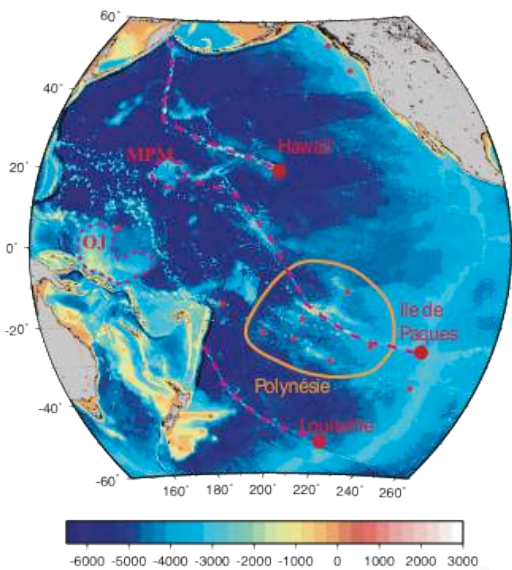


# Intertwined evolution of piles, plumes and slabs in the deep mantle

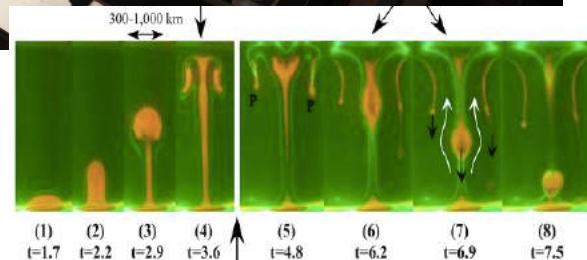
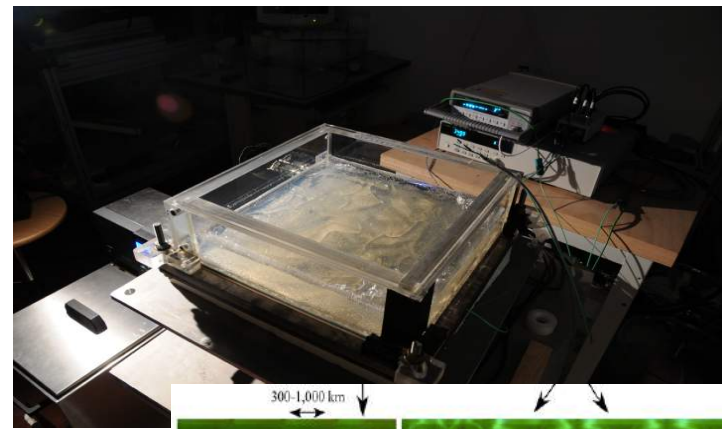
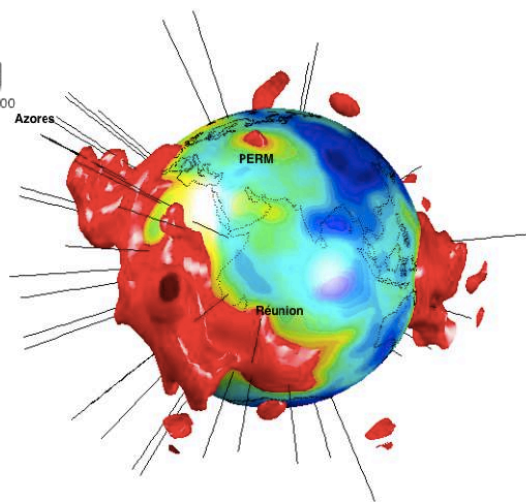
Anne Davaille, FAST, CNRS/Univ. Paris-Sud, Orsay, France

Ichrio Kumagai, Michael Le Bars, Sophie Androvandi, Judith Vatteville, Anna Massmeyer, Cecilia Cadio

K. Kurita, A. Limare, E. Stutzmann, J. Besse, V. Courtillot, A. Ismail-Zadeh, I. Panet, M. Diament



Observations  
+  
Fluid Mechanics

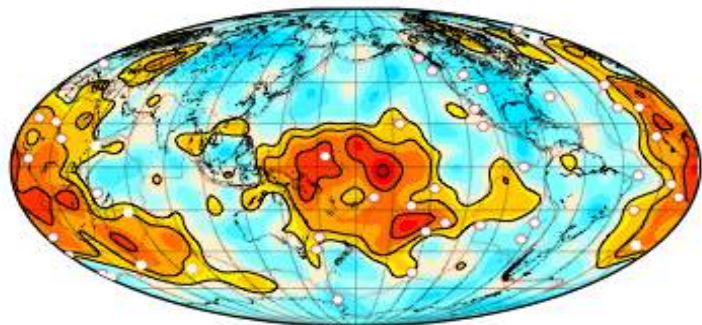


Morphology ?  
Time-dependence ?

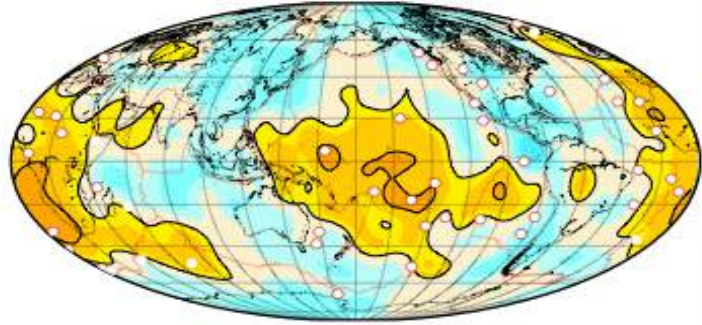
# 1- Observations: mantle « boxes »

S-wave velocity models at 2700 km depth

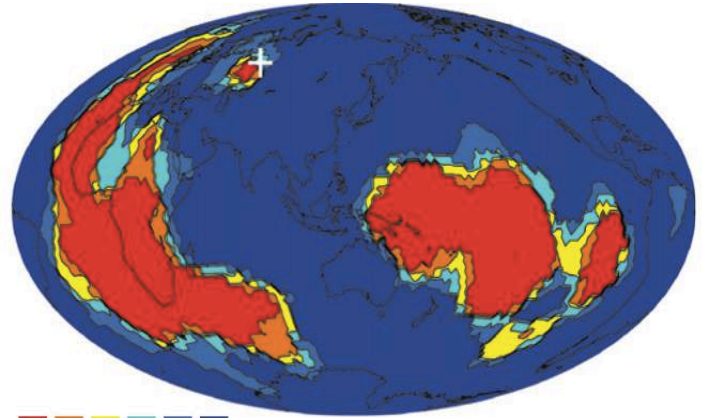
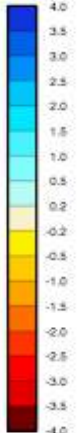
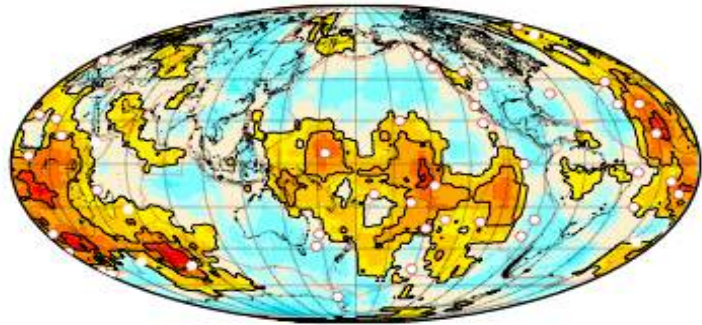
SAW24B16



S20RTS

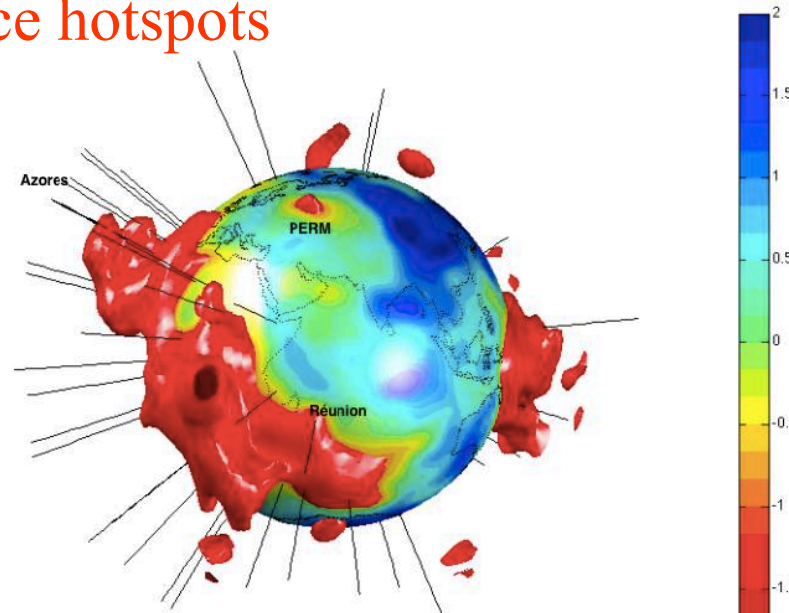


TXBW



( Lekic et al, 2012)

- 2 LLSVP + 1 + ...
- chemical heterogeneity
- old material (e.g. Jackson et al, 2010)
- surface hotspots

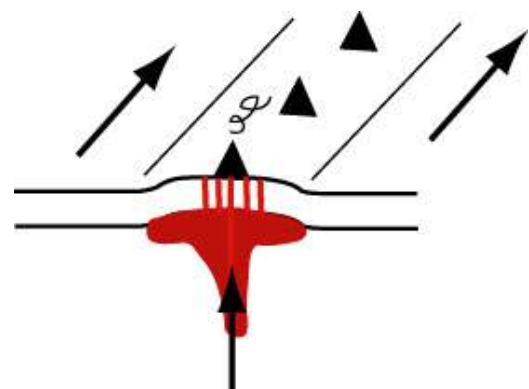


SMEAN -0.7% contour

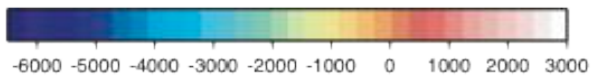
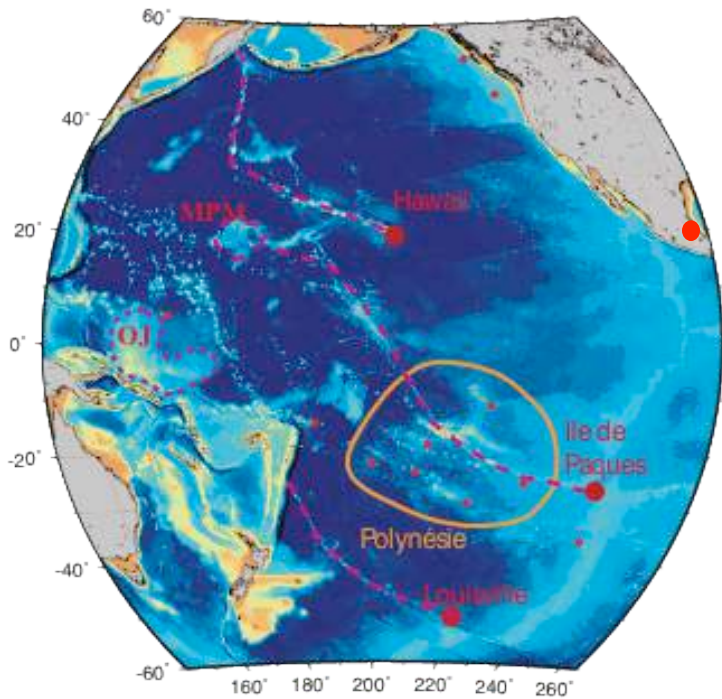
# 1- Observations: the hot spots « Zoo »

## -Surface Hot spot

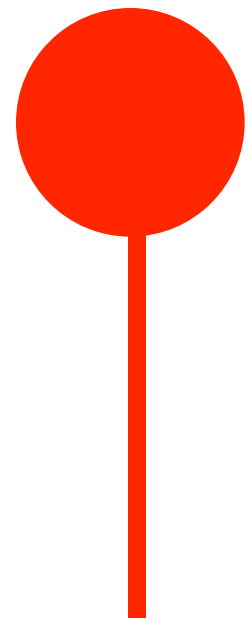
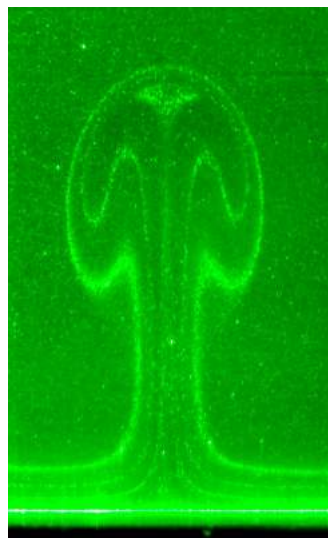
(Wilson 1963; Morgan 1971, 1972)



- Volcanism fixed / moving plate
- can start on by an oceanic plateau



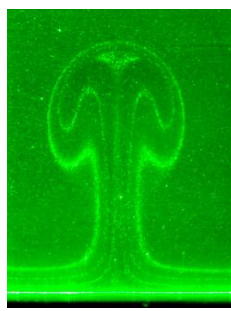
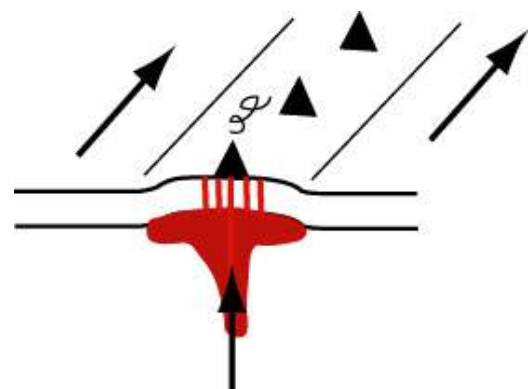
(Clouard & Bonneville 2000)



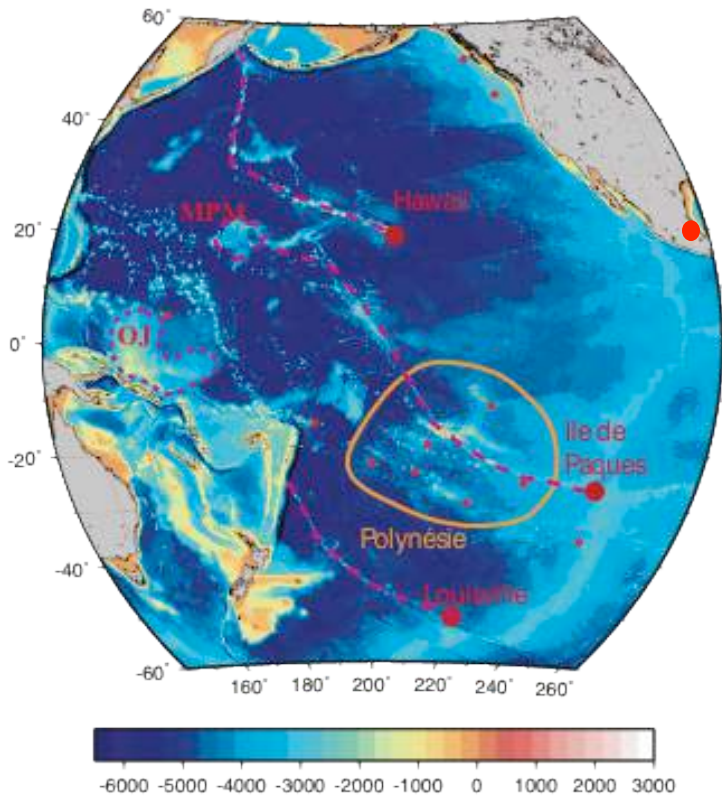
# 1- Observations: the hot spots « Zoo »

## -Surface Hot spot

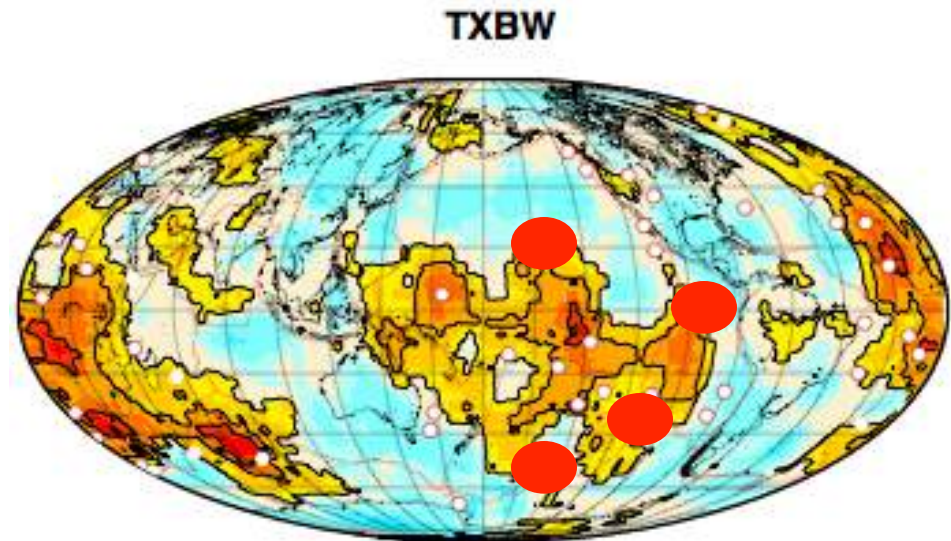
(Wilson 1963; Morgan 1971, 1972)



- Volcanism fixed / moving plate
- can start by an oceanic plateau
- Long tracks on the edges of slow anomaly



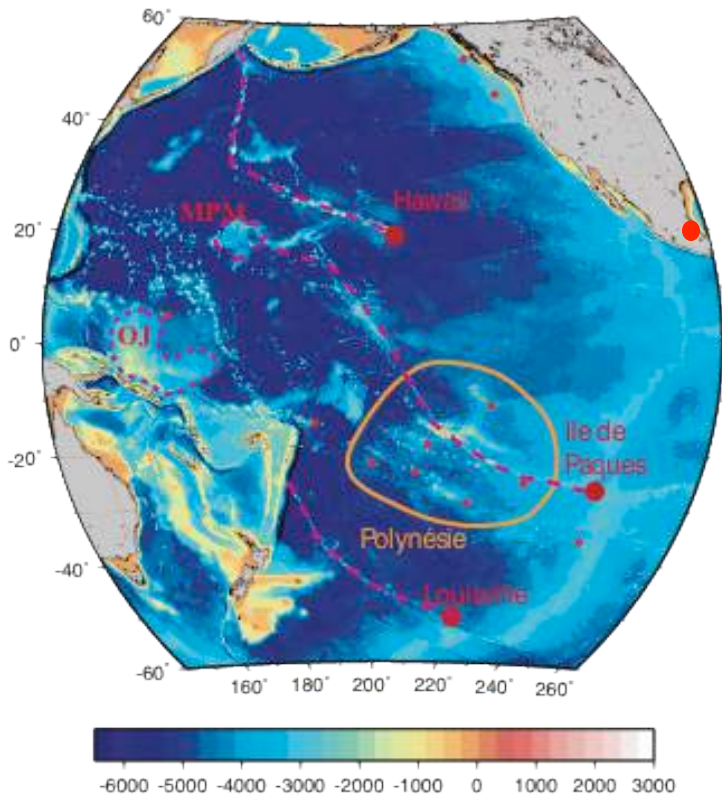
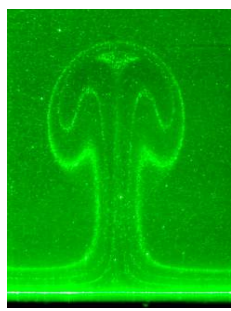
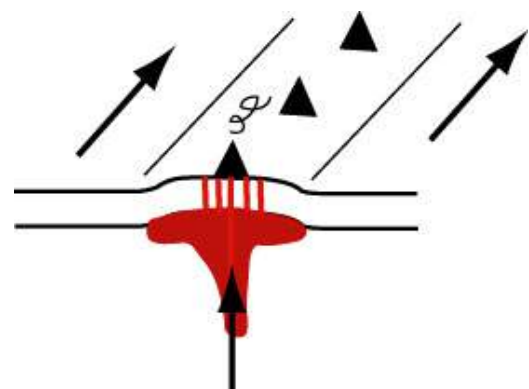
(Clouard & Bonneville 2000)



# 1- Observations: the hot spots « Zoo »

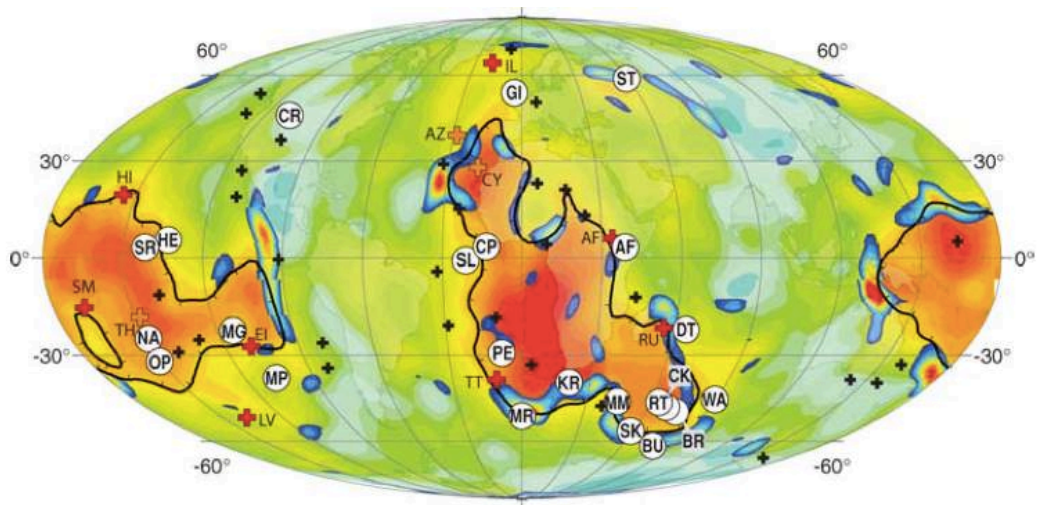
## -Surface Hot spot

(Wilson 1963; Morgan 1971, 1972)



(Clouard & Bonneville 2000)

- Volcanism fixed / moving plate
- can start by an oceanic plateau
- Long tracks/LIPs on the edges of LLSVPs over the last 200 Ma



(Torsvik et al, 2004)

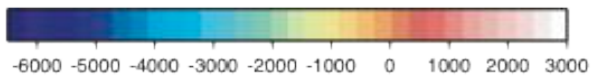
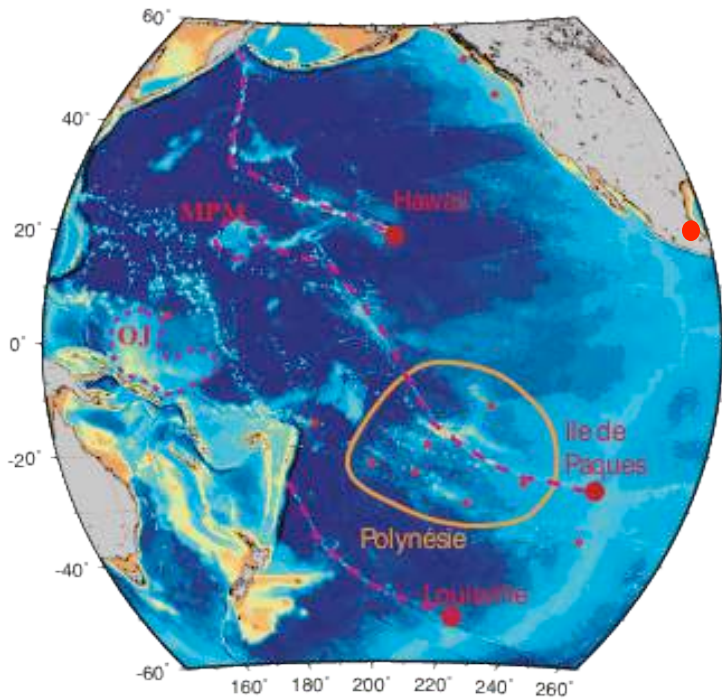
# 1- Observations: the hot spots « Zoo »

## -Surface Hot spot

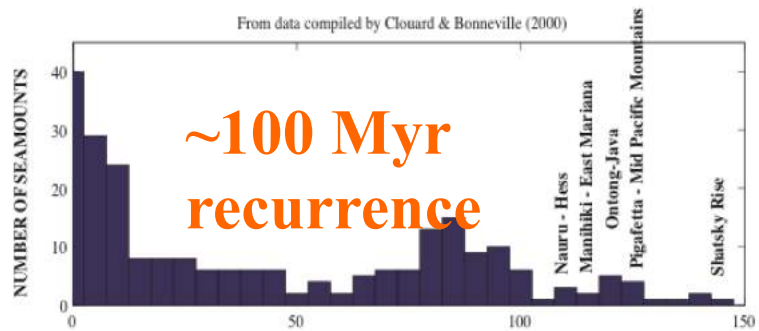
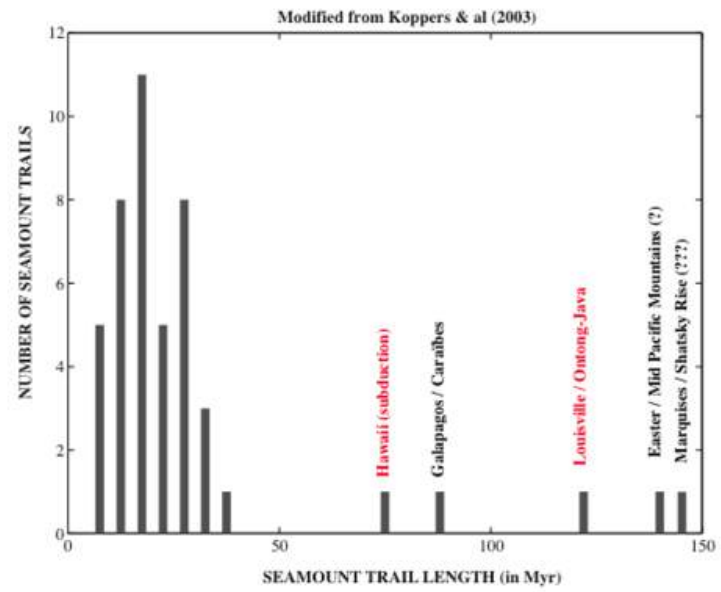
(Wilson 1963; Morgan 1971, 1972)

+ another type

= shorter tracks in « cluster »

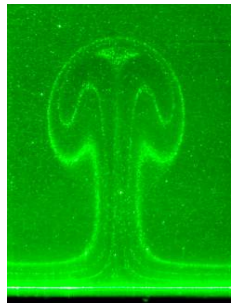
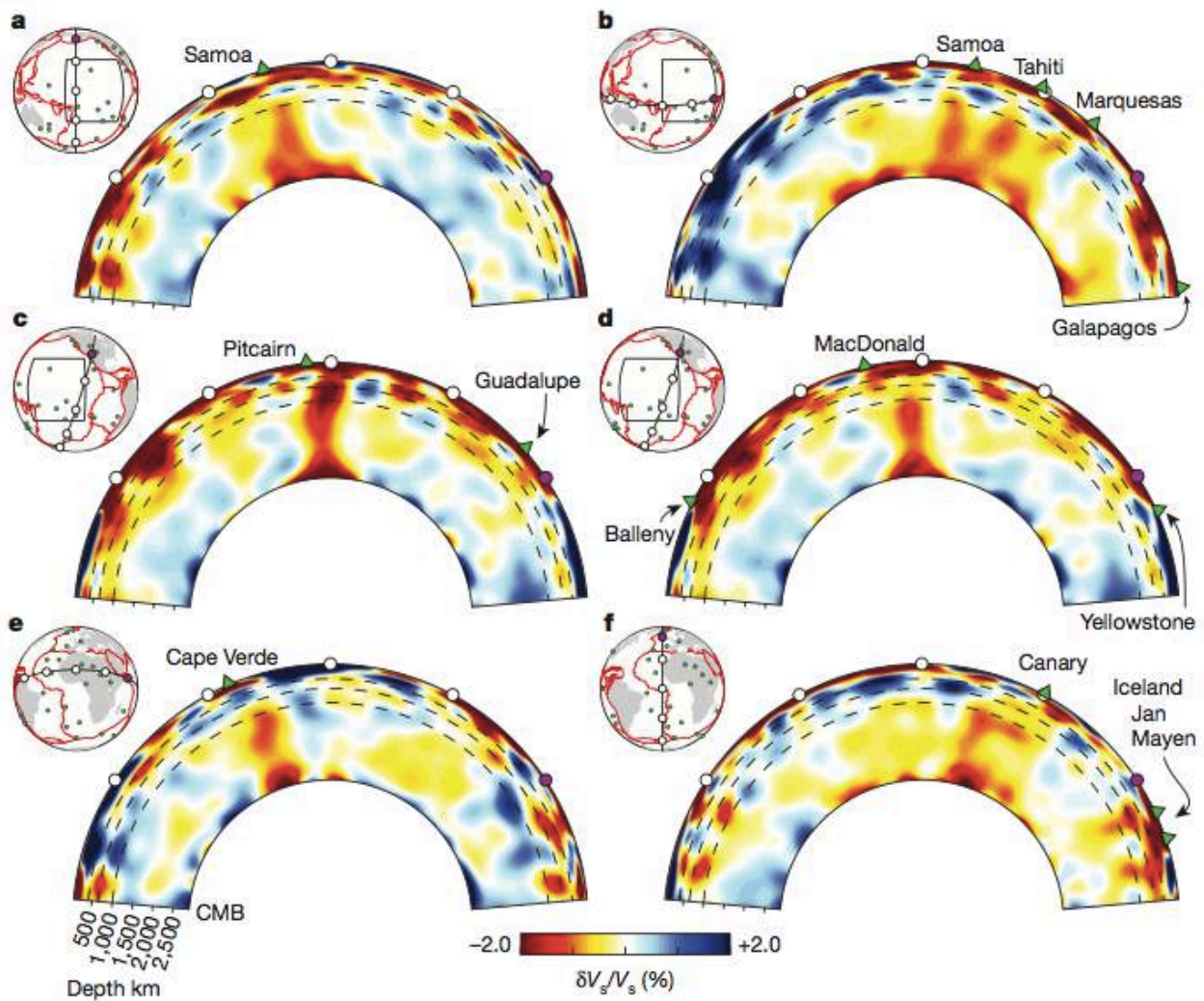


(Clouard & Bonneville 2000)



# 1- Observations: the hot spots « Zoo »

**mantle hot upwelling : A- something is happening around 1000 km**



**B- FAT plumes**

(French & Romanowicz, 2015; Montelli et al, 2006)

# QUESTIONS :

**A- Origin of mantle « boxes » ?**

**B- Upwellings**

- several types ?

- « fat » ?

- what happens around 1000 kms ?

**C- Do the mantle boxes / LLSVPs change shapes through time ?**



# QUESTIONS :

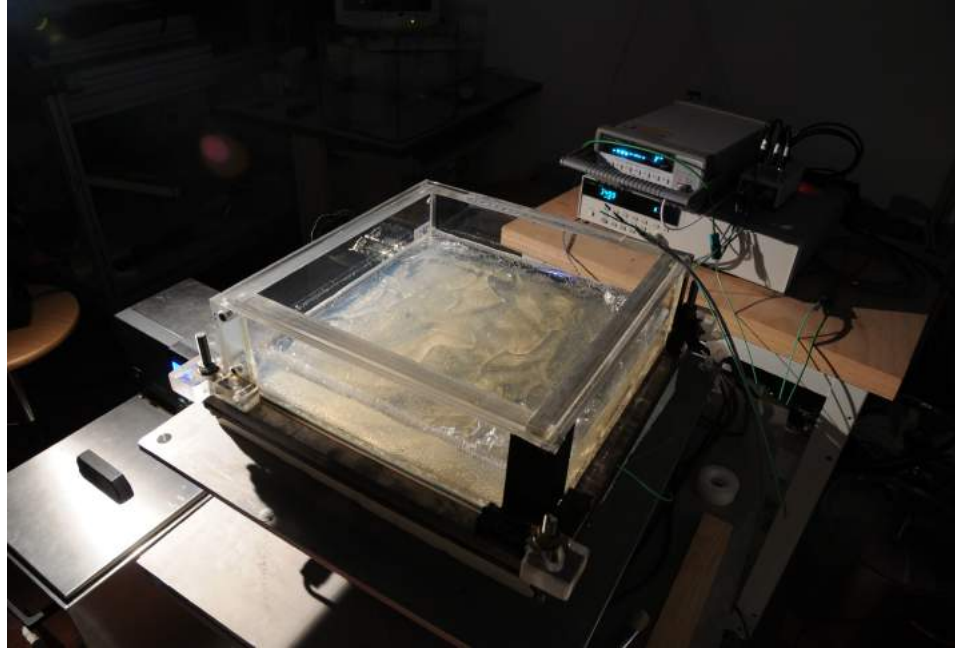
**A- Origin of mantle « boxes » ?**

**B- Upwellings**

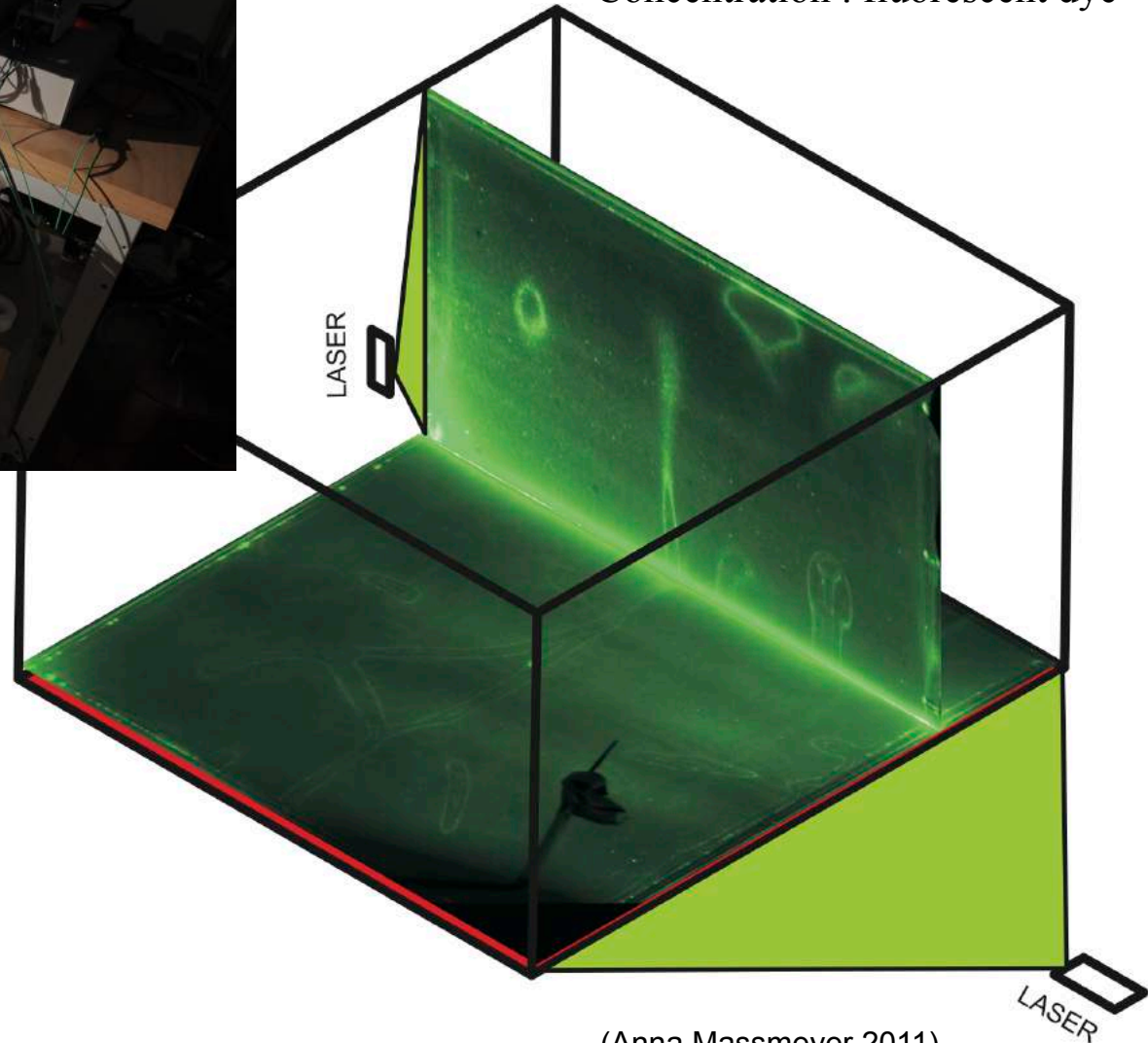
- several types ?
- « fat » ?
- what happens around 1000 kms ?

**C- Do the mantle boxes / LLSVPs change shapes through time ?**

## 2- Experiments : set up



- Plexiglass Tank 30x30 cm
- Aspect ratio: 2 to 5
- Velocity field : PIV
- Temperature isotherms : TLCs
- Concentration : fluorescent dye

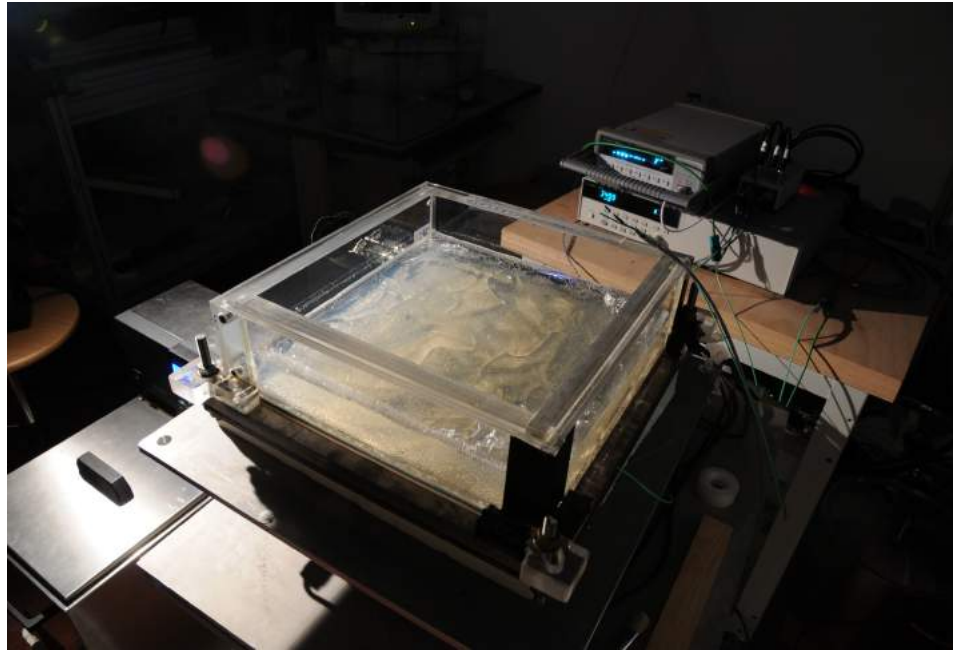


### FLUIDS :

- sugar : newtonian  $\eta(T)$
- polymers : shear thinning, viscoplastic
- colloids :  
newtonian (bulk)  $\rightarrow$  brittle (top)

(Anna Massmeyer 2011)

## 2- Experiments : set up

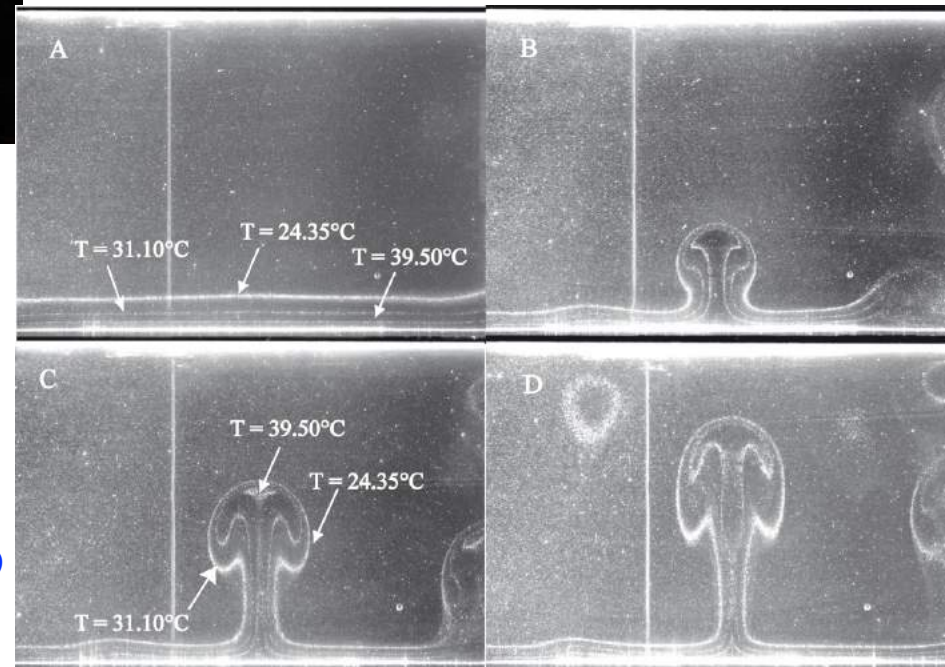


- Plexiglass Tank 30x30 cm
- Aspect ratio: 2 to 5
- Velocity field : PIV
- Temperature isotherms : TLCs
- Concentration : fluorescent dye

### FLUIDS :

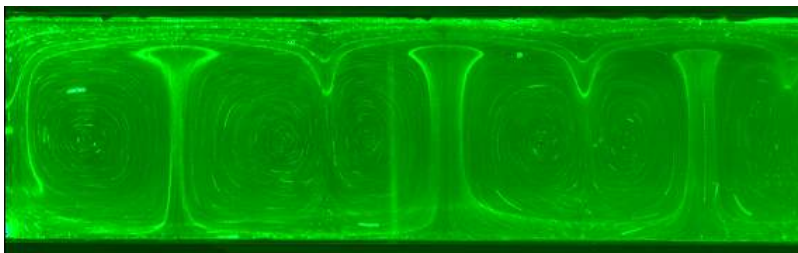
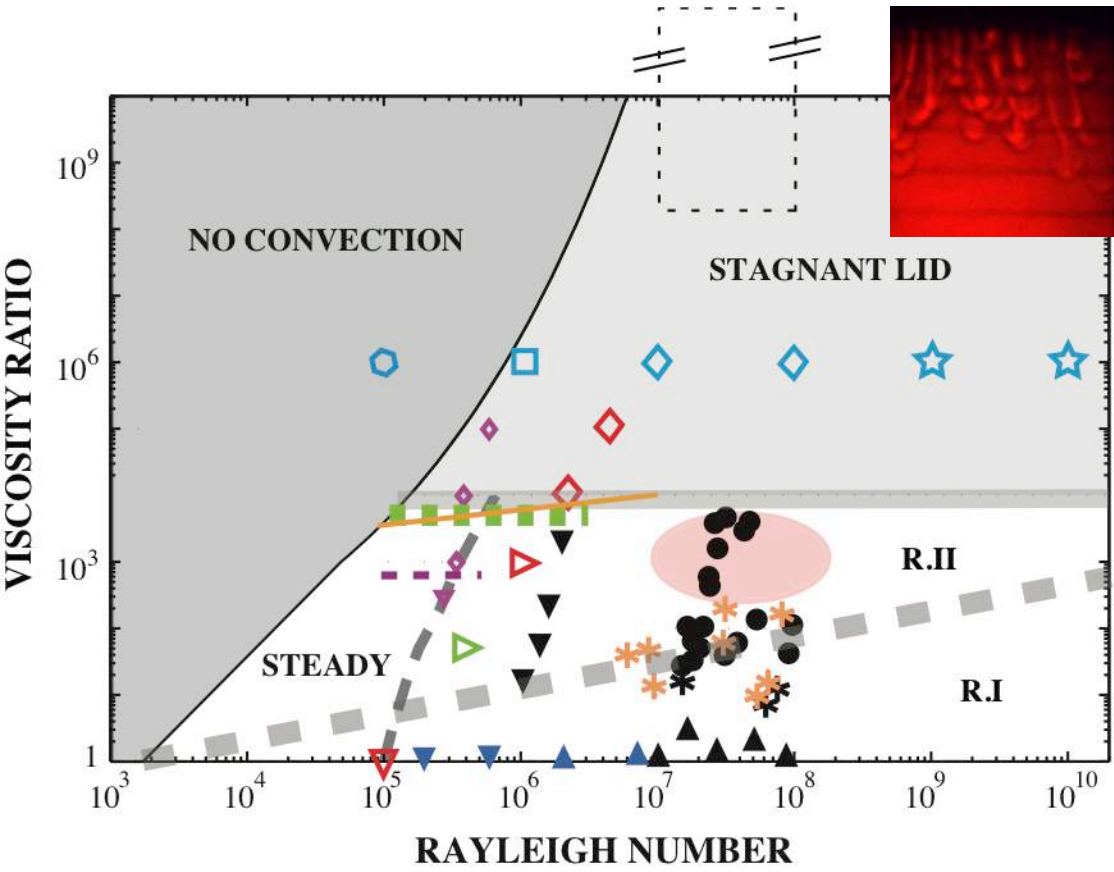
- sugar : newtonian  $\eta(T)$
- colloids :  
newtonian (bulk)  $\rightarrow$  brittle (top)

$$Ra = \alpha g \Delta T H^3 / \kappa \nu \sim 10^6 - 10^9$$



(Sophie Androvandi 2009)

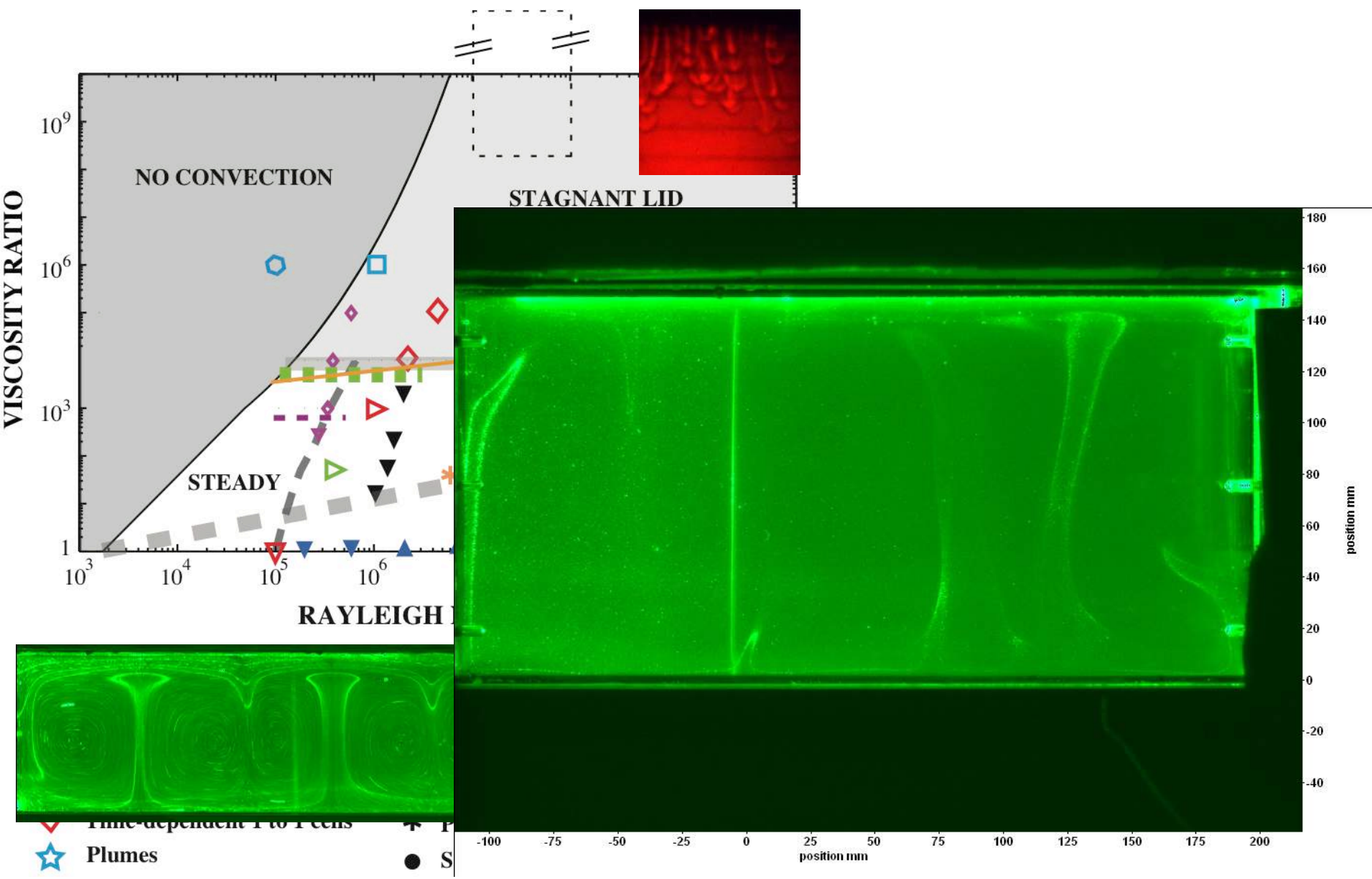
### 3. Convection in a fluid with Temperature-dependent viscosity



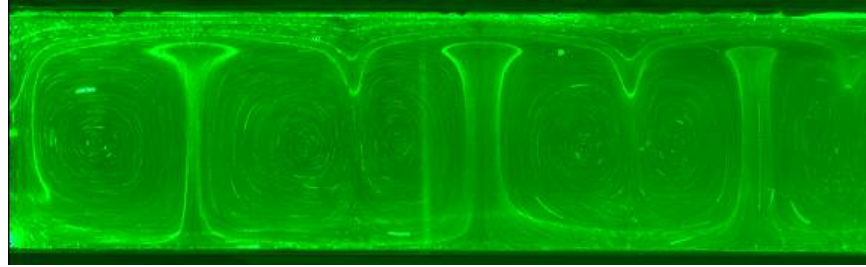
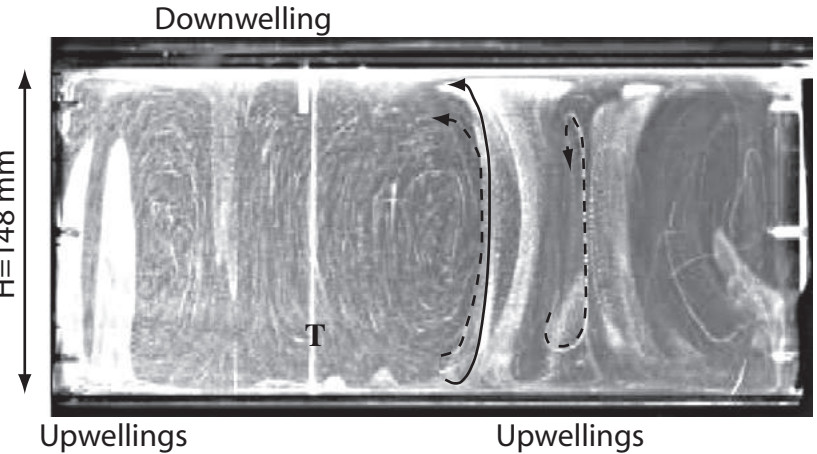
Temperature-dependent 1 to 1 cell  
 cold  
 1 to 3 cell

- ▽ Time-dependent 1 to 1 cells
- ★ Plumes
- ▲ Plumes, not cold
- Several hot plumes for 1 cold cell

### 3. Convection in a fluid with Temperature-dependent viscosity



### 3. Convection in a fluid with Temperature-dependent viscosity



#### Spacing:

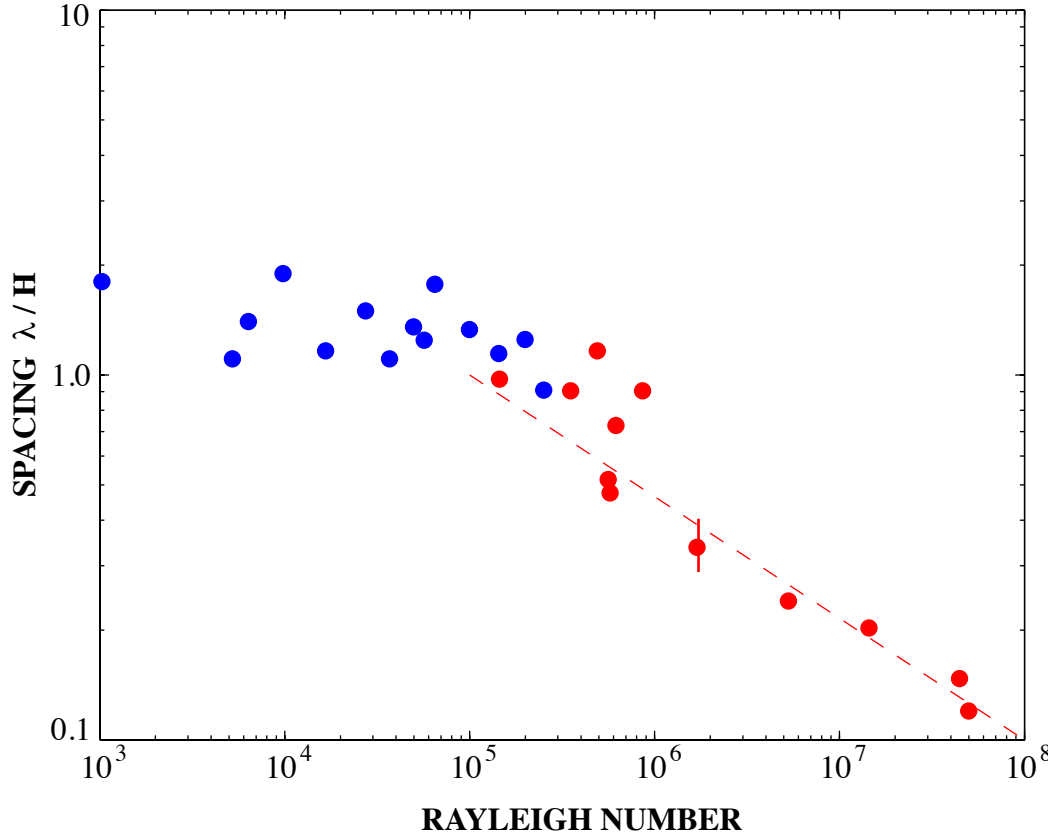
**-cold** : cell with  $l \sim 1-2 H$  (R-R)

Giannandrea & Christensen (F-F)  
 $\Rightarrow l \sim 3.5 H$

$\Rightarrow$  mantle  $\Rightarrow l = 10\ 000\ \text{km} \dots$

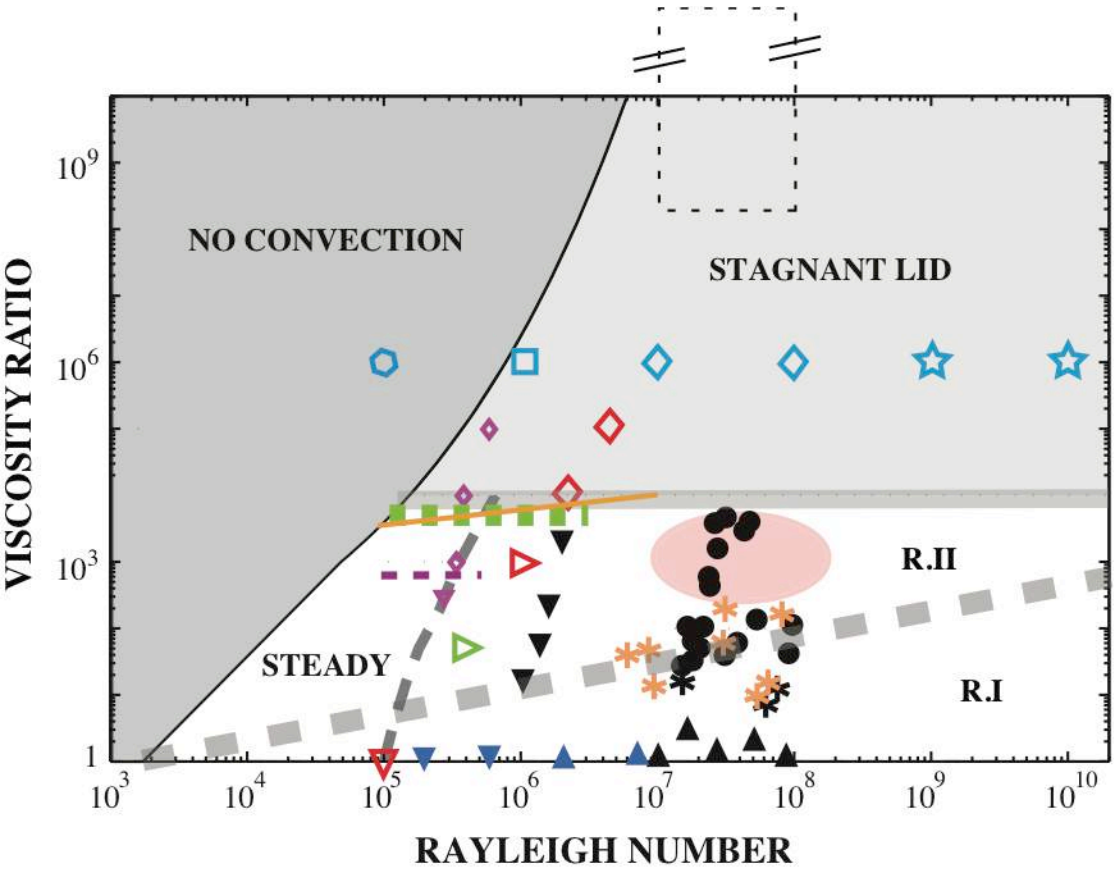
**-hot** : spacing  $\sim 2-4 H \cdot (Ra_c / Ra_{bot})^{1/3}$

$\Rightarrow \sim 200-1000\ \text{kms at CMB}$

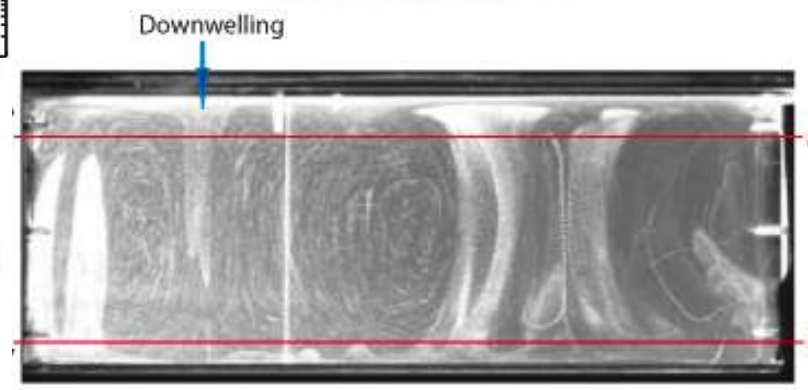
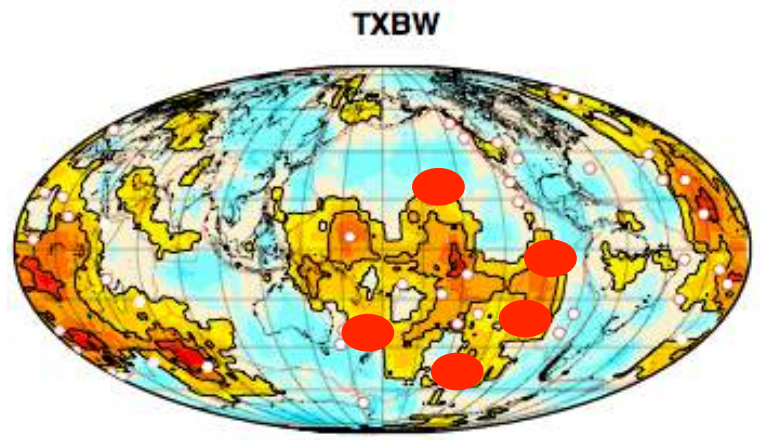


(Androvandi & al, 2009, 2011)

### 3. Convection in a fluid with Temperature-dependent viscosity



- No convection
- Steady 1 to 1 cells
- ◇ Time-dependent 1 to 1 cells
- ☆ Plumes
- ▼ Spokes = time-dependent 1 to 1 cell
- ▲ Plumes, hot = cold
- ▽ Spokes in L-L: 1 to 3 cell
- \* Plumes, hot ≠ cold
- Several hot plumes for 1 cold cell



High Ra: several scales

# QUESTIONS :

**A- Origin of mantle « boxes » ? Variable viscosity / Subduction**

**B- Upwellings**

- several types ?
- « fat » ?
- what happens around 1000 kms ?

**C- Do the mantle boxes / LLSVPs change shapes through time ?**



# QUESTIONS :

**A- Origin of mantle « boxes » ? Variable viscosity / Subduction**

**B- Upwellings : convection with density heterogeneities ?**

- several types ?

- « fat » ?

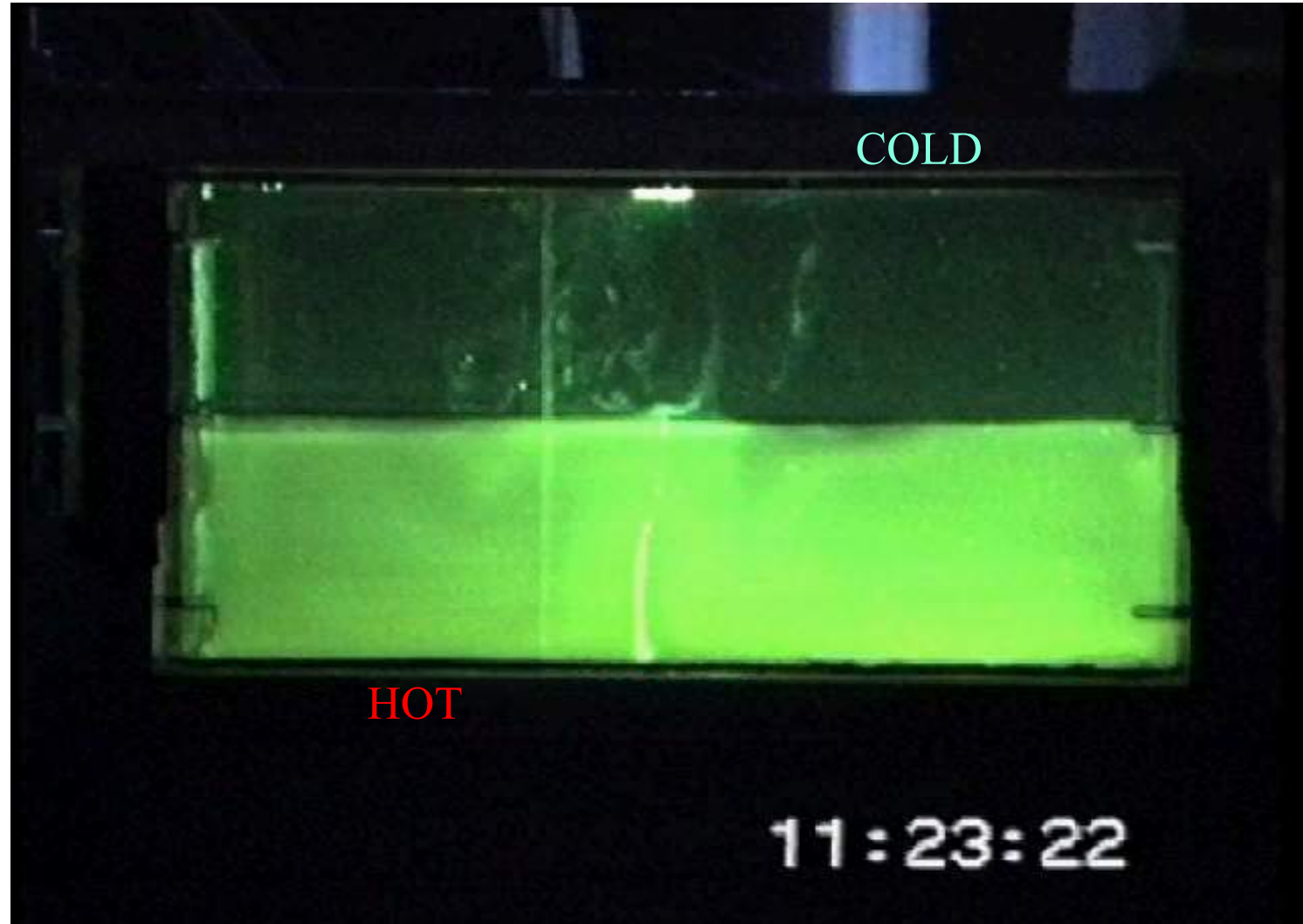
- what happens around 1000 kms ?

**C- Do the mantle boxes / LLSVPs change shapes through time ?**

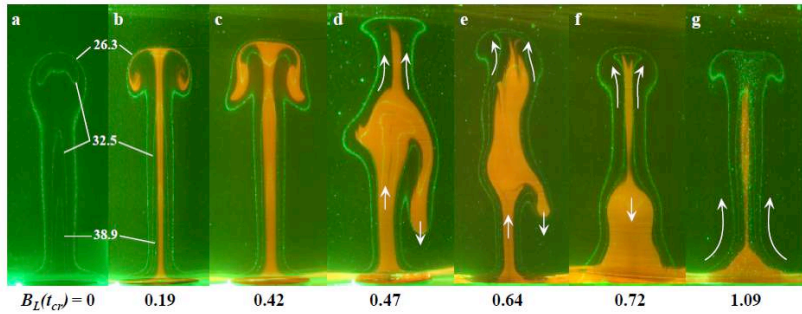
## 4. Interaction Convection / Denser Reservoir

Key parameter :  $B = \Delta\rho_c / \alpha \Delta T$  ( $\sim 0 - 10$ )

**WHOLE-LAYER :  $B < 1$  hot more viscous (polymers)**



# 4. Interaction Convection / Denser Reservoir

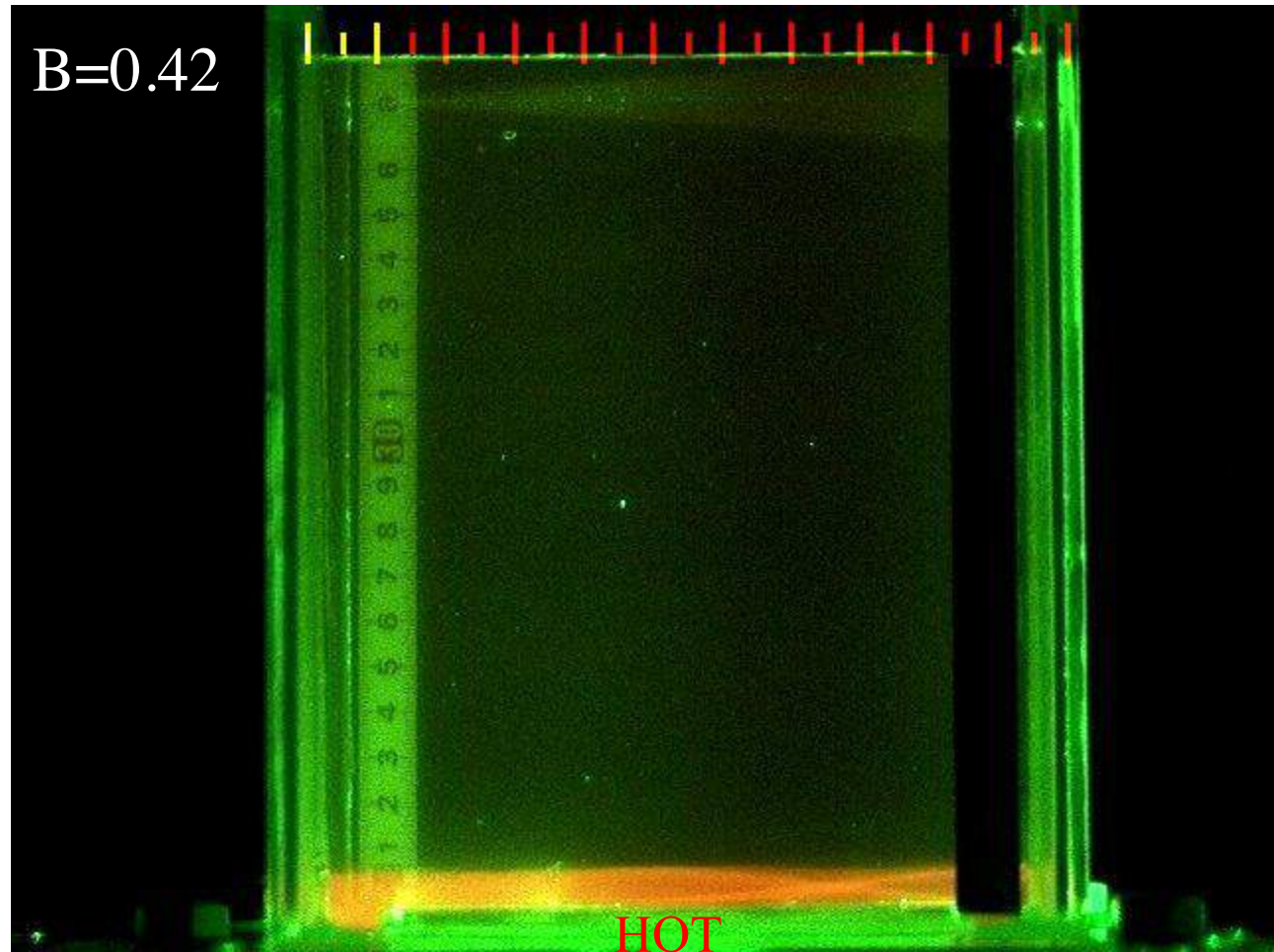


hot less viscous

(Ichiro Kumagai, Kei Kurita)

$$B = \Delta\rho_c / \alpha \Delta T$$

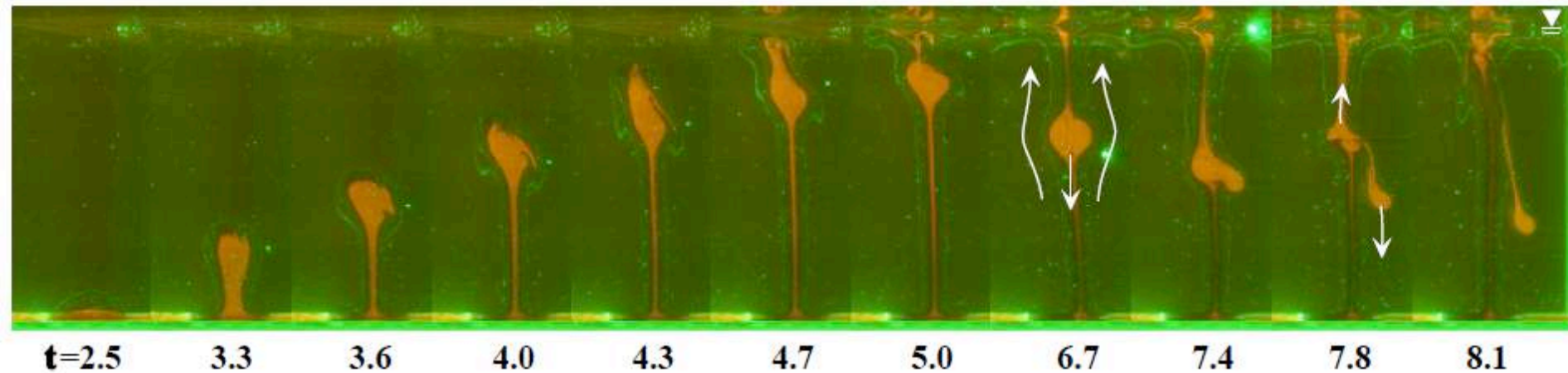
B=0.42



## 4. Interaction Convection / Denser Reservoir

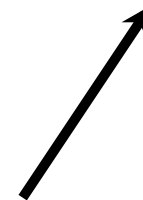
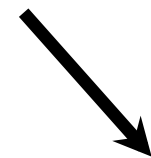
$B_L=0.59$

# Time-dependence



SUCCESSFULL...

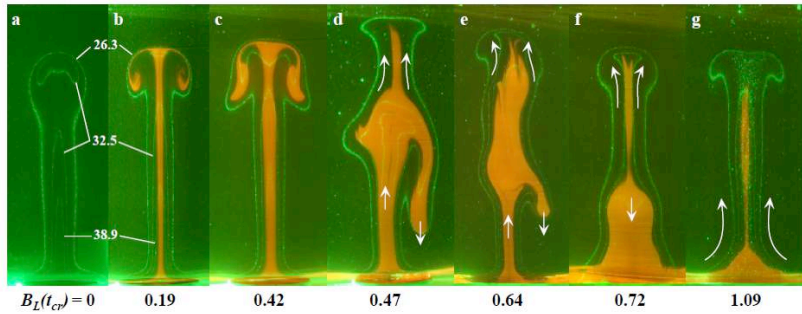
FAILING...



While going up, the thermochemical plume  
looses its heat by diffusion

(Kumagai & al, 2008)

# 4. Interaction Convection / Denser Reservoir

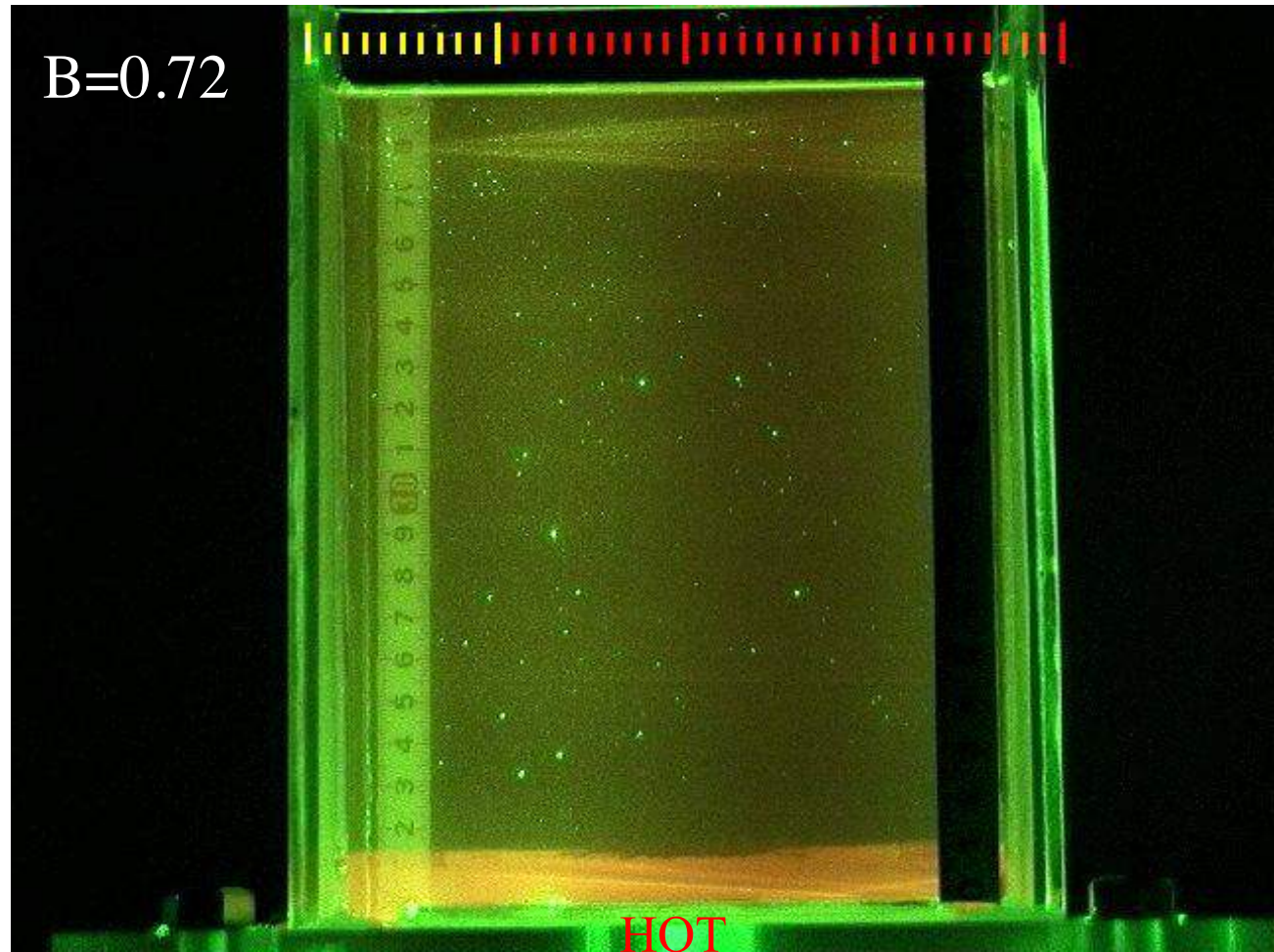


hot less viscous

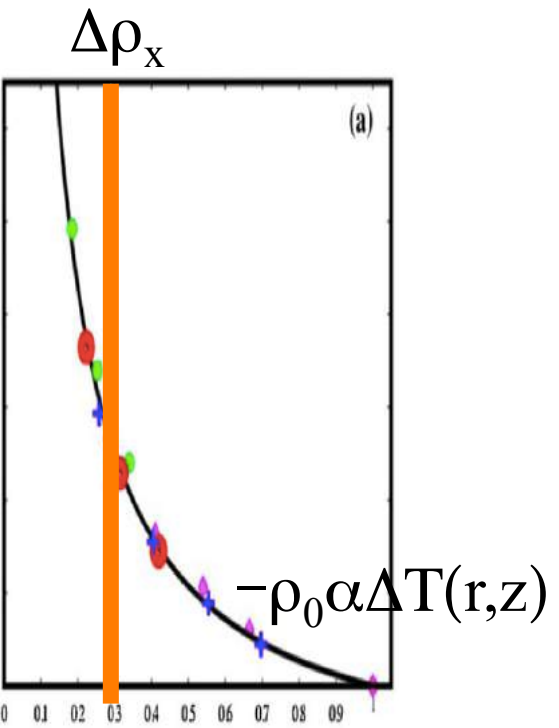
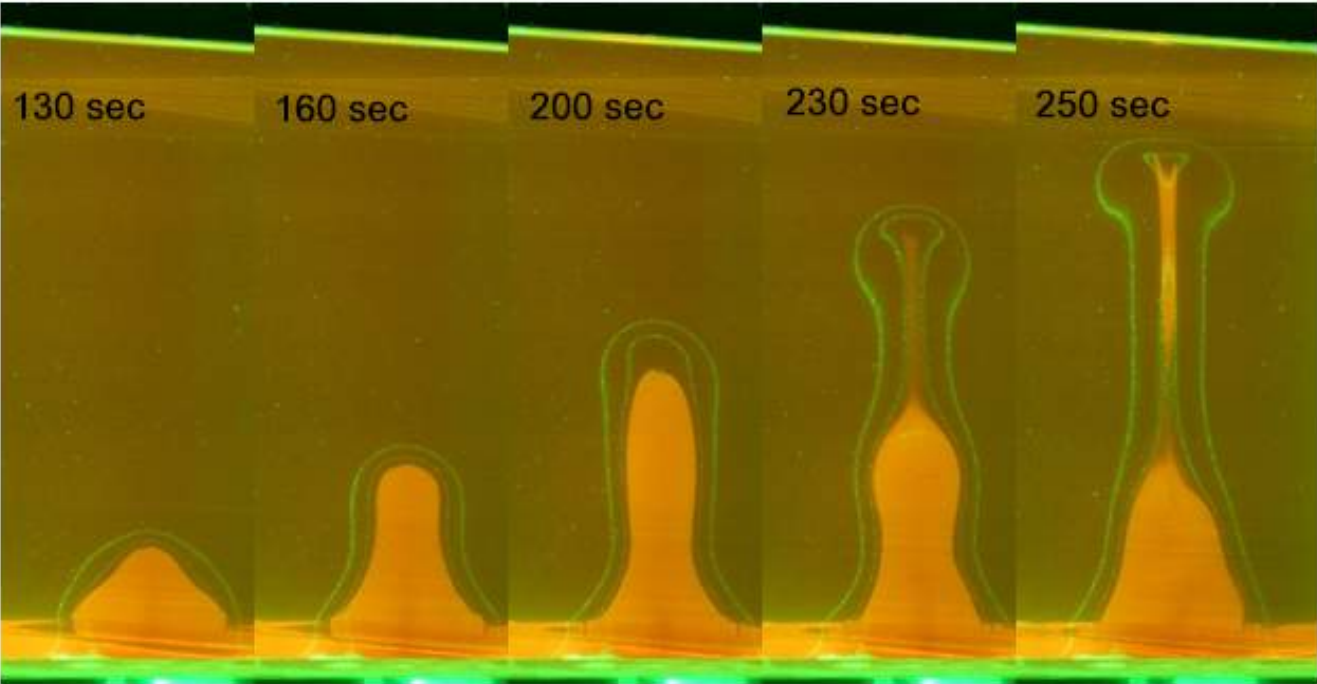
(Ichiro Kumagai, Kei Kurita)

$$B = \Delta\rho_c / \alpha \Delta T$$

$B=0.72$



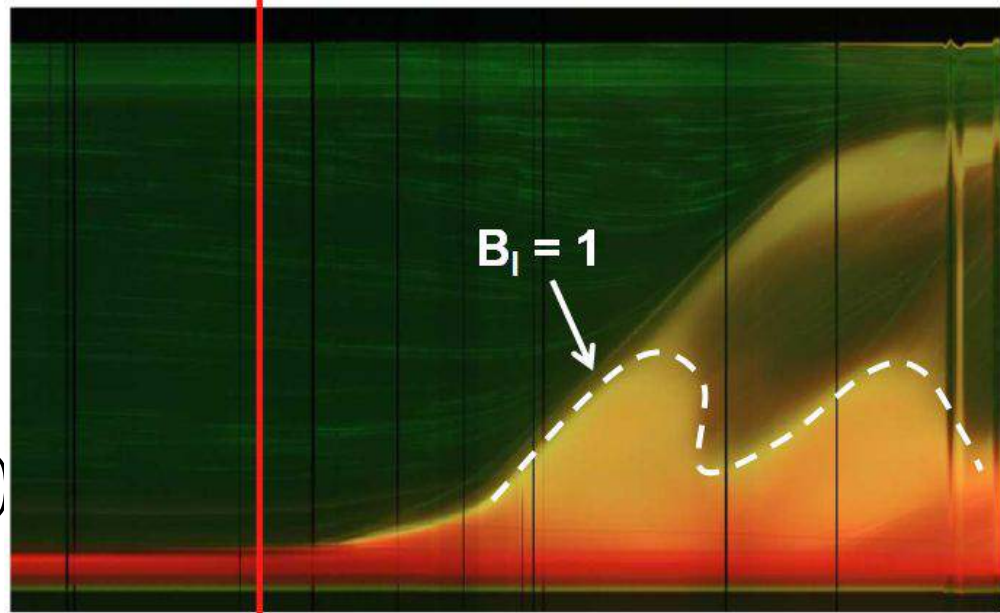
# 4. Interaction Convection / Denser Reservoir



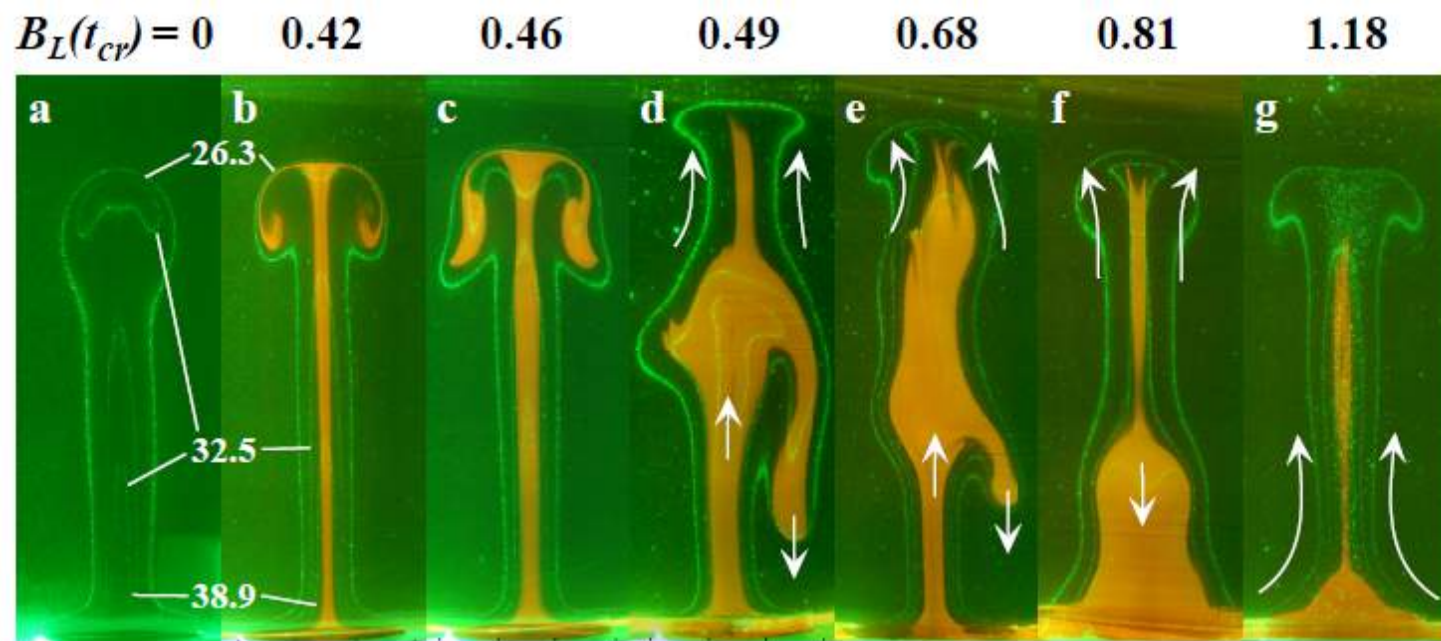
**B=0.72**

- Behaviour depends on time
- Dome + secondary plume

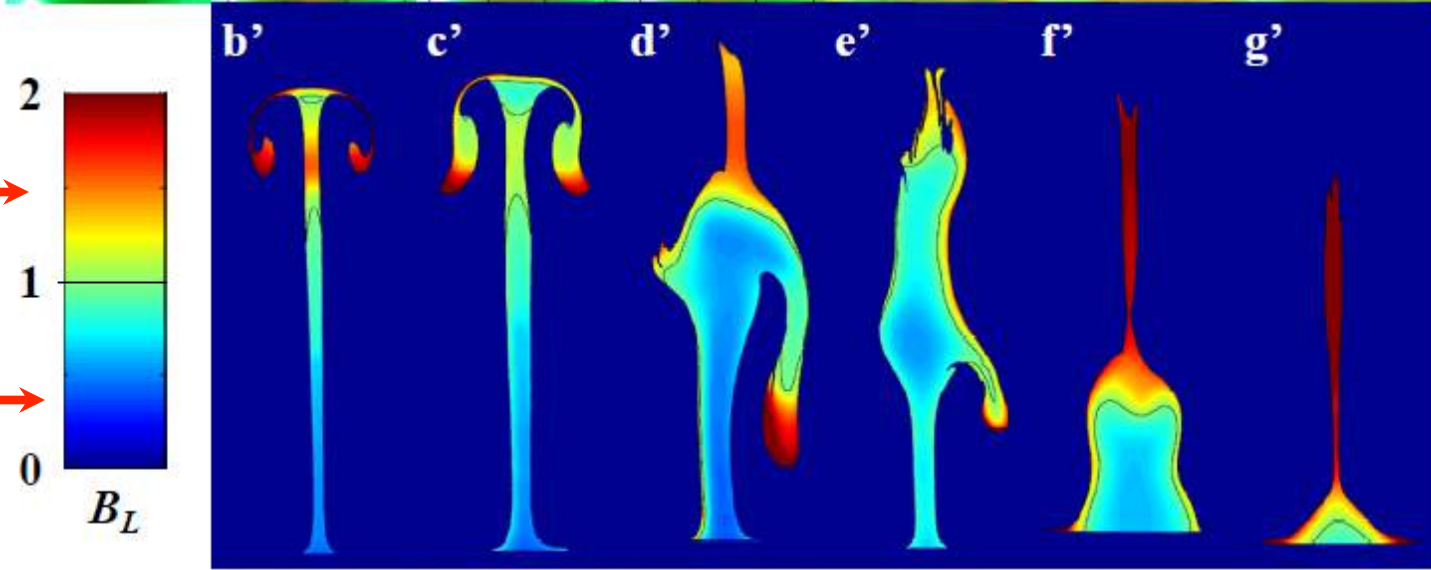
$$B_l = \Delta\rho_x / \rho_0 \alpha \Delta T(r,z)$$



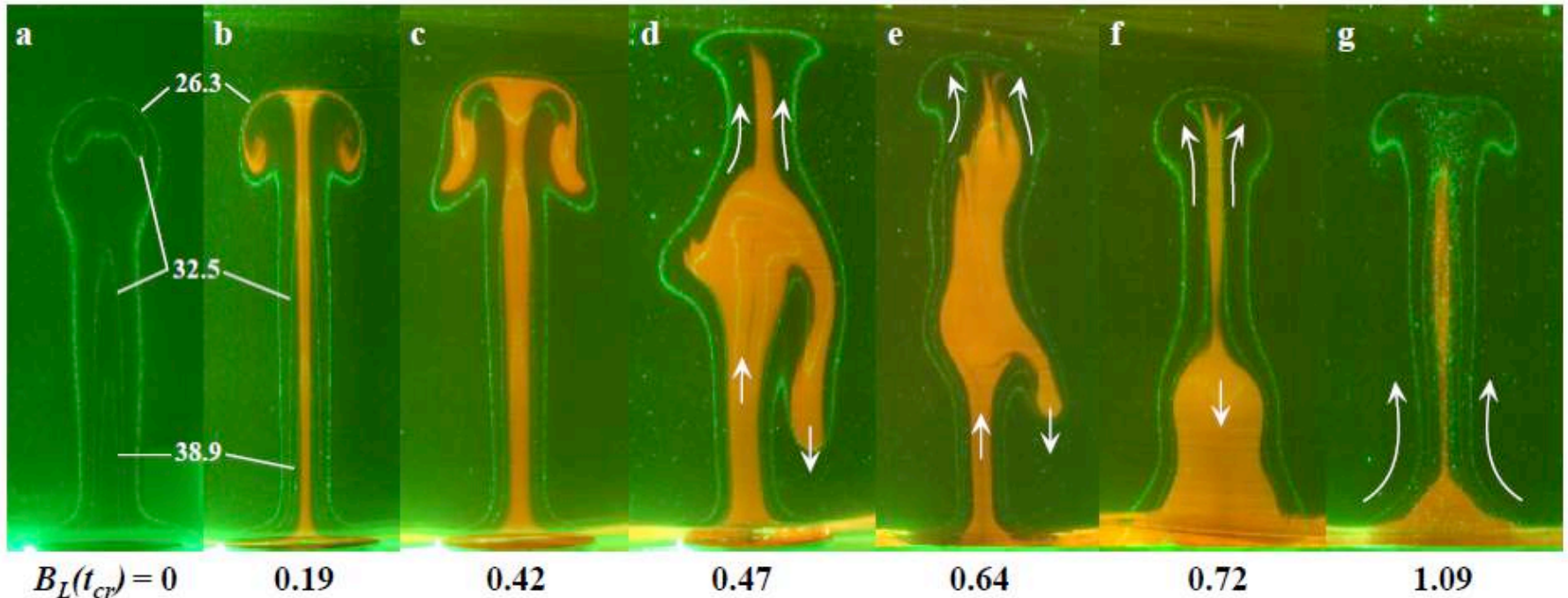
# 4. Interaction



KEY:  
 $B_1 > \text{or} < 1$

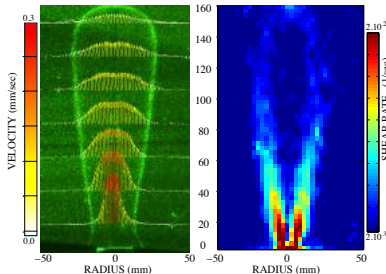


# 4. Interaction Convection / Denser Reservoir



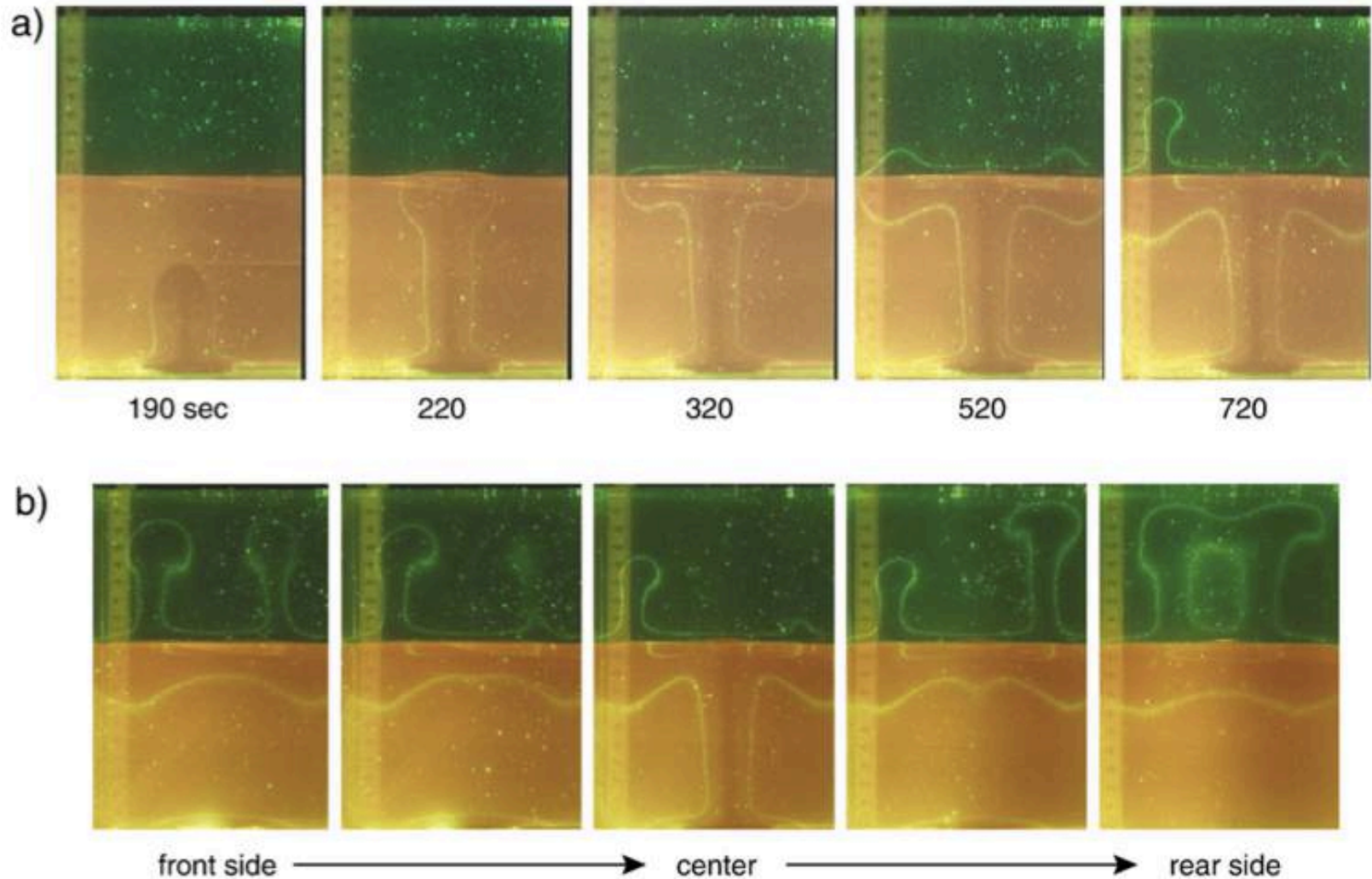
- . FAT upwellings -> circulation + viscosity
- . Complicated, time-dependent morphology
- . B increases => level of neutral buoyancy decreases

$\Delta\rho_x \geq 1\%$   
for  $\Delta T(r,z) \sim 300^\circ$





## 4. Interaction Convection / Denser Reservoir



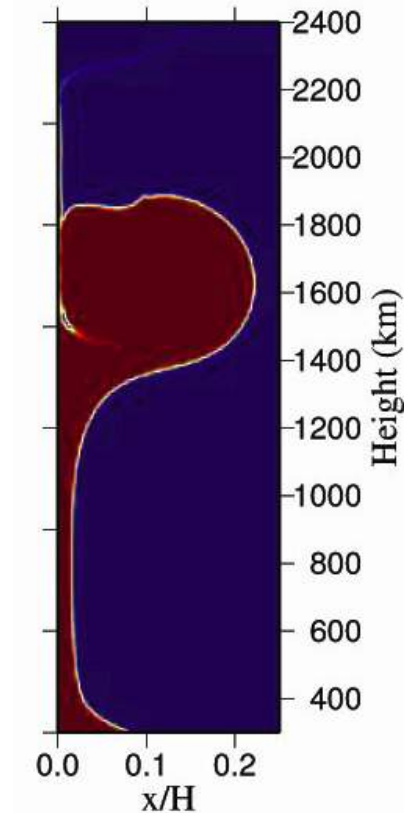
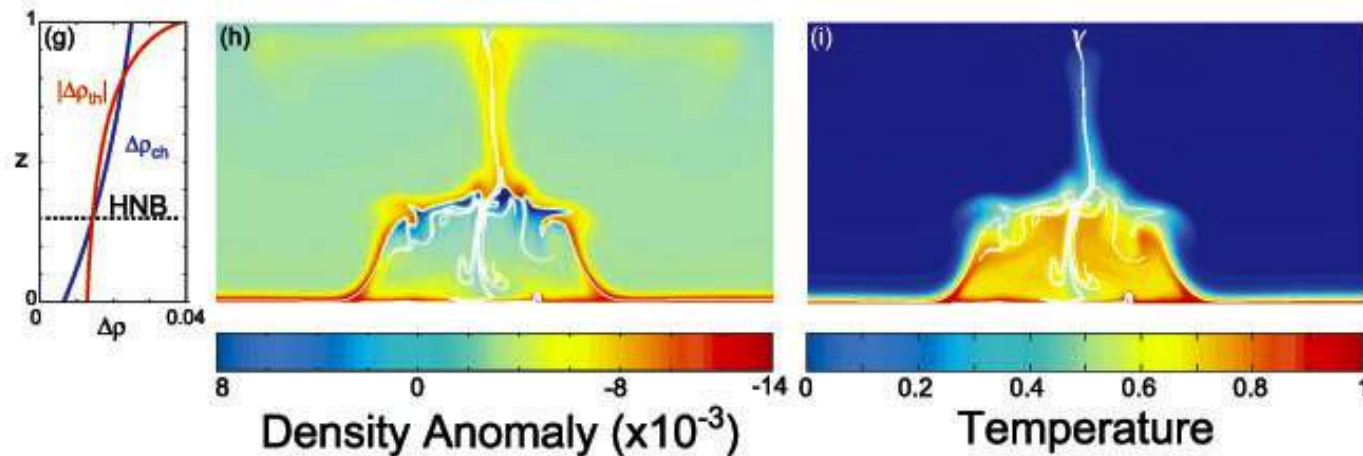
**Delay through a boundary**

**=> generation of a cluster of secondary plumes**

## 4. Interaction Convection / Denser Reservoir

- 'Real' Mantle => pressure/depth dependence of density, phase, thermal expansion, viscosity,...

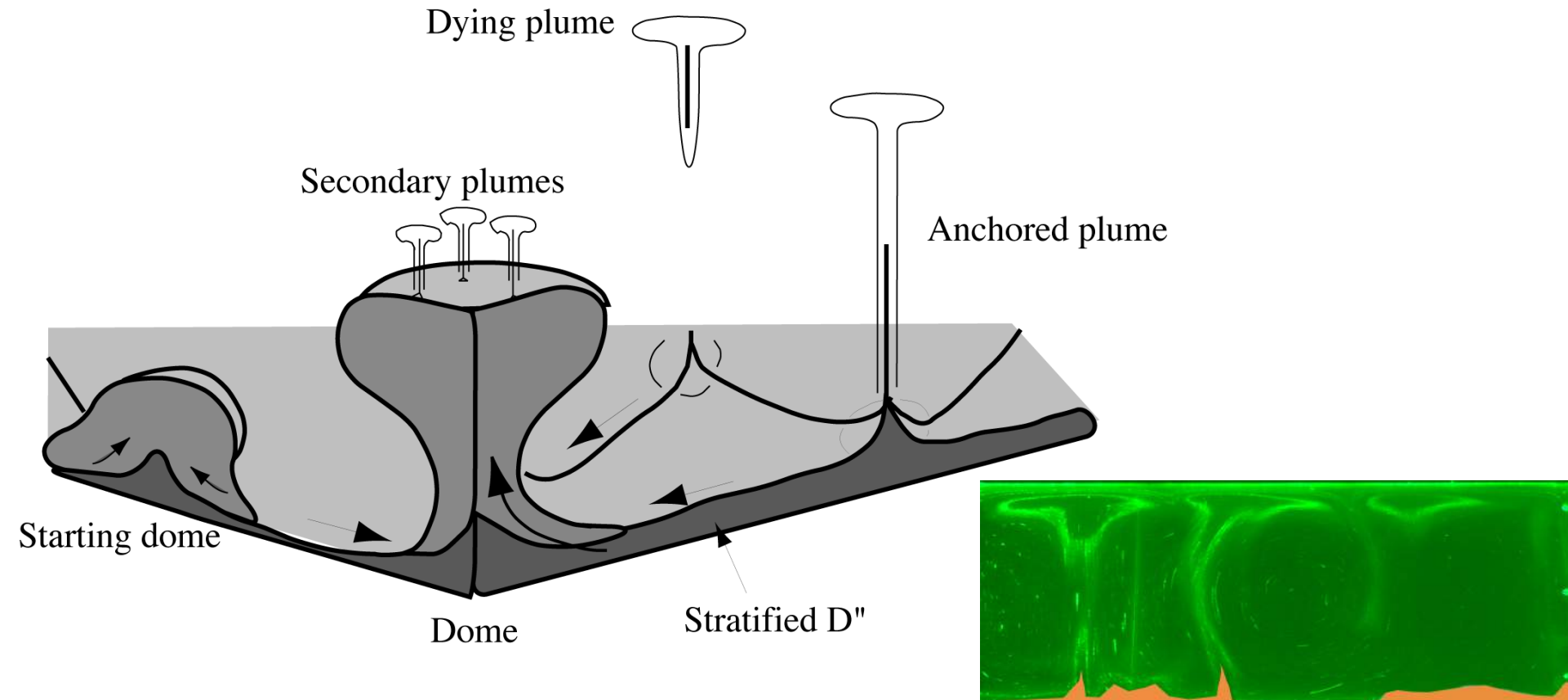
=> modify the level of neutral buoyancy



(Tan & Gurnis, 2008; Samuel & Bercovici, 2005; Ballmer et al, 2013; ...)

## 4. Interaction Convection / Denser Reservoir

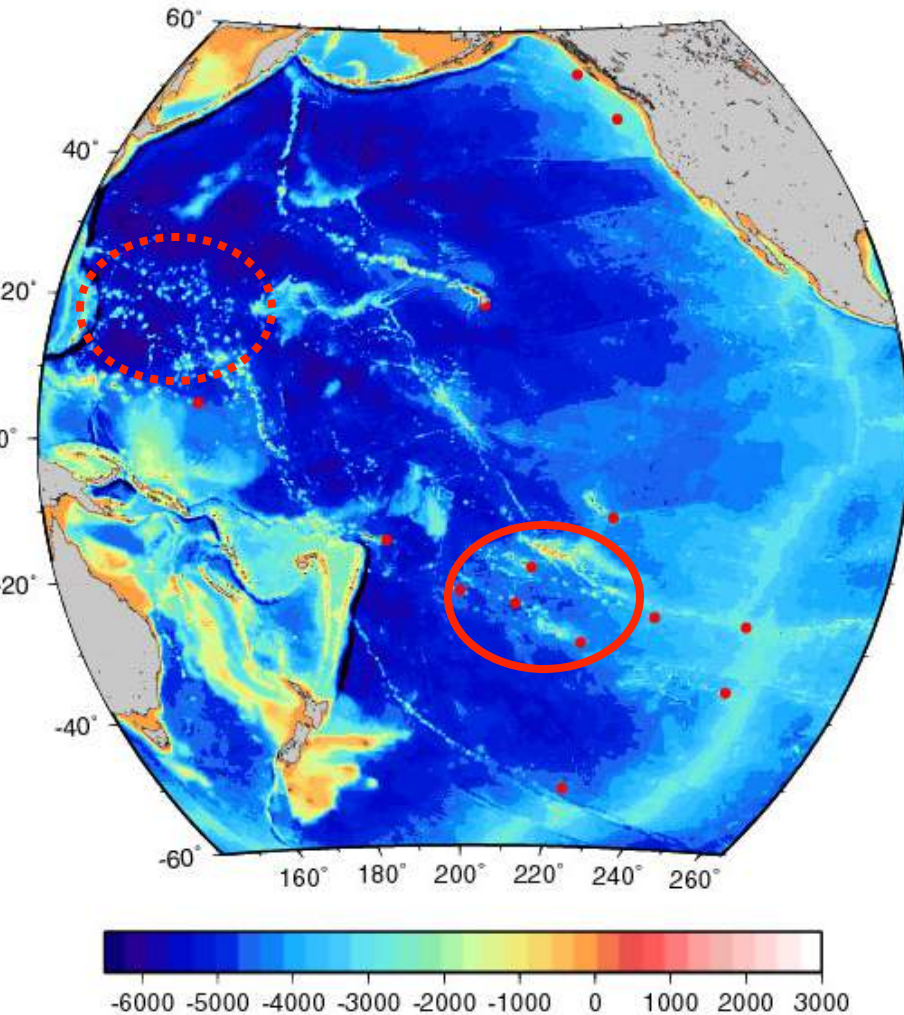
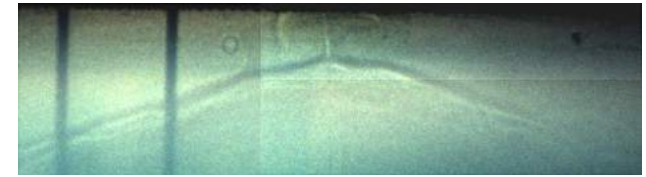
Coexistence of several types of hot instabilities



Laboratory experiments scalings  $\Rightarrow$  spacing  $\sim 2000-3500$  km at CMB

# 5. Inside the boxes: hot upwellings in the Pacific

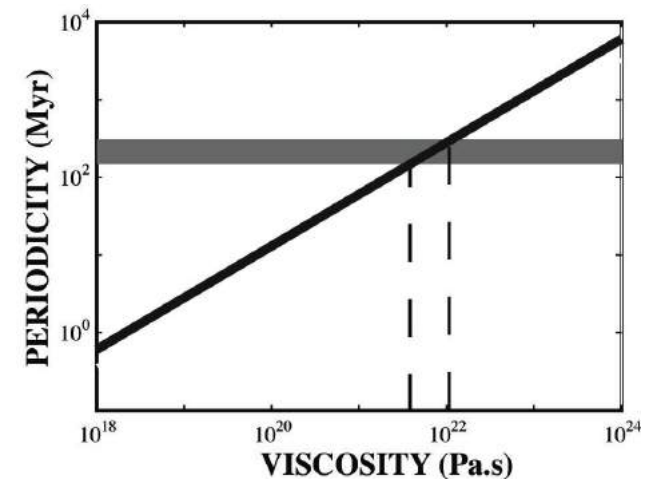
Polynesia = hot spot cluster (short tracks) + superswell



Darwin Rise = 100 Ma

⇒ = 1 complete pulsation of a thermochemical dome

if Viscosity  $\sim 2-6 \cdot 10^{21}$  Pa.s



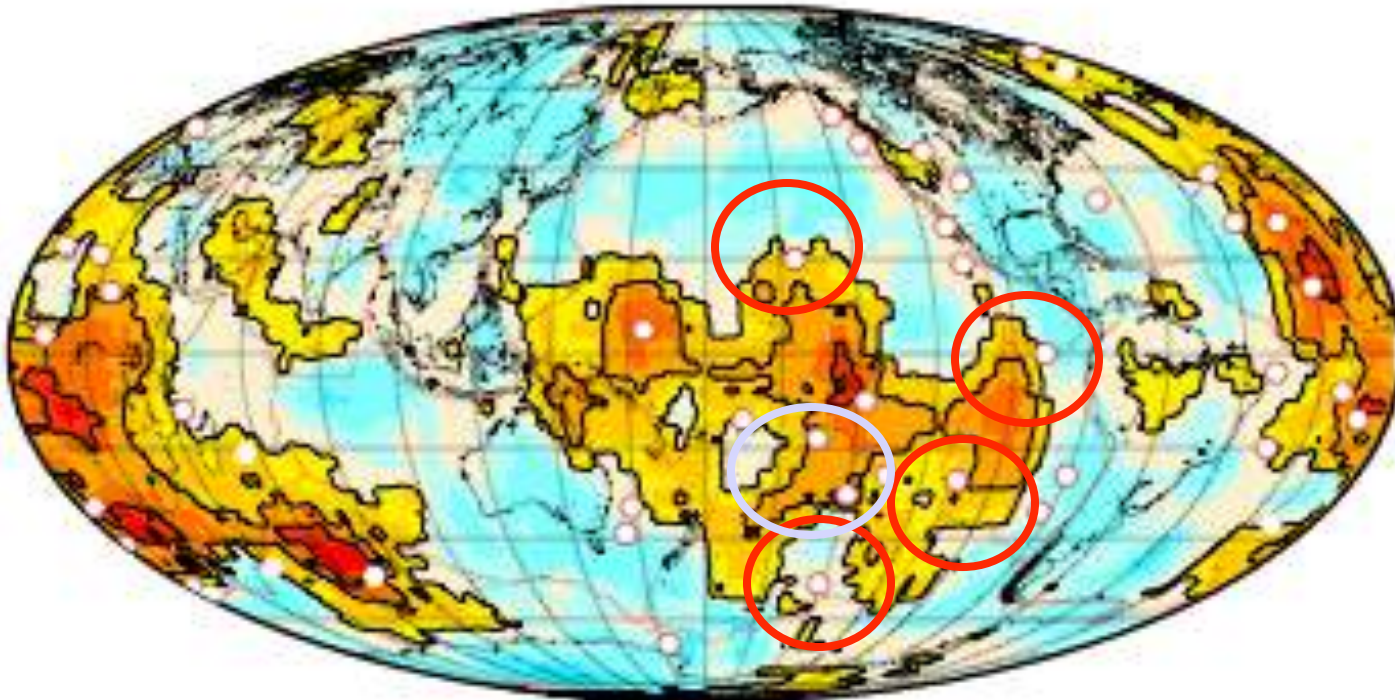
Pacific Bathymetry (Clouard & Bonneville, 2003)

(Davaille, 1999)

## 5. Inside the boxes: hot upwellings in the Pacific

Thermochemical instabilities => spacing ~2000-3500 km at CMB

TXBW



= > **missing** 2 or 3 instabilities

## 5. Inside the boxes: hot upwellings in the Pacific

missing instability

Explain the data

=> heavy dome => descending

=> + would have been under part of Line Island ~ 35 Myr ago

+ 140 Myr ago, Shatsky Rise was formed in the area

=> Another ~100 Myr cycle...

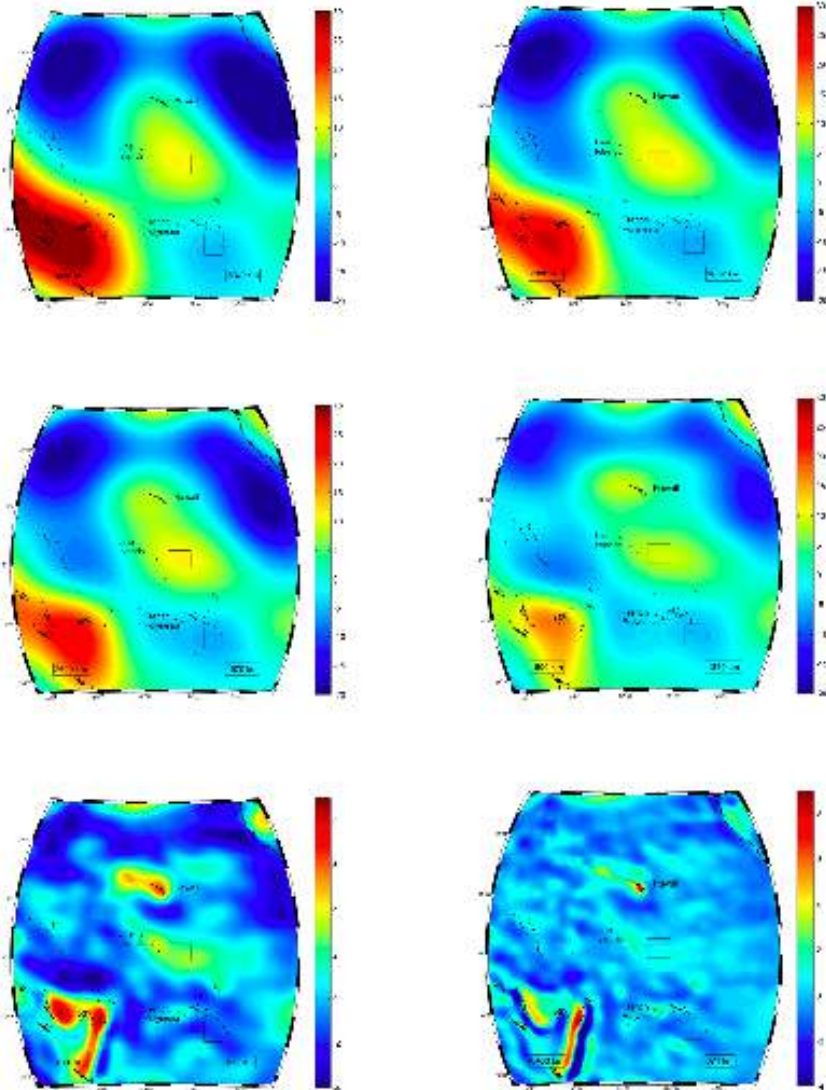
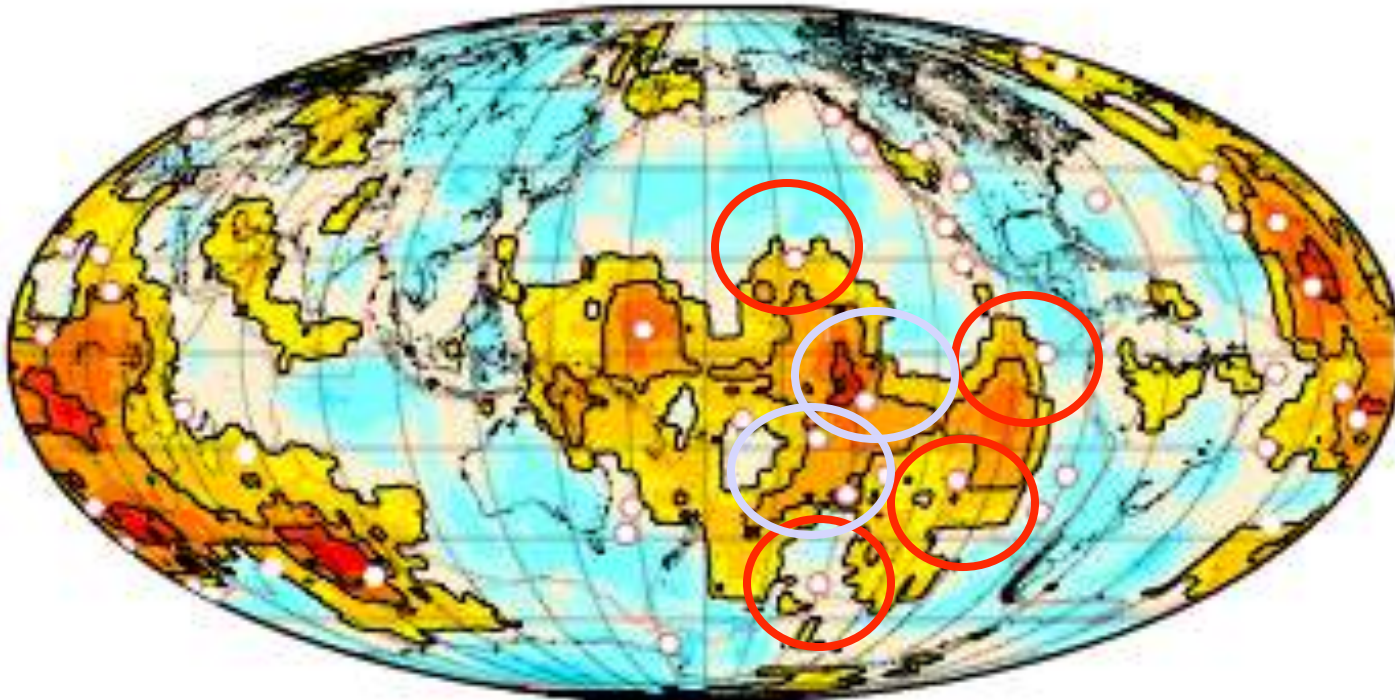


Figure 3. Continuous wavelet analysis of the EIGEN-GL04C geoid model in the Pacific Ocean at varying scales, from 3600 to 400 km. The analyzing wavelet scale and the

## 5. Inside the boxes: hot upwellings in the Pacific

Thermochemical instabilities => spacing  $\sim 2000-3500$  km at CMB

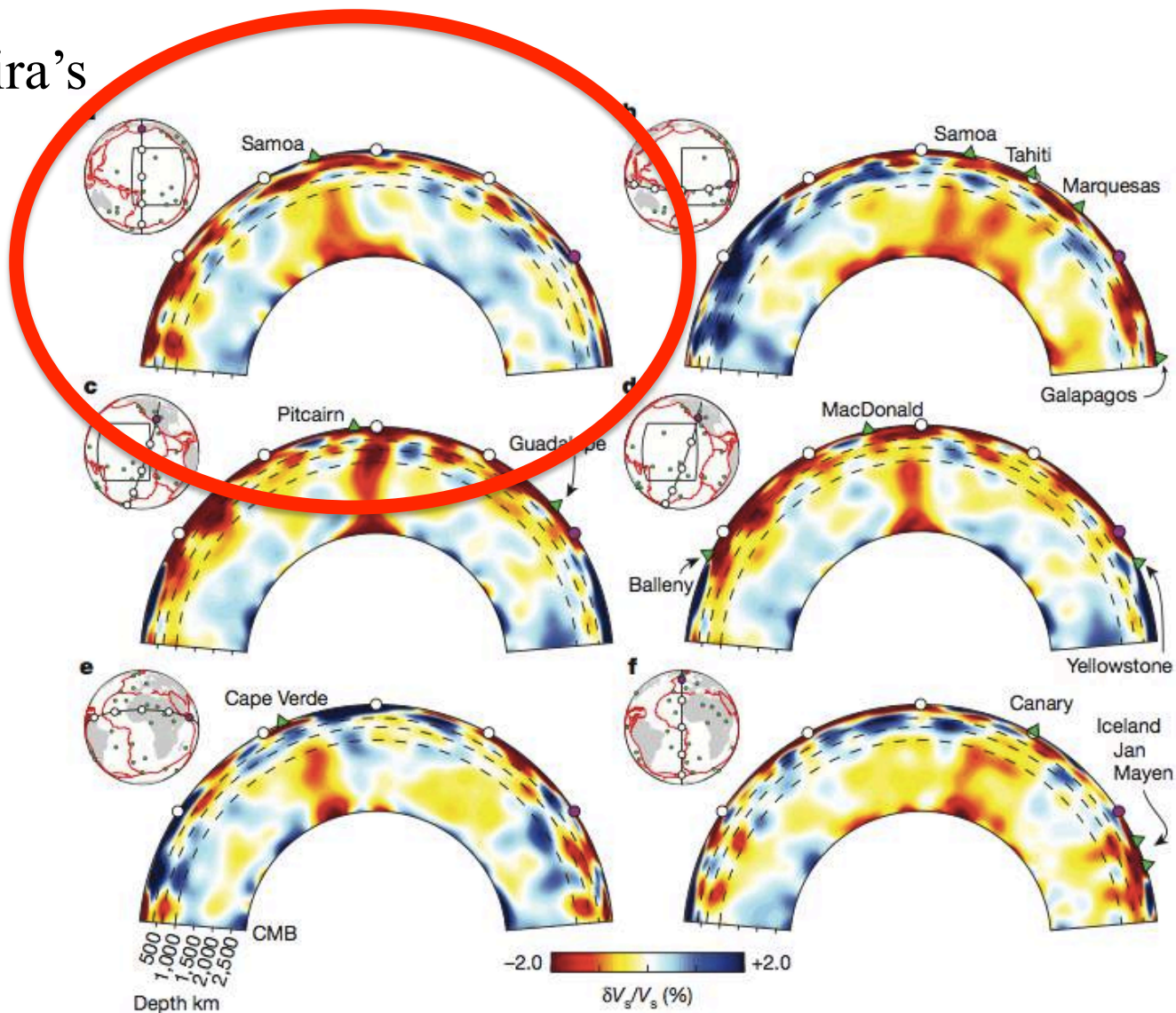
TXBW



= > **missing** 1 or 2 instabilities

# 5. Inside the boxes: hot upwellings in the Pacific

A. Ferreira's  
talk



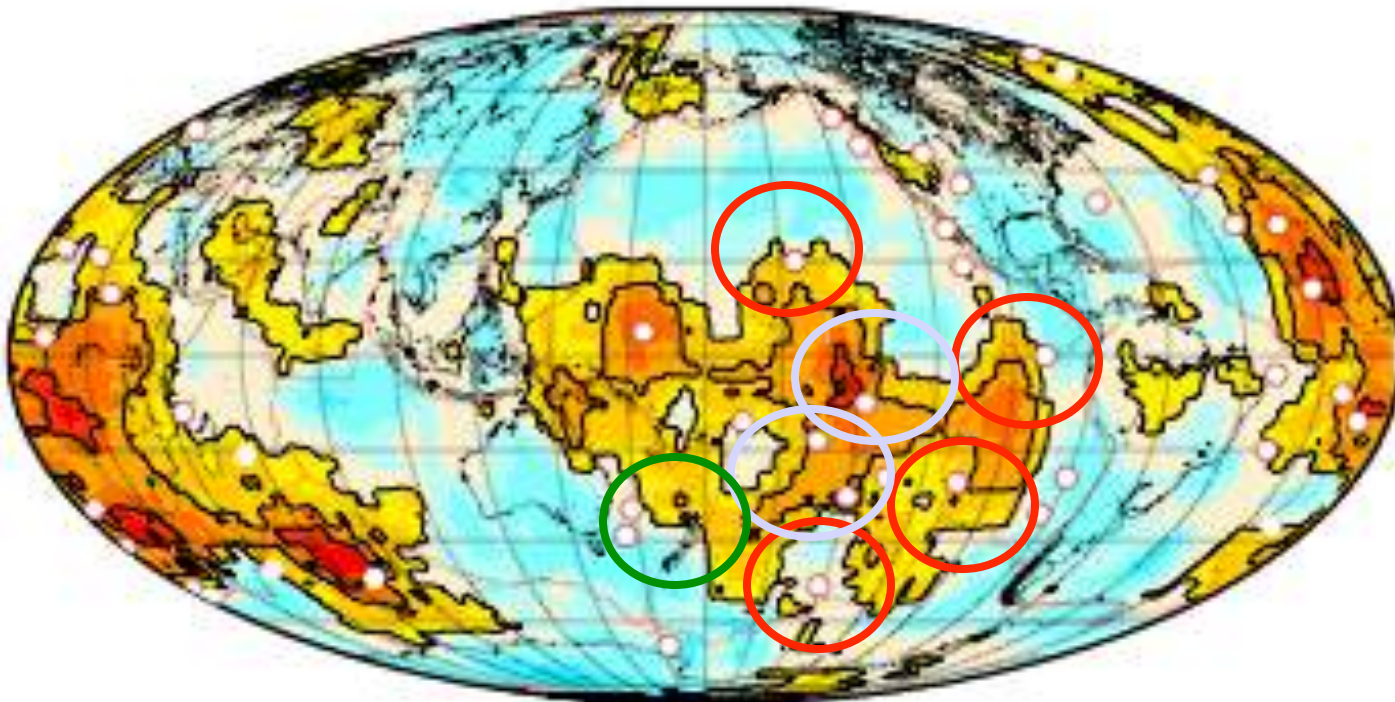
(French & Romanowicz, 2015)



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TXBW

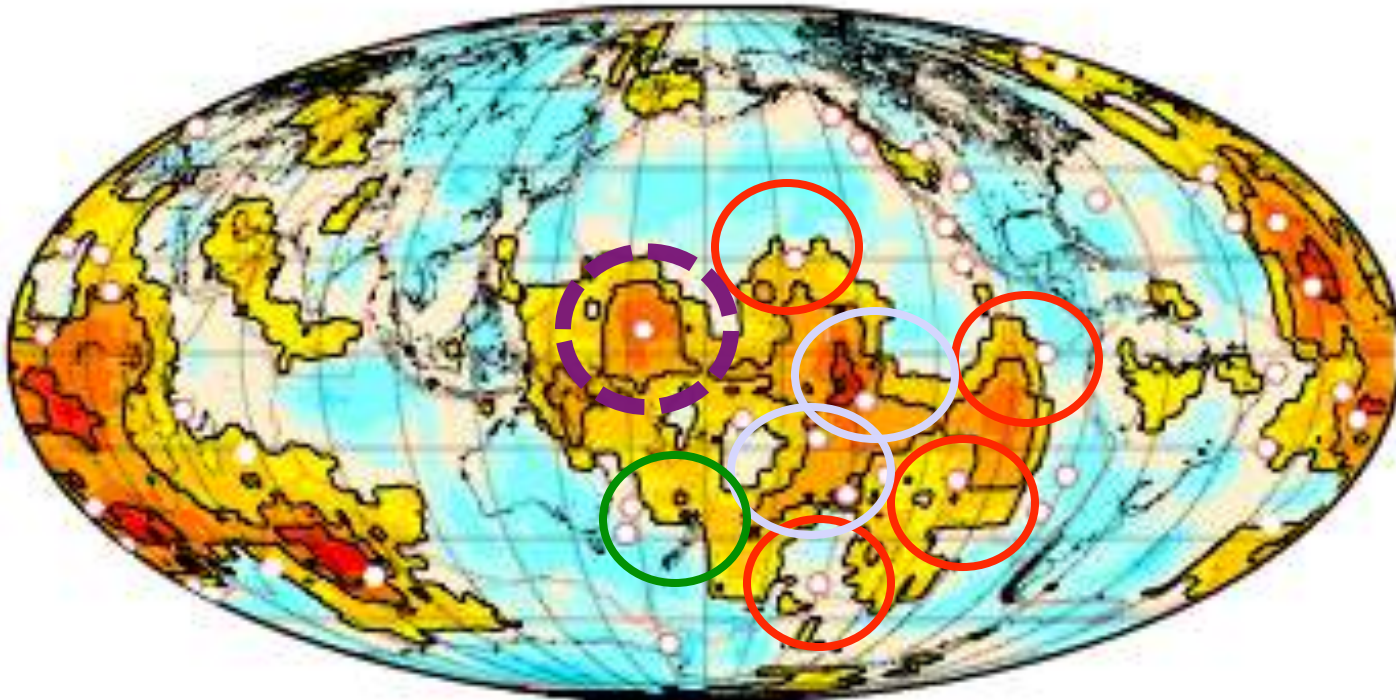


= > **missing** 1 instability

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Thermochemical instabilities => spacing  $\sim 2000$ - $3500$  km at CMB

TXBW

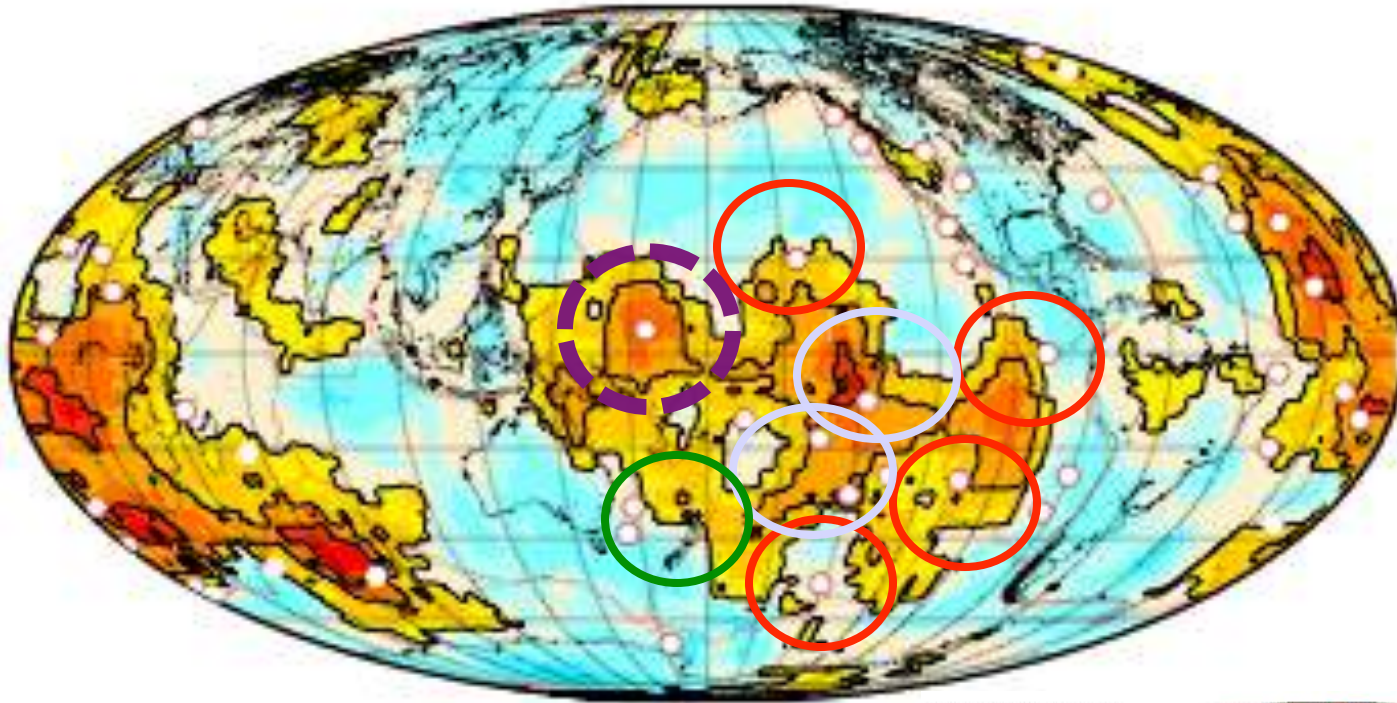


**MOVIE** => instability has not reached the surface yet

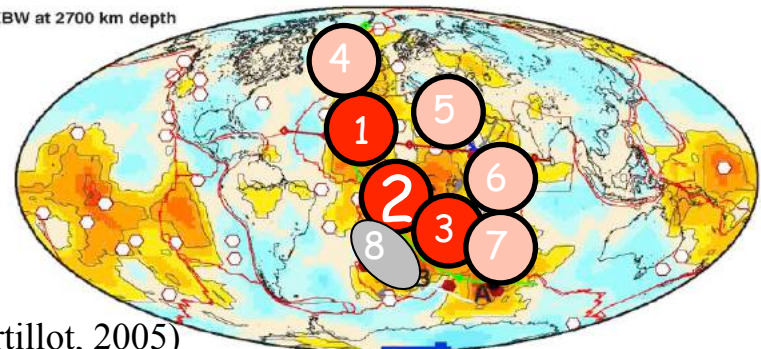
# 5. Inside the boxes:

Thermochemical instabilities => spacing ~2000-3500 km at CMB

**TXBW**



TXBW at 2700 km depth



# QUESTIONS :

**A- Origin of mantle « boxes » ? Variable viscosity / Subduction**

**B- Upwellings                      Compositional/density heterogeneities**

- several types ?

- « fat » ?      **Composition / viscosity /circulation within the plume**

- what happens around 1000 kms ?

**C- Do the mantle boxes / LLSVPs change shapes through time ?**

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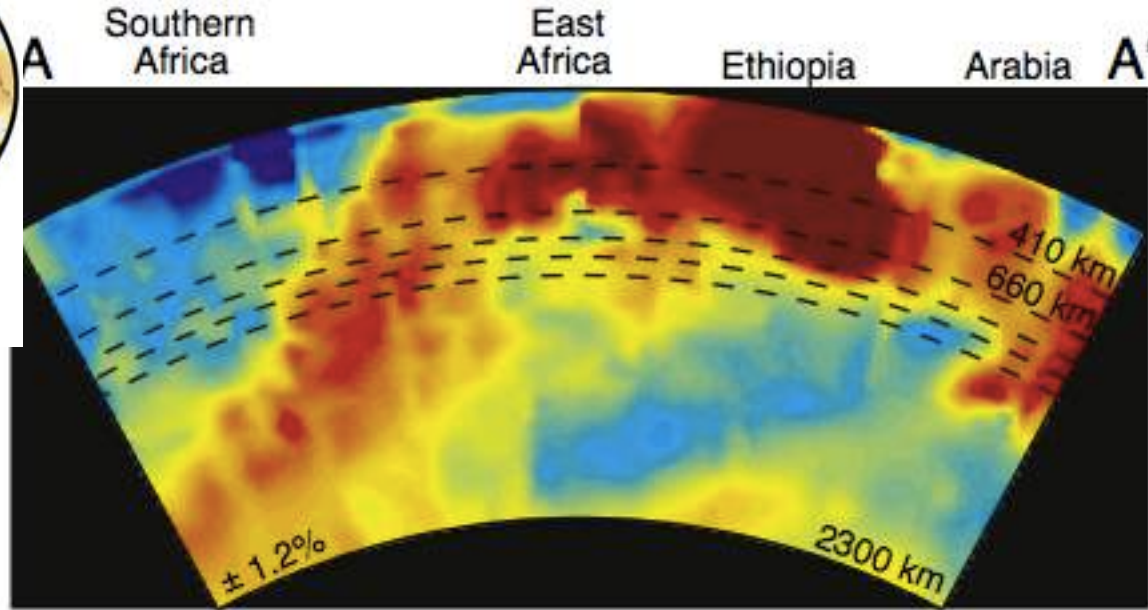
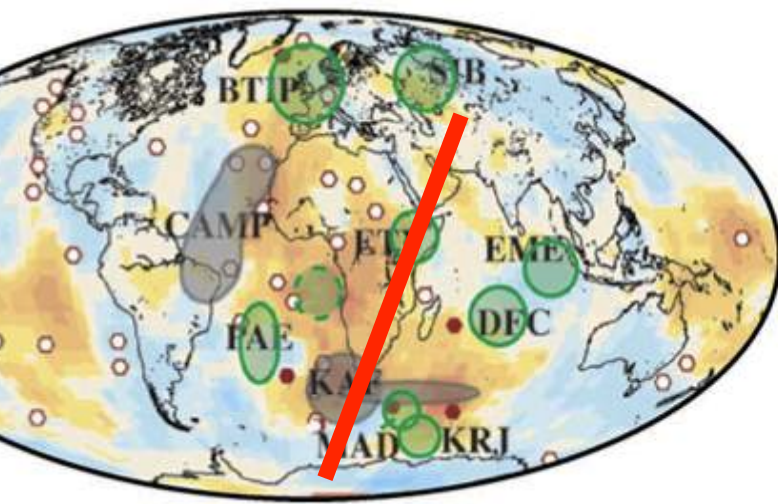
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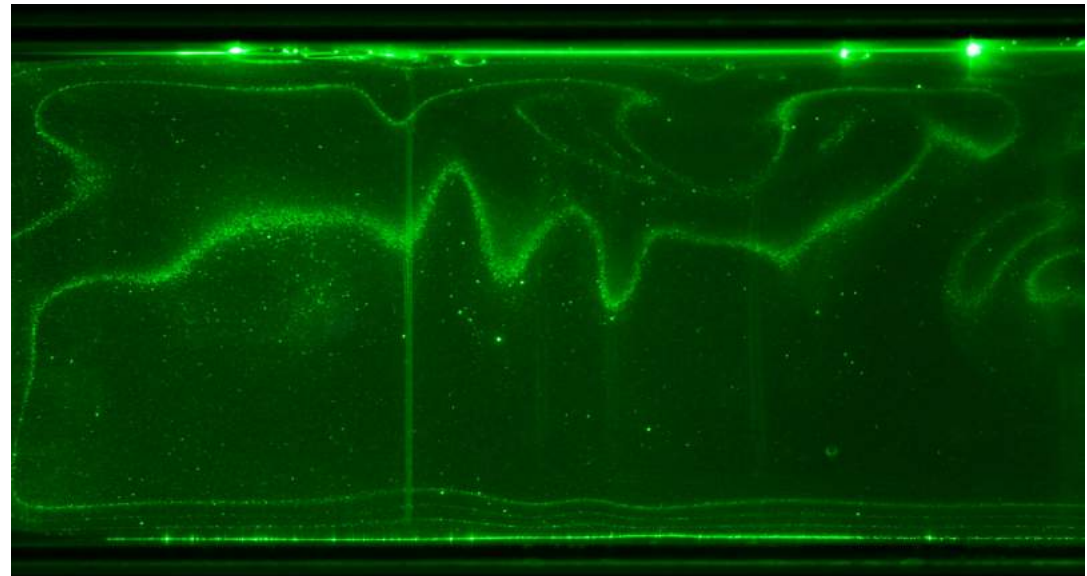
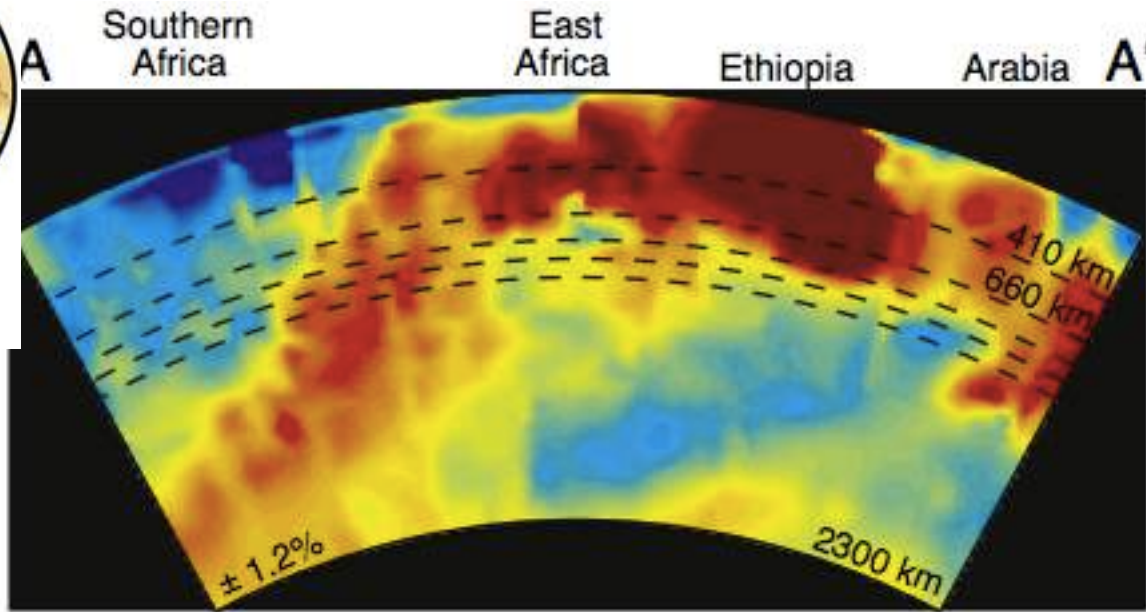
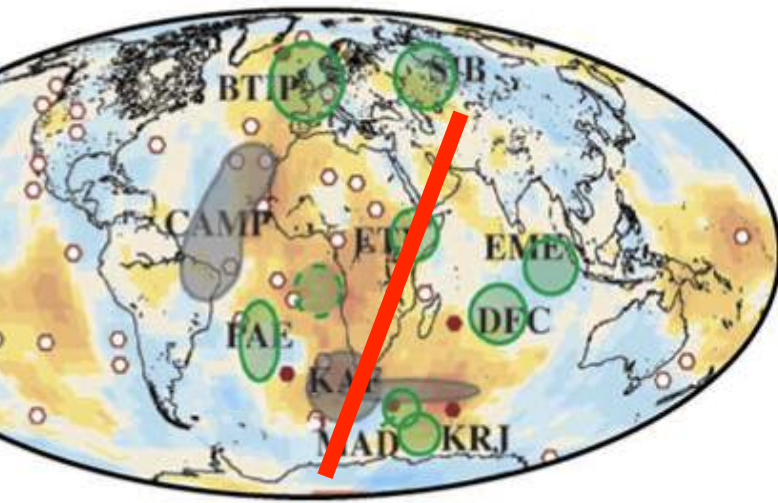
**C- Do the mantle boxes / LLSVPs change shapes through time ?**

## 6. Time-dependence:

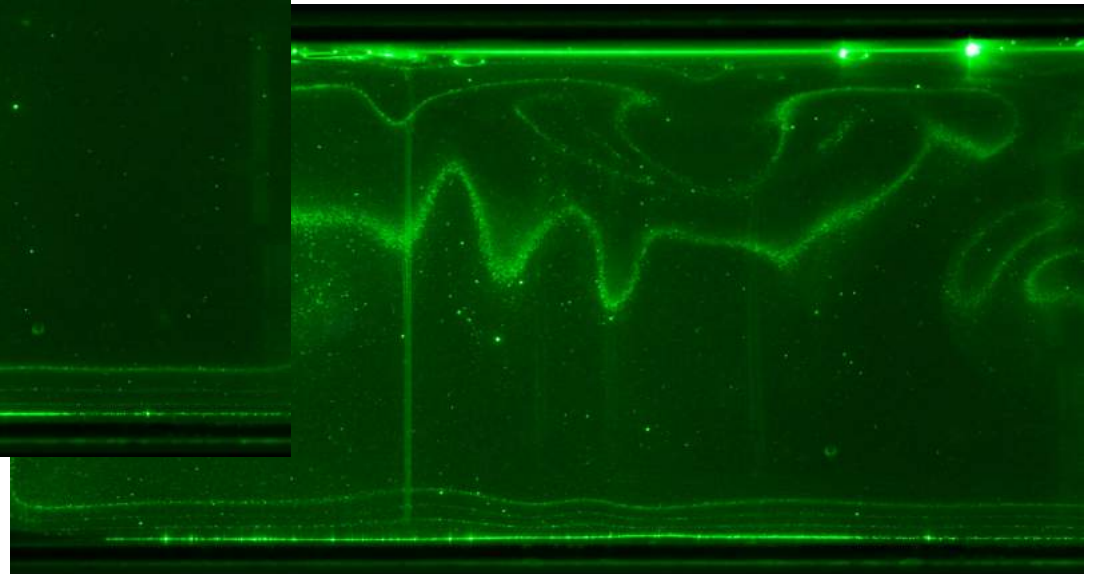
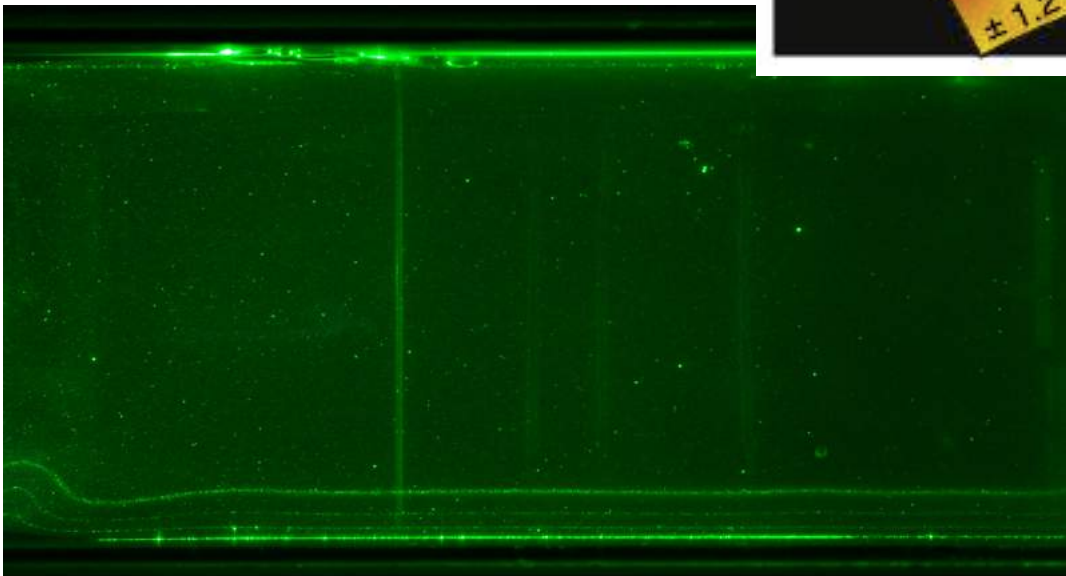
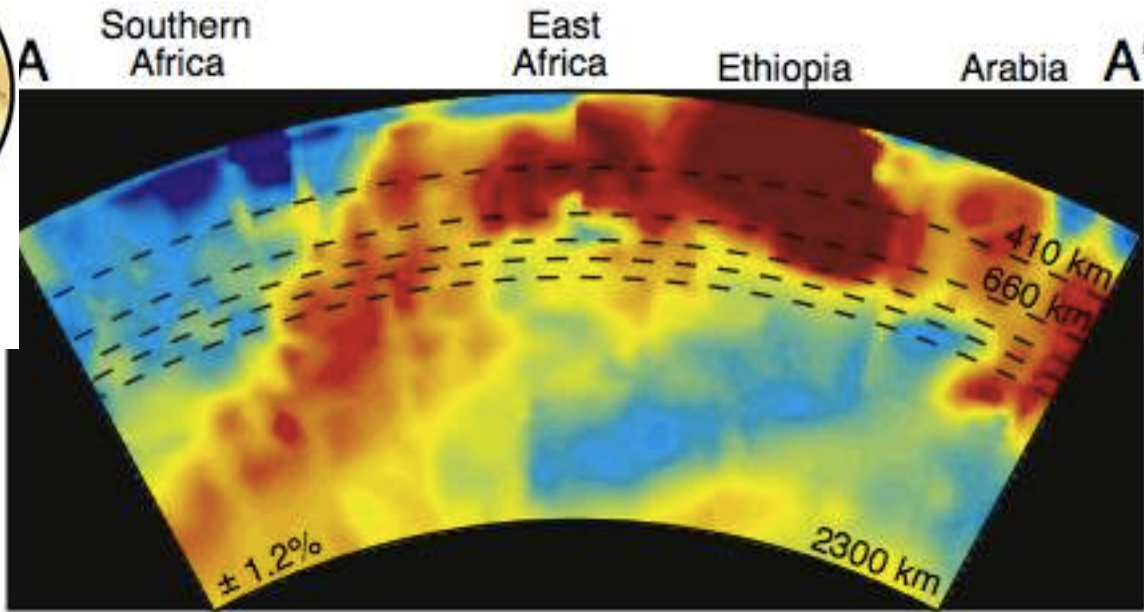
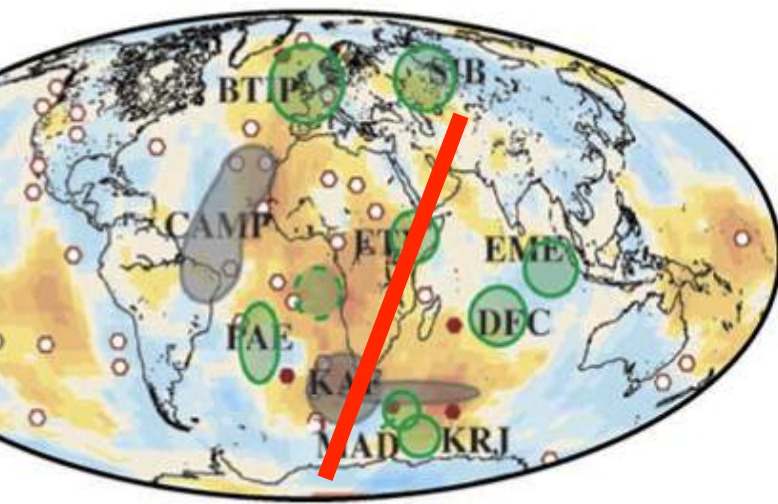


(Hansen et al, 2012)

## 6. Time-dependence:

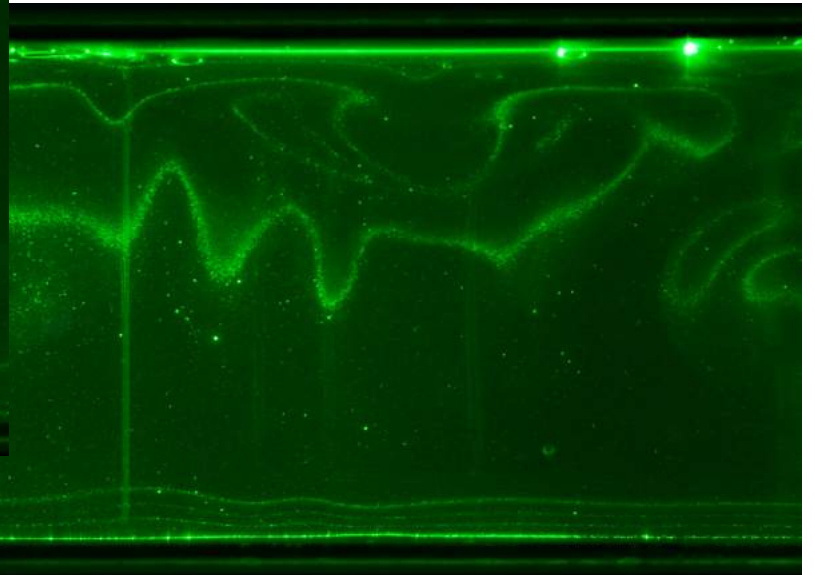
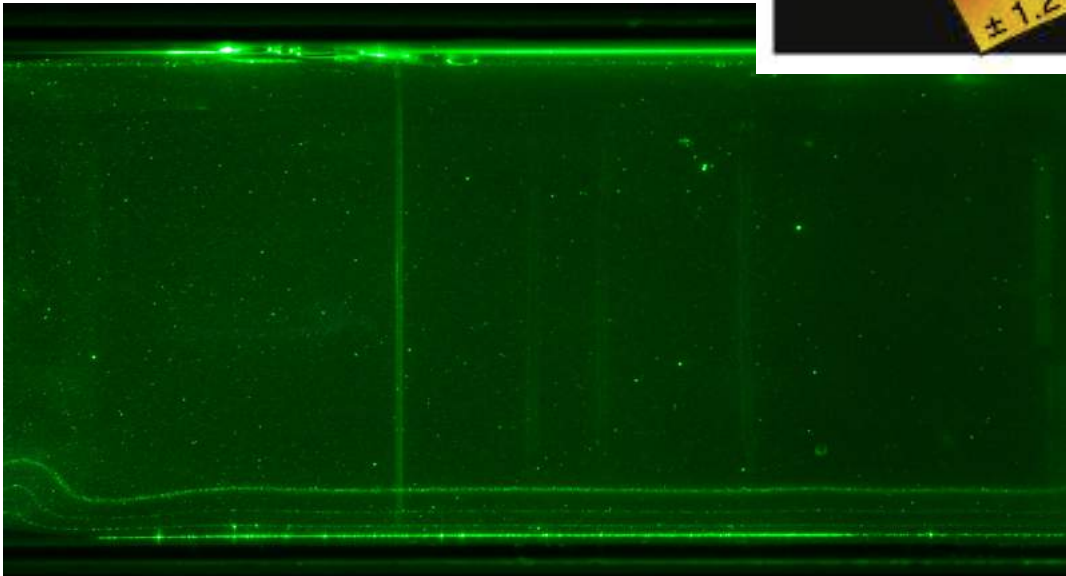
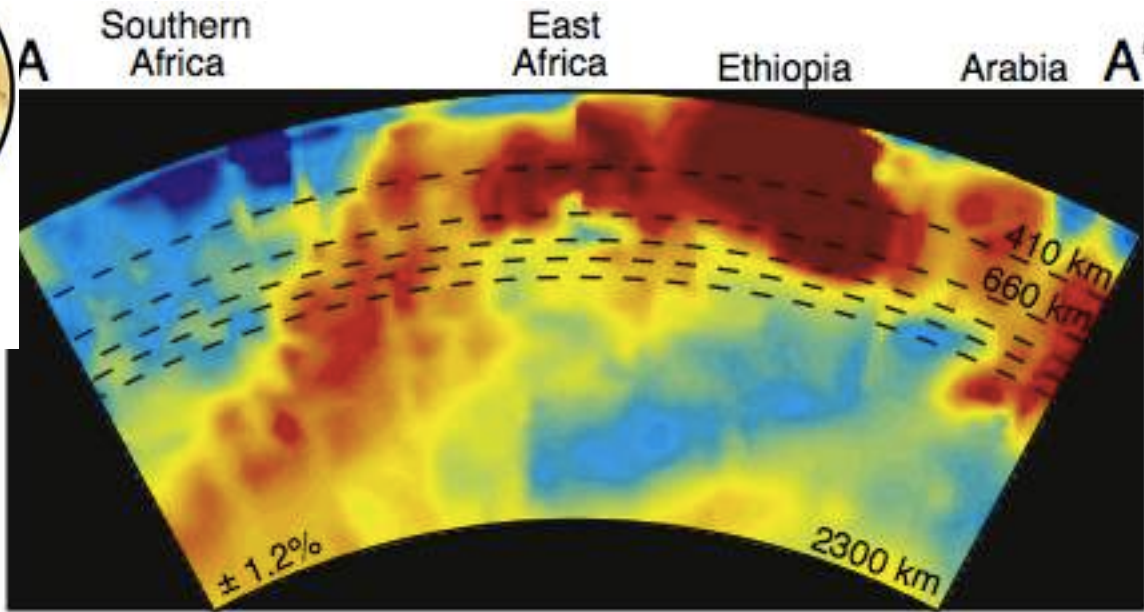
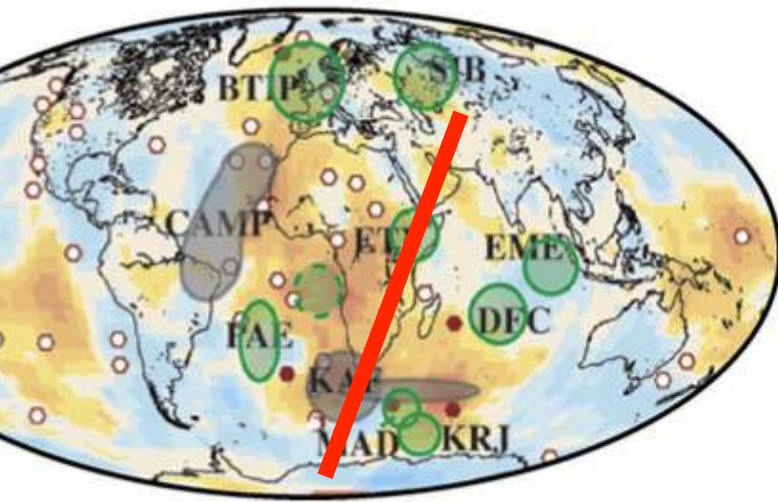


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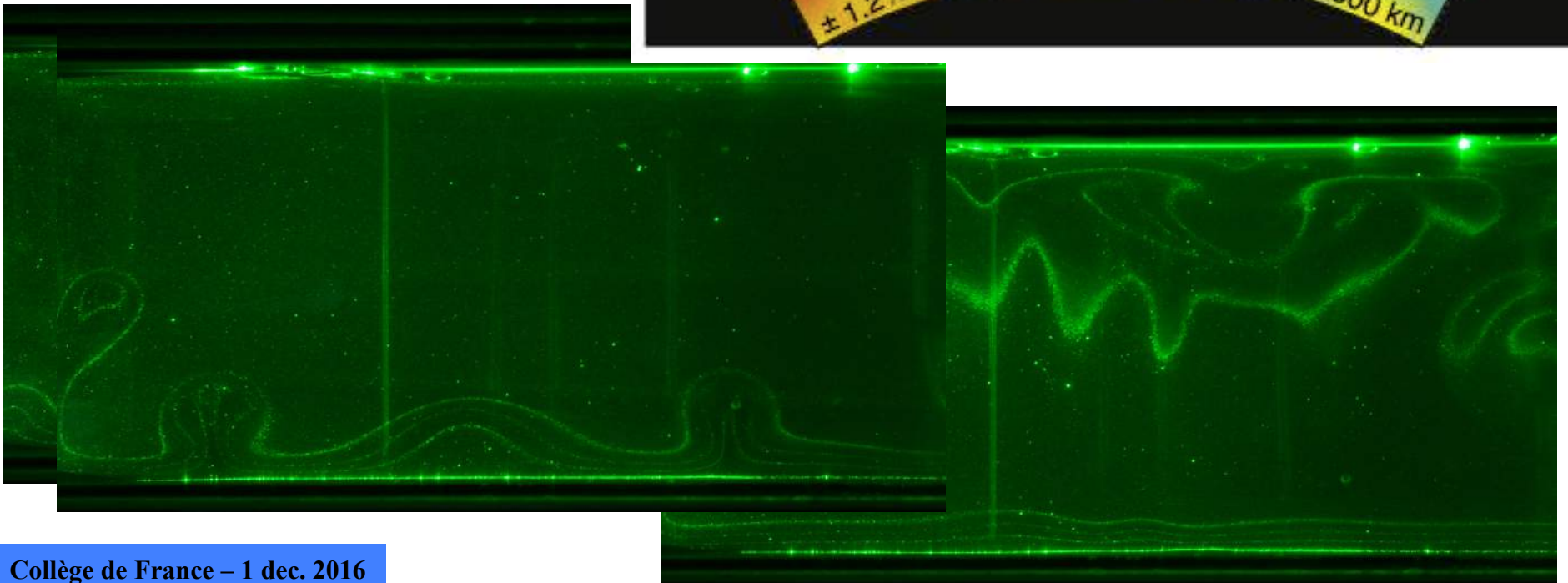
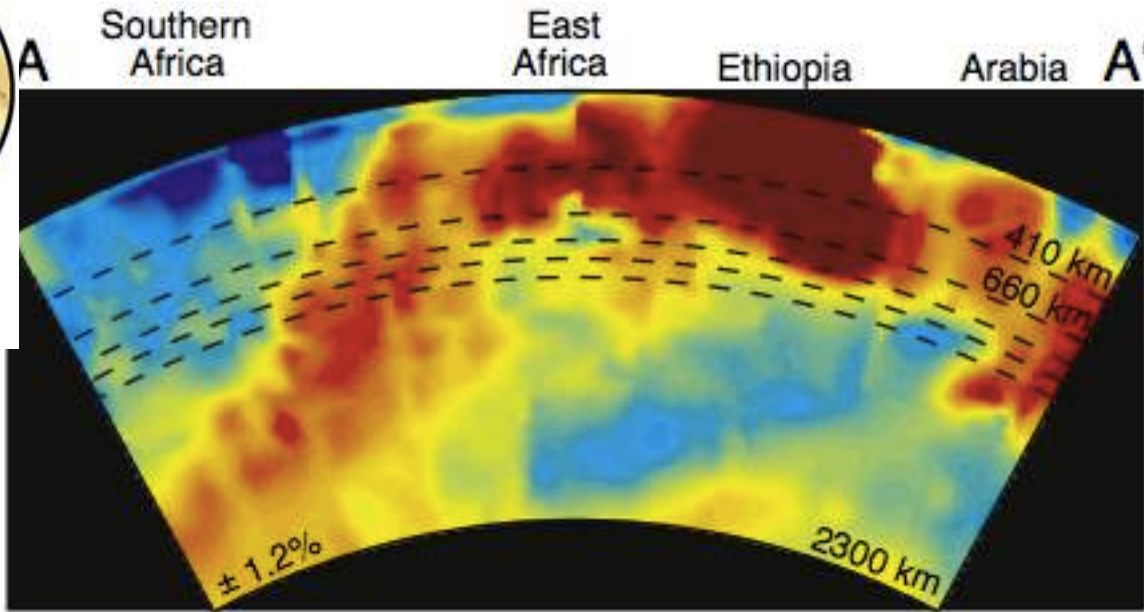
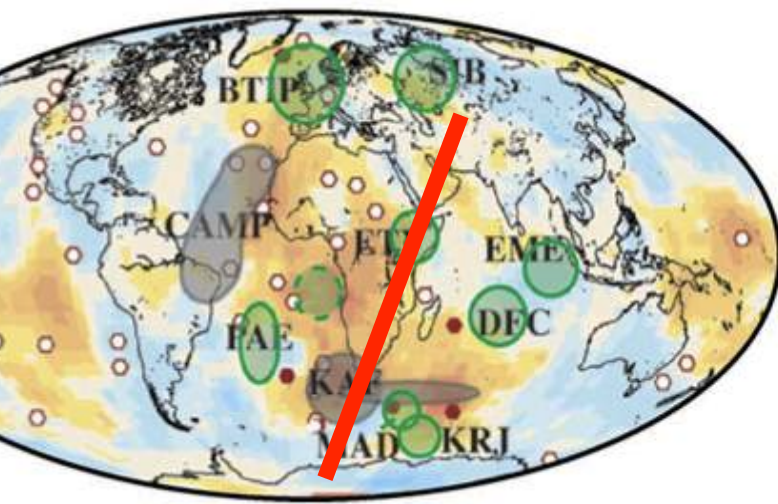




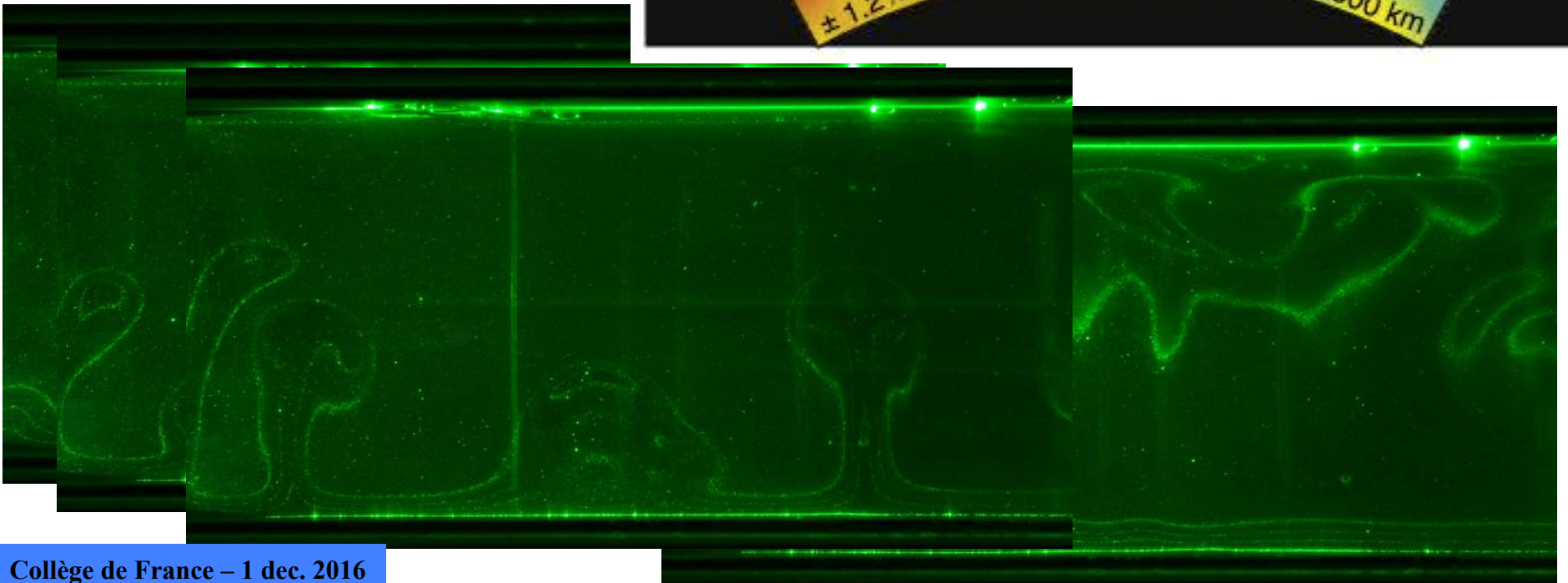
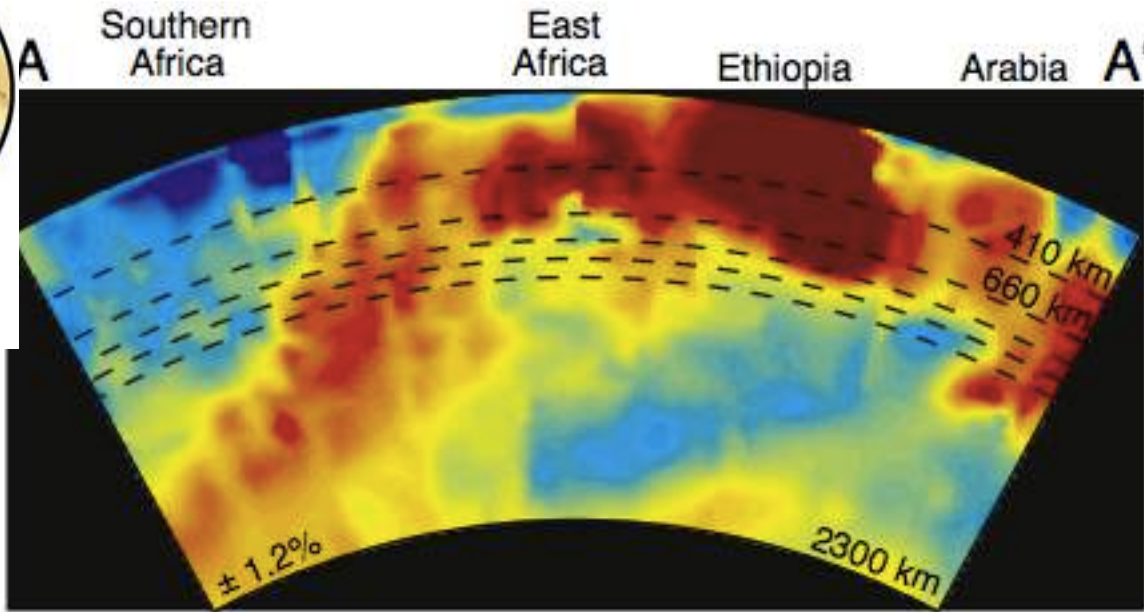
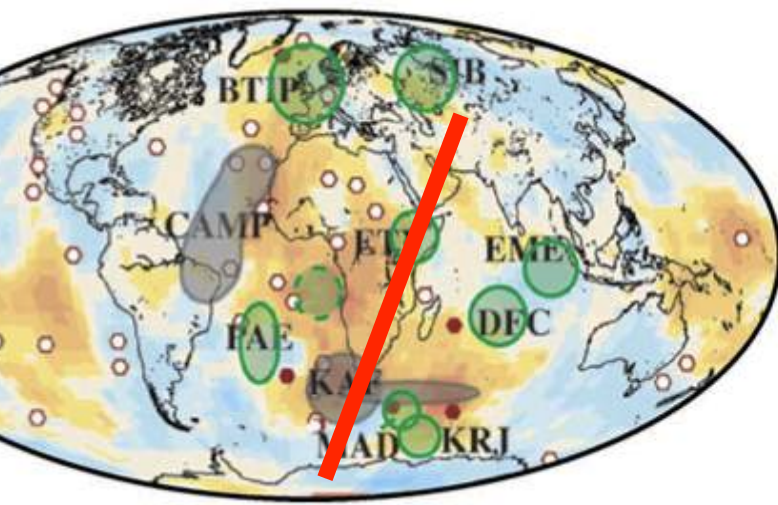
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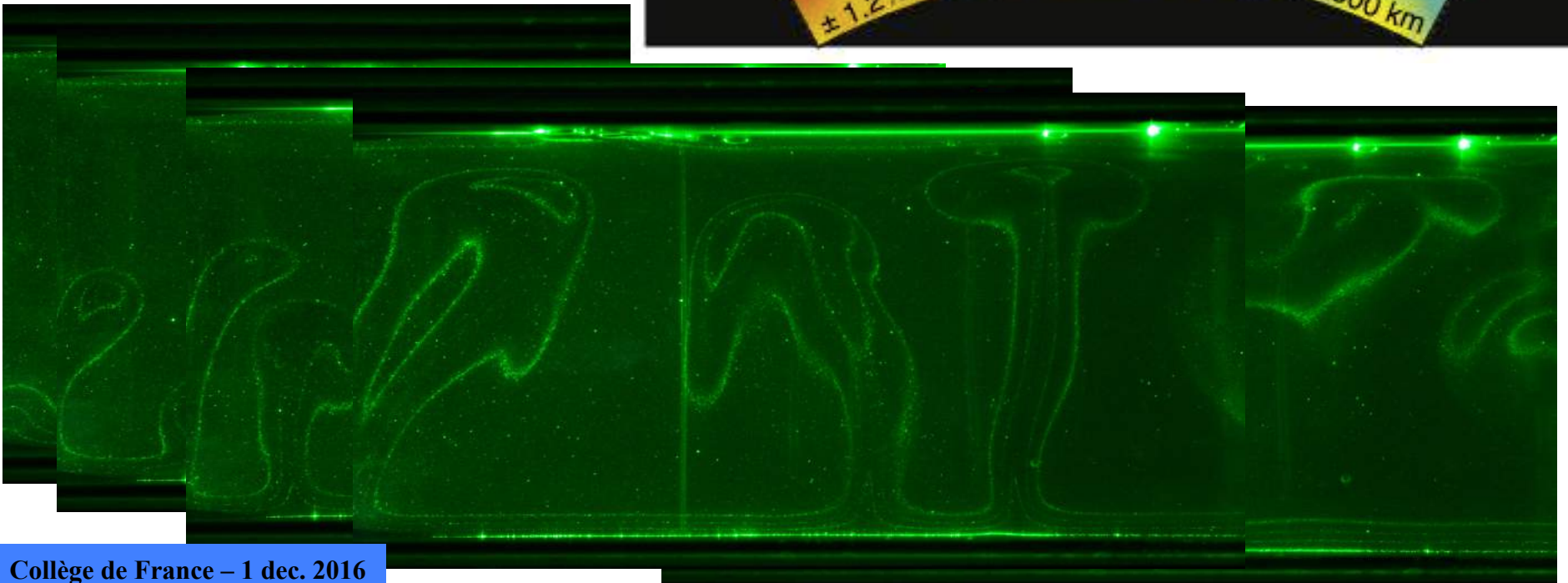
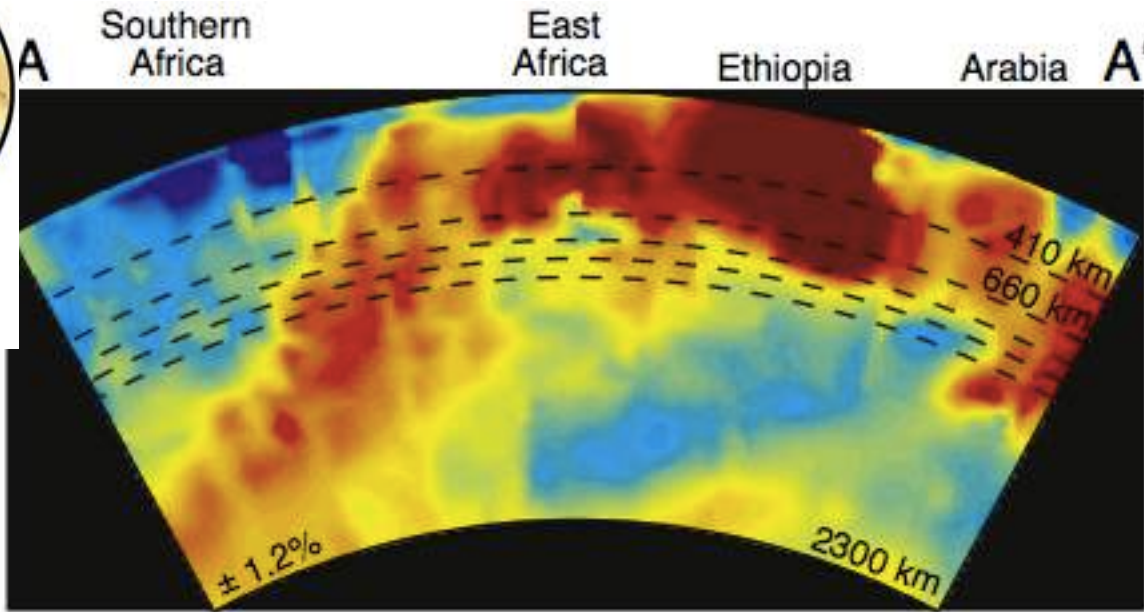
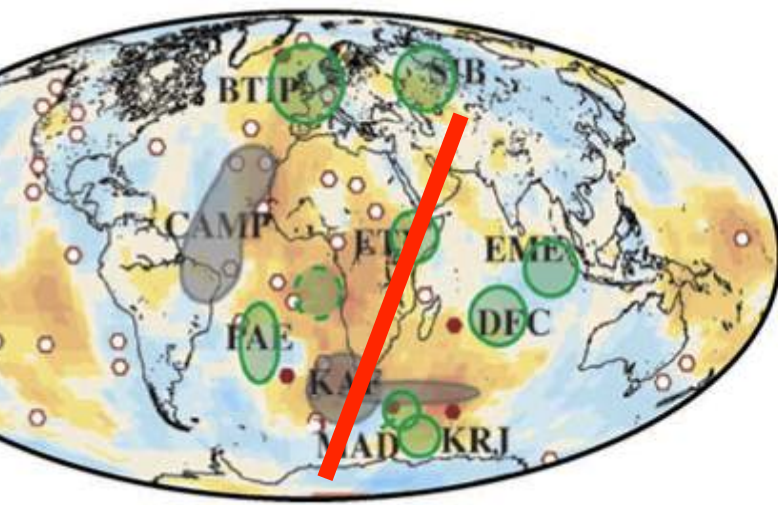
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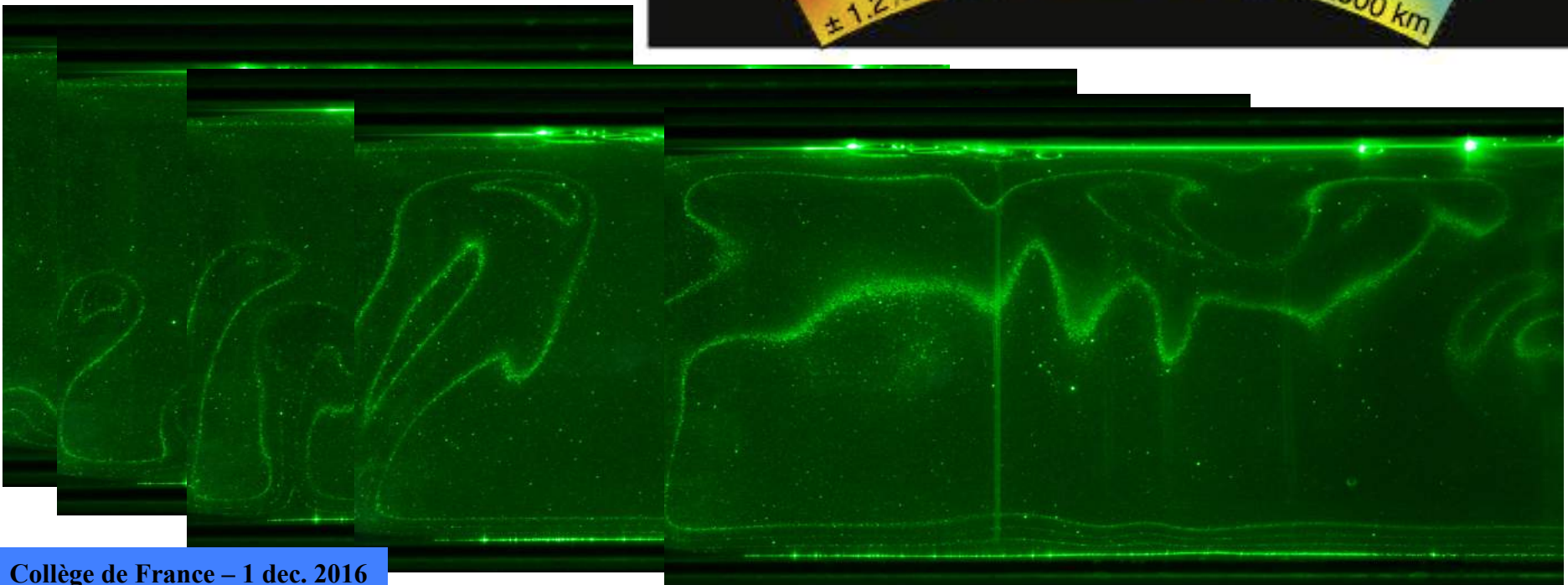
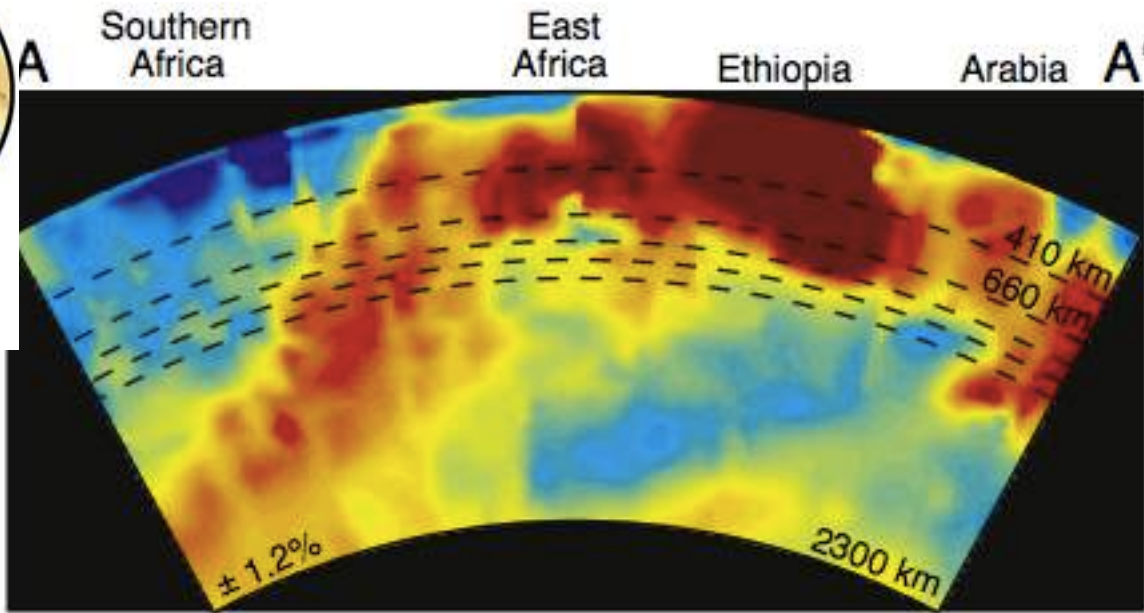
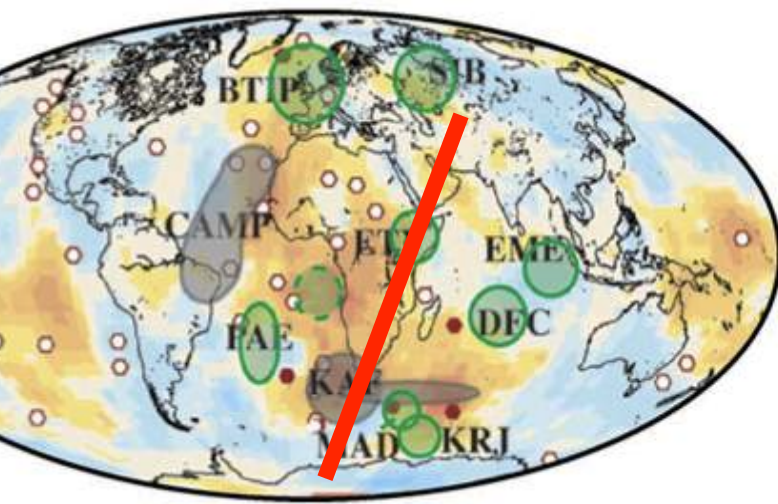
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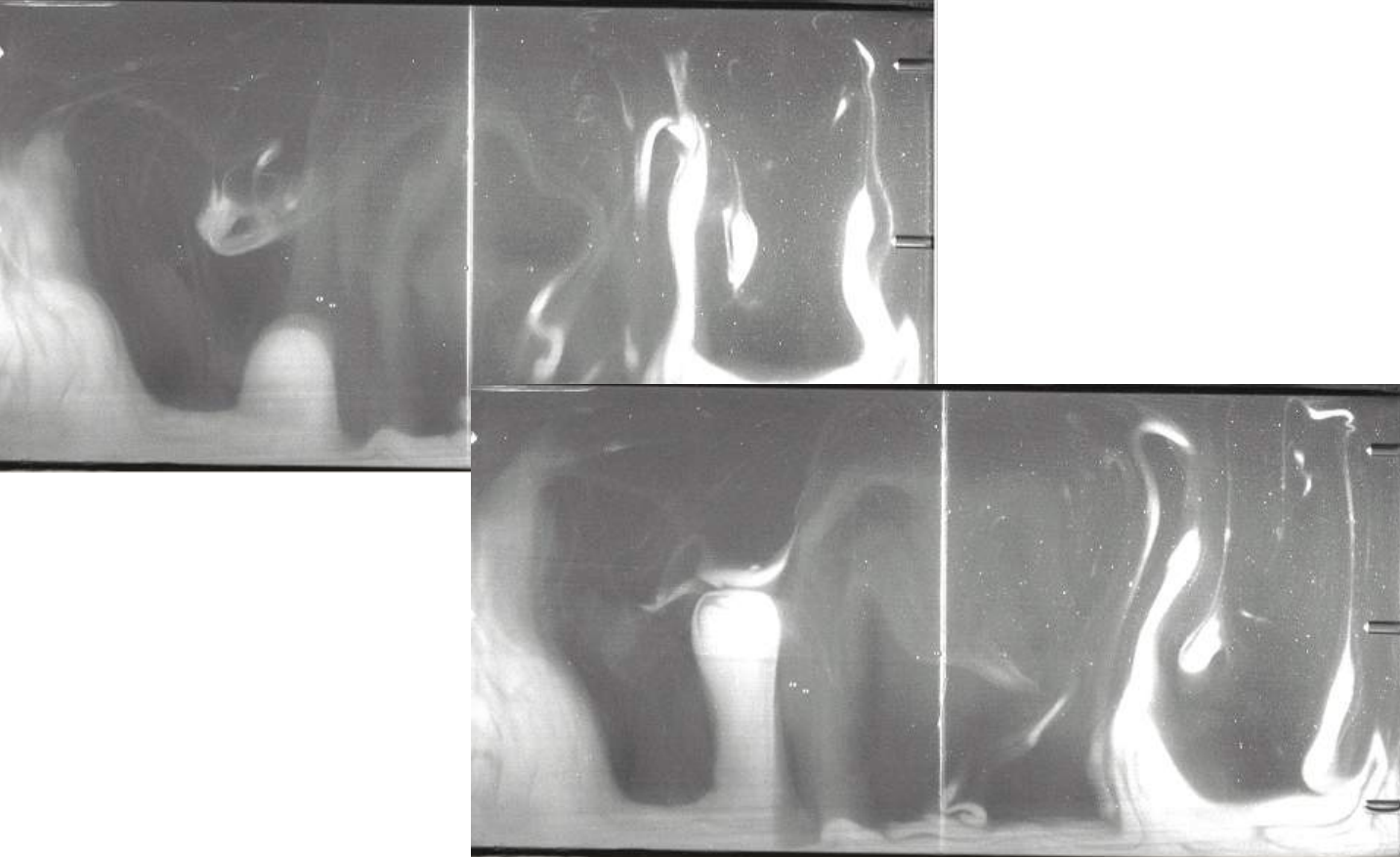
## 6. Time-dependence:

- Convection carries fast hot material from bottom to top  
Same for thermochemical instabilities



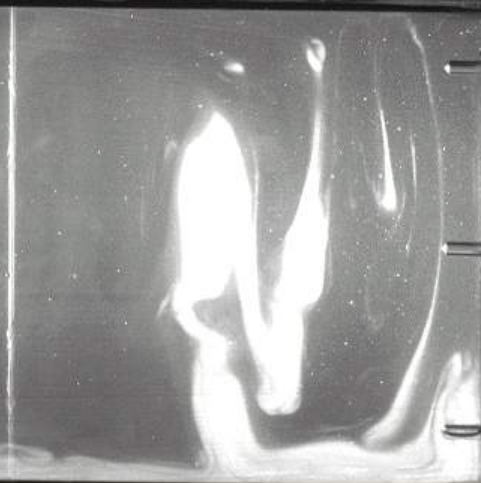
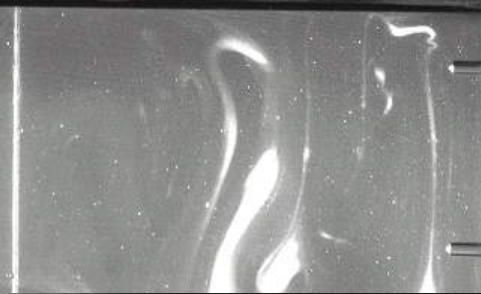
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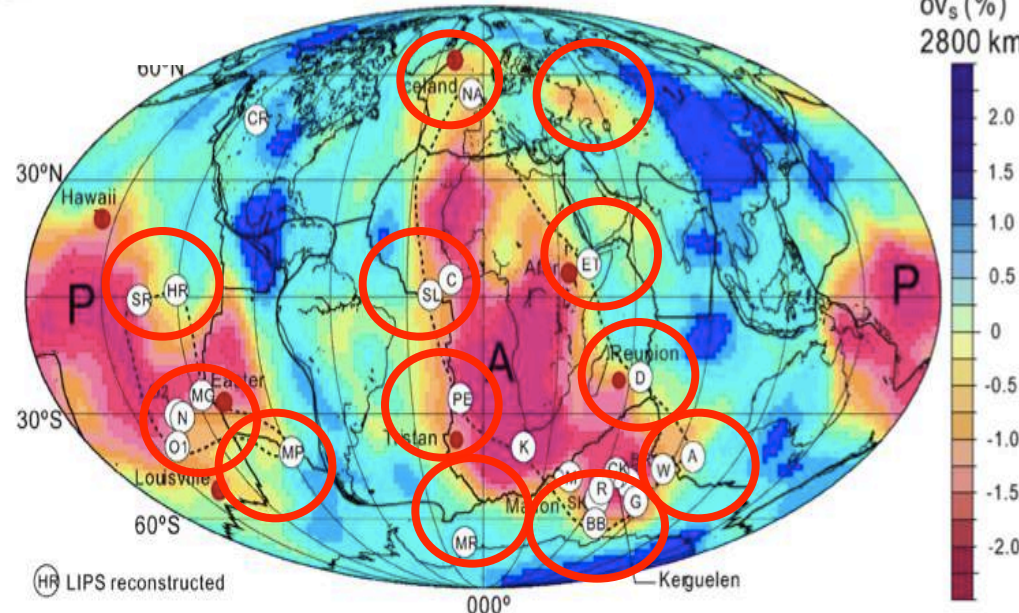
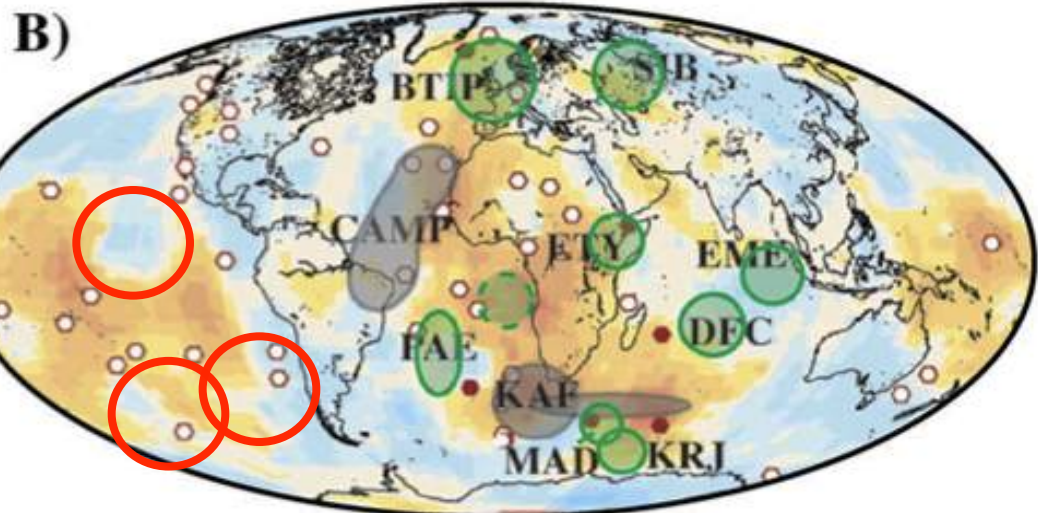




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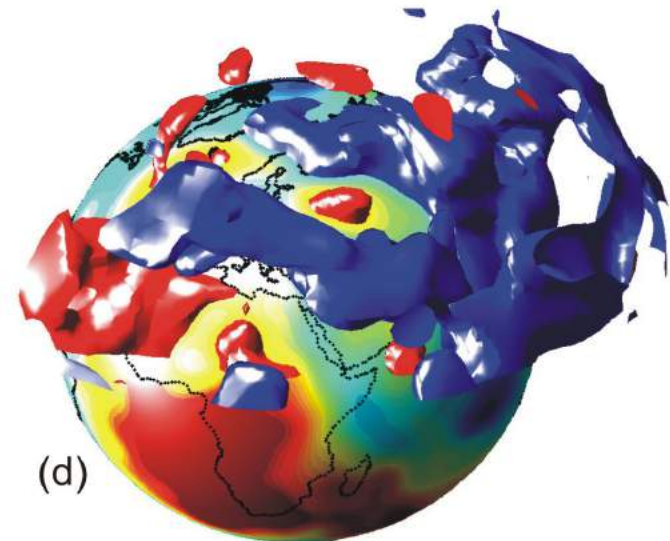
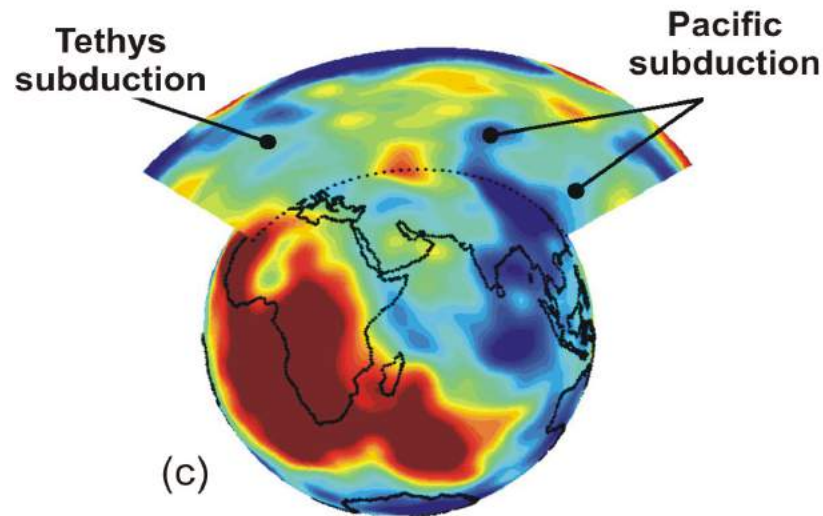
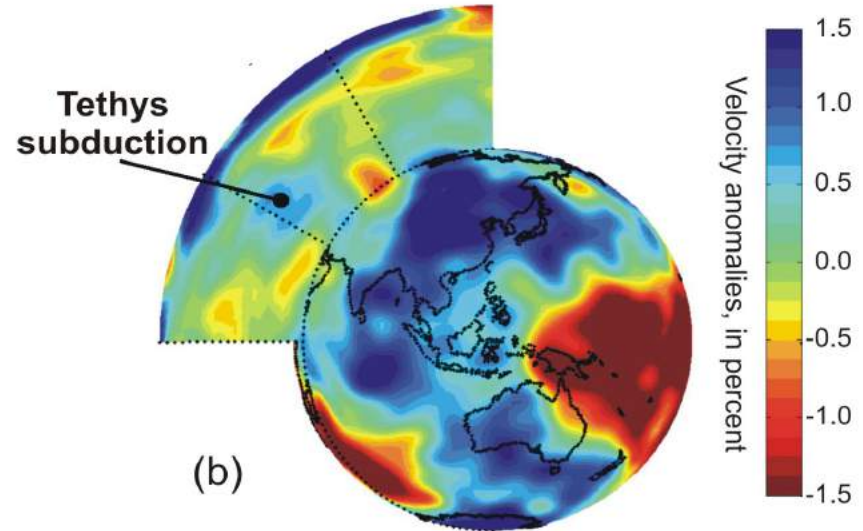
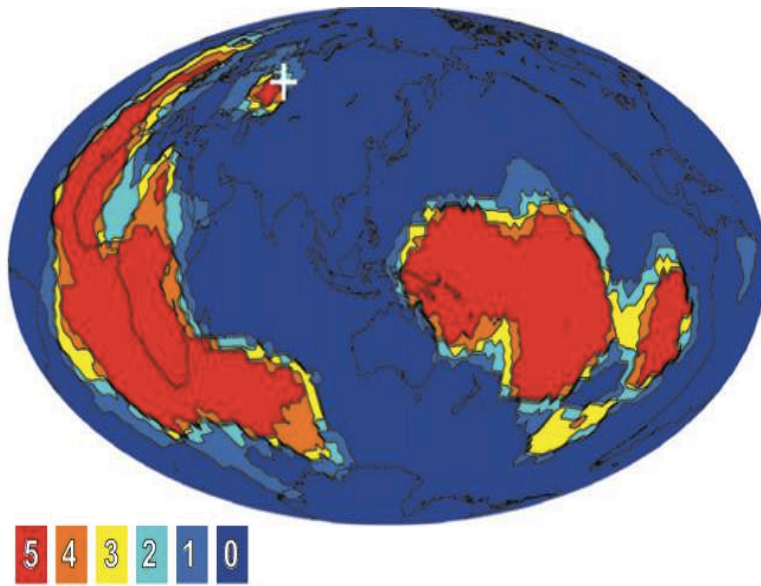
- Convection carries fast hot material from bottom to top  
=> LLSVP area decreases at each plume head (LIP) event

1 every ~ 30 Myr



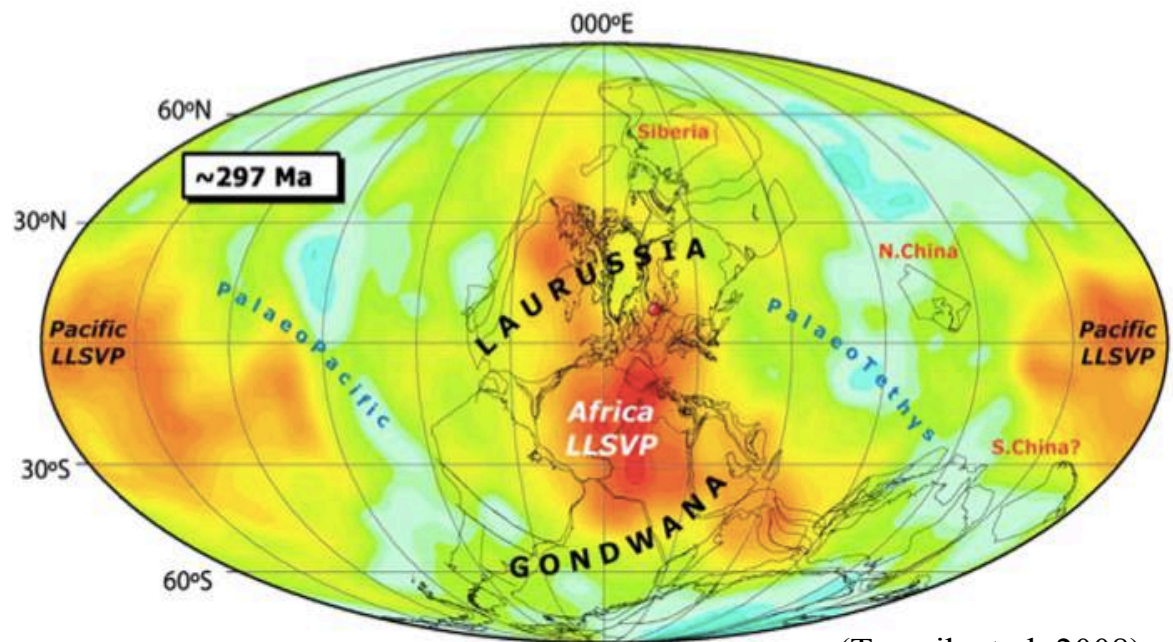
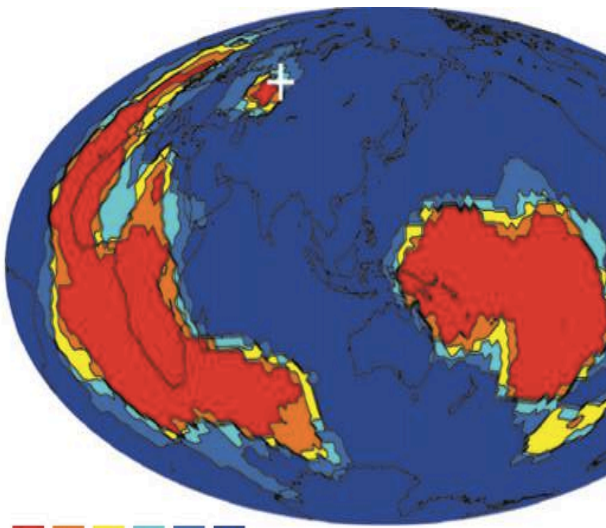
## 6. Time-dependence:

### Existence of a THIRD mantle box ?

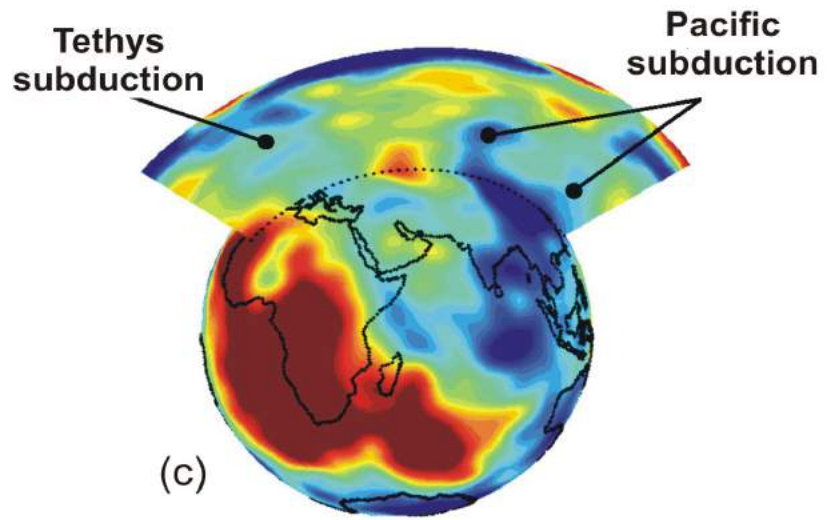


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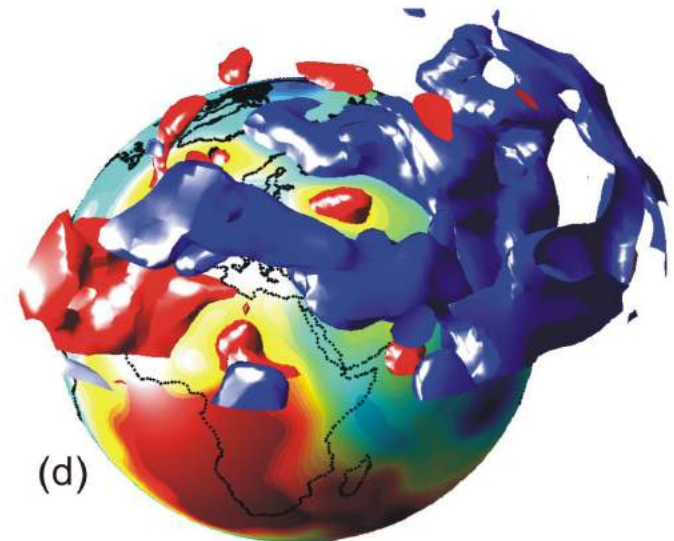
## Existence of a THIRD mantle box ?



(Torsvik et al, 2008)



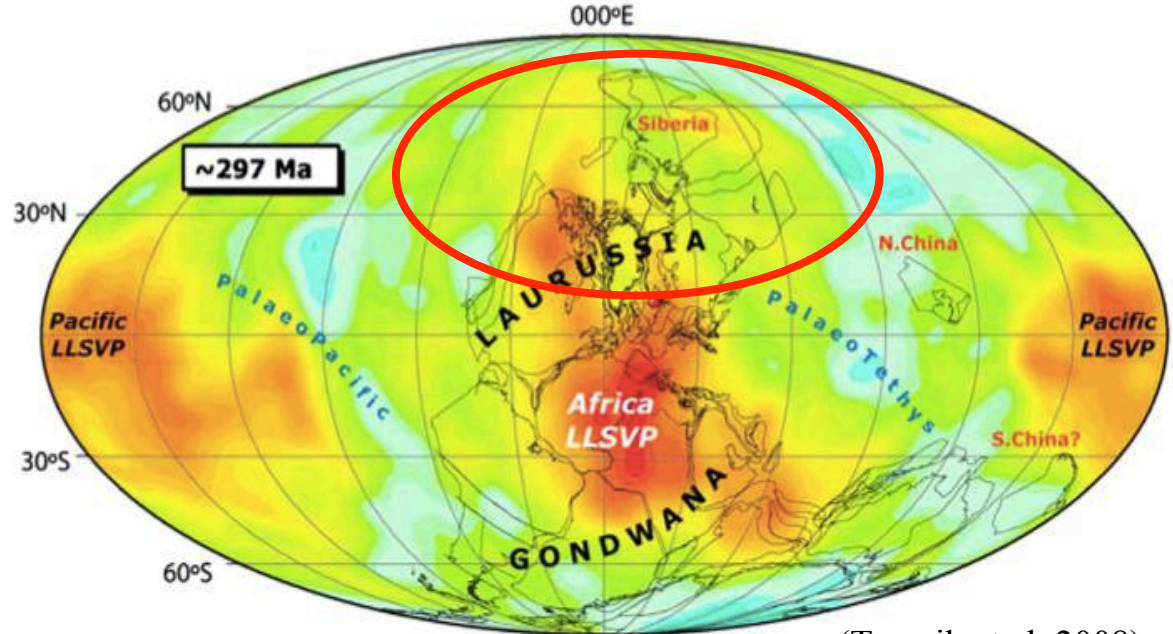
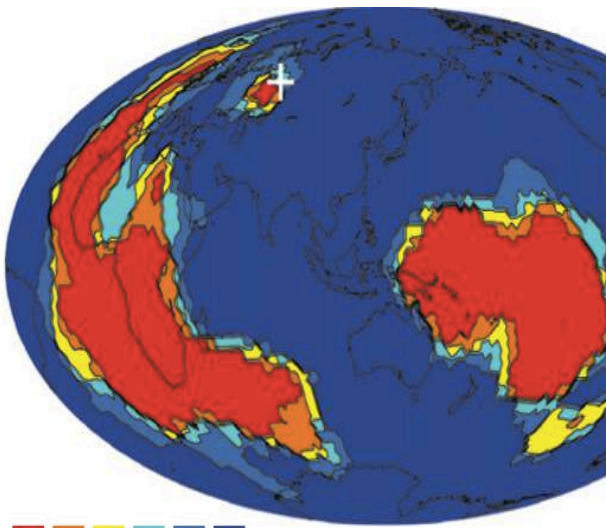
(c)



(d)

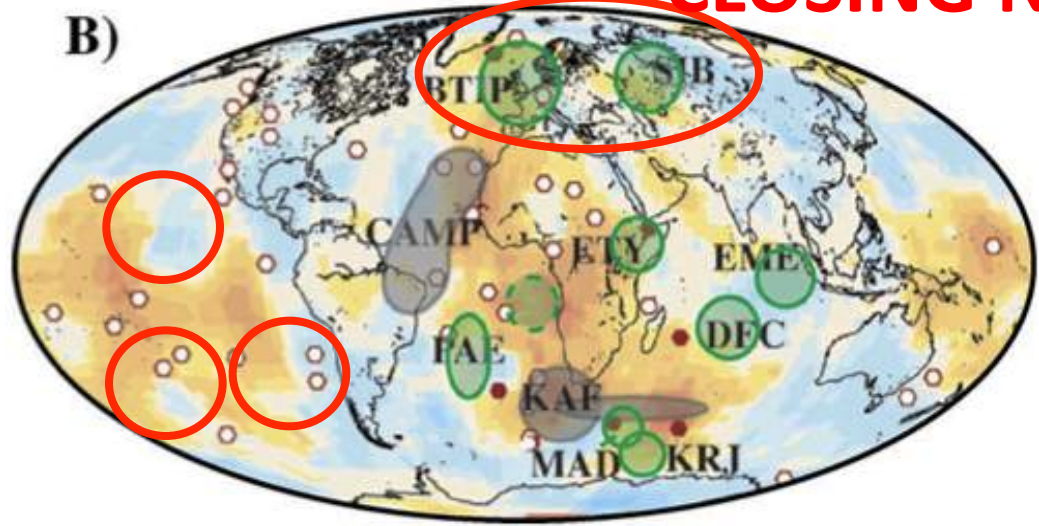
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(Torsvik et al, 2008)

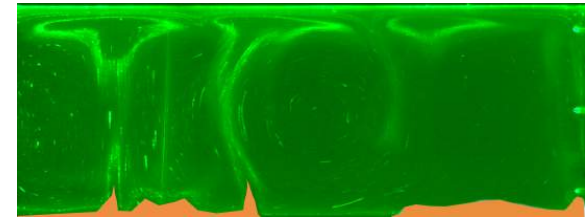
## CLOSING NOW ?



(d)

(A.D, A. Ismail-Zadeh, J. Besse 2013, 2016)

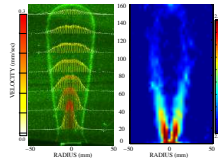
# CONCLUSIONS :



**A- Origin of mantle « boxes » ? Variable viscosity / Subduction**

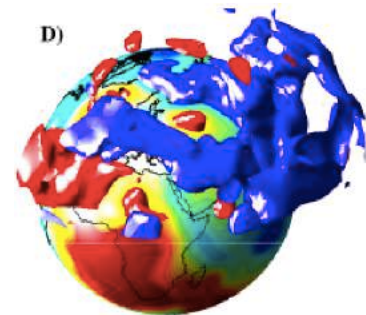
**B- Upwellings                      Compositional/density heterogeneities**

- several types ?
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- what happens around 1000 kms ?



**C- Do the mantle boxes / LLSVPs change shapes through time ?**

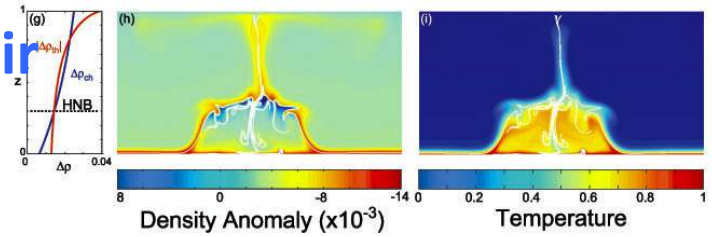
- **LLSVP morphology shaped by :**
  - . **plume head emission** (~30 Myr)
  - . **subduction location** (~ 100Myr - 1 Gyr)
  - . **heating up of old subducted material** (~ Gyr)
- **LLSVP = old material but CHANGING morphology**
- **3<sup>rd</sup> box around Siberian Trap => closing ?**



**THANK YOU !**



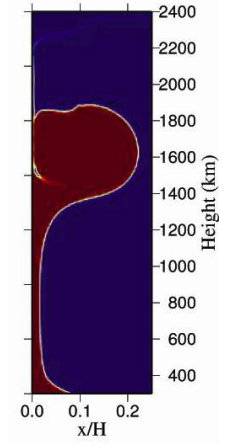
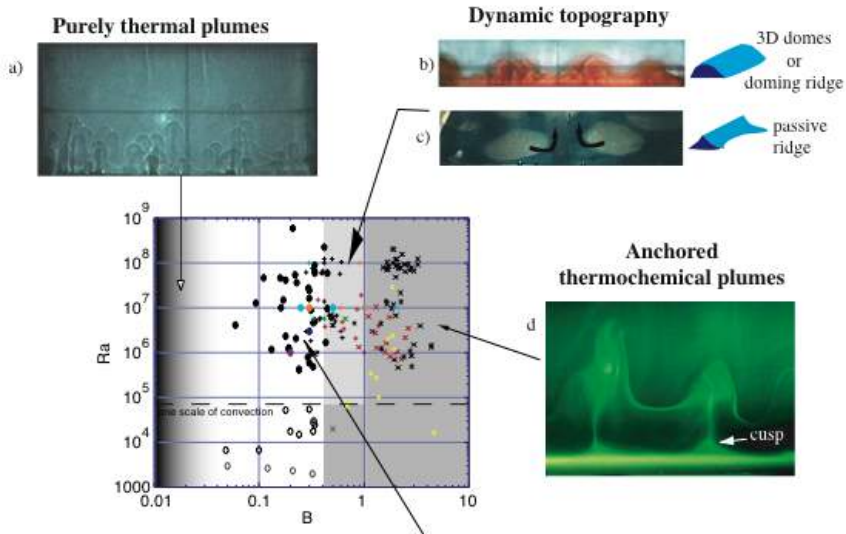
# 4. Interaction Convection / Denser Reservoir



## Stability:

-  $Ra = \frac{\alpha \Delta T \rho g L^3}{\eta \kappa}$

-  $B = \frac{\Delta \rho_x / \rho}{\alpha \Delta T}$



Tan & Gurnis, 2005;  
Samuel & Bercovici, 2006

## Morphology:

-  $\gamma = \eta_1 / \eta_2$

-  $a = d_1 / d_2$

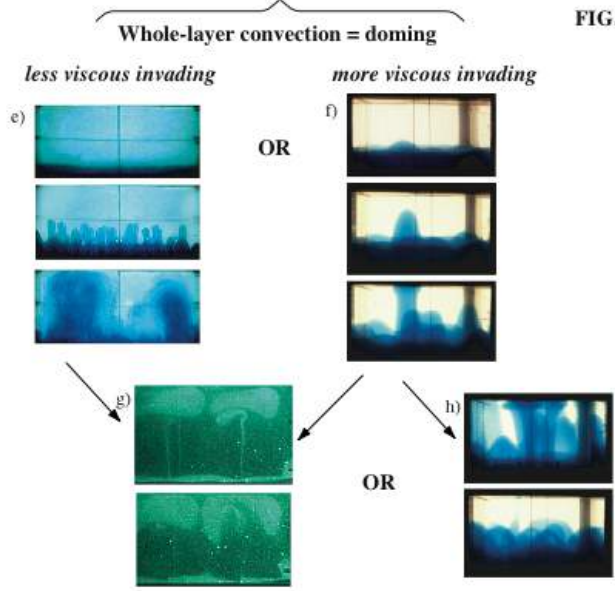


FIG.5

Richter & McKenzie, 1981;  
Olson, 1984;  
Schmeling, 1988;  
Olson & Kincaid, 1991;  
Tackley, 1998; 2002;  
Davaille, 1999; Davaille & al, 2002;  
Kellogg & al, 1999;  
Montague & Kellogg, 2000;  
Hansen & Yuen, 2000;  
Le Bars & Davaille, 2002, 2004;  
Jellinek & Manga, 2002, 2004;  
Samuel & Farnetani, 2002, 2005;  
McNamara & Zhong, 2004, 2005  
Lin & Van Keken, 2005