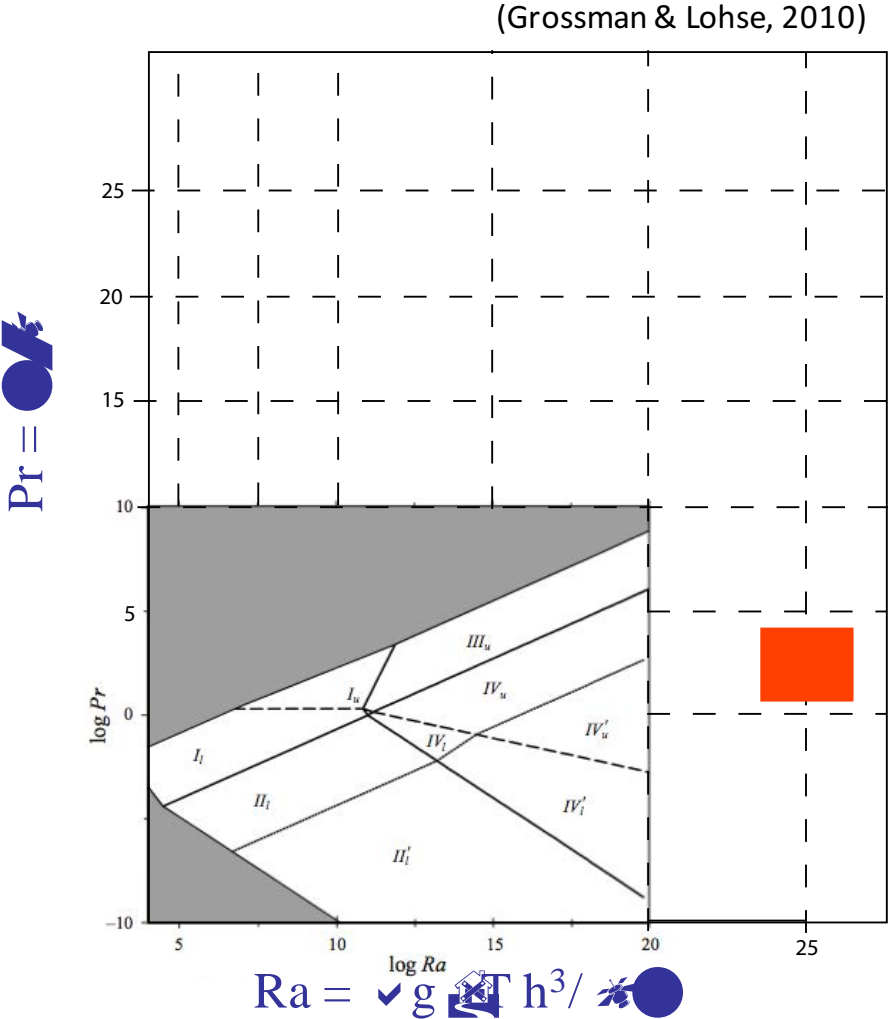
$$\frac{dT}{dt} + \vec{u} \cdot \vec{\nabla} T = \vec{\nabla} \cdot (k \cdot \vec{\nabla} T)$$

.  $\text{Pr} = \frac{\rho c_p \kappa}{\nu}$

.  $\text{Ra} = \frac{\alpha g \Delta T d^3}{\nu \kappa}$



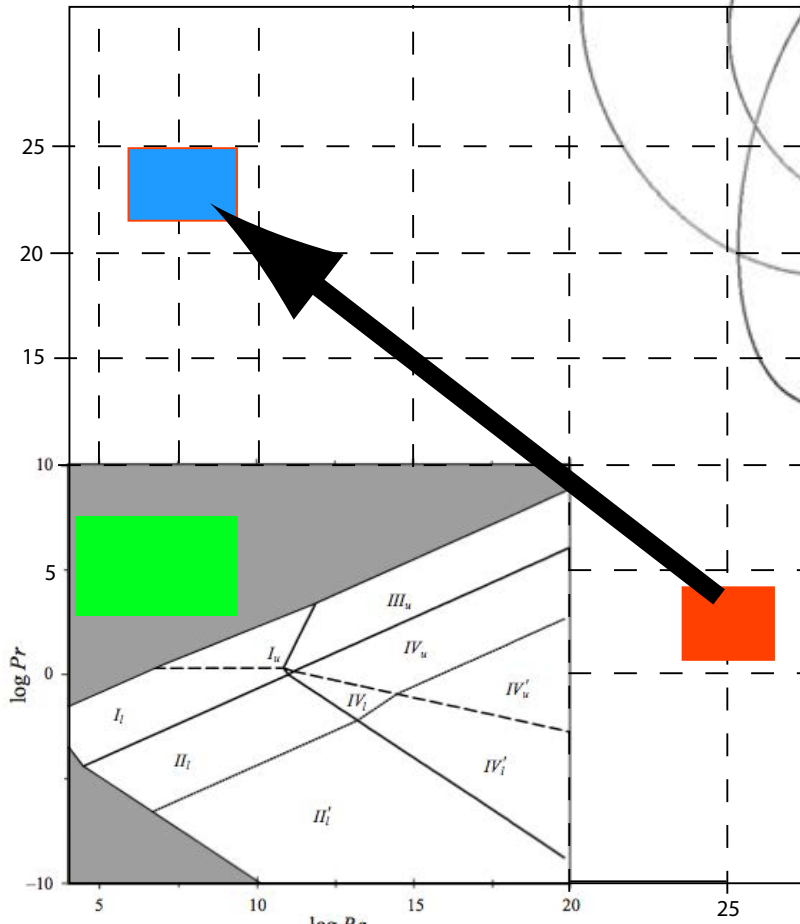
# 3-Thermal convection:



- Primitive Earth (4.5 Gyr ago):
- Magma Ocean  $Ra \sim 10^{25}$
  - Rotation (Ek=viscous/Coriolis)  
 $Ek \sim 10^{-13} - 10^{-14}$

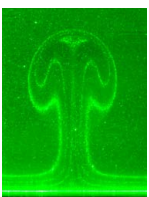
# 3-Thermal convection:

(Grossman & Lohse, 2010)

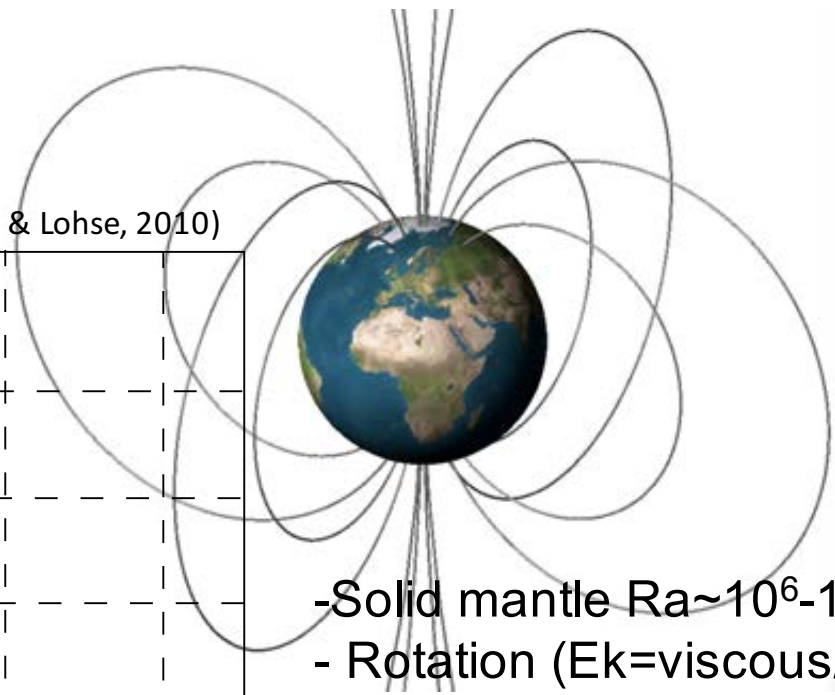


$Pr = \frac{\nu}{\kappa}$

LAB



$Ra = \frac{g \beta \Delta T h^3}{\nu \kappa}$



TODAY

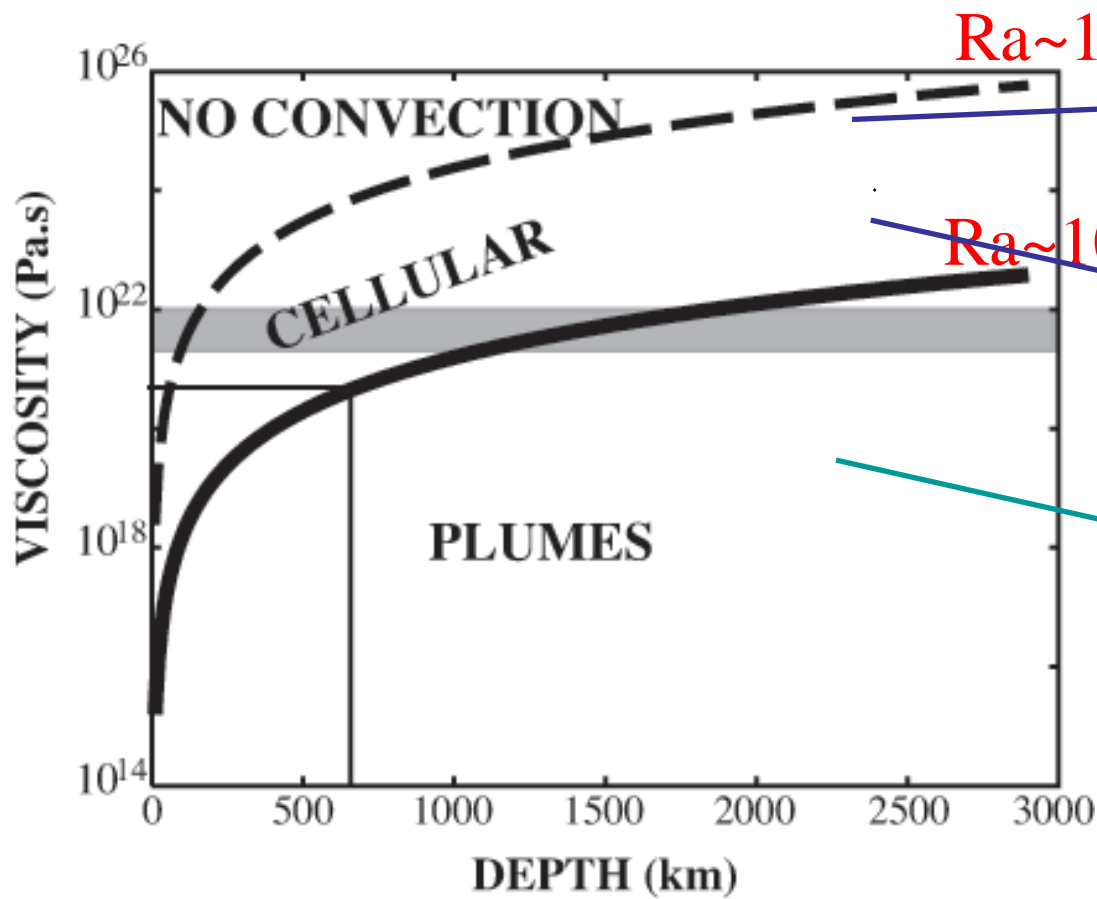
- Solid mantle  $Ra \sim 10^6 - 10^9$
- Rotation ( $Ek = \text{viscous}/\text{Coriolis}$ )  
 $Ek \sim 10^{10}$



4.5 Gyr ago

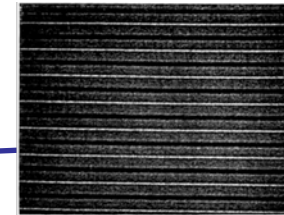
### 3.B. Convection with nearly constant viscosity

$$Ra = \frac{\alpha \cdot g \cdot \Delta T \cdot h^3}{\kappa \cdot \eta / \rho}$$

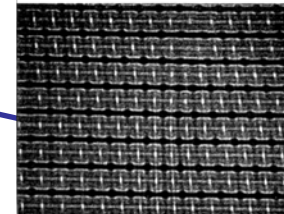
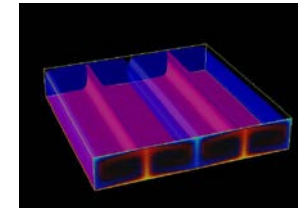


$Ra \sim 10^3$

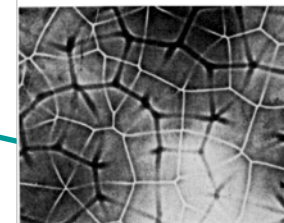
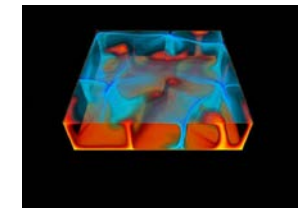
$Ra \sim 10^6$



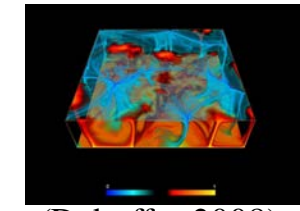
(a)



(d)

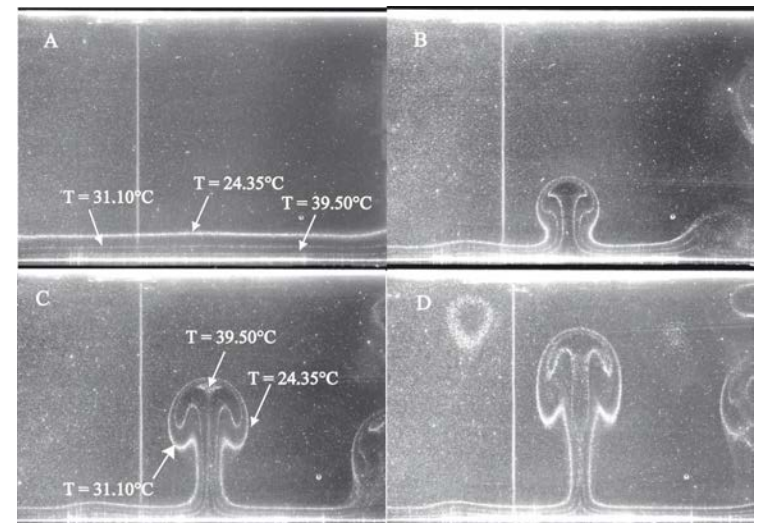
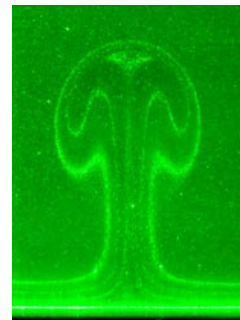


(e)

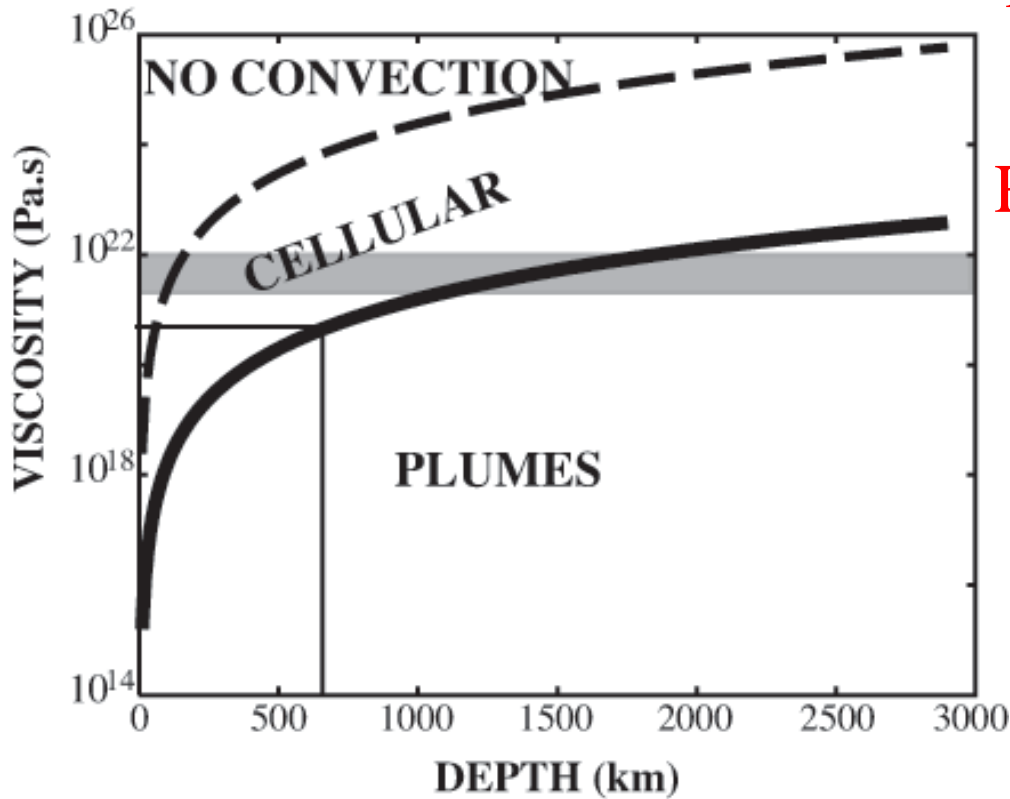


(Busse & Whitehead, 1971)

(Dubuffet, 2008)

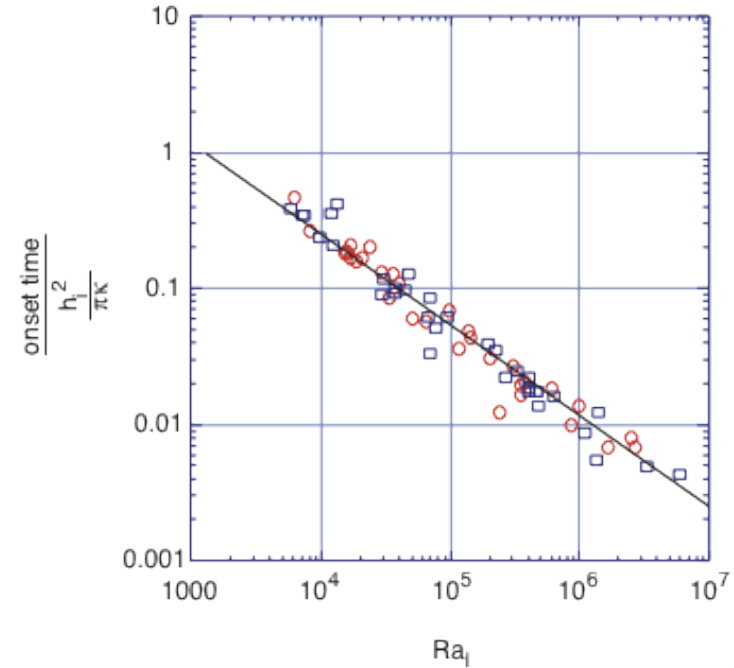


### 3.B. Convection with nearly constant viscosity



$Ra \sim 10^3$

$Ra \sim 10^6$



Hot **narrow** plumes in the mantle:  
**transient**

$$\tau_c = (H^2 / \kappa) \cdot (Ra_c / Ra_{visq})^{2/3}$$

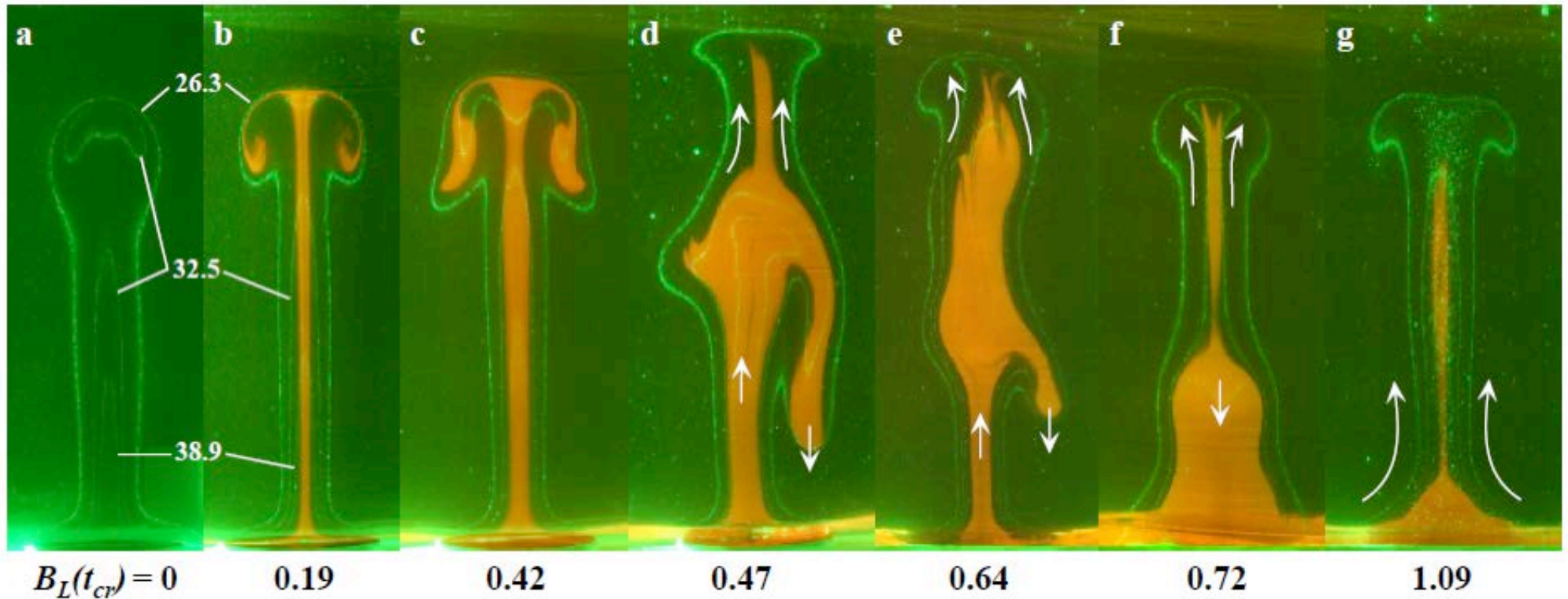
$\Rightarrow \tau_c \sim 10, 40, 200 \text{ Ma}$

for  $\eta \sim 10^{19}, 5 \cdot 10^{20}, 10^{22} \text{ Pa.s}$

**BUT** no cold plates, no LLSVPs

## 4. Interaction Convection / Denser Reservoir

# Initial Buoyancy ratio



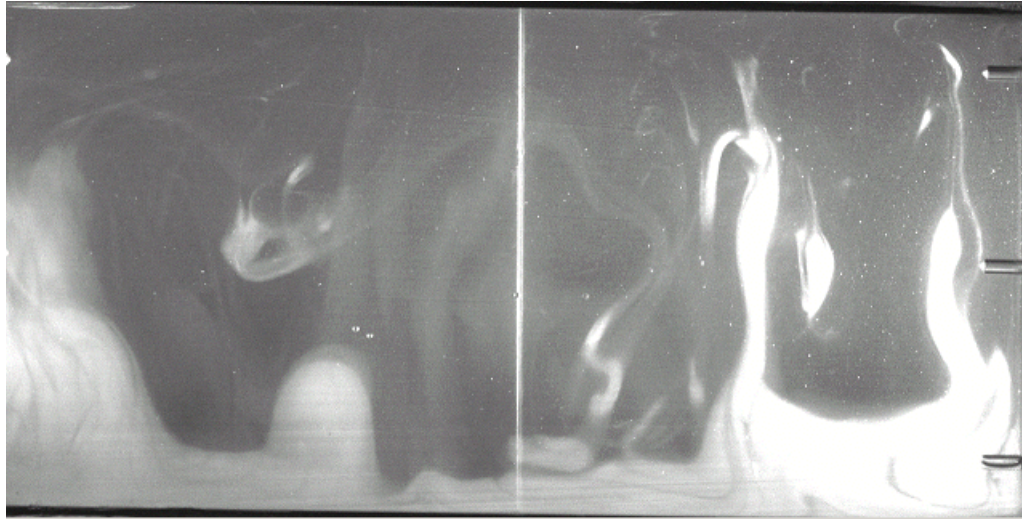
- Fat plumes
- Contorted plumes
- Hot is not rising if too dense => piles

$$B_L = \frac{\rho_{hot} - \rho_{cold}}{\rho_{cold}} \frac{H}{L}$$

(Kumagai et al, 2008)

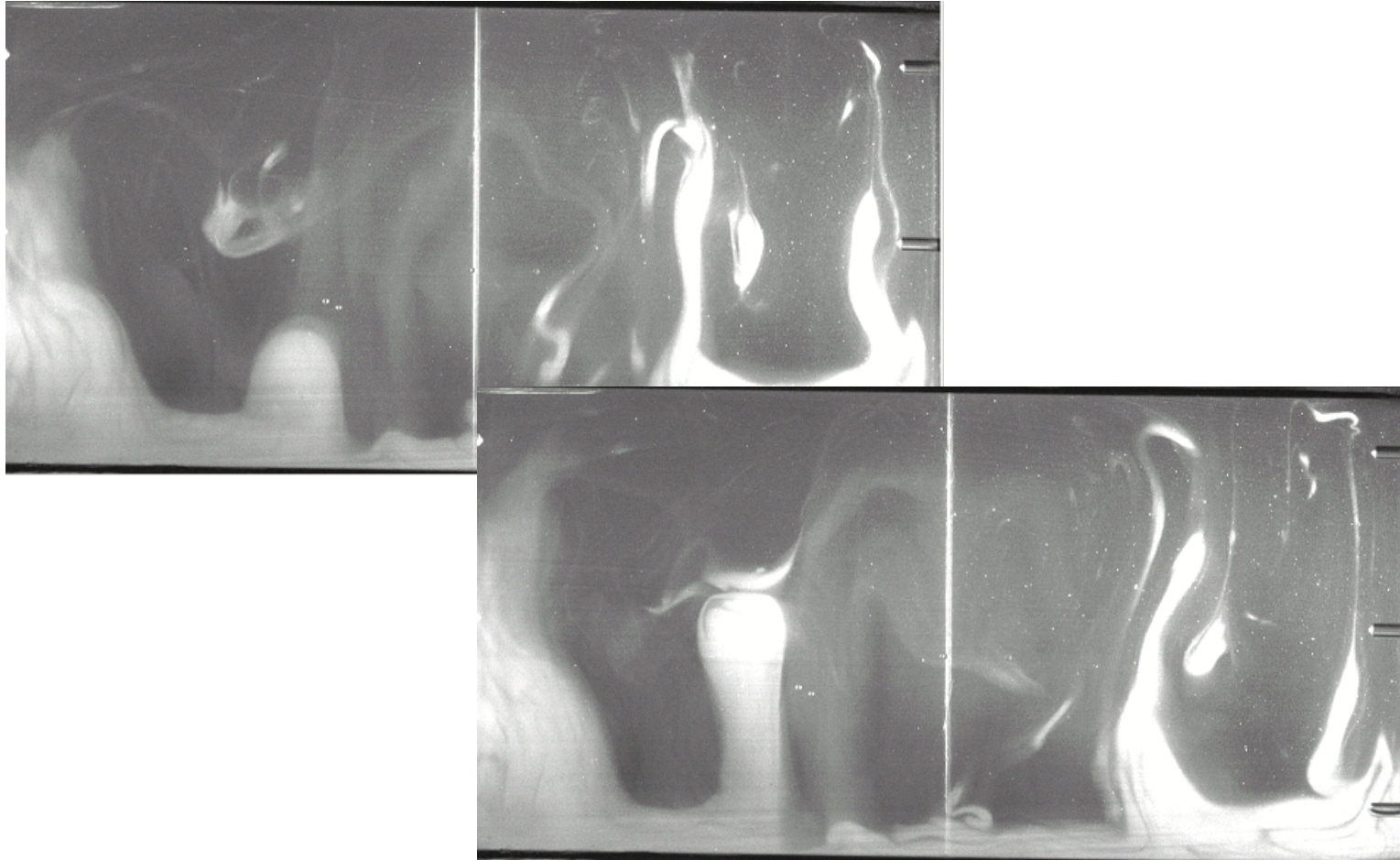
## 4. Interaction Convection / Denser Reservoir

- Convection carries fast hot material from bottom to top



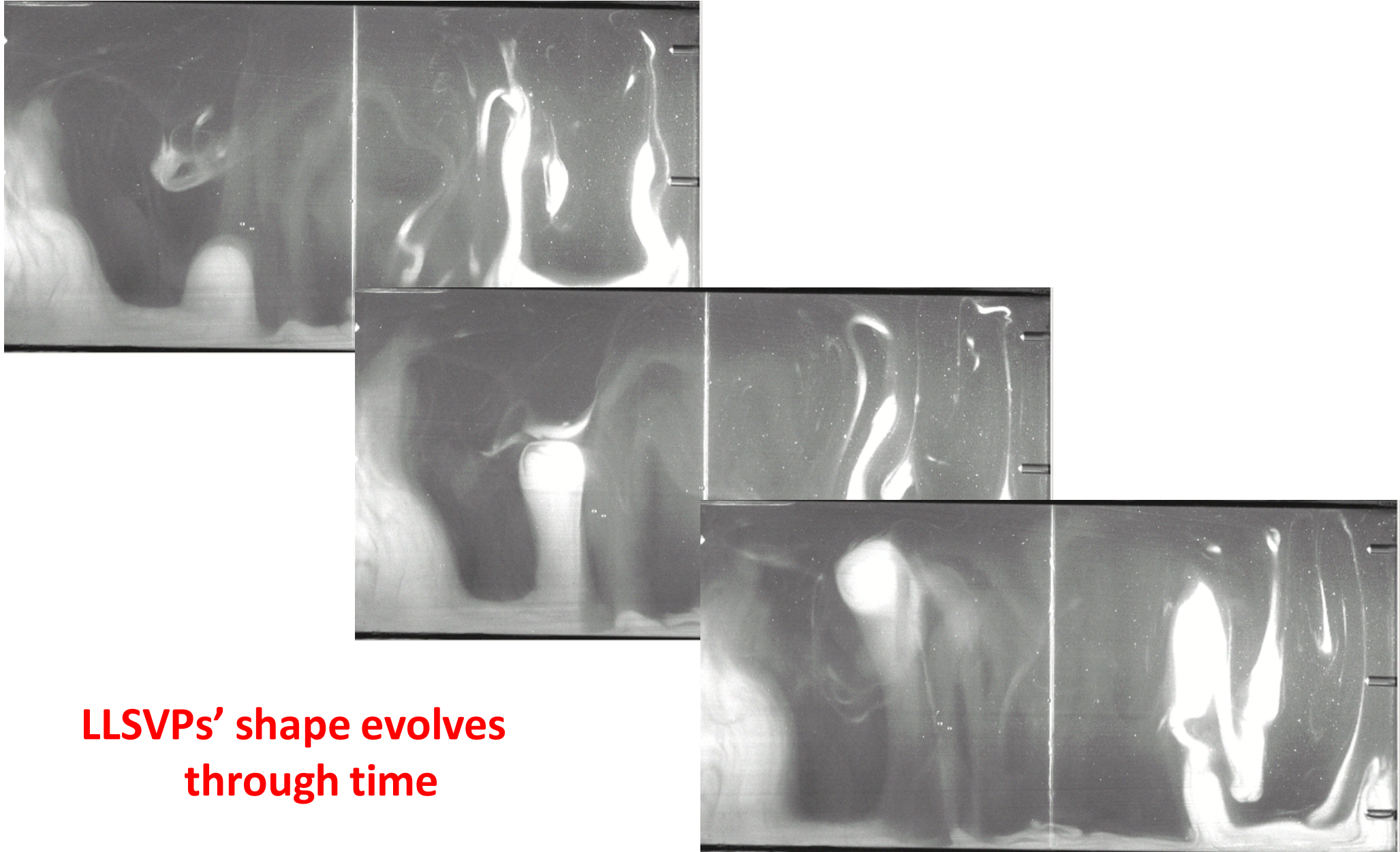
## 4. Interaction Convection / Denser Reservoir

- Convection carries fast hot material from bottom to top



## 4. Interaction Convection / Denser Reservoir

- Convection carries fast hot material from bottom to top





# 4. Interaction Convection / Denser Reservoir

## Stability:

$$- Ra = \frac{\sqrt{g \Delta \rho} d^3}{\nu \kappa}$$

$$- B = \frac{\Delta \rho_{ch}}{\Delta \rho}$$

## Morphology:

$$- a = d_1/d_2$$

$$a < 2\%$$

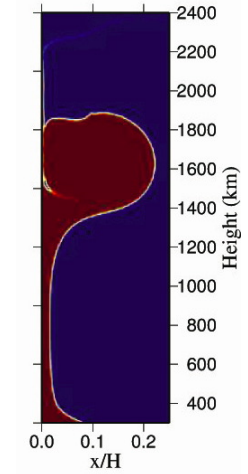
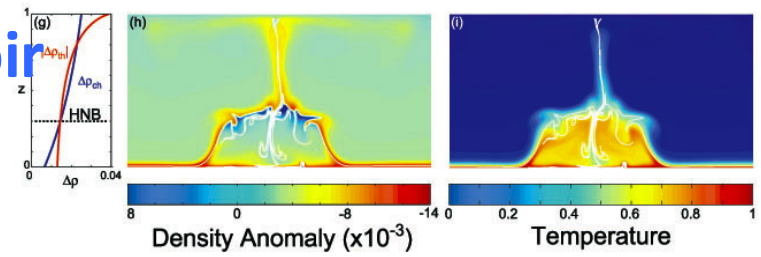
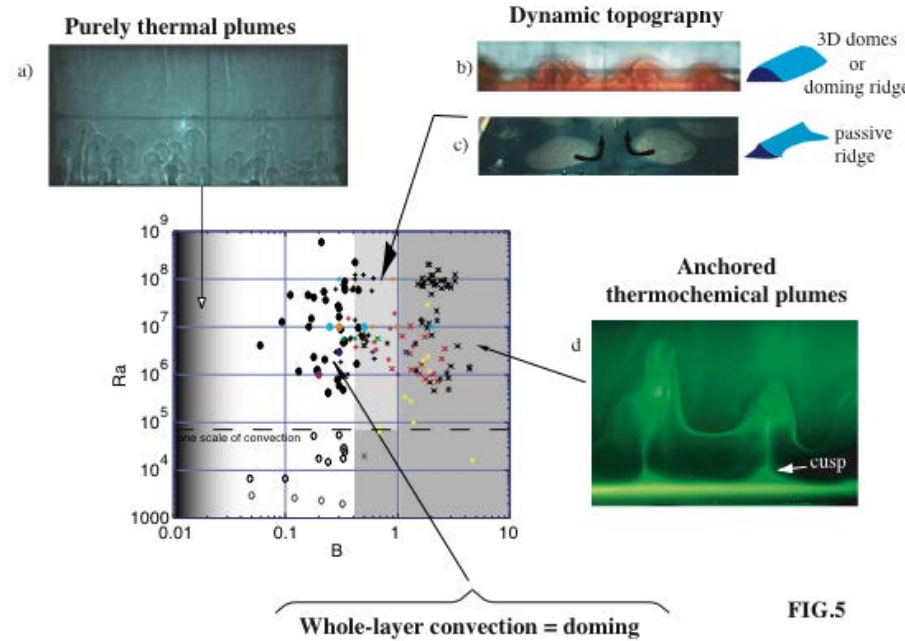
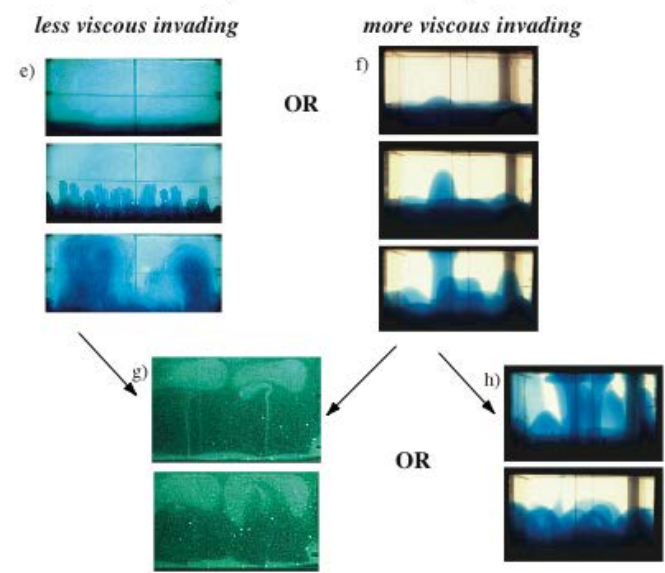


FIG.5

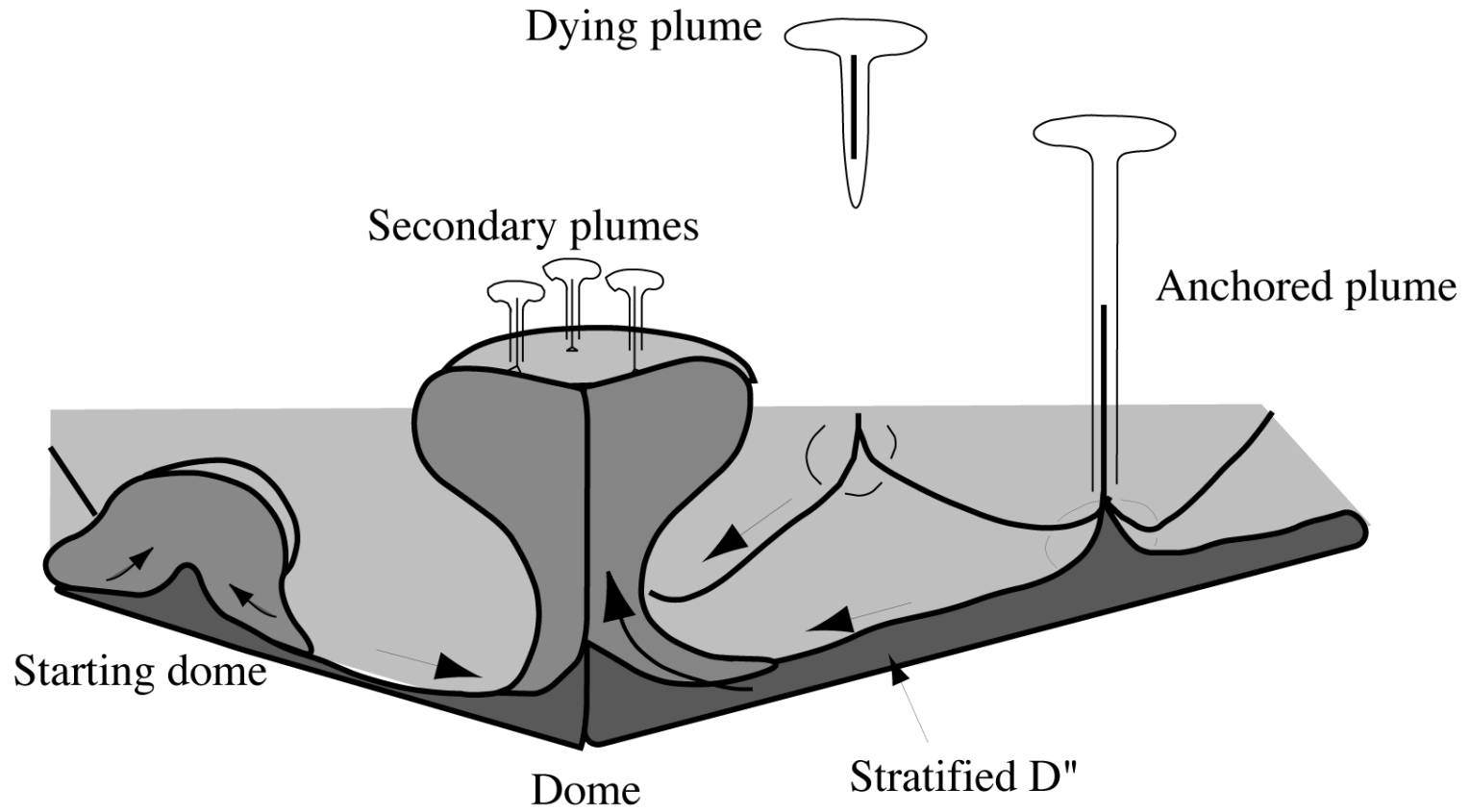
Tan & Gurnis, 2005, 2007;  
Samuel & Bercovici, 2006



(Davaille & al, 2005)

Richter & McKenzie, 1981;  
Olson, 1984;  
Schmeling, 1988;  
Olson & Kincaid, 1991;  
Tackley, 1998; 2002;  
Davaille, 1999; Davaille & al, 2002;  
Kellogg & al, 1999;  
Montague & Kellogg, 2000;  
Hansen & Yuen, 2000;  
Le Bars & Davaille, 2002, 2004;  
Jellinek & Manga, 2002, 2004;  
Samuel & Farnetani, 2002, 2005;  
McNamara & Zhong, 2004, 2005  
Lin & Van Keken, 2005

# Several types of plumes



# 4. Interaction Convection / Denser Reservoir

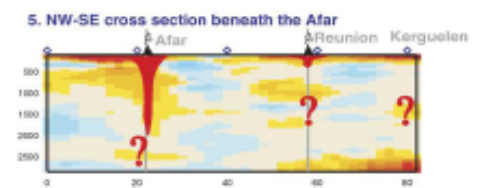
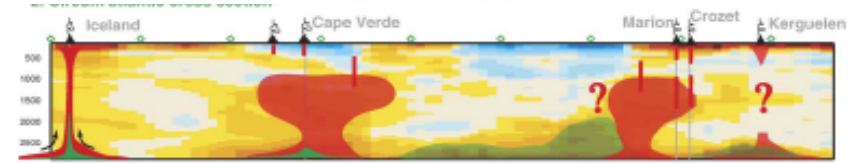
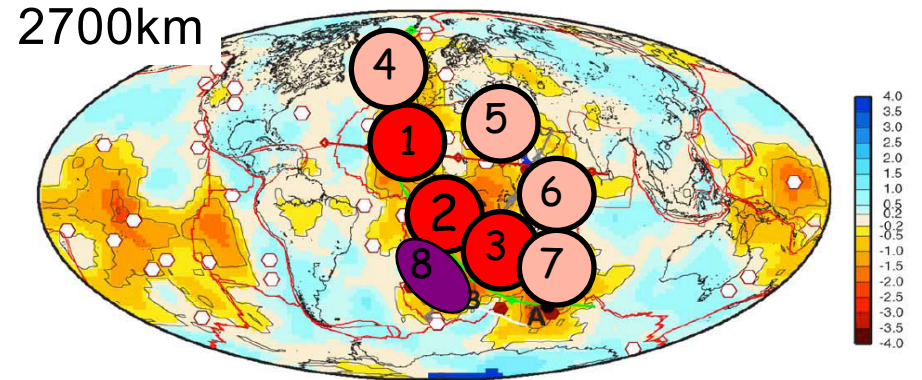
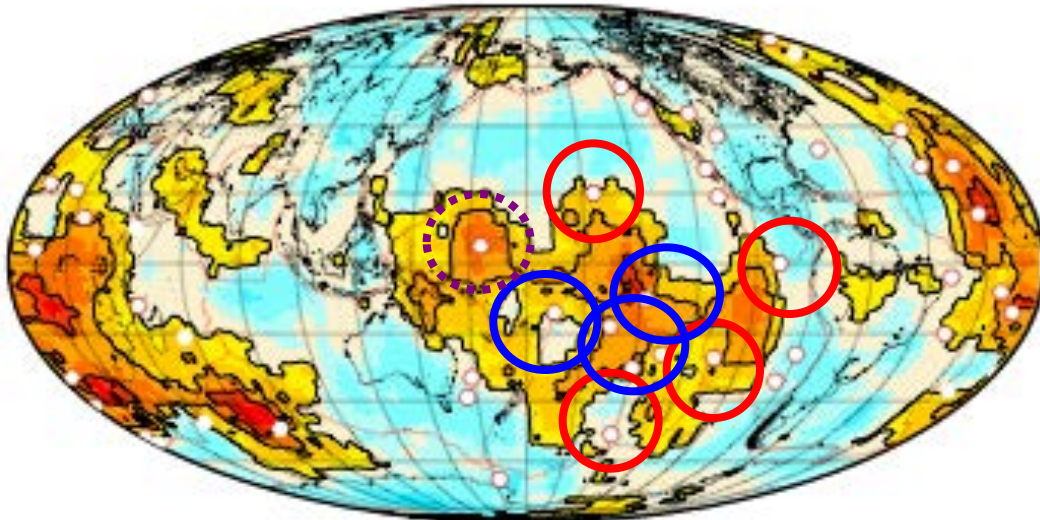
$-\Delta\rho_x/\rho \sim 0.1-0.6\%$

$-h_{\text{dense}} < 700 \text{ km}$

-spacing at CMB  $\sim 3000-4000 \text{ km}$

-recurrence time  $\sim 100-200 \text{ Myr}$

-5 to 9 thermochemical hot instabilities  
 long-lived hotspots, traps, superswells  
 large-scale slow seismic anomalies

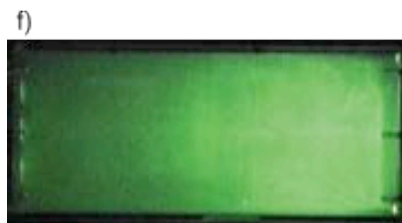
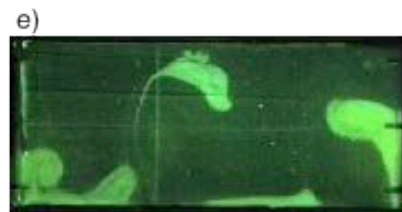
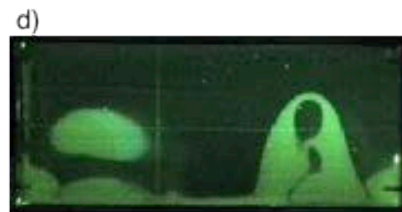
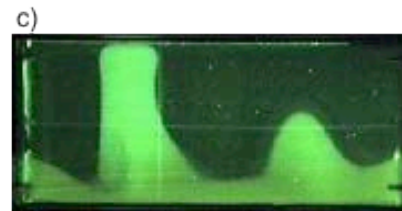
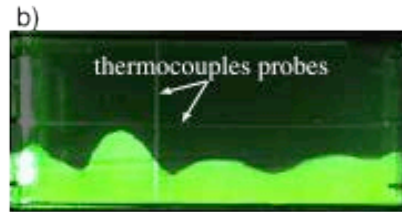
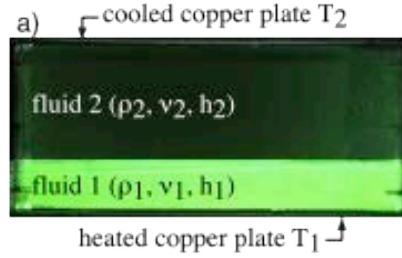


Davaille et al., 2005

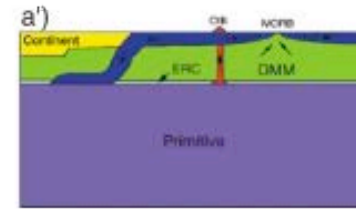
# 4. Interaction Convection / Denser Reservoir

MANTLE  
TEMPORAL  
EVOLUTION?

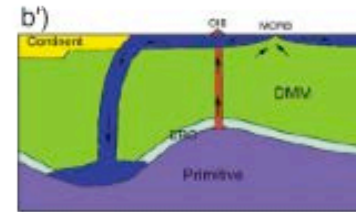
Today?



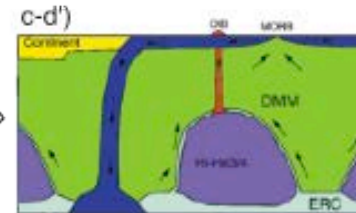
(LeBars and Davaille, JGR 2004)



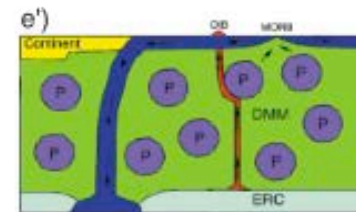
2-layer



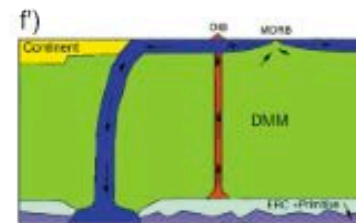
Kellogg & al, 1999



Tackley, 1998  
Davaille 1999

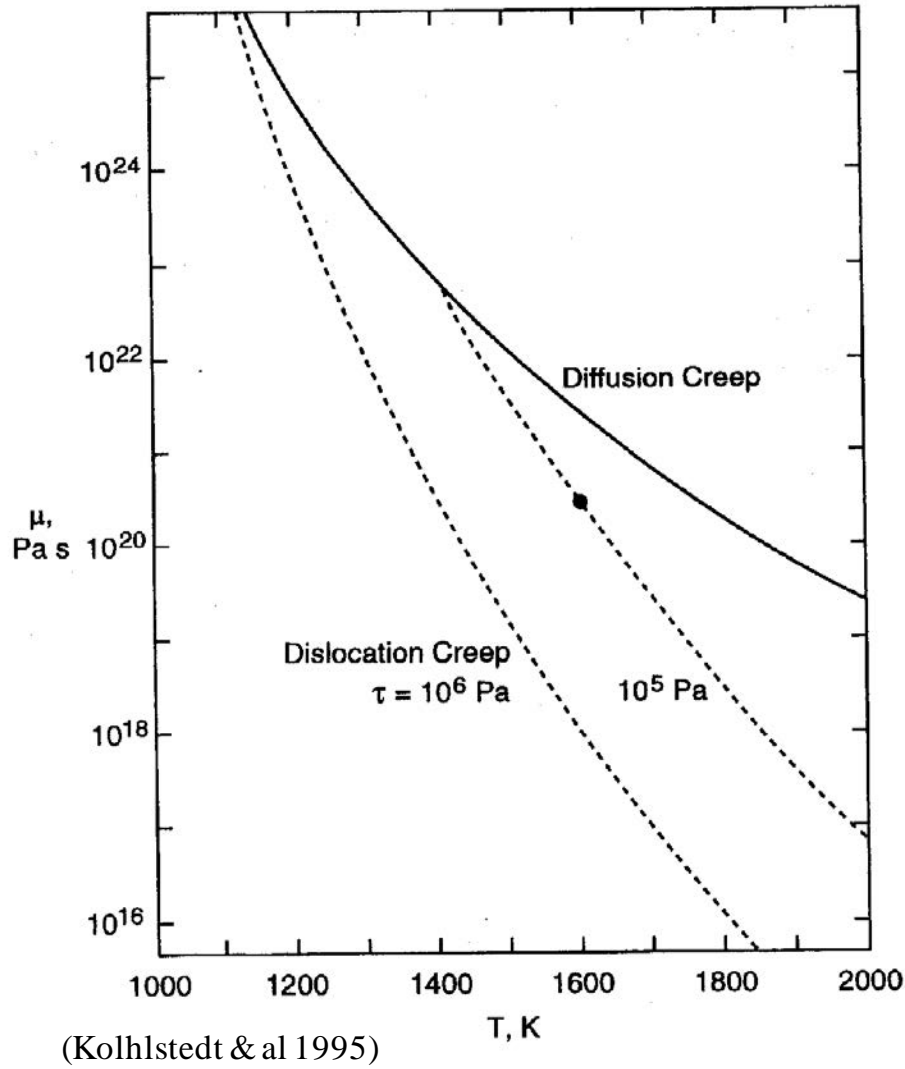


Davies 1984,  
Becker & al 1999



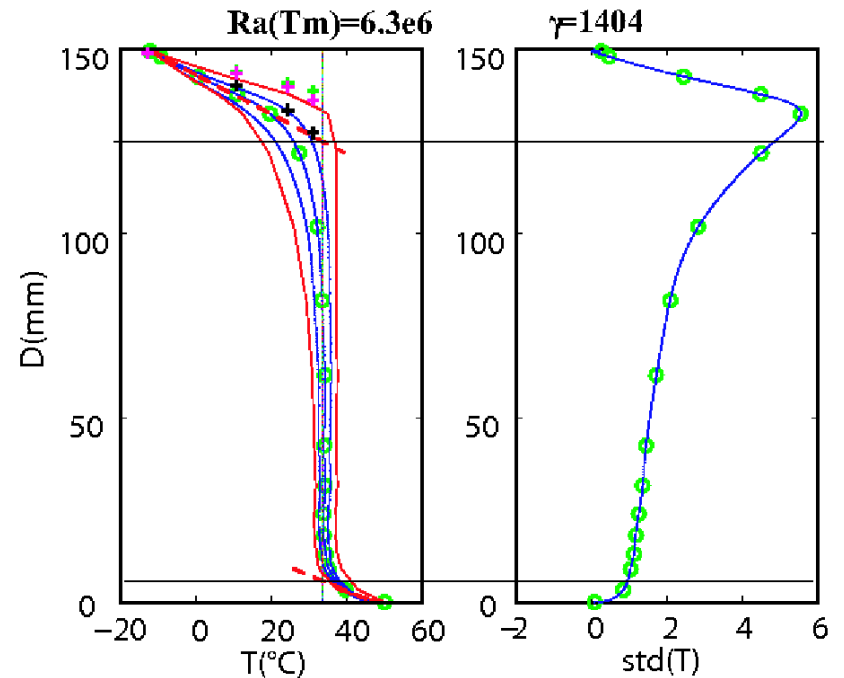
1-layer

# 5- Making a plate: Convection with strongly $T_p$ -dependent viscosity



**=> Asymmetry hot/cold**

 Viscosity Ratio

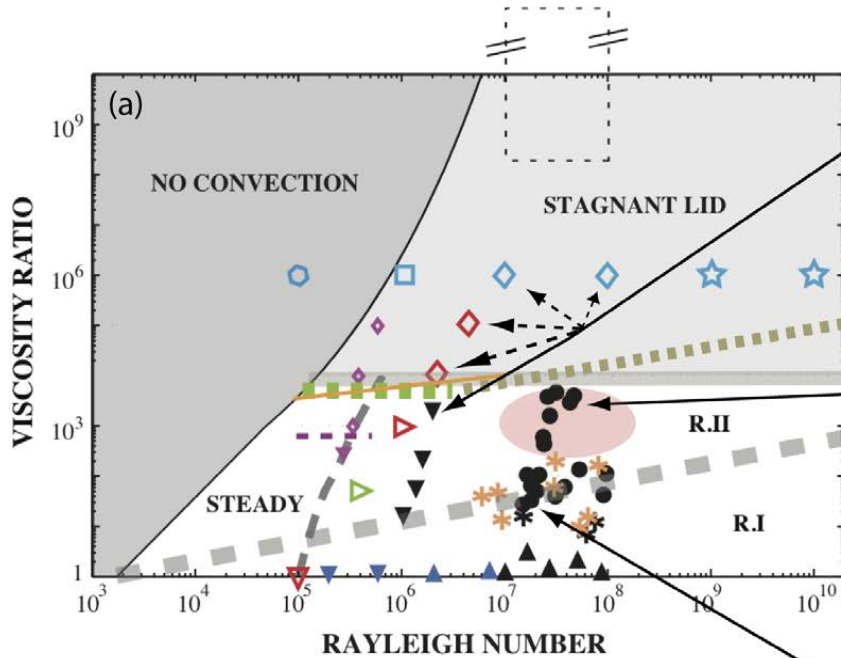
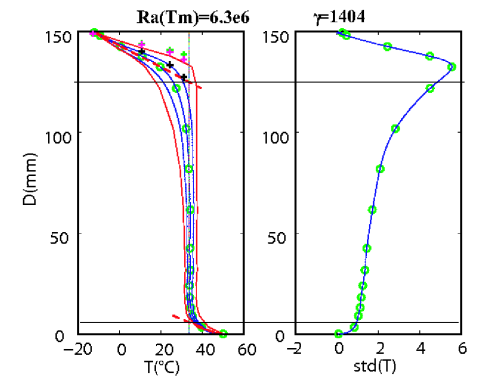


# 5- Making a plate: Convection with strongly $T_p$ -dependent viscosity

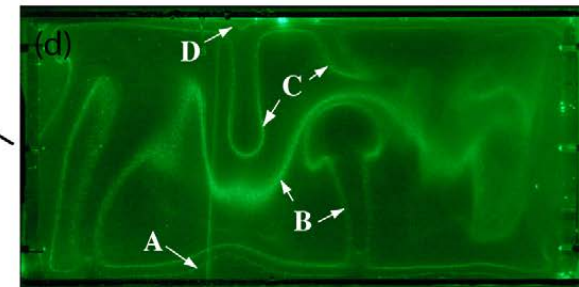
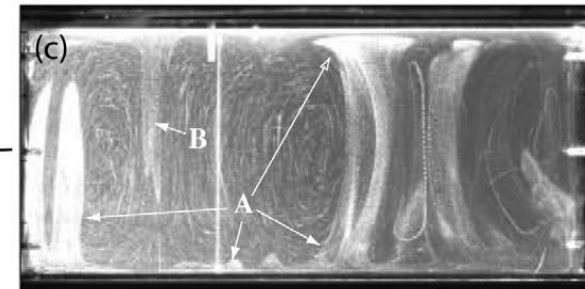
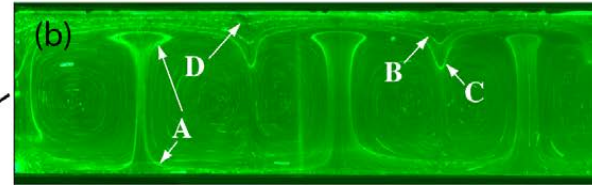
Richter, 1973, 1976  
 Booker, 1976  
 Stengel & al, 1982  
 Richter & al, 1983  
 White, 1988  
 Weinstein & Christensen, 1991  
 Davaille & Jaupart, 1993

Giannandrea & Christensen, 1993  
 Manga et al, 1998-2001  
 Shaeffer & Manga, 2001  
 Davaille et Vatteville, 2005  
 Androvandi et al, 2011  
 Christensen & Harder, 1991  
 Tackley, 1993, 1996

Ratcliff & al, 1997  
 Trompert & Hansen, 1998  
 Kageyama & Ogawa, 2000  
 Stemmer & al, 2006  
 Weber et al, 2017  
 Morris, 1984  
 Solomatov, 1995

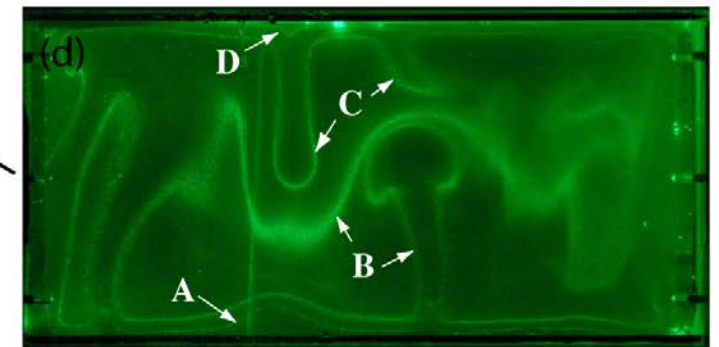
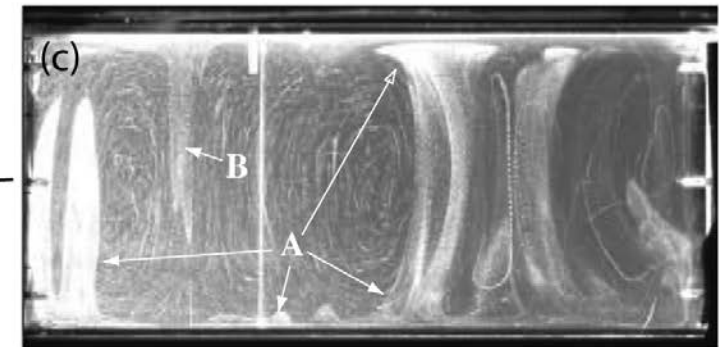
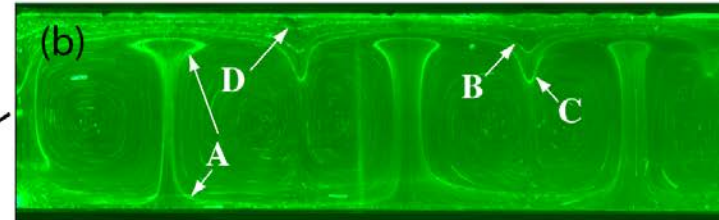
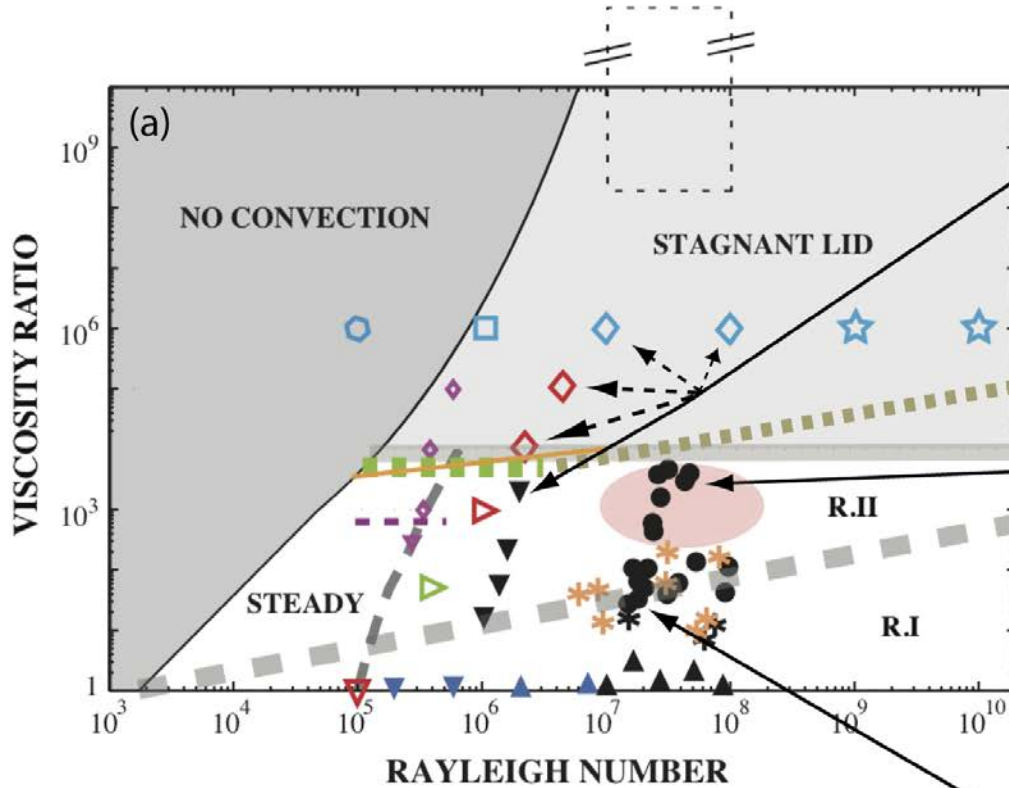


- No convection
- Steady 1 to 1 cells
- ◇ Time-dependent 1 to 1 cells
- ☆ Plumes
- ▼ Spokes = time-dependent 1 to 1 cell
- ▲ Plumes, hot = cold
- ▽ Spokes in L-L: 1 to 3 cell
- \* Plumes, hot ≠ cold
- Several hot plumes for 1 cold cell



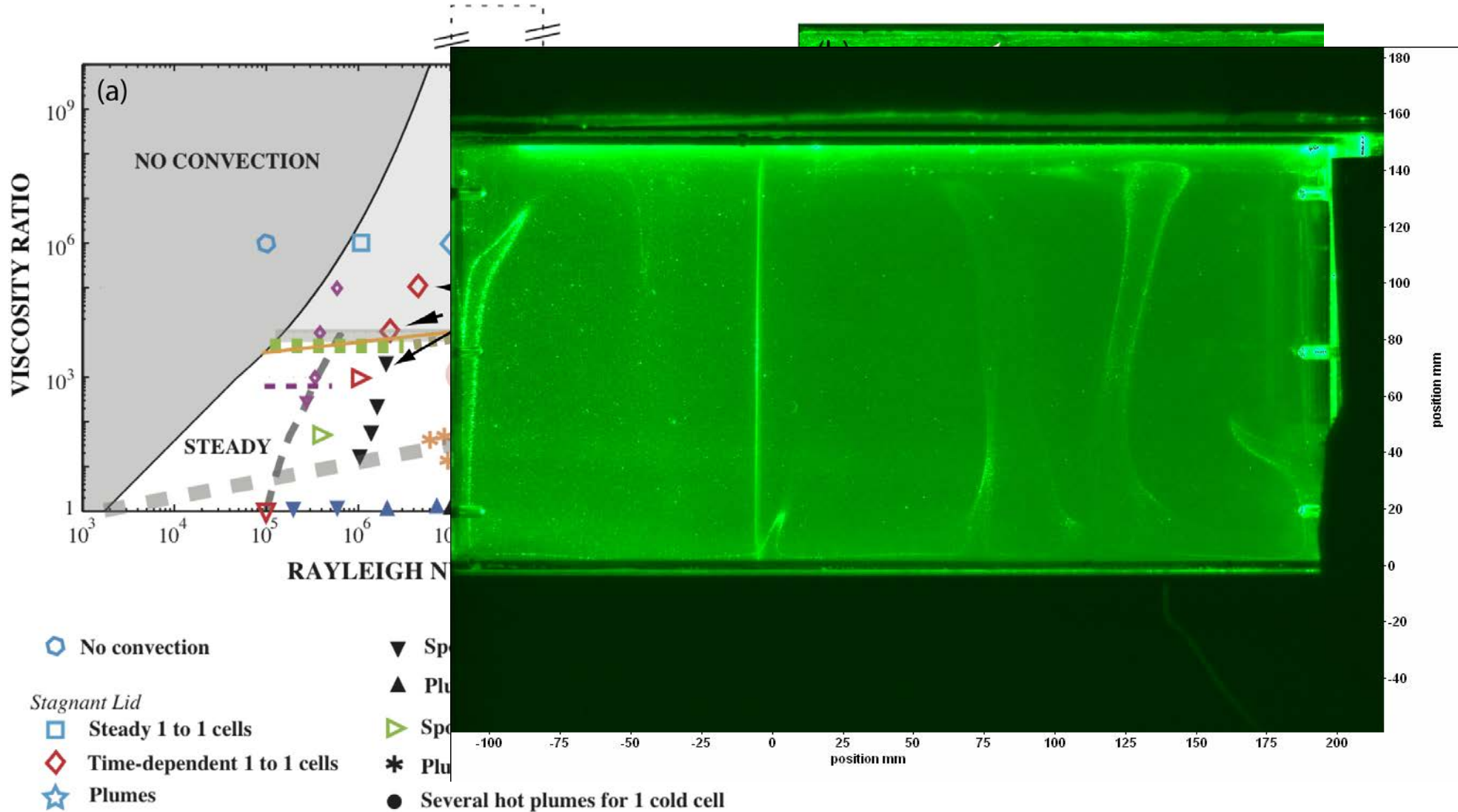
# 5- Making a plate: Convection with strongly $T_p$ -dependent viscosity

High viscosity ratio  $\Rightarrow$  Plate



- No convection
  - ▼ Spokes = time-dependent 1 to 1 cell
  - ▲ Plumes, hot = cold
  - Steady 1 to 1 cells
  - ◇ Time-dependent 1 to 1 cells
  - ☆ Plumes
  - ▽ Spokes in L-L: 1 to 3 cell
  - \* Plumes, hot  $\neq$  cold
  - Several hot plumes for 1 cold cell
- Stagnant Lid*

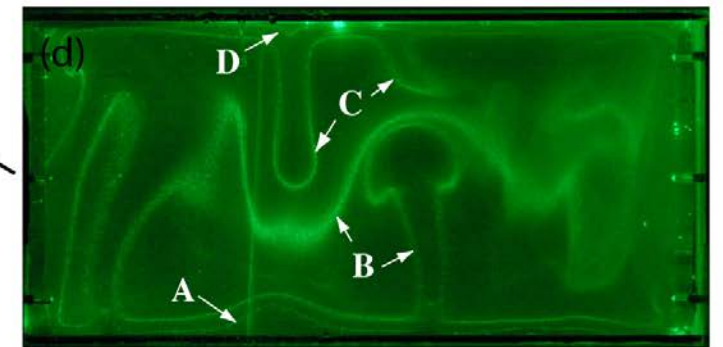
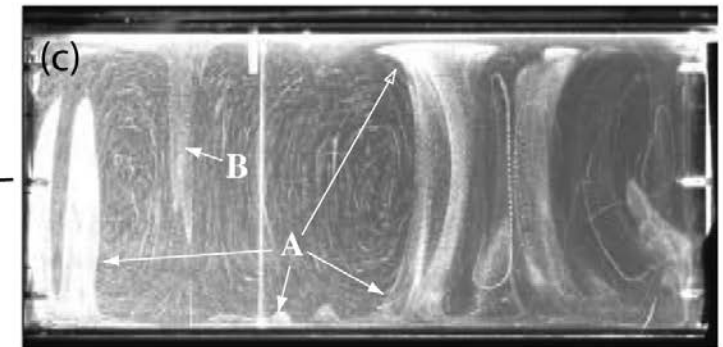
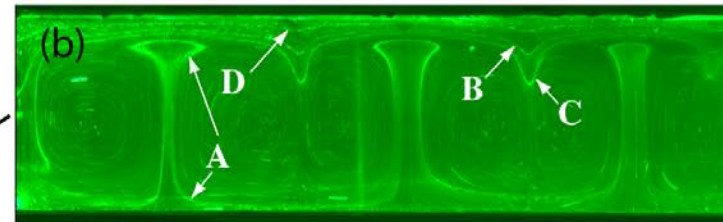
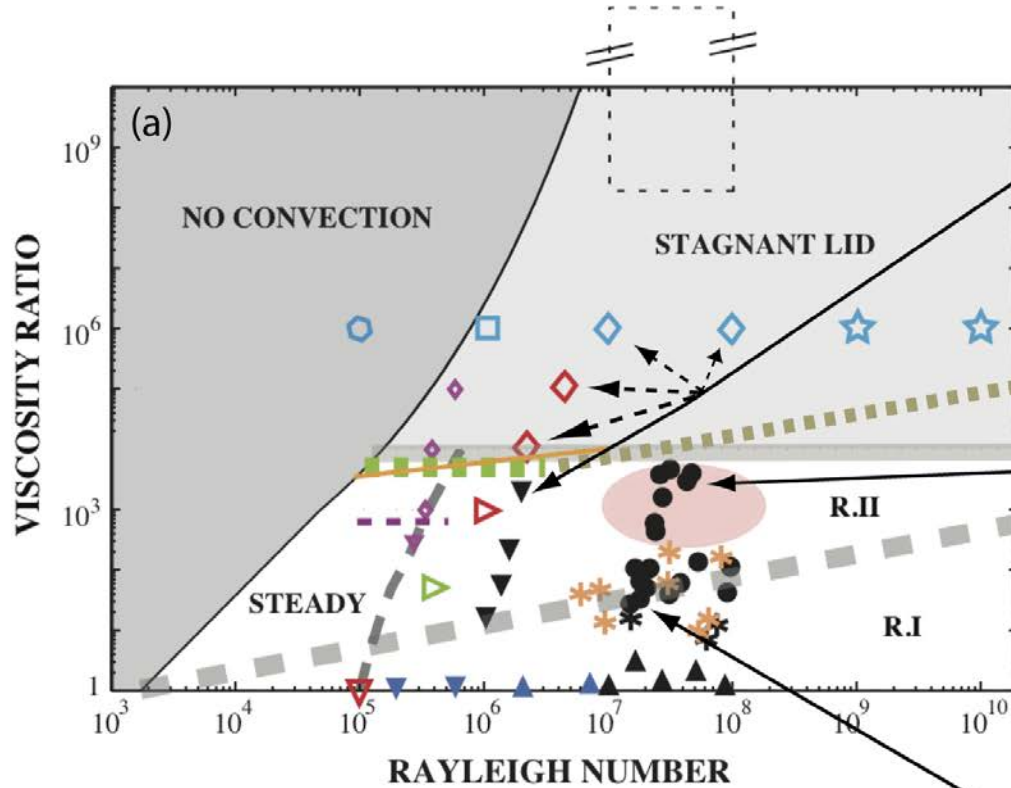
# 5- Making a plate: Convection with strongly $T_p$ -dependent viscosity





# 5- Making a plate: Convection with strongly $T_p$ -dependent viscosity

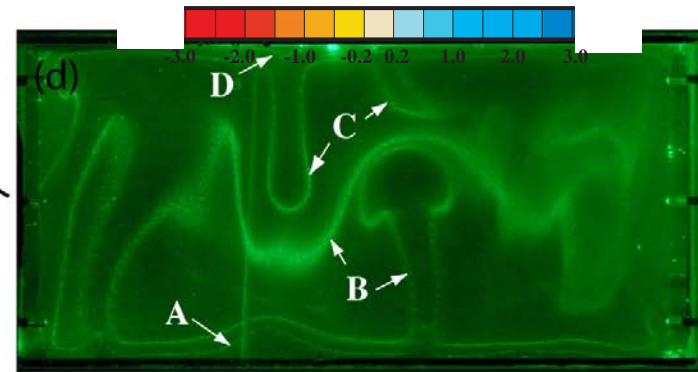
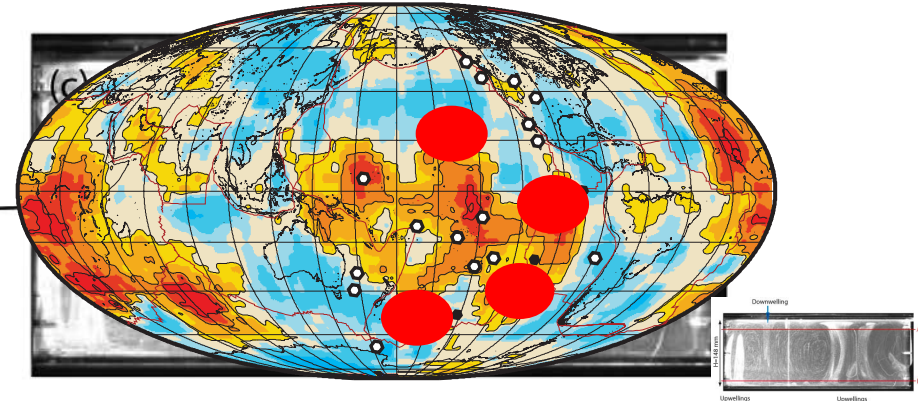
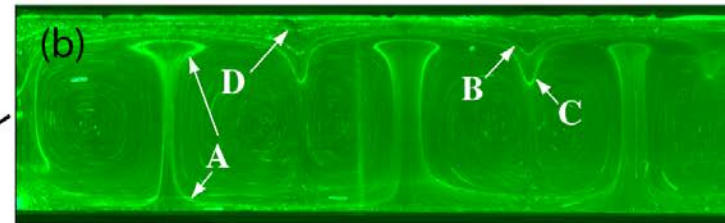
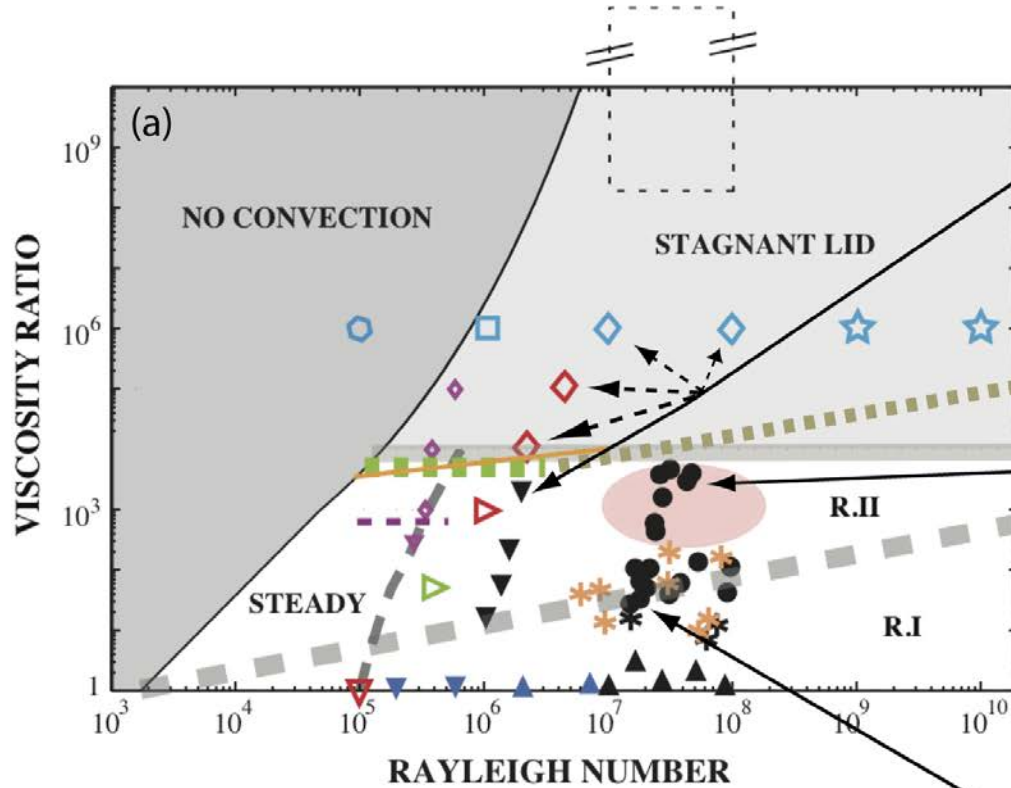
**Strong influence of downwellings  
on plume stability and generation area**



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- \* Plumes, hot  $\neq$  cold
- Several hot plumes for 1 cold cell

# 5- Making a plate: Convection with strongly $T_p$ -dependent viscosity

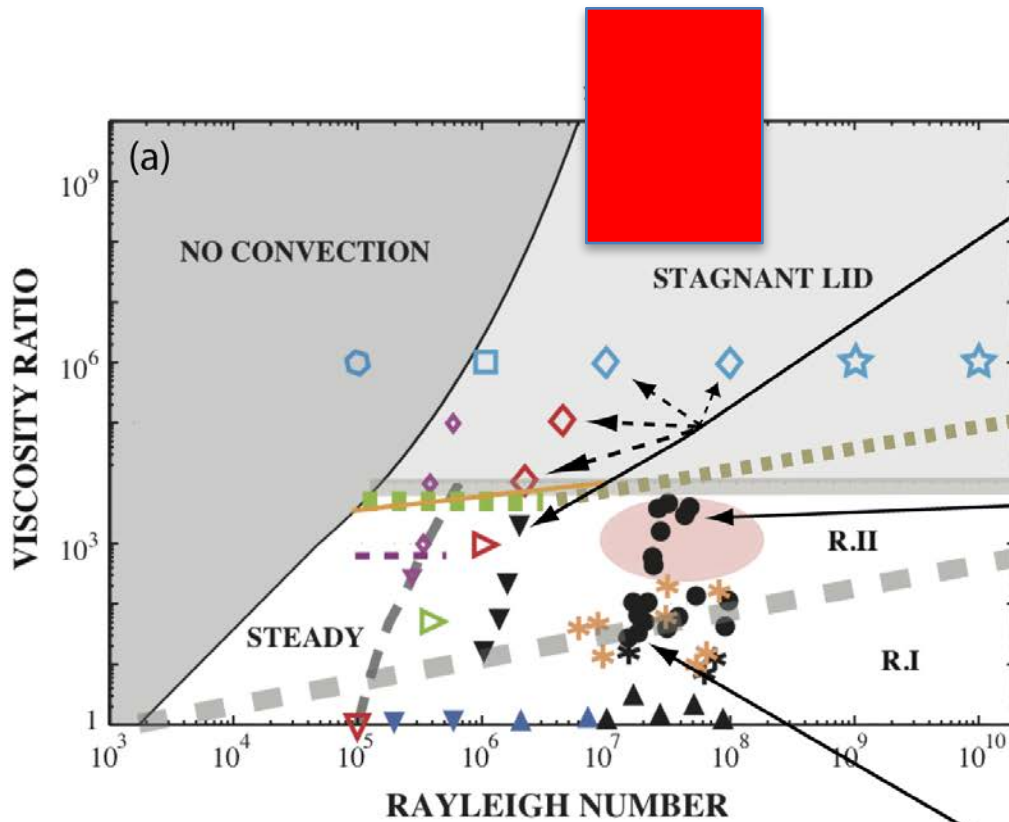
**Strong influence of downwellings on plume stability and generation area**



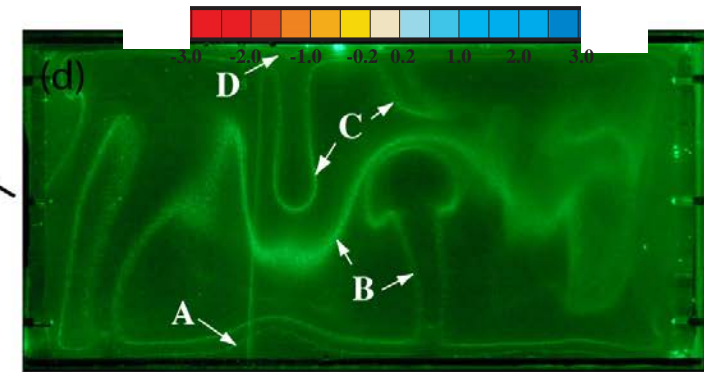
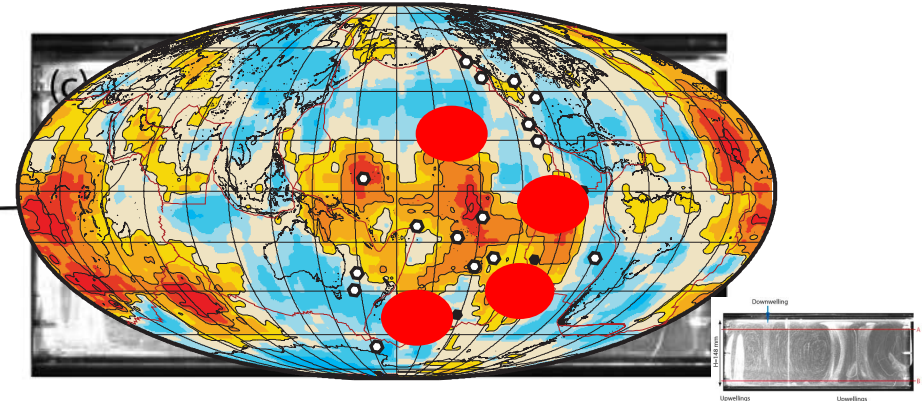
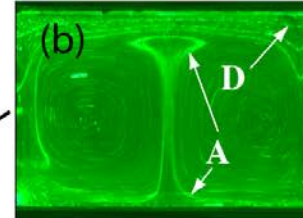
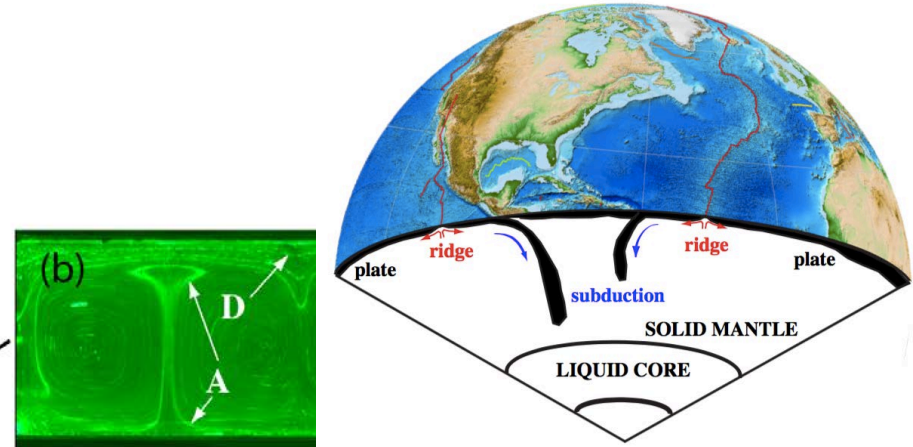
- ◊ No convection
- ◻ Steady 1 to 1 cells
- ◊ Time-dependent 1 to 1 cells
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- ▼ Spokes = time-dependent 1 to 1 cell
- ▲ Plumes, hot = cold
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- \* Plumes, hot  $\neq$  cold
- Several hot plumes for 1 cold cell

# 5- Making a plate: Convection with strongly $T_p$ -dependent viscosity

=> How to break the lid ?



- ◊ No convection
- ◻ Steady 1 to 1 cells
- ◊ Time-dependent 1 to 1 cells
- ☆ Plumes
- ▼ Spokes = time-dependent 1 to 1 cell
- ▲ Plumes, hot = cold
- ◄ Spokes in L-L: 1 to 3 cell
- \* Plumes, hot  $\neq$  cold
- Several hot plumes for 1 cold cell



## 6- Non-newtonian rheology: « Kick Hard to Break »

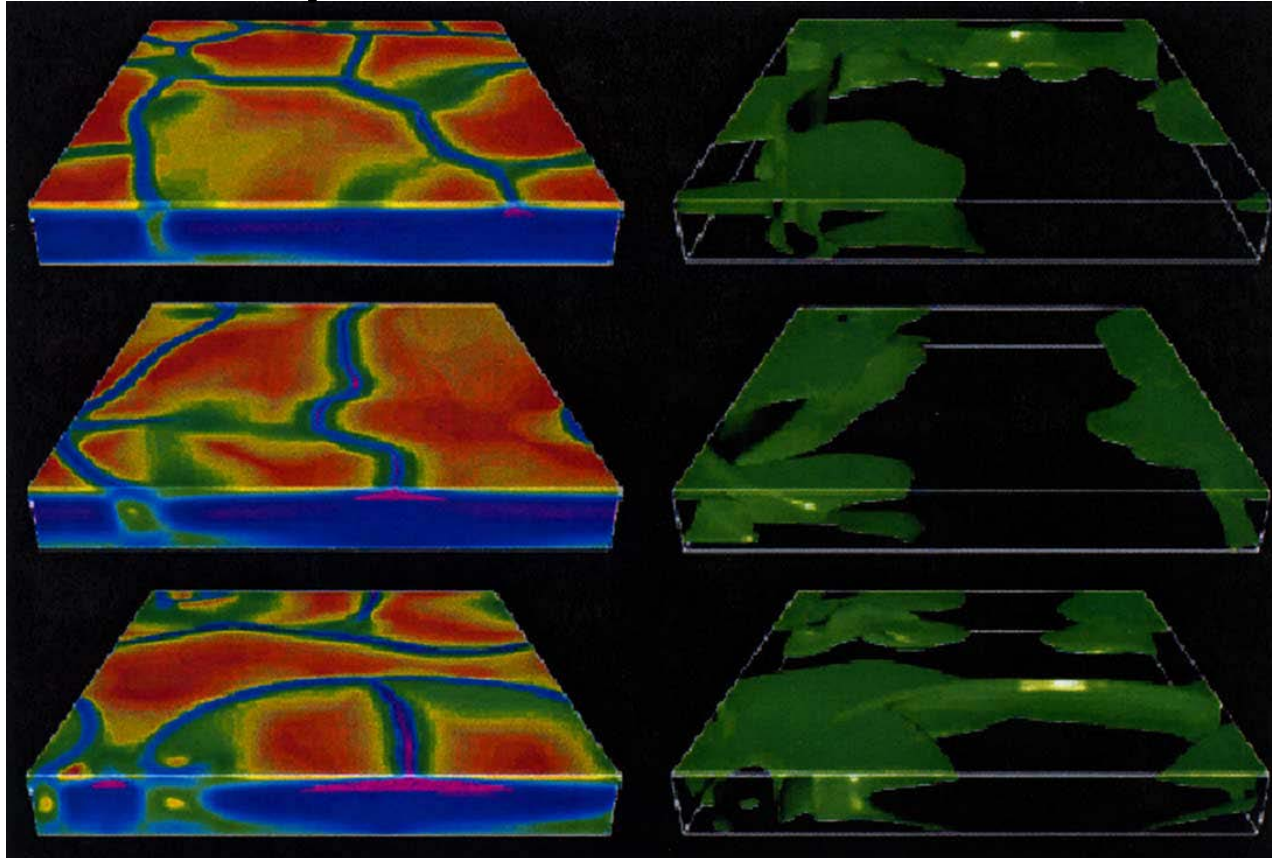
Ex: Yield stress + viscosity decrease under the plates due to melting

(Tackley 2000)

+ shear heating , + pre-existing weaknesses , + ...

Viscosity field

Cold isotherm



*Moresi and Solomatov, 1998*

*Regenauer-Lieb et al., 2001*

*Solomatov, 2004*

*Stein & al, 2004*

*Enns et al., 2005*

*Stegman et al., 2006*

*Ueda et al., 2008*

*Gerya et al., 2008*

*Thielmann & Kaus, 2014*

...

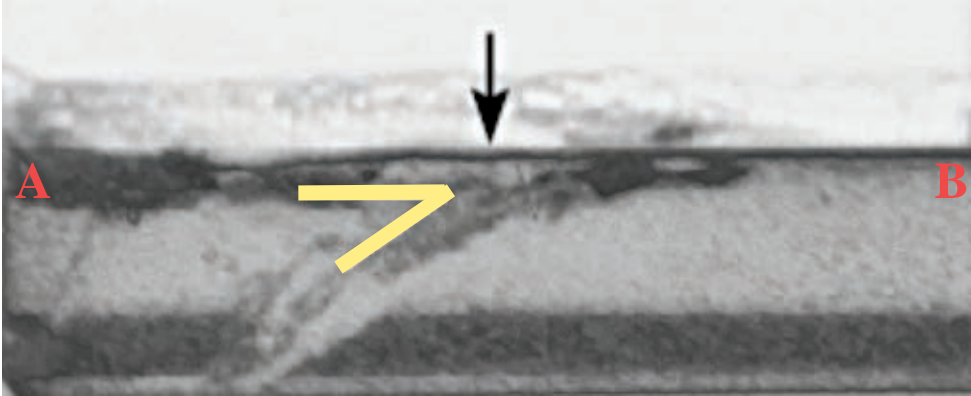
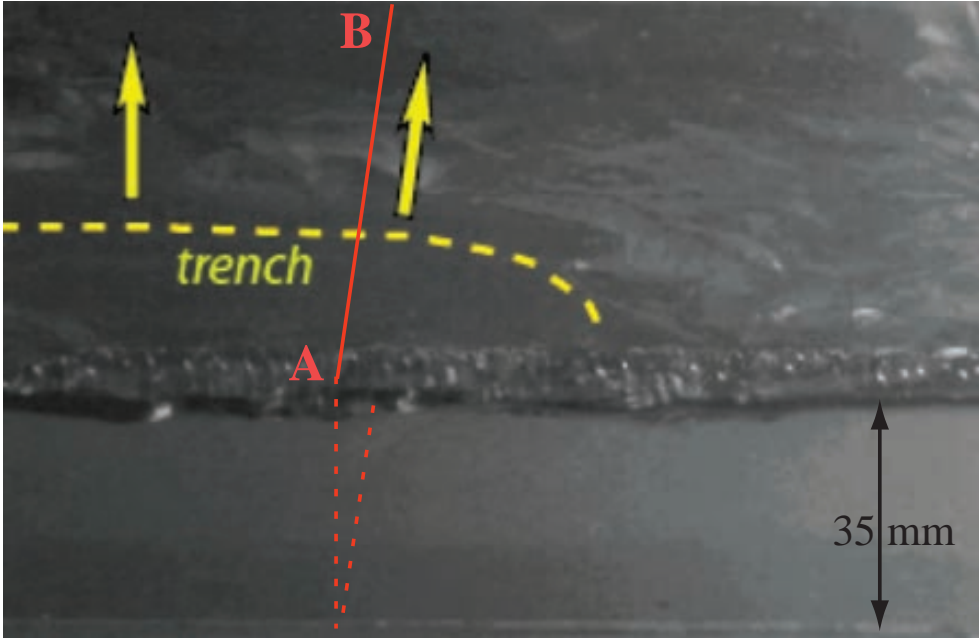
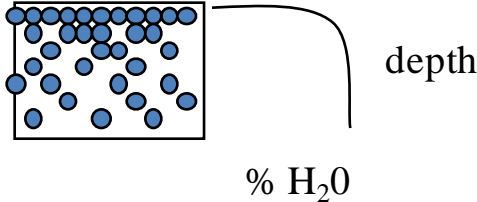
.Localize deformation => plates

.Instantaneous rheology

=> require yield stress much lower than measured...

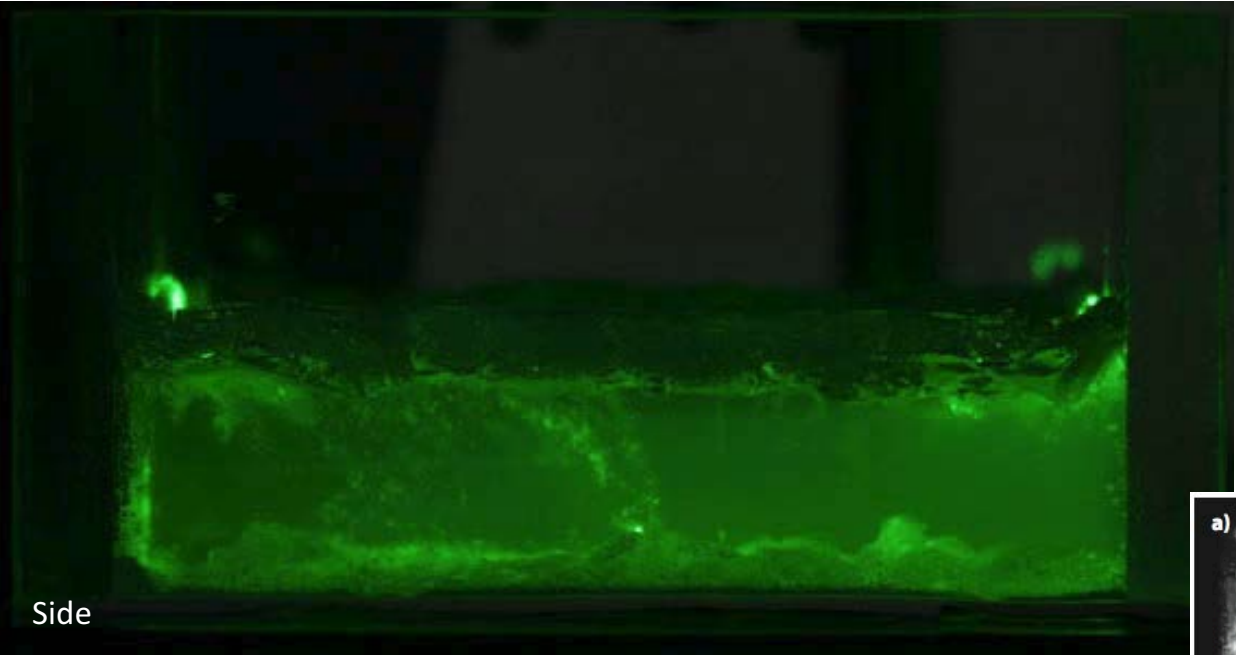
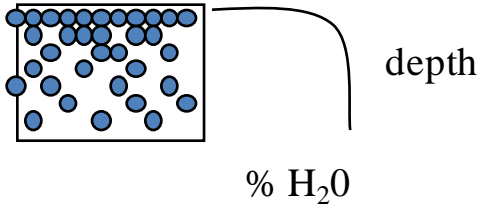
# 6- Non-newtonian rheology:

Can we learn from Soft Matter material ?

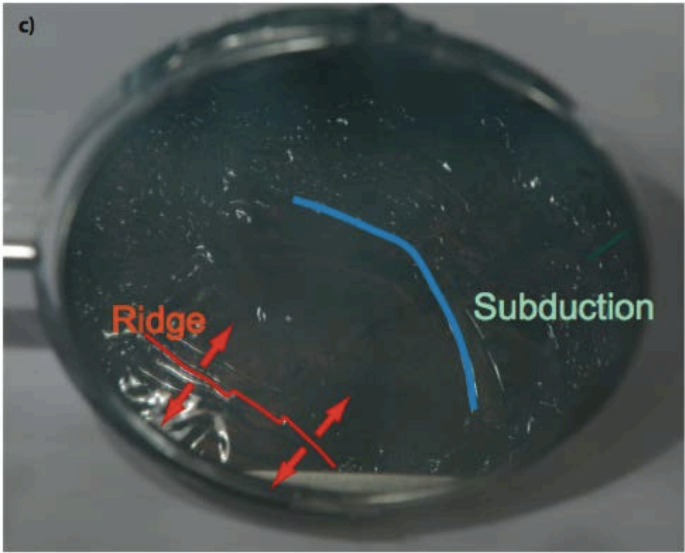
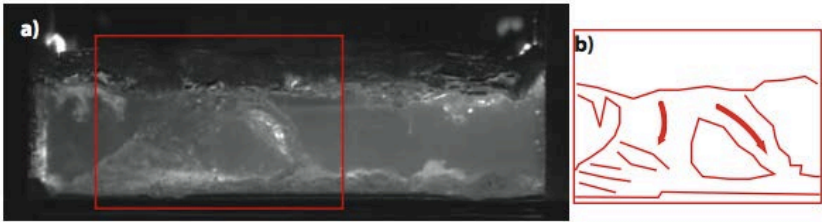


(Di Giuseppe et al, 2011)

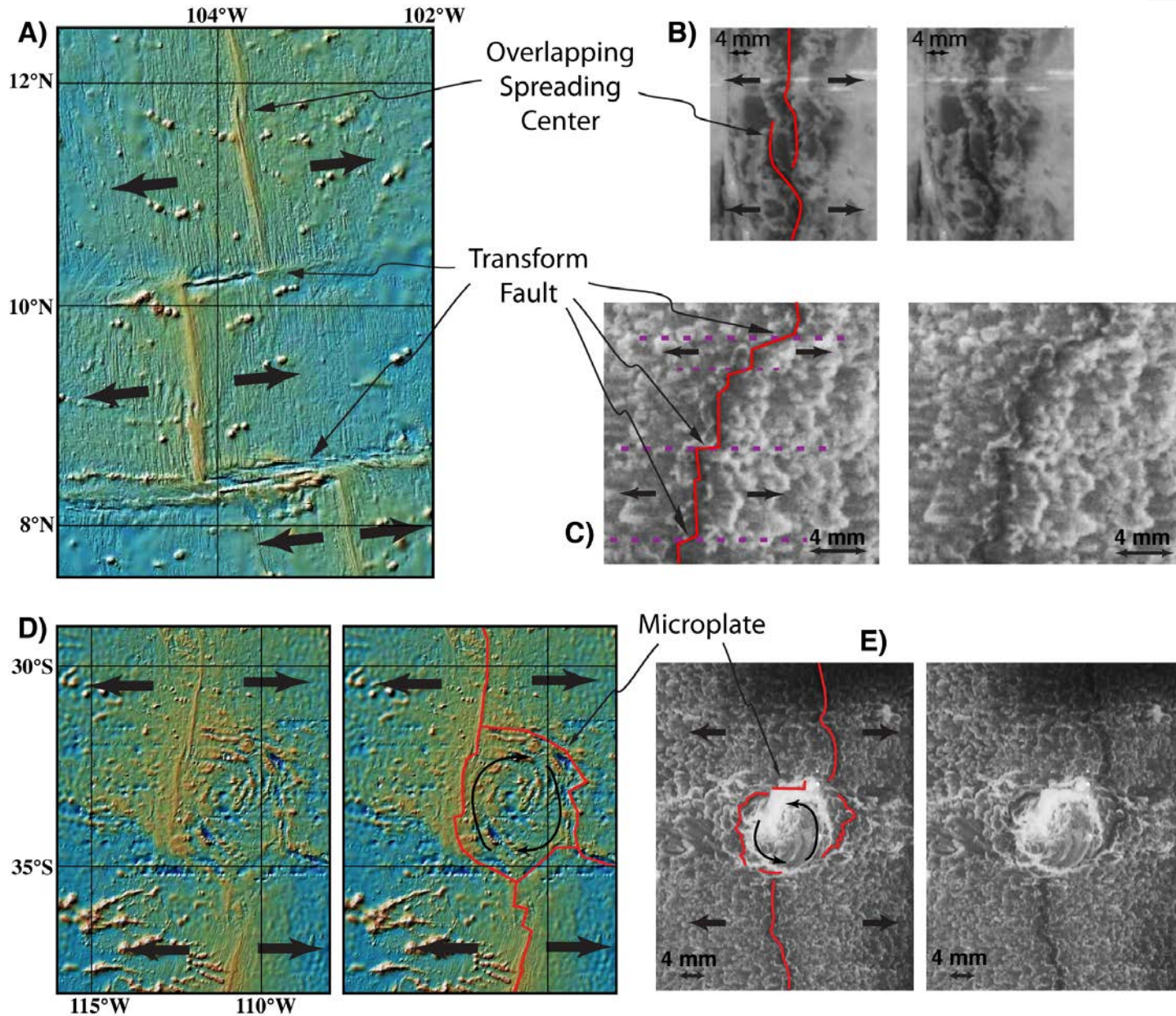
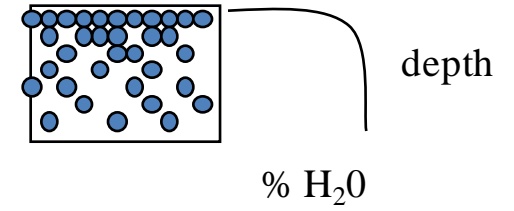
# 6- Non-newtonian rheology:



- Ridges and subduction.
- Several plates coexist.
- Plumes



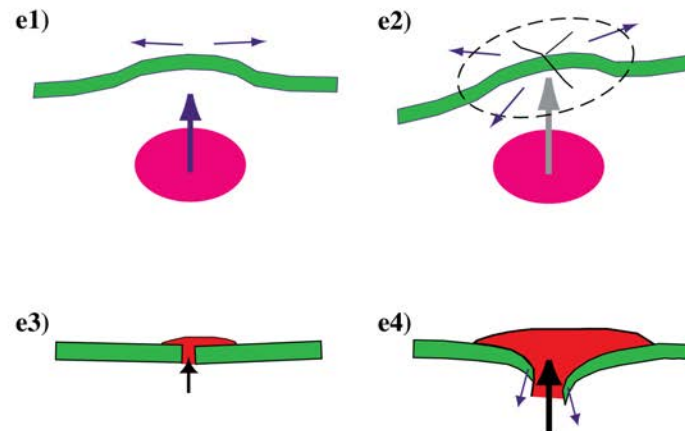
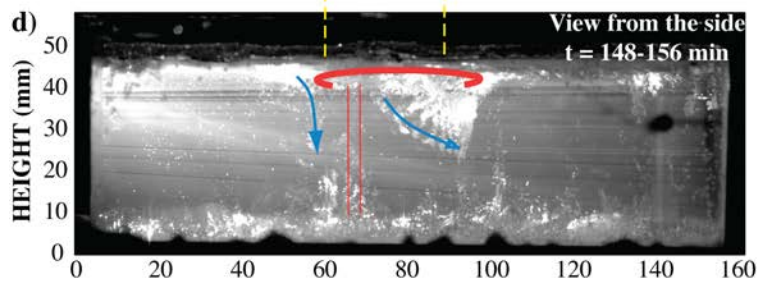
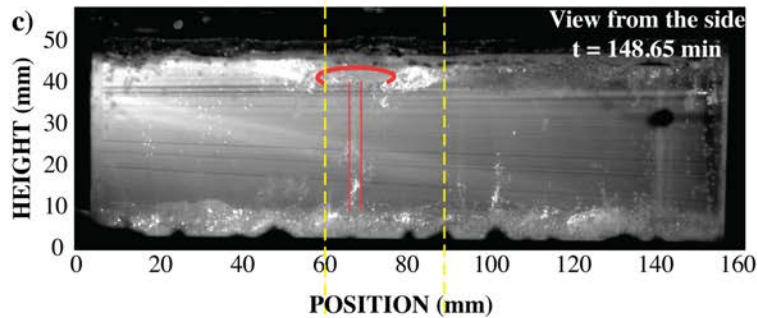
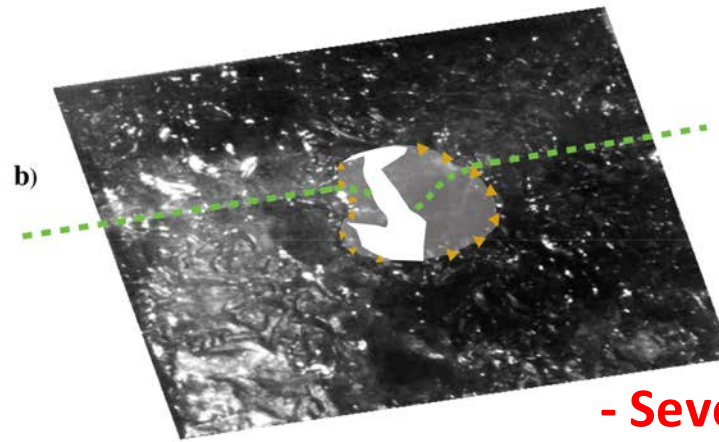
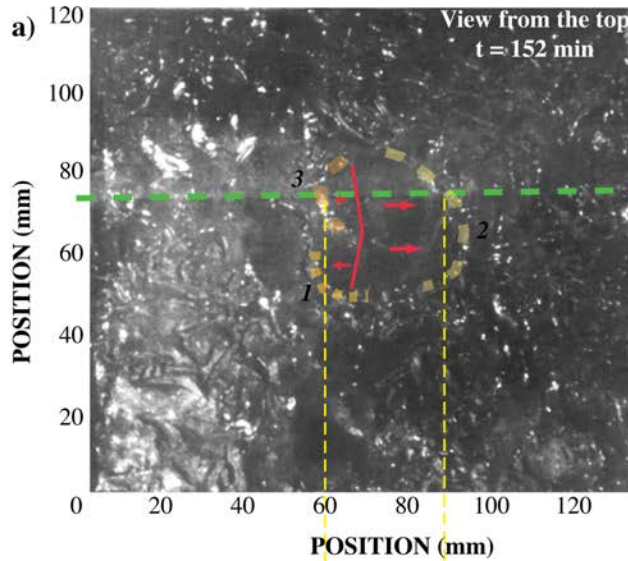
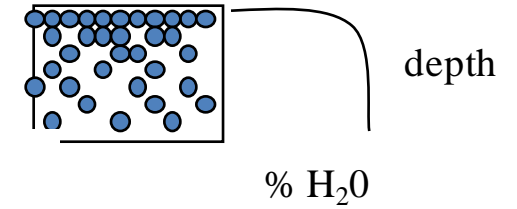
# 6- Non-newtonian rheology:



(Sibrant et al, 2018)

# 6- « Kick Hard to Break » :

## Plume-induced subduction



- Localized re-surfacing

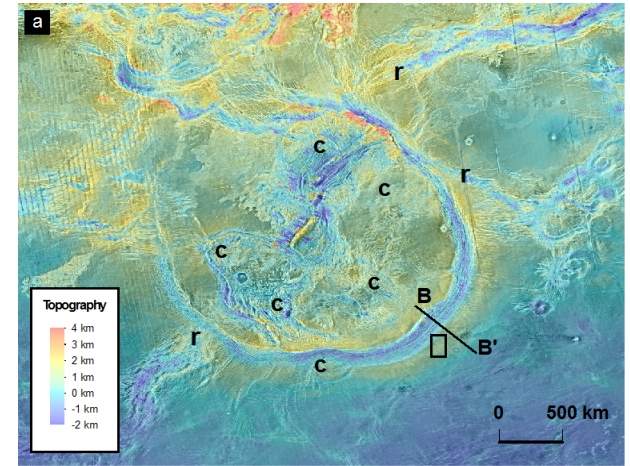
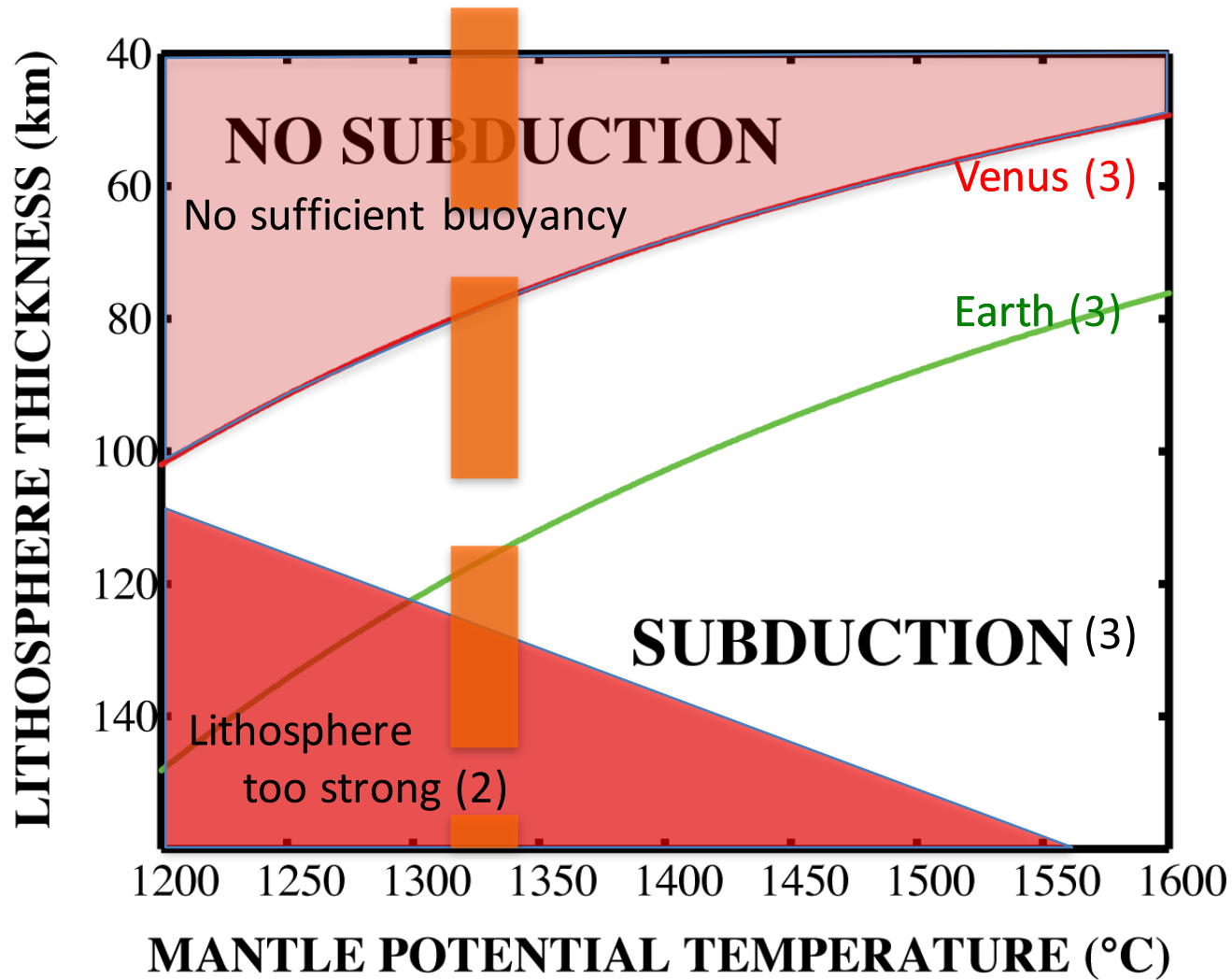
roll-back => tearing at slab edges

=> subduction stops when plume spreading stops

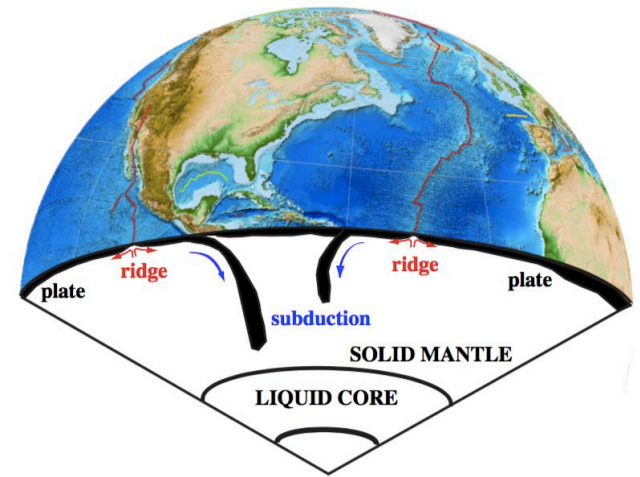


## 6- « Kick Hard to Break » :

### Plume-induced subduction



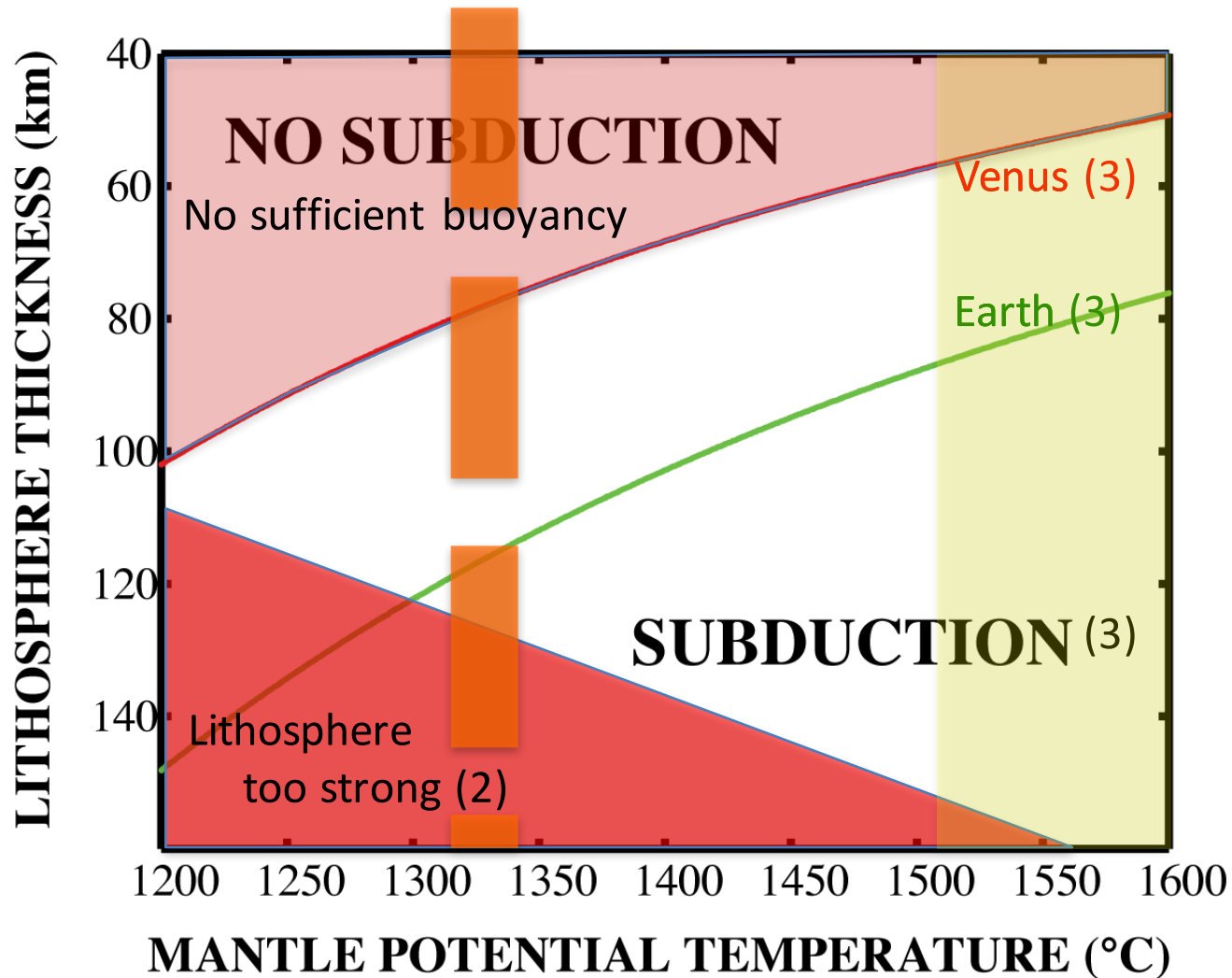
Artemis Coronae (Venus)



Carribean (Earth)

## 6- « Kick Hard to Break » :

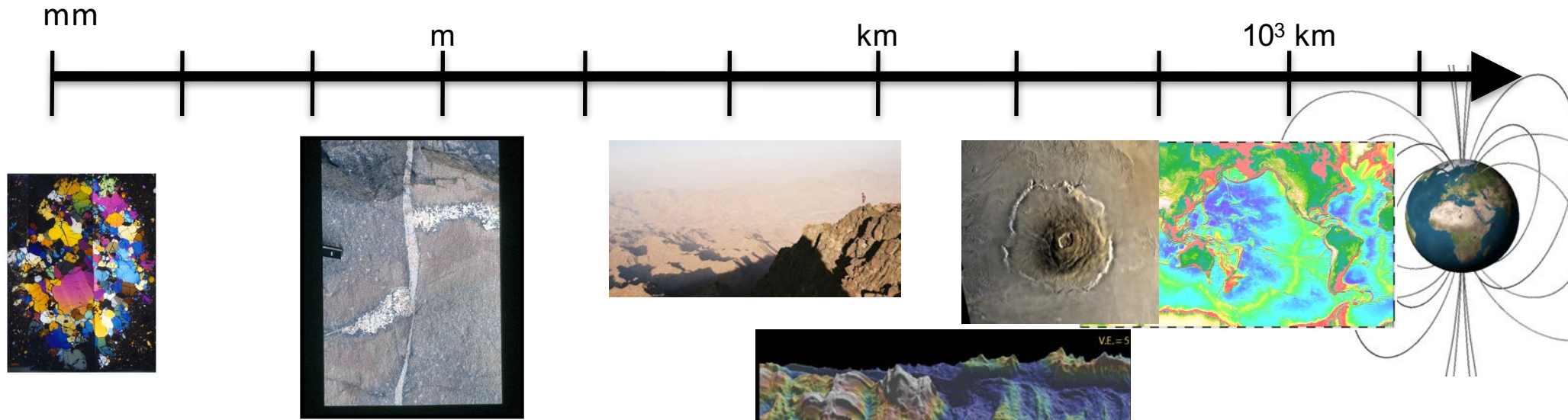
### Plume-induced subduction



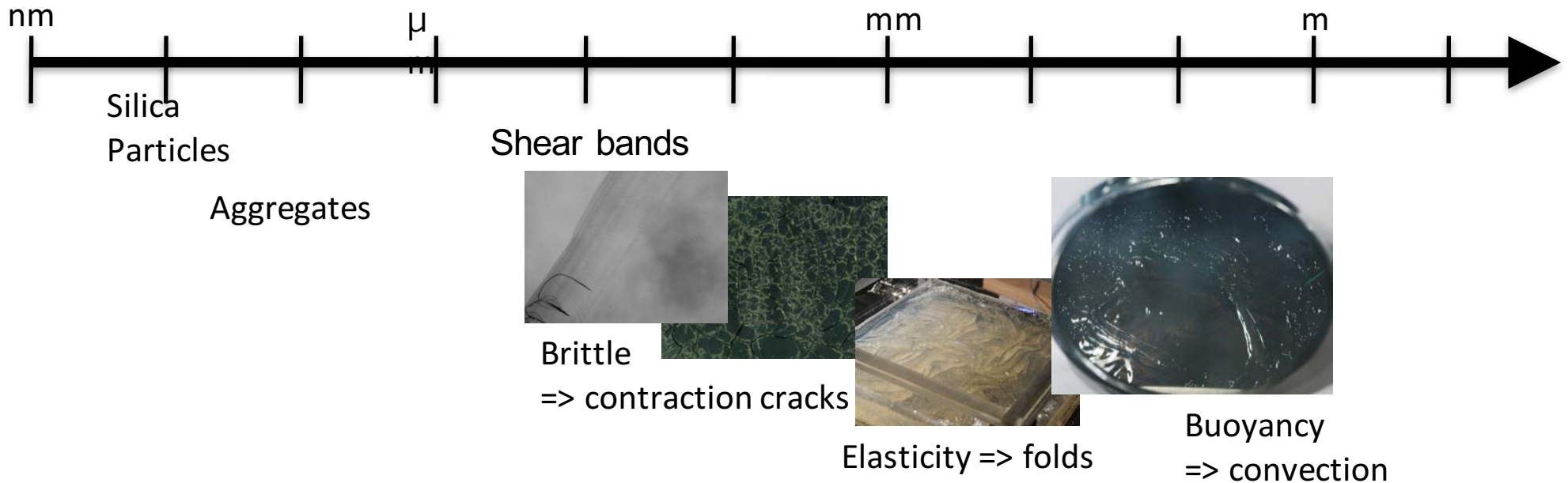
### Archean:

- Continent formation
- (1) Plume
- + (2) Subduction

# A visco-elasto-plastic « lithosphere » which has memory

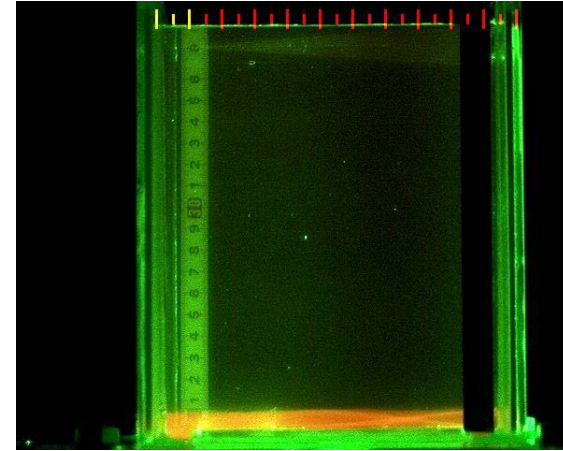


**Superposition of phenomena  
at different scales**

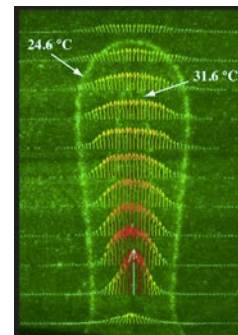


# CONCLUSIONS:

- 1- **NO dichotomy between plate tectonics and mantle convection**  
slabs = cold limbs of mantle convection
- 2- Mantle convection is **time-dependent**  
CMB= slab graveyard ; asthenosphere = plume graveyard
- 3- **Several types of hot spots/plumes**
- 4- **To make a plate** : strongly temperature-dependent viscosity  
**To break a plate** : two-phase mixture (solids, liquid) + memory
- 5- **Plate tectonics is only one regime out of many**



- NEXT:**
- find the « effective rheology » ?
  - map the different regimes
  - evolution of the Earth ?



**THANK YOU**